

1           **Aggregation of Slightly Buoyant Microplastics in Three-Dimensional Vortex Flows**

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19 **Abstract**

20 Although the movement and aggregation of microplastics at the ocean surface has been well  
21 studied, less is known about the subsurface. Within the Maxey-Riley framework governing the  
22 movement of small rigid spheres with high drag in fluid, aggregation of buoyant particles is  
23 encouraged in vorticity-dominated regions. We explore this process in an idealized model [that is](#)  
24 [qualitatively reminiscent](#) of a three-dimensional eddy with an azimuthal and overturning  
25 circulation. In the axially symmetric state, buoyant spherical particles that do not accumulate at  
26 the top boundary are attracted to a [loop](#) consisting of periodic orbits. Such a [loop](#) exists when  
27 drag on the particle is sufficiently strong. For small slightly-buoyant particles, this [loop](#) is located  
28 close to the periodic fluid parcel trajectory. If the symmetric flow is perturbed by a symmetry-  
29 breaking disturbance, additional attractors for small rigid slightly-buoyant particles may arise  
30 near periodic orbits of fluid parcels within the resonance zones created by the disturbance.  
31 Disturbances with periodic or quasi-periodic time dependence may produce even more attractors,  
32 with a shape and location that recurs periodically. However, not all such [loops](#) attract, and rigid  
33 particles released in the vicinity of one loop may instead be attracted to a nearby attractor.  
34 Examples are presented along with mappings of the respective basins of attraction.

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40 **Significance statement**

41 This paper investigates aggregation of small, [spherical](#), slightly buoyant, rigid particles in a  
42 [simple](#) three-dimensional vortex flow. Our goal was to gain insights into the behaviour of  
43 slightly buoyant marine microplastics in a flow that qualitatively resembles ocean eddies.  
44 Attractors are mapped out for the steady axisymmetric, steady asymmetric, and non-steady  
45 asymmetric vortices over a range of flow and particle parameters. Simple theoretical arguments  
46 are used to interpret the results.

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58 **I. Introduction**

59 Marine microplastic pollution has been a rising concern for the ocean environmental and for  
60 human health. Microplastics (scales < 5mm) and nanoplastics (scales < 1 µm) have been found  
61 in the tissues of marine animals, some of which are consumed by humans (Landrigan, et al.  
62 2023). This comes at a time when global production of plastics is projected to increase.

63 Observations of marine microplastics have been conventionally carried out using net tows and  
64 mostly occurred at or near the sea surface (van Sebille et al., 2015). However, the density of  
65 many types of microplastic particles, including high-density polyethylene, is sufficiently close to  
66 that of sea water that suspension within the water column for long periods of time is feasible. For  
67 the near-surface microplastics, Kukulka et al. (2010) and Kooi et al. (2016) present observational  
68 evidence for the fast decay in concentrations with depth over the top 5 – 20 m of the water  
69 column, with the vertical penetration of plastic particles dependent on the wind speed.  
70 Pabortsava and Lampitt (2020), on the other hand, show observational evidence for much deeper,  
71 below-the-mixed-layer subsurface peaks for three common types of microplastics in the Atlantic  
72 Ocean. Processes such as biofouling and bio-geo-chemical or photo degradation might increase  
73 the density of the plastic particles and eventually lead to the sinking of microplastics from the  
74 surface into the deeper part of the water column (Kaiser et al., 2017; Kreczak et al., 2021; Kvale  
75 et al., 2020). Consumption by biomass with the subsequent downward vertical transport is  
76 another vehicle for redistributing microplastics from the surface down. For example, Choy et al.  
77 (2019) suggest that this mechanism, specifically, consumption by pelagic red crabs and giant  
78 larvaceans, was responsible for the subsurface peaks in plastic particles concentrations observed  
79 at depths near 250 m in Monterey Bay. Thus, microplastics have been found well beneath the

80 | [ocean surface, but less is known regarding their spatio-temporal and size/density distributions](#)  
81 | [\(Shamskhany et al., 2021\)](#).

