



An overview of organic aerosols at an urban site in Hong Kong: insights from in-situ measurement of molecular markers

Hongyong Li^{1,2}, Xiaopu Lyu^{2*}, Likun Xue^{1*}, Yunxi Huo³, Dawen Yao³, Haoxian Lu⁴, Hai Guo^{3*}

¹Environment Research Institute, Shandong University, Qingdao, 266237, China

5 ²Department of Geography, Faculty of Social Sciences, Hong Kong Baptist University, Hong Kong

³Department of Civil and Environmental Engineering, The Hong Kong Polytechnic University, Hong Kong

⁴Marine Biological Resources Bank, Southern Marine Science and Engineering Guangdong Laboratory (Zhuhai), Guangdong, China

10 *Correspondence to:* Xiaopu Lyu (xiaopu_lyu@hkbu.edu.hk); Likun Xue (xuelikun@sdu.edu.cn); Hai Guo (hai.guo@polyu.edu.hk)

Abstract. Organic aerosol (OA) is a significant constituent of urban particulate matter (PM), and molecular markers therein provide information on sources and formation mechanisms of OA. With in-situ measurement of over 70 OA molecular markers at a bihourly resolution, this study focused on the temporal variations of representative markers and dynamic source contributions to OA at an urban site in Hong Kong. The levels of secondary OA (SOA) markers were markedly elevated in 15 continental and coastal air, and the primary markers were more of local characteristics. The diurnal patterns of 2-methyltetrols differed between scenarios, and their aqueous formation at night seemed plausible, particularly in the presence of troughs. Seven unambiguous sources were identified for the organic matters in submicron PM (PM₁-OM). Despite an urban site, the SOA contribution ($49 \pm 8\%$), primarily anthropogenic, was significant. Anthropogenic SOA dominated in continental and coastal air and in early afternoon. Local cooking and vehicle emissions became predominant in the scenario of marine air 20 without troughs. Even averaged over the study period, cooking emissions contributed up to 40% to PM₁-OM in the early evening. The study highlighted the need to control regional anthropogenic SOA and local cooking emissions to mitigate PM pollution in Hong Kong.

Keywords: Organic aerosol; Molecular marker; Source apportionment; Cooking emissions; Anthropogenic secondary organic aerosol

25 1 Introduction

Atmospheric particulate matter (PM) has been a widespread environmental concern, due to its significant climate and health effects (Ramanathan et al., 2001; Poschl, 2005; Yang et al., 2018). Organic aerosols (OA) make up 20-90% of fine PM (PM_{2.5}) by mass concentration (Kanakidou et al., 2005; Jimenez et al., 2009). While the inorganic PM compositions are somewhat being effectively controlled, reducing OA remains a major challenge for further mitigation of PM_{2.5} pollution in China (Huang 30 et al., 2014; Zhang et al., 2019). This is closely related to the lag and inadequacy of our knowledge of OA, especially the



insufficient understanding of time-resolved and unambiguous identification of OA sources and SOA formation mechanisms. Therefore, it becomes increasingly important to improve the understanding of OA with new theories and techniques. Molecular information records history of OA, including sources and chemical evolutions. Filter-based technique is a widely used method for measuring molecular compositions of OA (Yu et al., 2011; Ding et al., 2015; Fu et al., 2016). It generally takes half a day to a few days to collect PM on filters, followed by solvent or thermal extraction of target compounds and chemical analysis. However, during the collection cycle of each sample, emissions of OA and its precursors are not necessarily consistent, and chemical transformation may occur over several generations. For example, cooking and vehicle emissions vary significantly over the course of a day (Sun et al., 2018; Yao et al., 2021). With much higher time resolutions, Aerosol Mass Spectrometry (AMS) was developed and has been extensively applied in OA research (Sun et al., 2018; Liu et al., 2019; Yao et al., 2022). Nevertheless, AMS cannot provide molecular information of OA due to the application of hard ionization, which makes it difficult to identify unambiguous OA sources with AMS data (Sun et al., 2011; Lee et al., 2013). Applied earlier in the United States (Williams et al., 2006, 2010), in-situ measurement of OA molecular markers based on thermal desorption has become an emerging technique for chemical characterization of OA in China, *e.g.*, Beijing (Ren et al., 2019), Shanghai (Li et al., 2020; Wang et al., 2020), and Hong Kong (Lyu et al., 2020; Wang et al., 2022).

Hong Kong (HK) is a metropolis in East Asia with a developed economy and a high population density. Although the PM_{2.5} level continued to decrease in recent decades, it is still far from the World Health Organization guideline value (5 µg m⁻³). Led by the tertiary industry, HK has few high-emission industries. Street-level air pollution has long been a concern due to vehicle emissions (Lee et al., 2006; Huang et al., 2014; Liu et al., 2017a). Besides, mounting evidence demonstrates the considerable contributions of cooking emissions to OA (Lee et al., 2015; Sun et al., 2016). On the other hand, SOA formation in HK may also be significant, as mediated by the subtropical climate and coastal features. For example, studies indicated that aqueous processes accelerated SOA formation in spring with high relative humidity in HK (Li et al., 2013). In addition to local emissions and chemistry, transboundary transport also increases the atmospheric load of OA in HK (Huang et al., 2014; Lyu et al., 2017). With filter-based techniques, a range of molecular compositions of OA, such as polycyclic aromatic hydrocarbons (PAHs) (Guo et al., 2003; Liao et al., 2020), alkanes (Yao et al., 2004; Yu et al., 2011), dicarboxylic acids (Ho et al., 2006; Hu et al., 2013), and anhydrosugars (Zhang et al., 2012; Ho et al., 2014), have been well studied in HK. Further, OA sources have also been explored by introducing the molecular markers into receptor models (Hu et al., 2010; Huang et al., 2014; Chow et al., 2022). Cooking emissions, vehicle emissions, biomass burning, and secondary formation were identified as the main sources of PM in Hong Kong. AMS has also been applied in HK, advancing our understanding of the dynamic features of OA and its sources (Lee et al., 2015; Xing et al., 2022). In recent years, Thermal desorption Aerosol Gas chromatography coupled with mass spectrometry (TAG) allowed in-situ measurement of OA molecular markers and facilitated understanding of dynamic source contributions to OA at two background sites in HK (Lyu et al., 2020; Wang et al., 2022). In brief, it was found that the concentrations of SOA tracers were significantly elevated in PM_{2.5} pollution episodes, and transboundary transport led to a rise in biomass burning contributions to OA. Despite the previous studies, there is a lack of knowledge on time-resolved and unambiguous sources of OA in urban areas of HK.



65 This study presents the temporal variations of OA and a range of molecular markers therein based on a continuous measurement for ~1 month at an urban site in HK with TAG and AMS. Diurnal patterns of typical markers are discussed to indicate the emission patterns and possible formation mechanisms of OA. Source apportionment with the time-resolved data reveals the dynamic source contributions. The study enhances our understanding of urban OA in HK at bihourly resolution and molecular level, and provides a reference for PM pollution mitigation by intervention in OA emissions and formation.

70 2 Methods

2.1 Sampling campaign

The observation experiment was carried out in the campus of the Hong Kong Polytechnic University from June 6th to 26th in 2019. The instruments were deployed in a laboratory on 11th floor with a height of ~45 m above ground level. The sampling inlet protruded ~1.5 m from the wall. As shown in Figure 1, the site was in a mixed residential and commercial area with
75 intensive traffic, which was also one of the areas with the highest concentration of restaurants in HK. Besides, it is worth noting that the nearby funeral parlors and temples within 1 km to the east of the site often conducted the burning of worship supplies, ritual candles, and incenses.

