Impacts of atmospheric circulation patterns and cloud inhibition on aerosol radiative effect and boundary layer structure during winter air pollution in Sichuan Basin, China

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[1]School of Atmospheric Sciences, Nanjing University, Nanjing 210023, China
[2]School of Environment, Nanjing Normal University, Nanjing 210023, China
[3] Chongqing Institute of Meteorological Sciences, Chongqing 401147, China
[4]Chongqing Meteorological Observatory, Chongqing 401147, China
[5]Hubei Provincial Academy of Eco-environmental Sciences, Wuhan 430079, China
[6]Carbon monitoring and digital application technology center, Carbon peak and carbon neutralization strategy institute of Jiangsu Province, Nanjing 210023, China

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[4]Chongqing Meteorological Observatory, Chongqing 401147, China

[5]Hubei Provincial Academy of Eco-environmental Sciences, Wuhan 430079, China

[6]Carbon monitoring and digital application technology center, Carbon peak and carbon neutralization strategy institute of Jiangsu Province, Nanjing 210023, China

Correspondence to: Min Xie (minxie@njnu.edu.cn)

Abstract. Winter persistent aerosol pollution frequently occurs in the Sichuan Basin (SCB) due to its unfavorable weather conditions, such as low wind, wetness, and cloudiness. Based on long – term observational data analyses from 2015–2021, it has been found that the four representative stations in the SCB often simultaneously experience $PM_{2.5}$ pollution accompanied by variations in meteorological conditions above 850 hPa, which indicates a connection between regional winter air pollution in the SCB and large−scale synoptic patterns. The dominant 850 hPa synoptic patterns of winter SCB were classified into six patterns using T–model principal component analysis: (1) strong high pressure in the north, (2) east high west low (EHWL) pressure, (3) weak high pressure in the north, (4) weak ridge of high pressure after the trough, (5) low trough (LT), and (6) strong high pressure. Pattern 2 characterized with EHWL pressure system, and Pattern 5 featured with LT, were identified as key synoptic patterns for the beginning and accumulation of pollution processes. Pattern 1, characterized by a strong high pressure in the north, was the cleanest pattern associated with reduced $PM_{2.5}$ concentrations. The EHWL and LT patterns were associated with a remarkably high cloud liquid content, attributed to upper southerly winds introducing humid air. Clouds reduce solar radiation through reflection and scattering, resulting in more stable stratification and aerosol accumulation. This cloud radiation interaction (CRI) was more pronounced in the LT pattern due to denser isobaric lines and stronger southerly winds than in the EHWL pattern. Numerical simulation experiments utilizing WRF-Chem indicated that there is a upper-level heating during afternoon and surface cooling in the morning forced by the aerosol radiation interaction (ARI) under the EHWL and LT patterns. Additionally, strong surface cooling in the evening influenced by valley winds could be found. With wet and cloudy synoptic forcing, ARI directly affects the stability of the boundary layer and is modulated through CRI inhibition. For example, Chongqing exhibited lower $PM_{2.5}$ concentrations and stronger ARI compared to the western and southern SCB due to lower cloud liquid content and weaker CRI inhibition on ARI. The CRI inhibition caused a 50 % reduction in solar radiation and boundary layer height during the daytime under the LT pattern, which was larger than that under the EHWL pattern. This study comprehensively analyzed the spatial disparities in cloud inhibition on ARIs, their impacts on the boundary layer structure, and the discrepancies of these interactions under different synoptic patterns during pollution processes. The findings hold important implications for effective management of pollution processes in cloudy and foggy weather.

Key words: Synoptic patterns, Cloud radiation interaction inhibition, Aerosol radiation interaction, Boundary layer structure, Sichuan Basin.

1 Introduction

Particulate matter (PM) pollution has become a significant environmental concern in China (Xie et al., 2016a; 2016b; Che et al., 2019). High concentrations of aerosols not only worsen air quality and pose serious health risks to residents, but also have implications for weather and climate through their effects on radiation and clouds (Li et al., 2019; Zhao et al., 2020; Alexeeff et al., 2021; Yang et al., 2021). The interactions between aerosols and clouds present the largest uncertainty in anthropogenic radiative forcing of the Earth's climate (Liao et al., 2017; Haywood et al., 2021). Studying interactions among cloud, aerosol and radiation from an air quality perspective is crucial for a scientific understanding of relationship between weather and pollution.

Excessive emissions are the essential cause of air pollution, with primary aerosol and secondary aerosol formation playing significant roles in comprehending the complete picture of air pollution (Peng et al., 2021). Besides, meteorological conditions not only influence on the formation of secondary aerosols, but also govern the transportation and distribution of both primary and secondary aerosols, and thereby impact regional and long-range air pollution (Zhu et al., 2018; Luo et al., 2018; Nichol et al., 2020; Zhang et al., 2020; Jiang et al., 2021). PM and gaseous pollutants, primarily transported by the planetary boundary layer (PBL), are directly or indirectly influenced by various meteorological factors such as wind, relative humidity, PBL height (PBLH), and solar radiation. These factors contribute to the multi–temporal and spatial distribution characteristics through vertical and horizontal diffusion, physicochemical reactions, and dry and wet deposition (Park et al., 2017; Shu et al., 2017; Zhan et al., 2019; Huang et al., 2019). Large–scale synoptic forcing is considered the primary driving condition for meteorological factors, PBL structure, and the resulting distribution of atmospheric pollutants (Miao et al., 2019; Ning et al., 2019; Jiang et al., 2020; Li et al., 2021). Specific synoptic patterns can induce advection, which largely determines the local PBL structure and development. PBL, located at the bottom of the atmosphere, is responsible for the main exchange of heat, moisture, and matter between the surface and the free troposphere (Stull, 1988). The fate of pollutants emitted near the surface, a significant source of aerosols in the air, is largely controlled by the PBL (Garratt, 1994). The PBLH is often used as a metric to characterize the capacity and dilution of pollutants (Seidel et al., 2010). Synoptic patterns can directly determine the meteorological conditions of emitted pollutants and influence their transport by regulating PBL thermal stratification and mechanical turbulence (Stull, 1988; Ning et al., 2018; Zhan et al., 2019; Jiang et al., 2021; Zhang et al., 2022).

Unfavorable meteorological conditions play a significant role in contributing to aerosol pollution. When pollutants accumulate to a certain degree, aerosols can reduce surface solar radiation through backscattering or absorbing solar radiation, leading to surface cooling. This decrease in solar radiation and temperature near the ground weakens turbulent diffusion, suppresses the convective development of the PBL, and lowers PBLH, which in turn exacerbates aerosol pollution (Ding et al., 2016; Wang et al., 2018). Moreover, the increase in humidity caused by the decreased surface saturation vapor pressure and inhibited water vapor diffusion enhances aerosol hygroscopic growth accelerates liquid–phase and heterogeneous reactions, and contributes to aerosol pollution (Pilinis et al., 1989). The positive feedback between unfavorable PBL meteorology and increasing aerosols was found to be responsible for the majority of the increase in $PM_{2.5}$ during cumulative stages in various regions of eastern China affected by aerosol pollution, including the North China Plain, the Guanzhong Plain, the Yangtze River Delta, the Two Lakes Basin, the Pearl River Delta and the Northeast China Plain. But in the Sichuan Bain (SCB), the feedback is weak due to the suppression of the cloudy mid-upper layer (Zhong et al., 2018; Zhong et al., 2019). As for the aerosol-cloud interactions, arise from increasing aerosols acting as cloud condensation nuclei in cloud and translating into larger concentrations of smaller cloud droplets, leading to an increased cloud albedo reflecting more radiation back to space (Twomey, 1977; Lohmann and Feichter, 2005). Even a marginal increase in cloud droplets above pristine conditions in deep convective clouds causes more droplets to reach supercooled levels, which enhances latent heat release and invigorates convection (Rosenfeld et al., 2009; Possner et al., 2015). Further increases in cloud droplets result in direct radiative effects, reducing downward solar radiation, cooling the surface, and inhibiting convection (Scott et al., 2016).

The SCB is surrounded by high mountains with cloudy and wet weather conditions. The mean annual relative humidity in the SCB is around 75%, with cloud fraction exceeding 80%, and an average of 1200 hours of sunshine per year. The Chengdu–Chongqing city cluster in the SCB serves as the economic center of the upper reaches of the Yangtze River in China, accounting for approximately 10 % of the country's population. However, rapid industrialization and urbanization in this region have resulted in severe air pollution. The SCB is recognized as one of the most polluted regions in China, with high black carbon concentrations (Li et al., 2016; Cao et al., 2021). The Qinghai–Tibet Plateau on the western edge of the SCB significantly influences the transport and accumulation of pollutants through thermal and dynamic effects (Ning et al., 2017; Shu et al., 2021). In addition, the Qinghai–Tibet topography leads to higher cloud water content over the SCB than the other regions (Yu et al., 2004; Yang et al., 2012). Many studies have emphasized the importance of the interactions between cloud, aerosols and radiation in air pollution processes (Wang et al., 2018; Hu et al., 2021). High pollutant emissions, combined with the prevalence of cloudy and foggy weather, make these interactions in the SCB even more complex than those in other regions. The aerosol radiation interactions (ARI) can be inhibited by cloud in cities like Chengdu (Zhong et al., 2019). However, there is a lack of in-depth quantitative discussions regarding this aspects in the SCB. On one hand, the complex terrain in the SCB leads to differences in the meteorological conditions between them (Ning et al., 2017; Lu et al., 2022). For example, Chengdu is a typical basin city while Chongqing is a mountain city located on the basin slope, so they have markedly different climate conditions. It remains to be elucidated whether these conditions will result in spatial disparities in cloud inhibition on the ARI. On the other hand, synoptic forcing, as the primary driver of meteorological variations, undoubtedly play an unneglectable role in shaping cloud cover and boundary layer structures (Miao et al., 2020; Wang et al., 2022; Painemal et al., 2023). The discrepancies in cloud inhibition on ARI under different synoptic patterns also need to be revealed. Addressing these issues is crucial for understanding the persistent pollution processes and the intricate interactions between weather and pollution in the SCB. It holds important implications for the effective management of pollution processes in cloudy and foggy weather.

