

1 **Aggregation of Slightly Buoyant Microplastics in Three-Dimensional Vortex Flows**

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19 **Abstract**

20 Although the movement and aggregation of microplastics at the ocean surface has been well
21 studied, less is known about the subsurface. Within the Maxey-Riley framework governing the
22 movement of small rigid spheres with high drag in fluid, aggregation of buoyant particles is
23 encouraged in vorticity-dominated regions. We explore this process in an idealized model that is
24 qualitatively reminiscent of a three-dimensional eddy with an azimuthal and overturning
25 circulation. In the axially symmetric state, buoyant spherical particles that do not accumulate at
26 the top boundary are attracted to a loop consisting of periodic orbits. Such a loop exists when
27 drag on the particle is sufficiently strong. For small slightly-buoyant particles, this loop is located
28 close to the periodic fluid parcel trajectory. If the symmetric flow is perturbed by a symmetry-
29 breaking disturbance, additional attractors for small rigid slightly-buoyant particles may arise
30 near periodic orbits of fluid parcels within the resonance zones created by the disturbance.
31 Disturbances with periodic or quasi-periodic time dependence may produce even more attractors,
32 with a shape and location that recurs periodically. However, not all such loops attract, and rigid
33 particles released in the vicinity of one loop may instead be attracted to a nearby attractor.
34 Examples are presented along with mappings of the respective basins of attraction.

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40 **Significance statement**

41 This paper investigates aggregation of small, spherical, slightly buoyant, rigid particles in a
42 simple three-dimensional vortex flow. Our goal was to gain insights into the behaviour of
43 slightly buoyant marine microplastics in a flow that qualitatively resembles ocean eddies.
44 Attractors are mapped out for the steady axisymmetric, steady asymmetric, and non-steady
45 asymmetric vortices over a range of flow and particle parameters. Simple theoretical arguments
46 are used to interpret the results.

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58 **I. Introduction**

59 Marine microplastic pollution has been a rising concern for the ocean environmental and for
60 human health. Microplastics (scales < 5mm) and nanoplastics (scales < 1 μm) have been found
61 in the tissues of marine animals, some of which are consumed by humans (Landrigan, et al.
62 2023). This comes at a time when global production of plastics is projected to increase.
63 Observations of marine microplastics have been conventionally carried out using net tows and
64 mostly occurred at or near the sea surface (van Sebille et al., 2015). However, the density of
65 many types of microplastic particles, including high-density polyethylene, is sufficiently close to
66 that of sea water that suspension within the water column for long periods of time is feasible. For
67 the near-surface microplastics, Kukulka et al., 2010 and Kooi et al., 2016 present observational
68 evidence for the fast decay in concentrations with depth over the top 5 – 20 m of the water
69 column, with the vertical penetration of plastic particles dependent on the wind speed.
70 Pabortsava and Lampitt, 2020, on the other hand, show observational evidence for much deeper,
71 below-the-mixed-layer subsurface peaks for three common types of microplastics in the Atlantic
72 Ocean. Processes such as biofouling and bio-geo-chemical or photo degradation might increase
73 the density of the plastic particles and eventually lead to the sinking of microplastics from the
74 surface into the deeper part of the water column (Kaiser et al., 2017; Kreczak et al., 2021; Kvale
75 et al., 2020). Consumption by biomass with the subsequent downward vertical transport is
76 another vehicle for redistributing microplastics from the surface down. For example, Choy et al.
77 (2019) suggest that this mechanism, specifically, consumption by pelagic red crabs and giant
78 larvaceans, was responsible for the subsurface peaks in plastic particles concentrations observed
79 at depths near 250 m in Monterey Bay. Thus, microplastics have been found well beneath the

80 ocean surface, but less is known regarding their spatio-temporal and size/density distributions
81 (Shamskhany et al., 2021).

82 A potentially important aspect of the movement of plastics and microplastics is aggregation, a
83 process that occurs at the surface over large scales near the centers of the five major subtropical
84 gyres and has been attributed to Ekman drift, windage and inertia (Beron-Vera, 2021). Many
85 early models concentrated on the ocean surface, but Wichmann et al., 2019 has highlighted the
86 importance of resolving the full three dimensional circulation. If aggregation also occurs below
87 the surface, well beneath the direct influence of Ekman layers, the dynamics is likely to be
88 different. Indeed, modeling results by Wichmann et al., 2019, based on a framework created by
89 Lange and van Sebille, 2017 and Delandmeter and van Sebille, 2019, suggests that the large
90 scale accumulation associated with the garbage patches disappears below 60m depth.

91 To avoid confusion, we will refer to infinitesimal fluid elements as “fluid parcels”, and to rigid
92 plastic particles of finite size as “rigid particles”. Typically the position $x_p(t)$ of a rigid particle
93 is tracked according to

$$94 \quad x_p(t + \Delta t) = x_p(t) + \int_t^{t+\Delta t} u dt + dx_b,$$

95 where u is the fluid velocity and dx_b is an extra displacement due the non-fluid nature of the rigid
96 particle. The user can introduce custom schemes for calculating contributions to dx_b due to
97 factors such as windage and inertia (e.g. Beron-Vera et al., 2016), turbulent diffusion (e.g.
98 Kulkulka, 2012), wave induced Stokes drift (Onink et al., 2019), etc. Eulerian schemes in which
99 plastic particles are treated as concentrations, are rare, but Mountford and Morales Maqueda,
100 2019 developed an Eulerian model in which concentrations are advected by the fluid and are
101 subject to parameterized turbulence as well as sinking or rising according to a simple law

102 involving buoyancy and friction. In a similar fashion, Kvale et al., 2020 propose an Eulerian
103 model for the biological uptake and the resulting re-distribution of microplastics.

104 An alternative approach would be to use the Maxey-Riley equation (discussed below) to solve
105 for the rigid particle velocity, v , and then use the latter to compute the trajectory of that rigid
106 particle, i.e., $x_p(t + \Delta t) = x_p(t) + \int_t^{t+\Delta t} v dt$. This equation would account for the non-fluid-
107 following effects in a deductive way, however the resulting 6th-order system (for the three
108 components of velocity and position) would be computationally challenging. To better
109 understand the implications of the use of this approach while avoiding the computational burden
110 and complexity, we have elected to analyze the movement and aggregation of individual rigid
111 particles using a Maxey-Riley framework in connection with an idealized, analytically-
112 prescribed, 3D vortex flow that qualitatively resembles the geometry of the circulation in an
113 ocean eddy but is not a solution to any dynamical oceanographic equations of motion. As shown
114 by Pratt et al., 2014 and Rypina et al., 2015, kinematic models that reproduce the correct
115 geometry are able to also reproduce the important Lagrangian features of the flow. Even in our
116 simple flow, aggregation is non-trivial, often with multiple attractors present and lack of
117 attraction in some circumstances. Thus, we wanted to thoroughly explore this simple example
118 before investigating more realistic oceanic flows. We note that other idealized studies have been
119 carried out in connection with 2D wave fields and vortex flows (e.g. DiBenedetto 2018a,b and
120 Kelly et al., 2021).

121 Aggregation can be attributed to the presence of an attractor: here, an object with a dimension
122 less than three that is somehow set up by the fluid circulation patterns and towards which rigid
123 particle trajectories attract. As long as the fluid is incompressible, fluid parcels will not
124 experience attraction and will not aggregate, but plastic particles may do so. Note also that

125 because each attractor is generally associated with its corresponding basin of attraction, if rigid
126 particles are introduced outside of the basin of attraction, they will not be attracted and will not
127 aggregate towards this attractor.

128 In order to reach a better understanding of what leads to attraction and attractors in 3D flows, we
129 explore a simple canonical example in geophysical fluid dynamics, namely the flow in a rotating
130 cylinder. This flow resembles some of the characteristics of ocean eddies, including a horizontal
131 swirl and an overturning component in the vertical, but is much less complex than any realistic
132 oceanic eddy. Specifically, we use a simple analytically-prescribed phenomenological velocity
133 introduced by Rypina et al., 2015. The Lagrangian properties of this circulation have been
134 previously studied (Fountain et al., 2000; Pratt et al., 2014; Rypina et al., 2015) allowing us to
135 begin to investigate inertial rigid particles from an established base of knowledge. A prior theory
136 (Haller and Sapsis, 2008) governing the movement of rigid particles with high drag indicates that
137 accumulation is favored for slightly buoyant particles in flows dominated by vorticity, and this
138 also motivates our choice of background flow. Identification of the attractors that can arise in this
139 flow field, evaluating their reach and domains of attraction, and clarifying the circumstances that
140 lead to their formation are the primary objectives of this work. Although motivated by the
141 problem of marine microplastics, this study is, for now, mainly a curiosity-driven research
142 aiming to develop a basic understanding of the mechanisms that might lead to aggregation of
143 rigid particles in 3D flows. The hope is that with such basic understanding in hand, one could
144 later start investigating aggregation phenomena in more complex and more realistic ocean
145 mesoscale and submesoscale eddying flows.

146 **II. Methods**

147 The physics of the motion of a small, rigid sphere that moves with velocity $\vec{v}(t)$ through a fluid
 148 with pre-existing velocity distribution $\vec{u}(\vec{x}, t)$ has been the subject of investigation by Stokes,
 149 1851, Basset, 1888, Boussinesq, 1903, Faxen, 1922, Oseen, 1927, Tchen, 1947 and many others,
 150 and was put in a unifying framework by Maxey and Riley, 1983. More recent theoretical
 151 extensions include Beron-Vera et al., 2019 and Beron-Vera, 2021. We will use a form of the
 152 Maxey-Riley equation that has been extended to include constant frame rotation with angular
 153 velocity $\vec{\Omega}^*$:

$$\begin{aligned}
 154 \quad \frac{d\vec{v}}{dt} = & \frac{\rho_f}{\rho_p} \frac{D\vec{u}}{Dt} + \frac{\rho_f}{2\rho_p} \left(\frac{D\vec{u}}{Dt} - \frac{d\vec{v}}{dt} \right) - \frac{9\nu\rho_f}{2\rho_p d^2} (\vec{v} - \vec{u}) + \left(1 - \frac{\rho_f}{\rho_p} \right) \vec{g} + \frac{\rho_f}{\rho_p} \vec{\Omega}^* \times (\vec{u} - \vec{v}) \\
 155 \quad & + \frac{\rho_f}{\rho_p} 2\vec{\Omega}^* \times \vec{u} - 2\vec{\Omega}^* \times \vec{v} + \left(\frac{\rho_f}{\rho_p} - 1 \right) \vec{\Omega}^* \times \vec{\Omega}^* \times \vec{x}. \tag{1}
 \end{aligned}$$

156 The frame rotation was introduced into the non-rotating Maxey-Riley equation by replacing

$$157 \quad \vec{v}_s = \vec{v}_r + \vec{\Omega} \times \vec{x}_r, \quad \vec{u}_s = \vec{u}_r + \vec{\Omega} \times \vec{x}_r,$$

$$158 \quad \frac{D_s \vec{u}_s}{Dt} = \frac{D_r \vec{u}_r}{Dt} + 2 \vec{\Omega} \times \vec{u}_r + \vec{\Omega} \times \vec{\Omega} \times \vec{x}_r, \quad \frac{d_s \vec{v}_s}{Dt} = \frac{d_r \vec{v}_r}{Dt} + 2 \vec{\Omega} \times \vec{v}_r + \vec{\Omega} \times \vec{\Omega} \times \vec{x}_r,$$

159 where subscript “s” denotes stationary frame and subscript “r” – rotating frame. Alternatively,
 160 transformation into a rotating frame can be done following the variational method of Ripa, 1987.
 161 The subscripts “r” have then been dropped in Eq. (1) and all subsequent equations since all
 162 variables are now in the rotating frame. For non-spherical rigid particles, adjustments to the
 163 coefficients within the Maxey-Riley equations can be made to account for elliptical shapes (see,
 164 for example, DiBenedetto et al, 2018a,b and references therein) but at the cost of adding a third
 165 vector equation for the orientation of the ellipsoid. However, real microplastics often have

166 complex tangled-filament-like shapes which are poorly represented by an ellipsoid, and no
167 corrections for tangled filaments are currently available.

