# Retrieval of sea ice drift in the Fram Strait based on data from Chinese satellite HaiYang (HY1-D)

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**Abstract.** Melting of sea ice in the Arctic has accelerated due to global warming. The Fram Strait (FS) serves as a crucial pathway for sea ice export from the Arctic to the North Atlantic Ocean. Monitoring sea ice drift (SID) in FS provides insight into how Arctic sea ice responds to the climate change. The SID has been retrieved from Sentinel-1 SAR, AVHRR, MODIS and AMSR-E, and further exploration is needed for the retrieval of SID using optical imagery. In this paper, we retrieve SID in the FS using Chinese HaiYang1-D (HY1-D) satellite equipped with the Coastal Zone Imager (CZI). Multi-template matching technique is employed to calculate cross-correlation, and subpixel estimation is used to locate displacement vectors from the cross-correlation matrix. The dataset covering March to May 2021 was divided into hourly and daily intervals for analysis, and validation was performed using Copernicus Marine Environment Monitoring Service (CMEMS) SAR-based product and IABP buoy. A comparison with the CMEMS SID product revealed a high correlation with the daily interval dataset; however, due to the spatial and temporal variability of sea ice motion, differences are observed with the hourly interval dataset. Additionally, validation with IABP buoy yielded a velocity bias of -0.005 m/s and RMSE of 0.031 m/s for the daily interval dataset, along with a flow direction bias of 0.002 rad and RMSE of 0.009 rad respectively. For the hourly interval dataset, the velocity bias was negligible (0 m/s) with a RMSE of 0.036 m/s, while the flow direction bias was 0.003 rad with a RMSE of 0.010 rad. In addition, during the validation with buoys, we found that the accuracy of retrieving the SID flow direction is distinctly interrelated with the sea ice displacement.

## 1 Introduction

The Arctic, as one of the three poles of the Earth (Li et al., 2020) that stores 101,000 km$^3$ of fresh water (Finnish Meteorological Institute et al., 2022), is an important part of the cryosphere and plays a prominent role in global water resources and atmospheric cycles. Since the beginning of the twenty-first century, global warming has profoundly affected human production and activities and has become a strong threat to the stability of the climate system (Cook et al., 2014). Under the effect of Arctic amplification (Serreze et al., 2009), Arctic quick warming accelerates the melting of polar sea ice, leads to thinner sea

35  ice and accelerates sea ice transport (Krumpen et al., 2016; Maslanik et al., 2011). The acceleration of sea ice motion also indicates a reduction of the sea ice residence time in the Arctic (Sumata et al., 2023). Moreover, changes in sea ice feed back into the climate system, affecting energy transport (Dethloff et al., 2006; Döscher et al., 2014).

Sea ice drift (SID) is an important geophysical parameter to describe the dynamic of sea ice and the sea ice motion under the influence of winds, currents, and various external forces (gravity, Coriolis effect, etc.) (Encyclopedia of Ocean Sciences, 2022).

40  The primary SID circulation across the Arctic encompasses both the Beaufort Gyre (BG) and the Transpolar Drift (TPD) (Preller and Posey, 1989), and TPD transports large quantities of multiyear ice outward from the central Arctic toward the FS, Barents Sea and Baffin Bay (Colony and Thorndike, 1984; Martin and Augstein, 2000). The FS connects the Arctic and the North Atlantic (Sumata et al., 2022), and large quantities of sea ice are injected into the North Atlantic each year through the strait (Reimnitz et al., 1994). The long-term annual average ice outflow in the FS (1935-2014) is approximately 880,000 $km^3$,

45  accounting for 10% of the sea ice cover in the Arctic basin (Smedsrud et al., 2017) and the largest portion (90%) of the Arctic sea ice export volume (Sumata et al., 2022; Haine et al., 2015; Serreze et al., 2006). Sea ice, a mixture of ice and brine (Schwerdtfecer, 1963), which gradually disintegrates during outward transport in the FS. This process affects freshwater exchange and energy transport in the North Atlantic, which may alter the convective overturning of water masses and thermohaline circulation processes (Aagaard and Carmack, 1989). In addition, for shipping planning and scientific research,

50  the progression of polynyas and lead is heavily influenced by SID (Wagner et al., 2021). Therefore, observing SID in the FS is crucial to analyse the sea ice variation in the Arctic and the sea ice transport from the polar to sub-polar regions.

The development of satellites and remote sensing sensors promote satellite data as a prevailing trend in retrieving SID (Kwok, 2010). The SID retrieved from satellite-derived data possesses continuity in spatial and temporal, which helps researchers better evaluate the sea ice variation in the Arctic and provides a basis for climate forecasting and ship route planning. At present,

55  the primary data used in SID retrieval are radiometer, scatterometer, synthetic aperture radar (SAR) and optical imagery. SID products derived from radiometers and scatterometers inherently possess coarse spatial resolution owing to the characteristics of the sensors. Low resolution SID products exhibit large error on velocity in FS (Hwang, 2013). OSI SAF provides SID products retrieved from scatterometers and radiometers over the polar regions and its temporal coverage is from 2009 to now, however, the time interval of these products is greater than one day. Due to the complexity of sea ice dynamics, low temporal

60  resolution SID product may fail to depict the subdaily-scale variation of the sea ice motion. High spatiotemporal resolution SID products are valuable for capturing subtle SID patterns (Johansson and Berg, 2016). SAR image provides high spatial resolution, but radar backscatter is sensitive to the liquid water on sea ice surface, and the accuracy of retrieving SID during the melting season can be hampered (Stern and Moritz, 2002). This disadvantage constrains the application of SAR in retrieving SID. Optical imagery has been applied extensively in cryosphere observation. Many in orbit satellites equipped with optical

65  imagery and optical imagery have the advantages of wide swath and high resolution, which are beneficial for capturing the surface roughness and textural characteristics of sea ice. Optical imagery has the potential to retrieve SID in FS. While optical imagery has the advantages of high resolution and wide swath, this can also significantly increase the computational effort of the algorithm. To solve this problem, we applied the MCC (maximum cross-correlation) algorithm with improvements for SID

retrieval. The MCC has been widely used in retrieving SID with radiometer and scatterometer as well as with SAR (Girard-Ardhuin and Ezraty, 2012; Hollands and Dierking, 2011). However, when it comes to utilizing optical imagery for SID retrieval, many researchers tend to use different algorithms instead of the MCC. [Petrou et al.] designed an algorithm based on optical flow to retrieve SID from MODIS (Petrou and Tian, 2017). However, due to the absence of buoy in their study area, the accuracy of their algorithm is unevaluated. [Lopez-Acosta et al.] developed a complex ice tracking algorithm specifically designed for retrieving SID in FS but requiring multiple MODIS images to generate a complete SID field (Lopez-Acosta et al., 2019). [Fang et al.] discussed the potential of feature tracking with MODIS images for SID retrieval. The accuracy of their result is promising (Fang et al., 2023), but the spatial coverage needs to be further improved. [Wang et al.] designed an ice tracking algorithm to retrieve SID in the marginal ice zone (MIZ), but the algorithm cannot produce an intact SID field (Wang et al., 2021). The European Organization for the Exploitation of Meteorological Satellites (EUMETSAT) uses the MCC algorithm to retrieve SID from AVHRR with a spatial resolution of 20 km and a temporal resolution of 1 day (Dybkjaer, 2018). However, it has been found that the accuracy of the SID product retrieved from AVHRR presents low accuracy in East Greenland, with the Mean Absolute Error (MAE) of velocity reaching 10.40 km/day, which is even lower than that of the SID products retrieved from radiometer and scatterometer (Wang et al., 2022). In our study, we applied multi-template matching combined with a subpixel estimation for SID retrieval. The method of multi-template matching achieves a balance between computational efficiency and accuracy, and the subpixel estimation further improves the accuracy of the retrieval results while saving computational resources. We believe that the MCC algorithm with improvements can be utilized for retrieving SID effectively from optical imagery.

This paper focuses on designing an improved MCC algorithm using the Coastal Zone Imager (CZI) of Haiyang-1D (HY-1D) data to retrieve SID in the FS. Moreover, the retrieved results are evaluated via comparison with SAR-based SID products released by CMEMS and IABP buoy data. For sea ice motion, the average velocity of the sea ice motion in Arctic is 0.02 m/s (Colony and Thorndike, 1984), and the maximum velocity of the sea ice in the FS is 0.64 m/s (Lei et al., 2016). The multiyear ice (MYI) drifts from the Arctic basin and crushes in the central part of strait, which results in the fragmented ice in the southern part of the FS and along the eastern coast of Greenland. The fragmented ice drifts at a faster velocity near the FS compared to sea ice within the Arctic Basin. The sea ice characteristics described above increase the difficulty of retrieving SID in FS.

This paper is structured as follows. Section II describes the data used in this study, including the CZI image, CMEMS SID products and IABP buoy. Section III details the algorithm used to retrieve SID. Section IV compares the results with CMEMS SID products and buoy data for validation. Section V discusses the retrieved results. Finally, concluding remarks are provided in Section VI.

## 2 Data

### 2.1 HY-1D CZI data

100 Haiyang-1D (HY-1D) satellite was launched in 11 June 2020. The satellite is equipped with the Chinese Ocean Color and Temperature Scanner (COCTS), CZI, Ultraviolet Image (UVI), Satellite Calibration Spectrum, and Automatic Identification System. The revisit period of the CZI is three days, the resolution of the CZI image is 50 m, and the swath of the CZI image is nearly 950 km. CZI possesses four bands（VIR and NIR）, as shown in Table 1. The wide swath and high spatial resolution of CZI imagery make it suitable for the sea ice motion observation in FS.

105 Since the launch of China's first ocean satellite in 2002, concerted efforts have been undertaken to institute a comprehensive global operational ocean satellite observation system. Currently, the observation system consists of 10 satellites, which include three series: ocean color series satellites (HY-1), ocean dynamic environment series satellites (HY-2), and ocean surveillance and monitoring series satellites (HY-3) (Zeng et al., 2023). In this paper, SID retrieval is performed using the L1C data which is processed with radiometric calibration and geographic projection. Before retrieving, the CZI images are resampled to 300 110 m considering the algorithm's computational efficiency and the spatial resolution of the result.

**Table 1: Information on the HY-1D CZI data.**

| Band | Wavelength (μm) | Band name |
|------|-----------------|-----------|
| 1 | 0.42-0.50 | Blue |
| 2 | 0.52-0.60 | Green |
| 3 | 0.61-0.69 | Red |
| 4 | 0.76-0.89 | Near Infrared |

In the FS, the climate condition is harsh and sea ice changes rapidly in response to wind (Tsukernik et al., 2010), which is a potentially hazardous factor for shipping and scientific expeditions, achieving timely and accurate monitoring of SID is a crucial goal. The CZI image has the high spatial resolution and wide swath for providing accurate monitoring of the SID. In 115 our study, 111 CZI images from March to May 2021 were used to produce the intersection region (as shown in Fig. 1(a)) for the dataset, and 48 pairs of images were eventually produced for SID retrieval. Furthermore, to explore the influence of the image time interval on retrieval SID, the 48 pairs of image datasets are divided into 16 pairs of hours-level (less than 6 hours) dataset and 32 pairs of day-level (approximately 24 hours) dataset. Fig. 1(b) shows the time intervals of image pairs, and Fig. 1(c) plots the outer frame lines of the image pairs, where the white borders are the locations of the image pairs with hours-level 120 interval, and the darker blue borders are the locations of the image pairs with day-level interval.
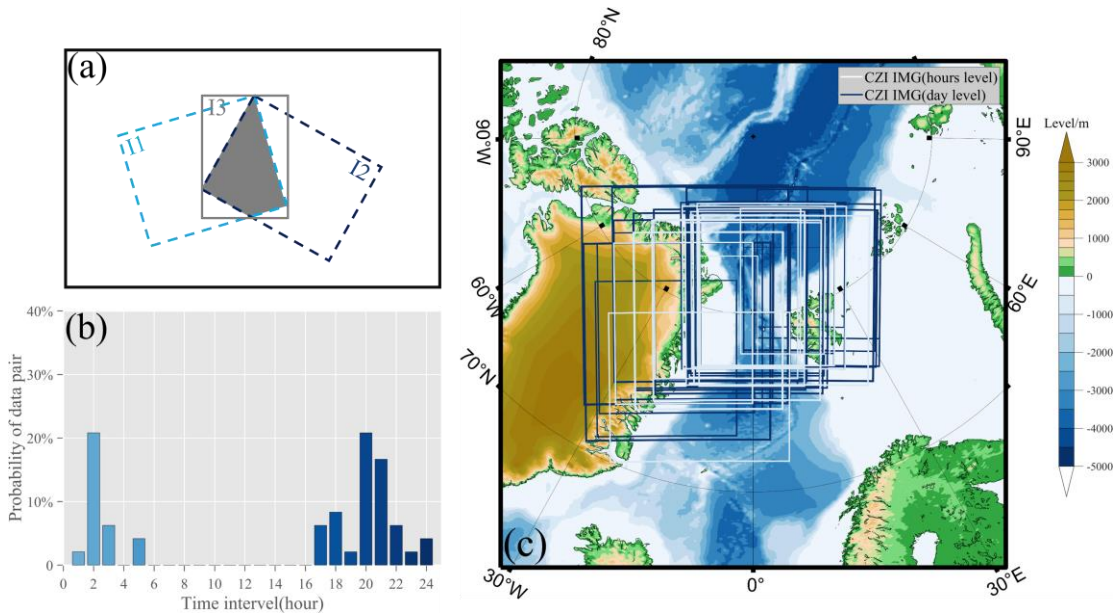
Figure 1: An illustration showing the creation of the dataset using CZI images (a). The time interval histogram of the dataset (b) and the location of image pairs (c).

