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Intra- and inter-annual changes in isoprene emission from central Amazonia

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Eliane Gomes Alves^{1,2*}, Raoni Aquino Santana³, Cléo Quaresma Dias-Junior^{4,2}, Santiago

3 Botía⁵, Tyeen Taylor⁶, Ana Maria Yáñez-Serrano^{7,8,9}, Jürgen Kesselmeier¹⁰, Efstratios 4

Bourtsoukidis¹¹, Jonathan Williams¹², Pedro Ivo Lembo Silveira de Assis¹³, Giordane 5

Martins¹³, Rodrigo de Souza¹⁴, Sergio Duvoisin Junior¹⁵, Alex Guenther¹⁶, Dasa Gu¹⁷, 6

- Anywhere Tsokankunku¹², Matthias Sörgel¹², Bruce Nelson¹⁸, Davieliton Pinto¹³, Shujiro 7
- Komiya¹, Diogo Martins Rosa¹³, Bettina Weber^{19,10}, Cybelli Barbosa^{10,19}, Michelle Robin¹, 8
- Kenneth J Feeley²⁰, Alvaro Duque²¹, Viviana Londoño Lemos²², Maria Paula Contreras²³, 9
- Alvaro Idarraga²⁴, Norberto López A.²⁴, Chad Husby²⁵, Brett Jestrow²⁵, Iván Mauricio 10
- Cely Toro⁴. 11
- 12
- 13 ¹Department of Biogeochemical Processes, Max Planck Institute for Biogeochemistry, Jena, Germany
- 14 ²Climate and Environment Department, National Institute of Amazonian Research, Manaus, Brazil
- 15 ³ Department of Atmospheric Sciences, Federal University of Western Para, Santarem, Brazil
- 16 ⁴Federal Institute of Para, Belem, Brazil
- 17 ⁵ Department of Biogeochemical Signals, Max Planck Institute for Biogeochemistry, Jena, Germany
- 18 ⁶ Department of Civil & Environmental Engineering, University of Michigan, USA.
- 19 ⁷ 7IDAEA-CSIC, 08034, Barcelona, Spain
- 20 ⁸CREAF, E08193 Bellaterra (Cerdanyola del Vallès), Catalonia, Spain
- ⁹Global Ecology Unit, CREAF-CSIC-UAB, E08193 Bellaterra (Cerdanyola del Vallès), Catalonia, Spain
- ¹⁰ Multiphase Chemistry Department, Max Planck Institute for Chemistry, Mainz, Germany
- ¹¹The Cyprus Institute, Nicosia, Cyprus
- ¹²Atmospheric Chemistry Department, Max Planck Institute for Chemistry, Mainz, Germany
- 21 22 23 24 25 26 27 28 29 30 31 ¹³ Department of Tropical Forest Sciences, National Institute for Amazonian Research, Manaus, Brazil
- ¹⁴ Meteorology Department, State University of Amazonas, Manaus, Brazil
- ¹⁵ Chemistry Department, State University of Amazonas, Manaus, Brazil
- ¹⁶Department of Earth System Science, University of California, Irvine, U.S.A.
- ¹⁷ Division of Environment and Sustainability, Hong Kong University of Science and Technology, Clear
- Water Bay, Hong Kong, China
- ¹⁸Coordination of Environmental Dynamics, National Institute of Amazonian Research, Manaus, Brazil
- ¹⁹ Institute for Biology, Division of Plant Sciences, University of Graz, Graz, Austria
- ²⁰Department of Biological Sciences, University of Miami, Coral Gables, FL, USA.
- 32 33 34 ²¹Departamento de Ciencias Forestales, Universidad Nacional de Colombia-Sede 19Medellín, Medellín,
- 35 36 Colombia
- ²²Department of Plant and Microbial Biology, University of Minnesota, USA
- 37 ²³Jardín Botánico de Cartagena "Guillermo Piñeres", Turbaco, Bolívar, Colombia.
- 38 ²⁴Fundación Jardín Botánico de Medellín, Antioquia, Colombia.
- 39 ²⁵Fairchild Tropical Botanic Garden, Miami, FL, USA
- 40 *egomes@bgc-jena.mpg.de
- 41

42 Abstract

43 Isoprene emissions are a key component in biosphere-atmosphere interactions, and the

44 most significant global source is the Amazon rainforest. However, intra- and inter-annual

- 45 variations in biological and environmental factors that regulate isoprene emission from
- 46 Amazonia are not well understood and, thereby, poorly represented in models. Here, with
- 47 datasets covering several years of measurements at the Amazon Tall Tower Observatory
- 48 (ATTO), in central Amazonia, Brazil, we (1) quantified canopy profiles of isoprene mixing
- ratios across seasons of normal and anomalous years and related them to the main drivers 49

of isoprene emission - solar radiation, temperature, and leaf phenology; (2) evaluated the 50 51 effect of leaf age on the magnitude of the isoprene emission factor (E_s) from different tree 52 species and scaled up to canopy with intra- and inter-annual leaf age distribution derived 53 by a phenocam; and (3) adapted the leaf age algorithm from MEGAN with observed 54 changes in E_s across leaf ages. Our results showed that the variability in isoprene mixing 55 ratios was higher between seasons (max. during the dry-to-wet transition seasons) than 56 between years, with values from the extreme 2015 El-niño year not significantly higher 57 than in normal years. In addition, model runs considering in-situ observations of canopy E_s 58 and the modification on the leaf age algorithm with leaf-level observations of E_s presented 59 considerable improvements in the simulated isoprene flux. This shows that MEGAN 60 estimates of isoprene emission can be improved when biological processes are 61 mechanistically incorporated into the model.

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1. Introduction

65 Isoprene dominates the emission of biogenic volatile organic compounds (BVOCs) into 66 the atmosphere, and its major global source is tropical vegetation (Guenther et al., 2012; 67 Sindelarova et al., 2014). In the atmosphere, isoprene is a short-lived (minutes to hours) 68 reactive BVOC species, and its photooxidation affects the atmospheric oxidation capacity 69 contributing to the formation of ozone (O_3) and secondary organic aerosols (SOA) 70 (Atkinson, 1997; Pöschl et al., 2010). With its high plant foliage biomass and rich plant 71 diversity (ter Steege et al., 2013), the Amazon Forest represents a key source of isoprene 72 to the atmosphere (Yáñez-Serrano et al., 2020). However, model estimates of isoprene 73 emission and its intra- and inter-annual variability in the Amazon still carry high 74 uncertainty, because only a few observational experiments have been conducted with 75 mechanistic and process-based approaches, which hinders further modeling optimization 76 (Alves et al., 2018; Yáñez-Serrano et al., 2020). One of the most critical knowledge gaps 77 is how plants' isoprene emission differs under extremely hot and dry conditions, such as in 78 El-niño years, and how this might affect atmospheric processes. As some studies have 79 indicated that extreme years will become more frequent and intense with climate change 80 (Nobre et al., 2016; Boulton et al., 2022), it is essential to understand the processes 81 mediated by isoprene in such years to improve model estimates (Yáñez-Serrano et al., 82 2020; Artaxo et al., 2022).

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84 Some reasons for uncertainties in isoprene model estimates are already known. The correct 85 determination of the magnitude of the isoprene source - or the emission factor at leaf standard conditions (1000 μ mol m⁻² s⁻¹ photosynthetically active radiation- PAR, 30 °C), 86 87 as it is conceptualized in models (e.g., Guenther et al., 1995) - is crucial to improve isoprene 88 modeling estimates. The Amazon plant biodiversity represents a considerable challenge for 89 determining the isoprene emission factor. Although previous studies suggested that $\sim 1\%$ 90 of tree species are hyperdominant - with their tree individuals responsible for half of all 91 tree stems, carbon storage, and productivity (ter Steege et al., 2013; Fauset et al., 2015) -, 92 it is still unclear which plant species can emit substantial amounts of isoprene (Monson et 93 al., 2013), how these isoprene emitters are distributed throughout the Amazon basin, and 94 how the isoprene emission factor varies seasonally and interannually as result of changes 95 in eco-physiological processes (Gomes Alves et al., 2022). Another source of uncertainty

is related to quantifying the main sinks of isoprene. Once emitted by plant foliage, isoprene 96 97 can undergo surface deposition onto plant canopy (Karl et al., 2004) and soil (Pegoraro et 98 al., 2006), can be oxidized at rates depending on the atmospheric concentration of other 99 gases such as NO_x, O₃ and OH (Atkinson, 1997), and can be transported into and out of the atmospheric boundary layer (Wei et al., 2018). Additionally, the rapid conversion of 100 101 isoprene photooxidation products can open a further sink for BVOCs in plants. This 102 chemical and biological processing of emitted compounds may affect vertical transport 103 processes, again influencing the biosphere (Kesselmeier et al., 2002; Canaval et al., 2020).