82 | A potentially important aspect of the movement of plastics and microplastics is aggregation, a  
83 | process that occurs at the surface over large scales near the centers of the five major subtropical  
84 | gyres and has been attributed to Ekman drift, windage and inertia (Beron-Vera, 2021). Many  
85 | early models concentrated on the ocean surface, but [Wichmann et al. \(2019\)](#) has highlighted the  
86 | importance of resolving the full three dimensional circulation. If aggregation also occurs below  
87 | the surface, well beneath the direct influence of Ekman layers, the dynamics is likely to be  
88 | different. Indeed, modeling results by Wichmann et al. (2019), based on a framework created by  
89 | Lange and van Sebille (2017) and Delandmeter and van Sebille (2019), suggests that the large  
90 | scale accumulation associated with the garbage patches disappears below 60m depth.

91 | [To avoid confusion, we will refer to infinitesimal fluid elements as “fluid parcels”, and to rigid](#)  
92 | [plastic particles of finite size as “rigid particles”](#). Typically the position  $x_p(t)$  of a [rigid particle](#)  
93 | is tracked according to

$$94 | \quad x_p(t + \Delta t) = x_p(t) + \int_t^{t+\Delta t} u dt + dx_b,$$

95 | where  $u$  is the fluid velocity and  $dx_b$  is an extra displacement due the non-fluid nature of the rigid  
96 | particle. The user can introduce custom schemes for calculating contributions to  $dx_b$  due to  
97 | factors such as windage and inertia (e.g. Beron-Vera et al., 2016), turbulent diffusion (e.g.  
98 | Kulkulka, 2012), wave induced Stokes drift (Onink et al., 2019), etc. Eulerian schemes in which  
99 | plastic particles are treated as concentrations, are rare, but Mountford and Morales Maqueda  
100 | (2019) developed an Eulerian model in which concentrations are advected by the fluid and are  
101 | subject to parameterized turbulence as well as sinking or rising according to a simple law

102 involving buoyancy and friction. [In a similar fashion, Kvale et al. \(2020\) propose an Eulerian](#)  
103 [model for the biological uptake and the resulting re-distribution of microplastics.](#)

104 An alternative approach would be to use the Maxey-Riley equation (discussed below) to solve  
105 for the rigid particle velocity,  $v$ , and then use the latter to compute the trajectory of that rigid  
106 particle, i.e.,  $x_p(t + \Delta t) = x_p(t) + \int_t^{t+\Delta t} v dt$ . This equation would account [for the non-fluid-](#)  
107 [following effects](#) in a deductive way, however the resulting 6<sup>th</sup>-order system (for the three  
108 components of velocity and position) would be computationally challenging. To better  
109 understand the implications of the use of this approach while avoiding the computational burden  
110 and complexity, [we have elected to analyze the movement and aggregation of individual rigid](#)  
111 [particles using a Maxey-Riley framework in connection with an idealized, analytically-](#)  
112 [prescribed, 3D vortex flow that qualitatively resembles the geometry of the circulation in an](#)  
113 [ocean eddy but is not a solution to any dynamical oceanographic equations of motion. As shown](#)  
114 [by Pratt \(2014\) and Rypina et al. \(2015\), kinematic models that reproduce the correct geometry](#)  
115 [are able to also reproduce the important Lagrangian features of the flow. Even in our simple](#)  
116 [flow, aggregation is non-trivial, often with multiple attractors present and lack of attraction in](#)  
117 [some circumstances. Thus, we wanted to thoroughly explore this simple example before](#)  
118 [investigating more realistic oceanic flows.](#) We note that other idealized studies have been carried  
119 out in connection with 2D wave fields and vortex flows (e.g. DiBenedetto 2018a,b and Kelly et  
120 al., 2021).

121 Aggregation can be attributed to the presence of an attractor: here, an object with a dimension  
122 less than three that is somehow set up by the fluid circulation patterns and towards which rigid  
123 particle trajectories attract. As long as the fluid is incompressible, fluid parcels will not  
124 experience attraction and will not aggregate, but plastic particles ~~with inertia, added mass, and~~