## 2.2 CMEMS SID product

125   As the CZI images possess a high spatial resolution, the SID products retrieved from the scatterometer and radiometer are not comparable for validation. Thus, the SAR-based SID product is chosen for comparison. The Global Ocean-High Resolution SAR Sea Ice Drift is a polar near real-time gridded sea ice drift product produced by the National Space Institute at the Technical University of Denmark (DTU space) (European Union-Copernicus Marine Service, 2015). The data is released by the Copernicus Marine Environment Monitoring Service (CMEMS) and distributed on the website of Copernicus Marine

130   Service. The composite product is updated every 12 hours covering 24 hours. Thus, two kinds of product are provided, with a nominal time intervals of 0:00 to 0:00 and 12:00 to 12:00, respectively. The grid resolution of the product is 10 km, and the product is available from both Arctic Ocean and South Ocean (Pedersen et al., 2015). The product has been validated using Woods Hole ice tethered profilers (ITPs). For the validation of 2021, the number of matched pair is 29180, the correlation coefficient between the product and ITP buoys is 0.99, and RMSD of dx and dy is 362.32 m (0.0042 m/s) and 339.81 m (0.0039

135   m/s), the BIAS of dx and dy is 4.64 m and 17.29 m and the BIAS of velocity is negligible (~0 m/s). The validation is performed with the 24-hour mean composite product. (European Union-Copernicus Marine Service, 2015). Sea ice motion changes rapidly in the FS, and SID results retrieved from images at different times is incompatible. Therefore, the CMEMS SID product which has the most temporal overlap with CZI images is chosen for comparison. Additionally, due to the disparity in spatial resolution, we resampled the retrieved SID to the resolution of the CMEMS SID product by linear interpolation.

**2.3 IABP buoys**

GPS position data from buoys are the most credible data for the retrieved SID validation. The International Arctic Buoy Programme（IABP）has deployed a network of drifting buoys in the Arctic to provide data for scientific research (Rigor et al., 2008). IABP has collected more than 500 months of data from more than 500 buoys deployed in the Arctic. Buoy data collected from IABP are an essential component of Arctic scientific data. These data have been used to validate retrieved SID (Hakkinen
145    et al., 2008; Fang et al., 2023). The drift trajectories of the 50 buoys used in our study are shown in Fig. 2 and these drift trajectories consist of consecutive GPS points. Among the buoys used for validation, 31 buoys from MOSAiC were included. The GPS position of MOSAiC buoy has an accuracy of ± 2.5 m (Qiu and Li, 2022) and the IABP buoy has an accuracy of approximately 300 m (Haarpaintner, 2006). We set a minimal distance of 4 km to search the buoys for validation. Because some of the buoys were deployed in the vicinity, we used the average of them to validate the result. Ultimately, 206 matched
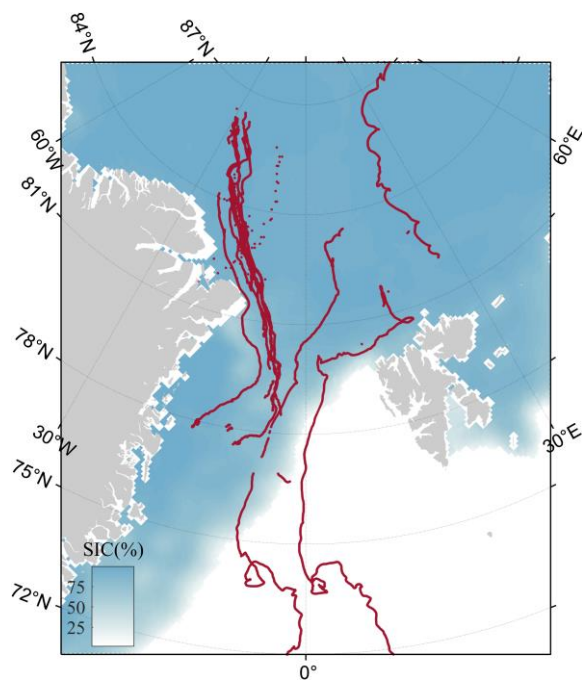150    drift vectors were produced for validation.



**Figure 2: Drifting trajectories of 50 IABP buoys from March to May 2021, SIC (sea ice concentration).**

**3 Improved SID retrieval method in the Fram Strait**

155    The flowchart of our study is shown in Fig. 3. We choose same orbit images of different moments to calculate the intersection and the intersection area is clipped to generate dataset. Data enhancement in preprocessing is used to distinguish between sea ice and seawater and highlight the textural details of the sea ice. Data enhancement is practical for retrieving SID using optical

imagery (Fang et al., 2023; Lopez-Acosta et al., 2019; Yan et al., 2023). After preprocessing, the sea ice displacement will be retrieved via multi-template matching and subpixel estimation. Since the optical images are susceptible to various factors, such as cloud and the angle of incidence of the light (Stow et al., 2004), the SID contains noise and low-quality result. We designed the quality control session to remove the low-quality result and noise from the SID. The penultimate step involves smoothing and filling to derive an intact SID field. Finally, we validate the results using buoys and the CMEMS SID product.
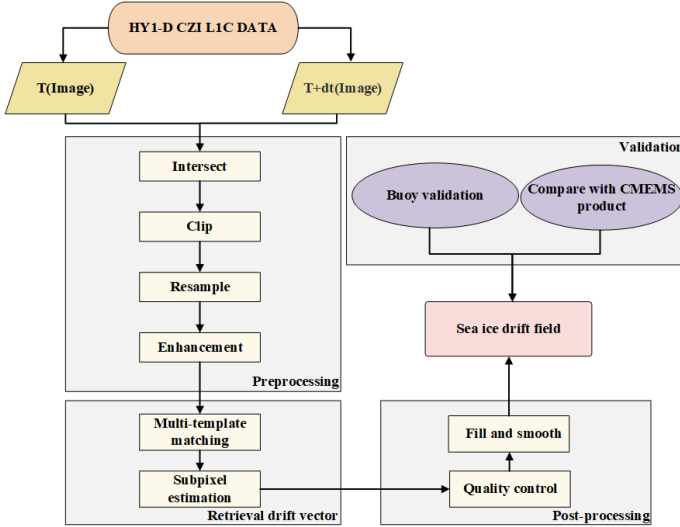


Figure 3: Flow chart of sea ice drift retrieval.

## 3.1 Data enhancement

Before preprocessing, the resampled image radiance values are converted to grayscale (0-255). Then, the data of different bands are stretched logarithmically. After reviewing study in which optical imagery was used to retrieve SID (Fang et al., 2023), we selected the average of the bands 2, 3 and 4 as the input data. Finally, histogram equalization and edge detection are applied to the image. The process and effect of data enhancement are explicitly described as follows.

In the polar regions, the lighting conditions are complex, and some images are captured at slight solar zenith angle (Stow et al., 2004). These disadvantages could lead to an uneven spatial distribution of image gray values, which affects the retrieval of SID. Logarithmic stretching is an image grayscale transformation that expands the low gray value portion and compresses the high gray value portion (Lupton et al., 2004). Equation (1) is the formula for logarithmic stretching, where $r$ is the gray value of the image and $c$ is the base. The larger the base is, the stronger the emphasis on the lower gray portions and the stronger the compression of the higher gray portions. As shown in Fig. 4(a) and (b), logarithmic stretching improves the uneven lighting and enhances the sea ice texture.

$$T(r) = c * log(1 + r), \tag{1}$$

In addition to logarithmic stretching, histogram equalization (HE) is also useful for adjusting the grayscale distribution of image (Krutsch and Tenorio, 2011). The adaptive histogram equalization (AHE) algorithm is a variant of histogram

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180    equalization (Pizer et al., 1987). The contrast is changed by calculating the local histogram of the image and then redistributing the value. The algorithm is an improvement over the traditional HE, but the disadvantage of AHE is that it introduces noise during grayscale stretching. Currently, contrast limited adaptive histogram equalization (CLAHE) has improved upon AHE (Zuiderveld, 1994). Therefore, we chose CLAHE to redistribute the gray values. The image histograms during grayscale transformation are shown in Fig. 4(d). Logarithmic stretching improves the gray value of the image (Fig. 4(b)), and CLAHE

185    enhances the texture and gap features in high sea ice concentration regions (Fig. 4(c)).

Cloud can lead to loss of contrast and diminished textural detail, affecting the quality of the result. Edge detection, the detection and characterization of significant intensity changes (Torre and Poggio, 1986), is valuable for reinforcing the texture differences between cloud and sea ice. In our study, Sobel edge detection is applied to enhance the texture, and we find that edge detection can also neutralize some disadvantageous effect of cloud. We will describe this aspect thoroughly in the

190    discussion section.
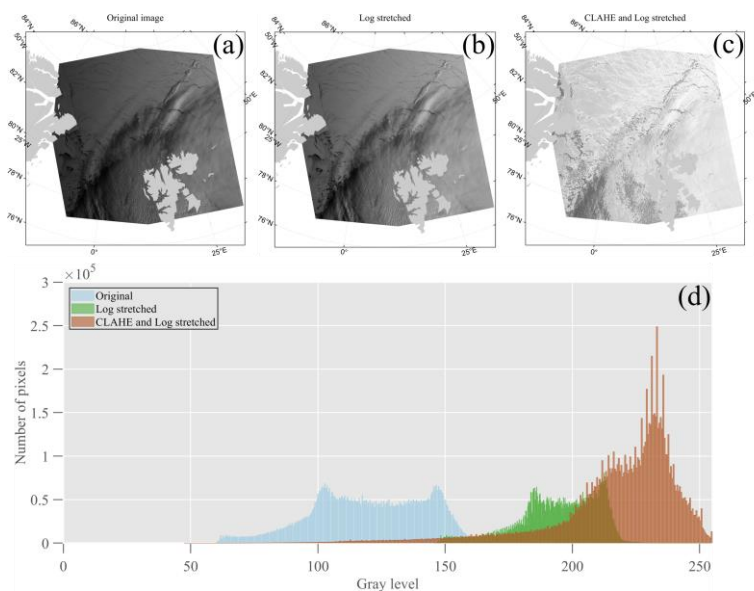


**Figure 4: An example of a CZI image (average of bands 2, 3 and 4) before and after data enhancement and its grayscale histogram. (a) is the original image and the color of its histogram is light blue, (b) is the image after logarithmic stretching and the color of its histogram is green, (c) is the image after logarithmic stretching and CLAHE and the color of its histogram is brown, and (d) is the**

195    **grayscale histogram of the above images. The date of the image is 07 April 2021, 07:16:35 LT.**

### 3.2 Sea ice drift retrieval

### 3.2.1 Multi-template matching

The traditional template matching requires searching in the vicinity of template to determine the location with the largest correlation coefficient, which consumes considerable time. Multi-template matching is performed with different template sizes,

