



## Bridging the spatial gaps of the Ammonia Monitoring Network using satellite ammonia measurements

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**Abstract.** Ammonia (NH<sub>3</sub>) is a key precursor to fine particulate matter (PM<sub>2.5</sub>) and a primary form of reactive nitrogen. The limited observations of NH<sub>3</sub> hinders further understanding of its impacts on air quality, climate, and biodiversity. Currently, NH<sub>3</sub> ground monitoring networks are limited in number across the globe, and even in the most established networks, large spatial gaps exist between sites and only a few sites have records that span longer than a decade. Satellite NH<sub>3</sub> observations can be used to discern trends and fill spatial gaps in networks, but many factors influence the syntheses of the vastly different spatiotemporal scales between surface network and satellite measurements. To this end, we intercompared surface NH<sub>3</sub> data from the Ammonia Monitoring Network (AMoN) and satellite NH<sub>3</sub> total columns from the Infrared Atmospheric Sounding Interferometer (IASI) in the contiguous United States (CONUS) and then performed trend analyses using both datasets. We explored the sensitivity of correlations between the two datasets to factors such as satellite data availability and distribution over the surface measurement period as well as agreement within selected spatial and temporal windows. Given the short lifetime of atmospheric ammonia and consequently sharp gradients, smaller spatial windows show better agreement than larger ones except in areas of relatively uniform, low concentrations where large windows and more satellite measurements improve the signal-to-noise ratio. A critical factor in the comparison is having satellite measurements across most of the measurement period of the monitoring site. When IASI data are available for at least 80% days of AMoN's 2-week sampling period within a 25 km spatial window of a given site, IASI NH<sub>3</sub> column concentrations and the AMoN NH<sub>3</sub> surface concentrations have a correlation of 0.74, demonstrating the feasibility of using satellite NH<sub>3</sub> columns to bridge the spatial gaps existing in the surface network NH<sub>3</sub> concentrations. Both IASI and AMoN show increasing NH<sub>3</sub> concentrations across CONUS (median: 6.8%·yr<sup>-1</sup> vs. 6.7%·yr<sup>-1</sup>) in the last decade (2008 - 2018), stressing the rising importance of NH<sub>3</sub> in terms of nitrogen deposition. NH<sub>3</sub>



35 trends for AMoN sites correlates with IASI NH<sub>3</sub> trend IASI and AMoN NH<sub>3</sub> trend ( $r = 0.66$ ) and show a similar spatial pattern,  
with the highest increases in the Midwest and eastern U.S., and NH<sub>3</sub> trend for AMoN sites correlates with IASI NH<sub>3</sub> trend ( $r$   
= 0.66). In spring and summer, increases of NH<sub>3</sub> were larger than 10%·yr<sup>-1</sup> in the eastern U.S. and Midwest (cropland  
dominated) and western U.S. (pastureland dominated), respectively. In terms of trend in NH<sub>3</sub> hotspots (defined as regions where  
the IASI NH<sub>3</sub> column is larger than the 95<sup>th</sup> percentile of 11-year CONUS map,  $6.7 \times 10^{15}$  molec/cm<sup>2</sup>), these largest emissions  
40 sources are also experiencing increasing concentrations over time with the median of NH<sub>3</sub> trend is 4.7% · yr<sup>-1</sup>. IASI data show  
large NH<sub>3</sub> increases in urban areas (8.1%·yr<sup>-1</sup>), including 8 of the top 10 most populous regions in the CONUS, where AMoN  
sites are sparse. The increasing NH<sub>3</sub> could have detrimental effects on nearby eco-sensitive regions through nitrogen deposition  
and on aerosol chemistry in the densely populated urban areas, hence needs immediate attention.

## 45 **1 Introduction**

Gas phase ammonia (NH<sub>3</sub>) is the most abundant alkaline gas in the atmosphere, mainly emitted from agricultural activities  
such as nitrogen fertilizer applications and livestock waste volatilization (Bouwman et al., 1997; Paulot et al., 2014). As a  
major precursor to fine particulate matter (PM<sub>2.5</sub>), NH<sub>3</sub> critically affects aerosol heterogeneous chemistry, air quality, visibility,  
human health, and climate (Hauglustaine et al., 2014; Hill et al., 2019; Lawal et al., 2018; Malm et al., 2004). Ammonia  
50 neutralizes sulfuric acid (H<sub>2</sub>SO<sub>4</sub>) and nitric acid (HNO<sub>3</sub>) in the atmosphere to form ammoniated aerosols, ammonium sulfate  
(NH<sub>4</sub>)<sub>2</sub>SO<sub>4</sub>) and ammonium nitrate (NH<sub>4</sub>NO<sub>3</sub>), which in total can contribute to more than 50 % of total PM<sub>2.5</sub> mass (Feng et  
al., 2020). NH<sub>4</sub>NO<sub>3</sub> is critical during wintertime haze periods because the cold and humid condition favor its formation (Shah  
et al., 2018; Zhai et al., 2021). Besides, NH<sub>3</sub> plays an important role in the nitrogen cycle. Wet deposition of NH<sub>4</sub><sup>+</sup> dominates  
the wet inorganic nitrogen deposition at nearly 70% of monitoring sites in the United States (Li et al., 2016). Total NH<sub>x</sub>  
55 (≡NH<sub>3</sub>(g) + NH<sub>4</sub><sup>+</sup>(aq)) deposition is expected to become even more dominant in the future because NO<sub>x</sub> emissions decrease  
under pollution control while NH<sub>3</sub> emissions are predicted to continue to increase with the rising global food demands (Erisman  
et al., 2008; Goldberg et al., 2021; Pinder et al., 2008). Excessive NH<sub>3</sub> deposition in the non-agricultural ecosystems can reduce  
biodiversity, result in soil acidification, and increase eutrophication, especially in the sensitive ecosystems (Ellis et al., 2013;  
Phoenix et al., 2006).

60 Although NH<sub>3</sub>'s importance has been well recognized, routine NH<sub>3</sub> observations are lacking even in countries with  
comprehensive monitoring networks, partly due to the difficulty of measuring gas phase NH<sub>3</sub> (von Bobruzki et al., 2010;  
Fehsenfeld et al., 2002). The Ammonia Monitoring Network (AMoN) (Puchalski et al., 2015) is the only routine set of NH<sub>3</sub>  
measurements in the United States, with 110 active AMoN sites in the contiguous United States (CONUS) in 2021, providing  
65 high-quality surface observations of NH<sub>3</sub>. AMoN data have been used widely for model evaluation and long-term trend  
analysis (Butler et al., 2016; Nair et al., 2019; Yao and Zhang, 2016, 2019). AMoN only provides bi-weekly NH<sub>3</sub> observations,



in contrast to monitoring networks for two other important gas phase precursors of  $PM_{2.5}$ ,  $SO_2$  and  $NO_2$ , which provide hourly or daily scale observations.  $PM_{2.5}$ ,  $SO_2$ , and  $NO_2$  are directly regulated as criteria pollutants, however contributions from  $NH_3$  emissions sources must be considered in State Implementation Plan (SIP) demonstrations for areas out of attainment for  $PM_{2.5}$ , which can be a challenge for areas lacking  $NH_3$  measurements (EPA 2023).

Population weighted  $PM_{2.5}$  concentrations are widely used to estimate the health effects of  $PM_{2.5}$ , however, the sparse number of  $NH_3$  sites with only biweekly or monthly resolution makes it difficult to derive population weighted  $PM_{2.5}$  precursor datasets. Gas phase  $NH_3$  is critical to determine the partitioning of the total  $NH_x$  (Hennigan et al., 2015), and the lack of gas phase  $NH_3$  observations hampers the evaluation of chemistry models. The ISORROPIA-II thermodynamic model has been extensively adopted to compute the equilibrium composition for the inorganic aerosol systems (Fountoukis and Nenes, 2007) and requires both gas and aerosol phase data as input to provide accurate and robust results (Hennigan et al., 2015). However, the limited number of  $NH_3$  ground monitoring sites currently prevents synthesizing the AMoN  $NH_3$  data with other ground monitoring networks, e.g., IMPROVE, as input for ISORROPIAII (Pan et al., 2020). GEOS-Chem implemented with ISORROPIA-II was found to significantly underestimate gas phase  $NH_3$  and overestimate  $NH_4^+$  in winter (Holt et al., 2015; Nair et al., 2019; Walker et al., 2012), with the normalized  $NH_4^+$  mean biases as high as 86% in January at sites for the Interagency Monitoring of Protected Visual Environments (IMPROVE) (Holt et al., 2015). The lifetime of  $NH_3$  ranges from hours to days, hence large spatiotemporal variability exists (Golston et al., 2020; Miller et al., 2015; Wang et al., 2021), and large spatial gaps exist in the current AMoN. Currently there are no AMoN sites in some states, e.g., North Dakota and South Dakota, and only 12 sites are within the characteristic length scale (12 km) of  $NH_3$  hotspots regions (Wang et al., 2021). Ten national parks in the U.S. are within 100 km of an  $NH_3$  hotspot, and more observations are needed to quantify the impacts of these hotspots on dry  $NH_3$  deposition in these regions (Pan et al., 2021). A lack of long-term AMoN data also hinders the possibility of investigating  $NH_3$  trends in the CONUS. Increasing  $NH_3$  concentrations are observed using AMoN data, yet all of the previous trend analyses are limited to fewer than 20 AMoN sites that may not be representative of  $NH_3$  trends in the CONUS (Butler et al., 2016; Yao and Zhang, 2016, 2019).

Satellite  $NH_3$  observations are on a global and daily basis, providing long-term trends and ubiquitous coverage. Instruments that measure  $NH_3$  include the Infrared Atmospheric Sounding Interferometer (IASI) on the MetOp satellites, Cross-track Infrared Sounder (CrIS) on NOAA and NASA Suomi National Polar-orbiting Partnership (S-NPP), Tropospheric Emission Spectrometer (TES) on NASA Aura satellite, Atmospheric Infrared Sounder (AIRS) on NASA EOS Aqua satellite, and Thermal and Near Infrared Sensor for Carbon Observations – Fourier Transform Spectrometer (TANSO-FTS) on the Greenhouse Gases Observing SATellite (GOSAT) (Clarisse et al., 2009; Shephard et al., 2011; Shephard & Cady-Pereira, 2015; Someya et al., 2020; Warner et al., 2016). Satellite  $NH_3$  data have been widely used to constrain  $NH_3$  emissions, estimate  $NH_3$  deposition, and analyze  $NH_3$  trends (Cao et al., 2020; Chen et al., 2020; Kharol et al., 2018; Van Damme et al., 2021). Van Damme et al. (2021) utilized 11-year IASI  $NH_3$  observations and found a worldwide  $NH_3$  increase ( $12.8 \pm 1.3$  %) from



2008 to 2018 with especially large increases in east Asia ( $75.7 \pm 6.3$  %) and North America ( $26.8 \pm 4.5$  %). Warner et al. (2017) used 14-year AIRS  $\text{NH}_3$  measurements and found statistically significant  $\text{NH}_3$  increase ( $2.61\% \cdot \text{yr}^{-1}$ ) in the U.S. from 2002 to 2016.

105 The global daily coverage and long-term data record make it possible for satellite observations to fill the spatial and temporal gaps of the current ground monitoring networks. Although limited in numbers, the validations of satellite  $\text{NH}_3$  observations with in-situ measurements provide confidence in integrating the two datasets (Guo et al., 2021; Sun et al., 2015). Sun et al. (2015) performed the first daily and pixel scale satellite  $\text{NH}_3$  validations using TES  $\text{NH}_3$  columns and airborne  $\text{NH}_3$  observations in the San Joaquin Valley of California, USA, showing that the differences between the total  $\text{NH}_3$  column and the in-situ total column were within 6 %. However, the validation included only 9 TES pixels, and TES is no longer in operation now. Guo et al. (2021) showed that IASI  $\text{NH}_3$  columns and  $\text{NH}_3$  columns derived from airborne and ground-based  $\text{NH}_3$  observations were indistinguishable from one another on daily and pixel bases in Colorado, USA, in summer. All of these validation works were performed in certain seasons and were limited to source regions with high  $\text{NH}_3$  concentrations (Guo et al., 2021; Sun et al., 2015; Warner et al., 2016). Ground-based FTIR  $\text{NH}_3$  observations provided a better temporal coverage for evaluating IASI and CrIS  $\text{NH}_3$  retrievals, however, low concentration sites were excluded from the evaluation and only ~ 10 sites were included across the globe (Dammers et al., 2016; Dammers et al., 2017). Furthermore, FTIR-based measurements also have not been directly validated against in-situ measurements of  $\text{NH}_3$  vertical profile themselves.