Despite an urban site, air pollution could also be influenced by regional transport. The 72-h backward trajectories of air masses reaching the site were calculated using the Hybrid Single Particle Lagrangian Integrated Trajectory model, which was driven
80 by the Global Data Assimilation System archive data at the horizontal resolution of $1^\circ \times 1^\circ$. The trajectories started from the location 50 m above the sampling site and were calculated every 2 hours. Based on their origins and paths, the air masses were classified into three types (Figure 1). Those from the South China Sea represented the marine air (blue trajectories in Figure 1) and accounted for ~87% of the total air masses. The rest included continental air (red) originating from and passing over the Mainland China (6%) and coastal air (orange) that arrived in HK along the coastlines of eastern and southern China (7%).
85 To analyze OA in small particles of interest, a submicron particulate matter (PM_{10}) cyclone was installed at the front of the inlet line made of pre-cleaned copper tubing. The cut-off size of $1\mu m$ also matched the design of the High Resolution – Time of Flight – Aerosol Mass Spectrometer (HR-ToF-AMS, referred to as AMS hereinafter) manufactured by Aerodyne Research Inc. The AMS couples size-resolved particle sampling and mass spectrometry analysis, enabling real-time measurement of non-refractory compositions of PM_{10} , including total organic matter (PM_{10} -OM), sulfate, nitrate, ammonium, and chloride. With
90 hard ionization and in absence of species separation, AMS only provides mass spectral fragmentation information of size-segregated particles that arrive at the ionizer at the same time. As an important addition, the molecular markers of PM_{10} -OM were measured every 2 hours using the first commercialized TAG jointly developed by Aerodyne Research Inc and University of California, Berkeley. The main difference between TAG and traditional gas chromatography – mass spectrometry (GC-MS) is that TAG automates sample collection, derivatization, transfer, and GC-MS analysis. Within each cycle, the air after the removal of gaseous organics by a charcoal denuder was drawn and impacted onto a collection and thermal desorption (CTD)
95 unit at a constant rate of ~10 L/min in the first 90 minutes. Afterwards, the CTD was heated according to a preset program. At



the same time, a helium flow saturated with N-Trimethylsilyl-N-methyl trifluoroacetamide was injected into the CTD to derivatize and purge the sample to a focusing trap (FT) that was maintained at room temperature. The FT allowed the volatiles to flow out and trapped the less-volatile compounds. Then, it was heated quickly, and the trapped substances were flushed to the chromatographic column head by helium. Lastly, the GC-MS analysis was performed, which however did not take any additional time, because it happened simultaneous with sample collection. The sample treatment, transfer, and some other steps took 30 minutes, thereby 2 hours for a complete cycle. With online derivatization and thermal desorption, the TAG was capable of analyzing a wide range of polar and nonpolar organic compounds. A mixture of deuterated compounds was used as internal standards (IS) to track and correct the changes in instrument sensitivity. More details about the AMS and TAG operations can be found in our previous studies (Lyu et al., 2020; Huo et al., 2022).

Besides, the mixing ratios of inorganic trace gases, including sulfur dioxide (SO₂), carbon monoxide (CO), nitric oxide (NO), nitrogen dioxide (NO₂), and ozone (O₃) were continuously measured in the sampling period. The instruments are listed in Table S1, and more operation and quality control details have been presented elsewhere (Guo et al., 2013). In this study, the average mixing ratios of SO₂, CO, NO, NO₂, and O₃ were 2.1 ± 0.1 ppbv (mean \pm 95% confidence interval, the same below), 282 ± 12 ppbv, 10.7 ± 1.9 ppbv, 20.2 ± 1.4 ppbv, and 9.3 ± 1.6 ppbv, respectively. We adopted the meteorological parameters measured by the Hong Kong Observatory at the King's Park Meteorological Station, which was ~870 m away from our sampling site. They include temperature (Temp), relative humidity (RH), wind speed (WS), wind direction (WD), and ultraviolet (UV). During the sampling period, the average Temp, RH, WS, and UV were 28.6 ± 0.2 °C, $82 \pm 1.1\%$, 2.4 ± 0.1 m/s, and 18.1 ± 2.8 W/m², respectively. The prevailing winds were roughly consistent with the origins of air masses indicated by the backward trajectories. Liquid water content (LWC) in PM₁ was calculated in the same way as that adopted by Yao et al. (2022), using the Extended Aerosol Inorganics Model with Temp, RH, and inorganic ions in PM₁ as inputs. It varied in the range of 0.2 – 62.0 µg/m³, with a mean of 7.5 ± 1.2 µg/m³.

2.2 Processing of TAG data

A total of 225 bi-hourly samples were analyzed by TAG throughout the experiment, and 74 OA markers in the volatility range equivalent to C₁₆ – C₃₆ *n*-alkanes were identified based on their retention time and mass spectrums, especially the base ion (Lyu et al., 2020), as shown in Table S2. In brief, we identified a species from the TAG chromatogram based on the retention time of the base ion of authentic standard and confirmed it by searching the ion fragments in an expanded NIST mass spectral library. For a few samples, the fatty acids and oleic acid were not fully derivatized, likely due to their exceptionally high concentrations in cooking hours. Peak fitting and integration were performed to both the derivatized and underivatized forms of them. The peak areas of the target compounds were scaled by those of the corresponding ISs with similar structures and close retention time. The pairing of target compounds and ISs are the same as that adopted in Lyu et al. (2020). In this way, the temporal variations of the 74 OA markers were not related to the changes in instrument sensitivity. However, we ran out of the external standards in this sampling campaign, so the concentrations of the detected species were not quantified. Instead, the internal standard scaled peak areas (IS-scaled PA) of the target compounds are used for analysis in this study.



130 For ease of analysis, we first classified the detected compounds based on their common origins and chemical structures. Further, correlation analysis was performed among them, as shown in Figure 2. Moderate to good correlations were identified for some species within 7 groups, including common SOA tracers, dicarboxylic acids (DCA) & hydroxyl dicarboxylic acids (OHDCA), sugars & sugar alcohols, *n*-alkanes, fatty acids, aromatic acids, and PAHs. The coefficient of determination (R^2) was higher than 0.65 for the correlations between any pairs of 4-/5-ring PAHs. The variations of levoglucosan and mannosan
135 were roughly consistent ($R^2 = 0.69$). Good correlations were identified within the subgroups of high molecular weight ($C_{25} - C_{30}$) and low molecular weight ($C_{20} - C_{24}$) *n*-alkanes, separately, but not between them. This likely resulted from the different sources of them, *e.g.*, vegetable wax and debris for the former and petroleum products for the latter (Kawamura et al., 2003; Simoneit et al., 2004; Kang et al., 2016a). Consistent variations were observed for OHDCA species. Some DCAs, such as succinic acid which is a known precursor of malic acid (Hu et al., 2013), also correlated well with OHDCA. The patterns of
140 isoprene SOA tracers formed in low NO_x environment, *i.e.*, 2-methyltetrols (2-MTs) and C_5 -alkenetriols (Claeys et al., 2004; Edney et al., 2005; Surratt et al., 2010), resembled each other ($R^2 > 0.74$). Besides, there also existed some correlations between species of different groups, mainly due to co-emissions from cooking (oleic acid, fatty acids, fructose isomers, and azelaic acid) and biomass burning (vanillic acid (VA), 4-hydroxybenzoic acid (4-OHBA), and levoglucosan). Based on the correlations, we selected 18 OA markers as representatives of the whole in the following analyses, as highlighted in Figure 2
145 and Table S2. All the chemical and meteorological data with time resolution higher than TAG data were converted to bihourly averages for matching purposes where necessary.