Characterized with high aerosol loadings and semi–permanent cloudy weather, the SCB provides an ideal region for studying the complex interactions between clouds, aerosols, and the PBL. This study objectively classifies the synoptic patterns influencing the SCB based on long–term data. By conducting an integrated analysis of pollutants and meteorological factors, the primary pollution sources and clean synoptic patterns are identified. To further investigate the inhibition of cloud radiation interaction (CRI) on ARI under different synoptic patterns in the SCB, WRF–CHEM simulation experiments are conducted. The results contribute to a deeper understanding of CRI, ARI, and the PBL interactions in regions influenced by plateau–basin topography with wet and cloudy weather. The data and methods are presented in Section 2, whereas Section 3 describes the synoptic patterns and their corresponding impacts on clouds, aerosols, radiation, and PBL. Finally, the conclusions are presented in Section 5.

2 Data and method

2.1 Observation data

Air quality monitoring data used in this study were obtained from air quality monitoring sites established by the Ministry of Ecology and Environment of China across the SCB. Hourly $PM_{2.5}$ observations from 18 stations in the SCB were collected during the winter period from 2015 to 2021 for data analysis and model verification purposes (Fig. 1b). The abbreviations CQ, CD, MY, DY, LS, MS, YA, ZY, ZG, YB, LZ, NJ, GA, NC, SN, GY, DZ, and BZ represent the following cities: Chongqing, Chengdu, Mianyang, Deyang, Leshan, Meishan, Yaan, Ziyang, Zigong, Yibin, Luzhou, Neijiang, Guangan, Nanchong, Suining, Guangyuan, Dazhou, and Bazhong, respectively.

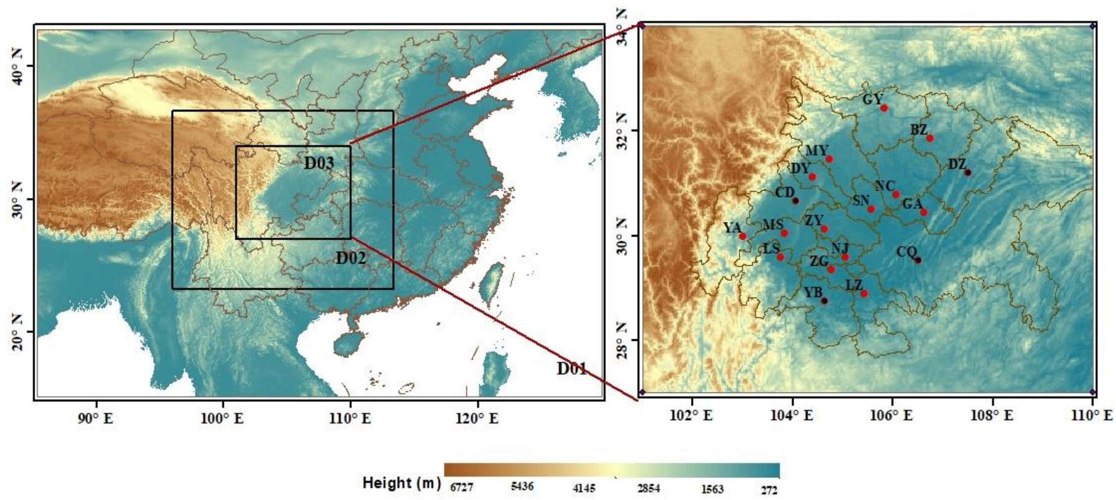

Figure 1. (a) Three layers of simulation domains in WRF-Chem with topography map as shading; (b) the locations of 18 air quality monitoring stations (red dots) and 4 sounding stations (black dots) in the Domain 3.

The SCB has four sounding stations: Wenjiang (CD), YB, DZ, and Shapingba (CQ), situated in the western, southern, northwestern, and eastern regions of the basin, respectively (Fig. 1b), and represent different pollution and meteorological conditions in different regions within the SCB. In all, the air pollution over the SCB exhibits a gradual decrease from southwest to northeast. Statistical analysis indicates that the western and the southern basin experience the most severe pollution. The western basin shows the highest pollution proportion, while the southern basin exhibits the highest occurrence of heavy pollution. In the northeastern basin, specifically in DZ, heavy pollution is more likely to occur during winter, which verifies it to be the third highest pollution zone outside the western and southern basin. This makes the spatial distribution during winter differs from the overall annual pollution pattern in the SCB (Lu et al., 2022; Qi et al., 2022). Regarding meteorological conditions, research reveals that DZ has the lowest ventilation coefficient during winter, while CQ has the highest. The SCB experiences frequent temperature inversions, with CD having a higher occurrence of inversions compared to the other three cities. CD also exhibits the strongest inversion intensity and is prone to multi-layer inversions. On the other hand, YB and CQ have greater inversion thickness, while CD has the smallest inversion thickness (Feng et al., 2020). The vertical distribution of the meteorological factors used in the study was obtained from an L–band sounding radar, collecting temperature, pressure, humidity, and wind data at 00:00 and 12:00 Coordinated Universal Time (UTC) on vertical levels every second from the surface up to 30 km. Ground observation data from the four cities, including temperature and dew point temperature, were used for meteorological factor simulation verification. All meteorological data were obtained from the China Weather Website Platform maintained by the China Meteorological Bureau. As for the calculation of PBLH, there are various methods to determine the PBLH, and differences in methods, data or threshold values may yield quite different PBLH results (Seibert et al., 2000; Eresmaa et al., 2006; Jiang et al., 2021). The bulk

Richardson number (*Ri*) method was adopted to calculate the PBLH with sounding data in the study by assuming that the PBLH is the height at which *Ri* reaches its critical value (*Rc*). *Ri* at a certain height *h* is calculated as follows:

$$Ri = \frac{(g/\theta_{v0})(\theta_{vh} - \theta_{v0})h}{u_h^2 + v_h^2}$$

Where g is the acceleration of gravity, θ_{v0} and θ_{vh} are the virtual potential temperature at surface and the height *h,* respectively, and u_h and v_h are the meridional and zonal wind components at *h*. We adopted the *Ri* method and *Rc* to be 0.25, because the EAR5 and YSU schemes use the same method and threshold value when calculating PBLH (Hong et al., 2006; ECMWF, 2017).

CD, YB, DZ and CQ were selected as representative cities for analysis in the study. These four representative cities are located in the western, southern, northwestern, and eastern regions of the basin, to capture diverse pollution and meteorological conditions within the SCB. These cities were chosen to represent the most polluted regions (Zhao et al., 2018; Lu et al., 2022), as well as typical basin and mountainous cities. Furthermore, there are only four sounding stations in the SCB available, which are located in these four representative cities. They can provide valuable vertical and surface meteorological observations, as well as pollution data, contributing comprehensive dataset used in this study.

ERA5 reanalysis data from the European Centre for Medium-Range Weather Forecasts (ECMWF), which assimilates comprehensive observation data, including ground observation, sounding data, aircraft observation data, and satellite observation data, were obtained for synoptic pattern classification and their impact on meteorological factors in four representative cities. The EAR5 data at the 850 hPa pressure level were collected for the synoptic pattern study. Additionally, cloud liquid water content and downward solar radiation derived from the EAR5 single–level datasets were obtained to assess the influences of synoptic forcing on CRI studies, while PBLH were adopted to conduct the simulation verification. Previous studies have demonstrated the reliability of ERA5 data in estimating cloud properties, including the cloud liquid content (Yao et al., 2019; Nandan et al., 2022; Ojo et al., 2023).

2.2 Synoptic pattern classification

The objective classification was conducted on the synoptic patterns of the SCB using ERA5 data, including geopotential height, u, and v components of winds at the 850 hPa pressure level. The analysis covered an area of 97–117° E and 24–37° N with a horizontal resolution of 0.25° × 0.25°. Given that PM pollution in the SCB is primarily prevalent during winter months (Zhao et al., 2018; Lu et al., 2022), the synoptic pattern classification was performed for winter seasons from 2015 to 2021 (December, January, and February) using the principal component analysis in the T–model (T–PCA) objective method. Compared with the subjective classification method, the objective method can process large amounts of data without relying on subjective experience (Huth et al., 2008; Miao et al., 2017). Among various classification methods, the T–PCA method accurately reflects the characteristics of the original synoptic circulations and exhibits spatial and temporal stability (Huth et al., 1996; Huth et al., 2008). Consequently, the T–PCA has been widely used in synoptic pattern classification researches (Ning et al., 2019; Miao et al., 2020; Li et al., 2021).

2.3 Model configuration and simulation experiments

To understand the combined effects of synoptic patterns and CRI inhibition on ARI and PBL, a series of parallel experiments were conducted on the simulation of a typical pollution episode using the Weather Research and Forecasting model with Chemistry (WRF–Chem v3.9.1) (Grell et al., 2005). The Advanced Research WRF (ARW) dynamics solver integrates the compressible, nonhydrostatic Euler equations, for example, the momentum equation, the continuity equation, the thermodynamic equation, the moisture equation and the ideal-gas equation of state (Skamarock et al., 2008). The model domain (Fig. 1a) was centered over the SCB and utilized three layers of nested grids with horizontal resolutions of 27, 9, and 3 km, respectively. A total of 32 vertical layers spanning from the surface to 100 hPa were defined. Initial and boundary meteorological fields were obtained from the National Centers for Environmental Prediction Final reanalysis data with a horizontal resolution of 1° × 1° and 6 h time interval. For chemical process simulations, anthropogenic emissions were sourced from the Multiresolution Emission Inventory for China (MEIC) in 2016, featuring a grid resolution of 0.25° × 0.25°. To address the empirically overestimated $PM_{2.5}$ emissions by the MEIC in the SCB (Zhan et al., 2023), the ensemble square root Kalman filter were implemented on the $PM_{2.5}$ emission during simulation (Wu et al., 2018; Lu et al., 2021). Biogenic emissions were calculated online using the Guenther scheme (Guenther et al., 2006). Table 1 provides a summary of the chosen physical and chemical parameterization schemes. The parameterization schemes employed in this study is the one used by the Chongqing Meteorological Bureau in the daily operational activities. The schemes have been obtained through multiple sets of control experiments and are considered suitable for the simulation in the SCB.