125 ~~drag~~ may do so. Note also that because each attractor is generally associated with its  
126 corresponding basin of attraction, if rigid particles are introduced outside of the basin of  
127 attraction, they will not be attracted and will not aggregate towards this attractor.  
128 In order to reach a better understanding of what leads to attraction and attractors in 3D flows, we  
129 explore a simple canonical example in geophysical fluid dynamics, namely the flow in a rotating  
130 cylinder. This flow resembles some of the characteristics of ocean eddies, including a horizontal  
131 swirl and an overturning component in the vertical, but is much less complex than any realistic  
132 oceanic eddy. Specifically, we use a simple analytically-prescribed phenomenological velocity  
133 introduced by Rypina et al. 2015. The Lagrangian properties of this circulation have been  
134 previously studied (Fountain, et al. 2000; Pratt et al. 2014; Rypina et al. 2015) allowing us to  
135 begin to investigate inertial rigid particles from an established base of knowledge. A prior theory  
136 (Haller and Sapsis, 2008) governing the movement of rigid particles with high drag indicates that  
137 accumulation is favored for slightly buoyant particles in flows dominated by vorticity, and this  
138 also motivates our choice of background flow. Identification of the attractors that can arise in this  
139 flow field, evaluating their reach and domains of attraction, and clarifying the circumstances that  
140 lead to their formation are the primary objectives of this work. Although motivated by the  
141 problem of marine microplastics, this study is, for now, mainly a curiosity-driven research  
142 aiming to develop a basic understanding of the mechanisms that might lead to aggregation of  
143 rigid particles in 3D flows. The hope is that with such basic understanding in hand, one could  
144 later start investigating aggregation phenomena in more complex and more realistic ocean  
145 mesoscale and submesoscale eddying flows.

## 146 **II. Methods**

147 The physics of the motion of a small, rigid sphere that moves with velocity  $\vec{v}(t)$  through a fluid  
 148 with pre-existing velocity distribution  $\vec{u}(\vec{x}, t)$  has been the subject of investigation by Stokes  
 149 (1851), Basset (1888), Boussinesq (1903), Faxen (1922), Oseen (1927), Tchen (1947) and many  
 150 others, and was put in a unifying framework by Maxey and Riley (1983). More recent  
 151 theoretical extensions include Beron-Vera et al. (2019) and Beron-Vera (2021). We will use a  
 152 form of the Maxey-Riley equation that has been extended to include constant frame rotation with  
 153 angular velocity  $\vec{\Omega}^*$ :

$$154 \quad \frac{d\vec{v}}{dt} = \frac{\rho_f}{\rho_p} \frac{D\vec{u}}{Dt} + \frac{\rho_f}{2\rho_p} \left( \frac{D\vec{u}}{Dt} - \frac{d\vec{v}}{dt} \right) - \frac{9\nu\rho_f}{2\rho_p d^2} (\vec{v} - \vec{u}) + \left( 1 - \frac{\rho_f}{\rho_p} \right) \vec{g} + \frac{\rho_f}{\rho_p} \vec{\Omega}^* \times (\vec{u} - \vec{v})$$

$$155 \quad + \frac{\rho_f}{\rho_p} 2\vec{\Omega}^* \times \vec{u} - 2\vec{\Omega}^* \times \vec{v} + \left( \frac{\rho_f}{\rho_p} - 1 \right) \vec{\Omega}^* \times \vec{\Omega}^* \times \vec{x}. \quad (1)$$

156 [The frame rotation was introduced into the non-rotating Maxey-Riley equation by replacing](#)

$$157 \quad \vec{v}_s = \vec{v}_r + \vec{\Omega} \times \vec{x}_r, \quad \vec{u}_s = \vec{u}_r + \vec{\Omega} \times \vec{x}_r,$$

$$158 \quad \frac{D_s \vec{u}_s}{Dt} = \frac{D_r \vec{u}_r}{Dt} + 2 \vec{\Omega} \times \vec{u}_r + \vec{\Omega} \times \vec{\Omega} \times \vec{x}_r, \quad \frac{d_s \vec{v}_s}{Dt} = \frac{d_r \vec{v}_r}{Dt} + 2 \vec{\Omega} \times \vec{v}_r + \vec{\Omega} \times \vec{\Omega} \times \vec{x}_{r2}$$

159 [where subscript “s” denotes stationary frame and subscript “r” – rotating frame. Alternatively,](#)  
 160 [transformation into a rotating frame can be done following the variational method of Ripa](#)  
 161 [\(1987\). The subscripts “r” have then been dropped in Eq. \(1\) and all subsequent equations since](#)  
 162 [all variables are now in the rotating frame. For non-spherical rigid particles, adjustments to the](#)  
 163 [coefficients within the Maxey-Riley equations can be made to account for elliptical shapes \(see,](#)  
 164 [for example, DiBenedetto et al, 2018a,b and references therein\) but at the cost of adding a third](#)  
 165 [vector equation for the orientation of the ellipsoid. However, real microplastics often have](#)



166 [complex tangled-filament-like shapes which are poorly represented by an ellipsoid, and no](#)  
167 [corrections for tangled filaments are currently available.](#)