To capitalize on the benefits of both surface and satellite observations and synthesize these datasets, a detailed understanding of the comparison between IASI  $\text{NH}_3$  column concentrations and AMoN  $\text{NH}_3$  surface concentrations is necessary. Here we focus on IASI  $\text{NH}_3$  measurements because it offers the longest data record (2008 - present) among the satellite  $\text{NH}_3$ -measuring instruments. The comparison between AMoN and IASI is complex because AMoN is a ground-based, point measurement integrated over fourteen days, whereas IASI is a space-borne volumetric measurement averaged over the pixel footprint at the instantaneous overpass time. There are several factors that need to be taken into consideration:

125 (1) The extent to which the IASI  $\text{NH}_3$  column represents the surface AMoN  $\text{NH}_3$  concentration: Knowledge of  $\text{NH}_3$  vertical profiles in the atmosphere are limited due to the lack of observational data, and model simulated  $\text{NH}_3$  vertical profiles are often biased compared with the airborne measurements (Schiferl et al., 2016). Ammonia is mostly concentrated in the planetary boundary layer (PBL) because of its short lifetime (~hours to days) and surface emission sources (Dentener & Crutzen, 1994; Guo et al., 2021; Sun et al., 2015; Seinfeld & Pandis, 2016). Sun et al. (2015) showed that  $\text{NH}_3$  was almost well mixed in the lower PBL, and the TES  $\text{NH}_3$  columns were strongly correlated ( $R^2 = 0.82$ ) with the median  $\text{NH}_3$  mixing ratios measured at the surface, demonstrating that satellite  $\text{NH}_3$  columns could represent the ground  $\text{NH}_3$  concentrations. Van Damme et al. (2015) converted IASI  $\text{NH}_3$  columns to surface  $\text{NH}_3$  concentrations using fixed  $\text{NH}_3$  profiles generated by GEOS-Chem, then performed monthly comparisons with ground monitoring networks. IASI derived surface  $\text{NH}_3$  observations are in fair



135 agreement with ground observations in Europe, China, and Africa, but are limited to a small number of sites in each region for  
a short time range, e.g., 27 sites in Europe in 2011 (Van Damme et al., 2015). Furthermore, the latest IASI NH<sub>3</sub> products have  
switched to a new algorithm and no longer use a fixed NH<sub>3</sub> profile (Whitburn et al., 2016; Van Damme et al., 2017).

(2) Optimal spatial window for comparing and integrating satellite pixels and AMoN sites: Previous comparisons of satellite  
140 NH<sub>3</sub> retrievals with observations from ground monitoring networks simply averaged the data from the monitoring site within  
a coarse model grid (~ 100 km) with the averaged modeling/satellite NH<sub>3</sub> concentration of the whole grid (Kharol et al., 2018;  
Nair et al., 2019; Van Damme et al., 2015). If NH<sub>3</sub> concentrations are uniformly distributed within the spatial window,  
increasing the spatial window will increase the number of IASI pixels and decrease the signal-to-noise ratio. However, the  
spatial heterogeneity of NH<sub>3</sub> is quite large near hotspots due to its short lifetime (Golston et al., 2020; Miller et al., 2015; Wang  
145 et al., 2021; Warner et al., 2016). The relationship between spatial window size and satellite/surface measurements agreement  
needs to be examined in more details.

(3) Temporal distribution of satellite measurements across the two-week AMoN sampling period: Previous comparisons of  
model or satellite products against surface observations did not consider the distribution of IASI measurements during the two-  
150 week sampling period (Kharol et al., 2018; Nair et al., 2019; Van Damme et al., 2015). AMoN measures continuously, whereas  
a series of cloudy days would preclude any valid satellite measurements. Therefore, any AMoN/satellite comparison is  
intrinsically biased towards clear sky days on the satellite side but includes all conditions for the AMoN site.

(4) Number of available IASI pixels in the comparison: Guo et al. (2021) has shown that, even at low column amounts, IASI  
155 NH<sub>3</sub> has no known biases. AMoN is an extremely sensitive measurement of NH<sub>3</sub>, far more precise than any satellite NH<sub>3</sub>  
product (NADP, 2023; Van Damme et al., 2017). Therefore, increasing the number of satellite measurements within a certain  
spatiotemporal window is expected to improve the signal-to-noise ratio in the satellite measurements and may lead to improved  
agreements with AMoN under clean conditions.

160 (5) Regional and seasonal variabilities: Different regional and seasonal patterns are expected to influence the comparison. The  
performances of thermal infrared sounders are highly affected by the thermal contrast between the surface air temperature and  
skin temperature (Clarisse et al., 2010). In winter, low thermal contrast results in low sensitivity, which explains the low  
number of IASI pixels in winter compared to summer (Clarisse et al., 2010; Guo et al., 2021). Kharol et al. (2018) showed that  
CrIS surface NH<sub>3</sub> concentrations had an overall mean CrIS–AMoN difference of ~+15%, however, they only averaged CrIS  
165 data over the warm season in 2013.

In this study, to demonstrate the capabilities of using IASI NH<sub>3</sub> observations to augment the ground monitoring network, we  
performed a comprehensive comparison between IASI and AMoN on weekly/seasonal scales. We directly compare the



170 correlation between IASI  $\text{NH}_3$  columns with AMoN surface  $\text{NH}_3$ . We avoided direct comparisons when converting column  
175  $\text{NH}_3$  into surface concentrations because of possible biases introduced by assuming vertical profiles, boundary layer heights at  
local sites, and gas phase - aerosol partitioning. The impacts of the different factors on the comparison are examined in the  
context of points raised above. After identifying the most optimal method for comparison, we examined  $\text{NH}_3$  trends over  
AMoN sites and the larger applicability of using satellite retrievals to discern  $\text{NH}_3$  trends over regions and seasons lacking  
AMoN data.

175

## 2 Data and methods

### 2.1 Satellite $\text{NH}_3$ observations

IASI is an infrared sounder launched on board of the MetOp-A, MetOp-B, and MetOp-C platforms in sun-synchronous orbits  
since October 2006, September 2012, and November 2018, respectively. IASI has a swath of 2200 km and provides global  
180 coverage twice per day at around 09:30 and 21:30 mean local solar time. At nadir, the IASI footprint has a 12-km diameter.  
The first IASI  $\text{NH}_3$  product was developed by Clarisse et al. (2009) by converting the brightness temperature differences into  
total  $\text{NH}_3$  columns. Later on, a flexible and robust retrieval algorithm based on an artificial neural network for IASI (ANNI)  
(Whitburn et al., 2016) was developed. The latest version consists of a reanalyzed dataset provided with the European Centre  
for Medium-Range Weather Forecasts Re-Analysis v5 (ERA5) as its meteorological input (Van Damme et al., 2017; Van  
185 Damme et al., 2021). Because the meteorological input for reanalysis data is coherent in time, it is the more appropriate dataset  
to be used to study trends. For the present analyses, we used IASI version 3.1 reanalysis (v3.1r) retrieval product data from the  
MetOp/A (2008-2018) and MetOp/B (2013-2018) satellites (limited to cloud fraction  $\leq 25\%$ ). Only the morning orbits were  
analyzed because of higher sensitivity than the evening overpasses (Clarisse et al., 2010).

### 190 2.2 Ground-based observations

AMoN is the only network providing a consistent, long-term record of  $\text{NH}_3$  gas concentrations across the United States. AMoN  
was established by the National Atmospheric Deposition Program (NADP) in October 2007 and expanded to 19 sites in 2010  
and 105 sites in 2018. AMoN deploys Radiello® passive samplers that rely upon diffusion theory, where gas phase  $\text{NH}_3$  is  
adsorbed onto a cylindrical interior filter and extracted as  $\text{NH}_4^+$  to be analyzed by Flow Injection Analysis (FIA). AMoN  
195 provides biweekly surface  $\text{NH}_3$  concentrations, and the network detection limit is  $0.083 \text{ mg NH}_4^+ \text{ L}^{-1}$  ( $\sim 0.078 \mu\text{g NH}_3 \text{ m}^{-3}$ ) for  
the 2-week samples in 2020 (NADP, 2023). The Radiello passive samplers were found to be biased low by 37% against  
denuders used as reference method (Puchalski et al., 2011). In this study, we are comparing the relative variations instead of  
absolute concentrations of IASI and AMoN, therefore the low bias of AMoN measurements is not as relevant to the outcome.



200 We incorporated data from all AMoN sites with one notable exception. Using satellite imagery, we identified that the AMoN  
site in Logan, Utah (UT01), is located only  $\sim 100$  m away from a livestock farm. Ammonia concentrations downwind of a  
beef/dairy feedlot at this distance are far above background levels and unrepresentative of those at the local-regional scales (1-  
10 km) (Golston et al., 2020; Miller et al., 2015; Sun et al., 2018). Concentrations at UT01 are expected to be strongly  
dependent upon the extent to which local winds blow directly from that farm to the AMoN site throughout the two-week  
205 integration period. Not surprisingly, the UT01 site has the highest annual mean concentration ( $16.2 \mu\text{g}/\text{m}^3$ ) in the entire AMoN  
network (three times higher than the next one). Furthermore, this AMoN site may be particularly susceptible to trends in animal  
operations or management practices at the farm. While it is possible the measurements of UT01 are representative of the local  
region, it is beyond the scope of this work to make such an assessment of its representativeness.

## 210 2.3 Trend analyses

### 2.3.1 Oversampled $\text{NH}_3$ maps

From 2008 to 2018, a  $0.02^\circ \times 0.02^\circ$  ( $\sim 2$  km) annual mean  $\text{NH}_3$  map in the CONUS was created each year based on a physical  
oversampling algorithm that represents the satellite spatial response functions as generalized 2-D super Gaussian functions  
(Sun et al., 2018). This algorithm weighs IASI measurements by their uncertainties which include varying sensitivities to  
215 thermal contrast as described in Sun et al. (2018) and Wang et al. (2021). For each year, seasonally averaged oversampling  
maps were also generated for spring (March, April, and May, MAM), summer (June, July, and August, JJA), fall (September,  
October, and November, SON), and winter (December, January, and February, DJF). For each season, we were able to achieve  
sufficiently overlapped IASI pixels through calculating the sum of the unnormalized spatial response function (SRF) of the  
oversampling results (Sun et al., 2018; Wang et al., 2021).

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### 2.3.2 Mann-Kendall test and Theil-Sen's slope estimator

We use the Mann-Kendall (MK) test and Theil-Sen's slope estimator for  $\text{NH}_3$  trend analyses. The non-parametric Mann-  
Kendall test and Theil-Sen's slope estimator are widely used in detecting trends of variables in meteorology and hydrology  
fields (Ahn and Merwade, 2014; Kendall, 1975; Yue and Wang, 2004). The Kendall rank correlation coefficient, commonly  
225 referred to as Kendall's  $\tau$  coefficient, is a statistic used to measure the rank correlation. An MK test is a non-parametric  
hypothesis test for statistical dependence based on the Kendall's  $\tau$  coefficient. The Theil-Sen's slope estimator is commonly  
used to fit a line to data points by calculating the median of the slopes of all lines through pairs of points.

Different from simple linear regression, the Mann-Kendall test and Theil-Sen's slope estimator do not require the data to  
230 follow normal distribution and therefore are more robust to any outliers (Yue and Wang, 2004). This method is computationally



efficient and is insensitive to outliers. For skewed and heteroskedastic data, the Theil-Sen estimator can be significantly more accurate than linear least squares regression. For normally distributed data, the Theil-Sen estimator competes well against the least squares in terms of statistical power (Yue and Wang, 2004).

### 3 IASI & AMoN comparison

#### 235 3.1 Sensitivity to spatial windows

For the initial analysis, we first show the simplest way of comparing the satellite measurements with ground observations. In other words, we center on each AMoN site, average all IASI observations within a given radius of the AMoN site for the sampling time frame (2 weeks) for comparison, and refer to that radius as a spatial window. If the distribution of NH<sub>3</sub> pixels is spatially uniform, increasing the spatial window may improve the correlation between the two because of a larger number of IASI pixels. Larger spatial windows include more IASI pixels than smaller spatial windows but at the expense of potentially not being representative of the AMoN site. In addition, a larger region is likely to encompass NH<sub>3</sub> spatial gradients. In contrast, small spatial windows may only include a limited number of IASI pixels, encompassing more inherent noise in the satellite measurements, especially if close to the detection limit. Each integrated 2-week AMoN measurement for each site was correlated with any relevant satellite data within the spatial window (total of 104 AMoN sites with 16,093 measurements). Correlations between IASI and AMoN for different spatial windows (15 km, 25 km, 50 km, and 100 km) are summarized in Table 1. The minimum spatial window radius of 15 km is based upon an approximate scale for NH<sub>3</sub> hotspots (Wang et al. 2021).

As the spatial window becomes larger, mean temporal coverage (defined as the percentage of days with available IASI data of the 2-week AMoN sampling period) and number of IASI pixels both have significant increases, but the Pearson's r coefficient only increases slightly from 0.35 at a 15 km spatial window to 0.44 at a 100 km spatial window. Indeed, doubling the spatial window from 50 km to 100 km yields an almost tripled mean number of IASI pixels, yet maintains the almost the same correlation with  $r = 0.45$  and  $r = 0.44$ , respectively. This indicates that including IASI pixels at longer distances from the AMoN site may not be representative of the AMoN site, especially near sources or regions with complex topography. The slightly increased r value over spatial window range may result from a tradeoff between averaging spatial gradients versus integrating a larger number of IASI pixels to improve the signal-to-noise ratio of the satellite measurements. To balance these competing effects, we select 25 km as the nominal spatial window for the further comparisons.