2.3 Application of Positive Matrix Factorization (PMF) model

PMF has been widely used as a quantitative tool to understand air pollution sources (Hu et al., 2010; Wang et al., 2018; Dai et al., 2020). In this study, the EPA PMF v5.0 was used to perform the source apportionment for PM_{1-OM} . A total of 13 OA
150 markers measured by TAG were used to indicate the sources, including 3 OHDCA species (MA, CMA, and 2-HGA, see Table S2 for the full names), 2 anthropogenic SOA tracers (DHOPA and phthalic acid), 3 isoprene SOA tracers (2-MT1, cis-2-MBT, and 2-MGA), and 5 primary organic aerosol (POA) species (levoglucosan, fructose-2, palmitic acid, norhopane, and oleic acid). The uncertainty file was compiled following the method described in Lyu et al. (2020), based on an error fraction of 10% and method detection limit (MDL) for individual species. The MDL for PM_{1-OM} was determined to be $0.201 \mu\text{g}/\text{m}^3$. The
155 MDLs of OA species were taken from Lyu et al., (2020) that used the same TAG ~8 months earlier. For the newly detected species, *e.g.*, phthalic acid, 2-MGA, cis-2-MBT, and norhopane, we assume 10% of the averages as the MDLs. Extreme values are generally difficult to be reproduced by PMF model, especially for data points with scaled residuals greater than 4. To address this, greater weight was given to the high values that are considered true in the PMF iteration process, as described by Wang et al. (2018).

160 The optimal number of PMF factors was determined according to the mathematical tests and physical interpretability. We examined the solutions with 3 – 10 factors and the corresponding Q/Q_{exp} ratios (Figure S1). The ratio decreased rapidly for 3 – 7 factors and the decline slowed down for 7 – 10 factors, demonstrating a minimum of 7 factors from a mathematical



perspective. Further inspections found that some tracers indicative of different sources (*e.g.*, DHOPA and MA) were not separated in the 7-factor solution, and some species were not well reproduced. The 9-factor solution however apportioned 165 tracers of a same source into multiple factors. We ended up with 8 factors, which provided reasonably good fit for all species with physically interpretable source profiles.

3 Results and Discussion

3.1 Overview of the sampling campaign

Figure 3 and Figure S2 show the time series of the selected OA markers, PM₁ compositions, trace gases, and meteorological 170 conditions throughout the sampling period. The origins of air masses are also labelled. The total organics, sulfate, nitrate, and ammonium in PM₁ were observed with the average concentration of 4.2 ± 0.5 , 2.9 ± 0.3 , 0.3 ± 0.1 and $1.1 \pm 0.1 \mu\text{g}/\text{m}^3$, respectively. The fraction of PM₁-OM ($47.1 \pm 2.2\%$) was slightly higher than that at an urban background site in HK (43.8%), but lower than that at a roadside site ~350 m away (57.7%) (Yao et al., 2021,2022). PM₁-OM had the highest concentration in the continental air, followed by that in coastal air and marine air, same for CO and O₃ (Table S3). However, the highest levels 175 of sulfate and ammonium were observed in coastal air, likely due to aqueous aging of regional air pollution (Huo et al., 2023). Although the marine air was supposed to be relatively clean, it contained elevated levels of SO₂ and NO_x, which correlated well with each other ($R^2 = 0.76$). This phenomenon was in line with that at the three roadside sites in HK, but was not identified at the general urban sites (Figure S3). Therefore, the close relationship between SO₂ and NO_x on the days with marine air were likely attributable to local vehicle emissions rather than ship emissions. Moreover, the concentrations of many air pollutants, 180 including both primary and secondary species, increased significantly on June 10th – 14th and 24th – 26th, when HK was affected by the troughs of low pressure with cloudy weather and low wind speed. This might be related to the unfavorable dispersion conditions and enhanced chemistry under special meteorological conditions, such as high RH. The effects also applied to some OA markers, such as malic acid and its homologues.

The cooking-related OA markers, such as oleic acid, palmitic acid, and fructose isomers, did not show obvious inclination 185 toward the trough periods and exhibited some periodic spikes in the marine air. Therefore, it seemed that local cooking emissions somewhat overwhelmed the effects of atmospheric mixing and transformation. The lower levels of cooking markers in continental and coastal air could be partially explained by the quicker chemical losses, due to the higher concentrations of oxidants, *e.g.*, O₃ (Table S3). Oleic acid with a carbon-carbon double bond in the molecule can be efficiently oxidized by O₃ and other oxidants (Zeng et al., 2020; Wang et al., 2021). Differently, levoglucosan, pyrene, and some odd *n*-alkanes (*e.g.*, *n*- 190 C₂₇) with a few spikes were more abundant in coastal air. Further, we found that high levels of levoglucosan were associated with easterly winds (Figure S4). As stated above, biomass product combustions were common within 1 km east of the site (Figure 1). Therefore, the observed variations of levoglucosan were mainly attributed to local emission patterns, different from the regional transport of biomass burning plumes from mainland to HK in cool seasons (Sang et al., 2011; Chow et al., 2015).



195 The wind dependence was not identified for others, except that the strong southwest winds caused low levels of many OA markers due to the dilution by marine air.

The temporal variations of secondary OA markers were quite different. The abundances of 2-MTs (*e.g.*, 2-MT1, see Table S2), C₅-alkenetriols (*e.g.*, *cis*-2-MBT), and 2-MGA increased significantly by 13.1, 18.4, and 13.8 times when the continental air arrived at the site. Consistent with that at an urban background site (Lyu et al., 2020), higher levels of 2-MTs (isoprene SOA in NO_x-lean environment) were observed at night and 2-MGA (isoprene SOA in NO_x-rich environment) exhibited the peak in
200 the afternoon. The temperature and UV were not exclusively high on the days with continental air, and we presume the same for biogenic emissions locally. Therefore, the notable levels of isoprene SOA tracers on these days suggested that the continental air transported biogenic SOA and/or the precursors to HK. Succinic acid correlated fairly well with 2-MGA in the continental air ($R^2 = 0.84$, compared to R^2 of 0.56 at other times), indicating that it was a degradation product of isoprene under high-NO_x conditions. Besides, the first day with continental air also experienced highest levels of DHOPA and phthalic acid,
205 two anthropogenic SOA tracers derived from aromatics (Bunce et al., 1997; Jang et al., 1997; Kleindienst et al., 2004). This could also be related to regional transport, given the notable rise in other anthropogenic air pollutants (*e.g.*, CO). In addition, it was likely that the enhancement of atmospheric oxidation capacity, as indicated by the highest level of O₃ over the sampling campaign, facilitated the formation of many SOA species on this day. These factors might also play a role in maintaining the high levels of OHDCA, such as malic acid and tartaric acid, which were elevated during a low-pressure trough in the previous
210 days with marine air.