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105 In addition, seasonal variation in isoprene emission from Amazon forests has been reported 106 by several in-situ studies, with the indication that isoprene seasonality is driven by intra-107 annual variation in solar radiation, temperature, and leaf phenology (Kuhn et al., 2004a, b; 108 Yáñez-Serrano et al., 2015; Alves et al., 2016, 2018; Wei et al., 2018; Langford et al., 109 2022). On a larger scale, satellite retrievals of isoprene oxidation products, like formaldehyde (Barkley et al., 2009; Bauwens et al., 2016), and direct retrieval of isoprene 110 111 (Fu et al., 2019; Wells et al., 2022) have given an initial view of the long-term Amazon isoprene emission, enabling not only seasonal but also inter-annual comparisons. Yet, there 112 remains a need to parameterize and evaluate the estimations with local and regional 113 114 measurements and to gain a better understanding of the main processes related to sources 115 and sinks of isoprene, since some studies have shown that satellite-derived isoprene 116 emission values are either overestimated (Alves et al., 2016) or underestimated (Gu et al., 117 2017), or even show maximum emissions in a different season when compared to in-situ 118 measurements (Alves et al., 2016, 2018).

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120 Here we report in-situ observations of isoprene mixing ratios during different seasons and 121 in consecutive years in central Amazonia to evaluate intra- and inter-annual variabilities in 122 two normal years (2013-2014) and one El-niño year (2015); in addition, we report 123 observations of leaf-level isoprene emission factor and leaf phenology monitoring. With 124 the intra- and inter-annual observations of isoprene at a central Amazonian site, this study 125 proposes to: (1) quantify the isoprene mixing ratios across seasons of normal and 126 anomalous years and compare them with the main drivers of isoprene emission - solar 127 radiation, temperature, and leaf phenology; (2) evaluate the effect of leaf age on the 128 magnitude of the isoprene emission factor from different tree species and scale up with 129 canopy intra- and inter-annual leaf age distribution; and (3) use the Model of Emissions of 130 Gases and Aerosols from Nature (MEGAN) to assess the effects of the observed changes 131 in the isoprene emission factor across leaf ages, by modifying the leaf age algorithm and 132 comparing simulations with observations at canopy-level.

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2. Methods

136 2.1 Amazon Tall Tower Observatory (ATTO)

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We performed measurements at the ATTO site located 150 km northeast of Manaus in the
 Uatumã Sustainable Development Reserve (USDR) in central Amazonia. The climate is

140 tropical humid, with two distinctive seasons – wet season (December-May) and dry season

141 (July-October) and transition seasons in between – and has a mean annual precipitation of 142 2380 mm (TRMM climatological average – 1998-2019; please see more details in Botía et 143 al., 2022). The vegetation in this area is considered mature, mostly non-flooded rainforest 144 (terra-firme), with a mean canopy height of 35 m, and predominantly occurs on plateaus at 145 a maximum altitude of approximately 130 m a.s.l. (Andreae et al., 2015). Air masses 146 arriving at the site predominantly come from the east (NE~20%, ENE~27%, E~33%, 147 ESE~19%) (Zannoni et al., 2020) and have passed through 1500 km of undisturbed terra-148 firme rainforest, with minor intrusion of air masses from Manaus (Pöhlker et al., 149 2019). Figure 1 shows seasonal variation in solar radiation, air temperature, precipitation, 150 and soil moisture from 2013 to 2019. Andreae et al. (2015) have more details on this 151 experimental site.

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153 2.2 Mixing ratios of isoprene – canopy level

Isoprene gradient mixing ratios were inferred by air samples collected from the INSTANT 155 156 tower (80 m height, coordinates: S 02°08.7520' W 58°59.9920') at eight heights in and above the canopy (0.05, 0.5, 4, 12, 24, 38, 53 and 79 m) during intensive campaigns across 157 different seasons from November 2012 to October 2015. Eight heated (50 C) and insulated 158 159 inlets (fluorinated ethylene propylene - FEP, OD ³/₈ in.) were connected to a quadrupole 160 Proton Transfer Reaction - Mass Spectrometer (PTRMS) (Ionicon Analytic GmbH, Austria) - using the primary ion H3O+ and operated under standard conditions (2.2 mbar 161 162 drift pressure, 600V drift voltage, 127 Td), which was housed in an air-conditioned 163 container 10 m from the INSTANT tower. The inlets were guided to a valve system, switching every 2 min between the different heights, completing a full profile in 16 min. 164 While an inlet was not sampled, it was flushed by a bypass pump at a flow rate of 16 lpm. 165 Humidity-dependent calibrations (using bubbled synthetic zero air to dilute the standard, 166 regulated as close as possible to ambient humidity conditions) were performed using a gas 167 168 cylinder containing isoprene (m/z 69). The dilution steps ranged from 22 to 0.8 ppb. To 169 determine the background signal for isoprene, a catalytic converter (Supelco, Inc. with 170 platinum pellets heated to >400 \circ C) was used to convert ambient VOC to CO₂ +H₂O. The 171 background signal was measured once every hour and then interpolated over the time of 172 the measurements. The detection limit (LOD) for isoprene varied between 0.09 (wet 173 season) and 0.1 (dry season) ppb. The mean total uncertainty of isoprene mixing ratios was 174 9.9 %, within the PTRMS measurement uncertainty ($\sim 10\%$). For more details on the 175 experimental setup, PTRMS conditions, and calibration, we refer the reader to Yãnez-176 Serrano et al. (2015)

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- 179 2.3 Flux of isoprene canopy level
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During a campaign in November 2015, eddy covariance fluxes of isoprene were measured for 11 days. Isoprene concentrations were obtained with the above-described PTRMS at a time resolution of 1 s and from a separate 3/8" inlet at 41 m height that sampled air at a flow rate of about 10 1 min⁻¹. A CSAT3 sonic anemometer (Campbell Scientific Inc., Logan, U.S.A.) measured the three-dimensional wind speed at high frequency (1 Hz) and

186 was placed at a distance of 0.5 m from the isoprene inlet. Fluxes were then calculated by

187 correlating fluctuations of the vertical wind vector to the fluctuations of isoprene 188 concentrations with the software package EddyPro® (LI-COR Inc., Lincoln, U.S.A.). A 189 method for despiking and raw data statistical screening was employed (Vickers and Mahrt, 190 1997). Half-hourly averaged fluxes were flagged according to a method of data quality 191 control (Mauder and Foken, 2004), and only data with the highest quality (flags 0 and 1) 192 was used for further analyses. Losses for sampling frequencies between 0.1 and 0.8Hz have 193 been observed as below 10% (Guenther and Hills, 1998; Spirig et al., 2005; Holst et al., 194 2010; Jensen et al., 2018). Footprints were calculated using a two-dimensional model for 195 a geographic domain of 2 x 2 km centered at the INSTANT tower (Kljun et al., 2015). The 196 Tovi Footprint Analysis Toolbox (LI-COR Inc., Lincoln, U.S.A.) was used to calculate 197 half-hourly footprints and to combine them for the measurement period. Mean daytime 198 uncertainties of eddy covariance isoprene flux were at most 15%. More details on the flux 199 measurements and data processing are given in Pfannerstill et al. (2018).

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201 2.4 Leaf Area Density – measurements with the Light Detection and Ranging sensor
 202 (LiDAR)

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204 Measurements of canopy leaf area density were carried out with a ground Light Detection 205 and Ranging sensor (LiDAR) at the ATTO site. These measurements aimed to give 206 information on the canopy structure around the INSTANT tower. Ground-LiDAR surveys 207 were conducted in October 2015 with a Riegl LD90-3100VHS-FLP system (Horn, 208 Austria), which generated a canopy profile map in vertical and horizontal directions. We 209 walked ten transects of 150 m in length with the ground-LIDAR system. The transects were 210 parallelly distributed at a distance of ~ 100 m from each other, with six transects to the 211 east/northeast, three transects to the west, and one transect to the south of the INSTANT 212 tower. Measurements were averaged every 15 m of each transect, summing up to ten 213 measurements per transect. Measurements of all ten transects were then averaged and 214 presented with the confidence interval (95%). More details about how the ground LiDAR 215 data were analyzed can be obtained from Stark et al. (2012).