260 **Table 1.** AMoN & IASI comparison results for different spatial windows

Spatial window	15 km	25 km	50 km	100 km
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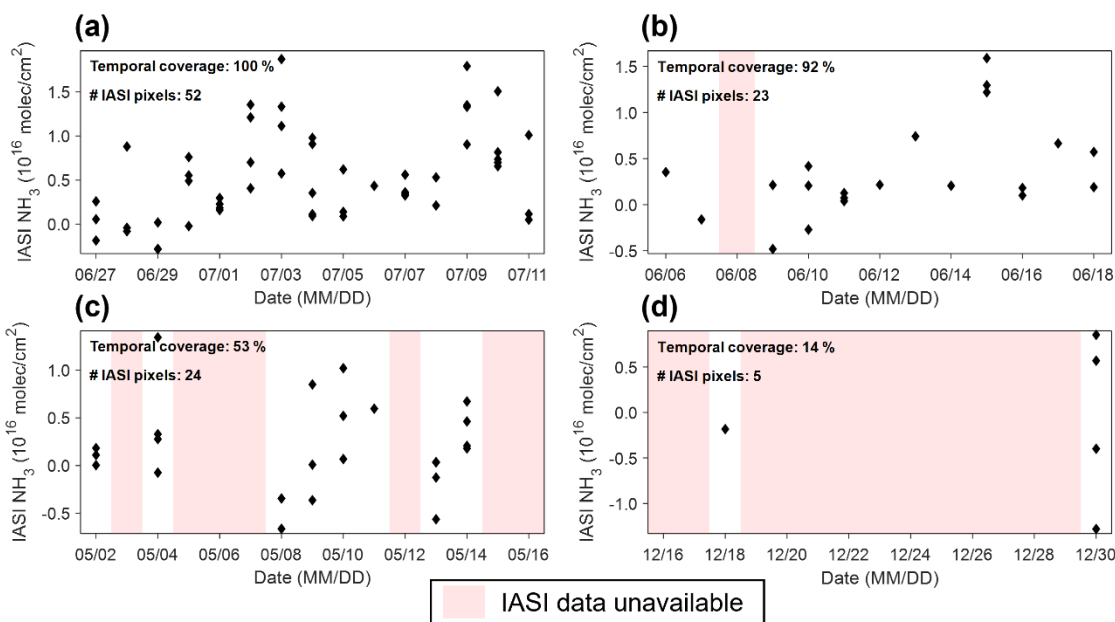




<b>Pearson's r</b>	0.35	0.41	0.45	0.44
<b>Mean temporal coverage per pair (%)</b>	31	44	57	71
<b>Mean # IASI pixels per pair</b>	7	17	69	278
<b># AMoN &amp; IASI pairs</b>	14734	15543	15933	16022

### 3.2 Sensitivities to temporal coverage and number of IASI pixels

NH<sub>3</sub> is a short-lived species with a complicated diurnal profile (Nair and Yu, 2020) and the potential for large day-to-day concentration changes because of the variability in emissions, wind speed, temperature, PBL height, and aerosol partitioning (Golston et al., 2020; Miller et al., 2015). Thus, the temporal distribution of satellite measurements within the AMoN measurement period may impact the comparison. Fig. 1 illustrates four examples where the number of IASI pixels, and their relative distribution throughout the 2-week AMoN integration period, may impact the comparison (25 km spatial window). An ideal comparison case would have a uniform number of IASI measurements on each day during the approximate 14-day AMoN measurement period, similar to the case shown in Fig. 1a. In this case, there is no specific day having more weight than the other when calculating the biweekly mean. More common, however, are cases where some days have no satellite measurements due to clouds or low thermal contrast. For example, Fig. 1b has one missing day (N=23 satellite measurements) but with an otherwise even distribution throughout the remainder of the period, while Fig. 1c (N=24) has nearly the same number of satellite measurements as Fig. 1b but clustered on only 8 of the 15 days. Finally, there are also many cases where selected day(s) have few or no IASI measurements at all (Fig. 1d). When neither temporal coverage nor the number of IASI pixels are high, one can still calculate the matched IASI NH<sub>3</sub> column for this AMoN sample, but the result is unlikely to be more representative than a more temporally distributed comparison.



280 **Figure 1.** Examples of IASI data temporal coverage over the biweekly AMoN sampling period: (a) several IASI measurements every day during the 2-week sampling period; (b) a few IASI measurements for most time of the 2-week sampling period; (c) many IASI measurements but only in several days during the 2-week sampling period; (d) sparse IASI measurements for only several days during the 2-week sampling period.

285 To this end, we explore the correlation with IASI data's temporal coverage of the 2-week sampling period and total number of IASI pixels within the 2-week AMoN sampling period using the 25 km spatial window. For example, the temporal coverages for Fig. 1 are 100%, 92%, 53%, and 14%, respectively, and the number of IASI pixels are 52, 23, 24, and 5, respectively. The impact of different temporal averaging and number of IASI pixels requirements are summarized in Table 2 and Table 3, respectively. Increasing temporal coverage and number of IASI pixels both yield higher  $r$  values than any of the simple spatial windows alone. Table 2 shows that the correlation improves to  $r = 0.74$  when the temporal coverage is  $\geq 80\%$ , suggesting a significant impact of temporal coverage of the IASI data. The IASI and AMoN correlations also increase over a simple spatial window with increasing numbers of IASI pixels, yet the impact is not as strong ( $r = 0.63$  for  $N \geq 40$ ) as the sensitivity to temporal coverage.

295

**Table 2.** The impact of IASI data's temporal coverage for the 2-week AMoN sampling period (25 km spatial window)

IASI temporal coverage per pair (%)	[0, 20)	[20, 50)	[50, 80)	[80, ∞)
-------------------------------------	---------	----------	----------	---------



<b>r</b>	0.17	0.29	0.47	0.74
<b>Mean # IASI pixels per pair</b>	3	13	26	38
<b># AMoN &amp; IASI pairs</b>	1766	7641	5137	999

**Table 3.** The impact of # IASI pixels (25 km spatial window)

<b># IASI pixels per pair</b>	<b>[0, 10)</b>	<b>[10, 20)</b>	<b>[20, 40)</b>	<b>[40, ∞)</b>
<b>r</b>	0.16	0.37	0.50	0.63
<b>Mean temporal coverage per pair (%)</b>	22	42	61	80
<b># AMoN &amp; IASI pairs</b>	4533	5025	5309	676

300

Because the temporal coverage and number of IASI pixels are not independent variables, additional analyses are conducted to study the sensitivity of these two effects using Monte-Carol method. First, the available dataset is filtered to cases when at least one of the fourteen days have multiple IASI measurements per AMoN measurement, at least 7 days of the 14-day sampling period had at least one IASI measurement, and the total number of IASI pixels is at least 20. The number of days with available IASI measurement is denoted by T. Two opposite approaches are explored for 104 qualified AMoN sites:

305

(1) Maximized temporal coverage (TC<sub>max</sub>): only one IASI pixel is randomly selected to represent that day, and the total number of IASI pixels equals T ( $T \leq 14$ ). In this case, the temporal coverage is maximized.

310

(2) Minimized temporal coverage (TC<sub>min</sub>): only days with the largest number of IASI pixels are selected until the total number of IASI pixels equals T ( $T \leq 14$ ). In this case, the temporal coverage is minimized, and the total number of selected IASI pixels is same with TC<sub>max</sub>.

315

For each AMoN site, we performed the two different sampling strategies for 100 times, then calculated the median r value to represent each site using the maximum and minimum coverage approaches. Fig. 2a shows the histogram and normalized fit of change in r ( $\Delta r = TC_{max} - TC_{min}$ ) for each site between the two scenarios with the number of bins determined by Sturge's

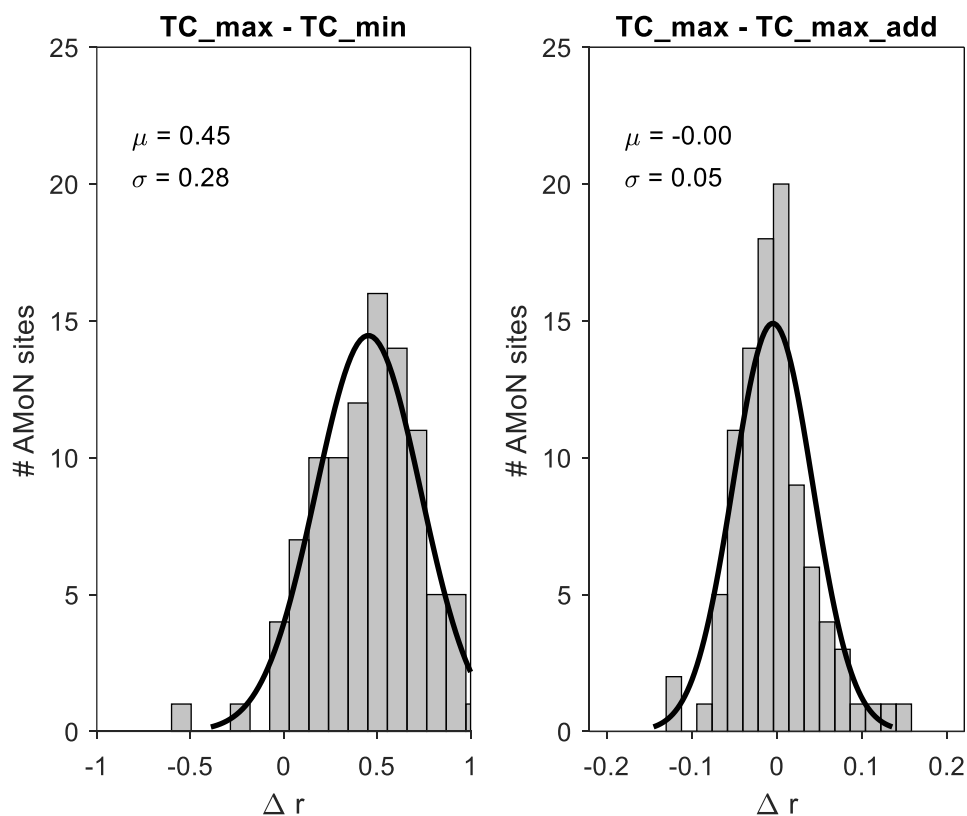


rule. The increased correlation of  $\Delta r = 0.45 \pm 0.28$  shows the large impact of temporal coverage. The total number of IASI pixels used for the two strategies were identical.

320 To further investigate the impact of including more IASI pixels after maximizing temporal coverage, we also test the process described in (1) and then randomly added (20-T) more IASI pixels from the remaining IASI pixels and referred to it as TC\_max\_add. Fig. 2b shows that the changes  $\Delta r$  between TC\_max and TC\_max\_add are small ( $-0.00 \pm 0.05$ ). For the TC\_max strategy, the initial number of IASI pixels was between 7 and 14, which means using TC\_max\_add strategy result in a 43 ~ 186 % increase in the number of IASI pixels compared to TC\_max alone. Adding more IASI pixels does not have a significant  
325 impact on the r values, indicating that maximized temporal coverage alone is the most important factor when comparing IASI to AMoN stations.

After applying a temporal coverage requirement (temporal coverage  $\geq 80$  %) to filter the overall dataset, we revisit the sensitivity of the agreement between spatial windows. The smaller spatial window now yields better agreement than the larger  
330 spatial windows (Table 4). Compared with Table 1 which has no filter for temporal coverage, the r values in Table 4 increase for all spatial windows. The correlations are clearly better for smaller spatial windows ( $r = 0.74$  for 25 km versus  $r = 0.48$  for 100 km). In this way, the use of a larger spatial window is indeed a tradeoff between the increasing temporal coverage versus incorporating a larger spatial gradient. The results further demonstrate that the IASI pixels far from the AMoN sites may not be representative to the AMoN site.

335



**Figure 2.** The change in  $r$  values for individual AMoN sites using different sampling strategies: **(a)** maximized temporal coverage (TC\_max); minimized temporal coverage (TC\_min) and **(b)** maximized temporal coverage & randomly adding more pixels (TC\_max\_add).

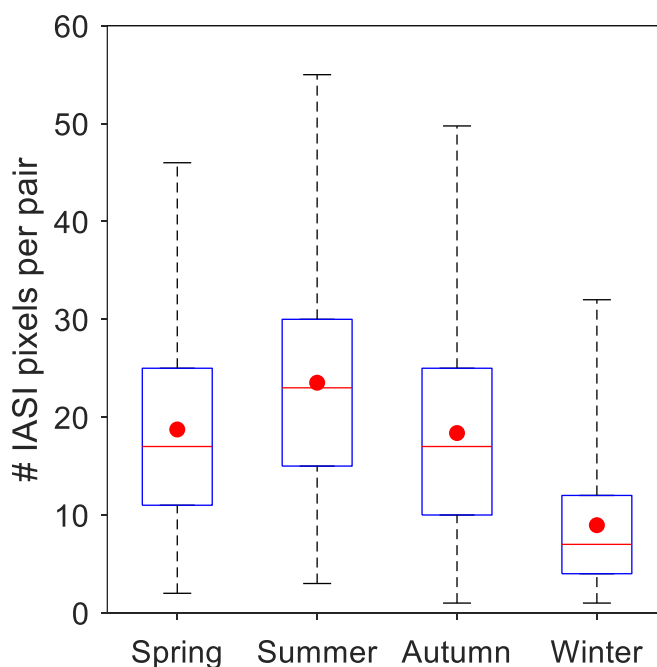
**Table 4.** AMoN & IASI comparison results for different spatial windows (temporal coverage  $\geq 80\%$ )

Spatial window	15 km	25 km	50 km	100 km
Pearson's $r$	0.76	0.74	0.58	0.48
Mean # IASI pixels per pair	19	38	119	392
# AMoN & IASI pairs	105	999	3138	6899



### 3.3 Sensitivity to seasons and temporal averaging

AMoN has similar numbers of measurements in spring (March, April, May), summer (June, July, August), autumn (September, October, November), and winter (December, January, February), while the mean number of IASI pixels (# IASI pixels) per pair in winter is only around half of other seasons (Fig. 3). In winter, low thermal contrasts result in low sensitivity of thermal infrared sounder, which explains the low number of IASI pixels in winter (Clarisse et al., 2010; Guo et al., 2021). The lower sensitivity of the infrared thermal sounder measurements in winter results in higher uncertainties, and thus comparisons between IASI and AMoN are especially important. When temporal coverage is at least 80%, IASI wintertime data still have good agreement with AMoN ( $r = 0.61$ ) although the comparison are limited to only a few AMoN & IASI pairs ( $N = 33$ ). IASI in general only provides a small number of pixels in winter, however, it indeed has the capability of reflecting surface  $\text{NH}_3$  variations even in winter.



