

# Temporal and spatial variations in dust activity in Australia based on remote sensing and reanalysis data sets

Yahui Che<sup>1</sup>, Bofu Yu<sup>1</sup>, Katherine Bracco<sup>1</sup>

<sup>1</sup>School of Engineering and Built Environment, Griffith University, Brisbane, 4111, Australia

Correspondence to: Bofu Yu (b.yu@griffith.edu.au)

**Abstract.** Spatial and temporal variations in the level of dust activity can provide valuable information for policy making and climate research. Recently, MODIS aerosol products have been successfully used for retrieving dust aerosol optical depth (DAOD), especially over bright dust source areas and MERRA-2 aerosol reanalysis provides DAOD, and additionally other dust aerosol-related parameters. In this study, spatial and temporal variations in dust activity in Australia were analyzed using MODIS and MERRA-2 combined (M&M) DAOD and MERRA-2 near-surface dust concentrations/estimated PM<sub>10</sub> for the period from 1980-2020. Validation results show that M&M DAOD has an expected error of  $\pm(0.016 + 0.15\tau)$  compared to the ground observations at the AERONET sites. MERRA-2 near-surface dust concentrations show a power law relationship with visibility data collected at meteorological stations with an  $r^2$  value from 0.18 to 0.44, and the estimated MERRA-2 PM<sub>10</sub> shows similar temporal variations and correlates with ground-based PM<sub>10</sub> data with an  $r^2$  value from 0.14 to 0.44 at six selected stations in Australia. Moreover, MERRA-2 horizontal dust flux shows the same major dust pathways as those in previous studies and similar dust emissions/deposition areas identified using ground-based observations. Dust events based on DAOD over eastern Australia are concentrated in the north in December, in the south in February, and can occur anywhere in January. Near-surface dust concentration was found to be the highest (over  $200\mu\text{g}/\text{m}^3$ ) over the center of Lake Eyre Basin in central Australia and radially decreased to the coast to below  $20\mu\text{g}/\text{m}^3$  via the two main pathways in the southwest and northeast. The ratio of near-surface dust concentration to PM<sub>10</sub> shows a similar spatial pattern. Total dust emission was estimated to be 40 MT (mega-tonnes) per year over the period 1980-2020, of which nearly 50% was deposited on land and the rest exported away from the Australian continent.

## 1. Introduction

Dust storms, as a natural hazard, occur frequently in Australia, especially in the central inland area, which is identified as the largest dust source in the Southern Hemisphere (Shao, 2009; McTainsh et al., 2011b; Ekström et

28 al., 2004; McTainsh et al., 2011a), contributing to approximately 5% of the global total dust emissions (Shao,  
29 2009; Wu et al., 2020; Chen et al., 2022). Dust sourced from the Lake Eyre Basin is not only deposited on the  
30 Australian continent but also transported to the Tasman Sea in the southeast and the Indian Ocean in the northwest  
31 (Strong et al., 2011; Bowler, 1976; Sprigg, 1982a; Speer, 2013; Ekström et al., 2004; Shao et al., 2007). The  
32 adverse impacts of dust storms on populated areas include incalculable economic loss in agriculture (Stefanski  
33 and Sivakumar, 2009) and household cleaning and associated activities (Tozer and Leys, 2013), human health  
34 issues such as respiratory problems (Roberts, 2013; Chen et al., 2007; Cowie et al., 2010; Goudie, 2014; Middleton,  
35 2017) and cardiovascular disease (Domínguez-Rodríguez et al., 2021; Zhang et al., 2016), and contamination of  
36 water sources (Middleton, 2017). Moreover, the Australian dust over the south-west Pacific Ocean strengthens the  
37 relationship between rainfall and the El Niño Southern Oscillation (ENSO) by driving ENSO-related anomalies  
38 in radiative forcing (Rotstayn et al., 2011), which is the direct dust feedback to climate (Shao et al., 2013).

39

40 The severity of dust activities can be indicated by a range of different approaches, including visibility-based dust  
41 event days/dust storm index (DSI) (Yu et al., 1992, 1993; McTainsh et al., 2011b; O’Loingsigh et al., 2017) and  
42 total suspended dust concentration (Shao et al., 2013; McTainsh et al., 2005; Tews, 1996; Baddock et al., 2014),  
43 aerosol optical depth (AOD)/DAOD (dust AOD) (Ginoux et al., 2010; Pu and Ginoux, 2018; Yu and Ginoux,  
44 2021; Ginoux et al., 2012; She et al., 2018) and dust index (Di et al., 2016; Yang et al., 2023; Bullard et al., 2008),  
45 simulated near-surface dust concentration (Prospero et al., 2020; Buchard et al., 2017), and PM10 (particles with  
46 a diameter of 10 micrometers or less) (Leys et al., 2011; de Jesus et al., 2020). The dust event database  
47 (DEDDB)/DSI has been widely used in dust and wind erosion research in Australia, benefiting from the long-term  
48 temporal data gathered at widely distributed Bureau of Meteorology (BoM) sites (McTainsh and Pitblado, 1987;  
49 McTainsh et al., 2011a; O’Loingsigh et al., 2014). Horizontal visibility has also been used for estimating the dust  
50 concentration and dust loading for large dust storms (McTainsh et al., 2005) and visibility-based dust  
51 concentration has been even used for exploring the climate forcing of dust at the global scale (Shao et al., 2013).

52

53 With the development of satellite remote sensing and numerical dust models, remote sensing and General  
54 Circulation Model (GCM) products have been increasingly applied to dust research with regard to spatial extent  
55 detection, columnar optical properties, and near-surface concentrations. Dust indices based on satellite images can  
56 be traced back to the detection of dust storms using Advanced Very High-Resolution Radiometer (AVHRR) data

57 (Ackerman, 1989), taking advantage of AVHRR's large spatial coverage. So far, several different dust indexes  
58 have been developed for regional or global dust detection and different sensors (Yang et al., 2023). Satellite data  
59 retrieved AOD/DAOD have been more frequently applied to quantitative dust research since AOD can be  
60 successfully retrieved over bright dust source areas (Hsu et al., 2004; Ginoux et al., 2010; Baddock et al., 2009).  
61 Benefiting from satellite providing dust source schemes (Ginoux et al., 2001), near-surface dust concentrations  
62 have been simulated for specific regions or on a global scale (Gelaro et al., 2017; Buchard et al., 2017; Shao et  
63 al., 2007; Wu et al., 2020).

64

65 Rare analyses of long-term AOD/DAOD data have been attempted for Australia. AOD/DAOD was mostly used  
66 for identifying the spatial extent of single dust events or as reference data for evaluating dust detection algorithms  
67 in Australia. For example, Baddock et al. (2009) assessed the performances of four detection algorithms based on  
68 Moderate Resolution Imaging Spectroradiometer (MODIS) L1 (Level 1) B data and MODIS Deep blue (DB)  
69 AOD for central Australia (i.e. the Lake Eyre Basin) on identifying airborne dust and mineral aerosols. There are  
70 a few dust studies on analyzing seasonal spatial variations of dust using a multi-year AOD/DAOD dataset. Ginoux  
71 et al. (2012) retrieved global DAOD using the MODIS DB aerosol dataset from 2003 to 2009 and analyzed major  
72 anthropogenic and natural dust emissions in Australia. Their results show that the contribution to total emissions  
73 by anthropogenic activities can be as high as 75% in Australia. Yu and Ginoux (2021) show the monthly MODIS  
74 DB DAOD and Multi-angle Imaging SpectroRadiometer (MISR) coarse mode AOD at 15 AERONET sites and  
75 the annual DAOD and coarse mode AOD in Australia from 2000 to 2019. A comparison with DSI shows that  
76 satellite AOD/DAOD presents the same dusty month/season as that by DSI at three AERONET sites in the main  
77 Australian dust source area. Yang et al. (2021) show similar AERONET coarse mode AOD variations and seasonal  
78 contribution of dust to total aerosols at nine AERONET sites and analyze seasonal DAOD from **the Modern-Era**  
79 **Retrospective analysis for Research and Applications, Version 2** (MERRA-2) aerosol reanalysis from the early  
80 2000s to 2020. There are limitations **to this type of investigation using MODIS and MERRA-2 data**. Firstly,  
81 MODIS DB retrieved DAOD shows much smaller coverage than the original DB AOD due to excluding low  
82 background DAOD, possibly resulting in an overestimation of dust activity severity. Secondly, MERRA-2 is very  
83 likely to have underestimated DAOD over 0.2, especially for severe dust storms such as those on the 23<sup>rd</sup> of  
84 October 2002 and the 23<sup>rd</sup> of September 2009 (Che et al., 2022).

85

86 PM10 is often taken as an effective indicator of dust severity for single dust storms (Leys et al., 2011; McGowan  
87 and Clark, 2008); however, long-term analysis of dust severity for Australia is of great difficulty using PM10 data.  
88 First, PM10 observations in each state mostly began after 2000 in populated urban areas while the dust source  
89 area in central Australia lacks PM10 observations. This spatial distribution of PM10 sites also causes difficulties  
90 in retrieving PM10 in Australia using satellite AOD. Second, little progress has been made in retrieving PM10 for  
91 large regions based on remote sensing products in Australia because 1) a reliable estimate of PM10 is difficult to  
92 obtain due to the relatively low dust concentrations approaching its retrieval uncertainty and 2) the inclusion of  
93 AOD and related predictors cannot improve the accuracy of simulated PM10 (Pereira et al., 2017). Therefore,  
94 trend analyses of PM10 concentrations in Australia have predominantly focused on site-based observations. For  
95 example, de Jesus et al. (2020) analyzed PM10 trends in major cities of Australia over the last two decades using  
96 site PM10 observations; however, they did not conduct a trend analysis specifically for dust concentrations  
97 because the dust component could not be accurately obtained from PM10, which consists of both dust and non-  
98 dust particles from multiple sources. In contrast to PM10, total near-surface suspended dust concentration is  
99 capable of indicating the severity of dust events with all ranges of particle sizes, while PM10 observations only  
100 include particles smaller than 10 $\mu$ m. For example, Love et al. (2019) analyzed a 17-year (1990-2007) near-surface  
101 dust concentration data collected by a high-volume air sampler (HVS) in Mildura. To provide a more specific  
102 analysis of dust particle concentrations, Prospero et al. (2020) removed soluble sea salt particles from collected  
103 HSV samples. Long-term dust analysis studies such as this are relatively few compared to those based on  
104 visibility-transferred dust concentrations. Nevertheless, the relationship between horizontal visibility and total  
105 suspended dust concentration varies among different studies. As noted by McTainsh et al. (2005), the relationships  
106 between visibility and dust concentration obtained in the United States of America (USA) (Chepil and Woodruff,  
107 1957) and Australia in the 1990s (Tews, 1996) have been inappropriate for estimating dust concentration over  
108 different areas of Australia, since visibility-based dust concentrations are strongly influenced by dust particle size.  
109 Considering that WMO weather stations are often not located in dust sources where previous visibility-dust  
110 concentration relationships were established, Baddock et al. (2014) developed an empirical model for calculating  
111 total dust concentration (in Mildura) 10-100km from the dust source (in Buronga).

112

113 The development of numerical dust models and GCMs provides dust cycle simulations for understanding the  
114 impacts of dust on the earth systems. Models such as the Georgia Tech/Goddard Global Ozone Chemistry Aerosol

115 Radiation and Transport (GOCART) (Ginoux et al., 2001), Mineral Dust Entrainment and Deposition (DEAD)  
116 (Zender, 2003), and Aerosol Species IN the Global AtmospheRe (MASINGAR) coupled with MRI/JMA 98 GCM  
117 (Tanaka and Chiba, 2006) have all proved capable of simulating the dust cycle for different regions around the  
118 world. However, dust emissions and depositions vary substantially among models (Chen et al., 2022; Wu et al.,  
119 2020). This would directly lead to a large discrepancy in conclusions based on different models. For example, the  
120 contribution of Australian dust to global dust emissions is estimated to vary from 0.02% to 27.8% using simulation  
121 outputs from 15 CMIP5 (Coupled Model Intercomparison Project Phase 5) models (Wu et al., 2020). Therefore,  
122 long-term analysis of the dust cycle in Australia needs a dataset with high accuracy and the capability to quantify  
123 long-term trends and variabilities, such as MERRA-2 aerosol product.

124  
125 Numerous validation studies have shown that MERRA-2 optical depth and near-surface concentrations could be  
126 used for temporal and spatial analysis and even long-term analysis of aerosols regionally or globally due to the  
127 high quality and long temporal coverage from 1980 to the present. This kind of study primarily focuses on the  
128 validation of MERRA-2 AOD with ground-based AOD from AERONET (Bucharad et al., 2017; Sun et al., 2019a;  
129 Che et al., 2022; Randles et al., 2017), SONET (Sun-Sky Radiometer Observation Network) (Ou et al., 2022),  
130 SKYNET (Sun et al., 2019b). For dust research in Australia, MERRA-2 AOD has been validated/evaluated with  
131 AERONET (Che et al., 2022; Mukkavilli et al., 2019) and MODIS DB dataset (Che et al., 2022), as well as its  
132 DAOD with the MACC (Monitoring Atmospheric Composition and Climate) simulation (Mukkavilli et al., 2019)  
133 and MODIS DB retrieved DAOD (Che et al., 2022). However, few studies were carried out to validate MERRA-  
134 2 near-surface dust concentrations.

135  
136 MODIS DB and MERRA-2 data products, dust aerosols, and near-surface dust concentration/PM10 observations  
137 were analyzed in this study for a better understanding of long-term dust entrainment and transport over Australia.

138 Objectives of this study include:

- 139 1. To develop a DAOD dataset using MERRA-2 aerosol reanalysis and MODIS DB aerosol datasets;
- 140 2. Validate MERRA-2 near-surface dust concentrations using ground-based visibility data sets, and  
141 MERRA-2 estimated PM10 with ground-based PM10 observations sourced from the New South Wales  
142 Air Quality Monitoring Network (NSW AQMN);
- 143 3. To corroborate MERRA-2 horizontal dust flux with major dust pathways identified in previous studies;

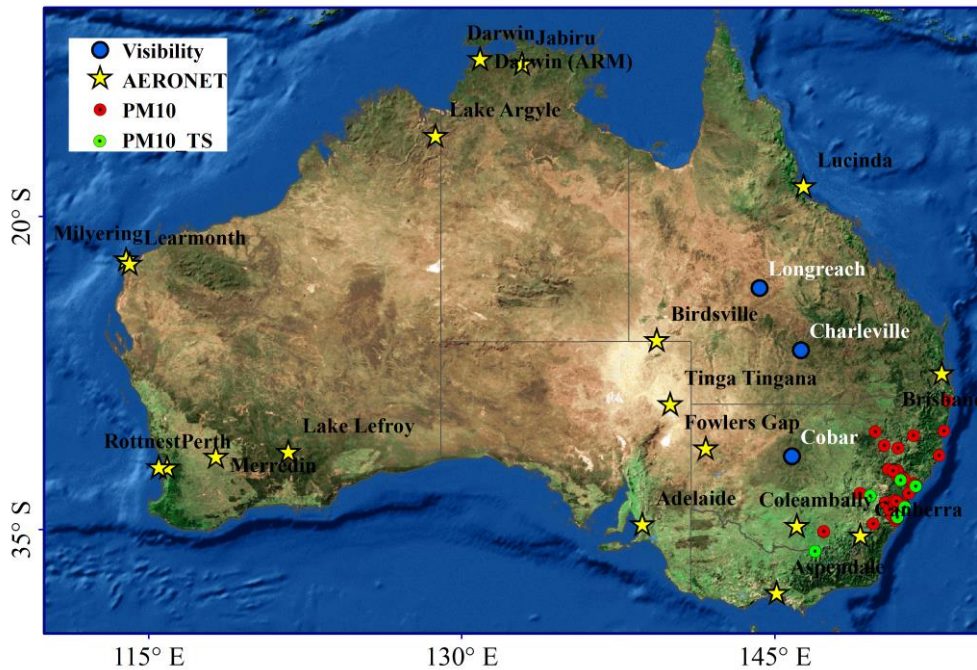
- 144 4. To map the seasonal MODIS and MERRA-2 (M&M) DAOD from 2002 to 2020 and seasonal MERRA-  
145 2 near-surface dust concentrations/PM10 over the period from 1980 to 2020;
- 146 5. To quantify the annual dust cycle for Australia over the period from 1980 to 2020, including dust emission  
147 in Australia using MERRA-2 emission data, dust import and export using MERRA-2 flux data, and dust  
148 deposition using MERRA-2 emission and flux data.

## 149 **2. Data and methodology**

### 150 **2.1 Ground-based PM10 and AERONET data**

151 AOD at 440nm and Ångström exponent (AE) at 440-675nm from AERONET v3 solar products were used for  
152 calculating AOD at 550nm, as well as Level 1.5 (L1.5) single scattering albedo (SSA) at 440nm from v3 inversion  
153 product to retrieve DAOD. The latest AERONET v3 solar product includes data at Level 1.0 (L1.0) (without data  
154 screening), L1.5 (with cloud screened and quality controlled), and Level 2.0 (L2.0) (quality assured)  
155 (<https://aeronet.gsfc.nasa.gov/>). Giles et al. (2019) reported that AOD from the AEROENT V3 product had a low  
156 uncertainty, suggested by a bias of +0.02 and one sigma uncertainty of 0.02. Since satellite AOD normally refers  
157 to that at 550nm, AE is necessarily used for spectrally interpolating AOD to this wavelength according to the  
158 dependence of AOD on wavelength (Angstrom, 1924). The Version 3 (V3) inversion product also includes data  
159 at three levels, L1.0, L1.5, and L2.0. The main difference between L1.5 and L2.0 is that the L2.0 inversion product  
160 is only made when the corresponding AOD is higher than 0.4 (Dubovik and King, 2000). This leads to the data  
161 volume of L2.0 SSA being far smaller than that of L1.5, especially over Australia predominated by low AOD  
162 conditions. L1.5 SSA data with a much larger data volume, therefore, was used for identifying dust-contaminated  
163 AOD. In this study, all L2.0 AOD and AE, L1.5 SSA, as well as fine mode fraction (FMF) in the Spectral  
164 Deconvolution Algorithm (SDA) database in Australia from 1997 to 2020 were used for retrieving DAOD at  
165 550nm (yellow stars in Fig.1).