Table 1 The main options of WRF–Chem

Items	**Contents**
Domains (x, y)	(155, 110), (184, 160), (320, 250)
Grid spacing (km)	27, 9, 3
Center	(29.1° N, 106.2° E)
Time step (s)	60
Microphysics	WRF Single–Moment 5 class (WSM5) scheme
Longwave radiation	RRTMG scheme (Iacono et al., 2008)
Shortwave radiation	RRTMG scheme (Iacono et al., 2008)
Planetary boundary layer	Younsei University scheme (Hong et al., 2006)
Land surface	United Noah land surface model (Tewari et al., 2004)
Cumulus parameterization	Grell–Freitas ensemble scheme

	(Grell et al., 2013)
Advection	fifth- and third-order differencing for horizontal and vertical advection respectively
Photolysis scheme	Fast-J photolysis (Fast et al., 2006)
Gas–phase chemistry	RADM2 (Stockwell et al., 1990)
Aerosol module	MADE/SORGAM (Schell et al., 2001)

To assess the impact of CRI inhibition on ARI under typical synoptic pollution patterns, four parallel experiments were conducted using simulation models. The selected simulation period for these experiments was January 1-7, 2017. The period was selected for two reasons: the Chinese government announced clean-air action in the year of 2013, aiming to reduce $PM_{2.5}$ concentrations in the next 5 year. Specifically, the year of 2017 was identified as a key year for assessing $PM_{2.5}$ pollution in China, as significant practical actions were implemented during the period (Wang et al., 2020) and the selected period encompassed both typical pollution and clean weather patterns.

The baseline experiment (BASE) included both CRI and ARI in the simulations. In contrast, the three sensitivity experiments focused on excluding either ARI or CRI. Experiment 1 (EXP1) did not consider ARI, Experiment 2 (EXP2) did not include CRI, and Experiment 3 (EXP3) omitted ARI when CRI was not included. The differences between BASE and EXP1 represented the disturbances caused by ARI, while EXP2 and EXP3 represented the influences of ARI without CRI inhibition. Detail differences between the experiments could be found in Table 2. The numerical experiments were initiated at 00:00 UTC on December 30, 2016, and ran until 00:00 UTC on January 8, 2017, with the first 48 hours designated as a model spin–up period.

Table 2 Four numerical simulation experiments are conducted in the study

Experiments	Description	Results	Meaning
BASE	Baseline simulation	BASE-EXP1	Disturbances by ARI
EXP1	Only shutting down ARI		
EXP2	Only shutting down CRI	EXP2-EXP3	Influences of ARI without CRI
EXP3	Shutting down both ARI and CRI		

*ARI: aerosol radiation interaction; CRI: cloud radiation interaction

3 Results and discussions

3.1 Relationships between synoptic patterns and $PM_{2.5}$ pollution in the SCB

Figure 2 illustrates the daily mean variations in $PM_{2.5}$ concentration and vertical distributions of potential temperature (PT) during winter period from 2015 to 2021, with a focus on the pollution episodes. The four sounding stations located in separate areas of the SCB (CD, YB, CQ, and DZ), consistently experienced pollution processes characterized by simultaneous changes in vertical thermal structures. For example, during the pollution events in January 2017 and December 2020, the $PM_{2.5}$ concentrations in all four cities reached their peak levels at the same time before rapid declining (Fig. 2). Interestingly, these pollution episodes were accompanied by warming in the upper layer atmosphere, while a decrease in $PM_{2.5}$ concentration correlated with cooling. Despite the significant distances between these cities (approximately 200–400 km), the synchronized changes in pollutant concentrations and vertical thermal structures could be attributed to large–scale synoptic patterns (Miao et al., 2020; Li et al., 2021). While the four cities with sounding stations were selected as representatives for vertical thermal structure analysis, other cities in the SCB also experienced pollution episodes and relevant physical processes, except for GY (Fig. S1). GY is located in the northern edge of the SCB, bordering Shaanxi and Gansu Provinces. The proportion of heavy $PM_{2.5}$ pollution in GY is the lowest in the basin, but the proportion of PM_{10} pollution is higher than other cities of SCB (Lu et al., 2022). Due to the lower $PM_{2.5}$ concentration, the two pollution processes in January 2017 in GY were not as significant as in other cities whithin the basin. However, the warming of upper air coincided with $PM_{2.5}$ increase could still be observed.

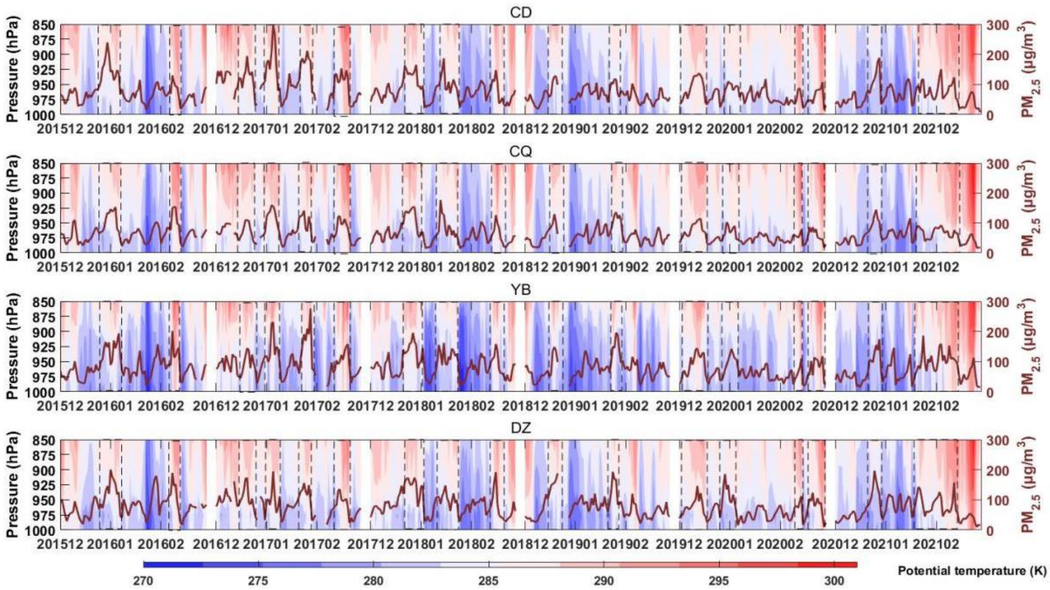

Figure 2. Time series of daily mean $PM_{2.5}$ and potential temperature derived from the sounding data during 2015-2021 winter months. The $PM_{2.5}$ pollution episodes are marked with black dotted boxes.

The time series of daily mean $PM_{2.5}$ from air quality monitoring sites and the accompanying vertical distributions of temperature, relative humidity, and wind in CD, CQ, YB, and DZ derived from the sounding stations are shown in Fig. S2, focusing on January 2017 as an example for analysis. During this month, two severe $PM_{2.5}$ pollution episodes occurred: one from January 1 to 7 and another from January 24 to 31 in 2017. These pollution episodes had a significant impact on air quality in all four cities. The highest daily $PM_{2.5}$ concentrations recorded during these episodes were 291.17 μg/m^3 in CD and 276.2 μg/m^3 in YB. Pollution in early January exhibited a gradual increase in $PM_{2.5}$ levels from January 1 to 3, with upper air warming and the emergence of an inversion above the PBL. Additionally, lower humidity and higher wind speeds above 1500 m were observed during the pollution accumulation period. Similarly, the late January pollution episode showed a rapid increase in $PM_{2.5}$, from January 24 to 27, together with warming, dryness, and high wind speed above 1500 m in all four cities. These consistent meteorological conditions during the pollution periods indicated significant synoptic forcing. The previous study has found that winter heavy pollution processes in the SCB are usually associated with abnormal warming above the 850 hPa (Lu et al., 2022). The warming is induced by strong southerly airflow above the basin. The southerly airflow in winter over the SCB originates from the Yunnan-Guizhou Plateau or the Indian Peninsula, characterized with high temperature, dryness, and high wind speed. The strong southerly airflow forms a warm lid over the basin, suppressing the vertical exchange of pollutants within the basin. As a result, pollutants accumulate rapidly, which may explain the phenomenon of rapid $PM_{2.5}$ growth accompanied by warming, dryness, and strong winds above 1500 m. Notably, the key layer for studying the connection between synoptic patterns and $PM_{2.5}$ pollution is approximately 850 hPa, corresponding to a height of approximately 1500 m within the PBL, where changes in specific meteorological conditions primarily affect surface–emitted pollutants.