**Figure 3.** Boxplot of number of IASI pixels per pair for spring, summer, autumn, and winter. The boxes denote the 25<sup>th</sup> and 75<sup>th</sup> percentiles, the whiskers denote the 1<sup>st</sup> and 99<sup>th</sup> percentiles, and the red dot denotes the mean.

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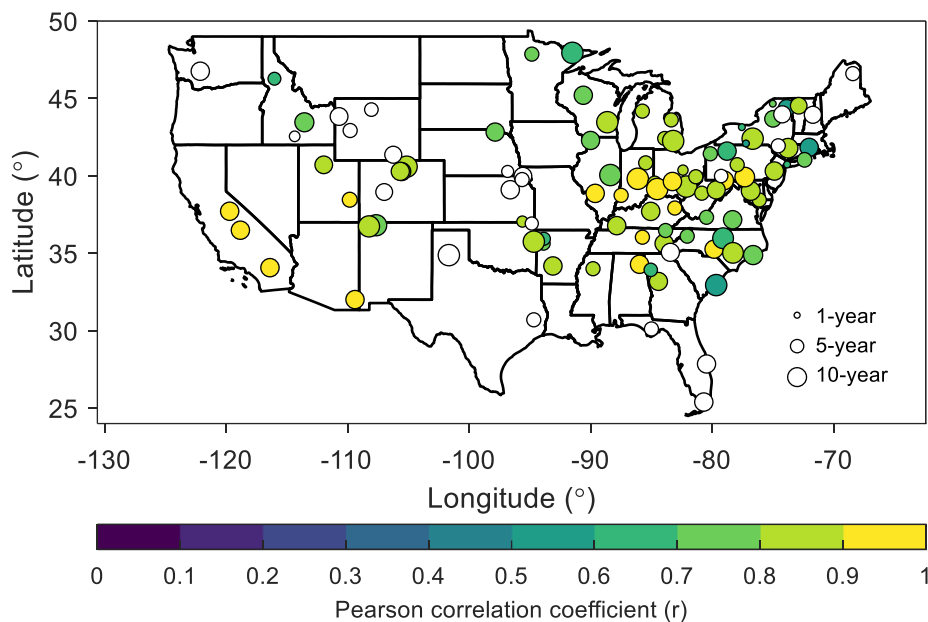
The results in 3.1 and 3.2 have already shown the importance of spatial window and temporal coverage. The temporal averaging, such as the tessellation oversampling and physical oversampling, is a common method to achieve a high spatial resolution map by sacrificing the temporal resolution (Sun et al., 2018; Van Damme et al., 2018; Wang et al., 2021). Here we neglect the interannual variability and calculate the multi-year averaged seasonal IASI  $\text{NH}_3$  concentrations using the 25 km



365 spatial window. By averaging the multi-year IASI data, the impacts of temporal coverage are alleviated because both temporal  
coverages and numbers of IASI pixels increase. Among the 101 AMoN sites with at least one full year data and available IASI  
v3.1r NH<sub>3</sub> data, 49 sites show strong agreement with IASI with  $r > 0.8$ , 29 sites have moderate agreement of  $0.5 < r \leq 0.8$ ,  
while 23 sites do not have statistically significant agreements (Fig. 4). If taking all data into consideration, the overall  $r$  value  
for the CONUS is 0.69. The AMoN sites with higher NH<sub>3</sub> concentrations tend to show better agreements between AMoN and  
370 IASI. The median AMoN NH<sub>3</sub> annual mean concentrations for all sites is 0.86  $\mu\text{g}/\text{m}^3$ . Most sites with no statistically significant  
agreements have a low NH<sub>3</sub> concentration (median: 0.48  $\mu\text{g}/\text{m}^3$ ). Currently, most AMoN sites are located in low or moderate  
NH<sub>3</sub> concentration regions with a lack of sites in the NH<sub>3</sub> hotspots (Wang et al., 2021) and urban areas, complicating the  
comparison between AMoN and IASI.

375 The above agreement demonstrates that IASI NH<sub>3</sub> column reflects the variation of the surface NH<sub>3</sub> concentration at seasonal  
resolution. For regions without any available ground measurements, IASI NH<sub>3</sub> observations can be used to help better  
understand the NH<sub>3</sub> variations. However, large differences exist among the relationships between IASI and AMoN NH<sub>3</sub>  
concentrations over different AMoN sites (an example of linear regression plot in Fig. 5b). Even for AMoN sites with excellent  
correlation ( $r > 0.8$ ), the slopes vary a lot, ranging from  $0.08 - 1.4 \times 10^{16}$  molec/cm<sup>2</sup> per  $\mu\text{g}/\text{m}^3$ . For instance, two AMoN sites  
380 in California, Joshua Tree National Park (CA 67) and Sequoia & Kings Canyon National Park (CA 83), both exhibit great  
seasonality agreements with IASI ( $r = 0.97$  and  $r = 0.99$ , respectively) but the slope for CA 83 is 44 % higher than CA 67. The  
difference between the slopes suggests that although IASI is able to capture the general seasonality, the relationship between  
NH<sub>3</sub> column and surface NH<sub>3</sub> is distinctly different due to complicated topography, meteorology, and other factors at different  
AMoN sites.

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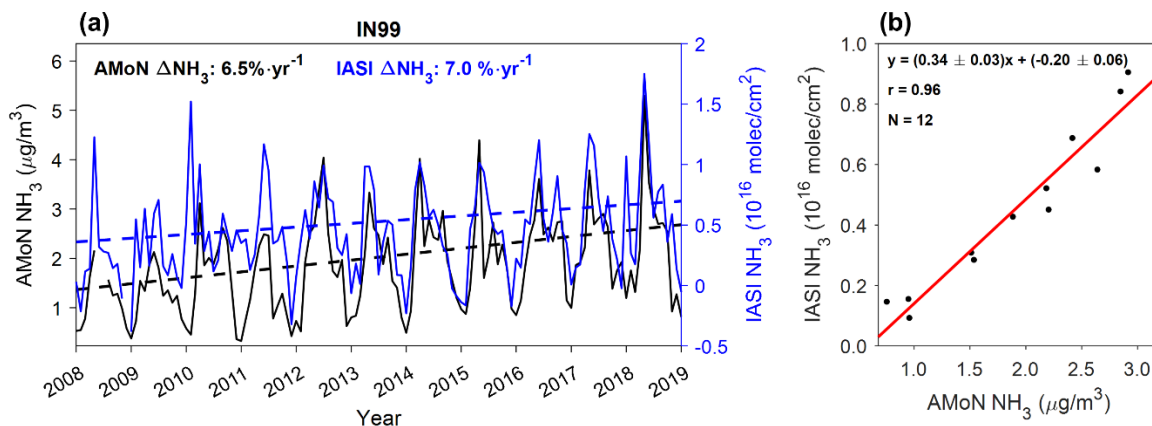
**Figure 4.** Multi-year averaged comparison results between AMoN sites and the IASI observations within 25 km of the AMoN sites at monthly resolution. Circles without filled color denote the AMoN sites with no statistically significant correlation with IASI. The circle sizes denote the length of AMoN data record.

## 4 Trend analysis

### 4.1 Trend in the CONUS

The methodology and comparison results in section 3 demonstrate that IASI  $\text{NH}_3$  can be used to estimate regional  $\text{NH}_3$  trends over the last decade. In this regard, satellite  $\text{NH}_3$  observations will be used to augment the AMoN observed  $\text{NH}_3$  trends in the CONUS over the last decade. We include AMoN trend analysis only for sites with full year coverage during 2008 - 2018 (N=13). Strong evidence of increasing  $\text{NH}_3$  concentrations in the U.S. comes from both ground-based observations and satellite measurements (Van Damme et al., 2021; Warner et al., 2017; Yao and Zhang, 2016; Yao and Zhang, 2019; Yu et al., 2018). Fig. 5a shows monthly IASI and AMoN timeseries in from Indianapolis, Indiana, USA (IN 99). The strong correlation ( $r = 0.96$ ) between the two measurements is shown in Fig. 5b. Although the  $\text{NH}_3$  seasonality remain consistent from 2008 to 2018 - namely spring maxima and secondary maxima in fall with lowest values in winter - both AMoN and IASI also show increasing trends of  $\text{NH}_3$  concentrations over the entire timeseries. AMoN shows a trend of  $6.5\% \cdot \text{yr}^{-1}$  while IASI shows a trend of  $7.0\% \cdot \text{yr}^{-1}$ .





**Figure 5.** (a) 2008 – 2018 monthly averaged NH<sub>3</sub> trends for AMoN site in Indianapolis, Indiana, U.S. (IN 99) and IASI NH<sub>3</sub> observations within 25 km of IN 99; (b) seasonality correlation between AMoN and IASI NH<sub>3</sub> for IN 99.

A long-term trend analysis was performed using AMoN and IASI data to examine the agreement between the datasets and explore any regional differences. IASI NH<sub>3</sub> columns smaller than the 5<sup>th</sup> percentile ( $0.5 \times 10^{15}$  molec/cm<sup>2</sup>) of the 11-year NH<sub>3</sub> average in the CONUS region were excluded to avoid spurious trend results caused by the higher noise in these measurements.

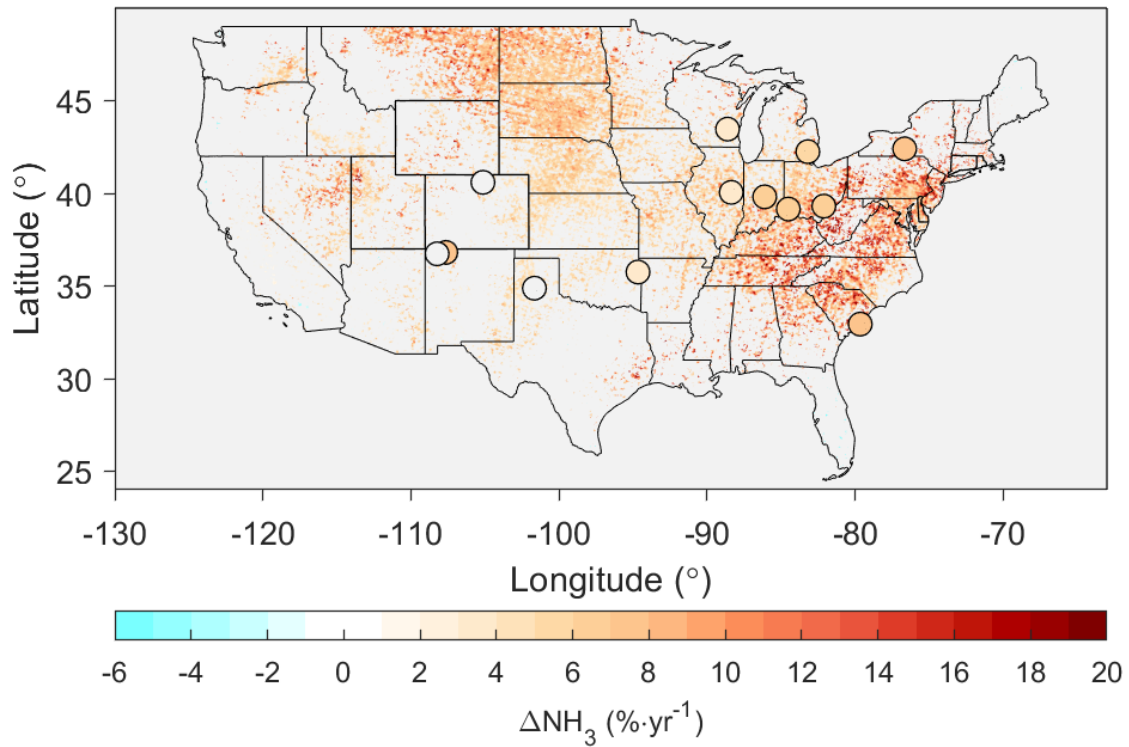
To perform the interannual trend analysis, we require each region or site to have at least one valid measurement in each season to alleviate the possible bias due to seasonal variations. Fig. 6 shows the annual percentage change for both IASI and AMoN. Most regions in the CONUS have increasing NH<sub>3</sub> concentrations based on the 11-year IASI observations (median:  $6.8\% \cdot \text{yr}^{-1}$ ), including eastern U.S., Midwest, and parts of the western U.S. 10 out of 13 AMoN sites have statistically significant NH<sub>3</sub> increases. AMoN data in general suggest similar increases (median:  $6.7\% \cdot \text{yr}^{-1}$ ). When plotting the trends of AMoN sites against the median of IASI trends within a 25 km spatial window (Fig. 7), a moderate correlation ( $r = 0.66$ ) was found between IASI and AMoN NH<sub>3</sub> trends. IASI in general suggested a higher NH<sub>3</sub> increase compared to AMoN (slope:  $1.26 \pm 0.51$ ) with the ratio larger than one for most sites. The absolute NH<sub>3</sub> change also is in correspondence with the previous study, with significant NH<sub>3</sub> increases across the CONUS regions, especially in the Midwest (Van Damme et al., 2021).