166



167

168

169

170

171

172

173

174

175

176

177

178

179

180

181

182

183

184

185

**Figure 1: Distribution of ground-based sites. Yellow stars are inland AERONET sites in Australia; Green and red circles indicate PM10 observation sites for validating MERRA-2 products, among them time series analysis was conducted at green sites; Circles in blue are visibility observation sites.**

Horizontal visibility records from the BoM are manually estimated and noted by a weather observer and visibility is also automatedly measured with a visibility meter at automatic weather stations (AWS) from the early 2000s to the present. All manual and AWS visibility records have an upper limit of 10km, which means that if the record shows a visibility of 10km the actual visibility could be 10km or greater. In addition to visibility, a synoptic code is recorded during notable weather events such as dust storms (Baddock et al., 2014) as well as a weather type for the weather observation immediately prior to the current observation (O’Loingsigh et al., 2010). There are 11 SYNOP (surface synoptic observations) codes for dust weather, including dust haze, raised dust or sand, dust whirls, thunderstorms with sand or dust storms, dust storms, and so on. Due to the visibility record only allocated one SYNOP code, the most important weather code (the higher the code number, the more important the code) was retained although it may have included several weather types (O’Loingsigh et al., 2010). The long-term visibility records with SYNOP codes have been widely used for wind erosion research (McTainsh et al., 1989, 1990, 2011a), dust event climatologies including the DEDB and DSI developed at Griffith University (McTainsh et al., 2011b; O’Loingsigh et al., 2014, 2010) and single dust events (Shao et al., 2007; McTainsh et al., 2005; Leys et al., 2011). Manual observations of dust activities (using horizontal visibility as a measure), span from the



186 early 19th century to present (O'Loingsigh et al. 2017). In this study, to ensure the consistency of data points to  
187 dust storm research, hourly BoM visibility observations of less than 10km with a dust SYNOP code (excluding  
188 thunderstorms with raised dust) were used. These observations were used to validate MERRA-2 near-surface dust  
189 mass concentration at three sites: Charleville, Cobar, and Longreach (Figure. 1) from 1980 to 2020

190  
191 PM10 concentrations are publicly downloadable from the NSW AQMN website (<https://www.dpie.nsw.gov.au/>).  
192 AQMN used a Tapered Element Oscillating Element instrument (TEOM), measuring atmospheric particles < 10  
193  $\mu\text{m}$ . The filter in the TEOM weighs collected samples every 2 seconds and the average value is reported hourly.  
194 Data quality control is applied to all PM10 databases according to Australian Standard 3580.9.8 (Leys et al., 2011).  
195 All PM10, gases, and climate observations data can be accessed using the AQMN web data download facility  
196 (<https://www.dpie.nsw.gov.au/air-quality/air-quality-data-services/data-download-facility>). In this study, monthly  
197 PM data from 62 AQMN urban sites (circles in green and red in Fig.1) were used for validating MERRA-2 near-  
198 surface dust concentrations from 2001 to 2020.

## 199 **2.2 MERRA-2 aerosol reanalysis**

200 MERRA-2 aerosol reanalysis provides long-term global aerosol parameter datasets from 1980 to the present  
201 (Randles et al., 2017; Gelaro et al., 2017; Buchard et al., 2017). MERRA-2 includes optical depth, near-surface  
202 mass concentrations, column mass concentrations, and horizontal mass flux in u/v-wind directions flux for each  
203 aerosol component, including sea salt, sulfate in ( $\text{SO}_4$  and  $\text{SO}_2$ ), organic carbon, dust, and black carbon. Due to  
204 the inclusion of the GEOS and GSI assimilation system, MERRA-2 aerosol simulations perform comparably with  
205 high-quality satellite-based datasets and are fairly close to aerosol observations. Benefiting from the incorporation  
206 of space-based and ground-based observations, the accuracy for the total AOD is guaranteed with physical models  
207 outputs constrained by the assimilation system  
208 ([https://gmao.gsfc.nasa.gov/research/science\\_snapshots/2015/MERRA2\\_global\\_aerosol\\_dist.php](https://gmao.gsfc.nasa.gov/research/science_snapshots/2015/MERRA2_global_aerosol_dist.php)). In this study,  
209 MERRA-2 dust (DAOD) and total AOD were used for providing the ratio of dust aerosols over total particulate,  
210 all near-surface aerosol mass concentrations for estimating PM10, horizontal and vertical dust flux for estimating  
211 dust loading import/export and dust emission/deposition for Australia, respectively.

212  
213 In MERRA-2, the horizontal dust flux ( $\text{mg}/\text{m}\cdot\text{s}$ ) is divided into northward and eastern components, which are  
214 represented by u and v-wind dust flux, respectively. Similar to determination of wind speed and wind direction,



215 the horizontal dust flux ( $F_{d,h}$ , see equation 1) and its direction ( $\phi_{d,h}$ , see equation 2) can be calculated using  
 216 MERRA-2 u/v-wind dust flux data. The angle between the direction of a dust flux pixel crossing the land borders  
 217 and land borders was used to determine whether the dust flux was imported or exported.

$$F_{d,h} = \sqrt{F_{d,u}^2 + F_{d,v}^2} \quad (1)$$

$$\phi_{d,h} = \text{mod} \left( 180 + \frac{180}{\pi} \text{atan2}(F_{d,u}, F_{d,v}), 360 \right) \quad (2)$$

218 where  $F_{d,u}$  and  $F_{d,v}$  represent dust flux in the u and v directions, respectively.

219 The total dust emission ( $E$ ) and deposition (both with a unit of Mt) in Australia were calculated using MERRA-2  
 220 vertical dust flux, including emission and dry/wet deposition flux. Equation 3 and equation 4 were used for  
 221 calculating  $E$ , as well as estimating dust deposition.

$$E = \sum_{m=1}^{12} E_m \quad (3)$$

$$E_m = \frac{1}{n} A \sum_{n=1}^N F_{m,n} T_m \quad (4)$$

222 where  $E_m$  and  $T_m$  represent the total dust emission in mass (Mt) and total seconds for the month  $m$ ,  
 223 respectively;  $F_{m,n}$  is the total flux of MERRA-2 dust emission flux ( $\text{kg}/\text{m}^2 \cdot \text{s}$ ) for all five bins (Table 1) for the  
 224 MERRA-2 pixel  $n$  in Australia during the month  $m$ ;  $A$  equals to 7.62 trillion  $\text{m}^2$ , standing for the total land  
 225 area of Australia excluding Tasmania.

226 Table 1. Dry size range ( $\mu\text{m}$ ) for dust aerosol simulation with 5 bins in MERRA-2

Bin	1	2	3	4	5
Radius ( $\mu\text{m}$ )	0.73	1.4	2.4	4.5	8.0
radius lower ( $\mu\text{m}$ )	0.1	1.0	1.8	3.0	6.0
radius upper ( $\mu\text{m}$ )	1.0	1.8	3.0	6.0	10.0

227 (<https://gmao.gsfc.nasa.gov/reanalysis/MERRA-2/FAQ/Dust.pdf>)

228

229 Limited by a lack of ground-based near-surface observations, most validation studies of MERRA-2 component  
230 concentrations focus on a single site or a few sites, especially the dust component. For example, MERRA-2 OC,  
231 BC, sulfate concentration data sets have been validated against ground-based observations in China in Nanjing  
232 (Zhao et al., 2021) and Jingsha and Lin'an (Ma et al., 2021), Beijing (only BC) (Qin et al., 2019; Ou et al., 2022)  
233 and over northern India (OC and BC) (Soni et al., 2021). As for the dust component, daily MERRA-2 dust  
234 concentrations have been validated in Barbados (daily product) (Buchard et al., 2017), and Cayenne, Northern  
235 South America (Prospero et al., 2020).

236

237 Although there are relatively few ground-based observations, validation results still show daily and monthly  
238 MERRA-2 surface mass concentrations in the forms of PM<sub>2.5/10</sub> or single components with relatively high  
239 accuracy. MERRA-2 near-surface component concentrations can be used for conversion to PM<sub>10</sub> using the  
240 equation below (Provençal et al., 2017; Ma et al., 2021):

$$[\text{PM}_{10}] = 1.375 \times [\text{SO}_4^{2-}] + 1.8 \times [\text{OC}] + [\text{BC}] + [\text{DU}_{10}] + [\text{SS}_{10}] \quad (5)$$

241 where  $[\text{SO}_4^{2-}]$ ,  $[\text{OC}]$ ,  $[\text{BC}]$ ,  $[\text{DU}_{10}]$ , and  $[\text{SS}_{10}]$  are concentrations of each aerosol component, namely sulfate,  
242 organic carbon, black carbon, dust, and sea salt aerosol, and the subscripts 10 indicates the particle diameter less  
243 than 10 $\mu\text{m}$ .  $[\text{SO}_4^{2-}]$  is multiplied by 1.375 under the assumption that  $\text{SO}_4$  is fully neutralized by ammonium in the  
244 form of  $(\text{NH}_4)_2\text{SO}_4$  (ammonium sulfate) and a scale factor of 1.8 for OC is included to take into consideration the  
245 organic compounds in the particulate organic matter. Equation 5 was developed for estimating MERRA PM<sub>10</sub>  
246 over Europe (Provençal et al., 2017), which may be inappropriate for PM<sub>10</sub> estimation over other regions. For  
247 example, Ma et al. (2021) considered the increasing trend of nitrate which was a large proportion of aerosols in  
248 China, and revised Equation 5 as:

$$[\text{PM}_{10}] = 1.375 \times [\text{SO}_4^{2-}] + 1.29 \times [\text{NO}_3^-] + 1.8 \times [\text{OC}] + [\text{BC}] + [\text{DU}_{10}] + [\text{SS}_{10}] \quad (6)$$

249 where  $[\text{NO}_3^-]$  is the concentration of nitrate. Considering that coarse mode aerosols take up 57%-71% of total  
250 aerosols over major cities (Chan et al., 2008) and nitrate emissions contribute much less than other aerosol species  
251 to the atmosphere (Bauer et al., 2007), especially during smoke events in the northern savanna (Desservettaz et  
252 al., 2017), in this study, the method developed by the Global Modeling and Assimilation Office (GMAO) (equation  
253 7) using MERRA-2 3-D aerosol mass mixing ratios was used for PM<sub>10</sub> estimation over Australia. Nitrate is  
254 missing due to its minor contribution to PM<sub>10</sub> concentrations.

$$[\text{PM}_{10}] = (1.375 \times [\text{SO}_4^{2-}] + [\text{BC}_{\text{phobic}}] + [\text{BC}_{\text{philic}}] + [\text{OC}_{\text{phobic}}] + [\text{OC}_{\text{philic}}] + [\text{DU}_{001}] + [\text{DU}_{002}] + [\text{DU}_{003}] + 0.74 \times [\text{DU}_{004}] + [\text{SS}_{001}] + [\text{SS}_{002}] + [\text{SS}_{003}] + [\text{SS}_{004}]) * \text{AIRDENS} \quad (7)$$

255 where the subscripts philic and phobic for [BC] and [OC] represent hydrophilic and hydrophobic BC and OC  
 256 aerosols, respectively, the numbers after [DU] and [SS] indicate the 4 size bins, i.e. 001 to 004 represent the bins  
 257 with radius from 001 for 0.1 to 1.0 $\mu\text{m}$ , 002 for 1.0 to 1.8 $\mu\text{m}$ , 003 for 1.8 to 3.0 $\mu\text{m}$ , and 004 for 3.0 to 6.0 $\mu\text{m}$  for  
 258 dust, and similarly for 0.03 to 0.1 $\mu\text{m}$ , 0.1 to 0.5 $\mu\text{m}$ , 0.5 to 1.5 $\mu\text{m}$ , and 1.5 to 5.0 $\mu\text{m}$  for sea salt, and the AIRDENS  
 259 means air density (<https://gmao.gsfc.nasa.gov/reanalysis/MERRA-2/FAQ/#Q5>)

### 260 **2.3 MODIS DeepBlue AOD dataset**

261 MODIS DB aerosol product provides nearly full global coverage of AOD, AE, and SSA datasets over EOS (Earth  
 262 Observing System) years (from 2000 to the present) (Sayer et al., 2017). MODIS DB has been widely used for  
 263 dust research over arid and semi-arid regions i.e., bright surfaces where the traditional dark target (DT) algorithm  
 264 is not applicable. For example, Ginoux et al. (2012) analyzed the global distribution of dust sources and emissions  
 265 using MODIS DB aerosol product, taking advantage of its coverage over bright surfaces and successful retrieval  
 266 of AE and SSA together with AOD. Similarly with other aerosol products released by NASA like DT (Levy et al.,  
 267 2013) and aerosol climate change initiative (aerosol\_CCI) such as AATSR (Advanced Along-Track Scanning  
 268 Radiometer) aerosol products (de Leeuw et al., 2015; Sundström et al., 2012; Kolmonen et al., 2016; Thomas et  
 269 al., 2009), all L2 MODIS DB aerosol datasets in the latest C61 product were produced with a spatial resolution of  
 270 10km with all MODIS radiance data. In this study, MODIS DB aerosol product for Aqua from 2002-2020 was  
 271 selected.

272  
 273 MODIS DB key parameters have been validated over Australia and globally, especially AOD. The MODIS DB  
 274 AOD dataset has been validated against AERONET data (Che et al., 2022; Sayer et al., 2019; Wei et al., 2019),  
 275 inter-compared with other AOD products for MODIS, such as MAIAC and DT (Shaylor et al., 2022), and even  
 276 evaluated with MERRA-2 AOD (Che et al., 2022) over Australia. These studies show that there is a high  
 277 probability of data points (MODIS DB and AERONET) within the expected envelope (EE) lines, which are  
 278 defined with two lines ( $\pm(0.03 + 0.15\tau)$ ) containing approximately 68% of data points (Che et al., 2022). The  
 279 latest MODIS DB aerosol product limits AE values from 0 to 1.8, and in low AOD conditions AE is set to < 1.0  
 280 over bright surfaces and AE is fixed to 1.5 over vegetated surfaces (Sayer et al., 2013; Hsu et al., 2013). Sayer et  
 281 al. (2019) tested the performance of MODIS AE in different conditions, including dust cases and fine mode cases.

282 Over vegetated surfaces, MODIS DB AE was overestimated systematically with a broad range of error from 0.5  
283 to 1 for dusty conditions, while over dry surfaces the performances of MODIS DB AE have improved in systematic  
284 overestimation but still a broad error range (Sayer et al., 2019). Therefore, dust detection by MODIS DB could be  
285 uncertain to some extent.

#### 286 **2.4 MODIS-MERRA (M&M) combined DAOD**

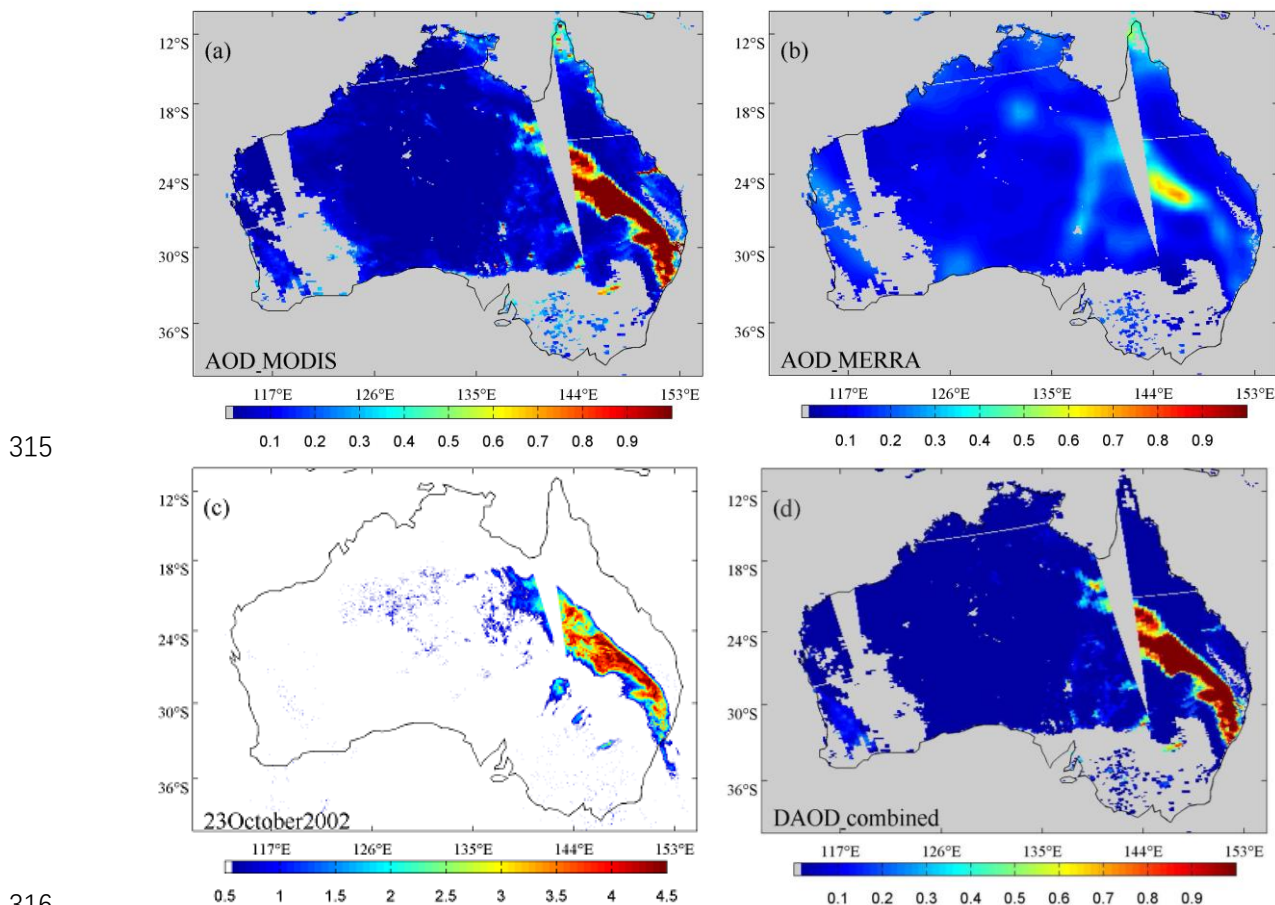
287 DAOD has been successfully retrieved based on MODIS DB product by a coarse mode fraction (CMF) in equation  
288 8 (Ginoux et al., 2010, 2012; Pu and Ginoux, 2017):

$$289 \text{CMF} = 0.98 - 0.5098\text{AE} - 0.05\text{AE}^2 \quad (8)$$

290 The early version of the MODIS DB DAOD dataset has been used for analyzing the global distribution of dust  
291 sources (Ginoux et al., 2012) and constructing a global DAOD climatology (Voss and Evan, 2020). The new  
292 version of MODIS DB DAOD dataset has been used for evaluating DAOD satellite remote sensing from GCMs  
293 and CALIOP, such as CMIP5 (Pu and Ginoux, 2018), MERRA-2 (Che et al., 2022), and CALIOP (Song et al.,  
294 2021). In Australia, the MODIS DB DAOD dataset was validated against AERONET data and results show that  
295 88% and 71% of data points for MODIS/Terra and MODIS/Aqua, respectively, fall within an EE of ( $\pm 0.05 + 0.15\tau$ )  
296 (Che et al., 2022). Although studies have shown MODIS DB DAOD dataset is of high quality, there are still  
297 several factors limiting its applications to dust research in Australia. First, the data coverage of the MODIS DB  
298 DAOD dataset only includes obvious dust plumes. Secondly, MODIS DB AE retrievals have a broad range of  
299 errors over both bright and vegetated surfaces, especially systematic overestimations over vegetated surfaces  
300 (Sayer et al., 2019), causing non-negligible uncertainty in the MODIS DB DAOD dataset. Thirdly, AE was fixed  
301 in low AOD conditions (Sayer et al., 2013; Hsu et al., 2013). MERRA-2 aerosol reanalysis is expected to make  
302 up for these deficiencies.