Using ERA5 reanalysis data for winter (December, January, and February) from 2015 to 2021, the 850 hPa synoptic patterns over the SCB were objectively classified into six types (Fig. 3). According to the relative positions of the high–pressure and low–pressure systems in the basin, these synoptic patterns could be described as follows: (1) strong high pressure in the north, (2) east high west low (EHWL) pressure, (3) weak high pressure in the north, (4) weak ridge of high pressure after the trough, (5) low trough (LT), and (6) strong high pressure. Patterns 1 and 3 exhibited high pressure in the northern SCB, which differed from the high–pressure intensity. With strong high pressure, the basin was primarily controlled by northerly airflow. Under weak high–pressure conditions, the basin was dominated by an easterly backflow. Patterns 2 and 5 had high and low pressures near the basin, forming a relatively dense isopotential altitude gradient and resulting in strong southerly winds over 850 hPa. Pattern 4 was a weak high–pressure ridge after a trough controlled the SCB with sparse isobaric lines and weak winds leading to static and stable weather conditions. During Pattern 6, the SCB was controlled by the cold high-pressure system, accompanying weak northerly airflow on the basin. Pattern 6 usually evolved from either Pattern 1 or Pattern 3.

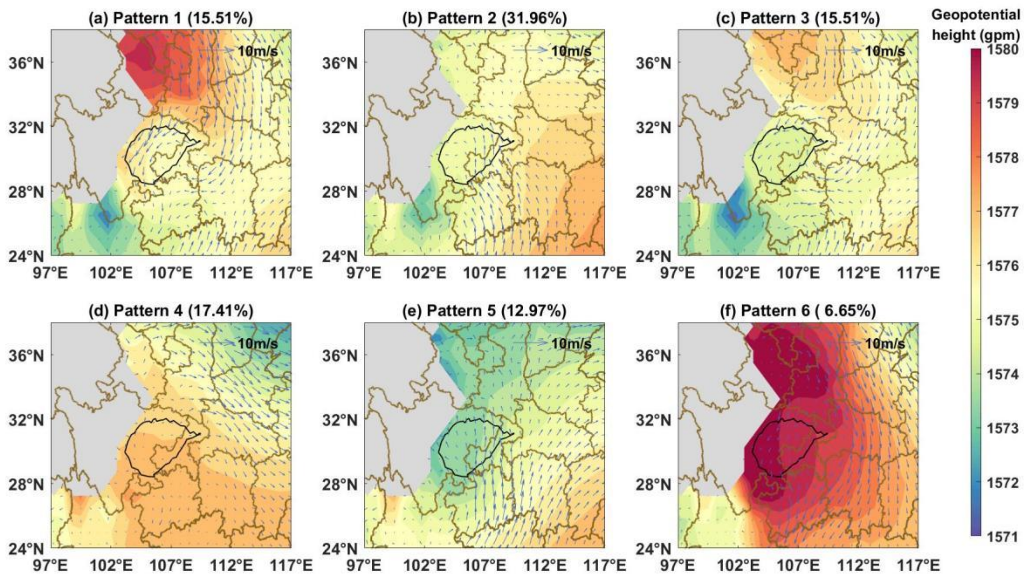

Figure 3. The 850hPa geopotential height field (shading) with wind vector fields (blue vectors), and frequency of occurrence for 6 synoptic patterns during 2015-2021 winter months. The SCB was outlined with an altitude contour of 750 m terrain height (black lines).

Patterns 2, 4, and 5 exhibited higher frequencies of pollution occurrence ($PM_{2.5}$ daily concentration ≥ 75 μg/m^3) according to statistical results from 18 cities in the SCB during the 2015–2021 winters (Fig. 4a). These patterns were associated with high $PM_{2.5}$ concentrations in 50–70 % days, including CD, DY, and MY in the northern SCB, 40–60 % for cities in the southern SCB, such as ZG and YB, and also 40–60 % of days for cities in the northern SCB, such as CQ, DZ, NC, and GA. Furthermore, the average $PM_{2.5}$ concentrations in the respective cities for the six synaptic patterns were calculated (Fig. 4b), aligning with the frequency of pollution occurrence. The days under Patterns 2, 4, and 5 exhibited higher average daily $PM_{2.5}$ concentrations. The average concentrations under these three synoptic patterns were 99.19, 103.43, and 111.97 μg/m^3 for CD, 95.44, 87.98, and 94.26 μg/m^3 for YB, 79.14, 83.96, and 74.77 μg/m^3 for CQ, and 91.02, 104.64, and 91.51 μg/m^3 for DZ, respectively. Regarding the impact of synoptic patterns on the accumulation or dispersion of $PM_{2.5}$, Fig. 4c illustrates the average daily changes in $PM_{2.5}$ concentration compared with the previous day for CD, CQ, YB, and DZ under the six synoptic patterns. Patterns 2 and 5 exhibited the most significant $PM_{2.5}$ accumulation under the influence of southerly airflow. The average $PM_{2.5}$ concentration under Pattern 1, 3 and 6 was lower in all cities of SCB than other three pollution patterns (Fig. 4a). Besides, the day to day $PM_{2.5}$ variations under Pattern 1, 3 and 6 exhibited negative growth trend in the four representative cities (Fig. 4c). As a result, Pattern 1,3 and 6 were identified as the "clean pattern". In addition, the pollution occurrence frequency of which was found higher for cities located in the eastern part of the SCB than other parts. Under Pattern 6, strongest northerly airflow affects the basin. The eastern part of the basin consists of parallel ridges and valleys, which reduces wind speed. The stronger the wind is, the more obvious the reduction of wind by terrain is. In contrast, the western part is relatively flat, which can result in higher surface wind speeds. The difference in wind impacted by terrain led to a weaker pollution removal effect in the eastern region, thus contributing to a higher proportion of pollution days under Pattern 6. Besides, differences in precipitation rates between eastern cities and other regions were not significant (the proportion of rainfall with a daily accumulated precipitation exceeding 10 mm in CD, CQ, YB and DZ under Pattern 6 were all less than 3%) , which might not the main reason why eastern cities in the SCB experience higher pollution frequency.

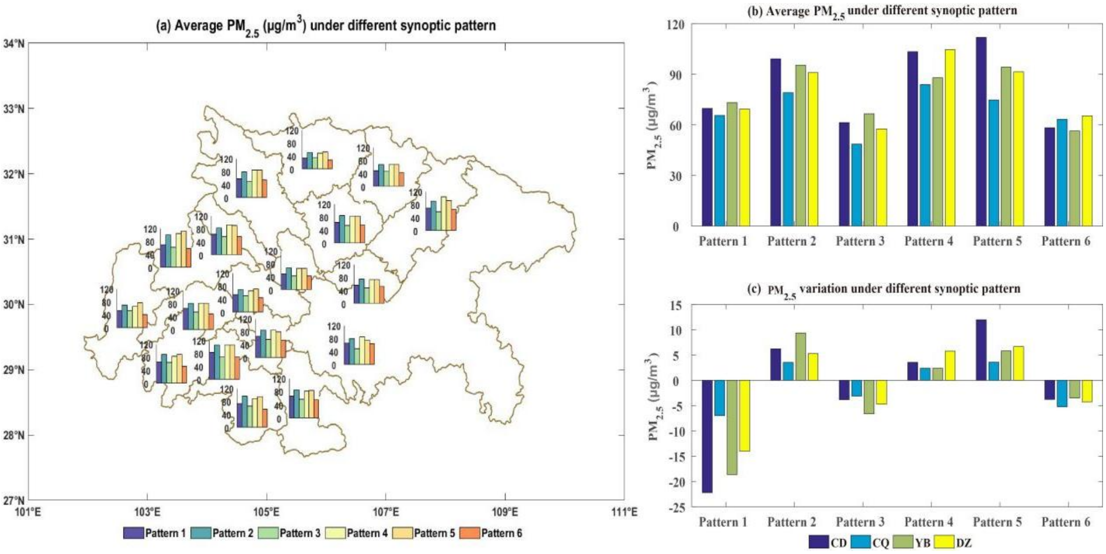

Figure 4. (a) The pollution occurrence frequency at 18 air pollution stations in SCB, (b) (c) average $PM_{2.5}$ concentrations and $PM_{2.5}$ day to day variations at 4 representative SCB cities, under 6 synoptic patterns.

The time series of daily mean $PM_{2.5}$ and the day-to-day classification of 850 hPa synoptic patterns are shown in Fig. S3, from December 2016 to January 2017. Six pollution episodes occurred during this period (December 03–12 and 16–26, 2016; January 1–7, 16–19, and 20–28, 2017; and February 14–23). It is observed that pollution episodes consistently began with Pattern 2 and ended with Pattern 1, accompanied by a rapid decline in $PM_{2.5}$. This finding suggests that Pattern 2 acted as a key synoptic forcing for the initation of pollution episodes. Additionally, statistical results revealed that Pattern 2 accounted for a high proportion of $PM_{2.5}$ increase during the six pollution episodes, reaching 48.48 %, while Pattern 5 had the second highest proportion of 21.88 %, with Patterns 2 and 5 combined accounting for more than 70 % of the pollution episodes. For example, during the two heavy pollution events in early and late January 2017, $PM_{2.5}$ rapidly accumulated with the interplay of Patterns 2 and 5. These two patterns represented a substantial proportion of 31.96 % and 12.97 %, respectively, during winters from 2015 to 2020 at 850 hPa level in the SCB (Fig. 3). Based on this analysis, Patterns 2, 4, and 5 were identified as synoptic pollution patterns, whereas Patterns 1, 3, and 6 were as clean patterns. In summary, Patterns 2 and 5 played crucial roles in the initiation and accumulation of $PM_{2.5}$ during pollution episodes.