The spatial consistency across the datasets differs significantly. Both AMoN and IASI suggest  $\sim 5\% \cdot \text{yr}^{-1}$  NH<sub>3</sub> increases in the Great Lake Region, while IASI suggests a higher NH<sub>3</sub> increase in the eastern US compared with AMoN. The IASI trend analysis results suggest a significant NH<sub>3</sub> increase in the northern Great Plains, e.g., North Dakota, South Dakota, and Montana, yet there are no AMoN sites in this region. Furthermore, the trends are consistent with the NH<sub>3</sub> emissions increases caused by increased N fertilizer usage in the northern Great Plains (Cao et al., 2020b). McHale et al. (2021) showed that wet-precipitation NH<sub>4</sub><sup>+</sup> concentrations based on NADP observations suggested the highest increases in the Great Plains, the Rocky Mountain Region, and the Great Lake Region from 2000 to 2017, which is geographically consistent with the NH<sub>3</sub> trends observed by



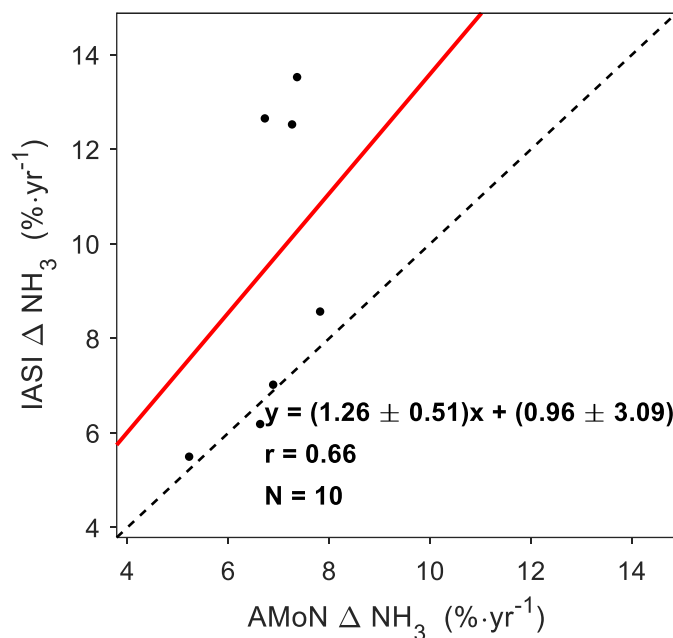
both AMoN and IASI. If considering the CONUS as a whole and calculating the annual mean  $\text{NH}_3$  for the whole CONUS during 2008 – 2018 to derive the overall trend in CONUS, the IASI  $\text{NH}_3$  change for 2008 – 2018 is  $(3.9 \pm 2.2) \% \cdot \text{yr}^{-1}$  and  $(1.3 \pm 0.8) \times 10^{14} \text{ molec/cm}^2 \cdot \text{yr}^{-1}$ , similar with the trend in the previous study  $(3.4 \pm 0.6) \% \cdot \text{yr}^{-1}$  and  $(1.1 \pm 0.4) \times 10^{14} \text{ molec/cm}^2 \cdot \text{yr}^{-1}$  (Van Damme et al., 2021).

In terms of trend in  $\text{NH}_3$  hotspots, which are here defined as regions where the IASI  $\text{NH}_3$  column is larger than the 95<sup>th</sup> percentile of 11-year CONUS map ( $6.7 \times 10^{15} \text{ molec/cm}^2$ ), the median of  $\text{NH}_3$  trend is  $4.7\% \cdot \text{yr}^{-1}$ , indicating that the regions of the largest emissions sources are also realizing increasing concentrations over time. Although the percent changes in the regions with the highest concentrations are smaller compared with the trend in CONUS median ( $8.0\% \cdot \text{yr}^{-1}$ ), in terms of the absolute changes, the median trend of  $\text{NH}_3$  columns over these  $\text{NH}_3$  hotspots are higher compared with the CONUS median ( $3.7 \times 10^{14} \text{ molec/cm}^2 \cdot \text{yr}^{-1}$  vs.  $2.8 \times 10^{14} \text{ molec/cm}^2 \cdot \text{yr}^{-1}$ ). The top 10  $\text{NH}_3$  hotspots in CONUS regarding column-areal weighting all exhibit increasing  $\text{NH}_3$  concentrations from 2008 to 2018 (Table 5). Within these hotspots, the central Great Plains experience the largest  $\text{NH}_3$  increase (median:  $5.0\% \cdot \text{yr}^{-1}$ ,  $4.0 \times 10^{14} \text{ molec/cm}^2 \cdot \text{yr}^{-1}$ ) while the San Joaquin Valley (median:  $2.0\% \cdot \text{yr}^{-1}$ ,  $1.6 \times 10^{14} \text{ molec/cm}^2 \cdot \text{yr}^{-1}$ ) and Imperial County, California (median:  $2.1\% \cdot \text{yr}^{-1}$ ,  $1.9 \times 10^{14} \text{ molec/cm}^2 \cdot \text{yr}^{-1}$ ) have a smaller change change.



445

**Figure 6.** Trend analysis for IASI NH<sub>3</sub> (2008 - 2018) and AMoN NH<sub>3</sub> measurements in the contiguous U.S. The gray color indicates no statistically significant change. The circle size denotes the length of AMoN data record.



450 **Figure 7.** Comparison between AMoN and IASI NH<sub>3</sub> trends (25 km spatial window) for AMoN sites with available nearby IASI trend data

**Table 5.** 2008 – 2018 IASI observed NH<sub>3</sub> trend in the top 10 NH<sub>3</sub> hotspots (column-areal weighting) in CONUS

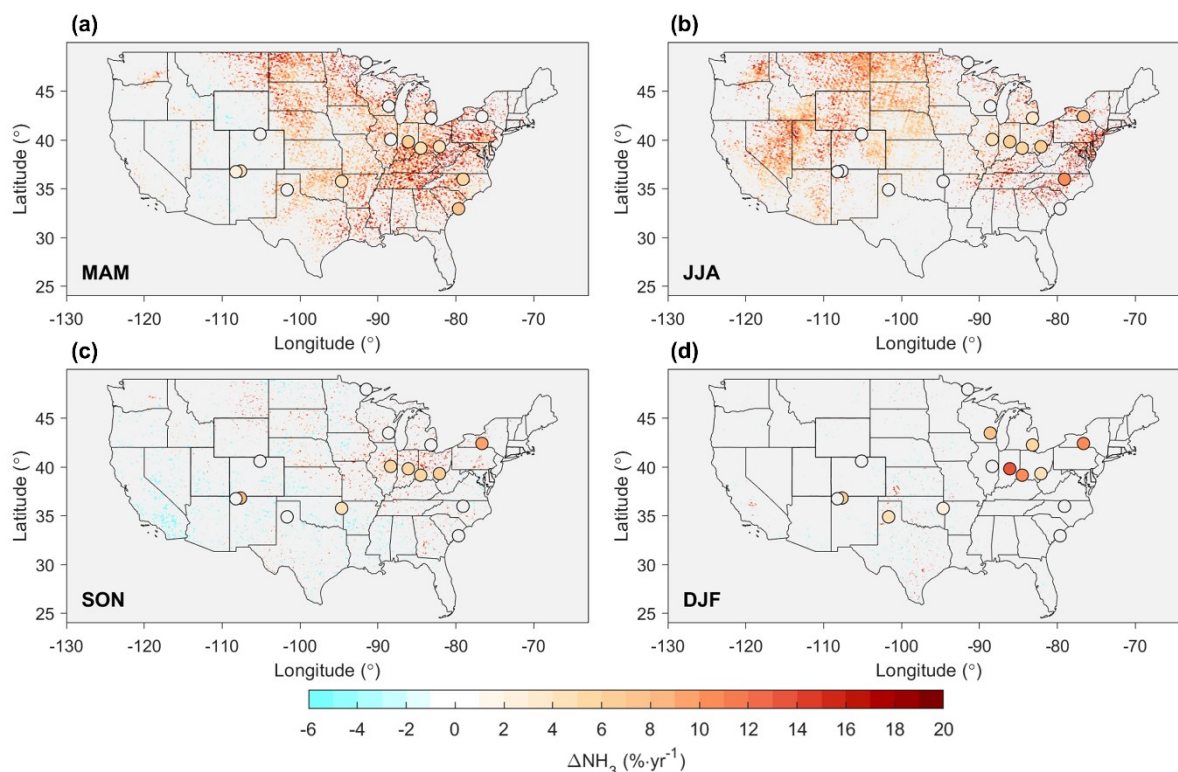
Hotspots	% · yr <sup>-1</sup>	10 <sup>14</sup> molec/cm <sup>2</sup> · yr <sup>-1</sup>
Central Great Plains	5.0	4.0
The San Joaquin Valley	2.0	1.6
North Oklahoma	3.9	2.9
Texas panhandle	3.6	2.8
Central Iowa	4.4	3.3
The Snake River Valley	3.8	3.3
Southeast Iowa	5.2	3.9



<b>Beadle County, South Dakota</b>	8.3	6.0
<b>Weld County, Colorado</b>	3.6	2.9
<b>Imperial County, California</b>	2.1	1.9

455 To provide a detailed insight of the increasing  $\text{NH}_3$  over the CONUS, we further perform trend analyses for different seasons  
 (Fig. 8). In spring, significant  $\text{NH}_3$  increases are found in the Midwest and Eastern US. In summer,  $\text{NH}_3$  increases shift to the  
 western US and part of the eastern US. AMoN and IASI seasonality clustering results show that the Midwest and eastern  
 United States, dominated by fertilizer  $\text{NH}_3$  emissions, have a broad, spring maximum of  $\text{NH}_3$ , while the western United States,  
 dominated by volatilization of livestock waste  $\text{NH}_3$  emissions, in contrast, show a narrower midsummer peak (Wang et al.,  
 460 2021). The spatial patterns of spring and summer  $\text{NH}_3$  trends are in agreement with the seasonality clustering results, indicating  
 that increasing  $\text{NH}_3$  emissions caused by agricultural activities may contribute to  $\text{NH}_3$  concentration increase. The increasing  
 wildfire activities in the western U.S. may also contribute to  $\text{NH}_3$  increases (Lindaas et al., 2021a, b). In fall and winter, most  
 regions in the U.S. do not have statistically significant IASI  $\text{NH}_3$  trends, and a decreasing  $\text{NH}_3$  trend is observed by IASI in  
 the Southwest US in fall. In contrast, AMoN data suggest a notable  $\text{NH}_3$  increase in Northeast and the Corn Belt region in  
 465 winter. Again, IASI data are susceptible to low thermal contrasts in winter, which to some extent explains the disagreement  
 between IASI and AMoN in winter as discussed in Section 3.3.

Wintertime  $\text{NH}_3$  plays an important role in haze episodes through the formation of aerosol phase  $\text{NH}_4\text{NO}_3$  (Shah et al., 2018;  
 Zhai et al., 2021), and increasing  $\text{NH}_3$  concentrations in winter may affect aerosol acidity and aerosol chemistry (Lawal et al.,  
 470 2018; Zheng et al., 2020). In the past decades,  $\text{NO}_x$  and  $\text{SO}_2$  emissions reductions have resulted in less  $\text{NH}_x$  partitioning into  
 particle phase  $\text{NH}_4^+$  (Shah et al., 2018), however, the partitioning alone is not able to fully explain the significant  $\text{NH}_3$   
 concentration increases (Yao and Zhang, 2019; Yu et al., 2018). The change of meteorological conditions, such as increasing  
 air temperatures may also contribute to the increasing  $\text{NH}_3$  trends (Warner et al., 2017; Yao and Zhang, 2019). No matter the  
 reason for increasing  $\text{NH}_3$  concentrations across the CONUS regions, the fact that both  $\text{NH}_3$  surface concentrations and  $\text{NH}_3$   
 475 column concentrations are increasing during the past decade will have significant impacts on air quality and nitrogen  
 deposition. EPA is reviewing the 2020  $\text{PM}_{2.5}$  National Ambient Air Quality Standard (NAAQS) currently set at  $12.0 \mu\text{g}\cdot\text{m}^{-3}$   
 and if the NAAQS is lowered,  $\text{NH}_3$  controls will become increasingly important for meeting the standard. Additionally, Pan  
 et al. (2021) demonstrates that  $\text{NH}_3$  transported from Colorado significantly increased the dry  $\text{NH}_3$  deposition the Rocky  
 Mountain National Park. Increasing gas phase  $\text{NH}_3$  may result in longer spatiotemporal scales for dry nitrogen deposition,  
 480 leading to adverse impacts on remote regions and sensitive ecosystems (Phoenix, et al., 2006). Reduction of  $\text{NH}_3$  emissions is  
 critical to protect human health and the biodiversity in sensitive ecosystems (Ellis et al., 2013, Hill et al., 2019).



**Figure 8.** 2008 – 2018  $\text{NH}_3$  trend for different seasons based on IASI  $\text{NH}_3$  measurements in the contiguous U.S. (a) spring (March, April, May); (b) summer (June, July, August); (c) autumn (September, October, November); (d) winter (December, January, February)

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#### 4.2 Trend in the urbanized areas

The short lifetime of  $\text{NH}_3$  leads to strong spatial variabilities of  $\text{NH}_3$  concentrations, and most AMoN sites are not located in highly populated urban regions (Wang et al., 2021). Fig. 9 shows population coverage of AMoN in the CONUS region. Population data were retrieved from the Gridded Population of the World, Version 4 (GPWv4) (Center for International Earth Science Information Network – Columbia University, 2018). More than half of the CONUS population is at least 100 km away from an AMoN site. As mentioned in the previous discussion of spatial windows, AMoN may best represent the  $\text{NH}_3$  variations for regions within  $\sim 10$  km radius, and less than 2% of CONUS population are within 10 km of an AMoN site. More urban AMoN sites are needed to represent the urban areas and better quantify  $\text{NH}_3$  emissions from mobile sources, trends in the urban areas. Satellite observations are the only dataset that can currently be used to investigate source contributions and trends in population centers (Cao et al., 2022).