200    ranging from coarse resolution to fine resolution, reducing calculation time while ensuring accuracy (Wang et al., 2008). As shown in Fig. 5, the time cost of multi-template matching for the 48 datasets is less than 150 s, and the average number of

image pixels uses for the retrieval is 5,208,298. This indicates that multi-template matching has high computational efficiency for wide-swath images. In addition, cross-correlation calculation is required during template matching, and the calculation is slow in the spatial domain (Thielicke and Stamhuis, 2014) and cannot yield a correlation matrix with equal size to the template. Therefore, we calculate the correlation in the frequency domain using Fast Fourier Transform (FFT).
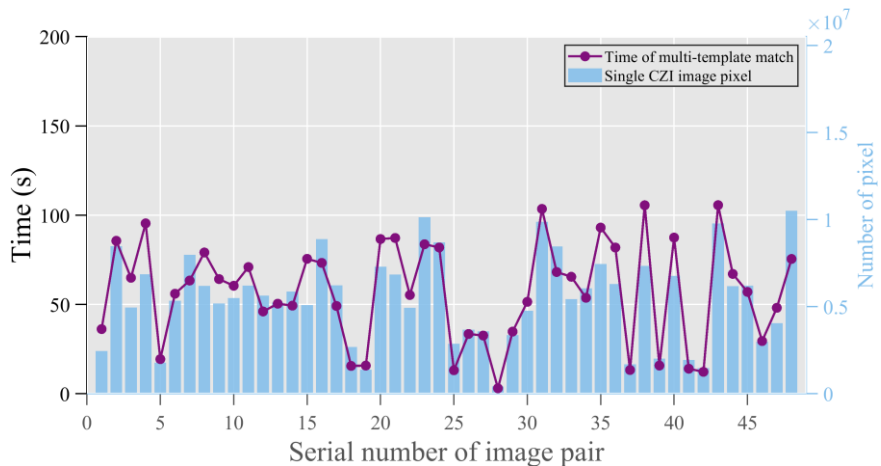


**Figure 5: Relationship between multi-template matching time and image size.**

### 3.2.2 Subpixel estimation

Identifying the peak of the correlation coefficient is an essential factor for improving the accuracy of retrieval. The integer displacements of the two templates can be determined directly from the peak positions of the correlation matrix. However, this approach introduces quantization error, and the difference of the minimum displacements will be multiples of the image resolution (Lavergne et al., 2010). To address this issue, subpixel precision refinement of the peak position with the biquadratic surface fitting method has been proposed to mitigate quantization errors (Kwok et al., 1998). Subpixel estimation could result in considerable computational costs when using sophisticated fitting (Hollands and Dierking, 2011). Therefore, for our study, we have opted for simple Gaussian fitting since this approach strikes a balance between computational efficiency and accuracy (Thielicke and Stamhuis, 2014).

### 3.3 Quality control

The retrieved SID obtained through multi-template matching and subpixel estimation is presented in Fig. 6(a). However, due to the influence of cloud and seawater, certain areas of SID have low confidence level. The algorithm may extract erroneous SID vectors, and it is necessary to filter those erroneous vectors. Therefore, we employ a series of methods to control the quality of the SID. By applying quality control, consistency within the SID is ensured while eliminating erroneous drift vectors. Subsequent sections provide detailed descriptions of the quality control procedures.

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### 3.3.1 Cross-correlation value and its derived parameters filter

The magnitude of the cross-correlation (R) in the MCC serves as an indicator of result quality (Haarpaintner, 2006). It
represents the correlation between two templates, making it an effective parameter for the initial quality control. Additionally,
we employ two parameters derived from the cross-correlation matrix to further control the quality. In the selection of peak,
incorrect peaks which are observed near the peak can be labeled as subpeaks. The ratio of peak to subpeak was used to measure
of the 'uniqueness' of the real peak. Furthermore, comparing the average of the correlation coefficient matrix with the peak
provides a measure of the signal-to-noise level of the peaks. Two parameters, the peak mean ratio (PMR) (Eq. 2) and the peak
second ratio (PSR) (Eq. 3) (Dybkjaer, 2018; Van Wyk de Vries and Wickert, 2021), are used to measure the quality of the
result.

$$PMR = \frac{C_{peak}}{mean(abs(C_{all-corr})}, \tag{2}$$

$$PSR = \frac{C_{peak}}{C_{subpeak}}, \tag{3}$$

In the above equation, $C_{peak}$ is the peak of cross-correlation, $C_{all-corr}$ is the cross-correlation matrix, and $C_{subpeak}$ is the second
peak of cross-correlation. Fig. 6(c) shows the result obtained using a neighborhood filter and co-filtering with the cross-
correlation value and its derived parameters filter. Compared to Fig. 6(b), which solely employs the neighborhood filter,
incorporating the cross-correlation and derived parameters filter effectively eliminates erroneous outcomes within open water
regions.
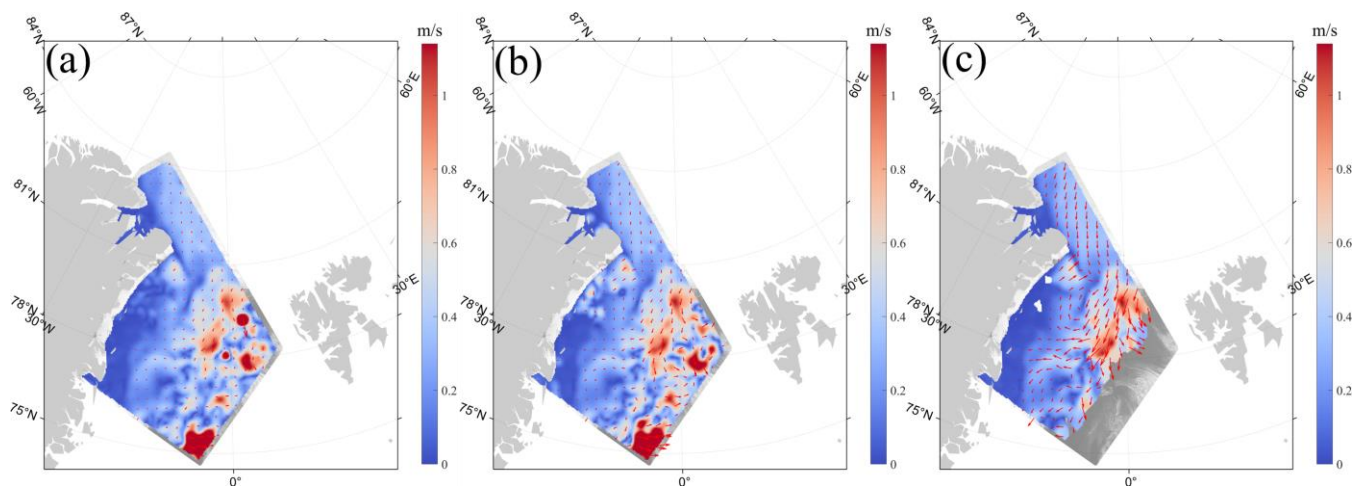


**Figure 6: An example of quality control applied to the retrieved SID. (a: SID with no quality control, b: SID after neighborhood filter (only), c: SID after using neighborhood filter and cross-correlation value and its derived parameters filter).**

### 3.3.2 Neighborhood filter

The neighborhood filter is commonly employed for postprocessing retrieved SID data by removing abrupt values indicative of incidental anomalous drifts that deviate from continuous sea ice movement patterns (Hyun and Kim, 2017; Girard-Ardhuin and Ezraty, 2012). In our study, two window sizes are designed for neighborhood filter. Fig. 6(b) shows the results of using the neighborhood filter only. Compared to Fig. 6(a), the neighborhood filtering removes the regions (upper part of the image) with significant velocity differences in the SID, but only using only the neighborhood filter is not incomplete. Combining the cross-correlation value and its derived parameters filter is necessary to comprehensively control the quality of the result.

### 3.4 Fill and smooth

The utilization of correlation coefficient and its derived parameters filter and neighborhood filter introduces missing values to the results. Given that the sea ice is spatially distributed continuously, adjacent values are employed to fill in these gaps and smoothen the outcomes during the final step of retrieval
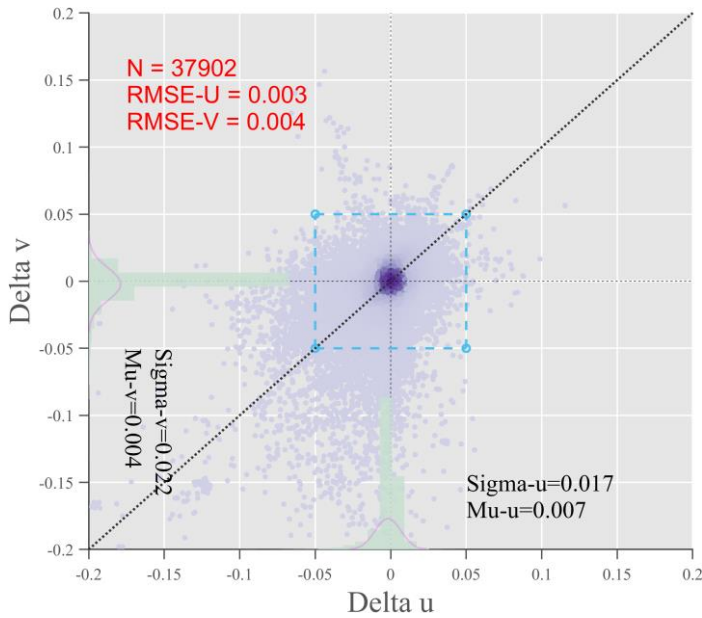
### 4 Results and validation

In this study, SID in FS is retrieved from the CZI images with a resolution of approximately 4 km. The results are categorized into day-level and hours-level datasets based on the time interval of images and compared with the IABP buoy and the CMEMS SID product.

### 4.1 Comparison of results with CMEMS products

### 4.1.1 SID with day-level time intervals

Higher-resolution SID is retrieved using CZI images. The SID products based on radiometer or scatterometer are incomparable with the retrieved SID due to the spatial resolution differences. Therefore, the SAR-based SID product (10 km) published by CMEMS is selected for comparison. The reliability of the CMEMS product has been verified in the released user manual (European Union-Copernicus Marine Service, 2015).

**Figure 7: The U-V difference between the retrieved SID (day-level) and the product. The histograms on the x and y-axes show the distribution of differences in the U and V directions, and the pink curves fit the histograms. The dashed blue box is 0.05 m/s, and the color shading reflects the density of the scatter distribution.**

Fig. 7 shows the U-V difference between the retrieved SID (day-level) and the CMEMS SID products, totaling 37,902 matched points. The U-V differences between the retrieved SID and the product are relatively slight. The majority of the U-V differences are less than 0.05 m/s and the majority of the scatter points are distributed in the center of the plot. The RMSE of U and V are 0.003 m/s and 0.004 m/s, respectively. A strong correlation between the product and the retrieved SID is shown by the histograms and accompanying fitted curves with a normal distribution on the X and Y axes.
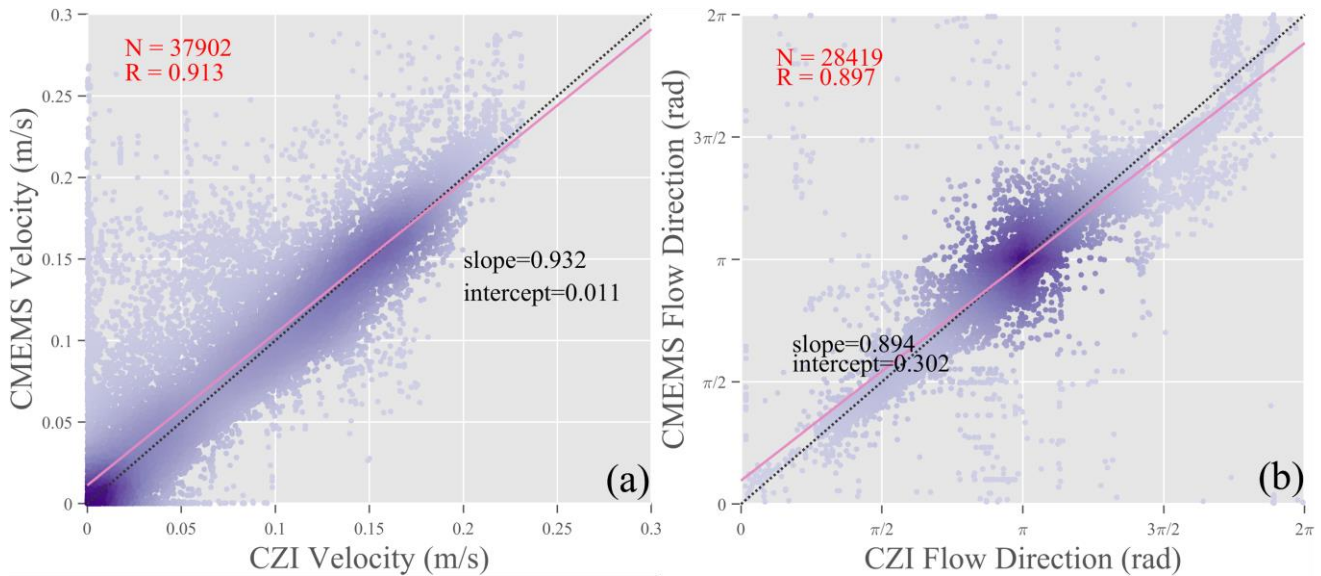
**Figure 8: The difference in velocity (a) and flow direction (b) between the retrieved SID (day-level) and the product. The dashed line denotes the ideal line, and the pink line denotes the fit line.**