168 In [Eq. \(1\), which is a](#) statement of Newton's second law for the rigid particle, the right-hand side  
169 represents, in order, the effects of inertia, added mass, drag, buoyancy, Coriolis acceleration  
170 associated with the added mass, the Coriolis acceleration associated with the particle mass,  
171 Coriolis acceleration associated with the fluid motion, and centrifugal acceleration. [A similar](#)  
172 [equation has been previously derived by Beron-Vera et al. \(2019\), though the centrifugal](#)  
173 [acceleration does not appear there explicitly, having been combined with the acceleration due to](#)  
174 [gravity in order to define an effective gravity and corresponding geopotential. Coordinates are](#)  
175 [then imagined to be aligned with geopotential surfaces, though standard spherical or Cartesian](#)  
176 [coordinates are usually used in practice \(Vallis, 2006\). Our explicit retention of the centrifugal](#)  
177 [acceleration will later allow absolute vorticity to arise naturally as a quantity of central](#)  
178 [importance.](#) We have omitted the lift force, the Basset history force, and the Faxen corrections  
179 (Gatignol, 1983). [Faxen corrections account for the variation of the flow across the rigid particle](#)  
180 [and are proportional to  \$a^2 \Delta u\$ . For a particle size that is much smaller than the typical length scale](#)  
181 [of the flow, these corrections are small and typically neglected \(Haller and Sapsis, 2008; Beron-](#)  
182 [Vera et al., 2019\). The history term, which is an integral along a particle path, accounts for the](#)  
183 [boundary layer effects that a particle leaves behind. It is typically ignored under the assumption](#)  
184 [that the chances of other particles crossing that localized boundary layer before it decays are](#)  
185 [small \(Beron-Vera et al., 2019; see also Langlois et al., 2015 and Daitche and Tel., 2011 for](#)  
186 [more info on the influence of the history term on the behavior of rigid particles\). Finally, the lift](#)  
187 [force arises when a particle rotates in a horizontally sheared flow. As shown in Beron-Vera](#)  
188 [2019, the inclusion of the lift force leads to the next-order,  \$O\(\tilde{\epsilon}^2\)\$  correction in the slow-](#)

189 | [manifold approximation, and thus can also be neglected for small  \$\tilde{\epsilon}\$ . In Eq. \(1\),  \$\rho\_p\$  and  \$\rho\_f\$  are](#)  
190 | densities of the rigid particle and the fluid,  $d$  is the particle radius,  $\nu$  is viscosity of the fluid,  $\vec{g}$  is  
191 | the gravity vector, and  $\frac{D\vec{u}}{Dt} = \frac{\partial\vec{u}}{\partial t} + \vec{u} \cdot \nabla\vec{u}$  is the fluid material derivative, evaluated for  
192 | undisturbed fluid velocity at the position of the center of the rigid particle. The position  $\vec{x}(t)$  of a  
193 | particle is determined by

$$194 \quad \frac{d\vec{x}}{dt} = \vec{v}(\vec{x}, t), \quad (2)$$

195 | and together [Eqs. \(1\) and \(2\)](#) compose a coupled, 6<sup>th</sup>-order system for computation of the particle  
196 | position and velocity as functions of time.

197 | If the velocities and lengths are nondimensionalized using characteristic scales  $U$  and  $L$  for the  
198 | background fluid flow, and  $L/U$  is used as a time scale, then [Eq. \(2\)](#) remains formally unchanged  
199 | while the nondimensional form of [Eq. \(1\)](#) is

$$200 \quad \frac{d\vec{v}}{dt} = \frac{3R}{2} \frac{D\vec{u}}{Dt} + \tilde{\epsilon}^{-1}(\vec{v} - \vec{u}) + \left(1 - \frac{3R}{2}\right) \vec{g}_r + 3R\vec{\Omega} \times (\vec{u} - \vec{v}) + 2\left(\frac{3R}{2} - 1\right) \vec{\Omega} \times \vec{v}, \quad (3)$$