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217 2.5 Leaf-level monitoring of leaf demography and phenology218

219 Leaf demography and phenology of 36 trees were monitored from March 2016 to 220 December 2017. Along 100 m of canopy walkways, canopy leaves were monitored 221 monthly to determine leaf ages and investigate how leaf age proportions vary during the 222 year. Ten branches of each tree were randomly selected and labeled with one iron ring at 223 their bottom end. All leaves attached from the bottom to the apical end were counted and 224 dated according to the day of observation. For the first observation, all leaves were assigned 225 with unknown age. In the following months, every time a new leaf was observed, the date 226 of observation was recorded for that specific leaf. For leaf age determination, the date of 227 the first observation of a new leaf was set back to 15 days before observation. The age was 228 calculated by the difference, in the number of days, between the first day and the last day 229 of observation, resulting in a number of days with a deviation of plus-minus 15 days. For instance, if a new leaf was observed on 1st July 2017, the flushing date of this leaf was 230 231 assigned for 17th June 2017 (+/- 15 days). Then, all subsequent measurements considered

17th June 2017 as a date for leaf flushing, and aging was counted based on the number of 232 233 days that this leaf stayed attached to the branch.

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235 2.6 Isoprene emission factor – leaf level

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237 Leaves of 21 canopy tree species, out of the 36 trees monitored for leaf demography and 238 phenology (described in section 2.5), were measured to determine the isoprene emission factor across different leaf ages (Table S1) from October to November 2017. The other 15 239 240 trees were unreachable with the sampling system and, therefore, not measured. Leaf-level 241 isoprene sampling was carried out in 2-3 leaves of each age class available for each tree 242 during the measurement period, using a commercial portable gas exchange system GFS-243 3000 (Walz, Effeithich, Germany). Each leaf was separately enclosed in the leaf chamber 244 at standard conditions – photosynthetic photon flux density (PPFD) set to 1000 µmol m⁻² 245 s^{-1} and leaf temperature to 30°C - until net assimilation, stomatal conductance and internal 246 CO₂ concentration were stable. The measurement stability criterion was assigned as one 247 standard deviation of the net assimilation mean. The airflow rate going into the leaf 248 chamber was 400 µmol s⁻¹ and CO₂ and H₂O concentrations were 400 µmol.mol⁻¹ and 21 249 mmol.mol⁻¹ (relative humidity of $\sim 60\%$), respectively. Air exiting the GFS-3000 leaf 250 chamber was routed to fill sorbent cartridges (stainless steel tubes filled with Tenax TA 251 and Carbograph 5 TD sorbents), and a downstream pump sampled the exiting air at a rate 252 of 200 sccm for 10 min. A hydrocarbon filter (Restek Pure Chromatography, Restek 253 Corporations, USA) was installed at the air inlet of GFS-3000 to remove isoprene from the 254 incoming ambient air, and all tubing in contact with the sampling air was made of PTFE. 255 Before each measurement, a blank sample was obtained from the empty leaf chamber.

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257 Isoprene content in the sorbent cartridges was determined in the laboratory at the 258 University of California (Irvine, U.S.A.). All cartridges were placed into a thermally 259 desorbing autosampler (TD-100, Markes International, Inc). The isoprene was pre-260 concentrated at 10 °C followed by injection into a gas chromatograph (GC, model 7890B, 261 Agilent Technologies, Inc) equipped with a time-of-flight mass spectrometer (Markes 262 BenchTOF-SeV) and a flame ionization detector (TD-GC-FID/TOF-MS) (Woolfenden 263 and McClenny, 1999; ASTM D6196-15, 2015). Internal standards tetramethylethylene 264 and decahydronaphtalene were injected into each sample after collection and before 265 analysis. The system was calibrated daily with a commercial isoprene standard from Apel 266 Riemer Environmental Inc. The external gas standard was prepared using a dynamic 267 dilution system, and the effluent was added to sorbent cartridges under conditions similar 268 to those used for sampling. Once the volume mixing ratio of isoprene (ppbv) was obtained, 269 leaf emission flux was determined using the Eq. (1):

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$$F = R_{ppbv} \times \frac{Q}{A} \tag{1}$$

where F (nmol.m⁻².s⁻¹) is leaf flux of isoprene emission; R_{ppby} (nmol.mol⁻¹) is isoprene 273 274 concentration of the sample (cartridge); *Q* is the flow rate of air into the leaf chamber (400 275 μ mol.s⁻¹); and A is the area of leaf within the chamber (0.08 m²). The isoprene emission rate was then calculated and converted to mg.m⁻².h⁻¹. For more details on tree species, leaf 276 277 age, and assigned leaf age class, see Table S1 in Supplementary Information.

324 2.7 Tower-camera derived leaf phenology and demography data

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326 Upper canopy leaf phenology was monitored with a Stardot RGB camera (model Netcam 327 XL 3MP) installed at 81m height on the ATTO INSTANT tower. For more details on the 328 camera setup, radiometric calibration, and detection of phenological stages, we refer the 329 reader to Lopes et al. (2016). Only images acquired near noon and under an overcast sky 330 (diffuse illumination) were selected for subsequent analysis. The camera (subsequently 331 called phenocam) monitored the upper crown surfaces of 194 trees from July 2013 to 332 November 2018. Images were analyzed to track the temporal trajectory of each tree crown 333 and assign them into one of three classes: "leaf flushing" (crowns that showed a strong 334 increase in greening), "leaf abscising" (crowns which showed a large increase in greying, 335 which is the color of bare upper canopy branches) or "no change". By counting the number 336 of individual trees per month for each category (flushing or abscission), we aggregated our 337 census to the monthly scale. Of the monitored trees, 69% (n = 134) had clear flushing and 338 abscission patterns, and, using the number of days after each flushing event, we determined leaf age classes and attributed a fraction of the upper canopy crowns to an age class at 339 340 monthly intervals. We defined the following leaf age classes: (i) young leaves (0-1 341 month), (ii) growing (1–2 months), (iii) mature leaves (3–6 months), and (iv) old leaves 342 (>6 months). Then, we partitioned the age classes into classes of leaf area index (LAI) 343 (i.e., young, growing, mature, and old LAI) by normalizing each leaf age class with the total LAI measured at ATTO. A constant LAI of 5.32 m² m⁻² was used for all months, 344 345 since the variability of this number throughout the year was not statistically significant 346 (unpublished results). For the normalization, we considered the total number of trees in 347 the camera frame (n = 194), assuming that the 31% that do not have clear flushing patterns 348 are part of the old age class. For more details on the methods and assumptions for 349 separating LAI into leaf age classes, see Wu et al. (2016). Datasets of flushing and 350 abscission (http://doi.org/10.17871/ atto.223.7.840) and the raw LAI age classes (http://doi. 351 org/10.17871/ atto.230.4.842).

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53 2.8 Isoprene emission trait – tree species level

355 To get more detailed information on the trees monitored with the camera, a total of 194 356 trees were taxonomically identified, and the isoprene emission trait was assigned. Isoprene 357 emission data were obtained from published data and new measurements for the study 358 species. New measurements were conducted at the ATTO research site (described in 359 section 2.6), and additional measurements were obtained using the PORCO method (Taylor 360 et al., 2021), a customized photoionization detection system, on trees in tropical botanical 361 gardens. Briefly, all PORCO measurements were made in situ on uncut 'sun' branches by 362 enclosing one-to-few leaves inside rigid leaf cuvettes, acclimating them to darkness, and 363 then exposing the leaves to photosynthetically active radiation controlled at 1000 µmol m⁻ 364 2 s⁻¹, and temperatures near 30°C, for 3.5 minutes of measurement time. Emission rates were corrected to a 30°C equivalent based on a standard temperature response curve 365 (Guenther et al., 1993). Emission rates exceeding 1 nmol $m^{-2} s^{-1}$ were considered positively 366 367 indicative of isoprene emissions. See the full method validation and a discussion of the 368 rarity of detection of other compounds as false positives for isoprene in Taylor et al. (2021). 369 Botanic gardens used for tree measurements were: A. Duque private collection, Retiro, 331