168 In Eq. (1), which is a statement of Newton's second law for the rigid particle, the right-hand side
169 represents, in order, the effects of inertia, added mass, drag, buoyancy, Coriolis acceleration
170 associated with the added mass, the Coriolis acceleration associated with the particle mass,
171 Coriolis acceleration associated with the fluid motion, and centrifugal acceleration. A similar
172 equation has been previously derived by Beron-Vera et al., 2019, though the centrifugal
173 acceleration does not appear there explicitly, having been combined with the acceleration due to
174 gravity in order to define an effective gravity and corresponding geopotential. Coordinates are
175 then imagined to be aligned with geopotential surfaces, though standard spherical or Cartesian
176 coordinates are usually used in practice (Vallis, 2006). Our explicit retention of the centrifugal
177 acceleration will later allow absolute vorticity to arise naturally as a quantity of central
178 importance. We have omitted the lift force, the Basset history force, and the Faxen corrections
179 (Gatignol, 1983). Faxen corrections account for the variation of the flow across the rigid particle
180 and are proportional to $a^2\Delta u$. For a particle size that is much smaller than the typical length scale
181 of the flow, these corrections are small and typically neglected (Haller and Sapsis, 2008; Beron-
182 Vera et al., 2019). The history term, which is an integral along a particle path, accounts for the
183 boundary layer effects that a particle leaves behind. It is typically ignored under the assumption
184 that the chances of other particles crossing that localized boundary layer before it decays are
185 small (Beron-Vera et al., 2019; see also Langlois et al., 2015 and Daitche and Tel, 2011 for more
186 info on the influence of the history term on the behavior of rigid particles). Finally, the lift force
187 arises when a particle rotates in a horizontally sheared flow. As shown in Beron-Vera, 2019, the
188 inclusion of the lift force leads to the next-order, $O(\tilde{\epsilon}^2)$ correction in the slow-manifold

189 approximation, and thus can also be neglected for small $\tilde{\epsilon}$. In Eq. (1), ρ_p and ρ_f are densities of
 190 the rigid particle and the fluid, d is the particle radius, ν is viscosity of the fluid, \vec{g} is the gravity
 191 vector, and $\frac{D\vec{u}}{Dt} = \frac{\partial\vec{u}}{\partial t} + \vec{u} \cdot \nabla\vec{u}$ is the fluid material derivative, evaluated for undisturbed fluid
 192 velocity at the position of the center of the rigid particle. The position $x(t)$ of a particle is
 193 determined by

$$194 \quad \frac{d\vec{x}}{dt} = \vec{v}(\vec{x}, t), \quad (2)$$

195 and together Eqs. (1) and (2) compose a coupled, 6th-order system for computation of the particle
 196 position and velocity as functions of time.

197 If the velocities and lengths are nondimensionalized using characteristic scales U and L for the
 198 background fluid flow, and L/U is used as a time scale, then Eq. (2) remains formally unchanged
 199 while the nondimensional form of Eq. (1) is

$$200 \quad \frac{d\vec{v}}{dt} = \frac{3R}{2} \frac{D\vec{u}}{Dt} + \tilde{\epsilon}^{-1}(\vec{v} - \vec{u}) + \left(1 - \frac{3R}{2}\right) \vec{g}_r + 3R\vec{\Omega} \times (\vec{u} - \vec{v}) + 2\left(\frac{3R}{2} - 1\right) \vec{\Omega} \times \vec{v}, \quad (3)$$

201 where $\tilde{\epsilon} = \frac{2\rho_f}{\rho_f + 2\rho_p}$, $\vec{g}_r = (\vec{g} - \vec{\Omega}^* \times \vec{\Omega}^* \times \vec{x}) / (L/U^2)$, $\vec{\Omega} = \frac{\vec{\Omega}^* L}{U}$ and $\tilde{\epsilon} = \frac{2}{9} \left(\frac{d}{L}\right)^2 \frac{UL}{\nu} \frac{1}{R}$ is the Stokes

202 number, the ratio of the adjustment time scale of a particle (due to drag) to the time scale of the
 203 background flow. For $\tilde{\epsilon} \ll 1$, viscous drag is the dominant force acting on the particle, implying
 204 that a particle with an initial velocity differing by an amount $> O(\tilde{\epsilon})$ from the local fluid velocity
 205 will be rapidly accelerated over a time scale $\tilde{\epsilon}$ to a velocity proximal to that of the fluid.

206 Thereafter the particle will undergo a slow evolution in which the weaker forces due to inertia,
 207 added mass, and buoyancy cause slight departures from the movement of the fluid itself.

208 The limit $\tilde{\varepsilon} \rightarrow 0$ constitutes a singular perturbation of Eq. (3), a problem that can be addressed
 209 using an approach due to Fenichel, 1979 that was originally formally developed for a steady
 210 background flow, but that has been extended by Haller and Sapsis, 2008 to include a time-
 211 varying background flow. In either case, it can be shown that following the initial viscous
 212 adjustment, the particle position and velocity tend toward a subspace or “slow manifold” on
 213 which the particle velocity is determined directly by the fluid velocity through an “inertial”
 214 equation, here extended to include frame rotation:

$$215 \quad \vec{v} = \vec{u} + \tilde{\varepsilon} \left(\frac{3R}{2} - 1 \right) \left[\frac{D\vec{u}}{Dt} + 2\vec{\Omega} \times \vec{u} - \vec{g}_r \right] + O(\tilde{\varepsilon}^2). \quad (4)$$

216 This result is the same as obtained by Beron-Vera et al., 2019, provided that their gravity vector
 217 is interpreted as our \vec{g}_r . The same authors also present more general cases, including those with
 218 the lift force and on the sphere. In Supplementary Material we present a simple derivation of Eq.
 219 (4) based on a multiple-scale expansion. It provides a quick, though less rigorous, alternative to
 220 the Fenichel approach.

221 A chief advantage of the slow manifold reduction is that the 6th order system given by Eqs. (2)
 222 and (3), in which particle velocity needs to be solved for, is reduced to a 3rd order system given
 223 by Eqs. (2) and (4), where the particle velocity is explicitly written as a function of fluid velocity
 224 and flow and particle parameters (and thus is known). The bracketed expression in Eq. (4), which
 225 determines the velocity of the rigid particle relative to the fluid, is nothing more than $\frac{\partial}{\partial x_j} \tau_{ij}$,
 226 where τ_{ij} is the stress tensor for the fluid. Thus the relative velocity of a rigid particle on the
 227 slow manifold is in the same direction as the net force that would act on a fluid parcel occupying
 228 the same space. Ordinarily, for a fluid parcel, that force would equate with an acceleration, but
 229 on the slow time scale, the relative particle velocity points in the same direction as the net fluid

230 force and its magnitude is proportional to $\tilde{\varepsilon} \left(\frac{3R}{2} - 1 \right) = \frac{2}{9} \frac{d^2}{L^2} \frac{UL}{\nu} \frac{(\rho_f - \rho_p)}{\rho_f}$. Since the aggregation of
 231 rigid particles requires departures of the particle velocity from the (divergence free) velocity
 232 field of the fluid, one can expect that aggregation will occur more slowly if d and $(\rho_f - \rho_p)/\rho_f$
 233 are small, or if ν is large. At the same time, the existence of attractors internal to the fluid may
 234 depend on $(\rho_f - \rho_p)/\rho_f$ being small: for example, a large density difference may mean that
 235 rigid particles simply sink to the bottom or rise to the surface (and are thus attracted to attractors
 236 external to the fluid interior).

237 As pointed out by Haller and Sapsis, 2008 (also see Beron-Vera et al., 2019), we can consider a
 238 continuous concentration of rigid particles with similar properties, and with smoothly varying
 239 velocity given by Eq. (4). The aggregation of such a concentration would appear to require that
 240 the divergence of that velocity be negative (though see an apparent counterexample in Fig 1c,
 241 presented later). Following Haller and Sapsis, 2008, consider the evolution of a material volume
 242 of rigid particles. The time rate of change of this volume is

$$243 \quad \frac{dV}{dt} = \oint \vec{v} \cdot \vec{n} dA_V = \iiint (\nabla \cdot \vec{v}) dV = \iiint \nabla \cdot \left[\vec{u} + \tilde{\varepsilon} \left(\frac{3R}{2} - 1 \right) \left(\frac{D\vec{u}}{Dt} + 2\vec{\Omega} \times \vec{u} - \vec{g}_r \right) \right] dV \quad (5)$$

244 where $\nabla \cdot \vec{u} = 0$ for an incompressible fluid. Shrinking V to an infinitesimal size allows the right-
 245 hand side to be approximated by V times the local value in the integrand, and the result may be
 246 integrated in time, yielding

$$247 \quad V(t) = V_0 \exp \left(\tilde{\varepsilon} \left(\frac{3R}{2} - 1 \right) \int_{t_0}^t \nabla \cdot \left(\frac{D\vec{u}}{Dt} + 2\vec{\Omega} \times \vec{u} - \vec{g}_r \right) ds \right)$$

$$248 \quad = V_0 \exp \left(-2\tilde{\varepsilon} \left(\frac{3R}{2} - 1 \right) \int_{t_0}^t [Q_r(x(s), s) + \vec{\Omega} \cdot \vec{\zeta}_r + |\vec{\Omega}|^2] ds \right)$$

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$$= V_0 \exp\left(-2\tilde{\varepsilon} \left(\frac{3R}{2} - 1\right) \int_{t_0}^t Q_a(x(s), s) ds\right). \quad (6)$$

250 Here $Q_r = \frac{1}{2}\left(\frac{1}{2}|\vec{\zeta}_r|^2 - |S|^2\right)$ is the three-dimensional Okubo-Weiss parameter (Okubo, 1970;
 251 Weiss, 1991), $\vec{\zeta}_r$ represents the relative vorticity vector for the fluid, $S = 1/2(\nabla\vec{u} + (\nabla\vec{u})^T)$ is
 252 the strain tensor, and $|S|$ is its Frobenius norm. The final step in Eq. (6) follows from
 253 introduction of the absolute vorticity vector

254
$$\vec{\zeta}_a = \vec{\zeta}_r + \overline{2\Omega} \quad (7)$$

255 and the corresponding function $Q_a = \frac{1}{2}\left(\frac{1}{2}|\vec{\zeta}_a|^2 - |S|^2\right)$. This result could have been anticipated
 256 from the fact that in the rest frame of the original Maxey-Riley equation, the vorticity and the
 257 absolute vorticity are the same.

258 We note that for a volume V of any size:

259
$$\frac{dV}{dt} = 2\tilde{\varepsilon} \left(\frac{3R}{2} - 1\right) \iiint Q_a dV = \tilde{\varepsilon} \left(\frac{3R}{2} - 1\right) \iiint \frac{\partial^2}{\partial x_i \partial x_j} \tau_{ij} dV = \frac{2}{9} \frac{d^2}{L^2} \frac{UL}{\nu} \frac{(\rho_f - \rho_p)}{\rho_f} \oint \frac{\partial}{\partial x_j} \tau_{ij} n_i dA_V,$$

260 (8)

261 where n_j denote the components of the outward unit vector normal to the bounding surface A_V .
 262 The first equality in Eq. (8) is a modest modification of Eq. (31) from Haller and Sapsis, 2008,
 263 and, as mentioned above, one could probably have guessed that our more general result could be
 264 obtained by replacing Q with Q_a . The remainder of the equation expresses volume changes in
 265 terms of the fluid stresses. Thus for buoyant particles, a volume $V(t)$ of any size will contract if
 266 the force normal to A_V due to the fluid stresses, integrated around A_V , is inward. In many cases,
 267 including quasigeostrophic eddies and gyres, internal waves, and the surface gravity waves
 268 considered by DiBenedetto et al., 2018a,b and all inviscid flows, the stress tensor is dominated

269 by pressure, i.e., $\frac{\partial}{\partial x_j} \tau_{ij} \cong -\frac{1}{\rho_f} \nabla p$, so the tendency to aggregate is determined entirely by the
 270 pressure field.

271 In general, Q_a can change sign along a particle trajectory, making it hard to predict whether the
 272 surrounding volume shrinks or expands with time. If a buoyant particle is trapped in a region in
 273 which Q_a is predominately positive, then this region is a good candidate for aggregation.

274 Persistent ocean eddies and other vortical structures are possibilities, not only because vorticity
 275 tends to dominate over strain, but also because such features have the ability to trap fluid for long
 276 periods of time. For dense particles, contraction occurs in areas dominated by strain, and it has
 277 been shown that aggregation of heavy particles can occur in strain-dominated filaments that arise
 278 in particle-laden turbulent flows, though the considered particle-to-fluid density differences tend
 279 to be quite large (see Brandt and Coletti, 2022 for a review). In our study, we will focus on
 280 vortex flows reminiscent of ocean eddies, and on lower dimension objects within such flows that
 281 can act as attractors for buoyant particles.