201 | where  $\tilde{\epsilon} = \frac{2\rho_f}{\rho_f + 2\rho_p}$ ,  $\vec{g}_r = (\vec{g} - \vec{\Omega}^* \times \vec{\Omega}^* \times \vec{x}) / (U^2)$ ,  $\vec{\Omega} = \frac{\vec{\Omega}^* L}{U}$  and  $\tilde{\epsilon} = \frac{2}{9} \left(\frac{d}{L}\right)^2 \frac{UL}{\nu R}$  is the Stokes  
202 | number, the ratio of the adjustment time scale of a particle (due to drag) to the time scale of the  
203 | background flow. For  $\tilde{\epsilon} \ll 1$ , viscous drag is the dominant force acting on the particle, implying  
204 | that a particle with an initial velocity differing by an amount  $> O(\tilde{\epsilon})$  from the local fluid velocity  
205 | will be rapidly accelerated over a time scale  $\tilde{\epsilon}$  to a velocity proximal to that of the fluid.  
206 | Thereafter the particle will undergo a slow evolution in which the weaker forces due to inertia,  
207 | added mass, and buoyancy cause slight departures from the movement of the fluid itself.

208 | The limit  $\tilde{\varepsilon} \rightarrow 0$  constitutes a singular perturbation of [Eq. \(3\)](#), a problem that can be addressed  
 209 | using an approach due to Fenichel (1979) that was originally formally developed for a steady  
 210 | background flow, but that has been extended by Haller and Sapsis (2008) to include a time-  
 211 | varying background flow. In either case, it can be shown that following the initial viscous  
 212 | adjustment, the particle position and velocity tend toward a subspace or “slow manifold” on  
 213 | which the particle velocity is determined directly by the fluid velocity through an “inertial”  
 214 | equation, here extended to include frame rotation:

$$215 \quad \vec{v} = \vec{u} + \tilde{\varepsilon} \left( \frac{3R}{2} - 1 \right) \left[ \frac{D\vec{u}}{Dt} + 2\vec{\Omega} \times \vec{u} - \vec{g}_r \right] + O(\tilde{\varepsilon}^2). \quad (4)$$

216 | [This result is the same as obtained by Beron-Vera et al. 2019, provided that their gravity vector](#)  
 217 | [is interpreted as our  \$\vec{g}\_r\$ .](#) The same authors also present more general cases, including those with  
 218 | the lift force and on the sphere. In Supplementary Material we present a simple derivation of Eq.  
 219 | (4) based on a multiple-scale expansion. [It provides a quick, though less rigorous, alternative to](#)  
 220 | [the Fenichel approach.](#)

221 | A chief advantage of the slow manifold reduction is that the 6<sup>th</sup> order system [given by Eqs. \(2\)](#)  
 222 | [and \(3\), in which particle velocity needs to be solved for, is reduced to a 3<sup>rd</sup> order system given](#)  
 223 | [by Eqs. \(2\) and \(4\), where the particle velocity is explicitly written as a function of fluid velocity](#)  
 224 | [and flow and particle parameters \(and thus is known\).](#) The bracketed expression in [Eq. \(4\)](#), which

225 | determines the velocity of the rigid particle relative to the fluid, is nothing more than  $\frac{\partial}{\partial x_j} \tau_{ij}$ ,

226 | where  $\tau_{ij}$  is the stress tensor for the fluid. Thus the relative velocity of a rigid particle on the  
 227 | slow manifold is in the same direction as the net force that would act on a fluid parcel occupying  
 228 | the same space. Ordinarily, for a fluid parcel, that force would equate with an acceleration, but  
 229 | on the slow time scale, the relative particle velocity points in the same direction as the net fluid

230 force and its magnitude is proportional to  $\tilde{\varepsilon} \left( \frac{3R}{2} - 1 \right) = \frac{2}{9} \frac{d^2}{L^2} \frac{UL}{\nu} \frac{(\rho_f - \rho_p)}{\rho_f}$ . Since the aggregation of  
 231 rigid particles requires departures of the particle velocity from the (divergence free) velocity  
 232 field of the fluid, one can expect that aggregation will occur more slowly if  $d$  and  $(\rho_f - \rho_p)/\rho_f$   
 233 are small, or if  $\nu$  is large. At the same time, the existence of attractors internal to the fluid may  
 234 depend on  $(\rho_f - \rho_p)/\rho_f$  being small: for example, a large density difference may mean that  
 235 rigid particles simply sink to the bottom or rise to the surface [\(and are thus attracted to attractors](#)  
 236 [external to the fluid interior\)](#).