303 A new DAOD dataset has been developed for Australia in this study using MODIS DB aerosol product and  
304 MERRA-2 aerosol reanalysis. Figure 2a shows MODIS DB AOD on 23<sup>rd</sup> October 2002 with hundreds of  
305 kilometers of dust plume in eastern Australia while the dust plume is seriously underestimated by MERRA-2  
306 (Figure 2b) compared with MODIS BTM in Fig.2c. In order to take advantage of MODIS DB in catching dust  
307 plumes and MERRA-2 in spatial coverage, DAOD equals to MODIS DB DAOD when it is available, otherwise,  
308 DAOD equals to MODIS DB AOD multiplied by the ratio of MERRA-2 DAOD to total AOD. This is based on  
309 the assumption that the dust fraction in MERRA-2 is shown to have high accuracy with LIVAS (Gkikas et al.,

310 2021) and AERONET (Che et al., 2022). Figure 2d shows that the final DAOD is capable of screening sea salt  
 311 AOD over the Cape York Peninsula and has the same spatial coverage as MODIS DB AOD. Although Sayer et al.  
 312 (2019) suggest that AE should be only used for discriminating coarse mode-dominated AOD from fine mode-  
 313 dominated AOD qualitatively, a smoke plume (Fig.2d) over the Australian east coast was effectively removed  
 314 from MODIS DB AOD.



317 **Figure 2: Development of DAOD using MERRA-2 and MODIS DB dataset. (a) MODIS DB AOD, (b)**  
 318 **MERRA-2 AOD, (c) MODIS BTD, and (d) MERRA-MODIS combined DAOD (M&M).**

319 **2.5 Gridded SILO monthly rainfall data**

320 The SILO datasets developed by the Queensland government are aimed at providing long-term continuous point  
 321 and full coverage gridded climate datasets for land areas of Australia from 1989 to the present  
 322 (<https://www.longpaddock.qld.gov.au/silo/>). The full coverage gridded rainfall dataset at a temporal resolution of  
 323 a day and a month was produced by interpolating BoM daily and monthly rainfall observations with a Kriging  
 324 method (Jeffrey et al., 2001). Validation results show that the accuracy of SILO data is typically higher around  
 325 areas with densely distributed BoM rainfall gauge sites (Jeffrey et al., 2001). Overall, except for parts of western

326 Australia with few BoM rainfall sites SILO rainfall shows a high accuracy with an  $R^2 > 0.8$  over most of the  
327 Australian land (Jeffrey et al., 2001).

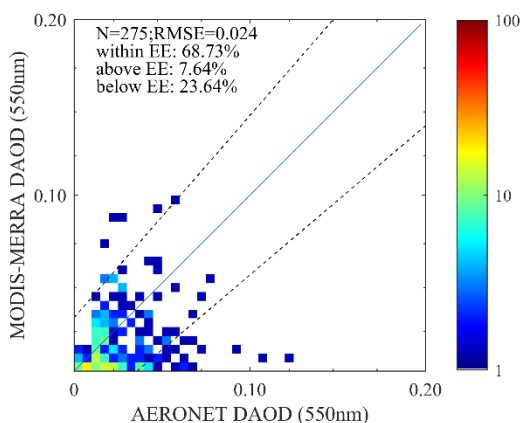
328

### 329 3. Results

#### 330 3.1 Validation of MERRA-2 surface mass concentration and DAOD

331 Figure 3 shows the validation result of MODIS (Aqua)-MERRA (M&M) combined DAOD over Australia from  
332 2002-2020. The average AERONET DAOD for all collocated data points is only 0.03 for Australia. When AOD  
333 is low, remote sensing AOD retrievals are likely to be close to the margin of uncertainty and hence subject to large  
334 relative bias, especially over the arctic region (Mei et al., 2013b, a), Qinghai-Tibet Plateau (Che et al., 2018, 2016),  
335 and Australia (Che et al., 2022; Sayer et al., 2019). The ratio of RMSE (root mean square error) to the mean  
336 AERONET DAOD for MODIS-MERRA DAOD was 0.8, indicating that the uncertainty is close to MODIS-  
337 MERRA DAOD. The EE that contains 68% of data points is  $\pm(0.016 + 0.15\tau_{Aero})$  for MODIS-MERRA  
338 DAOD to AERONET DAOD over Australia. The intercept in EE is 0.016 is much smaller than for the MODIS  
339 DB AOD over Australia (0.03) (Che et al., 2022), suggesting a high level of accuracy of this MODIS-MERRA  
340 DAOD dataset with a smaller absolute error.

341



342

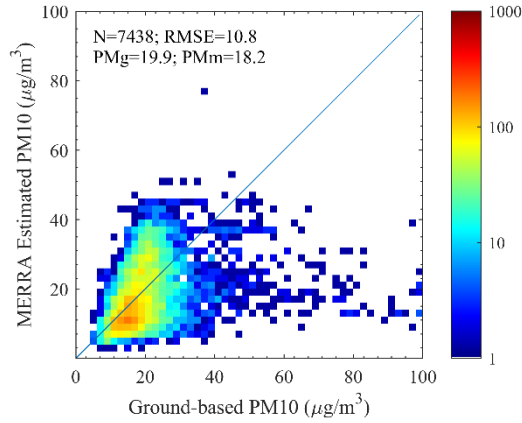
343 **Figure 3: Comparison of MERRA-MODIS DAOD with AERONET DAOD. The dashed lines denote an EE**  
344 **of  $\pm(0.016 + 0.15\tau_{Aero})$  which contains 68% of the data points.  $\tau_{Aero}$  represents AERONET AOD.**

345

346 Figure 4 shows the validation results of MERRA-2 estimated monthly PM10 with ground-based observations at  
347 62 AQMN stations for the period 2001-2020. There are 7438 data points in total for MERRA-2 PM10 validation

348 over eastern NSW. These selected PM10 observations are located downwind areas of inland dust sources such as  
349 the Lake Eyre, South Simpson lakes, and the Channel Country (O’Loingsigh et al., 2017). PM10 at these sites,  
350 therefore, could represent dust activities in southeast Australia during dust seasons. The mean monthly PM10 for  
351 all data points is  $19.9\mu\text{g}/\text{m}^3$ , indicating a clear atmosphere over eastern NSW on average. When PM10  
352 observations are greater than  $40\mu\text{g}/\text{m}^3$ , almost all the data points are below the 1-1 line, suggesting that MERRA-  
353 2 is incapable of catching high PM10 events (dust or other pollutions). Due to the relatively dense spatial  
354 distribution, AQMN sites are likely to observe the same dust events with similar PM10 observations and thus a  
355 similar extent of underestimation. Time series plots (Fig.5) show similar severe underestimation in Newcastle  
356 ( $13.5\mu\text{g}/\text{m}^3$  vs.  $106.9\mu\text{g}/\text{m}^3$ ), Randwick ( $12.4\mu\text{g}/\text{m}^3$  vs.  $84.7\mu\text{g}/\text{m}^3$ ), and Wollongong ( $13.0\mu\text{g}/\text{m}^3$  vs.  $65.2\mu\text{g}/\text{m}^3$ )  
357 in September 2009 and Bathurst ( $18.9\mu\text{g}/\text{m}^3$  vs.  $104.8\mu\text{g}/\text{m}^3$ ), Bulga ( $30.4\mu\text{g}/\text{m}^3$  vs.  $78.9\mu\text{g}/\text{m}^3$ ), and  
358 Newcastle ( $37.1\mu\text{g}/\text{m}^3$  vs.  $50.8\mu\text{g}/\text{m}^3$ ) in December 2019. In September 2009, a severe dust storm swept the  
359 Australian continent, causing a jump in PM10 for overpass areas (Leys et al., 2011). High PM10 (higher than  
360  $300\mu\text{g}/\text{m}^3$ ) lasted for approximately 12 hours and reached as high as  $15388\mu\text{g}/\text{m}^3$  at Bathurst. Similarly, high PM10  
361 concentrations were recorded at Bulga and Newcastle (Fig.5). Due to rainfall deficiency and high temperatures in  
362 November and December 2019, NSW experienced the longest bushfire season when more than five million  
363 hectares were burned (BBC News, 2020). The NSW AQMN stations, therefore, had recorded high PM10  
364 concentrations during the bushfire season. These underestimations by approximately 5 were also a major reason  
365 why the ground-based mean PM10 ( $19.9\mu\text{g}/\text{m}^3$ ) was higher than that of MERRA-2 ( $18.2\mu\text{g}/\text{m}^3$ ). When PM10 is  
366 less than  $20\mu\text{g}/\text{m}^3$ , data points show a slight bias of MERRA-2, and when PM10 is between  $20\mu\text{g}/\text{m}^3$  to  $40\mu\text{g}/\text{m}^3$ ,  
367 the bias in MERRA-2 reduces. This is shown by evenly distributed data points around 1-1 line but less association  
368 occurs between the two. Severe underestimations of MERRA-2 PM10 (Fig.5) show a strong seasonality at some  
369 sites like Albury and Bathurst in summer when PM10 is greater than  $40\mu\text{g}/\text{m}^3$ . This suggests that MERRA-2 is  
370 very likely to underestimate dust severity in summer because dust events mainly occur in summer throughout the  
371 year in NSW (Che et al., 2022). In spite of the underestimations, MERRA-2 is capable of tracing the seasonal  
372 variations of PM10 at six sites with an  $r^2$  value from 0.14 to 0.44. Overall, a RMSE is  $10.8\mu\text{g}/\text{m}^3$  for all monthly  
373 MERRA-2 and ground-based PM10 data.

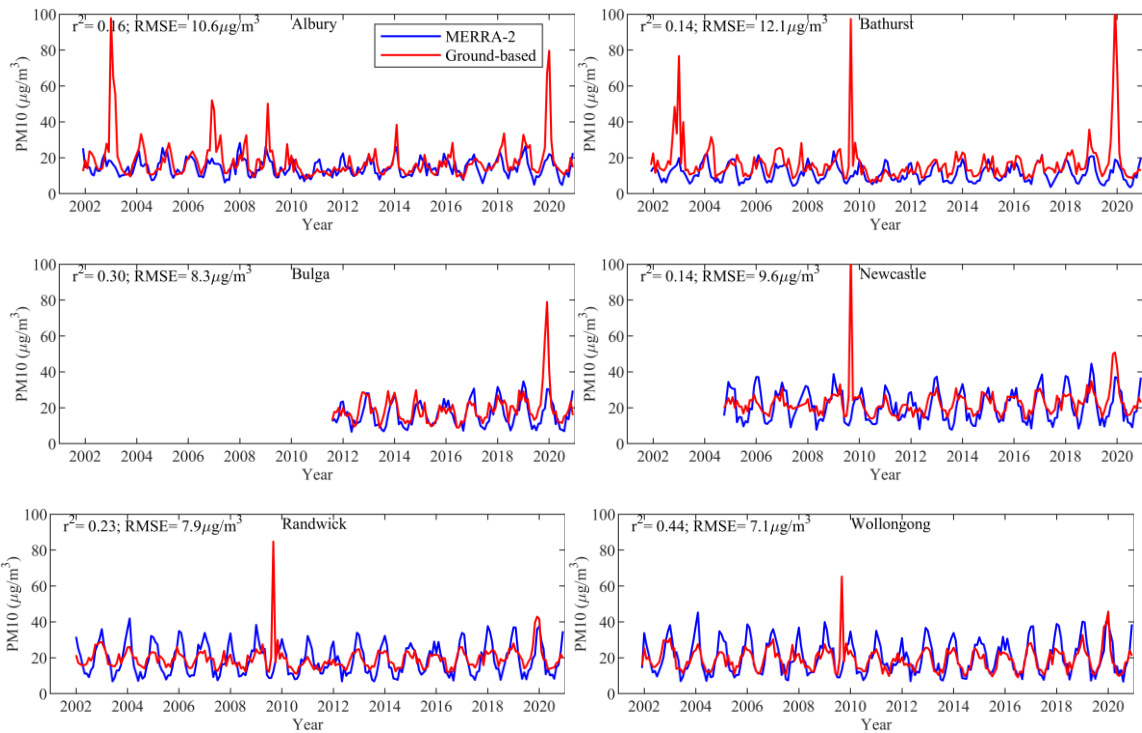




374

375 **Figure 4: Comparison of monthly MERRA-2 estimated PM10 with ground-based PM10 observations at 62**  
 376 **AQMN stations.**

377



378

379

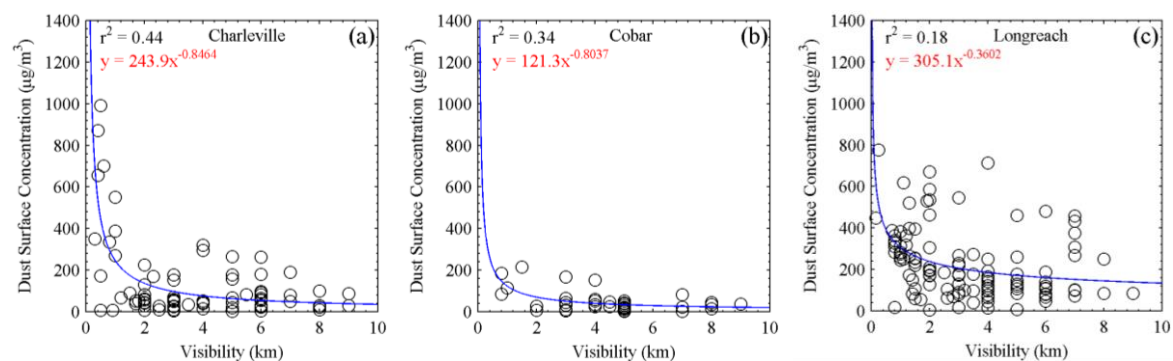
380

381 **Figure 5: Time series for MERRA-2 estimated PM 10 with ground-based PM observations at 6 sites in NSW.**

382

383 Figure 6 shows the relationships between MERRA-2 near-surface dust concentrations with horizontal visibility  
 384 with a dust type based on SYNOP code at Charleville, Cobar, and Longreach. The relationships are similar at  
 385 Charleville and Cobar in that MERRA-2 near-surface dust concentrations follow a power function relationship  
 386 with horizontal visibility. The  $r^2$  values for two sites of 0.44 and 0.34, respectively, also suggest a relatively robust  
 387 relationship between the two datasets. At Longreach, a low  $r^2$  value of 0.18 shows a weak relationship between

388 MERRA-2 near-surface dust concentrations and visibilities. Longreach is known to be in a region of frequent  
 389 local wind erosion activities with extensive tracts of clay soils in Eastern Australia (McTainsh et al., 1990).  
 390 Alluvial sediments and sandy clays, therefore, would be important sources of local dust events in Longreach (Rust  
 391 and Nanson, 1989; McTainsh et al., 1990). These clay aggregates exhibit lower optical extinction compared to  
 392 fine clay. However, due to their larger mass, they can still contribute to a high concentration of dust near the  
 393 surface in high visibility conditions. This likely explains why the  $r^2$  in Figure 6c differs from the other two.  
 394 Previous studies showed that near-surface dust concentrations/total suspended particle concentrations agree  
 395 statistically well with the visibility data (Baddock et al., 2014; Chepil and Woodruff, 1957; Shao et al., 2003; Tews,  
 396 1996) and visibility-defined DSI (O’Loingsigh et al., 2014), and horizontal visibility have often been used for  
 397 calculating dust concentrations (Leys et al., 2011; McTainsh et al., 2005). Similar power function relationships  
 398 between the two suggest the acceptable accuracy of MERRA-2 near-surface dust concentrations to an extent.  
 399