282 A simple example of aggregation is given by Haller and Sapsis, 2008, who argue that the
 283 elliptical center of a steady, non-divergent 2d eddy, with $\vec{g} = |\vec{\Omega}|=0$, acts as an attractor for
 284 buoyant particles. Here Q_a (now $=Q_r$), is ostensibly positive near the elliptical center of the
 285 eddy, corresponding to contraction of the phase space (which in our case coincides with the
 286 physical space) of the rigid particle motion. Since the central fixed point of the velocity field of
 287 the eddy is also a fixed point of the slow manifold particle velocity (Eq. (4)), buoyant particles
 288 initiated about the center should migrate towards the center. If the eddy is inviscid and its
 289 streamlines are circular, then the pressure and azimuthal velocity are related by the cyclostrophic

290 balance $\frac{1}{\rho_f} \frac{\partial p}{\partial r} = \frac{u_\theta^2}{r}$ so that $2Q_r = \frac{1}{\rho_f} \left(\frac{1}{r} \frac{\partial p}{\partial r} + \frac{\partial^2 p}{\partial r^2} \right)$, and for an eddy in solid body rotation ($u_\theta =$

291 $\Gamma_s r$), $2Q_r = \frac{1}{\rho_f} \left(\frac{1}{r} \frac{\partial p}{\partial r} + \frac{\partial^2 p}{\partial r^2} \right) = 2\Gamma_s^2$. As suggested in Figure 1a, a small concentration of rigid
 292 particles indicated by the cross hatched area shrinks as it moves towards the center of the eddy.
 293 The contraction is partially due to the geometric effect of movement towards smaller radius
 294 (term $\frac{1}{r} \frac{\partial p}{\partial r}$) but also due to the fact that the pressure gradient decreases to zero as the center is
 295 approached and thus the inner edge of the path moves more slowly inward than the outer part
 296 (term $\frac{\partial^2 p}{\partial r^2}$). In the case of solid body rotation the two terms contribute equally. A second
 297 example (Fig. 1b) is of an eddy with an azimuthal velocity given by $u_\theta = \Gamma_c r^{1/2}$. Here $\frac{\partial^2 p}{\partial r^2} = 0$
 298 and $2Q_r = \frac{1}{\rho_f} \left(\frac{1}{r} \frac{\partial p}{\partial r} \right) = \Gamma_c^2 / r > 0$, so the contraction of the patch is entirely due to the geometric
 299 effect of its movement towards smaller radius. The most curious case is that of a point vortex:
 300 $u_\theta = \Gamma_p r^{-1}$, for which $2Q_r = \frac{1}{\rho_f} \left(\frac{1}{r} \frac{\partial p}{\partial r} + \frac{\partial^2 p}{\partial r^2} \right) = \frac{\Gamma_p^2}{r^4} - \frac{3\Gamma_p^2}{r^4} < 0$. Here the vorticity is zero away from
 301 the eddy center and the velocity field is dominated by strain. The pressure gradient *increases* as
 302 the center of the vortex is approached, meaning that the inner part of the patch moves towards
 303 the center more rapidly than the outer portion (Fig. 1c) and this tendency (quantified by the
 304 factor $-\frac{3\Gamma_p^2}{r^4}$) surpasses the tendency towards geometrical contraction (quantified by the factor
 305 $\frac{\Gamma_p^2}{r^4}$). The area of the patch thus expands as rigid particles are drawn towards the center of the
 306 vortex. Note, however, that a patch surrounding the center of the vortex can only shrink. This
 307 behavior is made possible by the singularity at the center, and although this feature is artificial,
 308 point vortices are often used in idealized models of fluid flow and will act as sinks or “black
 309 holes” for buoyant particles even though $2Q_r < 0$.

310 The sign of Q_a is clearly not the whole story and does not encompass the effects of boundaries.
311 For example, consider the fate of heavy ($\rho_f < \rho_p$) particles in the eddy show in Fig. 1a. The
312 particles will migrate outward in each case, and no interior attraction will occur unless the eddy
313 is surrounded by a boundary, which would then act as an attractor.

314 In the next section, we will consider a more general, 3D, eddy-like circulation: one that has both
315 vertical and horizontal components of vorticity, time dependence, and a variety of vortical
316 structures that act as candidates for attraction. Our model is based on the incompressible flow in
317 a rotating cylinder (Greenspan, 1986), which has been studied in many configurations by
318 numerous authors as a model of ocean circulation (Hart and Kittelman, 1996; Pedlosky and
319 Spall, 2005), ocean eddies (Pratt et al., 2014; Rypina et al., 2015), or industrial processes and
320 engineering applications (Lopez and Marques, 2010 and references therein), and can be easily
321 set up in the laboratory setting (Fountain et al., 2000; Lackey and Sotiropoulos, 2006). In its
322 original configuration the cylinder rotates about a vertical axis at a constant (positive) angular
323 velocity ($\vec{\Omega} = \Omega \vec{k}$), and the lid, which is in contact with the fluid, rotates with a slightly greater
324 angular speed. The differential rotation sets up an azimuthal circulation in the horizontal and an
325 overturning circulation in the vertical. (Overturning is observed in ocean eddies as well and
326 Ledwell et al., 2008 present an example.) The steady, axially symmetric state of the rotating
327 cylinder flow that is established will be our first object of investigation. A steady but
328 asymmetrically-perturbed variant can be established by moving the axis of rotation of the lid
329 away from the axis of rotation of the cylinder, and this offset can also be varied in order to
330 induce time dependence. Fountain et al., 2000 set a similar situation up in a laboratory cylinder
331 using a submerged impeller that can be tilted, rather than the differentially rotating lid that can be
332 shifted, to establish an asymmetric disturbance flow. The authors discussed the Lagrangian

333 characteristics of the undisturbed flow and demonstrated the existence of secondary vortical
334 structures generated when the flow is perturbed. Pratt et al., 2014 reproduced similar structures
335 using a primitive equation simulation and explored the rich assembly of chaotic regions and non-
336 chaotic vortical structures as functions of the Ekman and Rossby numbers of the flow. The time-
337 dependent version of the rotating cylinder flow and a theory describing the resulting vortical
338 structures were discussed by Rypina et al., 2015, who based their examples on a
339 phenomenological model that reproduced many of the qualitative features of the numerically-
340 obtained velocity field. In dimensionless Cartesian coordinates, the model velocity field is given
341 by

$$342 \quad u^{(x)} = -bx(1 - 2z)\frac{r_o-r}{3} - ay(c + z^2) + \varepsilon \left[y(y - y_o + \gamma \cos(\sigma t)) - \frac{r_o^2-r^2}{2} \right] (1 - \beta z), \quad (9a)$$

$$343 \quad u^{(y)} = -by(1 - 2z)\frac{r_o-r}{3} + ax(c + z^2) - \varepsilon x(y - y_o + \gamma \cos(\sigma t))(1 - \beta z), \quad (9b)$$

$$344 \quad u^{(z)} = bz(1 - z)\frac{2r_o-3r}{3}, \quad (9c)$$

345 in which $r = (x^2 + y^2)^{1/2}$ and r_o is the cylinder radius. The velocity field consists of a steady,
346 axially symmetric flow of strength a with an overturning circulation of strength b . To this
347 symmetric state one can add an asymmetric, possibly unsteady and depth dependent, perturbation
348 of amplitude ε (not to be confused with the Stokes number $\tilde{\varepsilon}$). The perturbation is quantified by
349 an offset parameter y_o that introduces axial asymmetry in the velocity field, a frequency σ , and
350 an amplitude β for linear depth dependence and an amplitude γ for the time dependence. For the
351 case of axially symmetric, steady flow ($\varepsilon = 0$) the horizontal velocity field, in cylindrical
352 coordinates, becomes

$$353 \quad u^{(r)} = -br(1 - 2z)\frac{r_o-r}{3} \quad (10a)$$

354 and

$$355 \quad u^{(\theta)} = ar(c + z^2), \quad (10b)$$

356 where θ is the azimuthal angle. Table 1 lists the parameter values used for each numerical
357 experiment.

358 We now review the main features of the Lagrangian circulation in the rotating cylinder flow. In
359 the steady, symmetric configuration, each fluid trajectory is confined to the surface of a torus as
360 it winds around the cylinder. The typical torus is associated with quasi-periodic trajectories and
361 any such trajectory, followed for a sufficient length of time so that it completes many
362 overturning and azimuthal rotations around the cylinder, will sketch out the torus in 3D. Fig. 2b
363 contains several examples of such tori and Fig. 2a shows the corresponding Poincare map, made
364 by marking the crossing points of trajectories through a vertical slice through the cylinder. After
365 a large number of crossings each quasi-periodic trajectory traces out the cross section of the torus
366 on which it lives. The tori are nested within each another, with a single, horizontal, periodic
367 trajectory located at the center of the nest. Certain tori contain periodic trajectories, and these
368 will show up as a finite number of dots on the Poincare map. Because of this geometry, the
369 motion of fluid parcels is most naturally described in terms of action-angle-angle variables,
370 where the action, I , acts a label for a particular torus and is constant following each trajectory,
371 and the two angle variables, $\tilde{\theta}$ and ϕ , define the location of a parcel on the torus. Here $\tilde{\theta}$ is an
372 azimuthal angle that differs from the above cylindrical coordinate θ in how its origin is defined,
373 while the ‘poloidal’ angle ϕ wraps around the cross-section of each torus. The coordinates are
374 non-orthogonal but are defined in such a way that the angular velocities, $\Omega_{\tilde{\theta}}$ and Ω_{ϕ} , are also

375 constant following a trajectory. The explicit transformations to the action-angle-angle variables
376 are given in Mezic and Wiggins, 1994.

377 When the symmetric RC flow is perturbed by a small, steady, symmetry-breaking perturbation,
378 as controlled by the parameters ε and y_o in Eq. (9), the tori that are populated by periodic orbits
379 potentially become resonant and break up, resulting in chaotic motion of fluid parcels in the
380 vicinity (Fig. 2d-i). Tori with quasiperiodic orbits deform but stay intact. Examples are discussed
381 by Fountain et al., 2000 and Pratt et al., 2014, and the latter found that chaos generally dominates
382 in a large region that includes the central axis of the cylinder and extends around the boundaries
383 of the cylinder. Away from this region the space is occupied by tori that have survived the
384 perturbation, and these are sandwiched between tori that have broken up and created braided
385 regions of chaos. The breakup of a torus also gives rise to new tori that appear as islands in the
386 Poincare maps (Fig. 3d and 3g) and these contain non-chaotic trajectories. The number of islands
387 can be predicted by a theory that decomposes the symmetry-breaking perturbation into Fourier
388 modes, written in the $(I, \tilde{\theta}, \phi)$ coordinates, with wave numbers n and m in the $\tilde{\theta}$ and ϕ direction.
389 If the angular velocities $\Omega_{\tilde{\theta}}$ and Ω_{ϕ} characterizing the trajectories on a particular torus satisfy
390 the resonance condition $n\Omega_{\tilde{\theta}} + m\Omega_{\phi} = 0$ for some n and m , equivalent to the trajectories on that
391 torus being periodic, then that torus will break up and a new set of invariant tori (islands) will
392 form. Running through the center of the islands will be a periodic trajectory that will execute n
393 azimuthal cycles to every m poloidal (overturning) cycles. In the case shown in Fig. 3a, $n =$
394 $m = 1$, so the periodic trajectory circles the cylinder horizontally once for each overturning
395 cycle: a so-called 1:1 resonance.

396 If the symmetry breaking perturbation is quasi-periodic in time, with underlying frequencies σ_i ,
397 the resonance condition for the breakup of a torus becomes $n\Omega_{\bar{\theta}} + m\Omega_{\phi} + l_i\sigma_i = 0$, where l_i 's
398 are integers (Rypina et al., 2015). Unlike the resonance condition for the steady perturbation,
399 which is only satisfied on tori foliated by periodic trajectories, this new resonant condition may
400 be satisfied on tori that have quasi-periodic orbits, and the resonant islands that form will have a
401 shape and location that vary in time. An example (Fig. 2g,h) of the case of a resonance with a
402 single-frequency (i.e., time-periodic) perturbation shows a number of resonant islands. These
403 features vary in time, recovering their shape and location periodically, and the snapshots shown
404 are obtained by strobing the trajectories in 3D and at the forcing frequency. The green and blue
405 islands in Fig. 2h have resulted from the breakup of tori with quasiperiodic trajectories, and
406 center of the island corresponds to a closed material curve that is populated with quasiperiodic
407 trajectories.

408 Note that the resonance condition above and our results in general are applicable to quasi-
409 periodic disturbances with finite number of frequencies, rather than only periodic disturbances.
410 (We only show numerical simulations for the time-periodic case for simplicity.) Because any
411 broad-spectrum function can be arbitrary closely represented by a quasi-periodic function with a
412 finite number of frequencies, this could be applicable to some oceanic flows, especially those
413 with pronounced peaks in the spectrum. However, for flows with truly broadband spectrum, this
414 approach is probably poorly applicable and/or at least impractical because of the very large
415 number of discrete frequencies needed. This is similar in its utility/applicability to other
416 Kolmogorov-Arnold-Moser—based and resonance—based arguments used in prior papers by
417 many authors (including both us and the reviewer), see, for example, Rypina et al., 2007 and
418 Beron et al., 2008; 2010.

419 **III. Results**

420 Aggregation of rigid particles will occur in presence of an attractor, an object with a dimension
421 < 3 to which particles tend asymptotically in time. We are most interested in attractors that
422 occur in the interior of the rotating cylinder, and are set up by the background circulation, as
423 opposed to the physical boundaries of cylinder. We will see that a closed material contour
424 consisting of periodic orbits near the core of the nested tori in the steady symmetric case act as
425 an attractor for slightly buoyant particles, and that similar material contours consisting of
426 periodic or quasiperiodic orbits near the centers of the resonant islands in the asymmetric cases
427 can play the same role. We will explore three cases in increasing complexity, beginning with
428 steady flows with axial symmetry, and proceeding to steady, asymmetric flows and finally
429 unsteady asymmetric flows.

430 The search for attractors is motivated by the hypothesis that for cases of strong drag, where the
431 rigid particle velocity lies close to the fluid velocity, a periodic orbit for the rigid particle motion
432 will exist in the vicinity of a periodic trajectory for the fluid parcel motion, and that if $Q_a > 0$ in
433 a region surrounding the latter, that it should attract buoyant particles. For the time-dependent
434 case, we extend the search to included closed loops that contain recirculating rigid particles and
435 that vary periodically in time.