237 As pointed out by Haller and Sapsis (2008) (also see Beron-Vera et al. 2019), we can consider a  
 238 continuous concentration of rigid particles with [similar](#) properties, and with smoothly varying  
 239 velocity [given by Eq. \(4\)](#). The aggregation of such a concentration would appear to require that  
 240 the divergence of that velocity be negative (though see an apparent counterexample [in Fig 1c](#),  
 241 presented later). Following Haller and Sapsis (2008), consider the evolution of a material  
 242 volume of rigid particles. The time rate of change of this volume is

$$243 \quad \frac{dV}{dt} = \oint \vec{v} \cdot \vec{n} dA_V = \iiint (\nabla \cdot \vec{v}) dV = \iiint \nabla \cdot \left[ \vec{u} + \tilde{\varepsilon} \left( \frac{3R}{2} - 1 \right) \left( \frac{D\vec{u}}{Dt} + 2\vec{\Omega} \times \vec{u} - \vec{g}_r \right) \right] dV \quad (5)$$

244 where  $\nabla \cdot \vec{u} = 0$  for an incompressible fluid. Shrinking  $V$  to an infinitesimal size allows the right-  
 245 hand side to be approximated by  $V$  times the local value in the integrand, and the result may be  
 246 integrated in time, yielding

$$247 \quad V(t) = V_0 \exp \left( \tilde{\varepsilon} \left( \frac{3R}{2} - 1 \right) \int_{t_0}^t \nabla \cdot \left( \frac{D\vec{u}}{Dt} + 2\vec{\Omega} \times \vec{u} - \vec{g}_r \right) ds \right)$$

$$248 \quad = V_0 \exp \left( -2\tilde{\varepsilon} \left( \frac{3R}{2} - 1 \right) \int_{t_0}^t [Q_r(x(s), s) + \vec{\Omega} \cdot \vec{\zeta}_r + |\vec{\Omega}|^2] ds \right)$$

249 
$$= V_0 \exp\left(-2\tilde{\varepsilon} \left(\frac{3R}{2} - 1\right) \int_{t_0}^t Q_a(x(s), s) ds\right). \quad (6)$$

250 Here  $Q_r = \frac{1}{2}\left(\frac{1}{2}|\vec{\zeta}_r|^2 - |S|^2\right)$  is the three-dimensional Okubo-Weiss parameter (Okubo, 1970;  
 251 Weiss, 1991),  $\vec{\zeta}_r$  represents the relative vorticity vector for the fluid,  $S = 1/2(\nabla\vec{u} + (\nabla\vec{u})^T)$  is  
 252 the strain tensor, and  $|S|$  is its Frobenius norm. The final step in [Eq. \(6\)](#) follows from  
 253 introduction of the absolute vorticity vector

254 
$$\vec{\zeta}_a = \vec{\zeta}_r + \overline{2\Omega} \quad (7)$$

255 and the corresponding function  $Q_a = \frac{1}{2}\left(\frac{1}{2}|\vec{\zeta}_a|^2 - |S|^2\right)$ . We note that for a volume  $V$  of any  
 256 size:

257 
$$\frac{dV}{dt} = 2\tilde{\varepsilon} \left(\frac{3R}{2} - 1\right) \iiint Q_a dV = \tilde{\varepsilon} \left(\frac{3R}{2} - 1\right) \iiint \frac{\partial^2}{\partial x_i \partial x_j} \tau_{ij} dV = \frac{2}{9} \frac{d^2}{L^2} \frac{UL}{\nu} \frac{(\rho_f - \rho_p)}{\rho_f} \oint \frac{\partial}{\partial x_j} \tau_{ij} n_i dA_V, \quad (8)$$

259 where  $n_j$  denote the components of the outward unit vector normal to the bounding surface  $A_V$ .

260 [The first equality in Eq. \(8\) is a modest modification of Eq. \(31\) from Haller and Sapsis \(2008\),](#)  
 261 [and one could probably have guessed that our more general result could be obtained by replacing](#)  
 262 [Q with  \$Q\_a\$ . The remainder of the equation expresses volume changes in terms of the fluid](#)  
 263 [stresses.](#) Thus for buoyant particles, a volume  $V(t)$  of any size will contract if the force normal  
 264 to  $A_V$  due to the fluid stresses, integrated around  $A_V$ , is inward. [In many cases, including](#)  
 265 [quasigeostrophic eddies and gyres, internal waves, and the surface gravity waves considered by](#)  
 266 [DiBenedetto et al. \(2018a,b\) and all inviscid flows,](#) the stress tensor is dominated by pressure,  
 267 i.e.,  $\frac{\partial}{\partial x_j} \tau_{ij} \cong -\frac{1}{\rho_f} \nabla p$ , so the tendency to aggregate is determined entirely by the pressure field.