275 Fig. 8(a) shows the velocity difference between the retrieved SID and the CMEMS SID product, involving a total of 37,902 matched points. The retrieved SID demonstrates a strong correlation with the product in terms of velocity, as indicated by a high correlation coefficient of 0.913 and a fitted line that closely aligns with the ideal line. The statistics of the difference between the retrieved SID and the product are presented in Table 2, revealing RMSE of 0.003 m/s for the U component, 0.004 m/s for the V component, and 0.005 m/s for the velocity. These slight differences reflect the consistency between the retrieved

280 SID and the product.

The size of the template and search area is associated with the spatial resolution of result. To retrieve high spatial resolution SID in the FS, limited template size and search area is set in our study. Therefore, the maximum velocity of day-level result is lightly smaller than the product. The effect of template size and search area on retrieving SID will be explored in future research. Fig. 8(b) shows the flow direction difference between the retrieved SID and the product using a total number of 28419 matched

285 points. The correlation of the flow direction between the retrieved SID and the product is great, as indicated by a correlation coefficient of 0.897. The statistics of the difference in flow direction between the retrieved SID and the product are shown in Table 2.

**Table 2: Validation of the retrieved SID (day-level) with the product.**

| Delta | Number of match point | BIAS | MAE | STD | RMSE |
|---|---|---|---|---|---|
| U (m/s) | | -0.003 | 0.009 | 0.017 | 0.003 |
| V (m/s) | 37902 | -0.004 | 0.013 | 0.024 | 0.004 |
| Velocity (m/s) | | -0.005 | 0.014 | 0.027 | 0.005 |
| Flow direction (rad) | 28419 | 0.009 | 0.147 | 0.369 | 0.009 |

Regarding of flow direction, the flow direction of sea ice motion in the FS predominantly ranges from "$\pi/2$" to "$3\pi/2$" (with

290 the positive y-axis as the starting point in polar stereographic projection). The most abundant flow direction is "$\pi$", which indicates the southern part of the FS. In the comparison of flow direction, it is observed that the shelf ice on the east coast of Greenland has a slower velocity (less than 500 m/day) than the drift ice in the strait. Due to the potential inaccuracy associated with short sea ice displacement, a shelf ice mask was been established to exclude the slow-moving shelf ice. The impact of sea ice displacement on retrieval accuracy is analyzed in the discussion section.

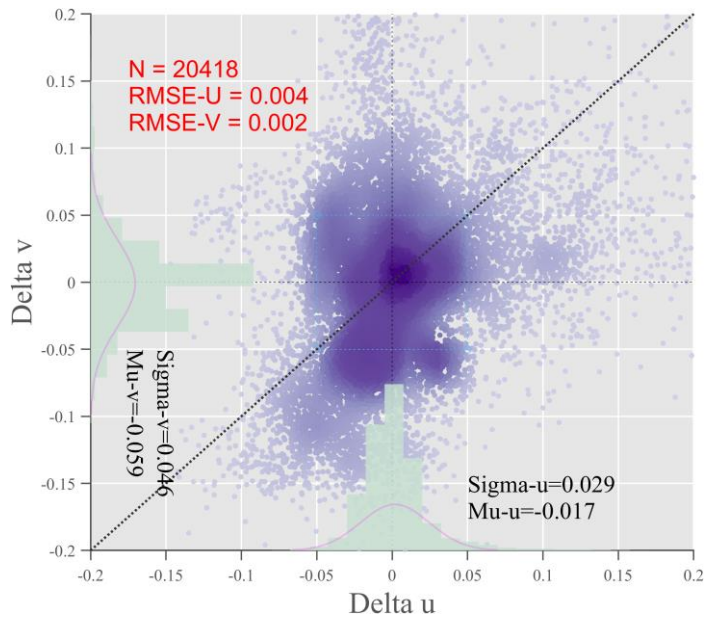**4.1.2 SID with the hours-level time interval**



**Figure 9: The U-V difference between the retrieved SID (hours-level) and the product. The histograms on the x and y-axes show the distribution of differences in the U and V directions, and the pink curves are a fit to the histograms. The dashed blue box is 0.05 m/s, and the color shading reflects the density of the scatter distribution.**

300 Fig. 9 illustrates the difference in U-V between the retrieved SID (hours-level) and the CMEMS SID product, with a totaling 20,418 matched points. The scatter points distribution of the U-V differences in Fig. 9 appears to be significantly more dispersed than that in Fig. 7, suggesting a notable disparity between the retrieved SID and the product. The apparent difference may be associated with the spatial and temporal variability of sea ice motion and will be explicitly analyzed in the discussion section.

**4.2 Comparison of results with IABP buoys**
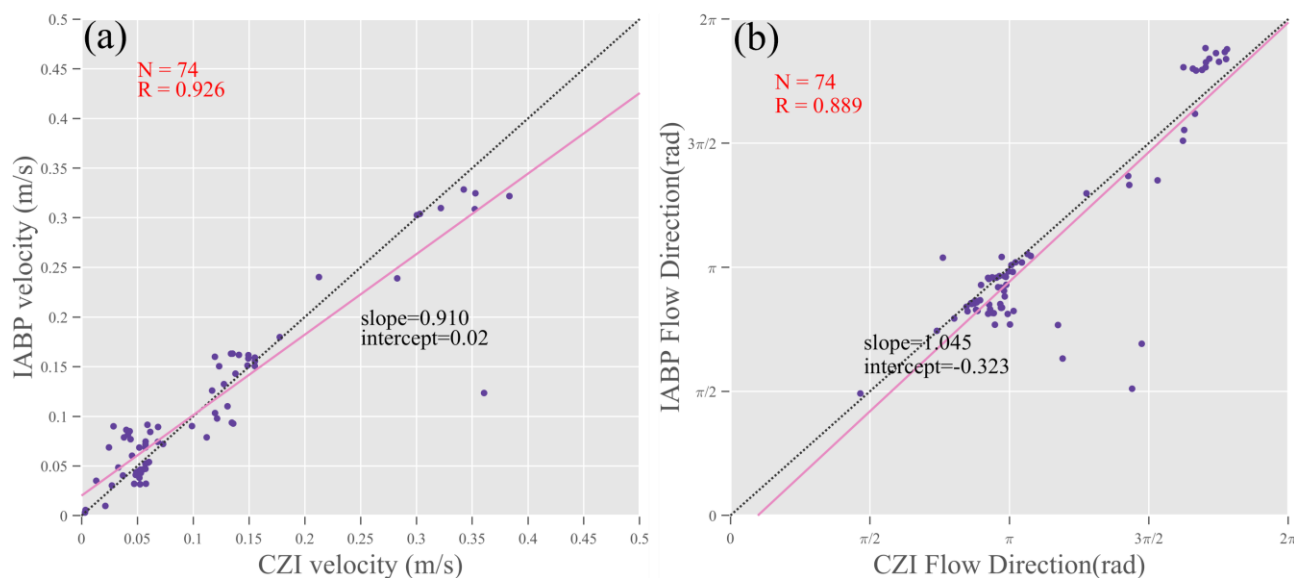
**4.2.1 SID with the day-level time interval**
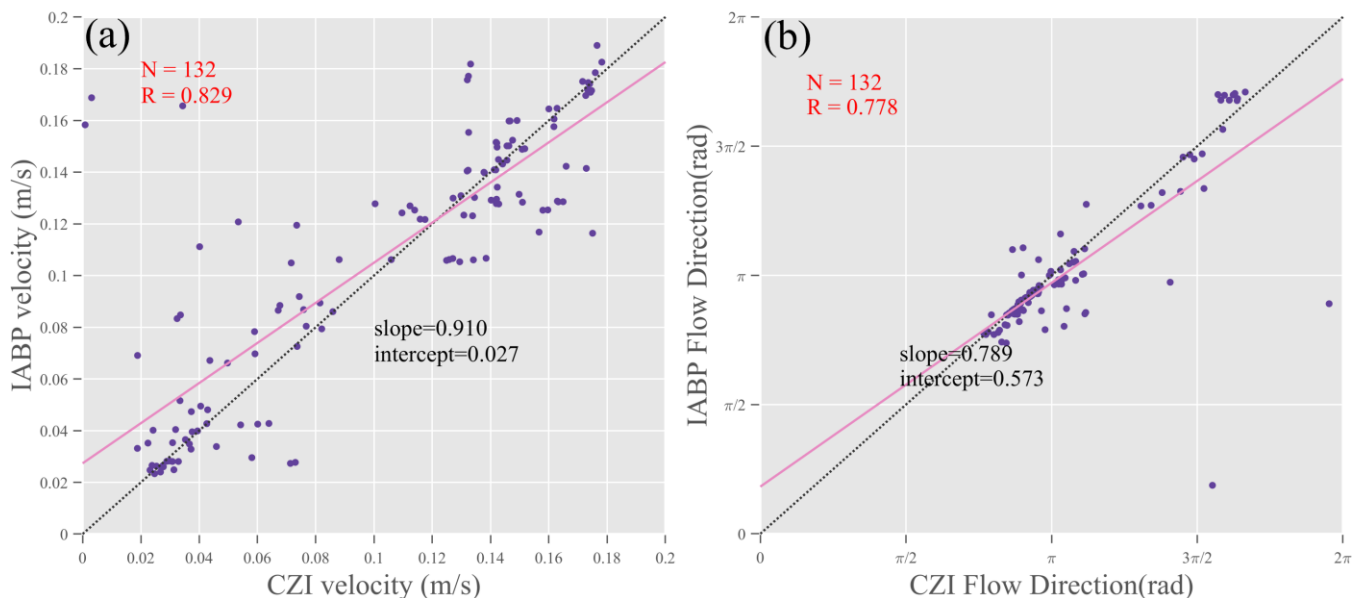


**Figure 10: Comparisons between the retrieved SID (day-level) and IABP buoys in velocity (a) and flow direction (b); the dashed line denotes the ideal line, and the pink line denotes the fit line.**

The GPS position of the buoys serves as a crucial reference for validating the accuracy of the retrieved SID. Fig. 10 illustrates the difference in velocity (a) and flow direction (b) between the result (day-level) and buoys, with a total of 132 matched points. The correlation coefficient between the retrieved SID and the buoy of velocity is approximately 0.829, while it is approximately 0.778 for the flow direction. With the time interval of images being approximately 24 hours (day-level), there is a strong correlation between the retrieved SID and the buoy.

Table 3 shows the statistics of the retrieved SID and the buoy, which indicate that the bias is approximately -0.005 m/s and the RMSE is approximately 0.031 m/s for velocity; moreover, for the flow direction, the bias is 0.002 rad and the RMSE is 0.009 rad.