3.2 Diurnal patterns of representative OA markers

An advantage of TAG measurement is the time-resolved data of OA markers. Figure 4 presents the average diurnal profiles of 9 representative species, and some others with lower priorities are shown in Figure S5. As discussed above, the high levels of levoglucosan were related to the combustion of biomass products nearby. Here, we notice a small peak at 10:00 and a marked
215 increase during 18:00 – 20:00 for levoglucosan. The pattern, especially the early evening peak, roughly coincided with the timing of smokes from the funeral parlor chimneys, according to our observations. Some other species, including 4-OHBA, VA, mannosan, xylose, hydroxybutyric acid, and protocatechuic acid showed consistent diurnal patterns (Figure S5), indicating that they were mainly derived from biomass burning. Besides, C₂₆ – C₃₁ *n*-alkanes were observed with distinct peaks at 8:00 and 20:00 and might also be influenced by nearby sacrificial activities. Other sources, *e.g.*, vehicle emissions, might
220 also contribute to the morning peak that was more pronounced than levoglucosan. As an indicator of vehicle emissions, norhopane exhibited higher levels in daytime than at night, and were slightly elevated in the morning and evening rush hours. Similar diurnal patterns were observed for C₂₀ – C₂₃ *n*-alkanes (Figure S5), which could also be derived from vehicle emissions. Oleic acid was obviously enhanced at noon and in early evening, when intensive cooking activities took place in restaurants and at home, same for fatty acids, azelaic acid, and fructoses (Figure S5). Pyrene, as a representative of PAHs, got the highest
225 level at 20:00. Interestingly, it correlated well with levoglucosan ($R^2 = 0.88$) at this point of time (Figure S6) and exhibited no correlation with oleic acid which also had the highest levels in the early evening. At the other times excluding 14:00 – 16:00



when photochemical reactions might consume a large fraction of NO_x , the variation of pyrene was consistent with that of NO_x ($R^2 = 0.70$). Therefore, the exhaust fume from nearby funeral parlors were responsible for the peak values of PAHs in the early evening, while the lower levels of PAHs were likely associated with vehicle emissions.

230 Contrary to the common understanding, the 2-MTs showed higher levels between midnight and dawn, same for C_5 -alkenetriols. This however was in line with previous observations at two background sites in HK conducted by us and another team (Lyu et al., 2020; Wang et al., 2022). Day-by-day inspections found inconsistent diurnal patterns for 2-MTs, which were grouped into four categories, as shown in Figure 5. As speculated in previous studies (Lyu et al., 2020), the higher levels of 2-MTs in continental and coastal air might result from transboundary transport and diurnal evolution of boundary layer. On the days

235 with marine air, two distinct patterns were identified. Similar night time enhancement was observed in the presence of troughs. Interestingly, 2-MTs exhibited the highest levels in the afternoon without the influence of the troughs. These days were accompanied by higher temperature and UV intensity, but lower RH and LWC (Figure 5). Therefore, the afternoon peak was likely attributed to the more intensive emissions and photochemical oxidation of isoprene. Moreover, we uncovered some clues that nocturnal chemistry might regulate the diurnal patterns of 2-MTs. According to Noziere et al. (2011), the ratio of 2-

240 methylerythritol to the sum of 2-MTs (2-MT1/2-MTs) indicates the sources of 2-MTs: primary emissions (0.35), NO_x -lean photooxidation (0.61), NO_x -rich photooxidation (0.76), and aqueous phase oxidation (0.90). Under the assumption that the instrument sensitivity to the two 2-MT isomers were the same, the average 2-MT1/2-MTs ratio was 0.70 in the marine air without troughs. However, it was much higher (0.81) in presence of troughs, especially at night, in line with the patterns of LWC. Hence, the higher levels of 2-MTs at night might be related to aqueous chemistry. Nevertheless, it could not be the sole

245 reason for the counterintuitive patterns of 2-MTs in continental and coastal air, where the average 2-MT1/2-MTs ratio was 0.76 and 0.77, respectively. The influence of other factors, such as transport and atmospheric mixing, deserves study. Conversely, 2-MGA exhibited a typical diurnal pattern of photochemical pollution, increasing from morning to afternoon, except an exclusively high value at ~10:00 on one day with continental air. Daytime enhancements were also observed for DHOPA and phthalic acid, suggesting that photochemistry was involved in the formation of these SOA tracers. It is noteworthy

250 that the peak of phthalic acid appeared at 12:00, ~4 hours earlier than the peaks of 2-MGA and DHOPA, which might be due to differences in formation mechanisms and/or oxidation states. For example, Yao et al. (2022) found that the peak of more oxidized oxygenated OA was several hours later than that of less oxidized OA, and aqueous processes were suspected between the two peaks. Despite a slight rebound at 12:00 – 14:00, the diurnal variation of malic acid was small, much less pronounced than that at an urban background site (Lyu et al., 2020). This pattern also applied to the other OHDCAs species, which might

255 be partially attributed to the lower atmospheric oxidation capacity, as reflected by the difference in odd oxygen ($\text{O}_x = \text{O}_3 + \text{NO}_2$) levels, *i.e.*, 29.5 ppbv at this site *vs.* 49.8 ppbv at the background site.

3.3 Source apportionment of $\text{PM}_{1\text{-OM}}$

Figure 6 shows the source profiles of the 8 factors resolved by PMF. Levoglucosan was assigned to multiple factors in the base run, although the factor of biomass burning accounted for the highest fraction. Constraint was applied to reduce the



260 levoglucosan loadings in other factors, thereby lower rotational ambiguity. Moreover, we also performed simulations without
PM₁-OM that was measured by AMS. This was to examine if it was appropriate to mix the components measured by two
different instruments in the source apportionment. As shown, the source profiles did not change much between these
simulations. Table S4 summarizes the R², slope, and intercept for the linear regression between the simulated and observed
values of individual species, which are in reasonable ranges. The scaled residuals for all the species were within the range of
265 -3 to 3. These metrics indicated that the observed magnitudes and variations of the PMF species were well reproduced. To
evaluate the model stability and uncertainty of the results, Bootstrapping (BS) and displacement (DISP) analyses were
conducted. Based on BS tests, the factor mappings ranged from 85% to 100% except for an unresolved factor (77%), above
the threshold of 80% representing robust factors recommended by the PMF user guide. Namely, there was no observation that
disproportionately influenced the solution. Besides, no factor swap was found from the DISP analysis. The BS and DISP
270 results indicated the robustness of the 8-factor solution. The unresolved factor is discussed below.

The 8 factors were identified as cooking emissions, vehicle emissions, biomass burning, isoprene SOA, anthropogenic SOA
represented by phthalic acid (ASOA-1), anthropogenic SOA represented by DHOPA (ASOA-2), OHDCA-like SOA, and an
unresolved factor. Factor 1 was dominated by oleic acid, palmitic acid, and fructose isomers, indicating the source of cooking
emissions (To et al., 2000; Lyu et al., 2021; Huo et al., 2022). The second factor was regarded as vehicle emissions, due to the
275 high loading of norhopane (Zielinska et al., 2004). Factor 3 with the dominance of levoglucosan signified biomass burning
(Fabbri et al., 2009). Factor 4 was identified by 2-MT2 and cis-2-MBT, both of which were isoprene SOA tracers formed in
NO_x-lean environments (Claeys et al., 2004). We define the fifth and sixth factor as ASOA-1 and ASOA-2 based on the highest
percentage of phthalic acid and DHOPA, respectively, because phthalic acid and DHOPA have been shown to be oxidation
products of aromatics (Bunce et al., 1997; Jang et al., 1997; Kleindienst et al., 2004). Factor 7 was characterized by its high
280 contribution to MA and moderate contributions to CMA and 2-HGA, all of which were OHDCA species. Our previous studies
(Lyu et al., 2020; Huo et al., 2023) indicated that OHDCA in HK was attributed to photochemical and aqueous processes rather
than direct emissions, so this factor was termed as OHDCA-like SOA. With percentage contributions to all the species lower
than 35%, the last factor could not be recognized through any indicative species. It had the smallest diurnal variation with a
coefficient of variation of 0.75, compared to 1.17 – 4.42 for the other factors. Therefore, we presume this unresolved factor
285 might be related to the background residue. As discussed below, this factor only contributed 9 ± 1% to PM₁-OM.