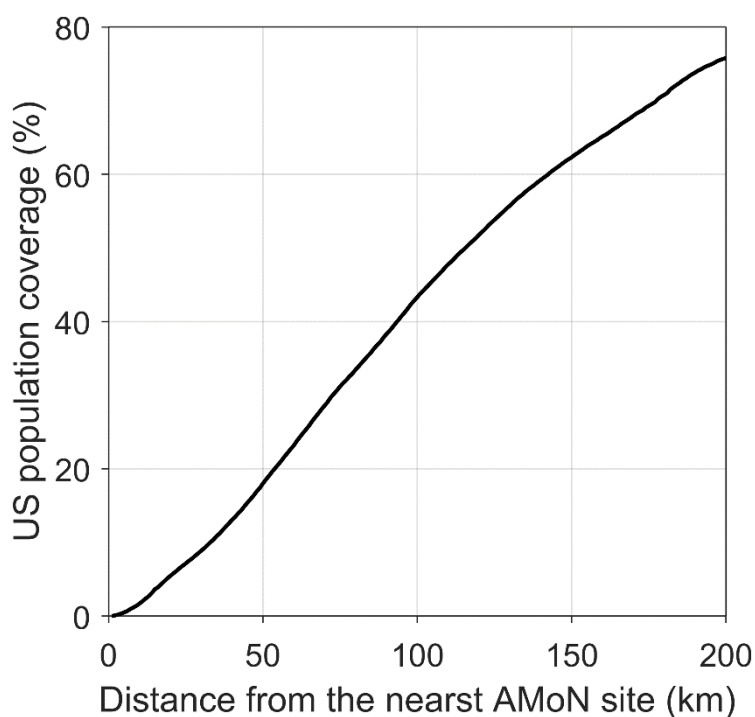
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We retrieved urban area data from the 2010 US Census, which includes two different types of urban areas: Urbanized Areas (UAs) of 50,000 or more people and Urban Clusters (UCs) of at least 2,500 and less than 50,000 people (U.S. Census Bureau,



2012). The urban areas have a similar  $\text{NH}_3$  trend compared with CONUS ( $8.1\% \cdot \text{yr}^{-1}$  vs.  $8.0\% \cdot \text{yr}^{-1}$ ), suggesting a simultaneous  
 500  $\text{NH}_3$  increase in both urban and rural areas. The top ten most populous urbanized areas almost all exhibit significant  $\text{NH}_3$   
 increases with the exception of Miami, Florida, which has a negative trend and Dallas, Texas, without any significant trend  
 (Table 6). These ten areas in total accommodate more than seventy million population, making up more than one fifth of the  
 total population in the CONUS. The  $\text{NH}_3$  increase in these densely populated areas and its impact on the aerosol chemistry  
 need to be further addressed.

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**Figure 9.** The population coverage of AMoN sites.

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**Table 6.** 2008 – 2018 IASI  $\text{NH}_3$  trend in the top 10 most populous urbanized areas

Urbanized Area	Population (million)	$\% \cdot \text{yr}^{-1}$	$10^{14} \text{ molec/cm}^2 \cdot \text{yr}^{-1}$
New York--Newark, NY--NJ--CT	18.0	10.8	2.0
Los Angeles--Long Beach--Anaheim, CA	12.0	4.3	2.1



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<b>Chicago, IL--IN</b>	8.6	5.2	2.5
<b>Miami, FL</b>	5.5	-25.2	-1.5
<b>Philadelphia, PA--NJ--DE--MD</b>	5.4	10.9	2.6
<b>Dallas--Fort Worth--Arlington, TX</b>	5.1	/	/
<b>Houston, TX</b>	4.9	7.9	2.0
<b>Washington, DC--VA--MD</b>	4.6	9.0	2.2
<b>Atlanta, GA</b>	4.5	9.4	2.2
<b>Boston, MA--NH--RI</b>	4.2	10.5	1.4

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## 515 **5 Implications**

Under favorable conditions, IASI NH<sub>3</sub> columns correlate with AMoN NH<sub>3</sub> surface concentrations even at the 2-week scale and for low concentration regions ( $r = 0.74$  when temporal coverage  $\geq 80\%$ ). IASI measurements' temporal coverage of AMoN's 2-week sampling period dominates the agreement presumably because of the larger day-to-day variability of NH<sub>3</sub>. The agreement demonstrates the strong potential for using IASI NH<sub>3</sub> columns to bridge the spatial gaps of the AMoN network.

520 The global coverage of satellite measurements enables IASI NH<sub>3</sub> product to serve as an alternative dataset in countries and regions that do not have any NH<sub>3</sub> monitoring networks, particularly in developing countries. For example, India is the second most populated country in the world with a sixth of the world's population, and recent study has shown that the unique role of NH<sub>3</sub> in forming massive chloride aerosols (up to 40  $\mu\text{g}/\text{m}^3$ ) in India (Gunthe et al., 2021). However, there are currently no long-term NH<sub>3</sub> ground monitoring networks in India, impeding the efforts to estimate and control NH<sub>3</sub> emissions (Beale et al.,

525 2022). IASI's low sensitivity to wintertime NH<sub>3</sub> shows the value of the more sensitive AMoN sites. Extra attention is needed when using IASI data in such circumstances.

The increasing NH<sub>3</sub> in the CONUS (median:  $6.8\% \cdot \text{yr}^{-1}$ ,  $2.8 \times 10^{14} \text{ molec}/\text{cm}^2 \cdot \text{yr}^{-1}$ ), including the hotspots region (median:  $4.7\% \cdot \text{yr}^{-1}$ ,  $3.7 \times 10^{14} \text{ molec}/\text{cm}^2 \cdot \text{yr}^{-1}$ ), highlights the more important role of NH<sub>3</sub> in PM<sub>2.5</sub> formation and nitrogen deposition

530 in the future. AMoN suggests a similar NH<sub>3</sub> increase ( $6.7\% \cdot \text{yr}^{-1}$ ) as well as a similar spatial pattern with IASI. Both IASI and





AMoN show largest  $\text{NH}_3$  increase in the Midwest and eastern U.S., with a moderate agreement for AMoN sites ( $r = 0.66$ ). More co-located measurements of  $\text{PM}_{2.5}$  mass and  $\text{NH}_3$  concentrations would help assess the impact increasing trends of  $\text{NH}_3$  will have on human health. The integrated satellite and ground-based measurements are already playing a role in our understanding of under-represented  $\text{NH}_3$  emissions sources in the inventories.  $\text{NH}_3$  already dominates the reactive nitrogen  
535 deposition in the majority areas in the U.S., with the continuing efforts on  $\text{NO}_x$  emission reductions,  $\text{NH}_3$  is expected to become the key species for nitrogen deposition (Li et al., 2016) and poses adverse impacts on the nearby ecosystem regions, e.g., the National Parks (Benedict et al., 2013; Pan et al., 2021). The changing partitioning of  $\text{NH}_x$  between  $\text{NH}_3$  and  $\text{NH}_4^+$  is likely to impact the lifetime of  $\text{NH}_x$  due to differences between the removal velocity of gas phase  $\text{NH}_3$  via dry deposition and particle phase  $\text{NH}_4^+$  wet deposition. The trends vary in different seasons, with  $\text{NH}_3$  increases mainly in spring in the Midwest and  
540 eastern U.S. (cropland dominated) while in summer in the western U.S. (feedlot dominated), suggesting the impacts from agricultural activities and the necessity of developing regionally-specific emission control strategies.

Because of the scarcity of the ground monitoring sites in the urban areas, satellite  $\text{NH}_3$  measurements are extremely valuable to characterize  $\text{NH}_3$  magnitude, seasonality, and trend in densely populated areas. Satellite observations suggests  $\text{NH}_3$  increases  
545 across the U.S. urban areas (median: 8.1%). New York—Newark, NY--NJ—CT alone has more than eighteen million population, experiencing an  $10.8\% \cdot \text{yr}^{-1}$   $\text{NH}_3$  increase. Measurements from satellites will help inform where ground based  $\text{NH}_3$  samplers could be located to better understand local air quality in overburdened communities that lack resources for continuous monitors. In addition,  $\text{NH}_3$  sources in the urban areas and the related atmospheric chemistry are both poorly understood (Gu et al., 2022; Sun et al., 2017) and could be constrained by satellite  $\text{NH}_3$  observations (Cao et al., 2022).  
550 However, satellite observations alone are not able to answer all questions under the complex urban atmospheric conditions. For instance, gas phase  $\text{NH}_3$  and  $\text{HNO}_3$  can nucleate directly to form  $\text{NH}_4\text{NO}_3$  particles in cold atmospheric conditions and is likely to result in rapid growth of new atmospheric particles in winter in urban areas (Wang et al., 2020). To provide accurate and fine spatial scale  $\text{NH}_3$  observations in the urban areas, more routine ground monitoring sites are needed both in urban areas and high  $\text{NH}_3$  emission source regions.

555

## 6 Data availability

The AMoN data were downloaded from the National Atmospheric Deposition Program/National Trends Network (NADP/NTN): <https://nadp.slh.wisc.edu/networks/ammonia-monitoring-network/>. The authors acknowledge the AERIS data infrastructure (<https://www.aeris-data.fr>) for providing access to the IASI Level 2  $\text{NH}_3$  data used in this study. Population data  
560 were retrieved from Center for International Earth Science Information Network, Columbia University: <https://sedac.ciesin.columbia.edu/data/collection/gpw-v4/>. The urban areas data are downloaded from the U.S. Census Bureau: <https://www.census.gov/geographies/mapping-files.html>.



### Author contributions

MAZ and RW designed the research; RW led the analysis; KS, DP, and XG contributed to data analysis; LC, MV, LP, and  
565 CC helped with the usage of IASI data; MP helped with the usage of AMoN data; and RW wrote the paper with contributions  
from all co-authors.

### Competing interests

Competing interests. The contact author has declared that none of the authors has any competing interests.

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### References

Ahn, K. H. and Merwade, V.: Quantifying the relative impact of climate and human activities on streamflow, *J. Hydrol.*, 515,  
257–266, <https://doi.org/10.1016/j.jhydrol.2014.04.062>, 2014.



- 590 Beale, C. A., Paulot, F., Randles, C. A., Wang, R., Guo, X., Clarisse, L., Van Damme, M., Coheur, P.-F., Clerbaux, C., Shephard, M. W., Dammers, E., Cady-Pereira, K., and Zondlo, M.: Large sub-regional differences of ammonia seasonal patterns over India reveal inventory discrepancies, *Environ. Res. Lett.*, <https://doi.org/10.1088/1748-9326/AC881F>, 2022.
- Benedict, K. B., Day, D., Schwandner, F. M., Kreidenweis, S. M., Schichtel, B., Malm, W. C., and Collett, J. L.: Observations of atmospheric reactive nitrogen species in Rocky Mountain National Park and across northern Colorado, *Atmos. Environ.*, *64*, 66–76, <https://doi.org/10.1016/j.atmosenv.2012.08.066>, 2013.
- 595 von Bobruzki, K., Braban, C. F., Famulari, D., Jones, S. K., Blackall, T., Smith, T. E. L., Blom, M., Coe, H., Gallagher, M., Ghalaieny, M., McGillen, M. R., Percival, C. J., Whitehead, J. D., Ellis, R., Murphy, J., Mohacsi, A., Pogany, A., Junninen, H., Rantanen, S., Sutton, M. A., and Nemitz, E.: Field inter-comparison of eleven atmospheric ammonia measurement techniques, *Atmos. Meas. Tech.*, <https://doi.org/10.5194/amt-3-91-2010>, 2010.
- 600 Bouwman, A. F., Lee, D. S., Asman, W. A. H., Dentener, F. J., van der Hoek, K. W., and Olivier, J. G. J.: A global high-resolution emission inventory for ammonia, *Global Biogeochem Cycles*, *11*, 561–587, <https://doi.org/10.1029/97GB02266>, 1997.
- Butler, T., Vermeylen, F., Lehmann, C. M., Likens, G. E., and Puchalski, M.: Increasing ammonia concentration trends in large regions of the USA derived from the NADP/AMoN network, *Atmos. Environ.*, *146*, 132–140, <https://doi.org/10.1016/j.atmosenv.2016.06.033>, 2016.
- 605 Cao, H., Henze, D. K., Shephard, M. W., Dammers, E., Cady-Pereira, K., Alvarado, M., Lonsdale, C., Luo, G., Yu, F., Zhu, L., Danielson, C. G., and Edgerton, E. S.: Inverse modeling of NH<sub>3</sub> sources using CrIS remote sensing measurements, *Environ. Res. Lett.*, *15*, <https://doi.org/10.1088/1748-9326/abb5cc>, 2020a.
- Cao, H., Henze, D. K., Cady-Pereira, K., McDonald, B. C., Harkins, C., Sun, K., Bowman, K. W., Fu, T. M., and Nawaz, M. O.: COVID-19 Lockdowns Afford the First Satellite-Based Confirmation That Vehicles Are an Under-recognized Source of Urban NH<sub>3</sub> Pollution in Los Angeles, *Environ. Sci. Technol. Lett.*, *9*, 3–9, [https://doi.org/10.1021/ACS.ESTLETT.1C00730/ASSET/IMAGES/MEDIUM/EZ1C00730\\_M004.GIF](https://doi.org/10.1021/ACS.ESTLETT.1C00730/ASSET/IMAGES/MEDIUM/EZ1C00730_M004.GIF), 2022.
- 610 Cao, P., Lu, C., Zhang, J., and Khadilkar, A.: Northwestward cropland expansion and growing urea-based fertilizer use enhanced NH<sub>3</sub> emission loss in the contiguous United States, *Atmos. Chem. Phys.*, *20*, 11907–11922, <https://doi.org/10.5194/acp-20-11907-2020>, 2020b.
- 615 Chen, Y., Shen, H., Kaiser, J., Hu, Y., Capps, S., Zhao, S., Hakami, A., Shih, J.-S., Pavur, G., Turner, M., Henze, D., Resler, J., Nenes, A., Napelenok, S., Bash, J., Fahey, K., Carmichael, G., Chai, T., Clarisse, L., Coheur, P.-F., van Damme, M., and Russell, A.: High-resolution Hybrid Inversion of IASI Ammonia Columns to Constrain U.S. Ammonia Emissions Using the CMAQ Adjoint Model, *Atmos. Chem. Phys.*, 1–25, <https://doi.org/10.5194/acp-2020-523>, 2020.
- 620 Clarisse, L., Clerbaux, C., Dentener, F., Hurtmans, D., and Coheur, P.-F.: Global ammonia distribution derived from infrared satellite observations, *Nat. Geosci.*, *2*, 479–483, <https://doi.org/10.1038/ngeo551>, 2009.