332 For applying isoprene measurements from external datasets (botanic garden measurements 333 or published literature) to our study species, we followed the methods of Taylor et al., 334 (2018, 2019). We used data compiled from 12 literature sources (Bracho-Nunez et al., 335 2013; Geron et al., 2002; Harley et al., 2004; Keller & Lerdau, 1999; Klinger et al., 1998; 336 Klinger et al., 2002; Lerdau & Keller 1997; Padhy & Varshney, 2005; Tambunan et al., 337 2006; Taylor et al., 2018; Taylor et al., 2021; Varshney & Singh, 2003). Tree species 338 taxonomy was standardized by the Taxonomic Name Resolution Service (Boyle et al., 339 2013; Boyle et al., 2021). We assigned species data only in terms of the genetically 340 determined capacity to produce isoprene (Monson et al., 2013); we did not consider the 341 variability in the strength of emissions, for which data are more limited and potentially 342 confounded by method variation and species plasticity. A species-level emission statusemitter or non-emitter-was applied where available in external datasets; otherwise, genus-343 344 level information was used to impute the emission status to unmeasured species. The 345 proportion of measured species in a genus that emit isoprene was used as an estimate of 346 the probability (pIE) that any species sampled from the genus would be an emitter. For a 347 genus corresponding to one of our study species, for $pIE \le 1/3$, the species was estimated 348 to be a non-emitter, and for pIE $\geq 2/3$, the species was estimated to be an emitter. For values 349 1/3 < pIE > 2/3, the genus average was considered ambiguous and the species was excluded 350 from the analyses. Whereas there is some expected error in the assignment of emission 351 status to any given species, analyses of large numbers of species will tend toward the 352 correct answer due to the tendency of genera to predominate in emitting or non-emitting 353 species (Taylor et al., 2018). All species for which no emission data were available at the 354 genus level were excluded from the analyses. The imputed isoprene emission status and 355 associated information for each of our study species can be found in Table S2. The source 356 data (literature reference or present study metadata) for each species that informed the 357 imputation can be found in Table S3.

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359 2.9 Modeled isoprene flux estimates - Model of Emissions of Gases and Aerosols from 360 Nature (MEGAN)

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Isoprene fluxes were simulated using the MEGAN version 2.1 model in which the flux activity factor for isoprene (γ i) is proportional to the emission response to light (γ P), temperature (γ T), leaf age (γ A), soil moisture (γ SM), leaf area index (LAI), and CO₂ inhibition (γ CO₂) according to Eq. (2) (Guenther et al., 2012):

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$$\gamma i = C_{CE} LAI \gamma P \gamma T \gamma A \gamma SM \gamma CO_2$$
⁽²⁾

For this study, the canopy environment model of Guenther et al. (2006) was used with a canopy environment coefficient (C_{CE}) of 0.57. MEGAN was run accounting for variations in light, temperature, and LAI separated into leaf age classes. CO₂ inhibition and soil moisture activity factors were set equal to a constant of 1, assuming these parameters do not vary. For all simulations, we assumed no seasonal variation in soil moisture because the soil moisture observed in this site consistently exceeds the threshold for the isoprene drought response in MEGAN 2.1 (Guenther et al., 2012), which means that MEGAN would predict no variation in isoprene emission resulting from the observed changes in soil moisture (Fig. 1).

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379 Solar radiation (PPFD) and air temperature inputs for all model simulations were obtained 380 from measurements at the INSTANT tower. Air temperature at 36 m height above ground 381 level was measured with a temperature and relative humidity sensor (CS215-L, Campbell Scientific 382 Inc., Logan, Utah, USA). In cases where the air temperature measurement at 36 m height failed, 383 the missing data were gap-filled with air temperature data available at other heights (73 m, 55 m, 384 40 m, 12 m), measured with CS215-L sensors installed on the INSTANT tower, or with the air 385 temperature at 18 m above the ground measured with a thermocouple (Conatex, St. Wendel, 386 Germany), installed along one evergreen tree of the species *Buchenavia parvifolia* (*Combretaceae*), 387 located 95 m away from the INSTANT tower. In cases where all the air temperature sensors failed 388 for less than 4 hours, the missing air temperature at 36 m height was gap-filled by linear 389 interpolation, visually checking data quality. In cases where no air temperature measurement was 390 available for a long time (e.g., one day, 2 months etc.), confirmed several times in 2013, the missing 391 air temperature at 36 m height was gap-filled by a multiple regression model developed with three 392 predictor variables: half-hourly variation of the soil temperature at 10 cm depth, soil heat flux, and 393 volumetric soil water content at 40 cm depth. The model training period was from 2013 June to 394 2014 May because the three predictor variables were usually available through the one-year period. 395 The developed model was validated based on the observation dataset from June 2014 to May 2015, 396 which showed good agreement with observed air temperature data at 36 m height during the 397 validation period ($R^2 = 0.83$; RMSE = 1.21; n = 7473). The developed and validated model was applied to the three predictor variables measured in 2013 for gap-filling the long-term missing data 398 399 of air temperature at 36 m height. In cases where the predictor variables were unavailable in 2013, 400 the missing data were gap-filled using Akima interpolation with visual data quality checks. 401 Incoming and outgoing shortwave radiation was measured with a net radiometer (NR- Lite2, Kipp 402 & Zonen, the Netherlands) at 75 m above ground. In cases where the radiation measurement failed 403 for no more than 1 hour, the missing radiation data were gap-filled by linear interpolation, visually 404 checking data quality. In cases where radiation data were unavailable for more than 1 hour, the 405 missing data were gap-filled by the mean diurnal course (over ± 15 -day) method. Lastly, we used 406 a constant value (5.32) for the LAI and normalized it with monthly leaf age fractions 407 obtained from the phenocam observations to derive the canopy leaf age for each month 408 (see section 2.6). More details on model settings are found in Guenther et al. (2012).

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3. Results and Discussion

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- 413 *3.1 Observations of canopy isoprene mixing ratios*
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We observed intra- and inter-annual variability of isoprene mixing ratios in canopy profiles from nine intensive campaigns from Nov 2012 to Oct 2015 (Fig. 2a and Table 1). Figure 2b shows the leaf area density profile measured around the INSTANT tower in Oct 2015 and the mean canopy height. In general, isoprene mixing ratios were higher during the dryto-wet transition season (Nov 2012) and the dry season (Aug 2014 and Oct 2015/El-niño 420 year) than the wet season (Feb and Mar in 2013 and 2014) and the wet-to-dry transition 421 season (Jun 2013); with an exception for the Sep 2013-dry season that showed values 422 comparable to the 2014-wet season, although still higher than the 2013-wet season. 423 Interestingly, mean isoprene mixing ratios in Oct 2015 (El-niño dry season) were slightly 424 higher than those observed in Aug 2014 and Sep 2013 (both dry seasons) but not higher 425 than those observed in Nov 2012 (dry-to-wet transition) (although this was variable and 426 not significant). Seasonal changes in isoprene mixing ratios and fluxes from central 427 Amazonia have already been reported and were related to variations in temperature, light 428 availability at the surface, and leaf phenology (Yáñez-Serrano et al., 2015; Alves et al., 429 2016, 2018; Wei et al., 2018; Langford et al., 2022), but the assessment of inter-annual 430 variability of consecutive years including anomalous years was lacking. Considering the 431 increased air temperatures observed in the 2015-El-niño dry season (Fig. 1b) and the fact 432 that tropical plant species emit high amounts of isoprene at high temperatures (Harley et 433 al., 2004; Alves et al., 2014; Jardine et al., 2014, Garcia et al., 2019; Rodrigues et al., 2020), 434 one could expect considerably higher emission and thereby high air mixing ratios of 435 isoprene during this extreme year. However, the 2015-El-niño dry season might have been 436 stressful for plants, with the anomalous drought (see soil moisture reduction in Fig. 1 d) likely offsetting the high-temperature stimulus on isoprene emission. This finding can be 437 438 supported by two studies performed on this study site. Firstly, isoprene emission measured 439 in hyperdominant tree species showed a reduction in emission from the wet to the dry 440 season with a compensating increase in emission of monoterpenes and sesquiterpenes that 441 have both temperature-dependent emissions, indicating that the reduction in isoprene 442 emission and the shift toward heavier compounds resulted from abiotic stresses (e.g., 443 drought) during the dry season (Gomes Alves et al., 2022), which might be substantially 444 higher in an extreme El-niño year. Secondly, the anomalous post-drought leaf flush 445 observed in Feb-Mar 2016 suggested that trees flushed out new leaves to recover from the 446 stress suffered during the 2015-El-niño dry season (Gonçalves et al., 2020).