Further, the source contributions to PM₁-OM were determined. As shown in Figure 7a, they varied noticeably over time,
underlining the value of source apportionment based on in-situ measurement of OA markers. Cooking emissions, vehicle
emissions, ASOA-2, OHDCA-like SOA and ASOA-1 were the main sources of PM₁-OM, with the average contribution of 19
± 3%, 18 ± 4%, 16 ± 4%, 15 ± 2%, 15 ± 3%, respectively. Biomass burning and isoprene SOA only accounted for 5 ± 2% and
290 4 ± 1% of PM₁-OM. This demonstrated the significant role of anthropogenic air pollution, either primary or secondary, in
contributing to PM₁-OM at this urban site. However, biogenic contributions might be somewhat underestimated, because i)
we failed to separate a source represented by 2-MGA, an isoprene SOA tracer formed in NO_x-rich environment; ii) OHDCA-
like SOA could be formed through the oxidation of biogenic precursors (Hu et al., 2013; Huo et al., 2023); and iii) SOA tracers



295 derived from monoterpenes were not detected. The dominant sources, especially cooking and vehicle emissions, were different from those at the urban background site (Lyu et al., 2020), likely due to the depletion of primary emissions during the course of their airborne migration to the background site.

On the days with continental air, vehicle emissions and ASOA-2 with the contribution to PM₁-OM of 31 ± 18% and 32 ± 8% were the most predominant, while only 1 ± 1% of PM₁-OM was attributed to cooking emissions (Figure 7b). It is noteworthy that anthropogenic SOA, sum of ASOA-1 and ASOA-2, accounted for almost half (47 ± 12%) of PM₁-OM. The peaks were in sync with that of CO, suggesting that the enhancement of anthropogenic SOA was more likely due to transboundary transport. Similarly, vehicle emissions and ASOA-2 were also the leading sources in coastal air, contributing 30 ± 12% and 28 ± 6% to PM₁-OM, respectively. However, in the marine air scenario, cooking emissions (26 ± 5%, source contribution, the same below) ranked the largest source of PM₁-OM, followed by ASOA-1 (17 ± 4%), OHDCA-like SOA (15 ± 3%), and vehicle emissions (13 ± 2%). As discussed in section 3.1, some OA markers were elevated under the influence of the low-
305 pressure troughs when backward trajectories still indicated marine air. Correspondingly, the contributions of ASOA-1 and OHDCA-like SOA to PM₁-OM increased to 26 ± 5% and 20 ± 3%, respectively. It was a bit surprising to note the rise in SOA contributions in these periods, because of the lower temperature and weaker UV (Figure 3). The growth of OHDCA-like SOA might be facilitated by the high RH and LWC through aqueous processes (Huo et al., 2023). While it is unknown whether this also caused the increase of ASOA-1, moderate correlation (R²=0.41) was identified between phthalic acid and the product of
310 RH and UV, which was not found at other times (Figure S7). Thus, the high RH might be one of the leading factors, if not all, of SOA enhancements in presence of troughs. In contrast, cooking emissions became the largest source of PM₁-OM (41 ± 9%) in the scenario of marine air without troughs, and in second place was vehicle emissions (14 ± 3%). This was more reflective of local emissions with less chemical aging. Overall, the SOA contributions increased significantly with the rise in PM₁-OM concentration (Figure S8), consistent with the findings at a suburban site that PM_{2.5} pollution events witnessed a notable rise
315 in SOA tracer levels (Wang et al., 2022). Moreover, the anthropogenic SOA might also be underestimated, similar to the situation with biogenic SOA. For example, cooking emissions were expected to be relatively consistent across days. The notable decline in continental and coastal air was likely a result of enhanced aging, due to the higher concentrations of oxidants, such as O₃ (Table S3). However, we have not identified a source of aged cooking emissions, although some fractions of it might have been allocated to the identified SOA factors, such as OHDCA-like SOA.

320 Furthermore, the diurnal variations of source contributions were also evident (Figure 7c). Cooking emissions were responsible for up to 40 ± 7% of PM₁-OM during 18:00 – 20:00. Almost at the same time, biomass burning also contributed 10 ± 5% to PM₁-OM, although it was on average a small source. Therefore, short-term exposure to emissions of these sources might be a concern, especially due to the fact that the highest levels of PAHs were related to biomass burning as discussed above. The total percentage contribution of SOA was relatively high and stable between 0:00 and 16:00, and decreased in the early evening
325 as the cooking and biomass burning emissions intensified. The special diurnal pattern of total SOA contribution was a result of superposition of 4 SOA factors with different diurnal patterns, as shown in Figure S9. It is worth mentioning that the contribution of anthropogenic SOA could reach 40 ± 29% in the early afternoon (12:00 – 14:00).



4 Conclusion

As a ubiquitous composition of atmospheric fine PM, OA has been extensively studied. Nevertheless, it is still a major obstacle to further improving PM air quality in many places around the world, and also a meeting point in control of air pollution complex, such as PM_{2.5} and O₃ in photochemical smog.

Taking advantage of an emerging technique for in-situ measurement of OA molecular markers, we reported by far the highest-resolution temporal variations of dozens of OA markers in a summer month in urban HK. The time-resolved data allowed to discuss the effects of transboundary transport and weather patterns (*e.g.*, troughs) on primary and secondary OA markers. DHOPA, phthalic acid, and isoprene SOA tracers were more enriched in the continental air, probably due to more abundant precursors and higher atmospheric oxidation capacity. The high RH brought on by troughs might have facilitated the formation of some OA species, such as OHDCA. Further, the diurnal variations of OA markers indicated emission patterns and SOA evolutions, such as the most intense cooking emissions in early evening and potentially enhanced 2-MTs formation in presence of troughs. Regarding the sources of PM₁-OM, we uncovered the dominance of cooking and vehicle emissions in the scenario of marine air without troughs, which was the cleanest case. Even as a whole, cooking emissions explained ~40% of the PM₁-OM in early evening. Anthropogenic SOA and vehicle emissions were the main sources of PM₁-OM in the continental and coastal air, and the SOA contributions, mainly from anthropogenic components, substantially increased when the PM air quality deteriorated. The findings demonstrated the need to properly manage local cooking and vehicle emissions and to control aromatics-derived SOA with concerted efforts on a regional or even superregional scale.

While this study provides valuable insights, we acknowledge there are several limitations. First, the molecular markers were not quantified, making it impossible to compare with the measurements elsewhere or compare among the detected species for absolute concentrations. It also hinders reuse of the data by other studies, such as concentration-based health risk assessment for some species. Second, some SOA tracers, such as those derived from monoterpenes and sesquiterpenes, were not detected. This increased the uncertainty of OA source apportionment to some extent. Third, we reiterated the unexpected diurnal patterns of 2-MTs, but the analysis of the underlying reasons is speculative. Therefore, future work will be dedicated to improving the analytical performance of TAG and exploring complex mechanisms responsible for unique features of SOA tracers, including 2-MTs.

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Author contributions

360 HL performed the study and drafted the manuscript. XL designed the study, oversaw interpretation of the results and revised the manuscript. YH, DY, and HL assisted in field campaign and processed relevant data. LX and HG led the project and revised the manuscript.

Data availability

All raw data are available upon request from the corresponding author Dr. Xiaopu Lyu.