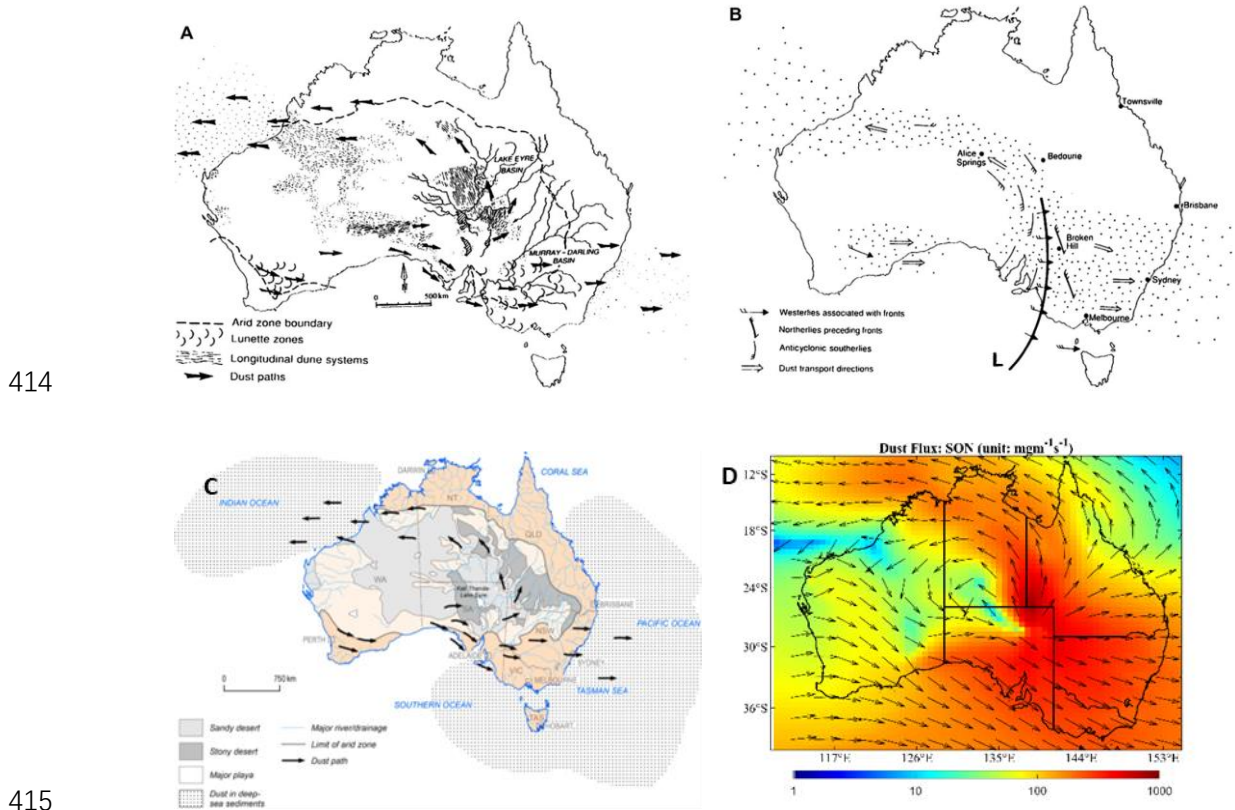
400

401 **Figure 6: Relationships between BoM horizontal visibility observations of MERRA-2 dust surface mass**  
 402 **concentrations.**

403

404 Figure 7 shows dust pathways and sink areas for Australia as identified in previous studies. Bowler (1976)  
 405 established the first dust pathways (Figure 7a) using trends of sand dune movement during the period of intense  
 406 dune building phases. Major dustfall areas lie mainly to the southeast and northwest of the Australian continent.  
 407 Sprigg (1982) proposed a conceptual model (Figure 7b) for describing how wind systems fed dust into the  
 408 pathways identified by Bowler using measured wind run, wind direction, and wind speed in desert areas. Blewett  
 409 (2012) adopted the dust pathways established by Bowler but provided a detailed classification of sand dunes and  
 410 accurately confirmed dustfall in the southeast offshore area. Figure 7d shows MERRA-2 horizontal dust flux over  
 411 Australia from 2002 to 2020. Compared to previous studies (i.e. Figure 7a-c), MERRA-2 horizontal flux

412 quantitatively shows the mean dust pathways and dustfall areas for Australia, providing independent support to  
 413 previous conceptual models.



416 **Figure 7: Comparison of dust pathways over Australia delineated in previous studies. (a) Bowler (1976),**  
 417 **(b) Sprigg (1982), (c) Blewett (2012), and (d) mean horizontal dust flux based on MERRA-2 from 2002-**  
 418 **2020.**

419 **3.2 Seasonal DAOD based on MODIS-MERRA, dust concentration and PM10 based on MERRA-2**

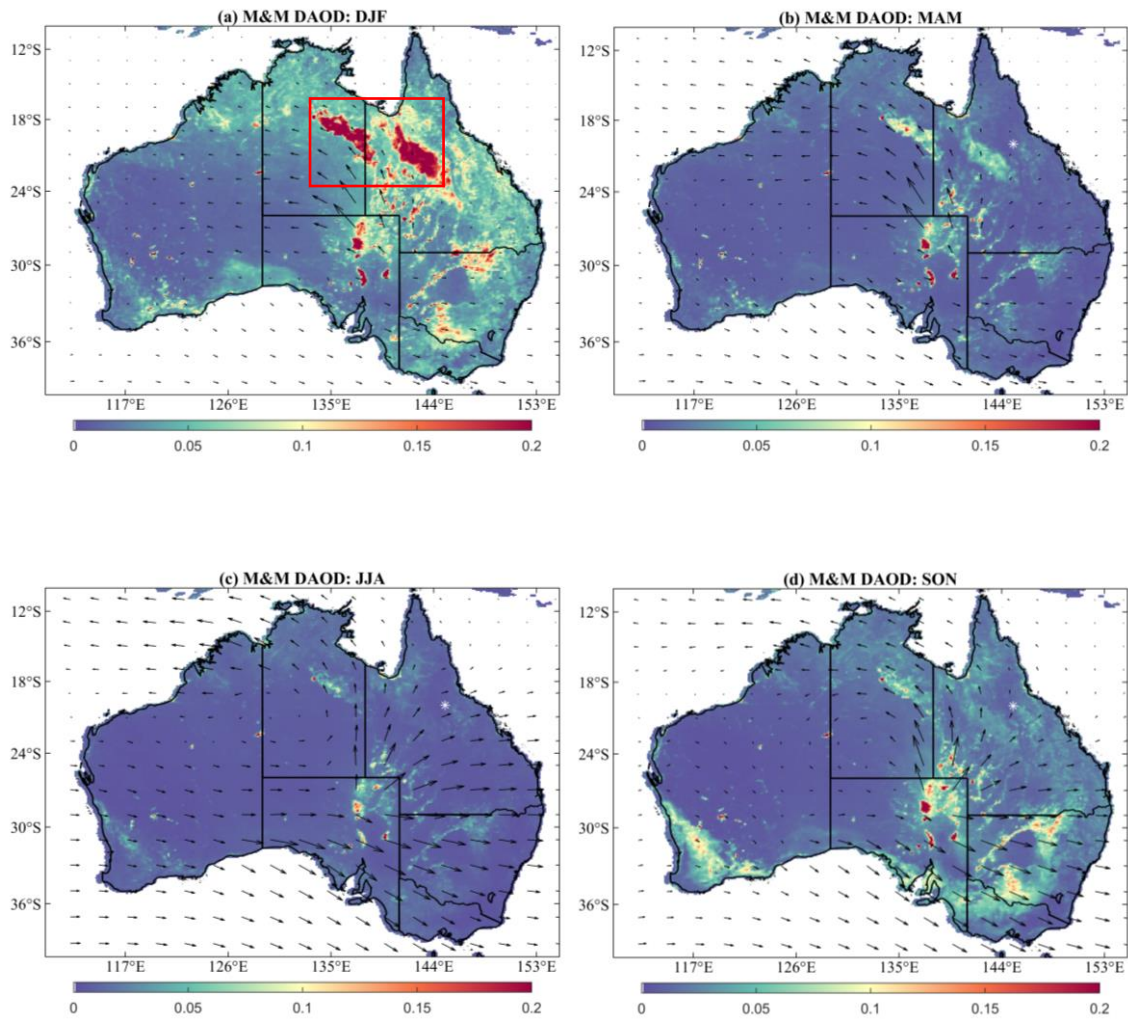
420 Figure 8 shows the mean seasonal MODIS-MERRA DAOD over Australia from 2002 to 2020. Spring (Sep-Nov)  
 421 and summer (Dec-Feb) are typically dust seasons in Australia, when DAOD is much higher than that in autumn  
 422 (Mar-May) and winter (Jun-Aug). This seasonal pattern is consistent with that shown by the monthly dust event  
 423 frequency based on synoptic observations (McTainsh et al., 1998). As dust activities would normally last until  
 424 March in southeastern Australia (McTainsh et al., 1998), DAOD in this area is relatively high compared to that in  
 425 Western Australia (Fig. 8b). High rainfall may inhibit occurrences of dust events, therefore, Figure 8c shows low  
 426 DAOD for most of Australia in winter.

427  
 428 High DAOD is shown over dust source areas and along main pathways. Dust and sand particles originated from

429 the inland area (around the Lake Eyre Basin) and were transported along two main dust transport pathways to the  
430 north and southeast (Bowler, 1976; De Deckker, 2019; McGowan et al., 2000; Sprigg, 1982a; Strong et al., 2011).  
431 In spring (Figure 8d), high DAOD regions are mainly concentrated around the Lake Eyre Basin in southeastern  
432 Australia, and in the southwest of Western Australia. Fig. 8d shows that the horizontal flux with dust entrained  
433 from the source region in Central Australia to the southeast is so far the strongest over the Australian continent.  
434 DAOD and horizontal dust flux are consistent with each other and both reflect the major dust pathway in  
435 southeastern Australia. High DAOD can be also found in the southwest of Western Australia, which has been  
436 identified as the starting point of the major pathways flow in the south in previous studies (Sprigg, 1982a). Due  
437 to onshore winds, high DAOD around this region is very likely to be generated from local dust sources. Also,  
438 DAOD in spring is much higher than in other seasons in this region. In summer (Figure 8a), the highest DAOD  
439 regions were found in the main pathway to the north as highlighted with a red box, and around the center of the  
440 Lake Eyre Basin. The high DAOD region in the north of NT and QLD is in the main pathway, with a spatial  
441 pattern consistent with prevailing wind directions. Another region with high DAOD in the middle of QLD shows  
442 a different spatial pattern with MERRA-2 horizontal dust flux, which may be caused by differences in data  
443 coverage between MODIS DB and MERRA-2 datasets. The second highest DAOD regions are concentrated  
444 around the southeast dust pathway in NSW. Meanwhile, the horizontal dust flux for these regions is much lower  
445 than those for northern regions. In autumn (Figure 8b), DAOD is an extension of that in summer that DAOD  
446 shows a very similar spatial pattern to that in summer but lower in value. In regions with high DAOD, the DAOD  
447 distribution is consistent with that of horizontal dust flux (i.e., high DAOD corresponds to large horizontal dust  
448 flux). In winter, dust emissions in the center of Lake Eyre Basin are the smallest and DAOD is the lowest among  
449 the four seasons. Dust is mainly transported to the east coast of Australia, deposited in eastern Australia and the  
450 ocean areas to the southeast and northwest.

451

452



453

454

455

456

457

458

459

460

461

462

463

464

465

466

467

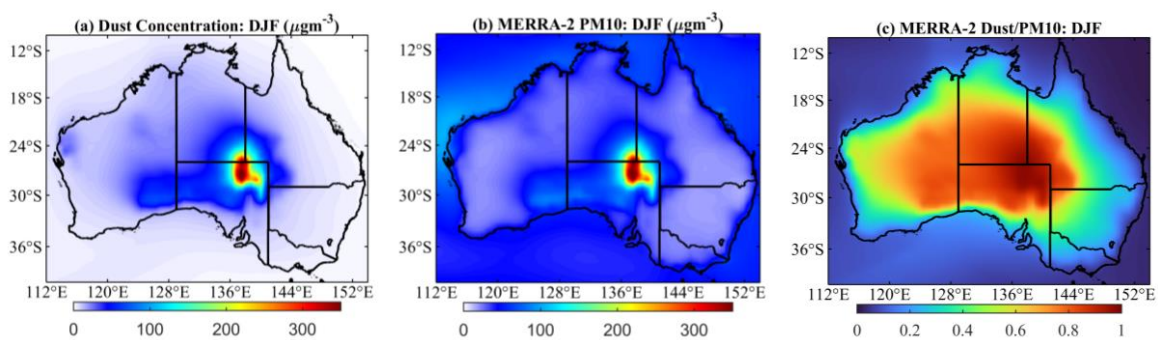
**Figure 8: Seasonal MERRA-MODIS DAOD from 2002-2020. (a) Sep-Nov, (b) Dec-Feb, (c) Mar-May, (d) Jun-Jul. Arrows represent horizontal dust flux (unit: kg/m·s) with the direction and magnitude. The red box outlines frequent dust events regions differing from those in previous studies based on the dust storm index.**

Figure 9 shows mean seasonal near-surface dust concentrations and PM10 based on MERRA-2, and the ratio of the two for Australia. The highest near-surface dust concentrations are mainly distributed over the Lake Eyre Basin which is the largest natural dust source in Australia, while high concentrations from 50 to 100 $\mu\text{g}/\text{m}^3$  are found around the Lake Eyre Basin, Great Victoria Desert, and Nullarbor Plain in four seasons (See Figure.9a, d, g, and j). In other regions, near-surface dust concentrations typically are less than 50 $\mu\text{g}/\text{m}^3$  and the concentration decreases towards the coastline. Low concentrations of less than 20 $\mu\text{g}/\text{m}^3$  of dust can be found over the ocean, particularly the Indian Ocean in the northwest. MERRA-2 PM10 (See Figure 9b, e, h, and k) shows that the spatial

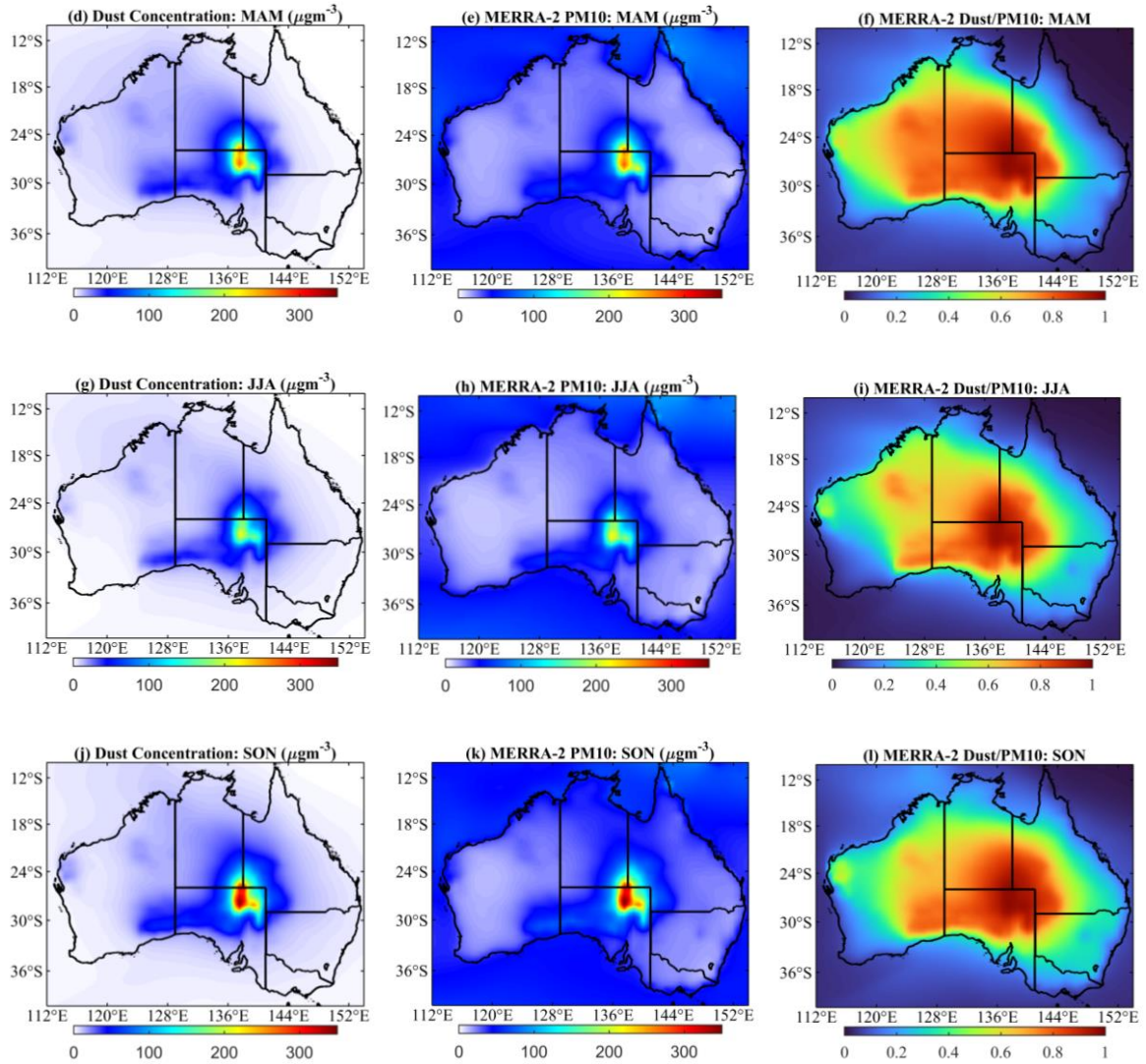


468 distribution is similar to that of the near-surface dust concentrations over the continent. Differences between the  
 469 two mainly occur in the offshore areas due to the influence of sea salt aerosols and carbonaceous aerosols in the  
 470 north in spring (Yang et al., 2021). Spatial distributions of dust concentration and PM10 are similar because PM10  
 471 accounts for the majority of dust particles in inland Australia (Figure 9c, f, I, and I). A high ratio of dust  
 472 concentration over PM10 of over 0.7 is mainly found in inland areas and the ratio decreases towards the coastline  
 473 which is consistent with coarse particles settling out of suspension more quickly compared to finer particles  
 474 (Fryrear et al., 1991). It should be noted that four states in eastern Australia are affected relatively less by dust  
 475 than other regions from the perspective of the ratio of dust to PM10. The least affected states are VIC (higher than  
 476 0.2), NSW (0.3), and QLD (higher than 0.3), respectively.

477  
 478 Dust sources and pathways play an essential role in determining the spatial pattern of near-surface dust  
 479 concentrations. Dust events occur frequently in Central Australia (McTainsh et al., 2011a), resulting in extremely  
 480 high near-surface dust concentrations throughout the year. In addition, the seasonal variation of the near-surface  
 481 dust concentration in such regions is more significant than other dust sources, with a difference exceeding  
 482  $100\mu\text{g}/\text{m}^3$  between spring/summer and autumn/winter. The near-surface dust concentrations over two major dust  
 483 pathways changed a little with season compared to those over the main dust source areas in the center of the  
 484 continent. Relatively high concentrations of near-surface dust are found to the north in the four seasons compared  
 485 to that along the dust pathway to the southeast. The wind systems responsible for dust pathways as described by  
 486 Spriggs (1982) explain that pre-frontal anti-cyclonic northerly winds are responsible for the main dust pathway in  
 487 the south. High dust concentrations are, however, not found in the main dust deposition area in the southeast.