436 (a) steady, axially-symmetric 3D flows

437 The fluid velocity field for this case is given by Eqs. (9c) and (10), and these indicate that the
438 location of the horizontal, periodic fluid parcel trajectory living at the center of the nested tori, is
439 given by $r = 2r_o/3$ and $z = \frac{1}{2}$. It is natural to ask whether a periodic trajectory for rigid particles

440 also exists nearby. In the slow-manifold approximation, the steady radial, azimuthal and vertical
 441 particle velocities are obtained by writing Eq. (4) in cylindrical coordinates, leading to

$$442 \quad v^{(r)} = u^{(r)} + \tilde{\varepsilon} \left(\frac{3R}{2} - 1 \right) \left[\left(u^{(r)} \frac{\partial}{\partial r} + u^{(z)} \frac{\partial}{\partial z} \right) u^{(r)} - u^{(\theta)} \left(2\Omega + \frac{u^{(\theta)}}{r} \right) - \Omega^2 r \right] \quad (11a)$$

$$443 \quad v^{(\theta)} = u^{(\theta)} + \tilde{\varepsilon} \left(\frac{3R}{2} - 1 \right) \left[\left(u^{(r)} \frac{\partial}{\partial r} + u^{(z)} \frac{\partial}{\partial z} \right) u^{(\theta)} + u^{(r)} \left(2\Omega + \frac{u^{(\theta)}}{r} \right) \right] \quad (11b)$$

$$444 \quad v^{(z)} = u^{(z)} + \tilde{\varepsilon} \left(\frac{3R}{2} - 1 \right) \left[\left(u^{(r)} \frac{\partial}{\partial r} + u^{(z)} \frac{\partial}{\partial z} \right) u^{(z)} + g \right] \quad (11c)$$

445 Position of attracting periodic orbit; approximate analytical expression on a slow manifold

446 Searching for points $r = r_c$ and $z = z_c$ for which $v^{(r)} = v^{(z)} = 0$, and that lie in the proximity of
 447 the horizontal trajectory of the flow, we introduce

$$448 \quad r_c = \frac{2r_o}{3} + \tilde{\varepsilon} \left(\frac{3R}{2} - 1 \right) \tilde{r} \text{ and } z_c = \frac{1}{2} + \tilde{\varepsilon} \left(\frac{3R}{2} - 1 \right) \tilde{z}.$$

449 Substituting into the right-hand sides of (11a,c) and setting both to zero results, after neglect of
 450 $O(\tilde{\varepsilon}^2)$ terms, in

$$451 \quad r_c = \frac{2r_o}{3} + \tilde{\varepsilon} \left(\frac{3R}{2} - 1 \right) \frac{g}{b} r_o \quad (12a)$$

452 and

$$453 \quad z_c = \frac{1}{2} + \frac{9}{2br_o} \tilde{\varepsilon} \left(\frac{3R}{2} - 1 \right) \left[\Omega^2 + a \left(c + \frac{1}{4} \right) \left(2\Omega + a \left(c + \frac{1}{4} \right) \right) \right]. \quad (12b)$$

454 For the parameters $a > 0$ and $b > 0$, circulation is cyclonic with upwelling in the center of the
 455 cylinder, and $(3R/2) - 1 > 0$ for buoyant particles, so the $O(\tilde{\varepsilon})$ corrections are positive and the
 456 periodic particle orbit lies at larger radius and elevation than the periodic fluid orbit. Note also

457 from Eq. (11b) that the azimuthal velocity component of the rigid particle on the periodic orbit is
458 equal to that of the fluid.

459 An explanatory sketch (Fig. 3) shows the position of the periodic orbit of the rigid particle
460 relative to that of the periodic orbit of the fluid. Since the rigid particle is buoyant, it can
461 maintain its level z only if it is situated in a region where the vertical fluid velocity is < 0 , here
462 to the right of the fluid periodic orbit. Also, the horizontal pressure gradients associated with the
463 centripetal acceleration associated with the frame rotation (term $\Omega^2 r$), the Coriolis acceleration
464 (term $2\Omega u^{(\theta)}$), and the centripetal acceleration due to the azimuthal velocity $u^{(\theta)^2}/2r$ are all
465 positive for this flow, so that low pressure exists at $r=0$ and the rigid particle is forced
466 horizontally inward. To remain stationary the particle must sit in a region where the radial
467 velocity of the fluid is outward. In this manner, the periodic trajectory exists at a location where
468 the forces of inertia, buoyancy and added mass can be countered by the drag due to the
469 background flow. If we fix all other parameters and increase Ω through positive values, the term
470 multiplying $\tilde{\epsilon}$ in Eq. (12b) will become dominated by the Ω^2 term and will grow without bound
471 and the periodic trajectory may cease to exist. At the same time, a periodic orbit for the rigid
472 particle can always be found close to that of the fluid, regardless of the magnitudes of the
473 parameters Ω , a , b etc., provided that the relative particle size d/L (and thus $\tilde{\epsilon}$), and/or the
474 relative density difference $\frac{(\rho_f - \rho_p)}{\rho_f}$ (and thus $\frac{3R}{2} - 1$) are made sufficiently small.

475 *Position of attracting periodic orbit; conditions for the loss of periodic orbit*

476 We have suggested that periodic orbits for rigid particles are encouraged when the $\tilde{\epsilon} \left(\frac{3R}{2} - \right.$
477 $\left. 1 \right) \ll 1$, and in the case of Run 1 the value is 0.0066. A cross-sectional plot of the radial and

478 vertical components of the slow manifold particle velocity in a vertical section through the
 479 cylinder (Fig. 4a) shows that the periodic orbit lies at $r = 0.369$ and $z = 0.504$ (as compared to
 480 the values $r_c = 0.338$ and $z_c = 0.502$ predicted by Eq. (12). (The convergence of the
 481 surrounding velocity field is too weak to be seen in the graphic.) If $\tilde{\varepsilon} \left(\frac{3R}{2} - 1 \right)$ is raised to the
 482 moderately small value 0.02, the position of periodic trajectory migrates towards larger radius
 483 (Fig. 4b), the reason being that the greater buoyancy (larger value of $\frac{3R}{2} - 1$) or smaller drag
 484 (larger $\tilde{\varepsilon}$) requires a larger downward fluid velocity for equilibrium. Since the maximum
 485 downward fluid velocity occurs at the outer cylinder wall (see Eq. (9c)) the position of the
 486 periodic orbit continues to migrate outward and is lost (Fig. 4c) when $\tilde{\varepsilon} \left(\frac{3R}{2} - 1 \right)$ exceeds a
 487 value close to 0.3.

488 *Position of periodic orbit in numerical simulations:*

489 The slow-manifold reduction yields to the prediction (Eq. (12)) of the position of the attracting
 490 material contour, or loop, for slightly buoyant particles. We can compare this prediction to what
 491 is observed in numerical simulations using the Maxey-Riley Eqs. (1) and (2) over a range of
 492 particle size d (and thus $\tilde{\varepsilon}$) and frame rotation Ω . As shown in Fig. 5, qualitative agreement with
 493 the slow-manifold prediction, and the sketch in Fig. 3, holds for a very small d (when $\tilde{\varepsilon}$ is small).
 494 Here the attractor in Fig. 5 is located close to the central periodic fluid parcel trajectory that lives
 495 at mid-depth, $z = 0.5$ and $r = \frac{2R}{3} \approx 0.33$. As d (and $\tilde{\varepsilon}$) increases, the attractor moves
 496 increasingly up and outward, and although the theory captures the trends, quantitative agreement
 497 with the numerical results worsens. Also, when frame rotation Ω is increased (panel c), the
 498 attractor responds by shifting up from mid-depth, again in qualitative but not quantitative
 499 agreement with the slow-manifold prediction in Eq. (12b).

500 Geometry of rigid particle trajectories and evidence of attraction in numerical simulations:

501 If in the neighborhood of the periodic rigid particle trajectory $Q_a > 0$, the phase space for
502 buoyant particles will contract and the periodic trajectory becomes a candidate for an attractor of
503 such particles. An example of the attraction towards the periodic orbit is shown in Figure 2c,
504 where a set of slightly buoyant particles ($\frac{\rho_p}{\rho_f} = 0.97$) has been initialized over the volume of the
505 cylinder, and Eqs. (1) and (2) have been integrated forward in time to determine their subsequent
506 trajectories. Each trajectory is shown using a unique color. It can be seen that the particles
507 aggregate within a ring-like structure of decreasing thickness in the general vicinity of the
508 periodic orbit of the fluid flow.

509 Basin of attraction – relationship to Q_a :

510 To map out the basin of attraction for the particle periodic orbit, we first consider the region over
511 which phase space contraction for the buoyant particles (i.e. $Q_a > 0$) occurs. This region is
512 shown in Fig. 6a for the current example, along with the streamlines of the fluid overturning
513 stream function. Much of the fluid flow recirculates entirely within the region of positive Q_a ,
514 whereas some of the outer streamlines cross the boundary (thick contour) between positive and
515 negative Q_a . If it were the case that rigid particles exactly followed streamlines of the fluid
516 overturning circulation, then net contraction or expansion of phase space along a rigid particle
517 trajectory would depend on the sign of the time-integrated value of Q_a along streamlines. The
518 $Q_a = 0$ contour, shown by a bold contour in each frame of Fig. 6, might then approximately
519 delineate the basin of attraction for buoyant rigid particles. In the slow-manifold approximation,
520 where rigid particle velocities lie close to the fluid velocities, the $Q_a = 0$ contour might continue
521 to do so.

522 To test this conjecture, we locate the basin of attraction in the numerical simulations by releasing
523 buoyant particles at various locations in the cross-section $0 < x < r_o$ and $0 < z < 1$, integrating
524 the subsequent trajectories over many overturning cycles, and recording the position (x_{final} and
525 z_{final}) of each particle where it crosses the same plane the final time (i.e., recording final
526 crossing with the Poincare section). We use the variable-step 4-th order Runge-Kutta integration
527 scheme, which we implemented in Matlab via the built-in function “ode45”. In our simulations,
528 the relative and absolute tolerances are set to the value of 10^{-9} to integrate particle trajectories
529 (Eqs. (2) and (3)) (our results were not sensitive to the further decrease in tolerance values).
530 Since the flow (Eqs. (9a,b,c)) is prescribed analytically and has no normal flow component at the
531 perimeter and top and bottom of the cylinder, no interpolation scheme is needed and no extra
532 boundary conditions are enforced during the integration. Integration of a trajectory is stopped
533 when a particle got within one particle radius from the cylinder walls or top/bottom. The values
534 of z_{final} as a function of initial particle position are mapped in Fig. 7a, where the large green
535 area corresponding to $z_{final} \cong 0.5$ indicates the region from which particles are attracted. Only
536 particles initiated near the central axis of the cylinder, and close to the cylinder boundaries lie
537 outside this region, and these rise to the surface of the cylinder, contact the upper lid, and are no
538 longer followed. It can be seen that the green area in Fig. 7a has an oval shape that somewhat
539 resembles the overturning streamlines at small x in the central part of the cylinder, but extends to
540 near the top, bottom and outer cylinder boundaries at larger x . Thus the $Q_a = 0$ contour provides
541 a rough indication of the size and shape of the basin of attraction, but misses some important
542 details.

543 Basin of attraction – dependence on Ω

544 We have seen that the location of the periodic orbit that acts as an attractor for buoyant particles
545 shifts up and out in response to increasing frame rotation Ω (Fig. 5c). In Fig. 8 we indicate the
546 corresponding changes in the extent of the basin of attraction with respect to changing Ω by re-
547 computing Fig. 8a with $\Omega = 0.3, 1, \text{ and } 10$. The two smaller Ω values (0.3 and 1) correspond
548 roughly to Rossby numbers $a/2\Omega$ of about 1 and 0.2, i.e., are representative of the ocean
549 submesoscale and mesoscale flows. The Q_a -functions for these cases are plotted in Fig. 6b-c.
550 Most submesoscale eddies are going to tend to have $u^{(\theta)}/r$ about the same magnitude as Ω
551 (except on the equator) and mesoscale eddies will have $u^{(\theta)}/r \ll \Omega$. The results in Fig. 8
552 suggest that, while the basin of attraction does shrink slightly with increasing Ω , this dependence
553 is weak. The main difference between the three numerical runs in Fig. 8 is in the associated
554 attraction time, which gets significantly shorter for larger values of Ω . This is explored in more
555 detail below.

556 Attraction time:

557 It follows from Eq. (6) that the attraction time towards the periodic orbit should scale as $T_a =$
558 $\left[2\tilde{\epsilon} \left(\frac{3R}{2} - 1\right) Q_a\right]^{-1}$ where $Q_a = \frac{1}{2}\left(\frac{1}{2}|\vec{\zeta}_a|^2 - |S|^2\right)$ with $\vec{\zeta}_a = \vec{\zeta}_r + \overline{2\Omega}$. Thus, for $\vec{\zeta}_r \geq 0$, as in
559 most of our numerical runs (except Experiment 1e), attraction time decreases with increasing Ω
560 for positive $\Omega \geq 0$. For negative $\vec{\zeta}_r$, which corresponds to the reversed direction of the flow in
561 our simulations (Experiment 1e), an increase in Ω will initially slow the attraction by decreasing
562 the magnitude of $\vec{\zeta}_a$ all the way to 0, at which point the periodic orbit will lose its attraction
563 properties, but then will speed up the attraction as Ω is further increased. This trend is confirmed
564 numerically in Fig. 9, where for the flow parameters corresponding to the “reversed flow” run in
565 Table 1 (Experiment 1e, with $\vec{\zeta}_r < 0$), we release a sample trajectory within the basin of

566 attraction and plot its z-coordinate as it winds around the can and eventually approaches the
567 attracting periodic orbit. As anticipated, the attraction time initially increases as Ω is increased
568 from 0 to 0.6, but then decreases as Ω is further increased to 2.