Competing interests

365 The authors declare no competing interests.

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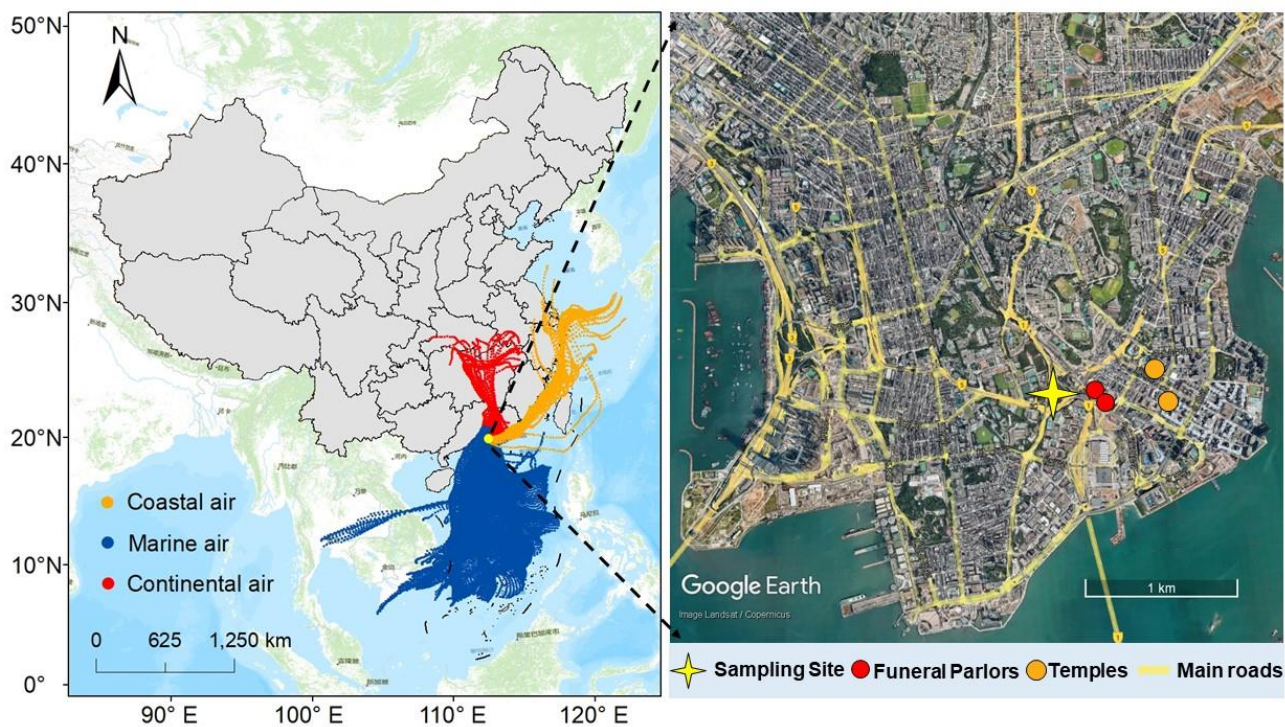


Figure 1: Location of the sampling site and the surrounding environment. Background