488



489

490

491

492 **Figure 9: Seasonal MERRA-2 near-surface dust concentration in (a) Dec-Feb, (d) Mar-May, (g) Jun-Jul,**  
 493 **and (j) Sep-Nov from 2002-2020; Seasonal MERRA-2 PM10 in (b) Dec-Feb, (e) Mar-May, (h) Jun-Jul, and**  
 494 **(k) Sep-Nov from 2002-2020; and the ratio of MERRA-2 near-surface concentration to MERRA-2 PM10**  
 495 **in (c) Dec-Feb, (f) Mar-May, (i) Jun-Jul, and (l) Sep-Nov from 2002-2020.**

496

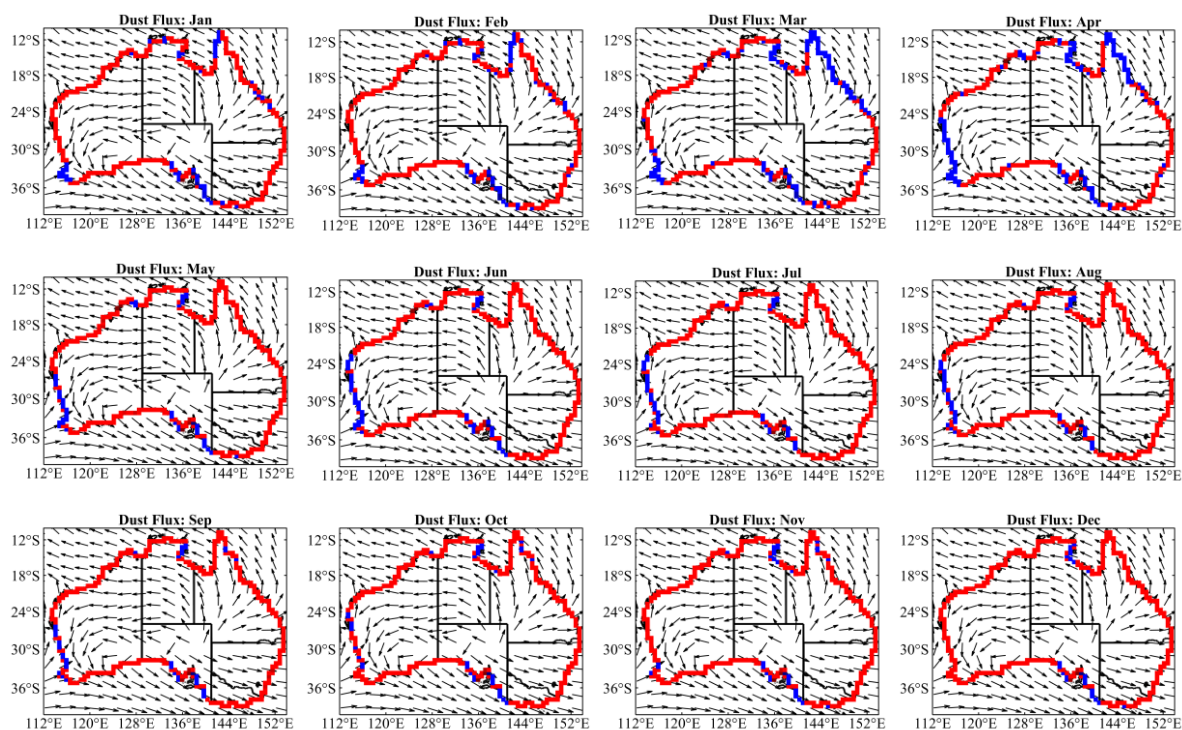
497 **3.3 Dust loading budget in the Australian continent**

498 Figure 10 shows the directions of horizontal dust flux using MERRA-2 u/-wind dust flux datasets. The colors of  
 499 the coastline indicate whether dust is transported from the land to the ocean (red) or from the ocean to the land  
 500 (blue). As the “loneliest” inhabited continent, Australia is located far away from other continents, with the largest  
 501 natural dust source in the southern hemisphere, the Lake Eyre Basin, and thus regarded as a main dust source  
 502 exporting dust. However, the blue border shows MERRA that areas of import can occur along the west, north, and south



503 coasts. Along the north coast, exported dust from QLD could be transported back to the Cape York Peninsula, part  
 504 of this dust would travel on to Arnhem Land and even travel back to the continent. This recirculated dust cannot  
 505 be defined as “true imported dust” because it originated from the Australian continent, and was transported over  
 506 the sea and back onto the continent. A similar situation can be found on the coastline in South Australia where  
 507 dust originated from the continent and was transported outwards from the Nullarbor Plain, across the Spencer  
 508 Gulf, and back to South Australia and VIC. Different from these two situations, dust imported from the west coast  
 509 is very likely to be from remote dust sources in South Africa. Firstly, MERRA-2 horizontal dust flux doesn’t show  
 510 bags of dust exported from the northwest coast are transported back to the continent. Secondly, dust originating  
 511 from the Mallee region is unlikely to be transported to the west coast crossing the Pacific Ocean (Bhattachan and  
 512 D’Odorico, 2014). Thirdly, forward trajectory analysis demonstrates dust originating in the Kalahari dust can be  
 513 transported over long distances to Australia (Bhattachan et al., 2012). The southwesterly winds are key to  
 514 transporting dust from South Africa to Western Australia due to their similar latitudes (Torre et al., 2022).  
 515 Therefore, the dust imported from 20°S to 36°S is therefore regarded as the only imported dust from an external  
 516 source for Australia in this study.

517



521 **Figure 10: Directions of dust flow in each month from 1980 to 2020. Blue and red borders indicate imported**  
 522 **and exported dust, respectively.**

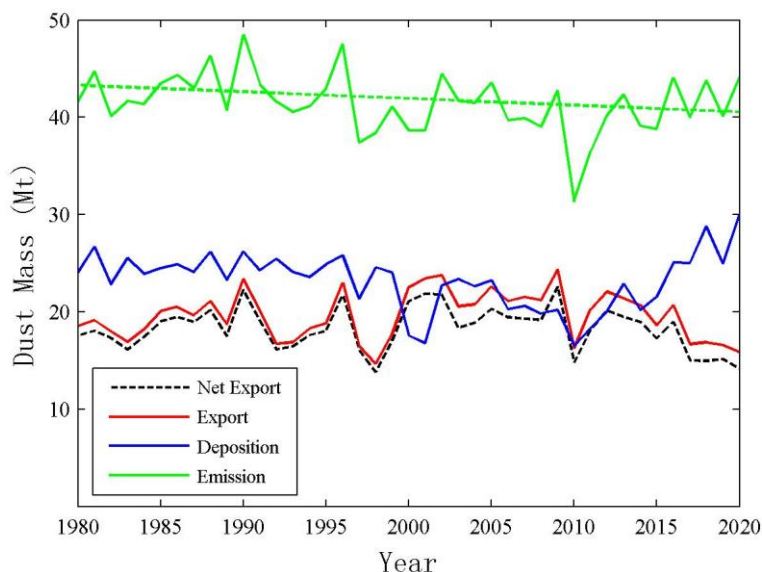
523  
524  
525  
526  
527  
528  
529  
530  
531  
532  
533  
534  
535  
536  
537  
538  
539  
540  
541  
542  
543  
544  
545  
546  
547  
548  
549  
550  
551

Figure 11 shows the annual dust budget for Australia from 1980 to 2020 using MERRA-2 aerosol reanalysis. The green line shows the annual total dust emission and suggests a slight decreasing trend over the past 42 years. Overall, the Australian continent emitted on average  $41.47 \pm 3.07$  Mt/yr into the atmosphere from 1980 to 2020, of which  $19.63 \pm 2.48$  Mt/yr is exported from Australia and  $23.19 \pm 2.97$  Mt/yr deposited over the land area. Additionally,  $1.34 \pm 0.55$  Mt dust is imported from non-Australia dust sources annually. The average annual dust emission over each decade from the 1980s to the 2010s are 42.74 Mt/yr, 42.25 Mt/yr, 40.99 Mt/yr, and 39.63 Mt/yr, respectively, showing an overall decreasing trend on a decade basis. The dust emission peaks were at 48.46 Mt/yr and 47.54 Mt/yr in 1990 and 1996, respectively. Although most of Australia was in drought during these two years, DSI based on visibility data and weather codes for these two years is around the average from 1965 to 2009 at 180 long-term stations (O’Loingsigh et al., 2014). The number of dust storms for these two years was around the average from 1960 to 2000 at both inland and coastal stations (Ekström et al., 2004). The discrepancy of this from the previous studies is caused by a lack of BoM weather observations in central Australia where dust emission is concentrated (Fig.12). Since 1996, a sharp reduction of 10.16 Mt/yr occurred in dust emission in 1997 and the change from 1998 to 2020 is relatively small (with a standard deviation of 2.25Mt/yr excluding 2010) to that for the period from 1980 to 1997. Dust emission reached its minimum at 31.6Mt/yr in 2010. This may be strongly related to high rainfall in Australia in 2010 when the annual rainfall was 687.3mm, which is only about 17mm lower than the highest rainfall of 710.6mm in 2000 during the period of 41 years (rainfall data can be found at [http://www.bom.gov.au/web01/ncc/www/cli\\_chg/timeseries/rain/0112/aus/latest.txt](http://www.bom.gov.au/web01/ncc/www/cli_chg/timeseries/rain/0112/aus/latest.txt)). The black dashed line in Figure 11 represents the net dust export that equals to the dust leaving the Australian coastline (total net export) minus that imported from west coastal (total net import) line from 20°S to 36°S on an annual basis. The blue line shows the total annual dust deposition over the continent, and was calculated using the mass balance equation (equation 9). The annual dust deposition shows a similar general trend to dust emission while opposite trend to dust export from 1980 to 2020. As annual dust deposition decreases the dust export increases (from 1980 to 2009). After a low value (16.24 Mt/yr and 16.53 Mt/yr) for both dust deposition and export in 2010, the annual dust export began to decrease, while annual dust deposition started to increase from 2010 to 2020. On a decade basis, the annual dust export reached its maximum (22.18Mt/yr) in the 2000s, while annual dust deposition reached the minimum of 20.72Mt/yr over the same period. During this decade, dust exported from Australia was the closest to dust imported.

552 
$$E + I = X + D \tag{9}$$

553 where  $E$ ,  $I$ ,  $X$ , and  $D$  represent annual dust emission, dust import, dust export, and dust deposition, respectively.  $E$   
 554 was calculated using the MERRA-2 dust emission flux dataset for all particle bins, and  $I$  and  $X$  were calculated  
 555 using MERRA-2 horizontal dust flux datasets. With equation 9,  $D$  can be evaluated.

556  
 557



558  
 559 **Figure 11: Annual dust budget for Australia. Green: annual dust emission, green dashed: trend of annual**  
 560 **dust emission, red: dust export, blue: dust deposition, and black dashed: net dust export (net export-,**  
 561 **export-, deposition is the difference between emission and net export).**

562  
 563 Table 2 presents details of the annual dust loading for Australia in terms of clay and silt. Clay is a fine particle,  
 564 traditionally ranging from 0.1~1.0 $\mu$ m in radius, which corresponds to MERRA-2 dust bin001. Silt is a much  
 565 coarser particle with a broad size range of roughly 1.0 $\mu$ m to 25.0 $\mu$ m in radius. Although in the MERRA-2 dataset,  
 566 the sum of dust bin002 to bin005 only covers 1.0 $\mu$ m to 10 $\mu$ m, in this study, the sum of dust bin002 to bin005 was  
 567 regarded as silt particles. Generally, clay accounts for 6.63 $\pm$ 0.58% of the total dust emission in Australia and silt  
 568 for 93.36 $\pm$ 6.84%. Ratios exceeding 86% of silt particles in dust deposition suggest that fine particles are more  
 569 likely to be transported and exported from the Australian continent.

570  
 571  
 572

573 Table 2. Annual dust budget for Australia in terms of clay and silt. Deposition\* equals  $E + I - X$  (equation 9)

	Particle size (Radius: $\mu\text{m}$ )	Emission (Mt/yr)	Dry Deposition (Mt/yr)	Wet Deposition (Mt/yr)	Deposition* (Mt/yr)
Total	0.1~10.0	41.47 $\pm$ 3.07	2.63 $\pm$ 0.34	4.44 $\pm$ 0.76	23.19 $\pm$ 2.97
Clay	0.1~1.0	2.75 $\pm$ 0.24	0.35 $\pm$ 0.05	0.55 $\pm$ 0.10	2.86 $\pm$ 0.23
		6.63 $\pm$ 0.10%	13.42 $\pm$ 0.46%	12.27 $\pm$ 0.28%	12.45 $\pm$ 0.97%
Silt	1.0~10.0	38.72 $\pm$ 2.84	2.28 $\pm$ 0.30	3.90 $\pm$ 0.66	20.33 $\pm$ 2.78
		93.37 $\pm$ 0.10%	86.58 $\pm$ 0.46%	87.73 $\pm$ 0.28%	87.55 $\pm$ 0.97%

574 **4. Discussion**

575 There are differences in the spatial distribution of dust activity in Australia based on different indicators from  
576 multiple data sets. Meteorological visibility-based dust event database (DEDB) and DSI have often been used to  
577 indicate the level of dust activity in Australia over the past several decades (Ekström et al., 2004; McTainsh et al.,  
578 2011a; McTainsh and Pitblado, 1987; McTainsh et al., 1990, 1989; McTainsh and Boughton, 1993; McTainsh et  
579 al., 1998, 2011b; O’Loingsigh et al., 2014). Although these indicators are quite capable of identifying the type of  
580 dust events and the dust source areas, dust severity using meteorological observations is limited because 1)  
581 definitions of dust events change over time, 2) only the most important type was recorded, 3) there is inconsistency  
582 in records at different meteorological sites, 4) synoptic observations are subjective, 5) BoM sites are sparse in  
583 remote areas (McTainsh et al., 2011a; McTainsh and Pitblado, 1987; O’Loingsigh et al., 2010; Strong et al., 2011).  
584 DEDB/DSI can be a valuable reference dataset for assessing remote sensing and reanalysis products due to the  
585 long-term coverage and distribution of BoM weather stations. The spatial distribution of dust activities identified  
586 with DEDB/DSI differs from that based on M&M and MERRA-2, including:

587 1) M&M DAOD shows that dust activities are most severe over the main dust source area, the Lake Eyre Basin,  
588 and along major dust pathways over eastern Australia from 2002 to 2020 while the atmosphere is relatively  
589 clean over Western Australia. DEDB/DSI (McTainsh et al., 2011a) and MERRA-2 near-surface dust  
590 concentration show not only high dust concentration over dust source areas in central Australia but also  
591 elevated dust concentration over downwind areas in the southeast and northwest of Australia. The main  
592 difference between the latter two is that MERRA-2 is able to quantify dust activity with near-surface dust  
593 concentrations and its variation over the Lake Eyre Basin, Nullarbor Plain, and downwind areas; on the  
594 contrary, DEDB/DSI can only indicate dust activity at a few sites.

595 2) Although M&M DAOD shows high dust concentrations over eastern Australia in spring and summer, its  
596 spatial pattern is dissimilar to that of DEDB/DSI. The dusty season is spring for the northern part of Australia

597 and summer for the southern part using DEDB (McTainsh et al., 1998), while two regions with high DAOD  
598 were found over northern Australia in summer (as shown in red boxed in figure 8), with low  
599 photosynthetically active vegetation (figure 12c). In another study based on MODIS DB aerosol product  
600 conducted by Ginoux et al. (2012), these two regions (the Barkly Tableland and the lee side of the Great  
601 Dividing Range) were found in spring, which differs from this study but coincides with the work of McTainsh  
602 et al. (1998). This is probably because 1) McTainsh et al. (1998) use meteorological data from 1960-1987  
603 while the MODIS DB data from 2003 to 2009 (Ginoux et al., 2012), while MODIS DB data from Aqua from  
604 2002 to 2020 have been used in this study, 2) MODIS DB shows much higher AOD over these two regions  
605 than MAIAC AOD (Shaylor et al., 2022), and MODIS DB retains high AOD for thick dust plumes (Che et  
606 al., 2022). For example, for the most severe two dust storms over Australia in the last twenty years, MODIS  
607 DB shows high AOD retrievals for the main dust plumes, which are even higher than 3.0, and the closest  
608 AERONET AOD to satellite Aqua overpass time is much less than MODIS DB (Che et al., 2022). This  
609 difference also indicates that more validation is still needed for MODIS DB aerosol products in Australia.

610 3) Another difference between the M&M dataset and DEDB/DSI and MERRA-2 is that the former shows a high  
611 level of dust activity over the southwest of Western Australia in spring and summer while the latter two don't.