**Table 3: Validation of the retrieved SID (day-level) with IABP buoys.**

| Delta | Number of match point | BIAS (Relative deviation$_{mean}$) | MAE | STD | RMSE |
|---|---|---|---|---|---|
| Velocity (m/s) | | -0.005(-1.330%) | 0.018 | 0.031 | 0.031 |
| Flow direction (rad) | 132 | 0.002(0.149%) | 0.003 | 0.009 | 0.009 |

### 4.2.2 SID with the hours-level time interval



**Figure 11: Comparisons between the retrieved SID (hours-level) and IABP buoys in velocity (a) and flow direction (b); the dashed line denotes the ideal line, and the pink line denotes the fit line.**

Fig. 11 shows the difference in velocity (a) and flow direction (b) between the result (hours-level) and buoys, with a total of 74 matched points. The correlation coefficient of the velocity between the retrieved SID and the buoy is 0.926, and the correlation coefficient of the flow direction between the retrieved SID and the buoy is 0.889. When the time interval of the images is less than 6 hours (hours-level), there is a satisfactory correlation between the retrieved SID and the buoy.

Table 4 provides the statistics of the retrieved SID and the buoy, revealing negligible bias values (~0 m/s) along with a RMSE of 0 036 m/s for velocity. The bias is 0.003 rad and the RMSE is 0.010 rad for flow direction. [Wang et al.] evaluated six SID products over the East Greenland with buoys and found that the MAEs of velocity for these products are ranged from 0.016-0.120 m/s and the MAEs of flow direction are ranged from 0.300-0.973 rad (Wang et al., 2022). Combined with Table 3 and

Table 4, it can be seen that the velocity and flow direction MAEs of our result are smaller than SID products based on radiometer, scatterometer and AVHRR.

**Table 4: Validation of the retrieved SID (hours-level) with IABP buoys.**

| Delta | Number of match point | BIAS (Relative deviation$_{mean}$) | MAE | STD | RMSE |
|---|---|---|---|---|---|
| Velocity (m/s) | | 0.000(1.998%) | 0.021 | 0.036 | 0.036 |
| Flow direction (rad) | 74 | 0.003(0.154%) | 0.006 | 0.010 | 0.010 |

## 5 Discussion

### 5.1 Analysis of the differences between the retrieved SID and the CMEMS SID product

#### 5.1.1 Comparison of individual retrieved SID and the product
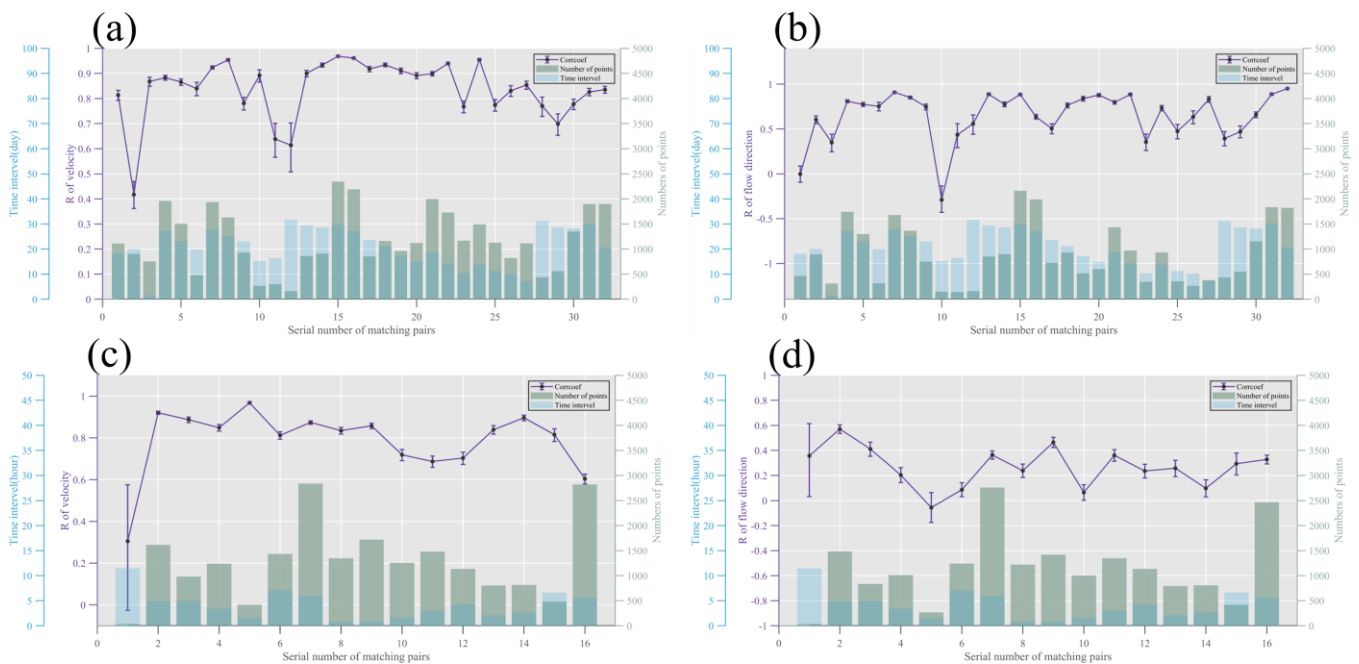


Figure 12: The correlation coefficient and 95% confidence level bar of correlation coefficient between the retrieved SID and the CMEMS SID product. (a), (b): The correlation coefficient between the retrieved SID (day-level) and the product in terms of velocity and flow direction. (c), (d): The correlation coefficient between retrieved SID (hours-level) and the product in terms of velocity and

**flow direction. The green histogram (right axis) shows the number of matched points between the retrieved SID and the product, and the blue histogram (the first left axis) shows the time interval between the retrieved SID and the product.**

The statistics of the retrieved SID and the CMEMS SID product are analyzed in section 4. In this section, we compare the

345   retrieved SID with the corresponding product in chronological order. Fig. 12(a) shows the velocity correlation coefficient curve between the day-level results and the product. Most of the correlation coefficients are greater than 0.7, the bars of 95% confidence level are idealistic. When the correlation coefficient is high, the bar of 95% confidence level is also behave relatively small. However, the $11^{th}$ and $12^{th}$ data pairs exhibited significantly lower correlation coefficients, primarily due to the involvement of fewer matched points (less than 500). Additionally, the correlation coefficient of the $2^{nd}$ data pair is much

350   lower than that of other data pairs due to the dense cloud in the CZI images, which affects the retrieval of SID. Fig. 12(b) shows the correlation coefficient curve of the flow direction between the day-level results and the product. The correlation coefficient of the $10^{th}$ pair is negative and considerably different from that of other data because there are fewer matched points (less than 500). Furthermore, the retrieved SID of $10^{th}$ is located on the northeast coast of Greenland. As the region where TPD encounters the Greenland coastline, the sea ice motion in the region is mutable. The sea ice motion trend in the area is highly

355   variable due to the influence of wind and other forces. The finding reveals that interactions between sea ice in the northern FS, where MYI congregates, can drastically affect the flow direction of the SID. Fig. 12(c) shows the velocity correlation coefficient curve between the hours-level results and the product. The significantly lower correlation coefficient of the $1^{st}$ pair of data can be attributed to the fact that there are not enough matched points (less than 200), and the confidence level bar of the $1^{st}$ pair is also unsatisfactory. Fig. 12(d) shows the correlation coefficient curve of the flow direction between the results

360   (hours-level) and the product. The correlation between the retrieved SID (hours-level) and the product in the flow direction is worse by comparing (b) and (d).

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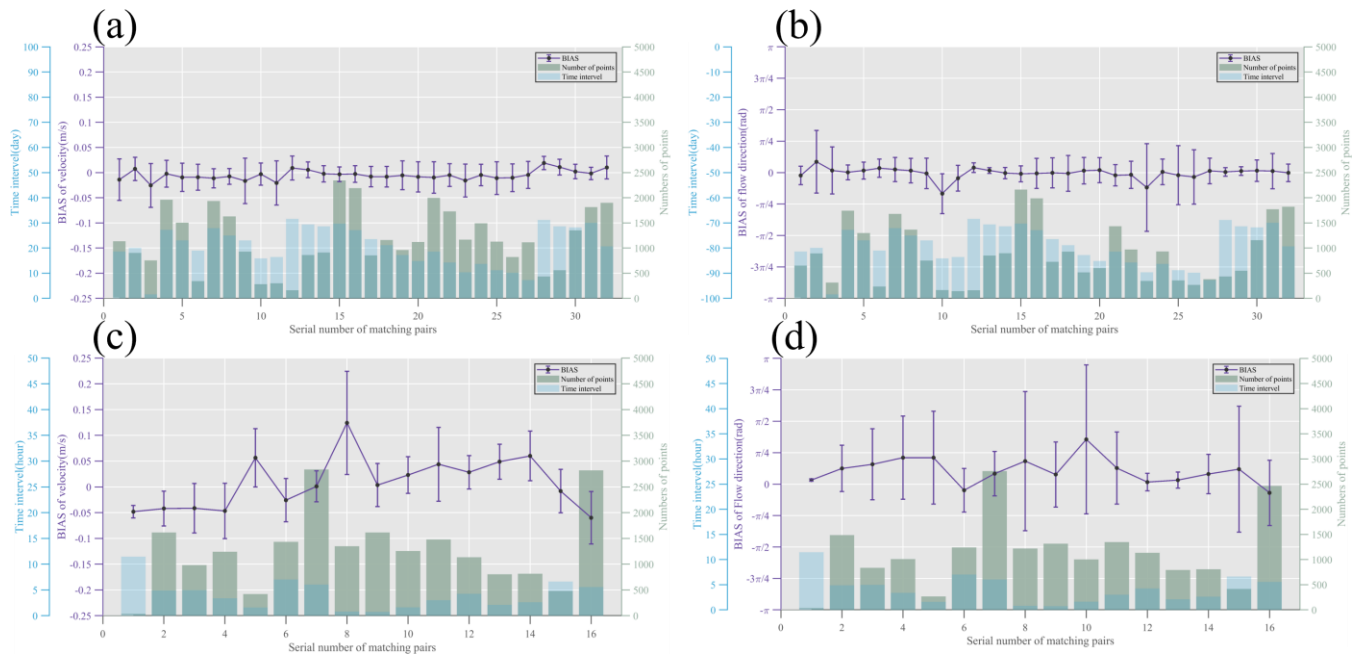
**Figure 13: The bias and standard deviation between the retrieved SID and the CMEMS SID product, (a), (b): The bias between the retrieved SID (day-level) and the product in terms of velocity and flow direction. (c), (d): The bias between retrieved SID (hours-level) and the product in terms of velocity and flow direction. The green histogram (right axis) shows the number of retrieved SID and the product matched points, and the blue histogram (the first left axis) shows the time interval between the retrieved SID and the product.**

The bias and standard deviation of the velocity between the day-level results and the CMEMS SID product are shown in Fig. 13(a). The velocity bias of the two SID datasets is less than ± 0.05 m/s, demonstrating a high degree of consistency. When the time interval is large (12th and 28th pairs) or there are fewer matched points (3rd, 10th, 11th and 13th pairs), the bias and standard deviation between our results and the products increase. Fig. 13(b) shows the bias and standard deviation of the flow direction between the day-level results and the product. The bias of the flow direction is slight. When there are fewer matched points (3rd, 10th and 23rd pairs), the bias and standard deviation of the flow direction become larger. Fig. 13(c) shows the velocity bias and standard deviation between the hours-level results and the product. Although most of the bias are smaller than 0.15 m/s, the velocity bias is significantly more substantial than that of the day-level result. The bias and standard deviation of the flow direction between the results (hours-level) and the product are shown in Fig. 13(d). By comparing (a), (b) and (c), (d) in Fig. 13, the day-level results are much less divergent from the product than the hours-level results.