maps are adapted from ArcGIS and © Google Earth.

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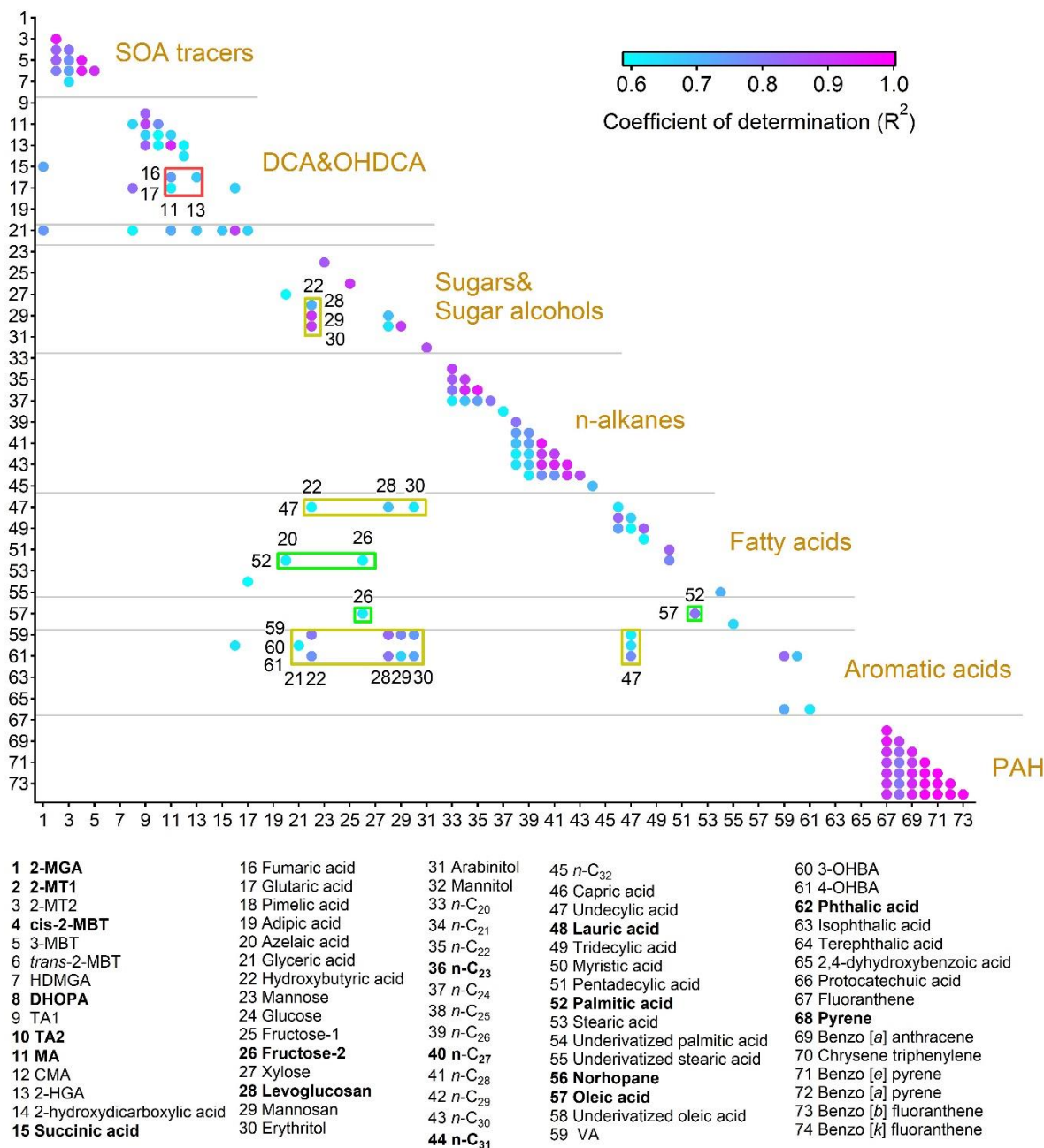
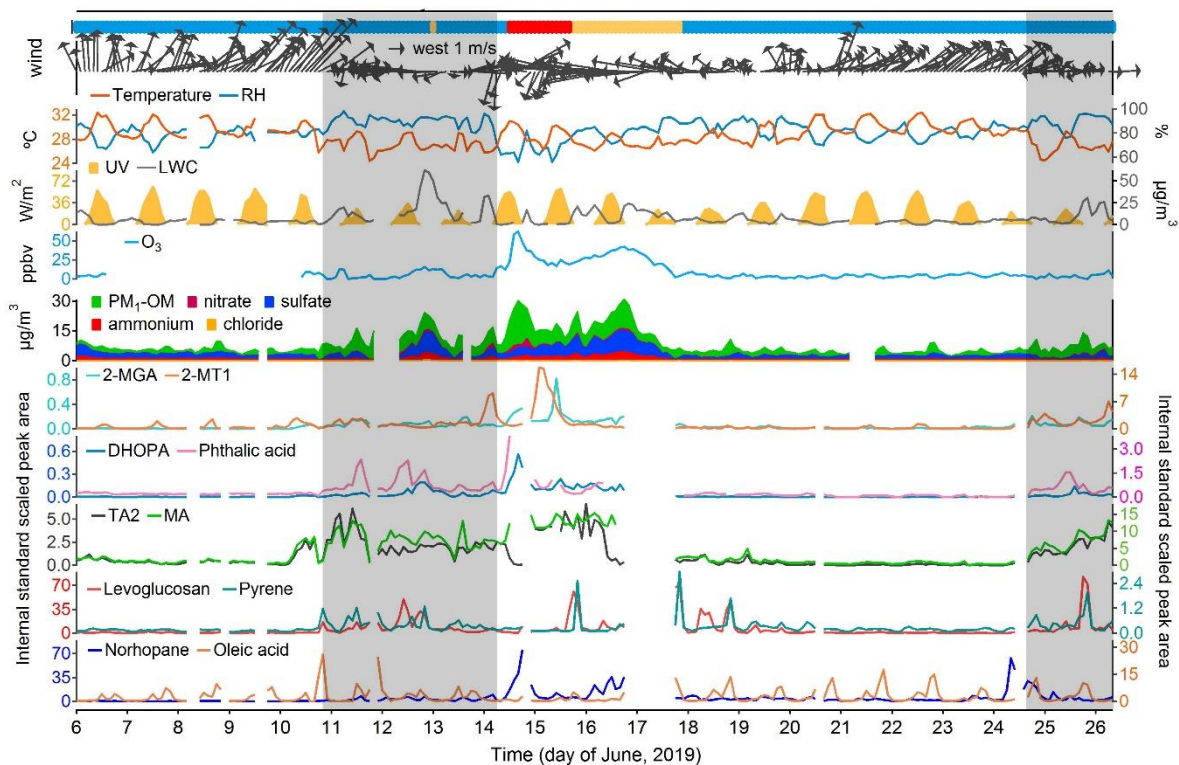
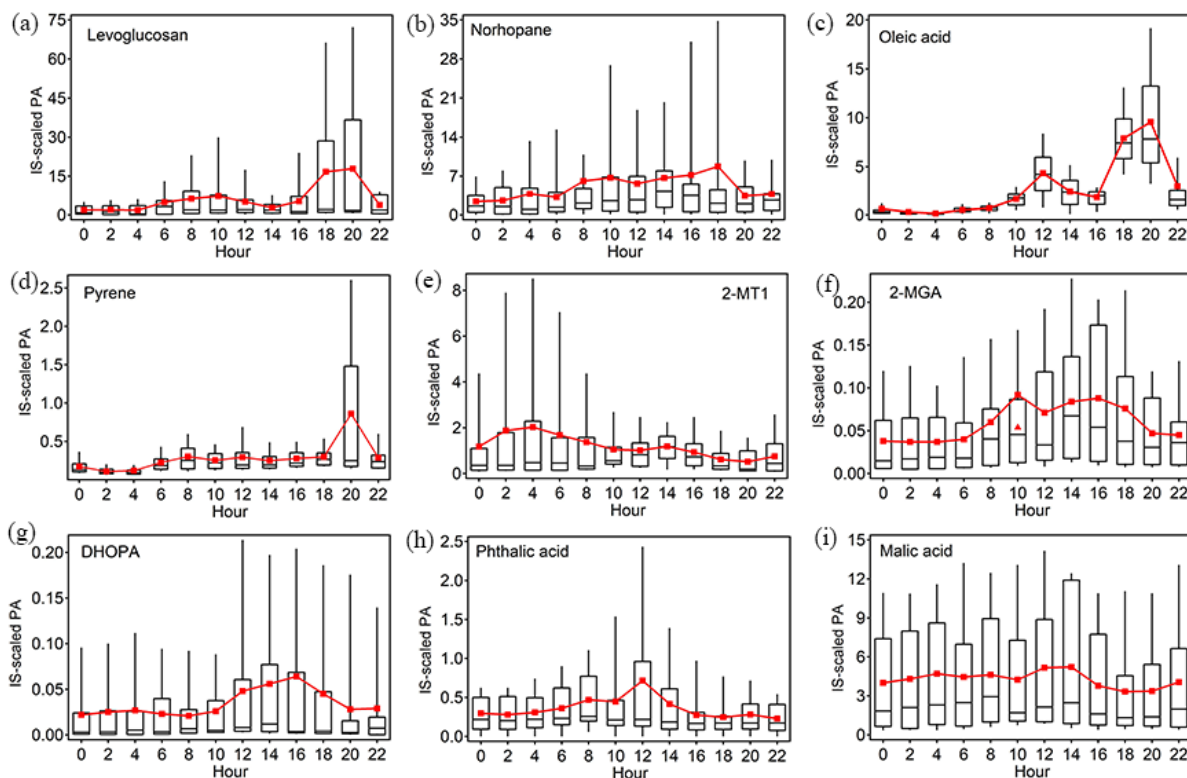