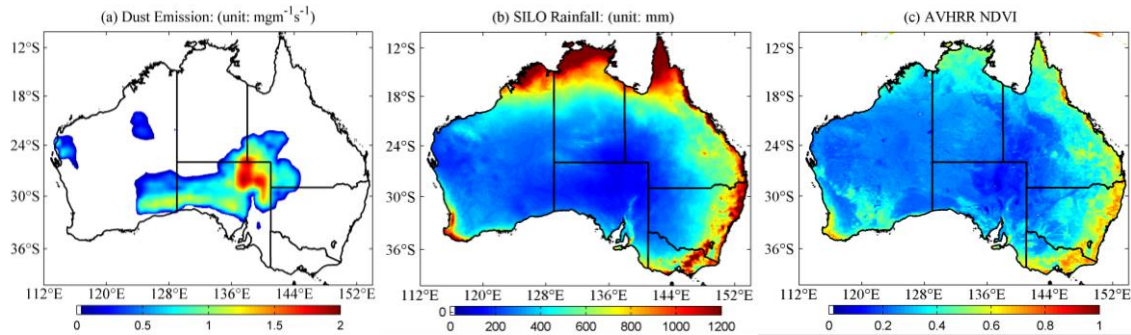
612  
613 Although the MERRA-2 dust dataset is well-constrained globally and even used as reference data to validate  
614 output from other models (Wu et al., 2020), the dust emission simulation for Australia in different studies varies  
615 considerably (Chen et al., 2022). MERRA-2 adopts the dust emission scheme proposed by Ginoux et al. (2001),  
616 which is based on soil wetness and surface wind speed. Discrepancy in the dust emission scheme may lead to  
617 different dust simulation outputs. Wu et al. (2020) compared MERRA-2 dust emission for different regions with  
618 15 CIMP models and showed that the annual dust emission for Australia varies from 0.6 Mt/yr to 2278 Mt/yr and  
619 only three out of 15 models (one with Ginoux's dust emission scheme) output similar dust emission to MERRA-  
620 2. Chen et al. (2022) compared annual dust emissions in nine studies and found that the annual dust emission for  
621 Australia in these studies varies within a relatively small range from 37 Mt/yr to 163 Mt/yr. Moreover, estimates  
622 of dust loading based on ground-based data for a single dust event reveal that MERRA-2 may underestimate dust  
623 emission in Australia. McTainsh et al. (2005) point out that published studies are very likely to overestimate dust  
624 loading under the assumption that dust concentration is uniform from the bottom to the top and recalculated dust  
625 loadings for dust storms on 8 February 1983 Melbourne, 20–30 May 1994 South Australia, 1 December 1987, and

626 23 October 2003, which were 1.23Mt, 3.3-6.4Mt, less than 3.35-4.85Mt, and 3.35-4.85Mt, respectively. These  
627 single dust events over a region produced close or even more dust than MERRA-2 monthly total dust emission  
628 from the entire Australia (3.58Mt, 2.35Mt, 4.76Mt, and 5.49Mt for corresponding months), indicating that  
629 MERRA-2 might underestimate dust emission in Australia. Additionally, AOD (Che et al., 2022) and PM10  
630 observations (Figure 5) during nationwide dust storms both provide evidence that MERRA-2 simulated dust  
631 concentrations are considerably lower than observations. As a negative result, annual dust emission estimates for  
632 high emission years, especially 2002, 2009, and 2019, were significantly underestimated, introducing uncertainty  
633 into dust emission analyses in Australia. For example, it is highly probable that the annual dust emission trend in  
634 the 2000s and 2010s is underestimated (figure 11).

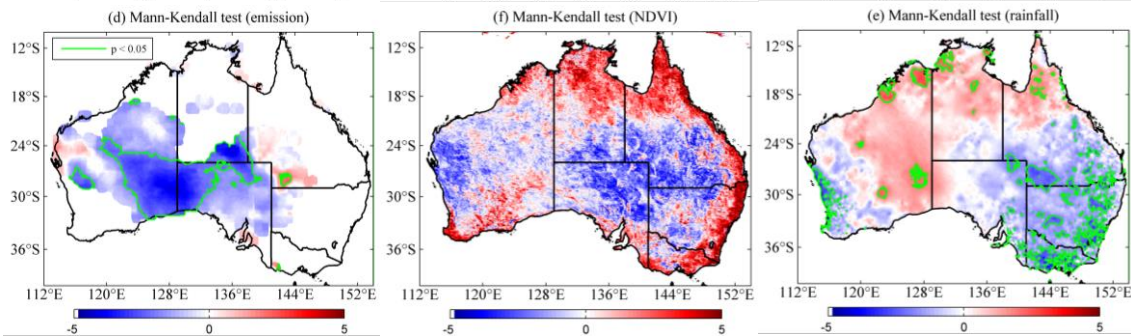
635  
636 As dust emissions are primarily determined by soil moisture content and surface wind speed (Ginoux et al., 2001),  
637 rainfall is strongly correlated with dust emissions (Figures 12a and b). Areas of dust emission, such as the center  
638 of the Lake Eyre Basin and the Nullarbor Plain, are typically associated with low rainfall, especially below  
639 150mm/yr. The Mann-Kendall (MK) tests conducted on annual dust emission and rainfall data from 1980 to 2020  
640 further highlight the substantial inhibitory impact of rainfall on dust emissions. A decreasing trend of dust emission  
641 occurred over the past 40 years with increased rainfall for almost all regions, especially in southwest of WA  
642 ((figure 12e). Northern Australia commonly shows an increasing trend of rainfall, while dust emissions remained  
643 essentially unchanged. This is because the highly vegetated surface in Northern Australia rarely emits dust  
644 particles. Conversely, with significant decreasing rainfall ( $p < 0.05$ ), the southwest of QLD, as a part of the Lake  
645 Eyre Basin, shows an increasing trend of dust emissions. Nevertheless, the impact of photosynthetically active  
646 vegetation on dust emissions was ignored in Ginoux's dust emission scheme (figure 12f), potentially resulting in  
647 uncertainties in dust emission estimates. Despite sharing a similar spatial pattern with rainfall trends, NDVI trends  
648 show an opposite trend to dust emissions in most of the Lake Eyre Basin. This indicates that the decreasing trend  
649 of photosynthetically active vegetation cover also contributes to the increasing trend of dust emissions in the  
650 southwest of QLD. It is essential to acknowledge and consider this factor in dust emission estimations.

651

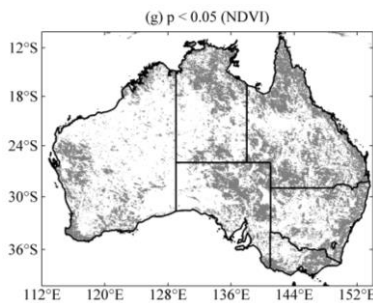
652



653



654



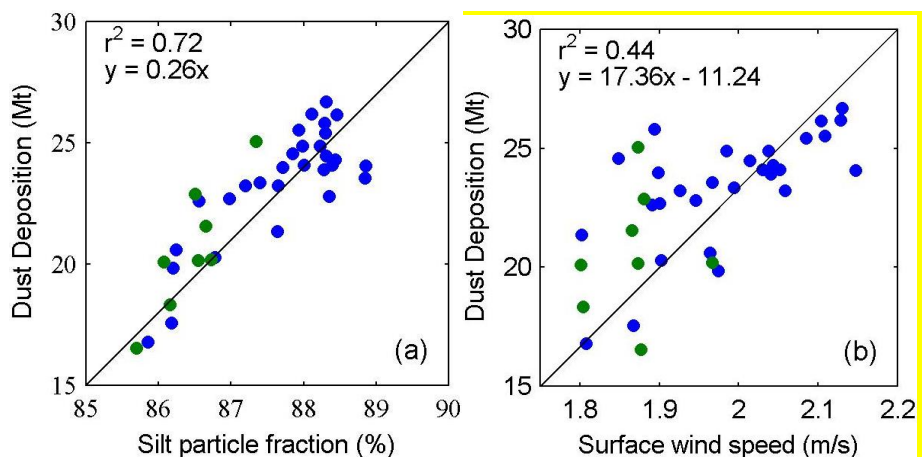
655 **Figure 12: Mean annual MERRA-2 dust emission (a), SILO mean annual rainfall (b), and (c) AVHRR**  
 656 **NDVI, and (d) Mann-Kendall (MK) test for annual dust emission, (e) for annual rainfall (f) for annual**  
 657 **NDVI from 1982 to 2019. Positives and negatives represent an increasing and decreasing trend, respectively.**  
 658 **A p-value < 0.05 for MK test is shown with green lines for dust emission (d) and rainfall (e) and gray colors**  
 659 **for NDVI in (g).**

660

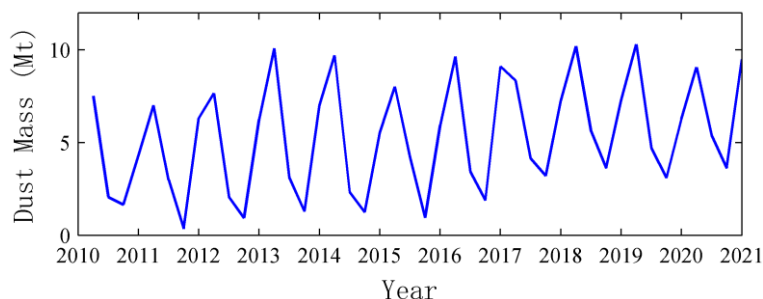
661 The post-2010 trend of dust deposition is likely caused by surface wind speed. Annual dust deposition shows  
 662 a strong correlation with the silt particle fraction (figure 13a) for the period from 1980-2010 (blue points) and  
 663 post-2010 (green points)., Large particles have a higher settling velocity. As more than 86% of dust was deposited  
 664 on the land and dust emission and import remained unchanged since 2000. A higher fraction of silt particles  
 665 indicates that deposited materials are made up of more silt particles. Nevertheless, as surface wind speed a key  
 666 factor in dust emission estimation, an increasing surface wind speed would entrain more coarse dust particles,  
 667 leading to more dust deposition on the land, as indicated by an  $r^2$  value of 0.44 between dust deposition with



668 surface wind speed (figure 13b). Additionally, figure 14 shows seasonal variations in the amount of dust deposition  
 669 in mainland Australia with the maximum deposition occurring in summer (Dec-Feb), and the minimum in autumn  
 670 (Sep-Nov). It is clear from figure 14 that the minimum deposition has increased from less than 1 Mt in 2011 to  
 671 more than 3 Mt over the past several years (2017-2020), while the peak deposition has not changed significantly  
 672 over the 11 years (Figure 14).



673  
 674 **Figure 13: Relationships between annual dust deposition and (a) silt particle fraction, and (b) surface wind**  
 675 **speed (McVicar, 2010) 1980-2016. Green points represent data since 2010; Indicators in each panel are**  
 676 **calculated using all data points.**



677  
 678 **Figure 14: Seasonal dust deposition on mainland Australia since 2000**

679  
 680  
 681 There is an inconsistency between dust emission and dust wet/dry deposition in Australia using the MEERRA-2  
 682 product. The imbalance between MERRA-2 dust emission and deposition could be caused by the incremental  
 683 update procedure in MERRA-2 and a lack of data assimilation for non-AOD parameters (Wu et al., 2020). Only  
 684 AOD in MERRA-2 was constrained by AOD observations from multiple sources (Randles et al., 2017), leading  
 685 to the bias in the dust component in the underlying aerosol model and in biased dust emission data that could not  
 686 be corrected. The average dust emission (41.47Mt/yr) plus dust import (1.34Mt/yr) is much higher than the

687 averaged dust dry/wet deposition (7.07Mt/yr) plus dust export (19.63Mt/yr) from 1980 to 2020. The difference of  
688 approximately 16.11Mt/yr or 38% suggests that the MERRA-2 dust deposition needs to be adjusted to balance  
689 with dust emission.

690

691 The long-term broad scale of dust activities in this study would provide useful information on dust activities in  
692 Australia. Most dust studies in Australia are based on ground-based observations. In the early 1990s, Yu et al.  
693 (1992, 1993) investigated the impact of rainfall in dust source areas in the previous autumn on dust event days in  
694 Mildura, Australia in the following summer. At a much larger scale, McTainsh et al. (1998) find that climatic  
695 drivers, i.e. wind run and soil moisture, affect dust storm frequencies in the north and south parts of eastern  
696 Australia in different ways during the dust season based on BoM meteorological observations. On the basis of  
697 McTainsh et al. (1998), Ekström et al. (2004) investigated the relationship between Australian dust storms and  
698 synoptic pressure distributions and find that spring-summer dust storms over central and southern Australia are  
699 most likely controlled by cold fronts with no precipitation and the summer-autumn dust storms are most likely  
700 controlled by the driest period of the year. Speer (2013) finds that westerly-induced dust storms that transport dust  
701 to the east coast tend to occur during El Nino years and positive and negative phases of the southern annular mode  
702 (SAM). Compared to ground-based observations, satellite remote sensing data provide dust related parameters at  
703 a broader scale. For example, with remote sensing products, Yu and Ginoux (2021) assess how ENSO and the  
704 Madden-Julian Oscillation (MJO) influence dust activities in Australia. They further show that during MJO phases  
705 dust activities are impacted by anomalies in convection and wind due to MJO and soil moisture and vegetation  
706 due to ENSO. This study provides the spatial pattern of dust activities in Australia using MERRA-MODIS  
707 combined DAOD and MERRA-2 PM10 and near-surface dust concentration. The findings in this study serve as  
708 an extension of previous studies, deepen our understanding of the spatial pattern of dust activities in Australia,  
709 and provide useful information on dust activities for future dust research in Australia.

## 710 5. Conclusion

711 On the basis of Che et al. (2022), this paper built a DAOD dataset based on MODIS DB and MERRA-2 aerosol  
712 datasets. Additionally, it has validated the MERRA-2 near-surface dust concentration, MERRA-2 estimated PM10,  
713 major dust pathways, MERRA-2-MODIS combined DAOD (M&M) with collected ground-based data and with  
714 these data sets analyzed the spatial and temporal distribution of dust activity over Australia from 1980-2020. The

715 M&M DAOD dataset was found to be of acceptable accuracy in Australia compared with AERONET data. M&M  
716 DAOD contributes to long-term dust research in Australia. A power law relationship (similar to previous studies)  
717 has been found between MERRA-2 hourly near-surface dust concentration and BoM manual horizontal visibility  
718 at three sites in eastern Australia. Monthly MERRA-2 estimated PM10 show similar variations with AQMN  
719 ground-based PM10 observations with an  $r^2$  value from 0.14 to 0.44 at 6 selected sites; however, MERRA-2 PM10  
720 is not sensitive to low PM10 in winter, peaks in summer, and is very likely to miss extreme high monthly PM10  
721 values. MERRA-2 horizontal flux in MERRA-2 aerosol reanalysis shows similar general dust pathways (Bowler,  
722 1976; Sprigg, 1982a) suggesting that both early simulations and MERRA-2 are all reliable in identifying dust  
723 pathways. Moreover, MERRA-2 further provides quantitative details on dust concentration and fluxes in a  
724 spatially consistent manner, with notable underestimation during high concentration dust events.

725

726 Dust events over Australia are shown to be concentrated in the north and southeast in spring (Sep-Nov) and occur  
727 anywhere to the east in summer, with the dust season finishing in autumn (based on M&M DAOD). Four main  
728 dust regions have been identified. These include the southwest of Western Australia and the north and south of  
729 eastern Australia. The center of the Lake Eyre Bains, which emits dust throughout the entire year, is identified as  
730 the dustiest region. Dust events over the southwest of Western Australia only span two months, starting in  
731 September and reaching their peak in October. Dust events to the north of eastern Australia start in October,  
732 gradually reaching a peak in December and January, ending in April.

733

734 Near-surface dust concentrations were found to be the highest (over  $200\mu\text{g}/\text{m}^3$ ) over the center of Lake Eyre Basin,  
735 and weakened radially according to distance from the center, decreasing to below  $20\mu\text{g}/\text{m}^3$  along the two main  
736 pathways to the southwest and northeast. The dust pathway in the southeast shows lower near-surface dust  
737 concentrations than the northeast, coinciding with the fact that dust entrained in central Australia is hardly  
738 transported to the east coast (Speer, 2013). This is also shown by the ratio of near-surface dust concentration to  
739 PM10, where high values are concentrated around central Australia, and relatively low ratios (below 0.5) are found  
740 in eastern Australia.

741

742 Total dust emission was estimated to be 40 Mt (mega-tonnes) per year over the period 1980-2020, of which nearly  
743 50% was deposited on land; the rest as net export from the Australian continent. The average annual dust emission

744 over each decade from the 1980s to the 2010s are 42.74 Mt/yr, 42.25 Mt/yr, 40.99 Mt/yr, and 39.63 Mt/yr,  
745 respectively, showing an overall decreasing trend on a decade basis, while uncertainties exist due to the  
746 underestimation of MERRA-2 during high concentration dust events. In the 2000s, more dust was exported than  
747 over other periods, 22.18 Mt/yr vs. 18.80 Mt/yr, and the closest to dust deposition (20.72 Mt/yr, the lowest);  
748 however, approximately 23.19Mt/yr was deposited over the land area over other periods. Among these particles,  
749  $6.63\pm 0.10\%$  ( $2.75\pm 0.24\text{Mt/yr}$ ) of emission was clay particles and almost all dust deposition ( $87.55\pm 0.87\%$ )  
750 consisted of silt. This indicates that exported dust from Australia is mainly composed of fine particles (clay).  
751 Additionally, dust import was identified from the north, south, and west coastlines using MERRA-2 flux data.  
752 Only dust across the coastline in the southwest of Western Australia was genuinely imported from other continents  
753 while other imported dusts are sourced and recycled from exported dusts from Australia across the north and south  
754 coastlines.

755

756 *Data availability.* The MERRA-2 and MODIS DB products are publically available from  
757 <https://search.earthdata.nasa.gov/search>. The PM10 data can be downloaded from  
758 <http://www.environment.nsw.gov.au/AQMS/search.htm> and the hourly horizontal visibility is available from the  
759 Australian Bureau of Meteorology. The SIIO rainfall data can be accessed from  
760 <https://www.longpaddock.qld.gov.au/silo/gridded-data/>.

761

762 *Author contributions.* Yahui Che: Term, Conceptualization, Data curation, Methodology, Software, Visualization,  
763 Investigation, Writing – original draft, preparation. Bofu Yu: Supervision, Writing – review & editing. Katherine  
764 Parsons: Writing – review & editing.”

765

766 *Conflict of interest statement.* The authors declare that they have no known competing financial interests or  
767 personal relationships that could have appeared to influence the work reported in this paper.

768

769 *Acknowledgement.* We are thankful to the Australian Bureau of Meteorology for observing and maintaining the  
770 horizontal data. The first author acknowledges Griffith University for the financial support provided through the  
771 Griffith University International Postgraduate Research Scholarship (GUIPRS) and the Griffith University  
772 Postgraduate Research Scholarship (GUPRS).