### 5.1.2 Analysis of the difference

The day-level results are compared to the CMEMS SID product in Fig. 7 and 8, revealing the slight difference in U, V, velocity, and flow direction. The statistics in Table (2) also illustrate the slight difference between our results and the product. However, Fig. 12 shows a significant difference between the hours-level result and the product. The retrieval methods used in our study

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differed slightly from the method of the product, but both algorithms are based on the MCC. This discrepancy may be attributed to either the time interval of SAR images used in the product or the time interval between the two kinds of SID.
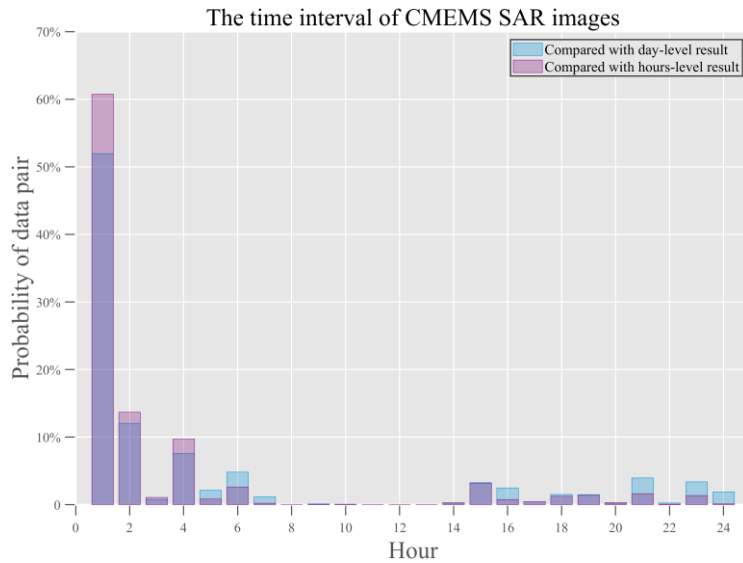


385 **Figure 14: The time interval of SAR images used by the CMEMS SID product during the comparison with two different results.**

Fig. 14 shows the time interval of SAR images used by the CMEMS SID product. The blue histogram represents the time interval of the SAR images used by the product, which is compared to the day-level result, and the purple histogram represents the time interval of the SAR images used by the product, which is compared to the hours-level result. Fig. 14 shows that the blue and purple histogram distributions are nearly identical. Most time intervals of SAR image used by the product are less

390 than 6 hours, and only a few are approximately 20 hours. The consistent distribution trend of the two kinds of histograms indicates that the time intervals of the SAR images used by the product are basically the same. The distribution of the histograms explains that the difference between the retrieved SID and the product is not relevant to the time interval of the SAR images used by the product.
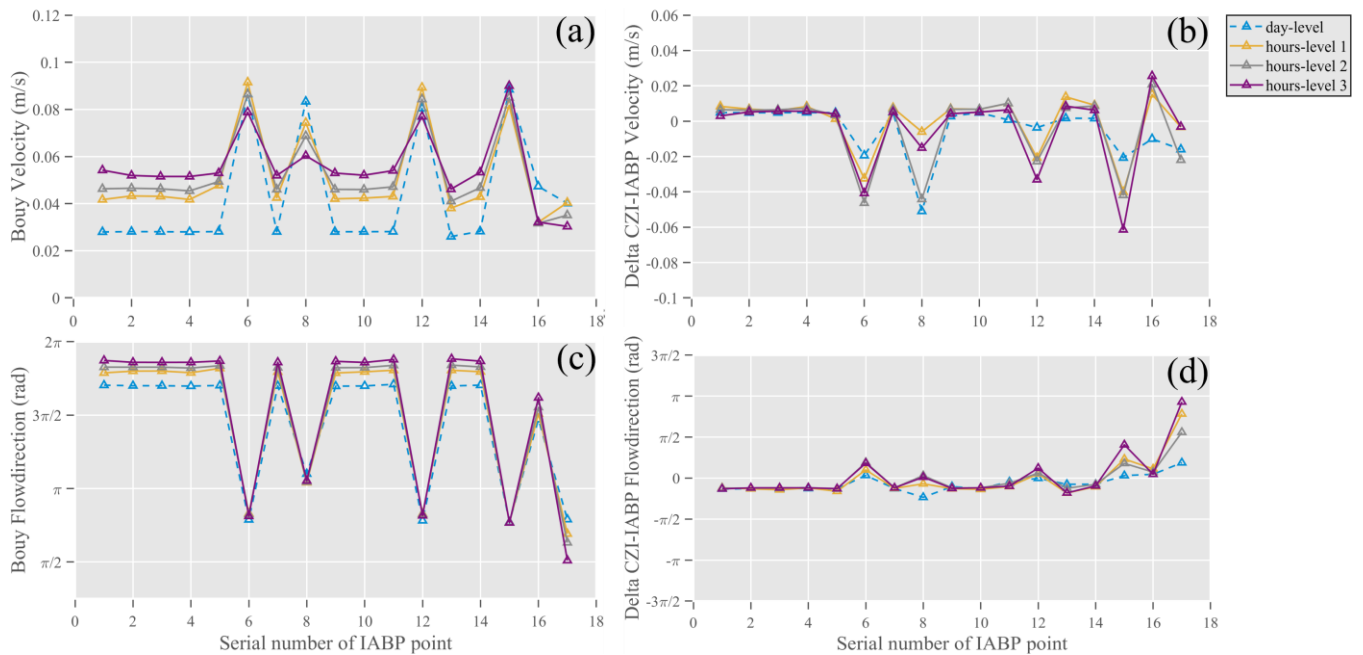
395 **Figure 15: The buoy velocity (a) and flow direction (c) and the difference (velocity (b), flow direction (d)) between the retrieved SID and buoys. The lines in blue represents SID and buoy vectors with day-level time interval and the lines in yellow, gray and purple represents SID and buoy vectors with hours-level time interval on the same day. The x-axis shows the different matched SID and buoys. The y-axis in (b) and (d) show the difference between SID and buoys.**

We attribute the difference between the hours-level result and the CMEMS SID product to the temporal and spatial variability

400 of sea ice motion. Fig. 15(a) and (c) show the variations in the velocity and flow direction of 17 different buoys with different time intervals. The date of day-level CZI images are 06 April 2021, 06:10:12 LT and 07 April 2021, 07:16:35 LT. The first date of the hours-level CZI images are 06 April 2021, 06:10:12 LT and 06 April 2021, 09:29:35 LT, the second date of the hours-level CZI images are 06 April 2021, 06:10:12 LT and 06 April 2021, 11:10:03 LT, and the third date of the hours-level CZI images are 06 April 2021, 09:29:35 LT and 06 April 2021, 11:10:03 LT. The triangles in Fig. 15(a) represent the velocity

405 of the buoys, and there is a significant difference in velocity between the day-level buoys and the hours-level buoys. The difference of velocity between hours-level buoys is less than the difference between hours-level buoys and day-level buoys, but the difference still exists. The same trend is also found in the difference of flow direction as shown in Fig. 15(c). Fig. 15(b) and (d) show the differences in velocity and flow direction between our results and the buoys, and the day-level result deviates less from the buoy. Additionally, different buoys show various differences from our results, which is related to the location of

410 the buoys.

The velocity and flow direction differences of the buoys with different time intervals suggests that SID has large spatial and temporal variability. There is a distinct variability in velocity and flow direction when the SID is retrieved using images with short time interval. When the drift vectors retrieved with longer time interval are compared with those vectors retrieved at shorter intervals, differences are noted. This phenomenon is a consequence of the spatial and temporal variability of SID. When

comparing the drift vectors retrieved with two different short time intervals, in extreme situations, a greater difference will be presented. The variability of sea ice motion is more complex and unpredictable on a small time scales. The complex spatial and temporal variability of SID explained the discrepancy between our results and the CMEMS SID product.

Despite the spatial and temporal variability of sea ice motion, it is still important to validate our retrieved results using the product. Furthermore, as long as a sufficiently large time interval is employed, the impact of SID variability diminishes, which results in consistent drift trend being observed. Utilizing SID products for evaluating the retrieved result remains valid while providing valuable references through consistency and difference ratings.

## 5.2 Effect of different time intervals on retrieving SID

In section 5.1, the consistency and discrepancy between the result and the CMEMS SID product is discussed, revealing the spatial and temporal variability of sea ice motion. In this section, we explore the impact of different time intervals on retrieving SID by analyzing the validation of the buoys and the statistics of the quality control parameters. The time interval of CZI images is a factor that affects the accuracy of retrieving SID. On the one hand, elaborate sea ice motion can be tracked by using images with short time interval; on the other hand, retrieving SID from images with longer time interval provides more stable result. Therefore, it is crucial to discuss the appropriate time interval for retrieving SID in FS.

Section 3.3 describes the three parameters used for quality control: R, PMR and PSR. Table 5 presents the statistics of these parameters with different time intervals. Each quality control parameter's mean value of the hours-level result is greater than that of the day-level result, which indicates that the hours-level result has the higher quality. However, the standard deviation of the quality control parameters exhibit opposite trend. The standard deviation of the quality control parameters for the day-level result is smaller than for the hours-level result. The statistics indicate that the day-level result remain relatively stable despite having slightly poorer quality.

**Table 5: Statistics of quality control parameters at different time intervals.**

| Parameter | MEAN | STD | Type |
|---|---|---|---|
| R | 0.372 | 0.199 | day-level |
| | 0.473 | 0.225 | hours-level |
| PMR | 9.338 | 4.857 | day-level |
| | 13.262 | 6.198 | hours-level |
| PSR | 1.620 | 0.838 | day-level |
| | 2.016 | 0.980 | hours-level |

In the validation with the buoy, the correlation of the day-level result is slightly lower than that of the hours-level result in velocity. In terms of flow direction, the correlation of the day-level result is also lower than that of the hours-level result. The retrieval of day-level SID confronts with the effects of large drift distances, wind change, light variation, cloud, etc., which are disadvantageous for retrieving. The statistics in Tables 3 and 4 also show that the hours-level result possesses better accuracy

than the day-level result. Images with short time interval provide a more stable scenes for retrieving SID, thus the hours-level result possesses fine accuracy.

In our study, the dataset was divided into two groups according to the image time interval. In comparison with the buoy, both results achieved great accuracy. According to the analysis of the day-level result and the hours-level result, we believe that retrieving SID using images with short time interval will provide SID with better accuracy. Our propositional time interval is shorter than the time interval proposed by Qiu, who suggested using images with 20-hours time interval to retrieve SID in FS from January to June (Qiu and Li, 2022). We attribute these divergent conclusions on suitable time interval due to the limitations of our dataset. In future studies, we are planning to incorporate HY1-C data and investigate the impact of varying time intervals on SID retrieval.