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448 Another interesting result was the seasonal variation in the shape of the isoprene mixing 449 ratio profiles (Fig. 2a). In general, all wet seasons (Feb-Mar 2013/2014) and the wet-to-dry 450 transition season (Jun 2013) data showed a constant profile with no clear vertical gradient 451 of isoprene. On the other hand, the dry seasons (Sep 2013, Aug 2014, and Oct 2015) 452 showed maximum mixing ratios between 12 m and 24m, and the dry-to-wet transition 453 season (Nov 2012) presented a well-defined peak at 24 m. This variation in the shape of 454 the isoprene mixing ratio profiles likely resulted from changes in isoprene emission across 455 seasons. Even though isoprene mixing ratio profiles are a combination of emission and air 456 mixing, when we analyzed the Bowen ratio at 24 m (figure S2) and the potential 457 temperature profiles (4-81 m; figure S1) across seasons, we observed that in-canopy air 458 mixing and the atmospheric stability were similar among seasons. This implies that 459 changes in isoprene mixing ratio profiles were predominantly attributed to the increase in 460 emission in certain layers, mostly at the upper canopy, during the dry and dry-to-wet 461 transition seasons. Furthermore, we suggest that the process that results in variation in the shape of isoprene mixing ratio profiles is a combination of variations in the canopy leaf 462 area density profile and canopy leaf age distribution throughout the year. The total amount 463 464 of LAI has a small variation over the year; still, the fractions of leaf ages that compose this 465 total LAI changes seasonally (Wu et al., 2016), as well as the shape of the canopy leaf area 466 density profile, with significant changes at the upper canopy (Martins Rosa, 2016). During the wet-to-dry transition season (May-Jun) and the dry season (Jul- Oct), upper canopy 467 468 trees presented leaf abscission and leaf flushing (Lopes et al., 2016, Gonçalves et al., 2020), 469 and the maturing process on the following months toward the beginning of the wet season 470 (Nov-Jan) might translate into higher leaf area density at the upper canopy (Martins Rosa, 2016) and higher gross primary productivity (GPP) (Botía et al., 2022). This implies that 471 472 two processes might be simultaneously occurring: one is that when there are more leaves 473 at the upper canopy, less light penetrates the canopy, which might induce the maximum 474 isoprene emission at the upper canopy as observed in Nov 2012; the other one is that leaves 475 at the upper canopy can have higher photosynthesis rates and, consequently, a higher 476 isoprene emission factor when they are mature (Alves et al., 2014), and more mature leaves 477 and higher GPP were observed in this study site during the dry-to-wet transition season and 478 beginning of the wet season (Lopes et al., 2016; Goncalves et al., 2020; Botía et al., 2022).

479

In addition, it has been suggested that seasonal variations in isoprene emissions could result from a variation in the isoprene emission factor with leaf aging, but there were not enough observational studies to support it in the Amazon (Alves et al., 2018). Therefore, in the next section, we show for the first time in-situ observations of isoprene emission factor at leaf-level with known leaf age and infer how this, together with variation in canopy leaf age distribution, likely affected intra- and inter-annual variability in emission during sequenced years.

487

488 3.2 Seasonal changes in the isoprene emission factor (E_s)

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The isoprene emission factor (E_s ; parameter measured at 1000 μ mol m⁻² s⁻¹ PAR, 30 °C) of 490 491 an ecosystem is determined by the fraction of species that emits this compound and by 492 variations in the E_s magnitude within species. Such variations may be conditioned by leaf 493 phenological status (e.g., young leaves have no or low emission, and old leaves emit less 494 isoprene than mature leaves) and environment (e.g., sun-leaves have higher E_s than shade-495 leaves) (Niinemets, 2016). We performed measurements of E_s from sun-adapted leaves 496 across different ages in 21 trees (from 20 tree species) located at the upper canopy and around the tower, and values ranged from 0 to 3.52 mg m⁻² h⁻¹ (see all species and emission 497 498 values in table S1). Of these 21 trees, 60 % had isoprene emission detectable by our 499 analytical system (TD-GC-TOFMS), while the other 40% did not. To evaluate whether 500 the E_s changes with leaf aging, we calculated the E_s ratios of mature (3–6 months) to young 501 (0-1 month), growing (1-2 months), and old (>6 months) leaves within the same tree 502 individual. We observed that, for some trees, E_s can be reduced by half when leaves are 503 older than six months (Fig. 3 and table S1), but the average of all trees combined showed 504 a statistically significant E_s reduction of 36% in old leaves compared to mature leaves 505 (paired t-test, p-value < 0.05).

506

As tropical species represent a mix of phenotypes with the predominance of non-deciduous plants, it was impossible to sample all leaf age classes for all tree species measured. Nevertheless, our dataset covers leave ages from 15 to 578 days (table S1), and we observed that all leaves measured at the young leaf age class did not show detectable isoprene

511 emission, and two leaves measured at the growing leaf age class showed emissions similar

512 to the mature leaf age class (Fig.3 and table S1). As our sampling did not cover a broad range of leaf ages below 60 days, especially among isoprene emitters, to improve the 513 514 robustness of our analysis, we added another species that had the E_s measured from the leaf 515 flushing day until the 30th day (young class) and at 226-227 days (old class) in the 516 southwestern Amazonia (Kuhn et al., 2004b). With this tree species added, we calculated 517 that the emission activity of E_s of young (0–1 month) and old (>6 months) leaves were, 518 respectively, 1% and 64% of the E_s observed in growing (1–2 months) and mature leaves 519 (3-6 months) (paired t-test, p-value <0.05), and that there was no statistically significant 520 difference between growing and mature leaves (paired t-test, p-value >0.05) (Fig. 3 and 521 table S1).

522

523 Furthermore, we observed that emitter species from our dataset could be combined into 524 two qualitative emission categories – medium emitter and low emitter –, given their E_s 525 magnitude compared to other leaf-level measurements in Amazonia (see a detailed 526 compilation in Yãnez-Serrano et al., 2020), and high emitter, with the data from the tree 527 species measured in southwestern Amazonia (Kuhn et al., 2004b) (Fig. 3). The maximum 528 E_s occurred in different leaf ages for each emitter category. Still, both high and medium 529 emitters had an E_s maximum before 150 days (mature). In contrast, the low emitter category 530 showed an E_s maximum in 295 days (old) for one species, but that was not statistically 531 significant when compared to all low emitter species (paired t-test, p-value >0.05). 532 Therefore, this indicates that species that emit considerable amounts of isoprene have 533 maximum E_s when their leaves are mature.

534

535 The variation of E_s across leaf ages is already known, also for tropical tree species (Kuhn 536 et al., 2004b; Alves et al., 2014); however, the quantification of these variations across 537 different species is still a challenge given the high biodiversity in the Amazonian rainforest, 538 and, although our results show the variation of E_s across leaf ages for more species than 539 previously reported, it is still necessary to further develop tools to upscale these results to the ecosystem level. Earlier studies indicated that the capacity to emit isoprene is more 540 541 common, and the E_s magnitudes are expected to be the highest in deciduous tree species 542 (Harrison et al., 2013; Dani et al., 2014). In fact, the high emitter (Fig. 3) is a tropical 543 deciduous tree species with a large range of variation in E_s within 30 days after leaf flushing 544 and with the maximum E_s observed in mature leaves at the end of the dry season (Kuhn et 545 al., 2004b). However, the number of deciduous trees that have regular leaf abscission and 546 leaf flushing during the dry season in central Amazonia may represent less than 15% of the 547 whole tree assembly (Gonçalves et al., 2020), which means that the effect of high 548 variability in the E_s with leaf aging from those trees might be low at the ecosystem level, 549 especially when we compare it with the other trees that showed less variability in the E_s 550 (Fig. 3, table S1).

551

552 Furthermore, for Amazonian tree species, the categorization of phenological habits goes 553 beyond evergreen and deciduous. Here, with a dataset of 194 trees (Fig. 4, and table S2) 554 monitored with a phenocam for leaf phenology and demography from 2013 to 2018, we 555 derived: (i) the camera-based canopy leaf area index (LAI) separated into four leaf age 556 classes - young (<=1 month), growing (1-2 months), mature (3-6 months), and old (>6 557 months) (Fig. 4a); and (ii) four classes of phenology (phenotypes) - evergreen, semi-558 evergreen, brevi-deciduous, and semi-brevideciduous (Fig. 4c), based on the frequency of 559 events of leaf abscission and leaf flushing (more details in Supplementary Information). 560 Then, we assigned the isoprene trait for these tree species with measurements and literature 561 data, and imputed the trait to non-measured species by following the method described in 562 Taylor et al. (2018) (Fig. 4 c). We observed that the isoprene trait did not have a higher 563 percentage within brevi-deciduous and semi-brevideciduous phenotypes, which have 564 regular and seasonal leaf abscission and leaf flushing. Instead, all phenotypes had a similar 565 fraction of isoprene emitters (Fig. 4c). This implies that leaf age is an important factor for 566 the magnitude of E_s regardless of phenotype.