773

774 **References**

- 775 Ackerman, S. A.: Using the radiative temperature difference at 3.7 and 11  $\mu\text{m}$  to track dust outbreaks, *Remote*  
776 *Sens. Environ.*, 27, 129–133, [https://doi.org/10.1016/0034-4257\(89\)90012-6](https://doi.org/10.1016/0034-4257(89)90012-6), 1989.
- 777 Angstrom, A.: Solar and terrestrial radiation. Report to the international commission for solar research on  
778 actinometric investigations of solar and atmospheric radiation, *Q. J. R. Meteorol. Soc.*, 50, 121–126,  
779 <https://doi.org/10.1002/qj.49705021008>, 1924.
- 780 Baddock, M. C., Bullard, J. E., and Bryant, R. G.: Dust source identification using MODIS: A comparison of  
781 techniques applied to the Lake Eyre Basin, Australia, *Remote Sens. Environ.*, 113, 1511–1528,  
782 <https://doi.org/10.1016/j.rse.2009.03.002>, 2009.
- 783 Baddock, M. C., Strong, C. L., Leys, J. F., Heidenreich, S. K., Tews, E. K., and McTainsh, G. H.: A visibility and  
784 total suspended dust relationship, *Atmos. Environ.*, 89, 329–336, <https://doi.org/10.1016/j.atmosenv.2014.02.038>,  
785 2014.
- 786 Bauer, S. E., Koch, D., Unger, N., Metzger, S. M., Shindell, D. T., and Streets, D. G.: Nitrate aerosols today and  
787 in 2030: a global simulation including aerosols and tropospheric ozone, *Atmos. Chem. Phys.*, 7, 5043–5059,  
788 <https://doi.org/10.5194/acp-7-5043-2007>, 2007.
- 789 BBC News: Australia weather: How much rain did it take to put out NSW fires?, 2020.
- 790 Bhattachan, A. and D’Odorico, P.: Can land use intensification in the Mallee, Australia increase the supply of  
791 soluble iron to the Southern Ocean?, *Sci. Rep.*, 4, 6009, <https://doi.org/10.1038/srep06009>, 2014.
- 792 Bhattachan, A., D’Odorico, P., Baddock, M. C., Zobeck, T. M., Okin, G. S., and Cassar, N.: The Southern Kalahari:  
793 a potential new dust source in the Southern Hemisphere?, *Environ. Res. Lett.*, 7, 024001,  
794 <https://doi.org/10.1088/1748-9326/7/2/024001>, 2012.
- 795 Blewett, R.: *Shaping a Nation: A Geology of Australia*, Commonwealth of Australia and ANU-E Press, Canberra,  
796 2012.
- 797 Bowler, J. M.: Aridity in Australia: Age, origins and expression in aeolian landforms and sediments, *Earth-Science*  
798 *Rev.*, 12, 279–310, [https://doi.org/10.1016/0012-8252\(76\)90008-8](https://doi.org/10.1016/0012-8252(76)90008-8), 1976.
- 799 Buchard, V., Randles, C. A., da Silva, A. M., Darmenov, A., Colarco, P. R., Govindaraju, R., Ferrare, R., Hair, J.,  
800 Beyersdorf, A. J., Ziemba, L. D., and Yu, H.: The MERRA-2 aerosol reanalysis, 1980 onward. Part II: Evaluation  
801 and case studies, *J. Clim.*, 30, 6851–6872, <https://doi.org/10.1175/JCLI-D-16-0613.1>, 2017.
- 802 Bullard, J., Baddock, M., McTainsh, G., and Leys, J.: Sub-basin scale dust source geomorphology detected using

803 MODIS, *Geophys. Res. Lett.*, 35, 15404, <https://doi.org/10.1029/2008GL033928>, 2008.

804 Chan, Y.-C., Cohen, D. D., Hawas, O., Stelcer, E., Simpson, R., Denison, L., Wong, N., Hodge, M., Comino, E.,  
805 and Carswell, S.: Apportionment of sources of fine and coarse particles in four major Australian cities by positive  
806 matrix factorisation, *Atmos. Environ.*, 42, 374–389, <https://doi.org/10.1016/j.atmosenv.2007.09.030>, 2008.

807 Che, Y., Xue, Y., Mei, L., Guang, J., She, L., Guo, J., Hu, Y., Xu, H., He, X., Di, A., and Fan, C.: Technical note:  
808 Intercomparison of three AATSR Level 2 (L2) AOD products over China, *Atmos. Chem. Phys.*, 16, 9655–9674,  
809 <https://doi.org/10.5194/acp-16-9655-2016>, 2016.

810 Che, Y., Xue, Y., Guang, J., She, L., and Guo, J.: Evaluation of the AVHRR DeepBlue aerosol optical depth  
811 dataset over mainland China, *ISPRS J. Photogramm. Remote Sens.*, 146, 74–90,  
812 <https://doi.org/10.1016/j.isprsjprs.2018.09.004>, 2018.

813 Che, Y., Yu, B., Parsons, K., Desha, C., and Ramezani, M.: Evaluation and comparison of MERRA-2 AOD and  
814 DAOD with MODIS DeepBlue and AERONET data in Australia, *Atmos. Environ.*, 277, 119054,  
815 <https://doi.org/10.1016/j.atmosenv.2022.119054>, 2022.

816 Chen, L., Mengersen, K., and Tong, S.: Spatiotemporal relationship between particle air pollution and respiratory  
817 emergency hospital admissions in Brisbane, Australia, *Sci. Total Environ.*, 373, 57–67,  
818 <https://doi.org/10.1016/j.scitotenv.2006.10.050>, 2007.

819 Chen, W., Meng, H., Song, H., and Zheng, H.: Progress in Dust Modelling, Global Dust Budgets, and Soil Organic  
820 Carbon Dynamics, *Land*, 11, 176, <https://doi.org/10.3390/land11020176>, 2022.

821 Chepil, W. S. and Woodruff, N. P.: Sedimentary characteristics of dust storms; Part II, Visibility and dust  
822 concentration, *Am. J. Sci.*, 255, 104–114, <https://doi.org/10.2475/ajs.255.2.104>, 1957.

823 Cowie, G., Lawson, W., and Kim, N.: Australian dust causing respiratory disease admissions in some North Island,  
824 New Zealand Hospitals, *N. Z. Med. J.*, 123, 87–88, 2010.

825 De Deckker, P.: An evaluation of Australia as a major source of dust, *Earth-Science Rev.*, 194, 536–567,  
826 <https://doi.org/10.1016/j.earscirev.2019.01.008>, 2019.

827 Desservettaz, M., Paton-Walsh, C., Griffith, D. W. T., Kettlewell, G., Keywood, M. D., Vanderschoot, M. V.,  
828 Ward, J., Mallet, M. D., Milic, A., Miljevic, B., Ristovski, Z. D., Howard, D., Edwards, G. C., and Atkinson, B.:  
829 Emission factors of trace gases and particles from tropical savanna fires in Australia, *J. Geophys. Res.*, 122, 6059–  
830 6074, <https://doi.org/10.1002/2016JD025925>, 2017.

831 Di, A., Xue, Y., Yang, X., Leys, J., Guang, J., Mei, L., Wang, J., She, L., Hu, Y., He, X., Che, Y., and Fan, C.:

832 Dust aerosol optical depth retrieval and dust storm detection for Xinjiang Region using Indian national satellite  
833 observations, *Remote Sens.*, 8, 702, <https://doi.org/10.3390/rs8090702>, 2016.

834 Domínguez-Rodríguez, A., Báez-Ferrer, N., Abreu-González, P., Rodríguez, S., Díaz, R., Avanzas, P., and  
835 Hernández-Vaquero, D.: Impact of Desert Dust Events on the Cardiovascular Disease: A Systematic Review and  
836 Meta-Analysis, *J. Clin. Med.*, 10, 727, <https://doi.org/10.3390/jcm10040727>, 2021.

837 Dubovik, O. and King, M. D.: A flexible inversion algorithm for retrieval of aerosol optical properties from Sun  
838 and sky radiance measurements, *J. Geophys. Res. Atmos.*, 105, 20673–20696,  
839 <https://doi.org/10.1029/2000JD900282>, 2000.

840 Ekström, M., McTainsh, G. H., and Chappell, A.: Australian dust storms: temporal trends and relationships with  
841 synoptic pressure distributions (1960-99), *Int. J. Climatol.*, 24, 1581–1599, <https://doi.org/10.1002/joc.1072>, 2004.

842 Fryrear, D. W., Stout, J. E., Hagen, L. J., and Vories, E. D.: Wind erosion: field measurement and analysis, *Trans.*  
843 *Am. Soc. Agric. Eng.*, 34, 155–160, <https://doi.org/10.13031/2013.31638>, 1991.

844 Gelaro, R., McCarty, W., Suárez, M. J., Todling, R., Molod, A., Takacs, L., Randles, C. A., Darmenov, A.,  
845 Bosilovich, M. G., Reichle, R., Wargan, K., Coy, L., Cullather, R., Draper, C., Akella, S., Buchard, V., Conaty,  
846 A., da Silva, A. M., Gu, W., Kim, G. K., Koster, R., Lucchesi, R., Merkova, D., Nielsen, J. E., Partyka, G., Pawson,  
847 S., Putman, W., Rienecker, M., Schubert, S. D., Sienkiewicz, M., and Zhao, B.: The modern-era retrospective  
848 analysis for research and applications, version 2 (MERRA-2), *J. Clim.*, 30, 5419–5454,  
849 <https://doi.org/10.1175/JCLI-D-16-0758.1>, 2017.

850 Giles, D. M., Sinyuk, A., Sorokin, M. G., Schafer, J. S., Smirnov, A., Slutsker, I., Eck, T. F., Holben, B. N., Lewis,  
851 J. R., Campbell, J. R., Welton, E. J., Korkin, S. V., and Lyapustin, A. I.: Advancements in the Aerosol Robotic  
852 Network (AERONET) Version 3 database - Automated near-real-time quality control algorithm with improved  
853 cloud screening for Sun photometer aerosol optical depth (AOD) measurements, *Atmos. Meas. Tech.*, 12, 169–  
854 209, <https://doi.org/10.5194/amt-12-169-2019>, 2019.

855 Ginoux, P., Chin, M., Tegen, I., Prospero, J. M., Holben, B., Dubovik, O., and Lin, S.-J.: Sources and distributions  
856 of dust aerosols simulated with the GOCART model, *J. Geophys. Res. Atmos.*, 106, 20255–20273,  
857 <https://doi.org/10.1029/2000JD000053>, 2001.

858 Ginoux, P., Garbuzov, D., and Hsu, N. C.: Identification of anthropogenic and natural dust sources using moderate  
859 resolution imaging spectroradiometer (MODIS) deep blue level 2 data, *J. Geophys. Res. Atmos.*, 115, 1–10,  
860 <https://doi.org/10.1029/2009JD012398>, 2010.



861 Ginoux, P., Prospero, J. M., Gill, T. E., Hsu, N. C., and Zhao, M.: Global-scale attribution of anthropogenic and  
862 natural dust sources and their emission rates based on MODIS Deep Blue aerosol products, *Rev. Geophys.*, 50,  
863 1–36, <https://doi.org/10.1029/2012RG000388>, 2012.

864 Gkikas, A., Proestakis, E., Amiridis, V., Kazadzis, S., Di Tomaso, E., Tsekeri, A., Marinou, E., Hatzianastassiou,  
865 N., and Pérez Garcíá-Pando, C.: ModIs Dust AeroSol (MIDAS): A global fine-resolution dust optical depth data  
866 set, *Atmos. Meas. Tech.*, 14, 309–334, <https://doi.org/10.5194/amt-14-309-2021>, 2021.

867 Goudie, A. S.: Desert dust and human health disorders, *Environ. Int.*, 63, 101–113,  
868 <https://doi.org/10.1016/j.envint.2013.10.011>, 2014.

869 Hsu, N. C., Tsay, S. C., King, M. D., and Herman, J. R.: Aerosol properties over bright-reflecting source regions,  
870 *IEEE Trans. Geosci. Remote Sens.*, 42, 557–569, <https://doi.org/10.1109/TGRS.2004.824067>, 2004.

871 Hsu, N. C., Jeong, M. J., Bettenhausen, C., Sayer, A. M., Hansell, R., Seftor, C. S., Huang, J., and Tsay, S. C.:  
872 Enhanced Deep Blue aerosol retrieval algorithm: The second generation, *J. Geophys. Res. Atmos.*, 118, 9296–  
873 9315, <https://doi.org/10.1002/jgrd.50712>, 2013.

874 Jeffrey, S. J., Carter, J. O., Moodie, K. B., and Beswick, A. R.: Using spatial interpolation to construct a  
875 comprehensive archive of Australian climate data, *Environ. Model. Softw.*, 16, 309–330,  
876 [https://doi.org/10.1016/S1364-8152\(01\)00008-1](https://doi.org/10.1016/S1364-8152(01)00008-1), 2001.

877 de Jesus, A. L., Thompson, H., Knibbs, L. D., Hanigan, I., De Torres, L., Fisher, G., Berko, H., and Morawska,  
878 L.: Two decades of trends in urban particulate matter concentrations across Australia, *Environ. Res.*, 190, 110021,  
879 <https://doi.org/10.1016/j.envres.2020.110021>, 2020.

880 Kolmonen, P., Sogacheva, L., Virtanen, T. H., de Leeuw, G., and Kulmala, M.: The ADV/ASV AATSR aerosol  
881 retrieval algorithm: current status and presentation of a full-mission AOD dataset, *Int. J. Digit. Earth*, 9, 545–561,  
882 <https://doi.org/10.1080/17538947.2015.1111450>, 2016.

883 de Leeuw, G., Holzer-Popp, T., Bevan, S., Davies, W. H., Descloitres, J., Grainger, R. G., Griesfeller, J., Heckel,  
884 A., Kinne, S., Klüser, L., Kolmonen, P., Litvinov, P., Martynenko, D., North, P., Ovigneur, B., Pascal, N., Poulsen,  
885 C., Ramon, D., Schulz, M., Siddans, R., Sogacheva, L., Tanré, D., Thomas, G. E., Virtanen, T. H., von Hoyningen  
886 Huene, W., Vountas, M., and Pinnock, S.: Evaluation of seven European aerosol optical depth retrieval algorithms  
887 for climate analysis, *Remote Sens. Environ.*, 162, 295–315, <https://doi.org/10.1016/j.rse.2013.04.023>, 2015.

888 Levy, R. C., Mattoo, S., Munchak, L. A., Remer, L. A., Sayer, A. M., Patadia, F., and Hsu, N. C.: The Collection  
889 6 MODIS aerosol products over land and ocean, *Atmos. Meas. Tech.*, 6, 2989–3034, <https://doi.org/10.5194/amt->

890 6-2989-2013, 2013.

891 Leys, J. F., Heidenreich, S. K., Strong, C. L., McTainsh, G. H., and Quigley, S.: PM10 concentrations and mass  
892 transport during “ Red Dawn” - Sydney 23 September 2009, *Aeolian Res.*, 3, 327–342,  
893 <https://doi.org/10.1016/j.aeolia.2011.06.003>, 2011.

894 Love, B. M., Leys, J. F., Strong, C. L., and McTainsh, G. H.: Dust climatology of Mildura, Victoria, Australia:  
895 transport direction, *Earth Surf. Process. Landforms*, 44, 1449–1459, <https://doi.org/10.1002/esp.4587>, 2019.

896 Ma, X., Yan, P., Zhao, T., Jia, X., Jiao, J., Ma, Q., Wu, D., Shu, Z., Sun, X., and Habtemicheal, B. A.: Evaluations  
897 of surface pm10 concentration and chemical compositions in merra-2 aerosol reanalysis over central and eastern  
898 china, *Remote Sens.*, 13, 1317, <https://doi.org/10.3390/rs13071317>, 2021.

899 McGowan, H. A. and Clark, A.: A vertical profile of PM10 dust concentrations measured during a regional dust  
900 event identified by MODIS Terra, western Queensland, Australia, *J. Geophys. Res. Earth Surf.*, 113, 2–03,  
901 <https://doi.org/10.1029/2007JF000765>, 2008.

902 McGowan, H. A., McTainsh, G. H., Zawar-Reza, P., and Sturman, A. P.: Identifying regional dust transport  
903 pathways: Application of kinematic trajectory modelling to a trans-Tasman case, *Earth Surf. Process. Landforms*,  
904 25, 633–647, [https://doi.org/10.1002/1096-9837\(200006\)25:6<633::AID-ESP102>3.0.CO;2-J](https://doi.org/10.1002/1096-9837(200006)25:6<633::AID-ESP102>3.0.CO;2-J), 2000.

905 McTainsh, G. H. and Boughton, W. C.: *Land Degradation Processes in Australia*, Longman Cheshire, Melbourne,  
906 389 pp., <https://doi.org/10.3/JQUERY-UIJS>, 1993.

907 McTainsh, G. H. and Pitblado, J. R.: Dust storms and related phenomena measured from meteorological records  
908 in Australia, *Earth Surf. Process. Landforms*, 12, 415–424, <https://doi.org/10.1002/esp.3290120407>, 1987.

909 McTainsh, G. H., Burgess, R., and Pitblado, J. R.: Aridity, drought and dust storms in Australia (1960-84), *J. Arid*  
910 *Environ.*, 16, 11–22, [https://doi.org/10.1016/s0140-1963\(18\)31042-5](https://doi.org/10.1016/s0140-1963(18)31042-5), 1989.

911 McTainsh, G. H., Lynch, A. W., and Burgess, R. C.: Wind erosion in eastern australia, *Aust. J. Soil Res.*, 28, 323–  
912 339, <https://doi.org/10.1071/SR9900323>, 1990.

913 McTainsh, G. H., Lynch, A. W., and Tews, E. K.: Climatic controls upon dust storm occurrence in eastern  
914 Australia, *J. Arid Environ.*, 39, 457–466, <https://doi.org/10.1006/jare.1997.0373>, 1998.

915 McTainsh, G. H., Chan, Y. C., McGowan, H., Leys, J., and Tews, K.: The 23rd October 2002 dust storm in eastern  
916 Australia: Characteristics and meteorological conditions, *Atmos. Environ.*, 39, 1227–1236,  
917 <https://doi.org/10.1016/j.atmosenv.2004.10.016>, 2005.

918 McTainsh, G. H., Leys, J., O’Loingsigh, T., and Strong, C. L.: Update of Dust Storm Index ( DSI ) maps for 2005

919 to 2010 and re-analysis and mapping of DSI for for the Australian Collaborative Rangeland Information System  
920 ( ACRIS ), Canberra: DSEWPaC, 1992–2008 pp., 2011a.