## 5.3 The potential of retrieving sea ice drift from image with cloud
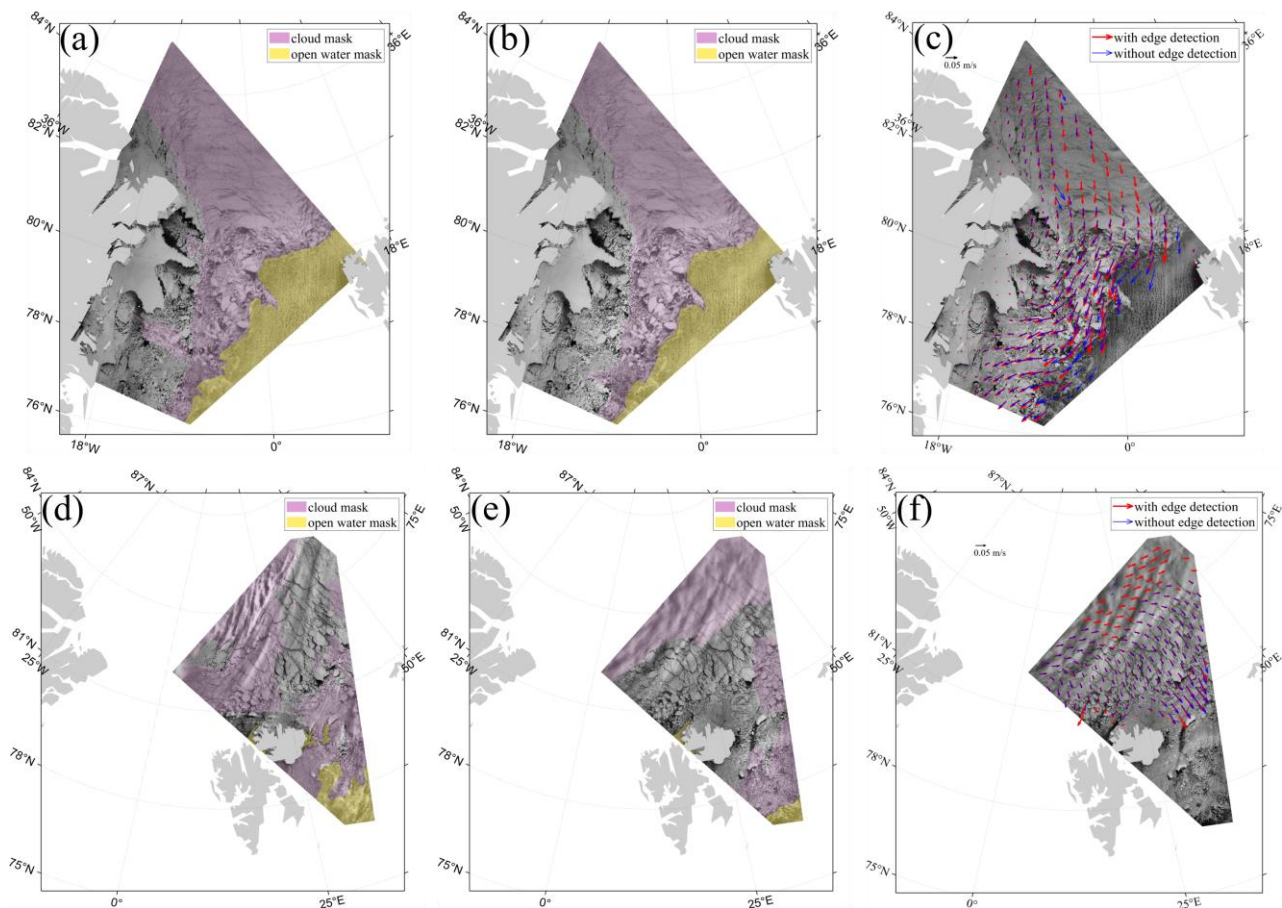


Figure 16: Two examples of retrieving sea ice drift from image with cloud. Images (a), (b), and (c) are a pair of results, and images (d), (e), and (f) are another pair of results. The arrows in (c) and (f) are diluted for direct viewing.

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Retrieving SID using optical imagery can be affected by factors such as cloud, lighting and other variables. Sobel edge detection is used to enhance the texture of sea ice. We found that SID result is also achieved for sea ice under cloud after using

455 Sobel edge detection. Fig. 16 shows two examples of SID retrieved from images with cloud. The first example uses hours-level time interval images for retrieval, with image (a) taken on 27 April 2021, 09:27:33 LT and image (b) taken on 27 April 2021, 11:07:20 LT. The second example uses day-level time interval images for retrieval, with image (c) taken on 02 May 2021, 03:20:35 LT and image (d) taken on 03 May 2021, 06:07:37 LT. In the second example, images (d) and (e) are also covered by thick cloud in addition to thin cloud, which more severely obscure the sea ice texture.

460 Fig. 16(a) and (b) show the cloud and open water by visual interpretation. Fig. 16(c) shows the SID vectors derived from the images, where the vectors with Sobel edge detection in the preprocessing are in red, and the vectors without Sobel edge detection are in blue. A comparison of these two types of vectors reveals a more consistent trend of SID with a homogeneous velocity distribution achieved by using Sobel edge detection, which demonstrates that edge detection promotes the retrieval algorithm and illustrates the potential for optical imagery to retrieve SID using images with thin cloud. Fig. 16(d), (e) and (f)

465 show another pair of result using CZI images with the day-level time interval. In contrast to the first pair of result, the CZI images (Fig. 16(d), (e)) have more cloud, and the variation in the cloud distribution is quite apparent. Comparing the vectors with two different colors, the red vector is better, even providing result in parts of the area under thick cloud. Thin cloud is distributed in lower right corner of the image, but the algorithm does not give credible result due to the strongly fragmented sea ice. Analytically, edge detection strengthens the ability of algorithm to retrieve SID from images with cloud. In addition,

470 Fig. 16 similarly shows the enhancement of the separating capacity for open water of algorithm by using detection. The red vectors in Fig. 16(c) and (f) are mainly distributed within the sea ice extent.
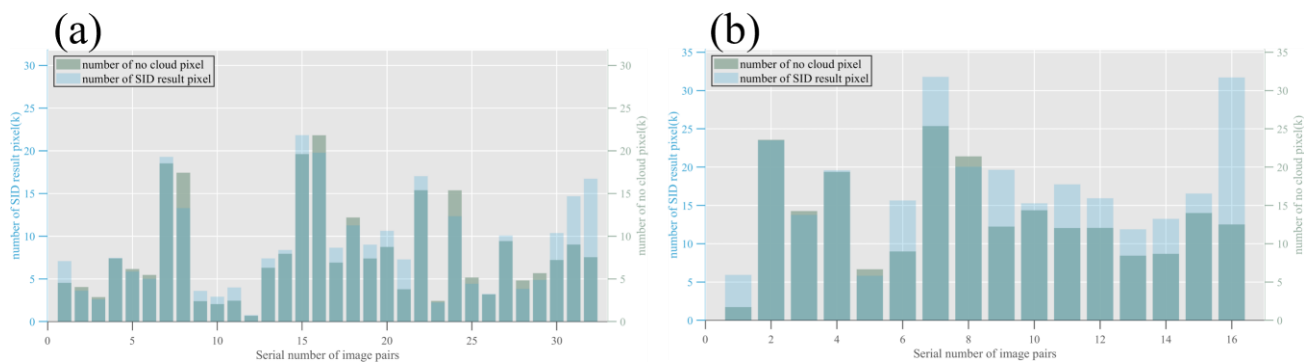


**Figure 17: The number of no-cloud pixels (green) and the retrieved SID pixels (blue) (a: day -level result, b: hours-level result).**

To quantify the potential of retrieving SID from image with cloud, we use the whiteness index (Gómez-Chova et al., 2008)

475 and Multiscale Edge-preserving Decompositions (MED) (Farbman et al., 2008) to detect the cloud pixels (Kang et al., 2018). Fig .17(a) and (b) show the number of no-cloud pixels and the number of retrieved SID pixels. For the day-level results, in addition to the influence of cloud pixels, dramatic shape changes of sea ice also hinder the retrieval of SID. As shown in Fig.

17(a), only a few ($1^{st}$, $15^{th}$, $21^{st}$, $30^{th}$, $31^{st}$ and $32^{nd}$) of the retrieved SID pixels are beyond the no-cloud pixels. For the $32^{nd}$ image pair, the number of SID pixels is 120% greater than the number of no-cloud pixels, possibly because the cloud in the images is quite thin. For the hours-level results, images with short time interval provide a stable scene for retrieving. Thus, many of the retrieved SID pixels ($1^{st}$, $6^{th}$, $7^{th}$, $9^{th}$ and $11^{th}$-$16^{th}$) are greater than the no-cloud pixels as shown in Fig. 17(b). For the $16^{th}$ image pair, the number of retrieved SID pixels is 150% greater than the number of no-cloud pixels, this is because the cloud is slight and the distributions of cloud in two images are similar. The average of SID pixels beyond the no-cloud pixels is 10.109% and 28.920% for the day-level result and the hours-level result, respectively. In conclusion, our approach has the potential to retrieve SID effectively with images containing cloud pixels. The result shows that with less strict restriction and appropriate preprocessing, images with cloud could be used to obtain credible and dense SID.

## 5.4 Factors affecting the accuracy of flow direction retrieval
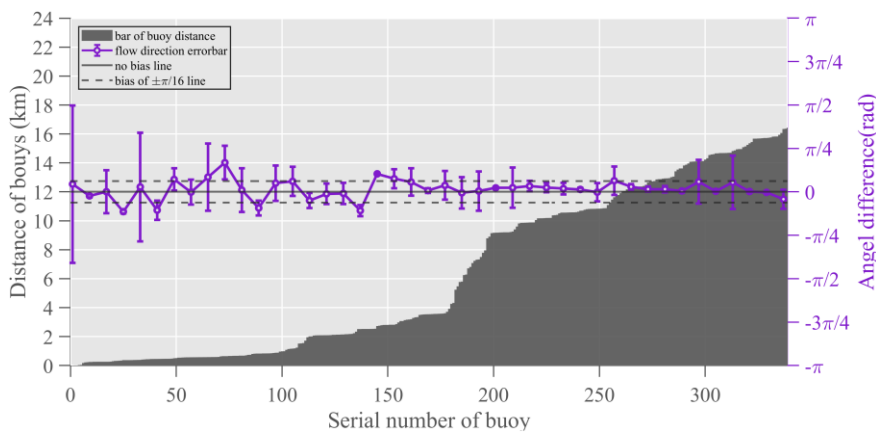


Figure 18: Relationship between the flow direction and the displacement of the buoys.

Qiu concluded that the flow direction accuracy of the retrieved SID is associated with the magnitude of the velocity (Qiu and Li, 2022). In our study, when dividing the dataset into hours-level and day-level by time interval for comparison with the buoys, inconsistent conclusion was found when analyzing the accuracy of flow direction. Compared with the buoy data, the statistics in flow direction of the SID at day-level is slightly smaller than the SID at hours-level. However, since the coverage areas of both datasets are repetitive, the magnitudes of velocity retrieved from both datasets should be the same. Thus, it is not the velocity magnitude of the SID that affects the accuracy of the flow direction retrieval.

We consider that the distance of sea ice motion affects the accuracy of the flow direction retrieval. Due to stochastic error, the identification of the highest correlation point deviates to the optimal point. Consequently, the retrieved flow direction is constrained to the vicinity of the optimal point. The variability of flow direction retrieval with longer distance is more stable than with shorter distance. Retrieving the flow direction precisely with small displacement is challenging, and the time interval needs to be increased to achieve longer sea-ice displacements, although this can also lead to other disadvantages (e.g., cloud, ice disintegration) for SID retrieval. Fig. 18 shows the relationship between the drift distance of the buoys and the difference

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between the retrieved SID and the buoy in the flow direction. We generated a total of 206 matched points and merged every 5 points to calculate the standard deviation. As the buoy displacement increases, the bias and standard deviation of the flow direction decrease, which indicates that the accuracy of the retrieved flow direction is affected by the distance of sea ice motion.

505     A larger displacements result in more precise flow direction retrieval but also require longer time interval images which will introduces disadvantageous factors for algorithm.

## 5.5 Factors affecting the accuracy of velocity retrieval
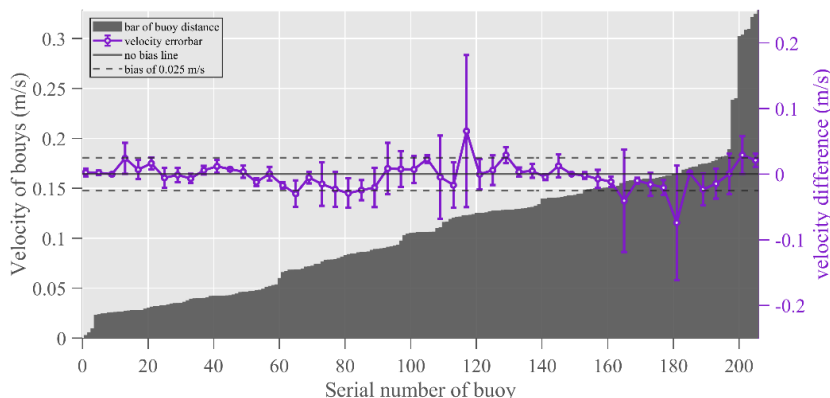


**Figure 19: Relationship between the velocity difference and the velocity of the buoys.**

510     The study area spans the Arctic basin to the Fram Strait, where the velocity of the SID generally presents complicated variations. The maximum magnitude of buoy-derived ice velocity can reach to 0.64 m/s in FS (Lei et al., 2016). Sea ice in the north of the FS or close to the land has a relatively low velocity. It is imperative to make further research on the accuracy of retrieving SID with different velocity.