The discussion above showed that pollution in the SCB tended to occur when southerly airflow controlled the upper-layer of the basin (Pattern 2 and 5), while the dispersion of pollutants was accompanied by northerly winds, which aligns with the findings of Lu et al. (2022). This study indicated that southerly airflow in the upper-layer could bring warm air, leading to warming above the basin and forming a “warm lid”. The surrounding mountains and plateau with the “warm lid” contributed to the formation of a relative enclosed space within the SCB, facilitating local circulations and allowing for the thorough mixing and secondary reactions of local emission and pollutants transported from outside. As a result, persistent and severe pollution often occurred under the influence of southerly airflow. When the northerly airflow began to dominate the SCB, the “warm lid” and local circulation were disrupted, leading to dispersion of pollutants through advection and vertical transport. Northerly winds were often associated with cold air and sometimes accompanied by weak precipitation, resulting to wet deposition and the removal of pollutants. Therefore, the arrival of northerly airflow often signified the ending of the pollution episode. The evolution of 850 hPa synoptic forcing and vertical meteorological conditions (Fig. 2 and 6) aligns with the study of Lu et al (2022). Therefore, there are also similar pollution change mechanisms.

Due to the convergence of air moving eastward across the Tibetan Plateau, the SCB experiences frequent wet and cloudy weather, with cloud cover fraction exceeding 80 % (Yu et al., 2004; Zhang and Lin, 1985). Clouds undoubtedly play an unneglectable role in the interactions of aerosols, radiation, and the PBL under typical synoptic forcing in this region. This study evaluated the average cloud liquid water content, downward solar radiation, and PBL under the influence of the six classified synoptic patterns in CD, CQ, DZ, and YB, using data from ERA5 (Fig. 5). The reanalysis data revealed significant higher cloud liquid water contents with Patterns 2 and 5, likely triggered by robust southerly air prevailing at 850 hPa over the SCB (Fig. 3). This southerly air brought warm and moist air, contributing to cloud formation. Dense clouds reduced solar radiation through reflection and scattering, resulting in surface cooling and inhibiting PBL development. The PBLH under Patterns 2 and 5 was approximately 900–1000 m, lower than that under the influence of clean synoptic Pattern 6 at 1500 m or Pattern 1 and 3 at 1200–1300 m (Fig. 5). In contrast, the clean synoptic Pattern 1 was characterized by a strong northerly flow at 850 hPa, resulting in lower cloud liquid water content over the basin and increased solar radiation, promoting PBL development. The lower PBLH with more stable stratification caused by the CRI in Patterns 2 and 5 could partially explain the rapid accumulation of $PM_{2.5}$ during these two pollution patterns.

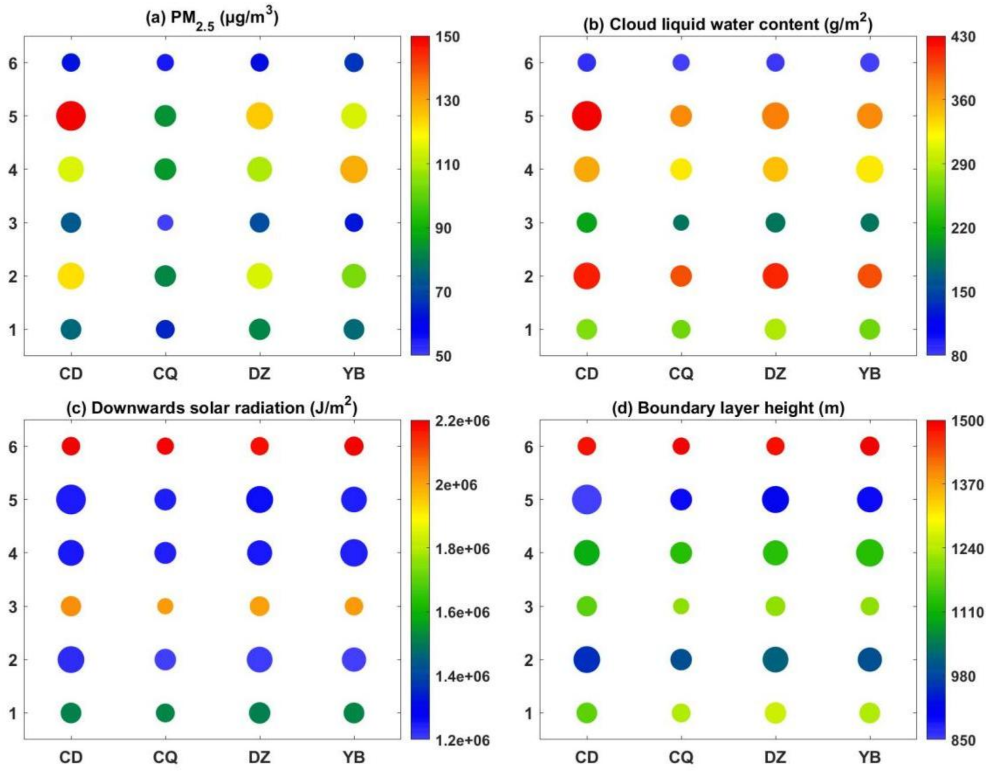

Figure 5. The averaged (a) $PM_{2.5}$ concentrations, (b) cloud liquid water contents, (c) downwards solar radiation and (d) boundary layer height derived from 2015-2021 winter months ERA5 reanalysis data, at 4 representative SCB cities under 6 synoptic patterns, the dot sizes represent $PM_{2.5}$ concentrations.

3.2 Integrate impacts of synoptic patterns and the CRI inhibition on ARI

Based on the above analysis, Patterns 2 and 5 were identified as the key pollution synoptic patterns accompanying dense clouds and, thus, strong CRI. However, the effects of pollution patterns on ARI and their interaction with CRI in the SCB remain unclear and warrant further investigation. A typical pollution episode from January 1–7, 2017, was selected to understand these complex processes and simulated using WRF–Chem. The BASE simulations were verified with observations to determine the accuracy and reliability of the simulation results. The simulation and observation of $PM_{2.5}$, T2, TD2 and wind speed values with some statistical metrics in CD from January 1–7 are shown in Fig. 6a-d. Similar information at CQ, YB and DZ can be found in the Fig. S4. The MB of the simulated and observed $PM_{2.5}$ concentrations were -15.59, -13.42, 2.10 and -13.11 μg/m^3, with NMB values of -4.12%, -4.22%, 6.01% and -0.68% at four cities, respectively, which are within the acceptable standards (NMB < ± 15%). The R of $PM_{2.5}$ were 78.91%, 57.23%, 61.15% and 62.86% for four representative cities, respectively. The statistical metrics for $PM_{2.5}$ are consistent with previous studies (Wang et al., 2020; Shu et al., 2021; Zhan et al., 2023), indicating that our model results for $PM_{2.5}$ are reasonable and acceptable. Regarding to the surface meteorological factors, low MB and high R for both temperature and dew point temperature suggested good simulation performance for these variables. However, the simulation results for wind speed were poor, which was expected under conditions of low wind and complex terrain. The high observed calm wind frequency, influenced by the starting speed of the anemometer, led to an overestimation in the simulation (Shu et al., 2021; Zhan et al., 2023). Additionally, it could be argued that unresolved topographic features introduce additional drag, beyond that generated by vegetation, which was not considered in the WRF model (Jimenez and Dudhia, 2012).

In addition, the temporal averaged and variations of vertical profiles for potential temperature, relative humidity and wind speed in the model were compared with the sounding data in CD (Fig. 6e-m). Model evaluation of vertical structures in CQ, YB and DZ can be found in Fig. S5. The SCB is characterized with cloudy and foggy conditions, which result in abundant water vapor and near 100% relative humidity above the nocturnal boundary layer. Models often underestimate the humidity above the boundary layer during night in the SCB (Shu et al., 2021). Furthermore, due to complex terrain and measurement bias of the anemometer for weak winds, the evaluation of simulation results for wind speed often exhibit certain deviations (Jimenez and Dudhia, 2021; Shu et al., 2021; Zhan et al., 2023). For the verification of PBLH, sounding data are commonly regarded as reliable vertical observation records, and PBLH calculated based on sounding data can be used as the true values to compare with other data for long-term validation (Guo et al., 2016). However, for short-term studies, due to limited availability of sounding data at only 00:00 and 12:00 UTC, the ERA5 data were also incorporated for the model evaluation of PBLH in this study (Fig. 6 and Fig.S5). The simulation PBLH showed a consistent trend with those calculated from ERA5 and sounding data. Overall, the simulation results can capture the meteorological and $PM_{2.5}$ variation trends. According to the simulation evaluation standards for the SCB in previous studies (Wang et al., 2020; Zhan et al., 2023), the results is acceptable and reasonable; thus, the simulation can be used for subsequent analysis and discussion.