Figure 2: Correlations between OA molecular markers with R^2 not lower than 0.6. Bold are the representative species that we focus on in this study. Full names of the abbreviations are given in Table S2.



565 **Figure 3: Time series of representative OA markers, PM₁ compositions, trace gases, and meteorological conditions. Blue, orange, and red bars at the top represent marine, coastal, and continental air, respectively. Shaded areas represent the periods with troughs. Missing data are due to instrument maintenance.**



570 **Figure 4: Average diurnal patterns of nine representative OA markers. Tip of the top (bottom) whisker is the 95th (5th) percentile. Top and bottom of the box denote the 75th and 25th percentiles, respectively. Line inside the box represents the median value, and the mean value is indicated by red square and triangle (excluding data on June 14th – 15th for 2-MGA only).**

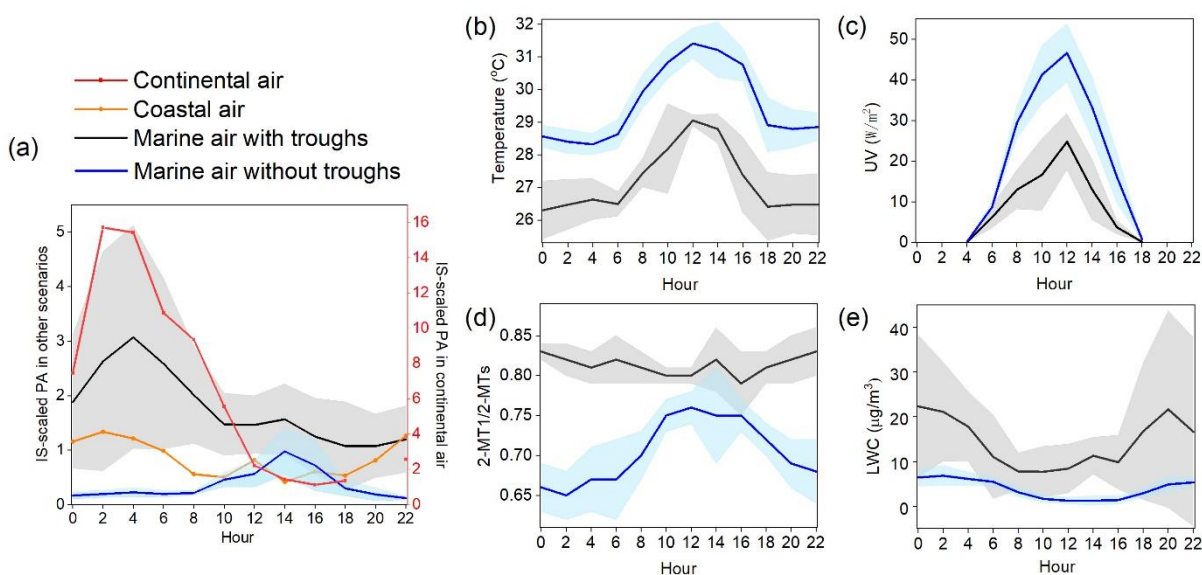
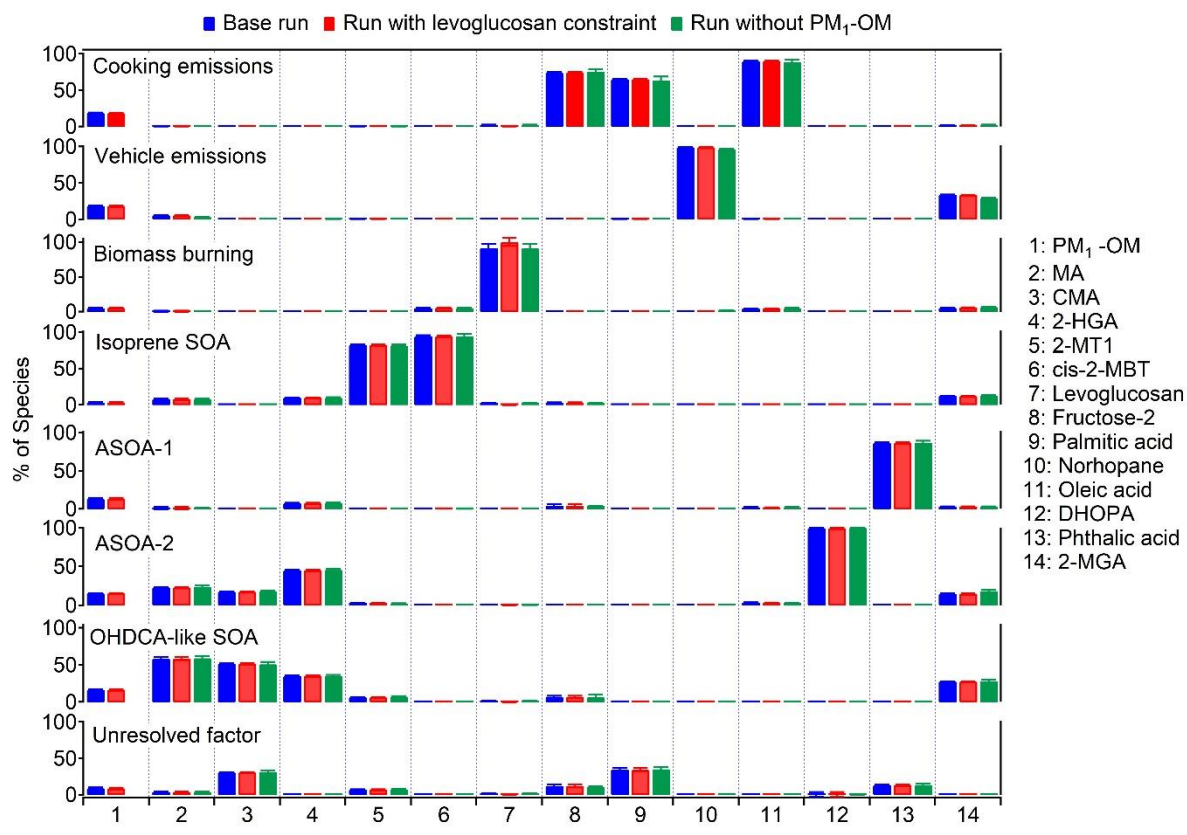


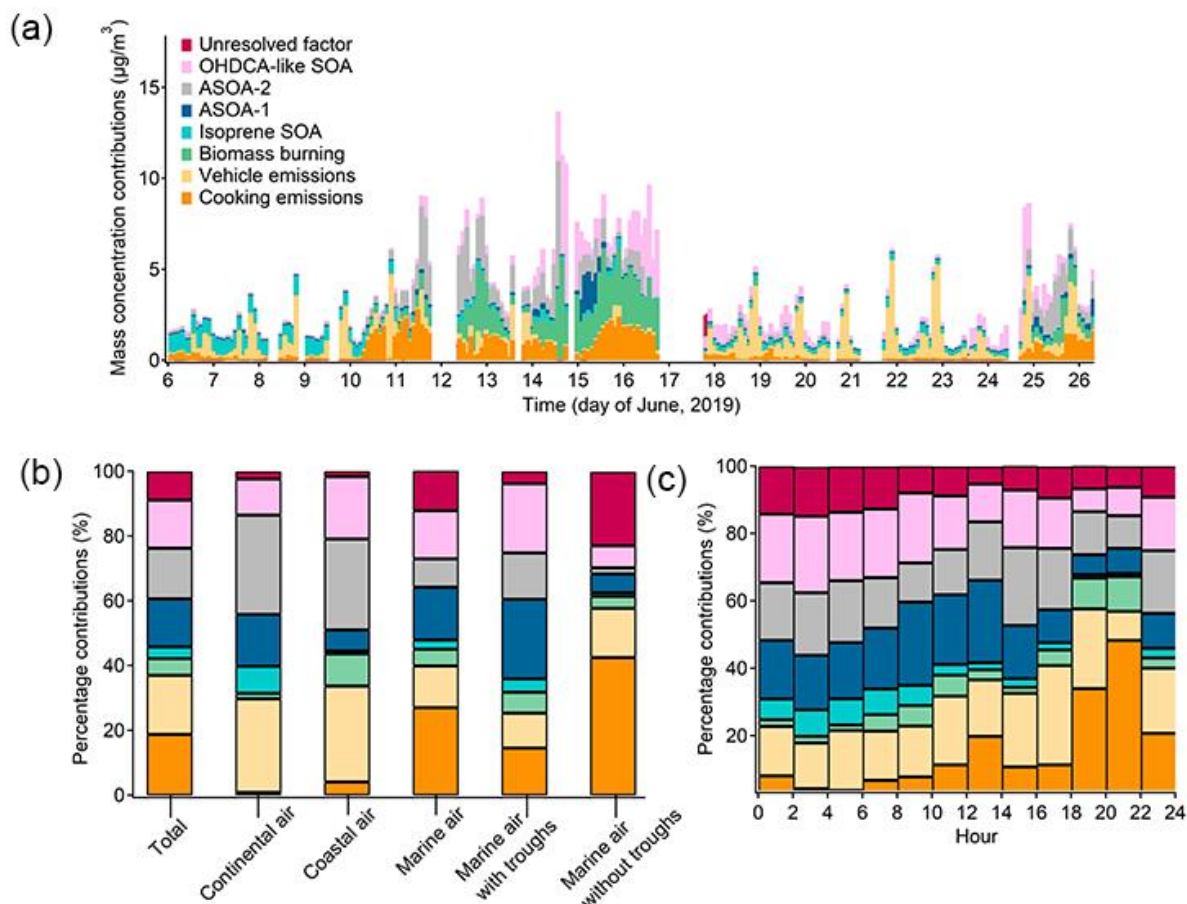


Figure 5: Diurnal patterns of 2-MT1 in different scenarios (a) and the comparisons of temperature (b), UV (c), 2-MT1/2-MTs ratio (d), and LWC (e) between the marine air with and without troughs. Shaded areas represent 95% confidence intervals.



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Figure 6: Average profiles of the eight factors resolved by PMF. The error bars represent 95% confidence intervals estimated by the BS method.



580 **Figure 7: Contributions of individual OA sources to PM_{1-OM} : time series of bihourly data (a); averages for the scenarios of different air masses (b); and diurnal variations (c).**