In our study, the data were from March to May 2021, and we generated a total of 206 matched points and merged every 4

515     points to calculate the standard deviation of the velocity difference. We further analyzed the relation between the accuracy of retrieved velocity and the buoy velocity. As shown in Fig 19, the bar shows that the velocity of buoys that we used in validation ranged from 0 m/s to 0.32 m/s, and the line of bias and standard deviation shows the accuracy of the retrieved velocity which is associated with the velocity of the buoys. As the velocity of the buoys increases to 0.1 m/s, both the bias and standard deviation also increase. Although the bias of retrieved velocity occurs within in an ideal range as the velocity of the buoys

520     increases, the standard deviation of the velocity difference becomes erratic. In general, sea ice motion with great velocity appears unstable in the images. Our algorithm still need improvement in retrieving SID with high velocity.
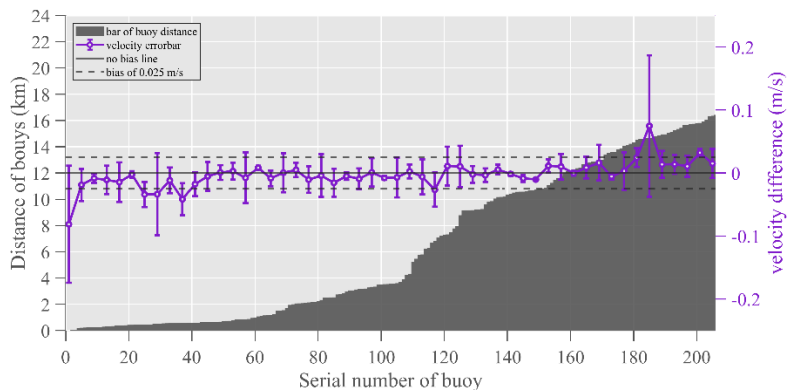
Figure 20: Relationship between the velocity difference and the displacement of the buoys.

Additionally, we delve into the correlation between the distance of sea ice motion and the accuracy of the retrieved velocity.
A significant challenge to the accuracy of velocity retrieval, as emphasized by Lavergne, is the presence of quantification noise
in template matching (Lavergne et al., 2010). Despite the acknowledged presence of quantification noise, our result surprisingly
does not reveal its discernible effects, and a lack of a statistically significant correlation between the accuracy of the retrieved
velocity and buoy distance is observed (As shown in Fig. 20).

Several explanations underpin this result. First, the spatial resolution of the resampled CZI imagery is 300m, which is a notable
improvement for the ice surface feature observation compared to the kilometer-level resolution of radiometer. The higher
spatial resolution is posited to be a key factor limiting the manifestation of quantification noise. The topographic features of
sea ice are often quantified by the surface roughness and form drag (Arya, 1973, 1975) and those observed surface features
have impact on ice drift speed (Zu et al., 2021). The form drag can be decomposed into the contributions from ridges, snow
dunes etc., and it is associated with the momentum exchange between sea ice and the lower atmosphere forced by wind (Zhang
et al., 2024). The topographic features (including height and spacing) within the optical imagery need to be described
quantitatively in the future work. Second, the strong sea ice motion in FS renders sea ice dynamics distinctly visible in the
image, indirectly suppressing the appearance of quantification noise. Finally, our application of subpixel estimation for the
precise localization of maximum correlation values, results in an exact determination of the maximum correlation location.
This refinement contributes to improve the accuracy of retrieving SID. In conclusion, high spatial resolution data and subpixel
estimation are able to suppress the negative effects caused by quantification noise.

## 5.6 The parameters for quality control

To validate the effectiveness of quality control parameters, a comprehensive examination was undertaken, involving a
comparative analysis between the quality control parameters and buoy validation of the SID. The red line in Fig. 21 shows the
R of the validation points, and to identify the relationships between R and the other parameters, the data are sorted in ascending
order by R. The bar in orange and bar in blue represent the PMR and PSR, respectively. The figure revealed a pronounced
congruence, wherein the R, PMR and PSR exhibited coherent consistency. A positive correlation was discerned, indicating

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that the increase in R proportionally accompanies the increase in the PMR and PSR. The magnitude of R has been extensively used as a quality control parameter for SID retrieval (Qiu and Li, 2022; Haarpaintner, 2006; Robert Ezraty et al., 2007). Given the observed synchronous trends of the PMR, PSR and R, it is evident that the PMR and PSR collectively demonstrate efficacy in quality control. In Fig. 21, the error-bar in purple shows the difference between the retrieved SID and buoys. The bias and standard deviation of the retrieved SID diminish as R, the PMR and the PSR increase, which indicates the effectiveness of these parameters for revealing the reliability of the result.
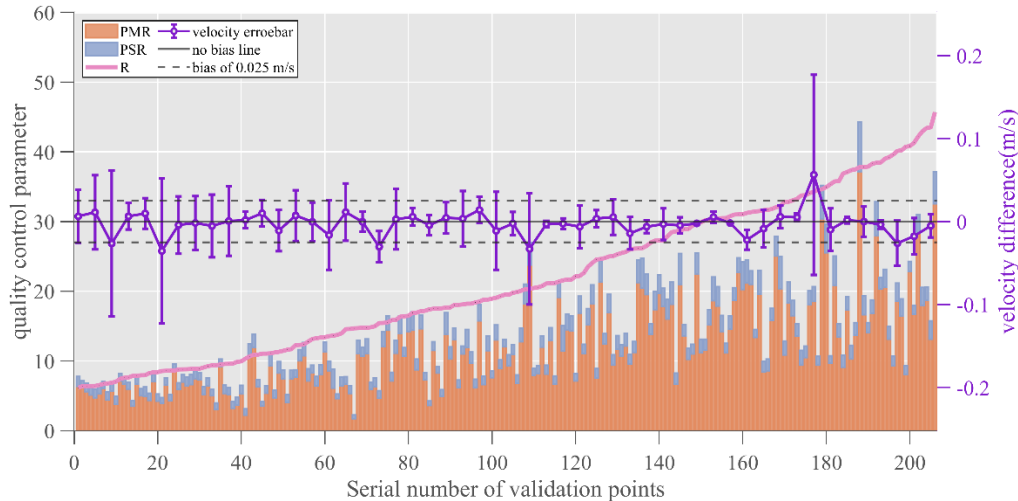


**Figure 21: Relationship between the velocity difference and quality control parameters.**

The coherence between the quality control parameters and the validation of the buoys is illustrated in Fig. 21, providing a foundation for the meticulous examination of the SID quality. As one of the main outlets of Arctic sea ice in the Atlantic Ocean, the sea ice concentration in the FS is inferior than the Arctic Basin (Peng and Meier, 2018; Wang et al., 2020), and the sea ice in the south and north of the strait render different morphologies in the images.

To investigate the quality of SID over the strait, we selected 80°N as the segmentation line and calculated the quality control parameters of SID in north and south of the segmentation line. The Fig. 22 shows the stacking bars of different parameters in different regions. For the sake of visualization, we normalize the PMR and PSR. The average values of these quality control parameters indicate that SID in the north of 80°N have higher quality. Comparison of quality control parameters for results with different time intervals reveals that the hours-level result possesses higher quality. Besides, the hours-level result also has greater mean value of quality control parameters in the south of 80°N where the sea ice is dispersive. High-concentration sea ice constrains the variability of sea ice motion and provides better spatial consistency, which is beneficial for retrieving SID. Similarly, the variability of the sea ice motion with short time interval is inconspicuous, providing a favorable scene for the algorithm. The sea ice kinematics in marginal ice zone are intricate, and the effects of wave on the fragmentary sea ice motion are evident (Williams et al., 2013, p.1), which increases the uncertainty of retrieval. The utilization of short time interval images for SID retrieval proves instrumental in enhancing the quality of monitoring sea ice motion within marginal ice zones.
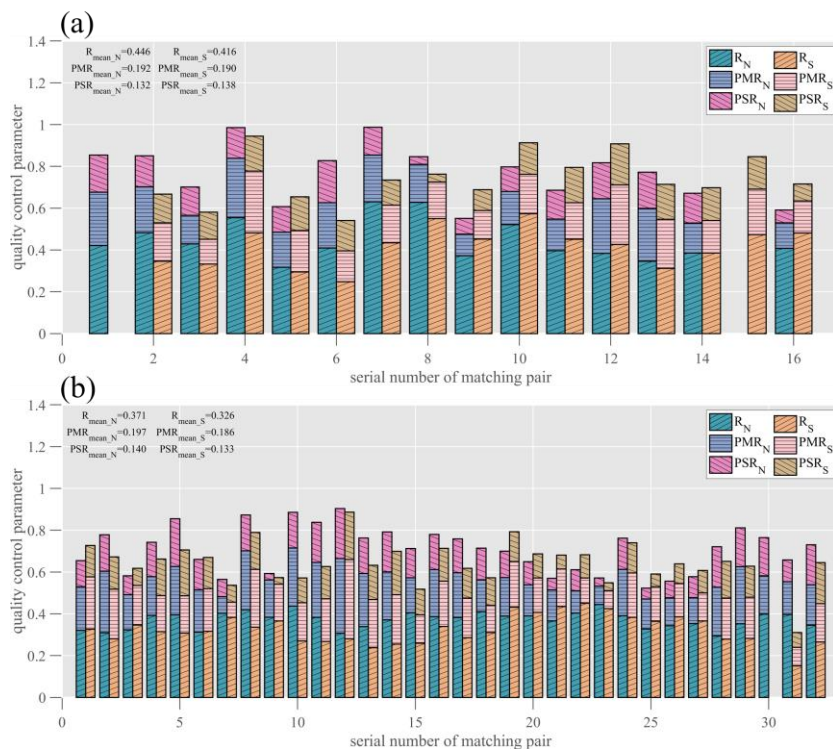
**Figure 22: Statistics of quality control parameters for individual retrieval SID. (a: hours-level, b: day-level). The suffix N and S mean north and south of 80°N in FS.**

## 6 Conclusion

The retrieval of SID in FS is complicated due to the geographic specificity. Sea ice drift within the Arctic basin is dominated by two trends, the BG and TPD. The FS is the outlet of TPD as well as the intersection of the Atlantic and Arctic. The heavily fragmented sea ice in the strait, coupled with its fast velocity, significantly impacts the retrieval of SID. Our method, utilizing the multi-template matching and subpixel estimation approach for retrieving SID in FS, demonstrates great accuracy. We use 111 scenes of CZI images and obtain 48 sets of image pairs for analysis. The dataset is categorized into hours-level and day-level according to the time interval. We found that the day-level result are consistent with the CMEMS SID product, the hours-level result is slightly different from the CMEMS SID product due to the spatial and temporal variability of the sea ice motion. Validation is also conducted using IABP buoys. For the velocity, the bias of the day-level result and hours-level result with the buoys are -0.005 m/s and 0.000 m/s, respectively, and the RMSE of the day-level result and hours-level result with the buoys are 0.031 m/s and 0.036 m/s, respectively. For the flow direction, the bias of the day-level result and hours-level result with the buoys are 0.002 rad and 0.003 rad, respectively, and the RMSE of the day-level result and hours-level result with the buoys are 0.009 rad and 0.010 rad, respectively. The accuracy of the retrieved SID surpasses that of the SID products retrieved from AVHRR.

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The method with multi-template matching and subpixel estimation for SID retrieval used in our study is an improvement of the traditional MCC algorithm. In our study, the reliability and accuracy of the algorithm are verified via comparison with the buoy. Additionally, the multi-template matching yields satisfactorily in terms of computational speed. Thus, we consider that

590   our method for SID retrieval balances the computational cost and accuracy.

Our study verified the feasibility and validity of using CZI images for SID retrieval in FS. The spatial resolution of the retrieved SID is 4 km, which is much higher than the SID products based on scatterometers, radiometers and AVHRR. For our result, the spatial resolution is suitable for capturing complex sea ice motion in the FS. The SID products from scattermeters and radiometers, which possess broader coverage, are more suitable for the study of sea ice mass balance and assimilation of

595   numerical prediction over the whole Arctic region. Our result also illustrates the high potential of the CZI for exploring the polar. At present, the HY-1C and HY-1D satellites are networked for observation, and the Chinese HY-1 series of satellites will be further utilized to advance polar research. Since 23 Dec 2021, the discontinuation of the Sentinel-1B satellite operation resulted in numerous data gaps in the CMEMS SID product which severely limited the long-term sea ice motion study in FS. Through leveraging expansions of HY1-C data in future, we anticipate that SID retrieved from CZI will serve as the

600   complement for the CMEMS SID product and facilitate the utilization of high spatial and temporal resolution product for sea ice dynamics research.

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