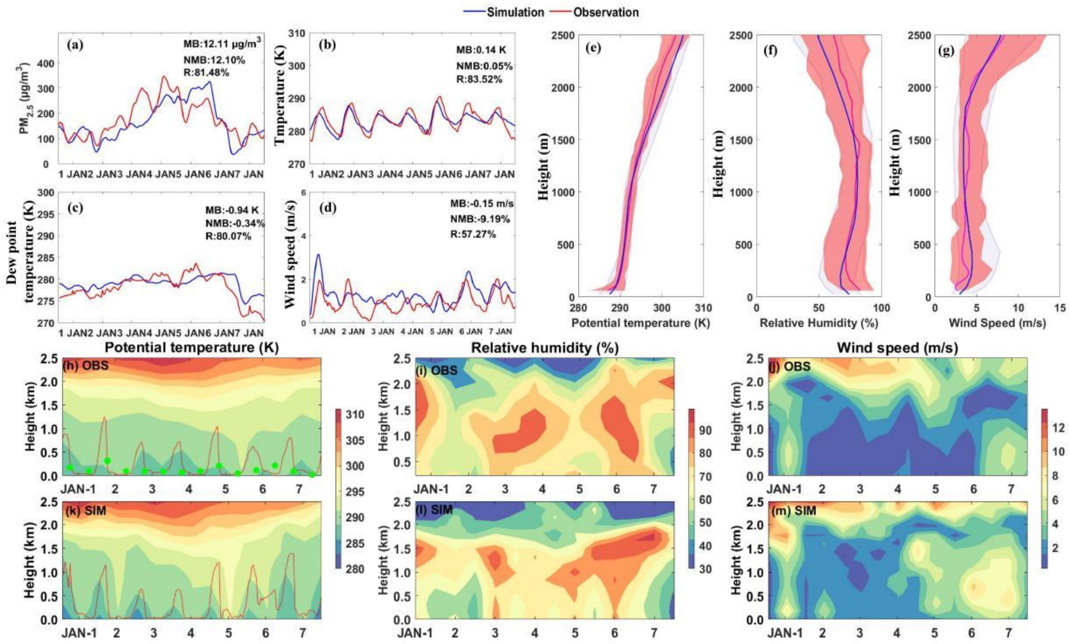

Figure 6. Time series of hourly simulated and observed (a) $PM_{2.5}$ concentration, (b) temperature at 2 m, (c) dew temperature at 2m and (d) wind speed near surface, and comparison of simulated and observed mean vertical profile of (e) potential temperature, (f) relative humidiy and (g) wind speed, the red and grey shaded areas represent deviations to the mean values of observation and simulation, respectively. The simulated and observed time-height sections of (h)(k) potential temperature, (i)(l) relative humidity and (j)(m) wind speed are also given, while the red lines in (h)(k) are time series the boundary layer heights derived from ERA5 data and simulation with green dots representing boundary layer heights calculated with sounding data. The above figures display information in CD. Additionally, the model verification information regarding CQ, YB and DZ can be found in Supplement Figure 3 and 4.

During the pollution episode that occurred from January 1 to 7, 2017, the pollution synoptic patterns controlled the SCB as follows: Pattern 2 from January 1 to 3, Pattern 5 from January 4 to 6, and Pattern 1 on January 7. Consequently, $PM_{2.5}$ pollution in the SCB occurred on January 1–6 and rapidly dissipated on January 7 (Fig. 7). The mean geopotential height at 850 hPa derived from the simulation of January 1–3 under Pattern 2 showed EHWL, with southerly flow prevailing over the SCB (Fig. 7a). The resulting upper air warming suppressed PBL development (Fig. 7d). During January 1–3 under Pattern 2, the average PBL heights were lower (Fig. 7d), acting as a lid above the SCB and hindering the airflow within the basin due to the surrounding mountains. Low wind speeds provided adverse diffusion conditions for pollutants emitted into the basin, resulting in severe pollution in the western and southern SCB (Fig. 7g). From January 4 to 6, the low pressure over the SCB evolved into a LT pattern, termed Pattern 5 in the previous analysis. Compared with Pattern 2, the isobaric lines were denser under the influence of the LT, leading to stronger southerly winds above the SCB (Fig. 7a–b). Lower average PBL heights were observed during January 4–6 under Pattern 5 compared with those of January 1–3 under Pattern 2 (Fig 7d–e), primarily due to stronger upper air warming and more stable stratification. Pollutants that accumulated during January 4–6 from the earlier pollution episode (January 1–3) further increased (Fig. 7g–i). On January 7, high pressure in the north dominated the SCB, with a prevailing northerly flow over the basin (Fig. 7c). The PBL height quickly increased due to upper-layer cold advection (Fig. 7f), resulting in a rapid decrease in $PM_{2.5}$ (Fig. 7i). Overall, synoptic patterns play a key role in the accumulation and diffusion of $PM_{2.5}$ during pollution episodes by modulating PBL development and stratification stability.

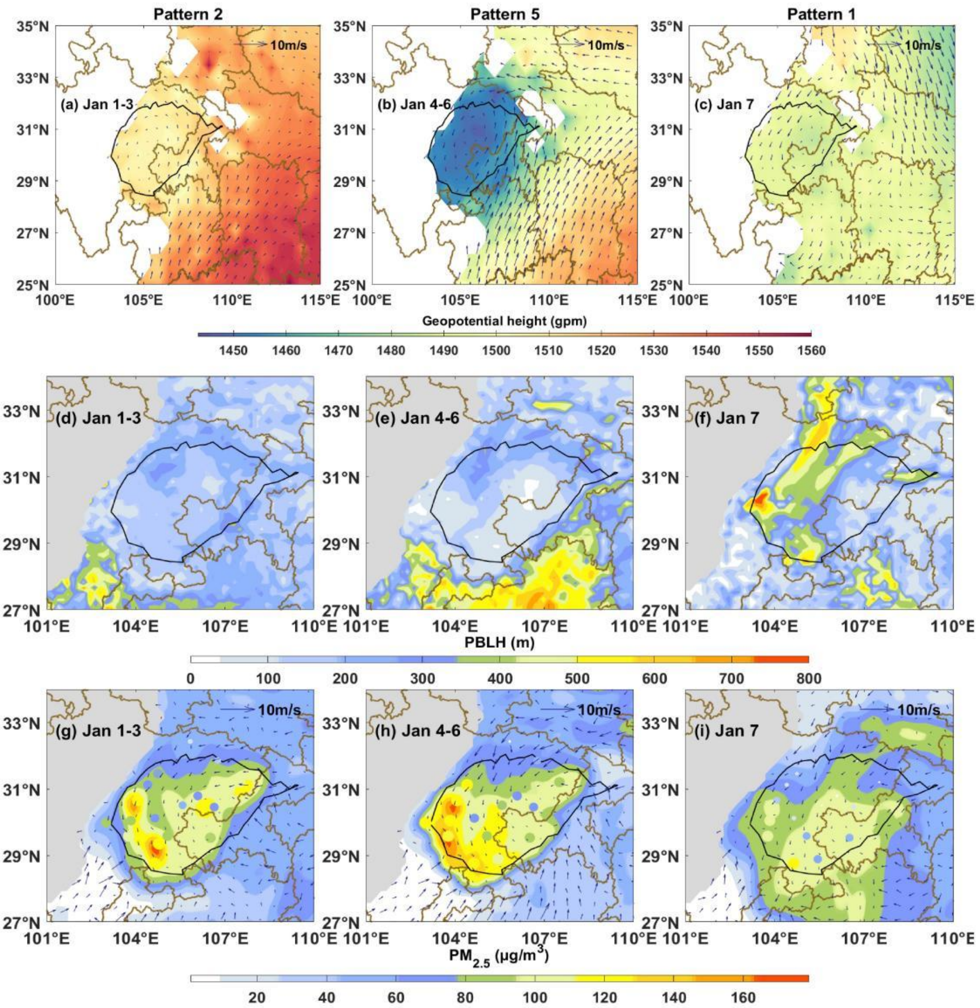

Figure 7. The simulated (a)-(c) 850 hPa geopotential height field (shading) with wind vector fields (blue vectors), (d)-(f) boundary layer height and (g)-(i) $PM_{2.5}$ concentrations (shading) and wind vector fields at 900hPa (blue vectors) for 1-3, 4-6 and 7 January. The size and color of scatters in (g)-(i) show corresponding observed $PM_{2.5}$ concentrations at 18 air quality monitoring stations. The SCB was outlined with an altitude contour of 750 m terrain height (black lines).

Pollutant accumulation can regulate the PBL structure through the ARI, further exacerbating pollution (Wang et al., 2018; Miao et al., 2020). In the SCB, this positive feedback is weaker than in the other regions and may be inhibited by cloud radiation (Zhong et al., 2019). A series of simulation experiments were conducted to investigate the aerosol radiation feedback in the SCB under the influence of two typical synoptic pollution patterns, as described in Section 2.4. BASE–EXP1 represents the perturbations caused by ARI, whereas EXP2–EXP3 demonstrates changes through ARI without CRI inhibition. Aerosols led to surface cooling through absorbing and scattering solar radiation, thereby inhibiting the development of the PBL, which in turn facilitated pollutant accumulation (Fig. 8). Compared with Pattern 2, the aerosol concentrations in Pattern 5 were higher, resulting in greater reduction of downward solar radiation reduction due to ARI, leading to more pronounced cooling near the ground and a lower PBLH. Overall, the ARI in Pattern 5 was more significant than that in Pattern 2, regardless of CRI inhibition (Fig. 8).