569 *Disappearance of the subsurface attractor when $\tilde{\varepsilon}$ becomes too large:*

570 Finally, to illustrate the disappearance of the subsurface attractor when $\tilde{\varepsilon}$ becomes too large, in
571 Fig. 10, we contrast 2 numerical simulations with the same flow parameters (corresponding to
572 the “slow overturn” run 1c in Table 1) but different particle diameters, $d = 10^{-3}$ vs $d = 5 \times$
573 10^{-4} . For larger d , the subsurface periodic orbit for rigid particles is no longer present within the
574 can, leading to all particles rising up to the surface (Fig. 10b). For smaller d , the periodic orbit is
575 still present and acts as an attractor for buoyant rigid particles over a significant portion of the
576 can (green region in Fig. 10a). We note that this run would be more qualitatively similar to the
577 oceanic mesoscale or submesoscale eddies, where the overturning component of circulation is
578 weak in comparison to the horizontal swirl.

579 (b) steady non-symmetrically perturbed case

580 We now consider a case in which the axial symmetry of the steady flow has been broken, here
581 through a change in the perturbation amplitude parameter ε from zero to 0.25, and in the offset
582 parameter y_0 from 0 to -0.2 in the Eqs. (9a,b). The fluid velocity field now contains something
583 like a stationary, “mode-1” azimuthal wave in the horizontal velocity field.

584 The resulting Lagrangian structure (Fig. 2d and e) has a sea of chaos that covers the near-axial
585 and outer regions of the cylinder, where no unbroken tori survive. Within this chaotic sea is a
586 region containing a nest of unbroken tori that surround a central periodic orbit. This orbit has
587 evolved from the central periodic orbit of the symmetry case and is now tilted. Within the nest of

588 unbroken tori there exist resonant layers, in which new tori have arisen, and the most prominent
589 is the “island” that is centered near $x = 0.4$ and $z = 0.2$ in the right-half (and near $x = 0.4$ and
590 $z = 0.2$ in the right half) of Fig. (2d). We further note that this center lies within the region of
591 positive Q_a (Fig. 6b). The island corresponds to the yellow tori in Fig. 3e and is produced by a
592 1: 1 resonance, so that the periodic trajectory running through its center executes one complete
593 azimuthal cycle and one overturning cycle before connecting back onto itself. Thus, in this
594 steady asymmetric configuration, we now have 2 periodic orbits of the fluid flow – the central
595 slightly-tilted periodic orbit near mid-depth (that evolved from the central horizontal periodic
596 orbit of the axisymmetric flow) and a new periodic orbit running through the center of the
597 resonant island (resulting from the break-up of the resonant torus satisfying $\Omega_{\bar{\theta}} + \Omega_{\phi} = 0$).

598 We speculate that for sufficiently small $\tilde{\epsilon}$ a periodic orbit for the rigid particle motion exists in
599 the vicinity of each of the 2 periodic orbits of the fluid flow. This conjecture is difficult to prove
600 due to a complex geometry, leading to centrifugal forces that act in different directions at
601 different locations along the particle path. For now we simply search for the supposed attractors
602 by releasing particles and following their trajectories.

603 As shown in Fig. 2f, separate attractors arise in the vicinity of two periodic orbits. The first
604 appears as a ring-like structure (purple core) lying near the center of the original nested tori and
605 the second is a similar feature with a red core near the center of the resonant island. The two are
606 chained together and each has its own basin of attraction (Fig. 7c): the first consisting of a
607 roughly elliptical patch (inner green region) in the x - z -plane, which corresponds of a slice
608 through a tube-like structure in 3D, and the second consisting on an annular (blue) region that
609 surrounds the green region and that occupies a relatively larger volume.

610 In order to check that attraction of slightly-buoyant rigid particles towards periodic orbits located
611 near the centers of the resonant islands in the perturbed flow is not limited to the case of the 1: 1
612 resonance, in an additional simulation (Fig. 11, experiment 2c in Table 1), we adjusted the
613 background flow parameter b in Eqs. (9), which is responsible for the overturning strength, to
614 create a 2: 1 resonance instead of a 1: 1 resonance, as in the original run. In this case, the
615 resonant torus breaks down giving rise to a 2-island chain on the corresponding Poincare section
616 (Fig. 11a), and the fluid periodic orbit that goes through the centers of both islands completes 2
617 full cycles in azimuth and 1 complete cycle in vertical before connecting onto itself. Also, as in
618 the original run, a second slightly-tilted periodic orbit still exists near mid-depth of the can.
619 When buoyant particles are released into this flow, two attractors arise, corresponding to the 2
620 periodic orbits of rigid particles – one near mid-depth (purple core in Fig. 11c) and another in red
621 near the center of the 2: 1 resonant island.

622 *Shift in position of the periodic orbit associated with a resonant island as a function of flow and*
623 *particle parameters, and frame rotation*

624 The position of the attracting periodic orbit for rigid particles that is located within the resonant
625 islands (we will refer to it as the resonant periodic orbit) in the asymmetrically-perturbed flow
626 depends both on the perturbation strength (via ε), on the flow and particle parameters (via $\tilde{\varepsilon}$), and
627 on the frame rotation Ω . Specifically, this resonant periodic orbit for the rigid particles will shift
628 away from the corresponding periodic trajectory of the fluid flow as $\tilde{\varepsilon}$ and Ω are increased. The
629 same is true for the slightly-tilted central attracting periodic orbit near mid-depth. This is
630 qualitatively similar to the shifting of the central periodic orbit up and out from $z = 0.5$,
631 $r = 0.34$ in the axisymmetric flow in response to changing $\tilde{\varepsilon}$ and Ω , which we explored in detail
632 the previous section both analytically (Eqs. (12)) and numerically (Fig. 3-5).

633 In order to numerically illustrate the shift in the position of the attracting periodic orbits, we
634 present (Figs. 12 and 13) numerical simulations in the steady perturbed flow configuration for 3
635 values of d (and thus $\tilde{\varepsilon}$) and 3 values of Ω . As both parameters increase, the attractors move
636 away from the corresponding periodic orbits of the fluid flow. This shift is evident from the
637 change in the color of the attraction basins in (a,d,g) and from the location of the yellow cloud of
638 dots in (c,f,i) in Figs. 12-13. Increases in $\tilde{\varepsilon}$ and Ω also lead to the shrinkage of the attraction
639 basins for both attractors and to a faster convergence rate, as is evident from the tighter cloud of
640 yellow dots in (c,f,i), as discussed in more detail below. The basin of attraction for the central
641 attractor – the green region in Fig. 12 – seems to shrink faster than the basin of attraction for the
642 resonant attractor (the blue-ish region) as d increases, so when d is increased from 2×10^{-3} to
643 3×10^{-3} , the central attractor vanishes, whereas the resonant attractor is still present (Fig. 12g).
644 On the other hand, the increase in Ω (Fig. 13) causes a faster shrinkage of the basin of attraction
645 for the resonant attractor than for the central attractor, so when Ω is increased from 2 to 5 in Fig.
646 13g, the resonant attractor disappears, whereas the central attractor is still present. Figs. 12g,h,i
647 (and Fig. 13g,h,i) show cases where this threshold has been exceeded, and one of the attractors
648 has been lost, whereas the other is still present.

649 Attraction time:

650 Similar to the unperturbed flow, the attraction time for attractors in the steady, perturbed flow
651 may still scale as $T_a = \left[2\tilde{\varepsilon} \left(\frac{3R}{2} - 1 \right) Q_a \right]^{-1}$, provided that Q_a is regarded as a typical value
652 within the corresponding basin of attraction. The predicted decrease in attraction time with
653 increasing $\tilde{\varepsilon}$ and Q_a is evident from the numerical simulations in Figs. 12-13, where in (c,f,i) we
654 color-coded trajectory crossings with the x-z Poincare plain by time, with blue/yellow
655 corresponding to initial/final time. For smaller values of $\tilde{\varepsilon}$ and Ω , we observe a wider and more

656 diffuse cloud of dots (because trajectories wind around the can many times before approaching
 657 the attractor), whereas as $\tilde{\epsilon}$ and Ω increase, the clouds at comparable times become denser and
 658 more compact around the attractors.

659 Basin of attraction

660 For the slightly-tilted central periodic orbit located within the central non-chaotic region near
 661 mid-depth in Fig. 2f, we observe that the basin of attraction – green region in Fig. 7b – extends
 662 roughly from the location of the periodic orbit to the edge of the central non-chaotic region (that
 663 is foliated by discretely sampled closed curves in Fig. 2d). Note that as $\tilde{\epsilon}$ increases, the attracting
 664 periodic orbit moves away from the center of this non-chaotic region towards its edge, leading to
 665 the shrinkage and eventual disappearance of the corresponding basin of attraction, shown by the
 666 green regions in Fig. 12a,d,g).

667 Similarly, in all of our numerical simulations, we observe that for the resonant attracting periodic
 668 orbit running through the resonant islands, the basin of attraction seems to cover the region
 669 between the orbit and the edge of the corresponding resonant island. An analytical expression for
 670 the width of the (non-degenerate) resonant island in the fluid flow (Pratt et al., 2014) predicts

671 that $\Delta I = \sqrt{\frac{\epsilon F_{nm}^0(I_0)}{\left(n \frac{d^j \Omega_\phi}{dI^j} + m \frac{d^j \Omega_\theta}{dI^j}\right)_{I_0}}}$, where ΔI is the deviation in the action coordinate away from I_0 , the

672 value of action at the resonant torus (i.e., at the center of the island). This width depends on the
 673 strength of the perturbation ϵ , the order of the resonance (via n and m in the resonance

674 condition), the background flow (via $\frac{d^j \Omega_\phi / \theta}{dI^j}$), and the structure of the perturbation (via $F_{nm}^0(I_0)$).

675 This expression could be used as an upper limit on the extent of the basin of attraction. However,
 676 because the attracting periodic orbit will move away from the center of the island towards its

677 edge as $\tilde{\epsilon}$ and Ω increase, the basin of attraction for the resonant attractor (blue region in Figs.
678 12a,d and 13a,d) becomes increasingly smaller than ΔI . One might speculate, then, that the
679 attractor will completely disappear when the attracting periodic orbit reaches the edge of the
680 resonant island. This is the case in Figs. 13g where the resonant attractor is no longer present.

681 (c) non-steady, non-symmetrically perturbed case

682 The final case that we will consider is one in which the perturbation is asymmetric and varies
683 periodically in time. The chosen perturbation frequency, $\sigma = 2\pi/9.1$, causes 2 strong additional
684 resonances (compared to the steady perturbed case) – one with $n = 0$, $m = 1$, and $l = 1$ (i.e.,
685 with a torus whose overturning frequency is equal to the perturbation frequency) that is shown in
686 blue in Fig. 2g,h and is located near the outer edge of the central non-chaotic region, and another
687 resonance, shown in green in Fig. 2g,h, with $n = 1$, $m = 1$, and $l = 1$, which is located between
688 the central non-chaotic region and the larger $n = 1$, $m = 1$ resonant island (that was present in
689 the steady case as well). Both of these new resonant structures are time dependent, their shape
690 and position recurring periodically. For example, the blue island, which looks like a crescent
691 moon pointing upward on the Poincare section at $t = 0$, becomes a crescent moon pointing
692 downward at time 4.55. The movement of the green island is more complex, as it turns both in
693 azimuth and vertical, making one complete loop over 9.1 time units. Because of the time-
694 dependence, trajectories must be strobed at the forcing frequency σ in order to capture
695 ‘snapshots’ of their forms as they recur at a particular phase in the time cycle. At the center of
696 each feature is a closed material curve that also varies periodically. Where the island has
697 emerged from the breakup of a torus with quasiperiodic orbits, the individual trajectories that
698 populate the material curves are themselves quasiperiodic.

699 Particle trajectory computations in this case confirm that the purple, red and green islands give
700 rise to attractors (Fig. 3i), whereas the blue island does not. In fact, slightly-buoyant rigid
701 particles that are released in the blue region converge towards the attractor that lies near the
702 purple region. This is also indicated by the basin of attraction of the central attractor extending
703 across the space occupied by the blue resonant island in Fig. 7c.

704 **IV. Discussion**

705 We have considered attraction phenomena for small, finite size, spherical, buoyant, rigid
706 particles in a three-dimensional rotating cylinder flow with azimuthal rotation and overturning,
707 and both with or without time dependence. The aim has been to gain insights into the behavior of
708 slightly buoyant microplastic particles in 3D vortex flows that qualitatively resemble ocean
709 eddies. The rigid particle motion is governed by a simplified version of the Maxey-Riley
710 equations (accounting for inertia, buoyancy and simplified quantification of drag and added
711 mass), and, approximately, by the slow-manifold reduction of these equations. We have
712 illustrated the possibility of aggregation of slightly-buoyant rigid particles in 3D vortex flows
713 towards closed loop attractors located subsurface within the interior of the flow. Even in our
714 idealized flow and for spherical particles with fixed radius and buoyancy, aggregation is non-
715 trivial, often with multiple attractors present and/or the lack of attraction in some circumstances.