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577

568 Although we do not have enough data to infer the phenotypes for the species monitored at 569 the branch level, we observed that the leaf age distribution of the 36 trees (Fig. 4b) was 570 similar to the 194 trees monitored with the phenocam (Fig. 4a); and that the fraction of 571 isoprene emitters was also similar when measured (21 trees -60% emitters; Fig. 3) and 572 non-measured (15 trees - 47% emitters) were combined (56% emitters) (Fig. 4d) and 573 compared to the phenocam trees (60% emitters) (Fig. 4c). Note that the tree species that 574 had no isoprene emission trait reported in the literature and did not fill the assumptions 575 necessary to input the trait, according to Taylor et al. (2018), were assigned with the 576 unknown flag (NA).

578 The similarity found in the seasonal leaf age distribution between the 194 trees monitored 579 by the phenocam and the 36 trees monitored at the branch level (Fig. 4) is in agreement 580 with the results presented by Gonçalves et al. (2020), which showed that the leaf phenology 581 and demography of the 194 trees are representative of the region of this study, by 582 comparing it to corresponding satellite vegetation indices retrieved from MODIS-MAIAC 583 (Multi-Angle Implementation of Atmospheric Correction). Also, this, together with the fact 584 that the isoprene trait distribution was similar among the scales (leaf level and upper 585 canopy), implies that the E_s variation with leaf age measured here can be used to optimize 586 model estimates for intra- and inter-annual isoprene emission.

- 587
- 588 3.3 Modeling of isoprene emission589

590 We used MEGAN to estimate isoprene emissions for the periods that we have in-situ 591 observations of isoprene and model inputs without considerable gaps, i.e., the years 2014 592 and 2015. We performed four different simulations (Fig. 5 and Table 2). For our first 593 simulation (S1), we applied MEGAN default settings for tropical vegetation (Fig. 5c,d), 594 which means that we used the E_s assigned to the broadleaf evergreen tropical tree and the 595 broadleaf deciduous tropical tree that is equal to 7 mg $m^{-2} h^{-1}$ (Guenther et al., 2012), half-596 hourly averages of air temperature and PPFD data measured at the same tower as the 597 isoprene observations (Fig. 5a,b), and no change in the leaf age algorithm. For the second 598 simulation (S2), we used a modified leaf age algorithm by adding the monthly distribution 599 of the LAI separated into leaf age classes (young, growing, mature, and old) as described 600 in the section above (Fig. 5c,d).

601

For a direct comparison between observations and model simulations, we performed eddy covariance (EC) isoprene flux measurements during 11 days during Nov 2015 and compared them with the simulations (Fig. 6). The isoprene emission sensitivity to the PPFD circadian cycle was well simulated by MEGAN when estimates were compared with EC 606 isoprene flux (r²=0.84, p-value <0.01) (Fig. 6 g). However, MEGAN simulations (S1 and 607 S2) overestimated the magnitude of emissions when compared with EC isoprene flux 608 around noontime (Fig. 6b); S1 and S2 had a daily average flux 2.71 and 2.68 times higher 609 than EC isoprene flux (p<0.01), respectively (Fig. 6h). This overestimation was a result of a high value for E_s in the model setup (7 mg m⁻² h⁻¹). To support this finding, we calculated 610 E_s from the observed EC isoprene flux data from 06:00 to 18:00 with the G93 algorithm 611 612 (Guenther et al., 1993), and E_s resulted in 3.21±1.76 mg m⁻² h⁻¹. We then ran a third 613 simulation (S3) with the corrected E_s (3.21 mg m⁻² h⁻¹) (Fig. 5c,d; Fig. 6b) and S3 estimates 614 presented a daily average flux 1.23 higher than EC isoprene flux (p=0.013) (Fig. 6b,h). The 615 mean E_s calculated from EC isoprene flux is in the same range as the E_s observed for the leaf level emissions of 21 trees (Fig. 3 and table S1), indicating that E_s from this study site 616 617 is lower than the one set in the model default.

618

619 Another modification in the model was done based on our leaf-level measurements. In 620 section 3.2, we present the E_s variation across leaf ages and suggest that the seasonal 621 variation in canopy leaf age distribution results in an emergent property to canopy seasonal 622 variation in E_s . With the LAI separated into leaf age classes (phenocam data) along the year 623 and the ratios of E_s (leaf level measurements) between mature and young leaves, mature 624 and old leaves, and mature and growing leaves, we modified the leaf age emission activity 625 factor of the leaf age algorithm in MEGAN. The modified leaf age emission activity factor 626 accounts for lower values of E_s in young and old leaves compared to mature and growing 627 leaves (Table 2). In our fourth simulation (S4) (Fig. 5c,d; Fig. 6b), we added the 628 modification in the leaf age emission activity factor, which led to a daily average 1.15 629 higher than EC isoprene flux (p=0.03) (Fig. 6 h).

630

631 To evaluate the effectiveness of our modifications in the model on intra- and inter-annual 632 timescales, we compared the isoprene mixing ratios observed at 38m height in all 633 campaigns performed in 2014 and 2015 with the four simulations. As our observations, 634 except for Nov 2015, are mixing ratios, it is only possible to indirectly compare with 635 MEGAN using an atmospheric model. However, considering that: air mixing and atmospheric stability were similar among the seasons (figures S2 and S3); isoprene 636 637 emission is primarily driven by changes in light, temperature, and leaf phenology (Alves 638 et al., 2018), and the variability of these factors was included in the model; we can still test 639 the comparability of the changes in the magnitudes from our measurements and simulations 640 that resulted from intra- and inter-annual variations. In figure 7, we show linear regressions 641 between observations and simulations. All datasets were filtered to the period between 12-642 15h, local time, to evaluate the time of the day with maximum emission and high mixing 643 in the surface layer and to reduce variability in photochemical isoprene loss rates. Figure 7 644 shows that, apart from the slope, all simulations were similarly and significantly 645 comparable to observations (r2=0.41 and r2=0.42, p<<0.01). However, it is important to 646 note that the finding of observed reduced E_s , compared to the model default settings, and 647 its changing across leaf ages may have an important effect on isoprene intra-annual 648 variation. Therefore, we expect that if more isoprene flux data, especially from long-term 649 measurements, were available for comparison with our simulations, we could have more 650 significant results in comparing observations and the simulations with all modifications in 651 MEGAN (S4). Additionally, as significant day-to-day isoprene variability was observed -

also over other Amazon regions, with isoprene concentrations of similar magnitudes occurring during both wet and dry seasons, likely resulting from the longer wet season lifetimes of isoprene (Wells et al., 2022), long-term flux measurements could help by offering a direct comparison between observations and modeling, and the possibility to evaluate atmospheric chemical processes.

657

658 In general, the modifications for the E_s (S3 and S4) and the leaf age activity factor (S4) 659 improved the estimates because they account for biological factors that have intra- and 660 inter-annual variations in this study site (Gonçalves et al., 2020), which represent a major 661 source of uncertainty in MEGAN (Niinemets et al., 2010). In this light, the main improvement presented here resulted from the E_s correction since our observations showed 662 that E_s was less than half of the value in the model default settings and that E_s varies 663 significantly among leaf ages. This is important because E_s is a crucial factor in determining 664 665 the magnitudes of emission of a specific site, which may vary substantially in Amazonia. 666 Although a long-term canopy flux measurement study in central Amazonia indicated that 667 E_s does not vary seasonally and argued that intra-annual changes in isoprene emission resulted only from micrometeorological and LAI variations (Langford et al., 2022), other 668 669 studies in central Amazonia have shown that emission varies substantially in a relatively 670 small spatial scale and on topographic gradients (Gu et al., 2017; Batista et al., 2019); and, 671 more recently, leaf-level measurements have shown that E_s varies within tree species both 672 seasonally and spatially, in particular when these species occur in different forest types and 673 topography (Gomes Alves et al., 2022).