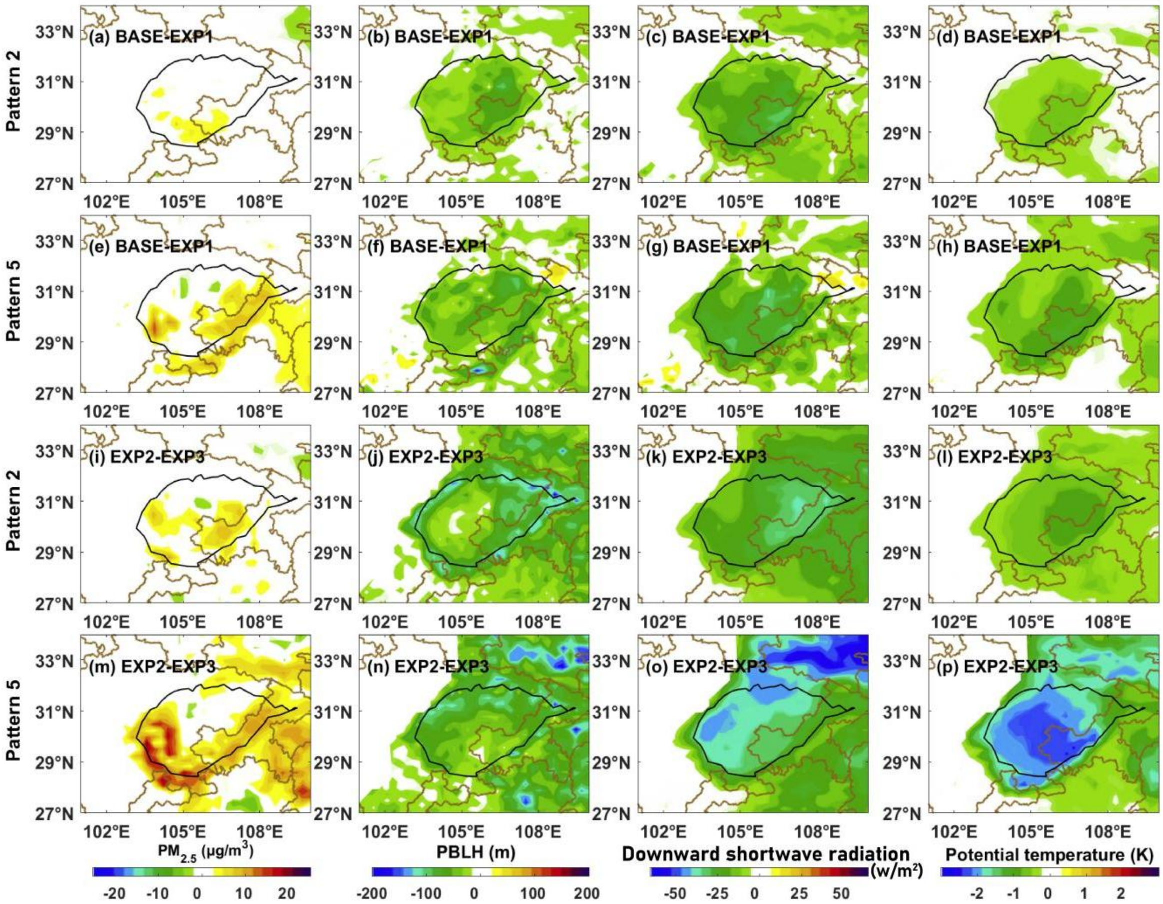

Figure 8. Spatial distribution of perturbations induced by (a)-(h) aerosol radiation interactions (ARI), and (i)-(p) ARI without the cloud radiation interaction (CRI) inhibition during 1-3 and 4-6 January representing Pattern 2 and Pattern 5 synoptic forcing, respectively. The SCB was outlined with an altitude contour of 750 m terrain height (black lines).

Furthermore, parallel simulation experiments revealed that the CRI significantly attenuated the ARI in the SCB under both pollution synoptic patterns. When the CRI was not considered, more solar radiation penetrated the PBL. Dense aerosols accumulated near the surface, intercepting more downward shortwave radiation, and resulting in stronger cooling near the ground. This suppressed the development of the PBL, and contributed to a more remarkable ARI (Fig. 8). Regarding the horizontal spatial distribution, a strong ARI was primarily observed in CQ, as well as in the western and southern SCB, despite CQ experiencing lower pollutant concentrations compared to the other two regions (Figs. S4 and 7). This weaker ARI phenomenon in the western SCB was also reported by Zhong et al. (2019) and was attributed to CRI inhibition on ARI. Considering the statistical results in Fig. 5, the average cloud liquid water content in CD and YB was significantly higher than that in CQ under the influence of Patterns 2 and 5. Consequently, a more remarkable CRI inhibition on the ARI would occur in the western and southern SCB compared to CQ, leading to a relatively weaker ARI distribution in these regions. Without considering the CRI, the ARI in the western and southern SCB would be much more pronounced than that in CQ. As for the northwestern SCB (DZ), the ARI in DZ is lower than in the other three regions. When the CRI is not considered, the ARI in DZ is higher than in CQ but lower than in CD and YB. This is because DZ has lower aerosol concentrations compared to CD and YB (Fig. 7), but exhibits higher cloud cover than CQ under Patterns 2 and 5 (Fig. 5).

Using the western SCB, which exhibited the highest pollution concentration, as an example, Fig. 9 illustrates the vertical diurnal variations in temperature and solar radiation caused by the ARI. The results in Fig. 9–11 were derived from the simulation experiments in CD, as CD is one of the most polluted cities with typical meteorological and geographical characteristics of the western SCB. The ARI caused surface cooling in the morning and upper–air warming in the afternoon. As local solar radiation increased from 8 am to 12 pm, the reduction in solar radiation caused by the ARI also increased. Surface cooling reached its peak at approximately 10 am to 12 pm, and gradually weakened in the afternoon. This diurnal variation might be attributed to the enhanced turbulence during morning PBL evolution (Wang et al., 2018). Afternoon surface cooling was partly compensated by the turbulent transport of warm air above the PBL. In addition, strong surface cooling between 5 pm and 8 pm in the SCB, was possibly influenced by remarkable valley wind circulations forced by the Qinghai–Tibet Plateau adjacent to the western SCB (Lu et al., 2022). The evening cooling of the plateau induced strong mountain winds, promoting surface cooling, while the upper–layer warming mainly occured around 1–1.5 km in the afternoon. In general, the ARI reduces solar radiation, causing surface cooling and upper air warming, thereby regulating the vertical atmospheric thermal structure, suppressing convection, and consequently decreasing PBL heights (Fig. 10).

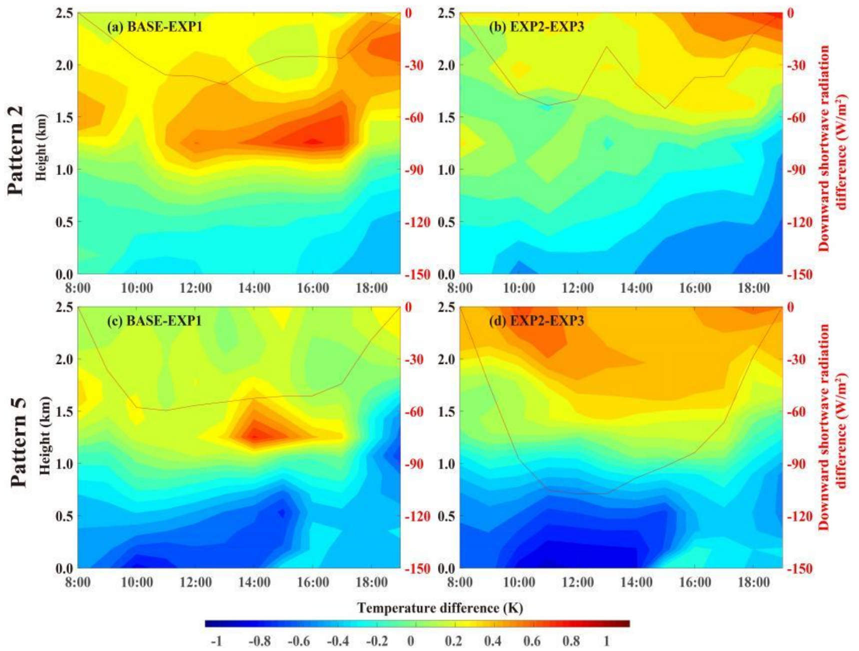

Figure 9. Diurnal variations of vertical temperature perturbations and downward solar radiation under influences of Pattern 2 and Pattern 5 induced by (a) (c) ARI and (b) (d) ARI without CRI inhibition.

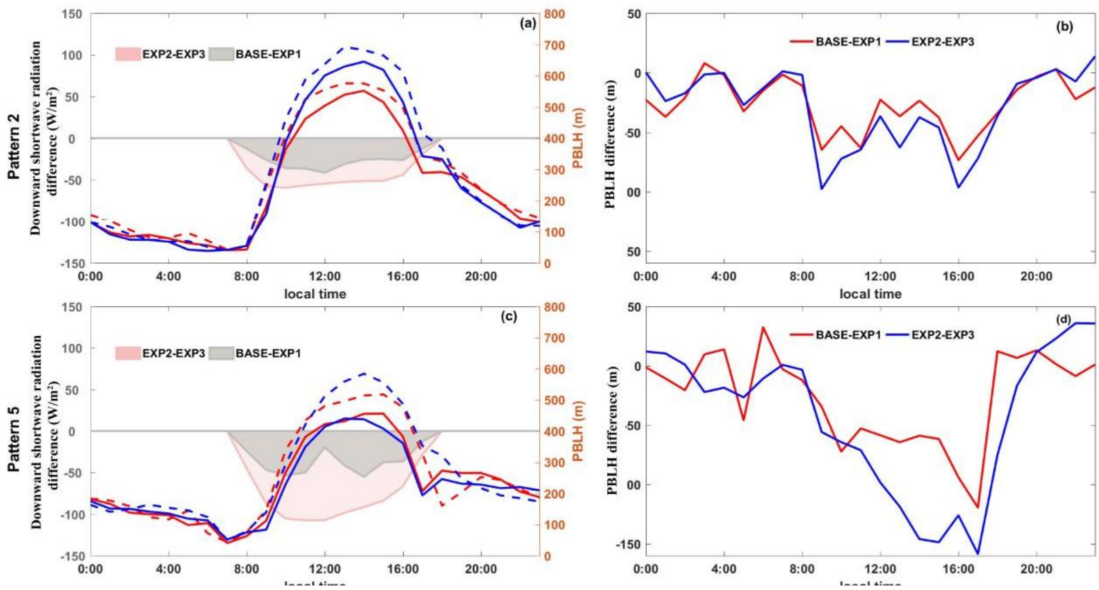

Figure 10. Diurnal variations of (a) (c) boundary layer height(lines) and downward solar radiation(shading), and (b)(d)the perturbations of boundary layer height induced by ARI and ARI without CRI inhibition, under Pattern 2 and Pattern 5 synoptic forcing respectively.