716 Our rotating cylinder model is much less complex than any real ocean eddy in many respects,
717 including the assumed quasiperiodic time dependence and the absence of decay and interaction
718 with the surroundings. Understanding aggregation in a simple periodic flow seems like a
719 reasonable first step towards understanding aperiodic, interacting, and decaying oceanic eddies.
720 This approach is common in applications of dynamical systems theory to oceanography and

721 meteorology. For example, arguments relating to the increased stability of jets due to the strong
722 Kolmogorov-Arnold-Moser stability near shearless trajectories have first been developed for
723 spatially-periodic and time-quasiperiodic flows and tested using idealized toy models, before
724 exploring these ideas in more realistic oceanic and atmospheric settings (see Rypina et al., 2007
725 and Beron et al., 2008; 2010). Note also that our results are applicable to quasi-periodic
726 disturbances with finite number of frequencies rather than just periodic disturbances (we only
727 show numerical simulations for the time-periodic case for simplicity), and a quasiperiodic
728 function might potentially be useful for approximating temporal variability in some oceanic
729 flows, especially those with pronounced peaks in the spectrum.

730 We have explored a steady axisymmetric rotating cylinder flow and a steady flow with its axial
731 symmetry broken. In all cases, we have observed emergence of subsurface attracting structures
732 that lead to the aggregation of buoyant particles towards them. We have linked these attractors to
733 the periodic orbits of rigid particles that exist in a region of net contraction of the phase space of
734 the particle motion. The slow manifold equations suggest that periodic orbits for rigid particles
735 exist near periodic orbits of the underlying fluid flow, provided the drag is sufficiently strong
736 (Stokes number $\ll 1$).

737 We have also explored one case of an axially asymmetric and time-periodic flow, with focus on
738 the resonant “islands” that arise due to the time-dependence. At the center of such islands are
739 closed material contours, or loops, composed of quasi-periodic orbits of the fluid flow. One such
740 structure has a nearby attractor, also a closed loop of quasiperiodic orbits for rigid particles,
741 while a second example does not. A detailed explanation awaits formulation of a quantitative
742 theory, something that is beyond the scope of the present paper and that will be presented in a
743 future work.

744 We have observed that the disappearance of an attractor, which can occur as the result of
745 increasing rigid particle size or frame rotation, coincides roughly with the displacement of the
746 position of the attractor to the outer edge of the resonant island from which it sprang. Whether
747 this purely geometric observation forms the basis for a general criterion for the loss of attraction
748 is unknown, as a dynamical justification is needed.

749 Marine microplastics can have complex non-spherical tangled-filament shapes, change their
750 physical and chemical properties in time due to aging and photo- or chemical-decay processes
751 (Andrady 2011), are subject to biofouling (see recent relevant work by Kreczak et al., 2021), and
752 may interact leading to the formation of clusters. None of these effects were considered in this
753 paper, and all will need to be taken into account for the realistic prediction of marine
754 microplastic evolution and re-distribution in the ocean. Real ocean eddies are also decaying in
755 time and are usually moving (translating) rather than stationary. Translation with a constant
756 velocity can be handled by considering the flow in a moving frame of reference, but decay and
757 interactions will likely change the geometry of the circulation and make the flow truly aperiodic.
758 Our simplified model cannot account for these effects, which will need to be explored separately
759 later.

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762 **Data Availability Statement:** No observational data was used. Details of the numerical
763 simulations using an analytical vortex model are provided in text.

764 **Author Contribution Statement:** IR led the overall effort and performed most of the numerical
765 simulations, LP contributed towards the theoretical understanding and interpretation of the
766 results, MD participated in the overall effort.

767 **Competing interests Statement:** no competing interests

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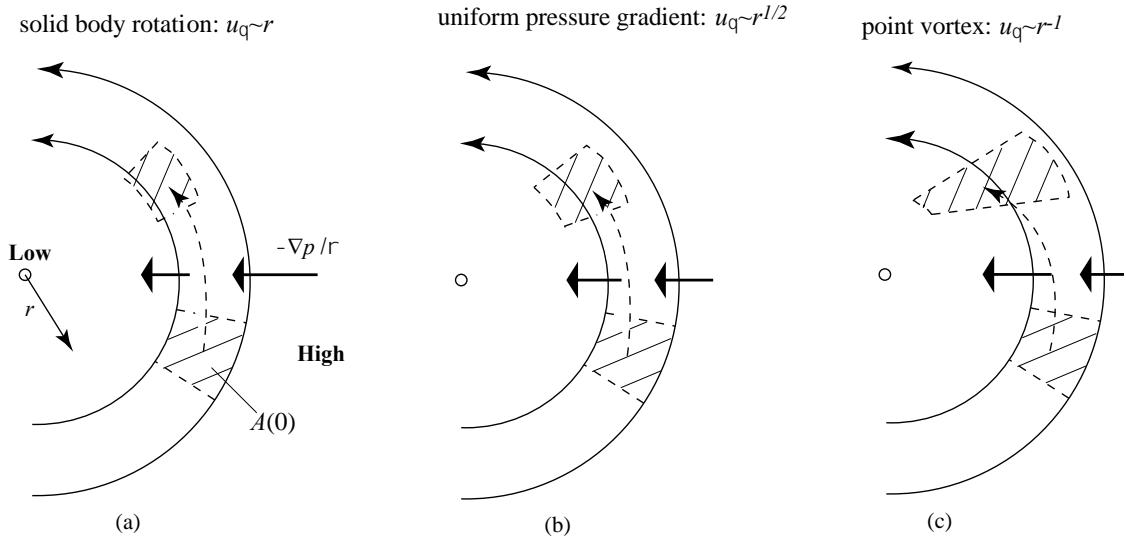
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Experiment	a	b	ε	y_o	σ	γ	β	Ω	d
1 – steady symmetric	0.62	7.5	0	0	0	0	0	0	10^{-3}
1a (small Ω)	0.62	7.5	0	0	0	0	0	0.3	10^{-3}
1b (large Ω)	0.62	7.5	0	0	0	0	0	1	10^{-3}
1c (slow overturn)	0.62	0.25	0	0	0	0	0	1	10^{-3} vs. 5×10^{-4}
1d ($z_{attractor}$ vs Ω)	0.62	7.5	0	0	0	0	0	Sweep 0 to 10	10^{-3}
1e (reversed flow)	-0.62	-7.5	0	0	0	0	0	0, 0.6, 2	10^{-3}
2 – steady asymmetric	0.62	7.5	0.25	-0.2	0	0	0	0	10^{-3}
2a (small Ω)	0.62	7.5	0.25	-0.2	0	0	0	0.3	10^{-3}
2b (large Ω)	0.62	7.5	0.25	-0.2	0	0	0	1	10^{-3}
2c (2:1 resonance)	0.62	3.8	0.25	-0.2	0	0	0	0	10^{-3}
3 - non-steady asymmetric	0.62	7.5	0.25	-0.2	$\frac{2\pi}{9.1}$	0.2	1	0	10^{-3}

924 Table 1: Dimensionless parameter values for numerical experiments. Fixed parameters in the
925 kinematic model (Eqs. 9a-c) are $c = 0.69$, and $r_0 = 1/2$ in all cases. Parameters that appear in
926 the nondimensional Maxey-Riley Eq. (3) are also nondimensional, with L , U , L/U as length,
927 velocity and time scales. Fixed parameter values based on $L = 1$ m and $U = 1$ m/s include
928 $\frac{\rho_p}{\rho_f} = 0.97$, $R = \frac{2\rho_f}{\rho_f + 2\rho_p} = 0.680$, $\frac{3R}{2} - 1 = .020$, $\vec{g}_r = \frac{gL}{U^2} = 10.0$, $\tilde{\varepsilon} = \frac{2}{9} \left(\frac{d}{L}\right)^2 \frac{UL}{\nu R} = 0.33$, and
929 $\tilde{\varepsilon} \left(\frac{3R}{2} - 1\right) = 0.0067$. Note that $\vec{\Omega} = \Omega \vec{k} = \frac{\vec{\Omega}^* L}{U}$. In dimensional units, our parameters
930 correspond to a 1 mm (or 0.5 mm in some simulations) particle in a rotating cylinder with a
931 diameter of 1 m.



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933 Figure 1. Three types of two-dimensional eddies with zero frame rotation and for which gravity
 934 is imagined to be zero: solid body rotation (a), constant pressure gradient (b), and point vortex
 935 (c). In each case, the cross hatched area represents a concentration of rigid particles with area
 936 $A(t)$.

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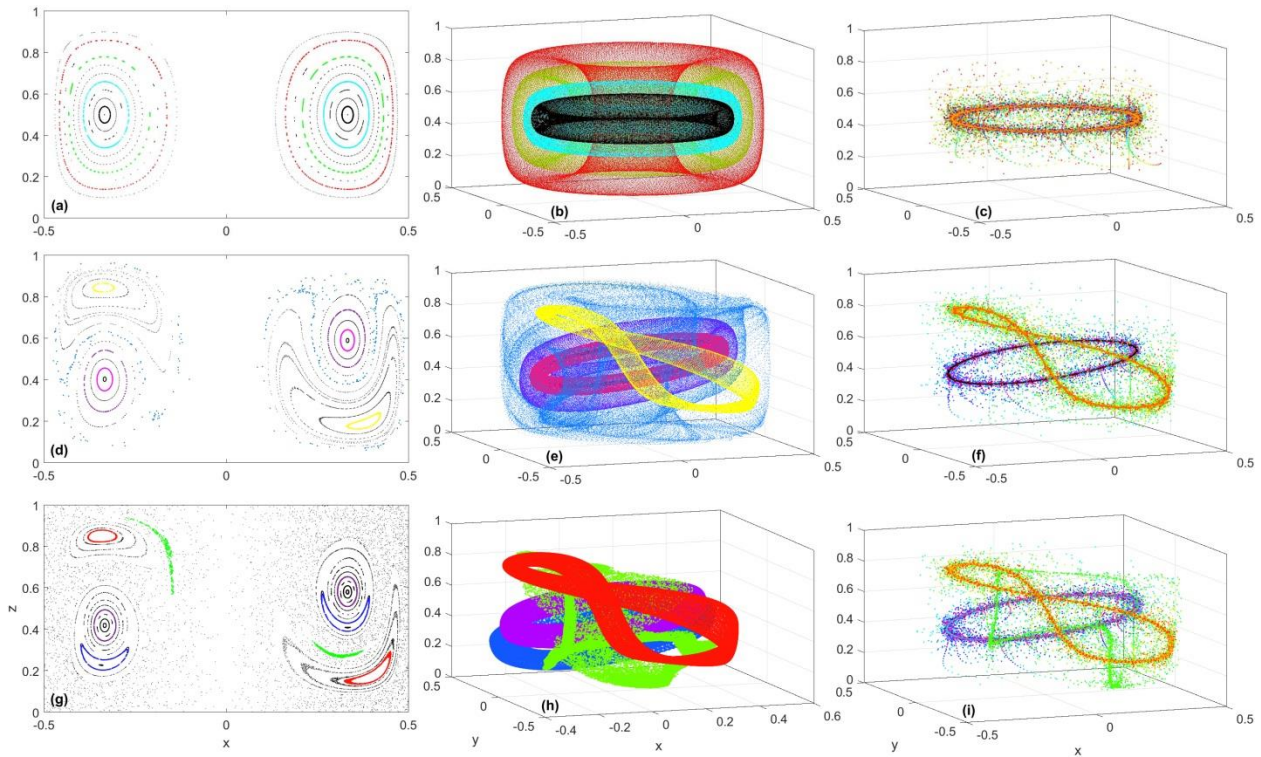
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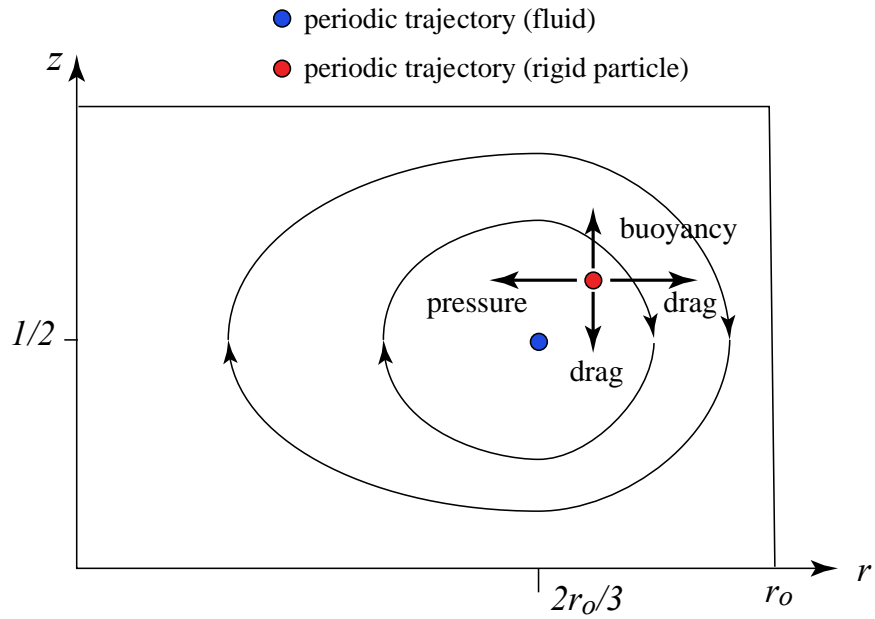
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946 Figure 2. (left) Poincare section, (middle) fluid parcels trajectories in 3D, (right) buoyant particle
 947 trajectories in 3D for a steady symmetric fluid flow (top row), steady asymmetric flow (middle
 948 row), and non-steady, asymmetric flow. Parameter setting are listed under Experiments 1, 2 and
 949 3 in Table 1. Colors in the left column of panels match the corresponding panel in the middle
 950 column, but the colors in the right column indicated time after release of the particles. Note the
 951 attraction of buoyant particles to a single attractor at mid-depth in panel (c), to 2 attractors in
 952 panel (f), and to 3 attractors in panel (i). Particles are released along a vertical line $x = 0.334$,
 953 $y = 0$, $0 < z \leq 0.6$ with initial velocity equal to that of the co-located fluid parcels.