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- 675 676

3.4 Implications of intra- and inter-annual variabilities in isoprene emission for modeling

677 Despite the high variability within seasons, our results showed significant changes between 678 seasons. Previous studies have shown the strong seasonality of isoprene emission in central 679 Amazonia, and we corroborate these studies that indicated changes in solar radiation, 680 temperature, and leaf phenology, as the important drivers of isoprene intra-annual 681 variability (e.g., Yáñez-Serrano et al., 2015; Alves et al., 2016, 2018). However, here we 682 further develop our understanding concerning the effect of leaf phenology, by suggesting 683 that there is seasonal variation in the ecosystem E_s resulting from changes in canopy leaf 684 age distribution, which may significantly contribute to the seasonality in the magnitude of actual emission rates. This is supported by our leaf-level E_s measurements, which showed 685 686 significant differences among leaf ages, with maximum values for mature leaves, and by 687 our results on canopy leaf age distribution changes. Furthermore, it is important to note 688 that leaf-level E_s from Oct-Nov 2017 showed maximum values for mature leaves, and those 689 were similar to the canopy E_s measured in Nov 2015. Oct and Nov (dry season and dry-to-690 wet transition seasons) are months with a substantially higher fraction of mature leaves in 691 the canopy compared to those from the wet and wet-to-dry transition seasons, meaning that 692 the highest values of E_s from mature leaves likely predominate the ecosystem E_s in Oct-693 Nov. In this sense, understanding how the E_s changes over seasons due to leaf age 694 composition within LAI will considerably improve model estimates of intra-annual 695 variations in isoprene. However, more long-term measurements of canopy isoprene flux 696 are needed to test it.

698 Surprisingly, inter-annual variabilities were less pronounced than intra-annual variability 699 when comparing normal years with the 2015-El-niño year. Our air temperature 700 measurements showed a significant increase during the dry season of 2015-El-niño year 701 compared to normal years. On a larger scale, regional land surface temperature retrieved by satellite showed an increase of up to $+4 \circ C$ from Oct to Dec 2015 in the Amazon basin 702 703 (Jiménez-Muñoz et al., 2016), and that was accompanied by a significant negative 704 maximum climatological water deficit in 43% of the whole Amazon rainforest (Aragão et 705 al., 2018). Such stresses were expected to provide a stimulus for isoprene emission, as it is 706 already largely known that isoprene emission can increase with increasing temperature and 707 that some studies have also shown that emissions increase after moderate drought (e.g., 708 Werner et al., 2021, Byron et al., 2022). However, our results did not show a significant 709 increase in isoprene mixing ratios in Oct 2015 compared to the dry seasons of previous 710 years, indicating that emissions were lower in Oct 2015, with the isoprene mixing ratio 711 profiles unlikely affected by in-canopy air mixing changes as suggested by the in-canopy 712 atmospheric stability analysis (figure S3). Understanding mechanisms of intra- and inter-713 annual variations in canopy emissions of isoprene is essential for predicting their influence 714 on atmospheric chemical-physical processes. For example, the contribution of isoprene as 715 a sink for hydroxyl radical (OH) varied seasonally (Nölscher et al., 2016); however, it did 716 not vary significantly when a normal year and the 2015-El-niño year were compared in this 717 study site (Pfannerstill et al., 2018), leading to the conclusion that these forests contributed 718 to the emission of other compounds to cope with the stress during the 2015-El-niño year, 719 resulting in an effect on the atmospheric oxidative capacity (Pfannerstill et al., 2021).

720

721 Some models predicted that higher temperatures and extended drought periods resulting 722 from climate change might increase global isoprene emissions (Pegoraro et al., 2006). 723 However, more recently, a synthesis of studies performed in the Amazon suggested that, 724 as the increase in temperature comes along with biomass loss given deforestation and forest 725 degradation, a decrease in isoprene emission from Amazonia may be expected (Yáñez-726 Serrano et al., 2020). Interestingly, although isoprene mixing ratios were not considerably 727 higher in the dry season of the 2015-El-niño year, previous studies observed higher 728 monoterpene mixing ratios compared to other dry seasons (Yáñez-Serrano et al., 2018) and 729 even higher monoterpene mixing ratios in drier and warmer days of the 2015-El-niño dry 730 season (Pfannerstill et al., 2018). In addition, another study conducted in central Amazonia 731 reported that the heat in 2015/16 led to a shift in plant emissions to more reactive 732 monoterpenes such as β -ocimene and that at high temperatures, monoterpene emissions 733 can be decoupled from photosynthesis (Jardine et al., 2017). Recently, leaf-level E_s 734 measurements in hyperdominant tree species in this study site showed that photosynthesis 735 and isoprene decreased while monoterpenes and sesquiterpenes proportionally increased 736 in the dry season, suggesting that plants might have emitted heavier compounds to cope 737 with the stress caused by high temperatures and potentially drought (Gomes Alves et al., 738 2022). We suggest that anomalies in isoprene emission during extreme years are less 739 expected than anomalies in emissions of monoterpenes and sesquiterpenes since plants may 740 also emit compounds from their storage pools when there is a limited carbon supply to 741 produce isoprene, as might be the case of plants reducing photosynthesis under heat and 742 drought stresses.

744 Summary and conclusions

745

746 Understanding mechanisms of intra- and inter-annual variations in canopy emissions of 747 isoprene from Amazonia is essential for predicting their influence on atmospheric 748 chemical-physical processes, especially when considering the role of Amazonia in the 749 global BVOC emission budget. Earlier studies presented seasonal isoprene emissions and 750 related them to the seasonality of temperature, solar radiation, and leaf phenology. 751 Nevertheless, to the best of our knowledge, this is the first study showing the E_s variation 752 across leaf ages for several Amazonian tree species and the first attempt to represent the 753 effect on seasonal isoprene flux with a model parameterization. Also, by comparing 754 observations of normal years to the extreme 2015-El-niño year, we were able to show that 755 isoprene emission does not substantially increase as a result of higher temperatures. We 756 suggest that the stress caused by elevated temperatures and droughts in extreme years might 757 reduce the isoprene temperature dependence, which is not currently well represented in 758 modeling.

759

760 Even though there are uncertainties related to measurements and model simulations, the 761 results presented here suggest that E_s varied seasonally and that this is a key factor in 762 improving model predictions. Additionally, previous studies showed that a distinguished 763 high monoterpene emission accompanies a non-pronounced increase in isoprene emission 764 in extreme years during the dry season at this study site, which is interesting to investigate 765 further since monoterpenes have higher reactivity in the atmosphere. Therefore, more 766 detailed and long-term measurements of isoprene and other BVOCs are encouraged to 767 improve our understanding of the intra- and inter-annual variability of BVOC emissions in 768 Amazonia, especially measurements that also account for biological factors that might 769 contribute to more mechanistic surface emission modeling and subsequently lead to better 770 predictions of atmospheric chemical-physical processes.

771

772 Data availability

773 Datasets are available upon request on <u>https://attodata.org</u>.

774 775

776 Authors' contributions

777 Eliane Gomes Alves has designed this study and performed the leaf-level measurements, 778 the statistical analysis of observational datasets, and the MEGAN simulations. Raoni 779 Santana and Cleo Quaresma have contributed to the analysis of the datasets of canopy 780 isoprene mixing ratios and of micrometeorology. Santiago Botía has contributed to the 781 analysis of the phenocam dataset and performed the MEGAN simulations. Tyeen Taylor 782 contributed new measurements of isoprene emissions from tropical tree species and the 783 imputation modeling of isoprene trait to the tree species monitored by the phenocam. Ana 784 Maria Yáñez-Serrano and Jürgen Kesselmeier have provided the canopy isoprene mixing 785 ratios dataset. Pedro Ivo Lembo Silveira de Assis and Giordane Martins have contributed 786 with the leaf age monitoring at the branch level. Rodrigo de Souza and Sergio Duvoisin 787 Junior contributed to the collection of isoprene samples measured at leaf-level. Alex 788 Guenther and Dasa Gu have contributed with the chemical analysis of isoprene samples 789 measured at leaf-level and the MEGAN simulations. Anywhere Tsokankunku, Matthias 790 Sörgel, Efstratios Bourtsoukidis, and Jonathan Williams contributed with the dataset of 791 eddy covariance isoprene flux. Bruce Nelson and Davieliton Pinto contributed to the 792 collection and the analysis of the phenocam dataset. Shujiro Komiya contributed to 793 analyzing the micrometeorology dataset to run the MEGAN simulations. Diogo Martins 794 contributed to the surface LiDAR data collection and analysis. Bettina Weber and Cybelli 795 Barbosa contributed with the temperature dataset to run the MEGAN simulations. Michelle 796 Robin contributed new measurements of isoprene emissions from tropical tree species. 797 Kenneth Feeley, Alvaro Duque, Viviana Lemos, Maria Contreras, Alvaro Idarraga, 798 Norberto Lopez, Chad Husby, and Brett Jestrow contributed expert guidance, specimen 799 curation, field assistance, and botanical identifications for isoprene measurements from 800 trees in botanic gardens and private collections. Iván Mauricio Cely Toro contributed to in-801 canopy micrometeorology analysis. All authors contributed to the writing of the 802 manuscript.