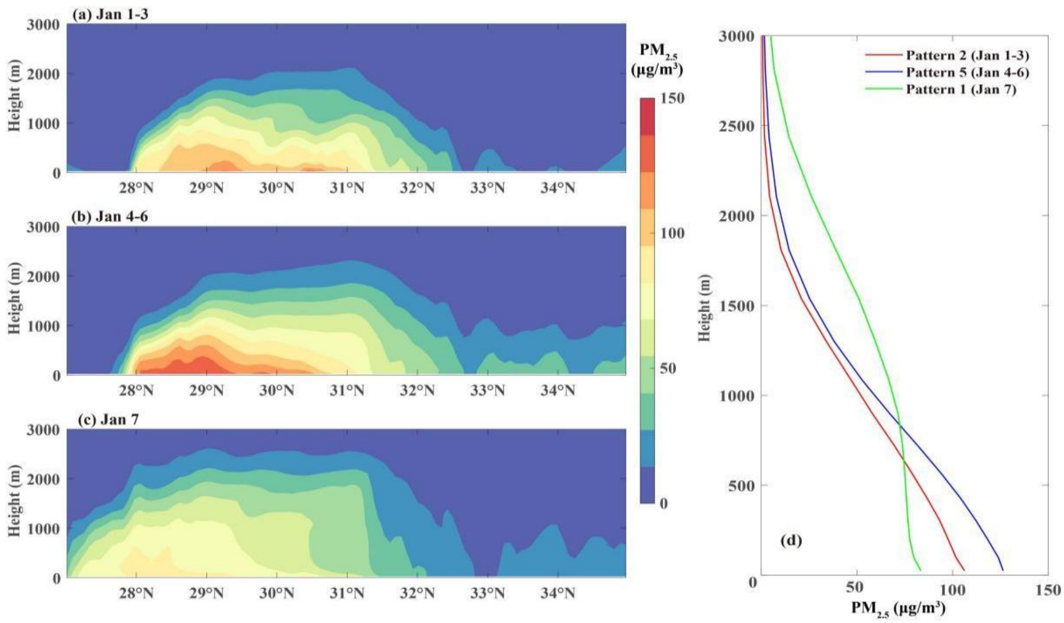

Figure 11. Meridional vertical distribution of averaged $PM_{2.5}$ between 104°E and 105°E under (a) Pattern 2 and (b) Pattern 5, and (c) average profiles of $PM_{2.5}$ within 28°N and 31°N.

Synoptic patterns play a role in the interaction between the ARI and PBL (Wang et al., 2018; Miao et al., 2020). Zonal average of $PM_{2.5}$ concentration between 104°E and 105°E was conducted, and the meridional vertical distribution of $PM_{2.5}$ between 27°N and 35°N was illustrated in Fig. 11a-b. Fig. 11(c) provides an average of $PM_{2.5}$ concentration within 28°N and 31°N, showing the vertical distribution profiles under Pattern 2 and 5. Due to the inhibition of "warm lid" above the SCB, the vertical exchange was not prominent under both Pattern 2 and 5, and $PM_{2.5}$ was more concentrated at the middle and lower levels. The $PM_{2.5}$ concentration under Pattern 5 was higher than Pattern 2 throughout the atmospheric column, indicating stronger aerosol radiative forcing and a more significant impact on the boundary layer structure under Pattern 5. During January 4–6, the surface cooling reached 1 K, with cooling layers higher than those observed on January 1–3. The differences in thermal structure modulations contributed to a lower diurnal PBLH in Pattern 5 than in Pattern 2 (Figs. 10a and c), indicating that Pattern 5 was more conducive to ARI. Based on the simulation experiments, this study further discussed the impact of synoptic forcing on the CRI inhibition of ARI. When the CRI was not considered, the solar radiation reduction at noon on January 4–6 by the ARI was nearly twice as high as when the CRI was considered. Correspondingly, surface cooling at noon was remarkably enhanced. In the evening, surface cooling occurred earlier and was stronger without the CRI (Fig. 9). The regulation of CRI on ARI was further reflected in changes in PBLH. Without the CRI, the diurnal PBLH increased significantly, with the PBLH decreased more with ARI without CRI inhibition. The PBLHs were decreased by the ARI during January 13–17 afternoon, reaching 2–3 times the decrease observed with CRI inhibition (Fig. 10). More significant CRI inhibition of ARI was revealed under Pattern 5 compared with that under Pattern 2, owing to the stronger ARI itself with higher aerosol concentrations in Pattern 5 and the more apparent CRI inhibition with denser cloud liquid water contents under the LT pattern (Fig. 5). Therefore, the intensity of CRI inhibition of ARI in the SCB was altered by synoptic forcing, with stronger effects under the influence of LT.

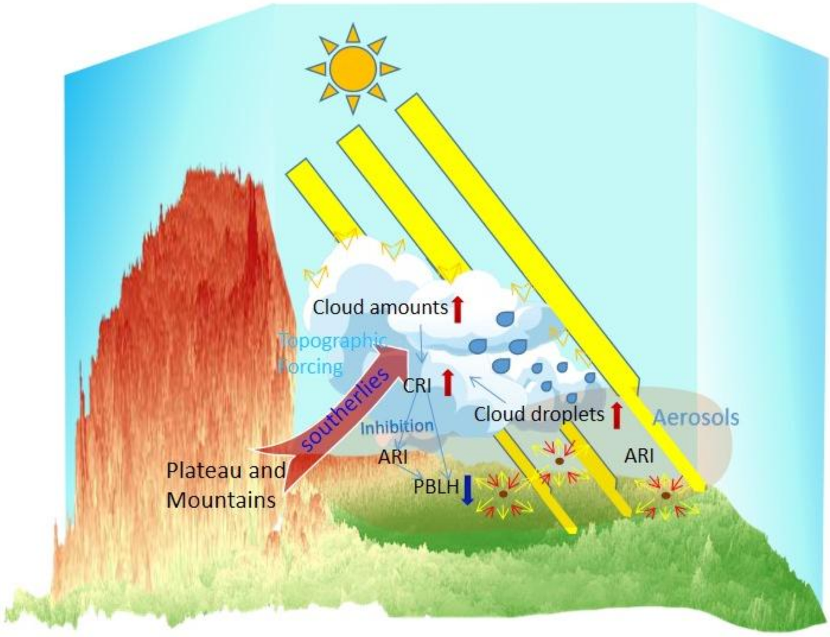

Figure 12. The aerosol radiative effect can be signigicantly inhibitted by cloud under influence of pollution synoptic patterns with dense cloud.

4 Conclusion

This study utilized synoptic classification and numerical simulation to gain insights in to the combined effects of synoptic patterns and CRI inhibition on ARI and PBL structures in the wet and cloudy SCB. Based on the long–term $PM_{2.5}$ observations and sounding data in the SCB, it was found that large-scale synoptic circulations at 850 hPa played crucial roles in the variations of $PM_{2.5}$ pollution. Synoptic classification was performed with the T-PCA method, which reveaed that Pattern 2 and 5 characterized with low pressure system and southerly airflow on 850 hPa were key synoptic patterns for onset and accumulation of $PM_{2.5}$, while Pattern 1 controlled by the northerly airflow represented a clean pattern associated with significant decrease in $PM_{2.5}$. Moreover, it was indicated that Pattern 2 and 5 exhibited denser cloud liquid water content and thus stronger compared to other patterns. Among these patterns, Pattern 5 exhibited the highest cloud liquid water content and CRI. This could be attributed to the robust southerly airflow induced by the dense isobaric lines, which brought warm and humid air masses into the region.

To illustrate the interactions among cloud, aerosol and PBL under pollution synoptic patterns, a pollution episode occurred from January 1 to 7, 2017, was simulated with using WRF-Chem. The simulation results showed that ARI remarkably reduced solar radiation was during the two pollution patterns. This reduction led to surface cooling in the morning and upper–air warming in the afternoon. Additionally, the enhanced evening surface cooling was impacted by the mountain–valley wind circulations forced by the plateau–basin topography of the SCB. This modulation in the vertical thermal structure by the ARI would then suppress the development of the PBL, favoring pollution outbreaks (Fig. 12). Furthermore, parallel simulation experiments indicated that CRI impacted stratification stability and modulated the vertical thermal structure by inhibiting ARI (Fig. 12). Regarding the spatial distribution, a stronger ARI appeared in Chongqing, despite lower $PM_{2.5}$ concentrations compared to the western and southern SCB. This was due to the lower cloud liquid water content and weaker CRI inhibition of ARI in Chongqing. When CRI inhibition was not considered, the ARI in the western and southern SCB was significantly stronger than that in Chongqing. In addition, under Pattern 5, the reduction in solar radiation and PBLH during the daytime due to ARI could be more than doubled when the CRI influence was neglected. This was primarily due to higher aerosol concentrations and cloud liquid water contents associated with a low trough in Pattern 5. This study provided insights into the interaction among aerosols, clouds, and PBL under different synoptic patterns, considering the complex terrain and foggy/cloudy climate of the SCB. The findings highlighted the significant role of CRI inhibition on ARI during wet and cloudy conditions, shedding light in the multi-scale atmospheric physical processes in the SCB.

Author contributions. HL and MX had the original idea for the study, designed the experiments, conducted the numerical simulation and prepared the initial draft manuscript. BL, YZ and KZ collected the data. TW and BZ helped perform the analysis with constructive discussions, reviewed and edited the manuscript. HL, MX, TW and BL acquired financial support for the project leading to this publication. SL and ML reviewed the manuscript.

Competing Interest: The authors declare no conflict of interest.

Data Available Statement

The ERA5 pressure layer and single layer data can be respectively downloaded from https://cds.climate.copernicus.eu/cdsapp#!/dataset/reanalysis-era5-pressure-levels and https://cds.climate.copernicus.eu/cdsapp#!/dataset/reanalysis-era5-single-levels?tab=form. The NCEP FNL data are available at https://rda.ucar.edu/datasets/ds083.2/. The MEIC data can be accessed in Zheng et al (2018) at https://doi.org/10.5194/acp-18-14095-2018. Air quality and meteorological monitoring data can be acquired from https://doi.org/10.7910/DVN/USX59F.

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