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958 Figure 3. Sketch showing the position in a vertical section of the periodic orbit (red dot) of the
 959 rigid particle relative to the periodic orbit (blue dot) of the fluid flow. The viewer sees one half
 960 of a vertical slide through the cylinder, with the azimuthal flow directed away from the viewer
 961 and the cylinder center at the left edge.

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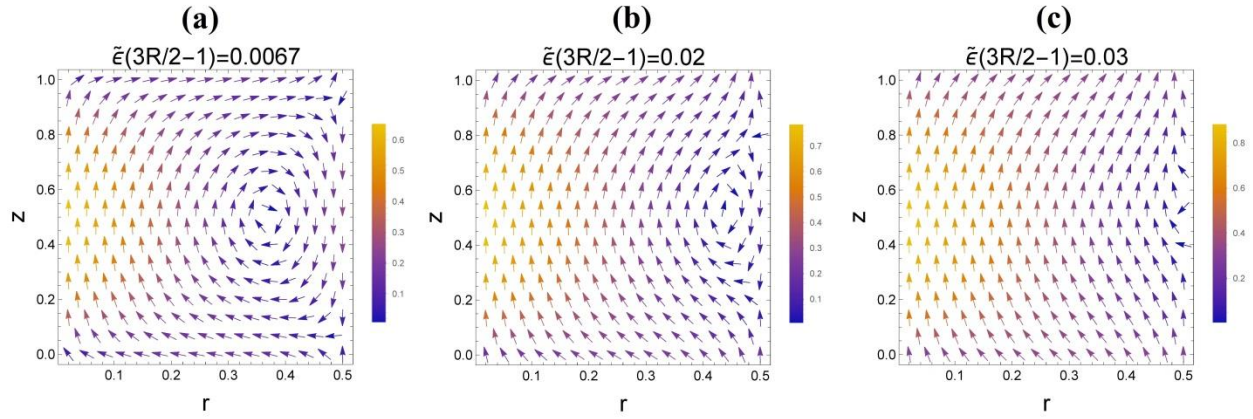
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970 Figure 4. The slow-manifold radial and vertical velocity components for the rigid particles,

971 plotted in the (r, z) plane for (a) $\tilde{\varepsilon} \left(\frac{3R}{2} - 1 \right) = 0.0067$, (b) = 0.02, and (c) = 0.03. Other

972 parameters are as listed for Experiment 1a in Table 1.

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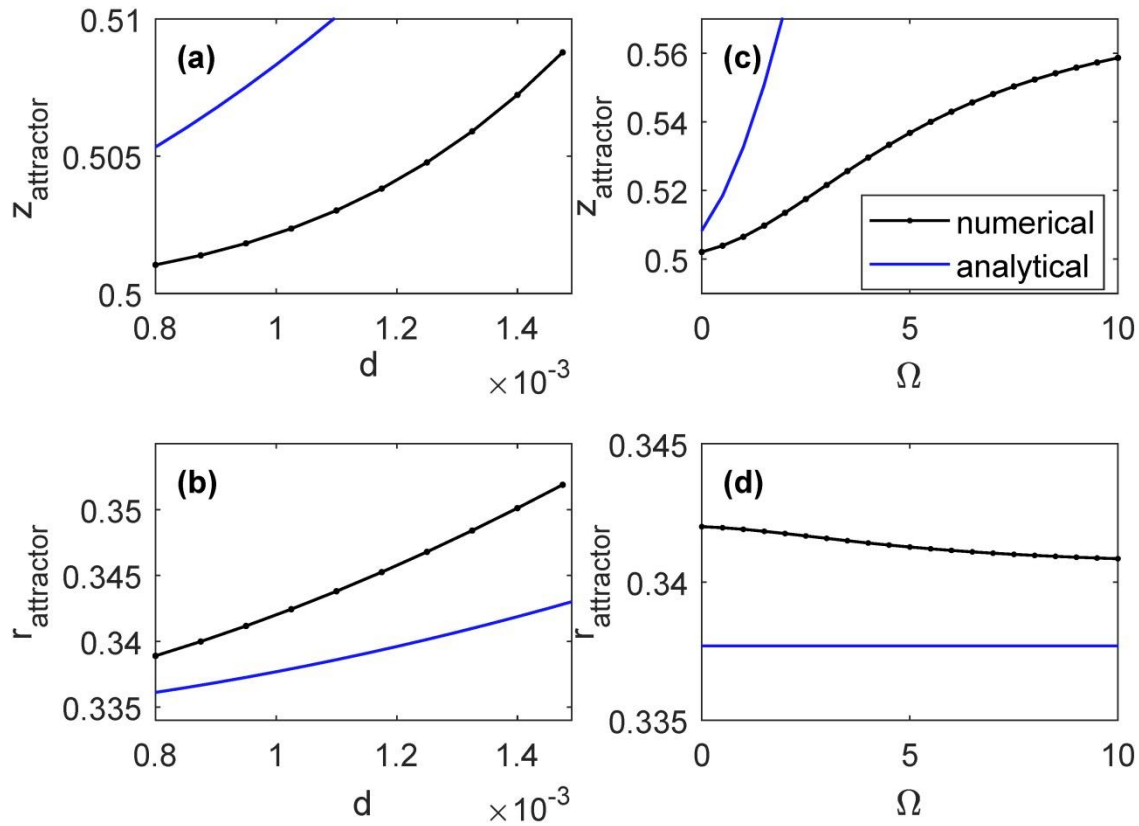
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984 Figure 5. For the steady symmetric rotating cylinder flow, the coordinates of the periodic orbit
 985 that acts as an attractor for buoyant particles as a function of particle diameter (a-b) and frame
 986 rotation (c). Flow parameters are listed in Table 1 and correspond to Experiment 1 for (a-b) and
 987 Experiment 1d for (c-d).

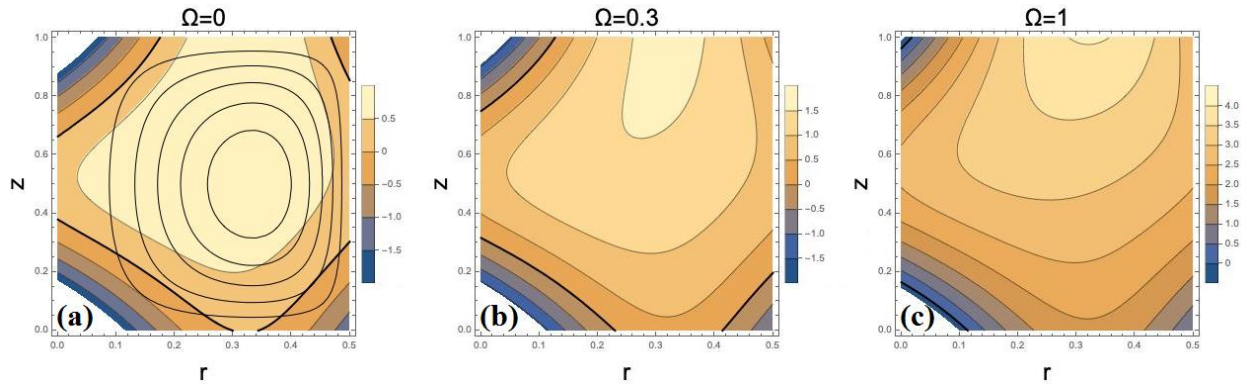
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994 Figure 6. (a): The Q_a function for the steady, axisymmetric, cylinder flow with the same
 995 parameter setting (see Experiment 1a) as for Figure 3a-c, and plotted in (x,z) along with the
 996 streamlines of the overturning circulation. The thick rigid curve corresponds to $Q_a = 0$. (b): The
 997 same parameter settings, except Ω has been raised from 0 to 0.3 (Rossby number $\cong 1$) (c):
 998 $\Omega = 1.0$. (Rossby number $\cong 0.2$).

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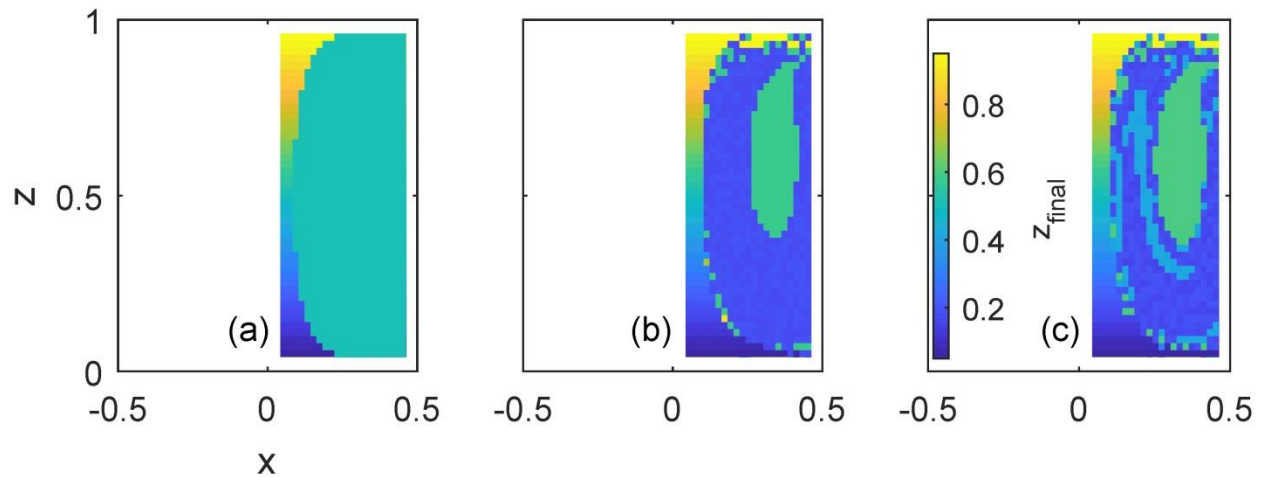
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1009 Figure 7. Domain of attraction for the attractors in (a) steady symmetric (Experiment 1 in Table
 1010 1), (b) steady asymmetric (Experiment 2 in Table 1), and (c) time-periodic asymmetric rotating
 1011 cylinder flow (Experiment 3 in Table 1). (These are the same 3 experiments that were used to
 1012 produce Fig. 2.) The color indicates the height (i.e., value of z-coordinate) of the final crossing of
 1013 a trajectory with the Poincare section, as a function of particle's release location. Particles
 1014 attracted to the same attractor thus correspond to same color.

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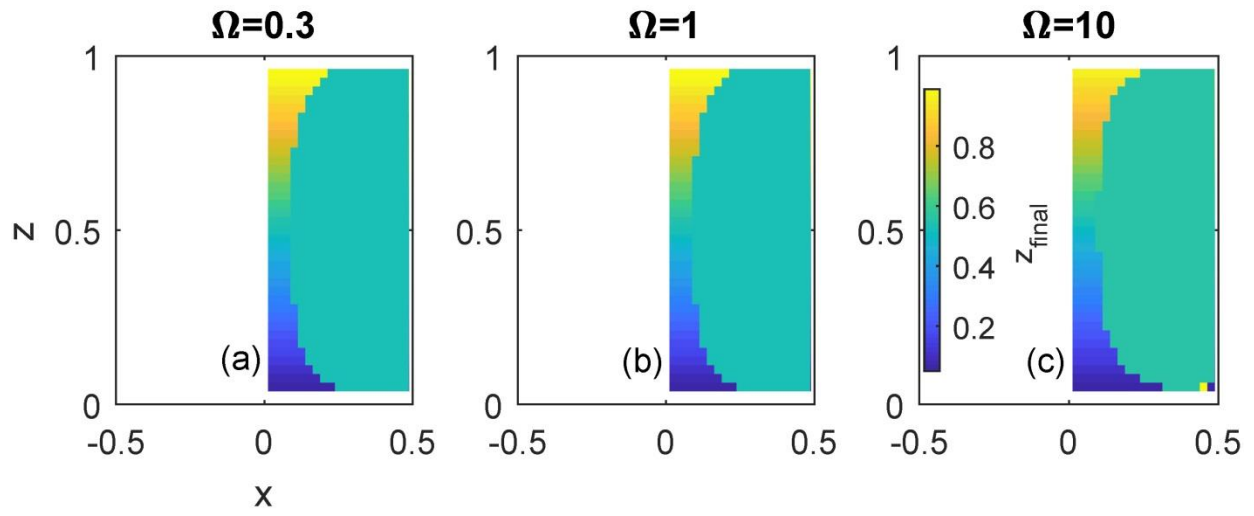
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1023 Figure 8. Same as in Fig. 7a but with frame rotation.