803

804 **Competing interests**

805 The authors declare that they have no conflict of interest

806

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Tables 1265

Table 1. Isoprene mixing ratios (ppbv) at 38 m for all field campaigns. Mixing ratios are
mean values of isoprene measured at 12:00-15:00, local time (UTC-4h). Values within
brackets are one standard deviation of the mean and the number of sampling days for each
campaign.

Year	Month	Season	Isoprene (ppbv) at 38 m
2012	November	dry-to-wet transition season	9.30 (4.90) (n=4 days)
2013	February	wet season	1.10 (0.66) (n=6 days)
2013	March	wet season	1.84 (1.44) (n=3 days)
2013	June	wet-to-dry transition season	1.83 (0.82) (n=5 days)
2013	September	dry season	5.02 (1.99) (n=8 days)
2014	February	wet season	5.92 (4.89) (n=3 days)
2014	March	wet season	2.92 (2.50) (n=11 days)
2014	August	dry season	7.76 (2.49) (n=15 days)
2015	October	dry season – <i>El-Niño</i> year	8.94 (1.41) (n=13 days)

	1 st model simulation (S1)	2 nd model simulation (S2)	3 rd model simulation (S3)	4 th model simulation (S4)
PPFD and air temperature	30 min averages – tower measurements	30 min averages – tower measurements	30 min averages – tower measurements	30 min averages – tower measurements
β^1	0.13	0.13	0.13	0.13
LDF ²	1	1	1	1
$C_{t1}{}^3$	95	95	95	95
${\rm C_{eo}}^4$	2	2	2	2
Isoprene emission factor (<i>E</i> _s)	7 mg m ⁻² h ⁻¹	7 mg m ⁻² h ⁻¹	3.21 mg m ⁻² h ⁻¹	3.21 mg m ⁻² h ⁻¹
LAI	5.32	5.32	5.32	5.32
Leaf age algorithm – LAI	default	Modified with leaf age classes derived from the phenocam: young leaves $(0-1 month)$, growing $(1-2 months)$, mature leaves $(3-6 months)$, old leaves (>6 months).	Modified with leaf age classes derived from the phenocam: young leaves $(0-1 month)$, growing $(1-2 months)$, mature leaves $(3-6 months)$, old leaves (>6 months).	Modified with leaf age classes derived from the phenocam: young leaves $(0-1 month)$, growing $(1-2 months)$, mature leaves $(3-6 months)$, old leaves (>6 months).
Leaf age emission activity factor	default $A_{new}=0.05$ $A_{gr}=0.6$ $A_{mat}=1$ $A_{old}=0.9$	default $A_{new}=0.05$ $A_{gro}=0.6$ $A_{mat}=1$ $A_{old}=0.9$	default $A_{new}=0.05$ $A_{gr}=0.6$ $A_{mat}=1$ $A_{old}=0.9$	modified according to leaf-level measurements: $A_{new}=0.01$ $A_{gro}=1$ $A_{mat}=1$ $A_{old}=0.64$

Table 2. Model parameters for all simulations for the years 2014 and 2015.

Note: Empirical coefficients are from Guenther et al. (2012)

1. Temperature empirical coefficient

1282 1283 2. Light-dependent fraction

3. Temperature empirical coefficient

- 4. Emission-class dependent empirical coefficient

Figures



1299 Figure 1. Seasonal variation of solar radiation (a), air temperature (b), precipitation (c), and soil moisture (d) during normal years (2013, 2014, 2017, 2018, and 2019), an El-niño (2015), and post-El-niño year (2016) - measured at the ATTO site. Boxplots present the median, the lower, and the upper quartiles, where the upper quartile corresponds to the 0.75 quantile and the lower quartile corresponds to the 0.25 quantile; whiskers connect the upper quartile and lower quartile to the maximum and minimum nonoutliers, respectively; and outliers are values that are more than 1.5*IQR (interquartile range) away from the top or bottom of the box.



1312 Figure 2. Mean isoprene mixing ratios for all field campaigns from Nov 2012 to Oct 2015, with one standard deviation - 12:00-15:00 local time, UTC-4h - a daytime period that isoprene emission is the highest; and mean canopy leaf area density profile with a confidence interval of 95% (b). The measurements of all intensive campaigns were collected at the same heights (0.05, 0.5, 4, 12, 24, 38, 53, and 79 m), but note that in the plot (a), the heights were shifted by 50 cm only for the better visualization of the error bars.



Figure 3. Isoprene emission factor (E_s) across leaf age classes and characterized into qualitative emission categories - low, medium, and high. Measured tree species were categorized into medium (blue) and low (black) emitters according to their E_s values, and different symbols represent different tree species. The high emitter category (red) represents a tropical species measured in Kuhn et al. (2004b). Values represent observations of individual trees and mean and one standard deviation for the categories medium and low emitters at mature and old leaf age classes. Shade areas represent the intervals of days for each leaf age class. The inset figure shows the mean E_s ratios of mature (3-6 months) to young (0-1 month), growing (1-2 months), and old (> 6 months) leaves calculated from the ratio of each individual tree.





EVERGREEN

Figure 4. Leaf phenology and demography and isoprene emission trait. Panel (a) shows the leaf age distribution separated into LAI that was observed with the phenocam, in 2017; and panel (b) shows the leaf age distribution observed at branch level for 36 trees, in 2017 - note that unknown age refers to leaves that were attached to the branch at the beginning of monitoring and therefore could not be assigned to an age class. Panel (c) shows the percentual distribution of the phenotypes assigned to the 194 trees observed with the phenocam – no massive flushing, evergreen, semi-evergreen, deciduous, and semi-brevideciduous -, and the emission trait assigned to each tree species within these phenotypes – emitters, non-emitters, and NA (NA=no data available). Panel (d) presents the percentual distribution of the isoprene trait estimated to the non-measured trees (red); and the isoprene emission trait within measured tree species (blue), with measured tree species being categorized in classes of medium emission, low emission and not detected emission.



1386 Figure 5. Simulated isoprene emission flux for 2014 and 2015. Monthly average of PPFD and air temperature (a, b) measured at the INSTANT tower. Simulations for 2014 (c) and 2015 (d) are: MEGAN simulation default, no change in the LAI code, emission factor equals to 7 mg m⁻² h⁻¹ and leaf age activity factor unmodified - S1; MEGAN simulation with change for LAI - leaf age classes, emission factor equals to 7 mg m⁻² h⁻¹ and leaf age activity factor unmodified - S2; MEGAN simulation with change for LAI - leaf age classes, emission factor equals to 3.21 mg m⁻² h⁻¹ and leaf age activity factor unmodified - S3; MEGAN simulation with change for LAI - leaf age classes, emission factor equals to 3.21 mg m⁻² h⁻¹ and leaf age activity factor modified - S4. Solid lines are means, and shaded areas represent one standard deviation of the mean.





Figure 6. Observation of isoprene flux (eddy covariance) and MEGAN simulation for 11 days in November 2015. Half-hourly averages of PPFD and air temperature (a); EC isoprene flux and MEGAN simulations (b); linear regression between EC isoprene flux and PPFD (c); quadratic regression between EC isoprene flux and air temperature (d); linear regression between simulations and PPFD (e); linear regression between simulations and air temperature (f); linear regression between EC isoprene flux and simulations (g); daily mean ratios between simulations and observation (h).





Figure 7. Correlation between isoprene mixing ratios observed at 38m during Feb and Mar
2014, Aug 2014, and Oct 2015, and the four simulations done for the respective periods.
Data represent hourly averages (12-15h, local time (LT)) of each day of measurements (a).