921 McTainsh, G. H., Leys, J. F., O’Loingsigh, T., and Strong, C. L.: Wind erosion and land management in Australia  
922 during 1940-1949 and 2000-2009, Canberra, 2011b.

923 McVicar, T.: Near-Surface Wind Speed.v10., <https://doi.org/10.25919/5c5106acbc02>, 2010.

924 Mei, L., Xue, Y., de Leeuw, G., von Hoyningen-Huene, W., Kokhanovsky, A. A., Istomina, L., Guang, J., and  
925 Burrows, J. P.: Aerosol optical depth retrieval in the Arctic region using MODIS data over snow, *Remote Sens.*  
926 *Environ.*, 128, 234–245, <https://doi.org/10.1016/j.rse.2012.10.009>, 2013a.

927 Mei, L., Xue, Y., Kokhanovsky, A. A., von Hoyningen-Huene, W., Istomina, L., de Leeuw, G., Burrows, J. P.,  
928 Guang, J., and Jing, Y.: Aerosol optical depth retrieval over snow using AATSR data, *Int. J. Remote Sens.*, 34,  
929 5030–5041, <https://doi.org/10.1080/01431161.2013.786197>, 2013b.

930 Middleton, N. J.: Desert dust hazards: A global review, *Aeolian Res.*, 24, 53–63,  
931 <https://doi.org/10.1016/j.aeolia.2016.12.001>, 2017.

932 Mukkavilli, S. K., Prasad, A. A., Taylor, R. A., Huang, J., Mitchell, R. M., Troccoli, A., and Kay, M. J.:  
933 Assessment of atmospheric aerosols from two reanalysis products over Australia, *Atmos. Res.*, 215, 149–164,  
934 <https://doi.org/10.1016/j.atmosres.2018.08.026>, 2019.

935 O’Loingsigh, T., McTainsh, G. H., Tapper, N. J., and Shinkfield, P.: Lost in code: A critical analysis of using  
936 meteorological data for wind erosion monitoring, *Aeolian Res.*, 2, 49–57,  
937 <https://doi.org/10.1016/j.aeolia.2010.03.002>, 2010.

938 O’Loingsigh, T., McTainsh, G. H., Tews, E. K., Strong, C. L., Leys, J. F., Shinkfield, P., and Tapper, N. J.: The  
939 Dust Storm Index (DSI): A method for monitoring broadscale wind erosion using meteorological records, *Aeolian*  
940 *Res.*, 12, 29–40, <https://doi.org/10.1016/j.aeolia.2013.10.004>, 2014.

941 O’Loingsigh, T., Chubb, T., Baddock, M., Kelly, T., Tapper, N. J., de Deckker, P., and McTainsh, G.: Sources  
942 and pathways of dust during the Australian “millennium drought” decade, *J. Geophys. Res.*, 122, 1246–1260,  
943 <https://doi.org/10.1002/2016JD025737>, 2017.

944 Ou, Y., Li, Z., Chen, C., Zhang, Y., Li, K., Shi, Z., Dong, J., Xu, H., Peng, Z., Xie, Y., and Luo, J.: Evaluation of  
945 MERRA-2 Aerosol Optical and Component Properties over China Using SONET and PARASOL/GRASP Data,  
946 *Remote Sens.*, 14, <https://doi.org/10.3390/rs14040821>, 2022.

947 Pereira, G., Lee, H. J., Bell, M., Regan, A., Malacova, E., Mullins, B., and Knibbs, L. D.: Development of a model

948 for particulate matter pollution in Australia with implications for other satellite-based models, *Environ. Res.*, 159,  
949 9–15, <https://doi.org/10.1016/j.envres.2017.07.044>, 2017.

950 Prospero, J. M., Barkley, A. E., Gaston, C. J., Gatineau, A., Campos y Sansano, A., and Panechou, K.:  
951 Characterizing and Quantifying African Dust Transport and Deposition to South America: Implications for the  
952 Phosphorus Budget in the Amazon Basin, *Global Biogeochem. Cycles*, 34, e2020GB006536,  
953 <https://doi.org/10.1029/2020GB006536>, 2020.

954 Provençal, S., Buchard, V., da Silva, A. M., Leduc, R., and Barrette, N.: Evaluation of PM surface concentrations  
955 simulated by Version 1 of NASA’s MERRA Aerosol Reanalysis over Europe, *Atmos. Pollut. Res.*, 8, 374–382,  
956 <https://doi.org/10.1016/j.apr.2016.10.009>, 2017.

957 Pu, B. and Ginoux, P.: Projection of American dustiness in the late 21st century due to climate change, *Sci. Rep.*,  
958 7, <https://doi.org/10.1038/s41598-017-05431-9>, 2017.

959 Pu, B. and Ginoux, P.: How reliable are CMIP5 models in simulating dust optical depth?, *Atmos. Chem. Phys.*,  
960 18, 12491–12510, <https://doi.org/10.5194/acp-18-12491-2018>, 2018.

961 Qin, W., Zhang, Y., Chen, J., Yu, Q., Cheng, S., Li, W., Liu, X., and Tian, H.: Variation, sources and historical  
962 trend of black carbon in Beijing, China based on ground observation and MERRA-2 reanalysis data, *Environ.*  
963 *Pollut.*, 245, 853–863, <https://doi.org/10.1016/j.envpol.2018.11.063>, 2019.

964 Randles, C. A., da Silva, A. M., Buchard, V., Colarco, P. R., Darmenov, A., Govindaraju, R., Smirnov, A., Holben,  
965 B., Ferrare, R., Hair, J., Shinozuka, Y., and Flynn, C. J.: The MERRA-2 aerosol reanalysis, 1980 onward. Part I:  
966 System description and data assimilation evaluation, *J. Clim.*, 30, 6823–6850, [https://doi.org/10.1175/JCLI-D-16-](https://doi.org/10.1175/JCLI-D-16-0609.1)  
967 0609.1, 2017.

968 Roberts, S.: Have the short-term mortality effects of particulate matter air pollution changed in Australia over the  
969 period 1993-2007?, *Environ. Pollut.*, 182, 9–14, <https://doi.org/10.1016/j.envpol.2013.06.036>, 2013.

970 Rotstayn, L. D., Collier, M. A., Mitchell, R. M., Qin, Y., Campbell, S. K., and Dravitzki, S. M.: Simulated  
971 enhancement of ENSO-related rainfall variability due to Australian dust, *Atmos. Chem. Phys.*, 11, 6575–6592,  
972 <https://doi.org/10.5194/ACP-11-6575-2011>, 2011.

973 RUST, B. R. and NANSON, G. C.: Bedload transport of mud as pedogenic aggregates in modern and ancient  
974 rivers, *Sedimentology*, 36, 291–306, <https://doi.org/10.1111/j.1365-3091.1989.tb00608.x>, 1989.

975 Sayer, A. M., Hsu, N. C., Bettenhausen, C., and Jeong, M. J.: Validation and uncertainty estimates for MODIS  
976 Collection 6 “deep Blue” aerosol data, *J. Geophys. Res. Atmos.*, 118, 7864–7872,

977 <https://doi.org/10.1002/jgrd.50600>, 2013.

978 Sayer, A. M., Hsu, N. C., Lee, J., Carletta, N., Chen, S. H., and Smirnov, A.: Evaluation of NASA Deep  
979 Blue/SOAR aerosol retrieval algorithms applied to AVHRR measurements, *J. Geophys. Res. Atmos.*, 122, 9945–  
980 9967, <https://doi.org/10.1002/2017JD026934>, 2017.

981 Sayer, A. M., Hsu, N. C., Lee, J., Kim, W. V., and Dutcher, S. T.: Validation, Stability, and Consistency of  
982 MODIS Collection 6.1 and VIIRS Version 1 Deep Blue Aerosol Data Over Land, *J. Geophys. Res. Atmos.*, 124,  
983 4658–4688, <https://doi.org/10.1029/2018JD029598>, 2019.

984 Shao, Y.: Physics and Modelling of Wind Erosion, *Phys. Model. Wind Eros.*, 452, <https://doi.org/10.1007/978-1->  
985 4020-8895-7, 2009.

986 Shao, Y., Yang, Y., Wang, J., Song, Z., Leslie, L. M., Dong, C., Zhang, Z., Lin, Z., Kanai, Y., Yabuki, S., and  
987 Chun, Y.: Northeast Asian dust storms: Real-time numerical prediction and validation, *J. Geophys. Res. Atmos.*,  
988 108, 2003JD003667, <https://doi.org/10.1029/2003JD003667>, 2003.

989 Shao, Y., Leys, J. F., McTainsh, G. H., and Tews, K.: Numerical simulation of the October 2002 dust event in  
990 Australia, *J. Geophys. Res. Atmos.*, 112, 8207, <https://doi.org/10.1029/2006JD007767>, 2007.

991 Shao, Y., Klose, M., and Wyrwoll, K.-H.: Recent global dust trend and connections to climate forcing, *J. Geophys.*  
992 *Res. Atmos.*, 118, 11,107-11,118, <https://doi.org/10.1002/jgrd.50836>, 2013.

993 Shaylor, M., Brindley, H., and Sellar, A.: An Evaluation of Two Decades of Aerosol Optical Depth Retrievals  
994 from MODIS over Australia, *Remote Sens.*, 14, 2664, <https://doi.org/10.3390/rs14112664>, 2022.

995 She, L., Xue, Y., Guang, J., Che, Y., Fan, C., Li, Y., and Xie, Y.: Towards a comprehensive view of dust events  
996 from multiple satellite and ground measurements: Exemplified by the May 2017 East Asian dust storm, *Nat.*  
997 *Hazards Earth Syst. Sci.*, 18, 3187–3201, <https://doi.org/10.5194/nhess-18-3187-2018>, 2018.

998 Song, Q., Zhang, Z., Yu, H., Ginoux, P., and Shen, J.: Global dust optical depth climatology derived from CALIOP  
999 and MODIS aerosol retrievals on decadal timescales: regional and interannual variability, *Atmos. Chem. Phys.*,  
1000 21, 13369–13395, <https://doi.org/10.5194/acp-21-13369-2021>, 2021.

1001 Soni, A., Mandariya, A. K., Rajeev, P., Izhar, S., Singh, G. K., Choudhary, V., Qadri, A. M., Gupta, A. D., Singh,  
1002 A. K., and Gupta, T.: Multiple site ground-based evaluation of carbonaceous aerosol mass concentrations retrieved  
1003 from CAMS and MERRA-2 over the Indo-Gangetic Plain, *Environ. Sci. Atmos.*, 1, 577–590,  
1004 <https://doi.org/10.1039/d1ea00067e>, 2021.

1005 Speer, M. S.: Dust storm frequency and impact over Eastern Australia determined by state of Pacific climate

1006 system, *Weather Clim. Extrem.*, 2, 16–21, <https://doi.org/10.1016/j.wace.2013.10.004>, 2013.

1007 Sprigg, R. C.: Alternating wind cycles of the Quaternary era and their influences on aeolian sedimentation in and  
1008 around the dune deserts of south-eastern Australia, in: *Quaternary Dust Mantles of China, New Zealand and*  
1009 *Australia*, 1982a.

1010 Sprigg, R. C.: Some stratigraphic consequences of fluctuating Quaternary sea levels and related wind regimes in  
1011 southern and central Australia, in: *Alternating wind cycles of the Quaternary era and their influences on aeolian*  
1012 *sedimentation in and around the dune deserts of south eastern Australia*, Wasson R, Australian National University,  
1013 Canberra, 211–240, 1982b.

1014 Stefanski, R. and Sivakumar, M. V. K.: Impacts of sand and dust storms on agriculture and potential agricultural  
1015 applications of a SDSWS, *IOP Conf. Ser. Earth Environ. Sci.*, 7, 012016, <https://doi.org/10.1088/1755->  
1016 [1307/7/1/012016](https://doi.org/10.1088/1755-1307/7/1/012016), 2009.

1017 Strong, C. L., Parsons, K., McTainsh, G. H., and Sheehan, A.: Dust transporting wind systems in the lower Lake  
1018 Eyre Basin, Australia: A preliminary study, *Aeolian Res.*, 2, 205–214,  
1019 <https://doi.org/10.1016/j.aeolia.2010.11.001>, 2011.

1020 Sun, E., Xu, X., Che, H., Tang, Z., Gui, K., An, L., Lu, C., and Shi, G.: Variation in MERRA-2 aerosol optical  
1021 depth and absorption aerosol optical depth over China from 1980 to 2017, *J. Atmos. Solar-Terrestrial Phys.*, 186,  
1022 8–19, <https://doi.org/10.1016/j.jastp.2019.01.019>, 2019a.

1023 Sun, E., Che, H., Xu, X., Wang, Z., Lu, C., Gui, K., Zhao, H., Zheng, Y., Wang, Y., Wang, H., Sun, T., Liang, Y.,  
1024 Li, X., Sheng, Z., An, L., Zhang, X., and Shi, G.: Variation in MERRA-2 aerosol optical depth over the Yangtze  
1025 River Delta from 1980 to 2016, *Theor. Appl. Climatol.*, 136, 363–375, <https://doi.org/10.1007/s00704-018-2490->  
1026 [9](https://doi.org/10.1007/s00704-018-2490-9), 2019b.

1027 Sundström, A. M., Kolmonen, P., Sogacheva, L., and de Leeuw, G.: Aerosol retrievals over China with the  
1028 AATSR dual view algorithm, *Remote Sens. Environ.*, 116, 189–198, <https://doi.org/10.1016/j.rse.2011.04.041>,  
1029 2012.

1030 Tanaka, T. Y. and Chiba, M.: A numerical study of the contributions of dust source regions to the global dust  
1031 budget, *Glob. Planet. Change*, 52, 88–104, <https://doi.org/10.1016/j.gloplacha.2006.02.002>, 2006.

1032 Tews, K.: *Wind erosion rates from meteorological records in eastern Australia 1960-92*, Griffith University, 1996.

1033 Thomas, G. E., Carboni, E., Sayer, A. M., Poulsen, C. A., Siddans, R., and Grainger, R. G.: Oxford-RAL Aerosol  
1034 and Cloud (ORAC): aerosol retrievals from satellite radiometers, *Satell. Aerosol Remote Sens. over L.*, 193–225,

1035 [https://doi.org/10.1007/978-3-540-69397-0\\_7](https://doi.org/10.1007/978-3-540-69397-0_7), 2009.

1036 Torre, G., Gaiero, D., Coppo, R., Cosentino, N. J., Goldstein, S. L., De Vleeschouwer, F., Roux, G. Le, Bolge, L.,  
1037 Kiro, Y., and Sawakuchi, A. O.: Unraveling late Quaternary atmospheric circulation in the Southern Hemisphere  
1038 through the provenance of Pampean loess, *Earth-Science Rev.*, 232, 104143,  
1039 <https://doi.org/10.1016/j.earscirev.2022.104143>, 2022.

1040 Tozer, P. and Leys, J.: Dust storms - What do they really cost?, *Rangel. J.*, 35, 131–142,  
1041 <https://doi.org/10.1071/RJ12085>, 2013.

1042 Voss, K. K. and Evan, A. T.: A new satellite-based global climatology of dust aerosol optical depth, *J. Appl.*  
1043 *Meteorol. Climatol.*, 59, 83–102, <https://doi.org/10.1175/JAMC-D-19-0194.1>, 2020.

1044 Wei, J., Li, Z., Peng, Y., and Sun, L.: MODIS Collection 6.1 aerosol optical depth products over land and ocean:  
1045 validation and comparison, *Atmos. Environ.*, 201, 428–440, <https://doi.org/10.1016/j.atmosenv.2018.12.004>,  
1046 2019.

1047 Wu, C., Lin, Z., and Liu, X.: The global dust cycle and uncertainty in CMIP5 (Coupled Model Intercomparison  
1048 Project phase 5) models, *Atmos. Chem. Phys.*, 20, 10401–10425, <https://doi.org/10.5194/acp-20-10401-2020>,  
1049 2020.

1050 Yang, L., She, L., Che, Y., He, X., Yang, C., and Feng, Z.: Analysis of Dust Detection Algorithms Based on FY-  
1051 4A Satellite Data, *Appl. Sci.*, 13, 1365, <https://doi.org/10.3390/app13031365>, 2023.

1052 Yang, X., Zhao, C., Yang, Y., and Fan, H.: Long-term multi-source data analysis about the characteristics of  
1053 aerosol optical properties and types over Australia, *Atmos. Chem. Phys.*, 21, 3803–3825,  
1054 <https://doi.org/10.5194/acp-21-3803-2021>, 2021.

1055 Yu, B., Neil, D. T., and Hesse, P. P.: Correlation between rainfall and dust occurrence at mildura, Australia: The  
1056 difference between local and source area rainfalls, *Earth Surf. Process. Landforms*, 17, 723–727,  
1057 <https://doi.org/10.1002/esp.3290170708>, 1992.

1058 Yu, B., Hesse, P. P., and Neil, D. T.: The relationship between antecedent regional rainfall conditions and the  
1059 occurrence of dust events at Mildura, Australia, *J. Arid Environ.*, 24, 109–124,  
1060 <https://doi.org/10.1006/jare.1993.1010>, 1993.

1061 Yu, Y. and Ginoux, P.: Assessing the contribution of the ENSO and MJO to Australian dust activity based on  
1062 satellite- And ground-based observations, *Atmos. Chem. Phys.*, 21, 8511–8530, [https://doi.org/10.5194/acp-21-](https://doi.org/10.5194/acp-21-8511-2021)  
1063 8511-2021, 2021.



1064 Zender, C. S.: Mineral Dust Entrainment and Deposition (DEAD) model: Description and 1990s dust climatology,  
1065 J. Geophys. Res., 108, 4416, <https://doi.org/10.1029/2002JD002775>, 2003.

1066 Zhang, X., Zhao, L., Tong, D. Q., Wu, G., Dan, M., and Teng, B.: A systematic review of global desert dust and  
1067 associated human health effects, Atmosphere (Basel)., 7, <https://doi.org/10.3390/atmos7120158>, 2016.

1068 Zhao, Q., Zhao, W., Bi, J., and Ma, Z.: Climatology and calibration of MERRA-2 PM2.5 components over China,  
1069 Atmos. Pollut. Res., 12, 357–366, <https://doi.org/10.1016/j.apr.2020.11.016>, 2021.

1070