- Clarisse, L., Shephard, M. W., Dentener, F., Hurtmans, D., Cady-Pereira, K., Karagulian, F., van Damme, M., Clerbaux, C., and Coheur, P. F.: Satellite monitoring of ammonia: A case study of the San Joaquin Valley, *J. Geophys. Res. Atmos.*, 115, 1–15, <https://doi.org/10.1029/2009JD013291>, 2010.
- 625 Dammers, E., Palm, M., Damme, M. van, Vigouroux, C., Smale, D., Conway, S., Toon, G. C., Jones, N., Nussbaumer, E., Warneke, T., Petri, C., Clarisse, L., Clerbaux, C., Hermans, C., Lutsch, E., Strong, K., Hannigan, J. W., Nakajima, H., Morino, I., Herrera, B., Stremme, W., Grutter, M., Schaap, M., Kruit, R. J. W., Notholt, J., Coheur, P.-F., and Erismann, J. W.: An evaluation of IASI-NH 3 with ground-based Fourier transform infrared spectroscopy measurements, *Atmos. Chem. Phys.*, 16, 10351–10368, <https://doi.org/10.5194/acp-16-10351-2016>, 2016.
- 630 Dammers, E., Shephard, M. W., Palm, M., Cady-Pereira, K., Capps, S., Lutsch, E., Strong, K., Hannigan, J. W., Ortega, I., Toon, G. C., Stremme, W., Grutter, M., Jones, N., Smale, D., Siemons, J., Hrpcek, K., Tremblay, D., Schaap, M., Notholt, J., and Willem Erismann, J.: Validation of the CrIS fast physical NH<sub>3</sub> retrieval with ground-based FTIR, *Atmos. Meas. Tech.*, 10, 2645–2667, <https://doi.org/10.5194/amt-10-2645-2017>, 2017.
- EPA, United States Environmental Protection Agency, Air Quality Implementation Plans, <https://www.epa.gov/air-quality-implementation-plans>, last access: January 2023.
- 635 Ellis, R. A., Jacob, D. J., Sulprizio, M. P., Zhang, L., Holmes, C. D., Schichtel, B. A., Blett, T., Porter, E., Pardo, L. H., and Lynch, J. A. Present and future nitrogen deposition to national parks in the united states: Critical load exceedances. *Atmos. Chem. Phys.*, 13, 9083–9095. <https://doi.org/10.5194/ACP-13-9083-2013>, 2013
- Erismann, J. W., Sutton, M. A., Galloway, J., Klimont, Z., and Winiwarter, W.: How a century of ammonia synthesis changed the world, *Nat. Geosci.*, 1, 636–639, <https://doi.org/10.1038/ngeo325>, 2008.
- 640 Fehsenfeld, F. C., Huey, L. G., Leibrock, E., Dissly, R., Williams, E., Ryerson, T. B., Norton, R., Sueper, D. T., and Hartsell, B.: Results from an informal intercomparison of ammonia measurement techniques, *J. Geophys. Res. Atmos.*, 107, <https://doi.org/10.1029/2001JD001327>, 2002.
- Fountoukis, C. and Nenes, A.: ISORROPIAII: A computationally efficient thermodynamic equilibrium model for K<sup>+</sup>-Ca<sup>2+</sup>-Mg<sup>2+</sup>-NH<sub>4</sub><sup>+</sup>-Na<sup>+</sup>-SO<sub>4</sub><sup>2-</sup>-NO<sub>3</sub><sup>-</sup>-Cl<sup>-</sup>-H<sub>2</sub>O aerosols, *Atmos. Chem. Phys.*, 7, 4639–4659, <https://doi.org/10.5194/acp-7-4639-2007>, 2007.
- 645 Goldberg, D. L., Anenberg, S. C., Lu, Z., Streets, D. G., Lamsal, L. N., McDuffie, E., and Smith, S. J.: Urban NO<sub>x</sub> emissions around the world declined faster than anticipated between 2005 and 2019, *Environ. Res. Lett.*, 16, 115004, <https://doi.org/10.1088/1748-9326/AC2C34>, 2021.
- 650 Golston, L. M., Pan, D., Sun, K., Tao, L., Zondlo, M. A., Eilerman, S. J., Peischl, J., Neuman, J. A., and Floerchinger, C.: Variability of Ammonia and Methane Emissions from Animal Feeding Operations in Northeastern Colorado, *Environ. Sci. Technol.*, 54, 11015–11024, <https://doi.org/10.1021/acs.est.0c00301>, 2020.
- Gu, M., Pan, Y., Sun, Q., Walters, W. W., Song, L., and Fang, Y.: Is fertilization the dominant source of ammonia in the urban atmosphere, *Sci. Total Environ.*, 838, 155890, <https://doi.org/10.1016/J.SCITOTENV.2022.155890>, 2022.



- 655 Gunthe, S. S., Liu, P., Panda, U., Raj, S. S., Sharma, A., Darbyshire, E., Reyes-Villegas, E., Allan, J., Chen, Y., Wang, X., Song, S., Pöhlker, M. L., Shi, L., Wang, Y., Kommula, S. M., Liu, T., Ravikrishna, R., McFiggans, G., Mickley, L. J., Martin, S. T., Pöschl, U., Andreae, M. O., and Coe, H.: Enhanced aerosol particle growth sustained by high continental chlorine emission in India, *Nat. Geosci.*, 14, 77–84, <https://doi.org/10.1038/s41561-020-00677-x>, 2021.
- 660 Guo, X., Wang, R., Pan, D., Zondlo, M. A., Clarisse, L., van Damme, M., Whitburn, S., Coheur, P. F., Clerbaux, C., Franco, B., Golston, L. M., Wendt, L., Sun, K., Tao, L., Miller, D., Mikoviny, T., Müller, M., Wisthaler, A., Tevlin, A. G., Murphy, J. G., Nowak, J. B., Roscioli, J. R., Volkamer, R., Kille, N., Neuman, J. A., Eilerman, S. J., Crawford, J. H., Yacovitch, T. I., Barrick, J. D., and Scarino, A. J.: Validation of IASI Satellite Ammonia Observations at the Pixel Scale Using In Situ Vertical Profiles, *J. Geophys. Res. Atmos.*, 126, <https://doi.org/10.1029/2020JD033475>, 2021.
- 665 Hauglustaine, D. A., Balkanski, Y., and Schulz, M.: A global model simulation of present and future nitrate aerosols and their direct radiative forcing of climate, *Atmos. Chem. Phys.*, 14, 11031–11063, <https://doi.org/10.5194/acp-14-11031-2014>, 2014.
- Hennigan, C. J., Izumi, J., Sullivan, A. P., Weber, R. J., and Nenes, A.: A critical evaluation of proxy methods used to estimate the acidity of atmospheric particles, *Atmos. Chem. Phys.*, 15, 2775–2790, <https://doi.org/10.5194/acp-15-2775-2015>, 2015.
- 670 Hill, J., Goodkind, A., Tessum, C., Thakrar, S., Tilman, D., Polasky, S., Smith, T., Hunt, N., Mullins, K., Clark, M., and Marshall, J.: Air-quality-related health damages of maize, *Nat. Sustain.*, 2, 397–403, <https://doi.org/10.1038/s41893-019-0261-y>, 2019.
- Holt, J., Selin, N. E., and Solomon, S.: Changes in inorganic fine particulate matter sensitivities to precursors due to large-scale us emissions reductions, *Environ Sci Technol*, 49, 4834–4841, <https://doi.org/10.1021/acs.est.5b00008>, 2015.
- 675 Kendall, M.: Rank correlation methods (4th edn.) charles griffin. San Francisco, CA, 1975.
- Kharol, S. K., Shephard, M. W., McLinden, C. A., Zhang, L., Sioris, C. E., O'Brien, J. M., Vet, R., Cady-Pereira, K. E., Hare, E., Siemons, J., and Krotkov, N. A.: Dry Deposition of Reactive Nitrogen From Satellite Observations of Ammonia and Nitrogen Dioxide Over North America, *Geophys. Res. Lett.*, 45, 1157–1166, <https://doi.org/10.1002/2017GL075832>, 2018.
- 680 Lawal, A. S., Guan, X., Liu, C., Henneman, L. R. F., Vasilakos, P., Bhogineni, V., Weber, R. J., Nenes, A., and Russell, A. G.: Linked Response of Aerosol Acidity and Ammonia to SO<sub>2</sub> and NO<sub>x</sub> Emissions Reductions in the United States, *Environ. Sci. Technol.*, <https://doi.org/10.1021/acs.est.8b00711>, 2018.
- Li, Y., Schichtel, B. A., Walker, J. T., Schwede, D. B., Chen, X., Lehmann, C. M. B., Puchalski, M. A., Gay, D. A., and Collett, J. L.: Increasing importance of deposition of reduced nitrogen in the United States, *Proc. Natl. Acad. Sci. U.S.A.*, 113, 5874–5879, <https://doi.org/10.1073/pnas.1525736113>, 2016.
- 685 Lindaas, J., Pollack, I. B., Garofalo, L. A., Pothier, M. A., Farmer, D. K., Kreidenweis, S. M., Campos, T. L., Flocke, F., Weinheimer, A. J., Montzka, D. D., Tyndall, G. S., Palm, B. B., Peng, Q., Thornton, J. A., Permar, W., Wielgasz, C., Hu, L., Ottmar, R. D., Restaino, J. C., Hudak, A. T., Ku, I. T., Zhou, Y., Sive, B. C., Sullivan, A., Collett, J. L., and



- 690 Fischer, E. v.: Emissions of Reactive Nitrogen From Western U.S. Wildfires During Summer 2018, *Journal of Geophysical Research: Atmospheres*, 126, <https://doi.org/10.1029/2020JD032657>, 2021a.
- Lindaas, J., Pollack, I. B., Calahorrano, J. J., O'Dell, K., Garofalo, L. A., Pothier, M. A., Farmer, D. K., Kreidenweis, S. M., Campos, T., Flocke, F., Weinheimer, A. J., Montzka, D. D., Tyndall, G. S., Apel, E. C., Hills, A. J., Hornbrook, R. S., Palm, B. B., Peng, Q., Thornton, J. A., Permar, W., Wielgasz, C., Hu, L., Pierce, J. R., Collett, J. L., Sullivan, A. P., and Fischer, E. v.: Empirical Insights Into the Fate of Ammonia in Western U.S. Wildfire Smoke Plumes, *Journal of Geophysical Research: Atmospheres*, 126, <https://doi.org/10.1029/2020JD033730>, 2021b.
- 695 Malm, W. C., Schichtel, B. A., Pitchford, M. L., Ashbaugh, L. L., and Eldred, R. A.: Spatial and monthly trends in speciated fine particle concentration in the United States, *J. Geophys. Res. Atmos.*, 109, n/a-n/a, <https://doi.org/10.1029/2003JD003739>, 2004.
- Miller, D. J., Sun, K., Tao, L., Pan, D., Zondlo, M. A., Nowak, J. B., Liu, Z., Diskin, G., Sachse, G., Beyersdorf, A., Ferrare, R., and Scarino, A. J.: Ammonia and methane dairy emission plumes in the San Joaquin valley of California from individual feedlot to regional scales, *J. Geophys. Res. Atmos.*, 120, 9718–9738, <https://doi.org/10.1002/2015JD023241>, 2015.
- 700 McHale, M. R., Ludtke, A. S., Wetherbee, G. A., Burns, D. A., Nilles, M. A., and Finkelstein, J. S.: Trends in precipitation chemistry across the U.S. 1985–2017: Quantifying the benefits from 30 years of Clean Air Act amendment regulation, *Atmos. Environ.*, 247, <https://doi.org/10.1016/J.ATMOSENV.2021.118219>, 2021.
- 705 NADP, National Atmospheric Deposition Program, the Ammonia Monitoring Network, <https://nadp.slh.wisc.edu/networks/ammonia-monitoring-network/>, last accessed: January 2023.
- Nair, A. A. and Yu, F.: Quantification of atmospheric ammonia concentrations: A review of its measurement and modeling, <https://doi.org/10.3390/atmos11101092>, 1 October 2020.
- 710 Nair, A. A., Yu, F., and Luo, G.: Spatioseasonal Variations of Atmospheric Ammonia Concentrations Over the United States: Comprehensive Model-Observation Comparison, *J. Geophys. Res. Atmos.*, 124, 6571–6582, <https://doi.org/10.1029/2018JD030057>, 2019.
- Pan, D., Mauzerall, D. L., Benedict, K. B., Wang, R., Golston, L., Collett, J. L., Jr., Tao, L., Sun, K., Guo, X., Schichtel, B. A., Ham, J. M., Prenni, A. J., Puchalski, M., Mikoviny, T., Müller, M., Wisthaler, A., and Zondlo, M. A.: A Paradigm Shift in Sulfate-Nitrate-Ammonium Aerosol Formation in the United States and its Implications for Reactive Nitrogen Deposition, *American Geophysical Union Fall Meeting 2020*, Online, 1-17 Dec. 2020, A074-06, <https://agu.confex.com/agu/fm20/meetingapp.cgi/Paper/679051>, 2020.
- 715 Pan, D., Benedict, K. B., Golston, L. M., Wang, R., Collett, J. L., Tao, L., Sun, K., Guo, X., Ham, J., Prenni, A. J., Schichtel, B. A., Mikoviny, T., Mü, M., Wisthaler, A., and Zondlo, M. A.: Ammonia Dry Deposition in an Alpine Ecosystem Traced to Agricultural Emission Hotspots, *Environ. Sci. Technol.*, 55, 7785, <https://doi.org/10.1021/acs.est.0c05749>, 2021.
- 720