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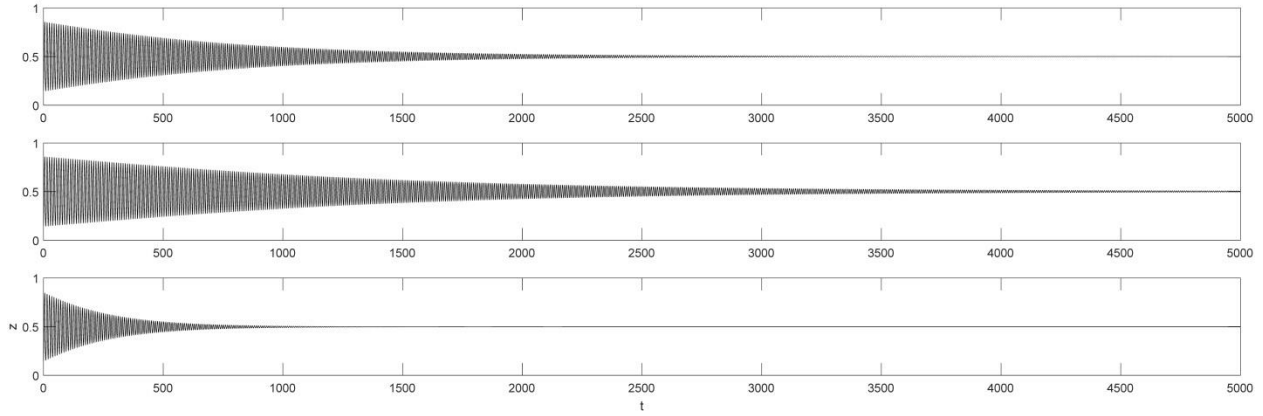
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1035 Figure 9. For the “reversed flow” experiment (Experiment 1e in Table 1), z-position of a sample
 1036 particle trajectory as function of time for 3 values of Ω : 0 (top), 0.6 (middle), and 2 (bottom).

1037 Time t is in dimensionless units (but since our scaling coefficient for time is equal to 1 sec, the
 1038 numbers on the x-axis can also be read as dimensional time in sec.)

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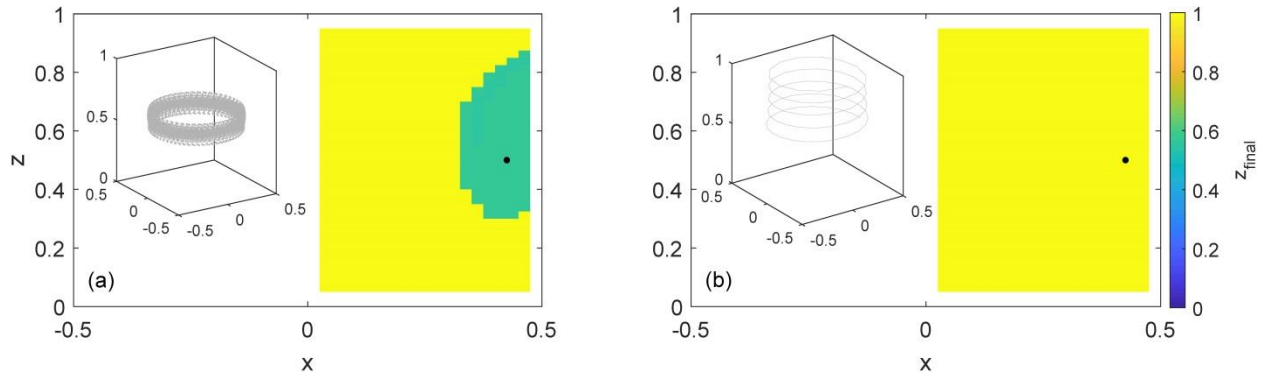
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1049 Figure 10. For the “slow overturn” Experiment 1c from Table 1, color indicates the final z-
 1050 coordinate of a particle’s trajectory at the end of integration time as a function of particle’s
 1051 release location for 2 values of d : (a) 5×10^{-4} and (b) 10^{-3} . Yellow corresponds to particles
 1052 rising up to the top, whereas green indicates the basin of attraction of the subsurface attracting
 1053 periodic orbit. The insets at the left side of each frame show a sample trajectory whose release
 1054 location is indicated by the black dot.

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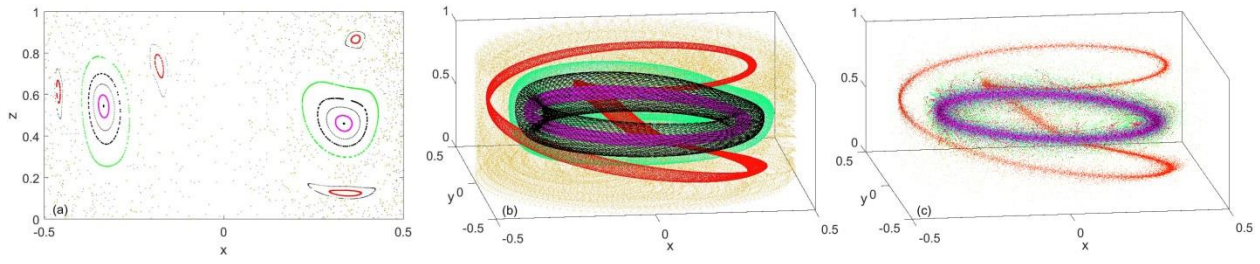
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1065 Figure 11. Same as Fig. 2(d-f) but with $b = 3.8$.

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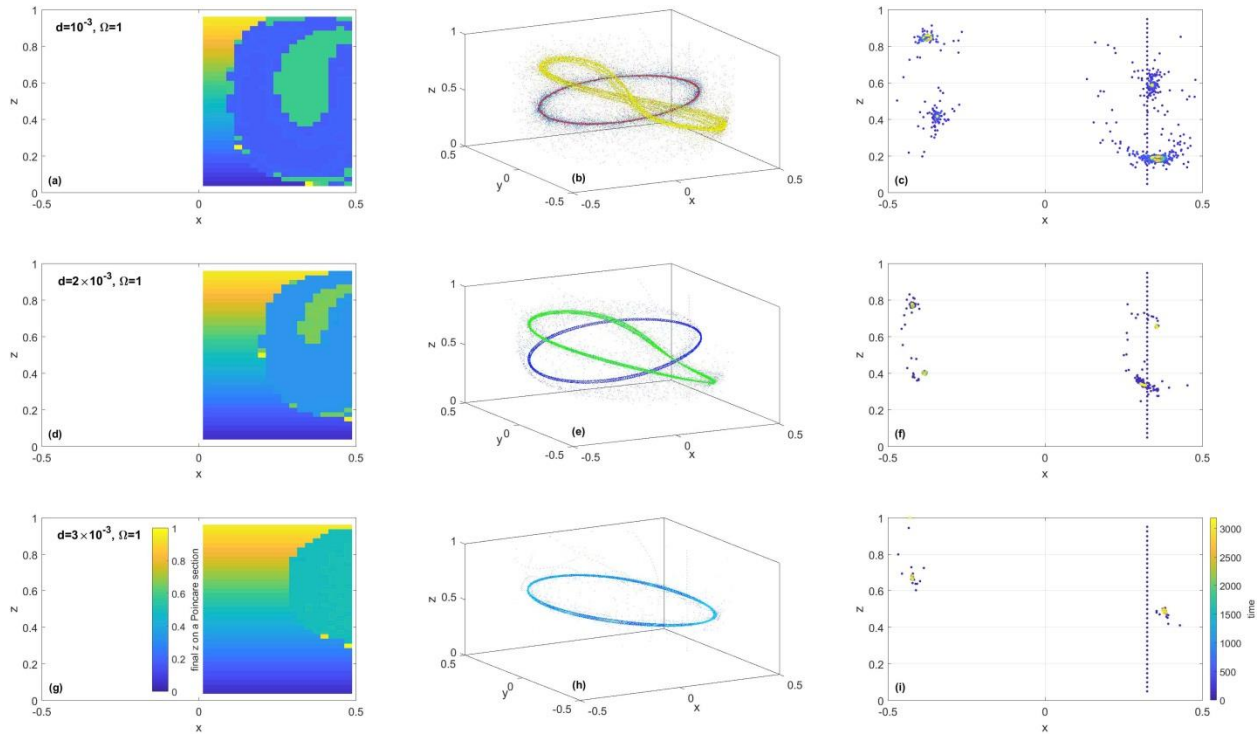
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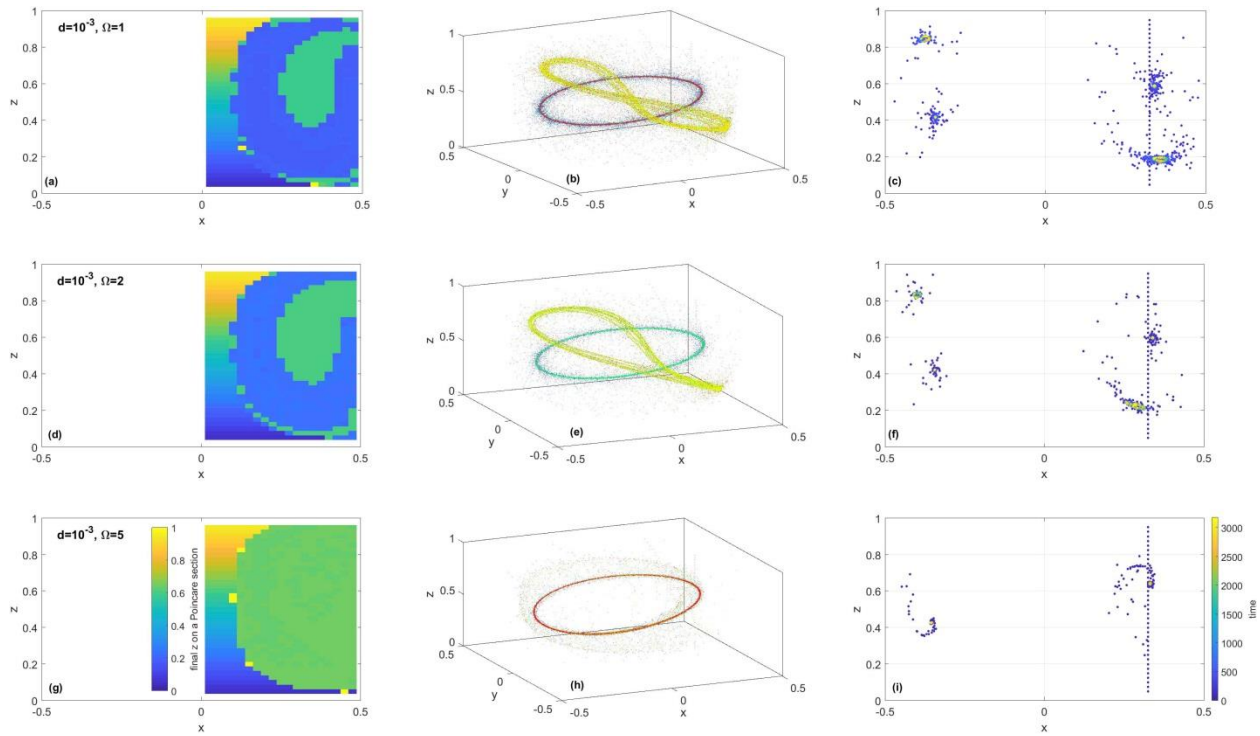
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1079 Figure 12. For the steady perturbed system (Experiment 2 in Table 1), changes in the location of
 1080 the attracting periodic orbits, basins of attractions, and time of attraction as a function of particle
 1081 diameter d (and thus $\tilde{\epsilon}$). (a,d,g) show z -coordinate of the last crossing of trajectory with the x - z
 1082 Poincare plane as a function of release location; flat regions are basins of attraction for the 2
 1083 attractors. (b,e,h) show 20 trajectories in 3d released along a vertical line at $y = 0, x = 0.334,$
 1084 $0.05 < z < 0.95$; denser cores indicate attractors. (c,f,i) show crossing of the same select 20
 1085 trajectories with the x - z Poincare plane, color coded by time; blue corresponds to release
 1086 location, yellow corresponds to final positions.

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1091 Figure 13. For the steady perturbed system (Experiment 2 in Table 1), changes in the location of
 1092 the attracting periodic orbits, basins of attractions, and time of attraction as a function of frame
 1093 rotation Ω . (a,d,g) show z-coordinate of the last crossing of trajectory with the x-z Poincare plane
 1094 as a function of release location; flat regions are basins of attraction for the 2 attractors. (b,e,h)
 1095 show 20 select trajectories in 3d released along a vertical line at $y = 0, x = 0.334, 0.05 < z <$
 1096 0.95 ; denser cores indicate attractors. (c,f,i) show crossing of the same 20 trajectories with the x-
 1097 z Poincare plane, color coded by time; blue corresponds to release location, yellow corresponds
 1098 to final positions.

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