- Paulot, F., Jacob, D. J., Pinder, R. W., Bash, J. O., Travis, K., and Henze, D. K.: Ammonia emissions in the United States, European Union, and China derived by high-resolution inversion of ammonium wet deposition data: Interpretation with a new agricultural emissions inventory (MASAGE\_NH3), *J. Geophys. Res. Atmos.*, 119, 4343–4364, <https://doi.org/10.1002/2013JD021130>, 2014.
- 725 Phoenix, G. K., Hicks, W. K., Cinderby, S., Kuylenstierna, J. C. I., Stock, W. D., Dentener, F. J., Giller, K. E., Austin, A. T., Lefroy, R. D. B., Gimeno, B. S., Ashmore, M. R., and Ineson, P.: Atmospheric nitrogen deposition in world biodiversity hotspots: The need for a greater global perspective in assessing N deposition impacts, *Glob. Chang. Biol.*, 12, 470–476, <https://doi.org/10.1111/j.1365-2486.2006.01104.x>, 2006.
- 730 Pinder, R. W., Gilliland, A. B., and Dennis, R. L.: Environmental impact of atmospheric NH<sub>3</sub> emissions under present and future conditions in the eastern United States, *Geophys. Res. Lett.*, 35, 1–6, <https://doi.org/10.1029/2008GL033732>, 2008.
- Puchalski, M. A., Sather, M. E., Walker, J. T., Lehmann, C. M. B., Gay, D. A., Mathew, J., and Robarge, W. P.: Passive ammonia monitoring in the United States: Comparing three different sampling devices, *Journal of Environmental Monitoring*, 13, 3156–3167, <https://doi.org/10.1039/C1EM10553A>, 2011.
- 735 Puchalski, M. A., Rogers, C. M., Baumgardner, R., Mishoe, K. P., Price, G., Smith, M. J., Watkins, N., and Lehmann, C. M.: A statistical comparison of active and passive ammonia measurements collected at Clean Air Status and Trends Network (CASTNET) sites, *Environ. Sci.: Process. Impacts*, <https://doi.org/10.1039/c4em00531g>, 2015.
- Schiferl, L. D., Heald, C. L., Damme, M. van, Clarisse, L., Clerbaux, C., Coheur, P., Nowak, J. B., Neuman, J. A., Herndon, S. C., Roscioli, J. R., and Eilerman, S. J.: Interannual variability of ammonia concentrations over the United States: sources and implications, *Atmos. Chem. Phys.*, 12, 12305–12328, <https://doi.org/10.5194/acp-16-12305-2016>, 2016.
- 740 Shah, V., Jaeglé, L., Thornton, J. A., Lopez-Hilfiker, F. D., Lee, B. H., Schroder, J. C., Campuzano-Jost, P., Jimenez, J. L., Guo, H., Sullivan, A. P., Weber, R. J., Green, J. R., Fiddler, M. N., Bililign, S., Campos, T. L., Stell, M., Weinheimer, A. J., Montzka, D. D., and Brown, S. S.: Chemical feedbacks weaken the wintertime response of particulate sulfate and nitrate to emissions reductions over the eastern United States, *Proc. Natl. Acad. Sci. U.S.A.*, <https://doi.org/10.1073/pnas.1803295115>, 2018.
- 745 Sun, K., Cady-Pereira, K., Miller, D. J., Tao, L., Zondlo, M. A., Nowak, J. B., Neuman, J. A., Mikoviny, T., Müller, M., Wisthaler, A., Scarino, A. J., and Hostetler, C. A.: Validation of TES ammonia observations at the single pixel scale in the San Joaquin Valley during DISCOVER-AQ, *J. Geophys. Res. Atmos.*, 120, 5140–5154, <https://doi.org/10.1002/2014JD022846>, 2015.
- 750 Sun, K., Tao, L., Miller, D. J., Pan, D., Golston, L. M., Zondlo, M. A., Griffin, R. J., Wallace, H. W., Leong, Y. J., Yang, M. M., Zhang, Y., Mauzerall, D. L., and Zhu, T.: Vehicle Emissions as an Important Urban Ammonia Source in the United States and China, *Environ. Sci. Technol.*, 51, 2472–2481, <https://doi.org/10.1021/acs.est.6b02805>, 2017.
- 755 Sun, K., Zhu, L., Cady-Pereira, K., Chan Miller, C., Chance, K., Clarisse, L., Coheur, P. F., González Abad, G., Huang, G., Liu, X., van Damme, M., Yang, K., and Zondlo, M.: A physics-based approach to oversample multi-satellite,



multispecies observations to a common grid, *Atmos. Meas. Tech.*, 11, 6679–6701, <https://doi.org/10.5194/amt-11-6679-2018>, 2018.

760 Van Damme, M., Clarisse, L., Dammers, E., Liu, X., Nowak, J. B., Clerbaux, C., Flechard, C. R., Galy-Lacaux, C., Xu, W., Neuman, J. A., Tang, Y. S., Sutton, M. A., Erisman, J. W., and Coheur, P. F.: Towards validation of ammonia ( $\text{NH}_3$ ) measurements from the IASI satellite, *Atmos. Meas. Tech.*, 8, 1575–1591, <https://doi.org/10.5194/amt-8-1575-2015>, 2015.

Van Damme, M., Clarisse, L., Whitburn, S., Hadji-Lazaro, J., Hurtmans, D., Clerbaux, C., and Coheur, P. F.: Industrial and agricultural ammonia point sources exposed, *Nature*, 564, 99–103, <https://doi.org/10.1038/s41586-018-0747-1>, 2018.

765 Van Damme, M., Whitburn, S., Clarisse, L., Clerbaux, C., Hurtmans, D., and Coheur, P. F.: Version 2 of the IASI  $\text{NH}_3$  neural network retrieval algorithm: Near-real-time and reanalysed datasets, *Atmos. Meas. Tech.*, 10, 4905–4914, <https://doi.org/10.5194/amt-10-4905-2017>, 2017.

770 Van Damme, M., Clarisse, L., Franco, B., Sutton, M. A., Erisman, J. W., Wichink Kruit, R., van Zanten, M., Whitburn, S., Hadji-Lazaro, J., Hurtmans, D., Clerbaux, C., and Coheur, P. F.: Global, regional and national trends of atmospheric ammonia derived from a decadal (2008–2018) satellite record, *Environ. Res. Lett.*, <https://doi.org/10.1088/1748-9326/abd5e0>, 1 May 2021.

Walker, J. M., Philip, S., Martin, R. v., and Seinfeld, J. H.: Simulation of nitrate, sulfate, and ammonium aerosols over the United States, *Atmos. Chem. Phys.*, 12, 11213–11227, <https://doi.org/10.5194/acp-12-11213-2012>, 2012.

775 Wang, M., Kong, W., Marten, R., He, X. C., Chen, D., Pfeifer, J., Heitto, A., Kontkanen, J., Dada, L., Kürten, A., Yli-Juuti, T., Manninen, H. E., Amanatidis, S., Amorim, A., Baalbaki, R., Baccharini, A., Bell, D. M., Bertozzi, B., Bräkling, S., Brilke, S., Murillo, L. C., Chiu, R., Chu, B., de Menezes, L. P., Duplissy, J., Finkenzeller, H., Carracedo, L. G., Granzin, M., Guida, R., Hansel, A., Hofbauer, V., Krechmer, J., Lehtipalo, K., Lamkaddam, H., Lampimäki, M., Lee, C. P., Makhmutov, V., Marie, G., Mathot, S., Mauldin, R. L., Mentler, B., Müller, T., Onnela, A., Partoll, E., Petäjä, T., Philippov, M., Pospisilova, V., Ranjithkumar, A., Rissanen, M., Rörup, B., Scholz, W., Shen, J., Simon, M., Sipilä, M., Steiner, G., Stolzenburg, D., Tham, Y. J., Tomé, A., Wagner, A. C., Wang, D. S., Wang, Y., Weber, S. K.,  
780 Winkler, P. M., Wlasits, P. J., Wu, Y., Xiao, M., Ye, Q., Zauner-Wieczorek, M., Zhou, X., Volkamer, R., Riipinen, I., Dommen, J., Curtius, J., Baltensperger, U., Kulmala, M., Worsnop, D. R., Kirkby, J., Seinfeld, J. H., El-Haddad, I., Flagan, R. C., and Donahue, N. M.: Rapid growth of new atmospheric particles by nitric acid and ammonia condensation, *Nature*, 581, 184–189, <https://doi.org/10.1038/s41586-020-2270-4>, 2020.

785 Wang, R., Guo, X., Pan, D., Kelly, J. T., Bash, J. O., Sun, K., Paulot, F., Clarisse, L., van Damme, M., Whitburn, S., Coheur, P. F., Clerbaux, C., and Zondlo, M. A.: Monthly Patterns of Ammonia Over the Contiguous United States at 2-km Resolution, *Geophys. Res. Lett.*, 48, <https://doi.org/10.1029/2020GL090579>, 2021.

Warner, J. X., Wei, Z., Larrabee Strow, L., Dickerson, R. R., and Nowak, J. B.: The global tropospheric ammonia distribution as seen in the 13-year AIRS measurement record, *Atmos. Chem. Phys.*, 16, 5467–5479, <https://doi.org/10.5194/acp-16-5467-2016>, 2016.





- 790 Warner, J. X., Dickerson, R. R., Wei, Z., Strow, L. L., Wang, Y., and Liang, Q.: Increased atmospheric ammonia over the world's major agricultural areas detected from space, *Geophys. Res. Lett.*, 44, 2875–2884, <https://doi.org/10.1002/2016GL072305>, 2017.
- Whitburn, S., van Damme, M., Clarisse, L., Bauduin, S., Heald, C. L., Hadji-Lazarou, J., Hurtmans, D., Zondlo, M. A., Clerbaux, C., and Coheur, P.-F.: A flexible and robust neural network IASI-NH 3 retrieval algorithm, *J. Geophys. Res. Atmos.*, 121, 6581–6599, <https://doi.org/10.1002/2016JD024828>, 2016.
- 795 Yao, X. and Zhang, L.: Trends in atmospheric ammonia at urban, rural, and remote sites across North America, *Atmos. Chem. Phys.*, 16, 11465–11475, <https://doi.org/10.5194/acp-16-11465-2016>, 2016.
- Yao, X. and Zhang, L.: Causes of Large Increases in Atmospheric Ammonia in the Last Decade across North America, *ACS Omega*, 4, 22133–22142, <https://doi.org/10.1021/acsomega.9b03284>, 2019.
- 800 Yu, F., Nair, A. A., and Luo, G.: Long-Term Trend of Gaseous Ammonia Over the United States: Modeling and Comparison With Observations, *J. Geophys. Res. Atmos.*, 123, 8315–8325, <https://doi.org/10.1029/2018JD028412>, 2018.
- Yue, S. and Wang, C. Y.: The Mann-Kendall test modified by effective sample size to detect trend in serially correlated hydrological series, *Water Resources Management*, <https://doi.org/10.1023/B:WARM.0000043140.61082.60>, 2004.
- Zhai, S., Jacob, D. J., Wang, X., Liu, Z., Wen, T., Shah, V., Li, K., Moch, J. M., Bates, K. H., Song, S., Shen, L., Zhang, Y., 805 Luo, G., Yu, F., Sun, Y., Wang, L., Qi, M., Tao, J., Gui, K., Xu, H., Zhang, Q., Zhao, T., Wang, Y., Lee, H. C., Choi, H., and Liao, H.: Control of particulate nitrate air pollution in China, *Nat. Geosci.*, 14, 389–395, <https://doi.org/10.1038/s41561-021-00726-z>, 2021.
- Zheng, G., Su, H., Wang, S., Andreae, M. O., Pöschl, U., and Cheng, Y.: Multiphase buffer theory explains contrasts in atmospheric aerosol acidity, *Science (1979)*, 369, 1374–1377, <https://doi.org/10.1126/SCIENCE.ABA3719>, 2020.