268 In general,  $Q_a$  can change sign along a particle trajectory, making it hard to predict whether the  
 269 surrounding volume shrinks or expands with time. If a buoyant particle is trapped in a region in  
 270 which  $Q_a$  is predominately positive, then this region is a good candidate for aggregation.  
 271 Persistent ocean eddies and other vortical structures are possibilities, not only because vorticity  
 272 tends to dominate over strain, but also because such features have the ability to trap fluid for long  
 273 periods of time. [For dense particles, contraction occurs in areas dominated by strain, and it has  
 274 been shown that aggregation of heavy particles can occur in strain-dominated filaments that arise  
 275 in particle-laden turbulent flows, though the considered particle-to-fluid density differences tend  
 276 to be quite large \(see Brandt and Coletti, 2022 for a review\). In our study, we will focus on  
 277 vortex flows reminiscent of ocean eddies, and on lower dimension objects within such flows that  
 278 can act as attractors for buoyant particles.](#)

279 A simple example of aggregation is given by Haller and Sapsis (2006), who argue that the  
 280 elliptical center of a steady, non-divergent 2d eddy, with  $\vec{g} = |\vec{\Omega}|=0$ , acts as an attractor for  
 281 buoyant particles. Here  $Q_a$  (now  $=Q_r$ ), is ostensibly positive near the elliptical center of the  
 282 eddy, corresponding to contraction of the phase space (which in our case coincides with the  
 283 physical space) of the rigid particle motion. Since the central fixed point of the velocity field of  
 284 the eddy is also a fixed point of the slow manifold particle velocity (Eq. (4)), buoyant particles  
 285 initiated about the center should migrate towards the center. If the eddy is inviscid and its  
 286 streamlines are circular, then the pressure and azimuthal velocity are related by the cyclostrophic  
 287 balance  $\frac{1}{\rho_f} \frac{\partial p}{\partial r} = \frac{u_\theta^2}{r}$  so that  $2Q_r = \frac{1}{\rho_f} \left( \frac{1}{r} \frac{\partial p}{\partial r} + \frac{\partial^2 p}{\partial r^2} \right)$ , and for an eddy in solid body rotation ( $u_\theta =$   
 288  $\Gamma_s r$ ),  $2Q_r = \frac{1}{\rho_f} \left( \frac{1}{r} \frac{\partial p}{\partial r} + \frac{\partial^2 p}{\partial r^2} \right) = 2\Gamma_s^2$ . As suggested in Figure 1a, a small concentration of rigid  
 289 particles indicated by the cross hatched area shrinks as it moves towards the center of the eddy.

290 The contraction is partially due to the geometric effect of movement towards smaller radius  
 291 (term  $\frac{1}{r} \frac{\partial p}{\partial r}$ ) but also due to the fact that the pressure gradient decreases to zero as the center is  
 292 approached and thus the inner edge of the path moves more slowly inward than the outer part  
 293 (term  $\frac{\partial^2 p}{\partial r^2}$ ). In the case of solid body rotation the two terms contribute equally. A second  
 294 example (Fig. 1b) is of an eddy with an azimuthal velocity given by  $u_\theta = \Gamma_C r^{1/2}$ . Here  $\frac{\partial^2 p}{\partial r^2} = 0$   
 295 and  $2Q_r = \frac{1}{\rho_f} \left( \frac{1}{r} \frac{\partial p}{\partial r} \right) = \Gamma_C^2 / r > 0$ , so the contraction of the patch is entirely due to the geometric  
 296 effect of its movement towards smaller radius. The most curious case is that of a point vortex:  
 297  $u_\theta = \Gamma_P r^{-1}$ , for which  $2Q_r = \frac{1}{\rho_f} \left( \frac{1}{r} \frac{\partial p}{\partial r} + \frac{\partial^2 p}{\partial r^2} \right) = \frac{\Gamma_P^2}{r^4} - \frac{3\Gamma_P^2}{r^4} < 0$ . Here the vorticity is zero away from  
 298 the eddy center and the velocity field is dominated by strain. The pressure gradient *increases* as  
 299 the center of the vortex is approached, meaning that the inner part of the patch moves towards  
 300 the center more rapidly than the outer portion (Fig. 1c) and this tendency (quantified by the  
 301 factor  $-\frac{3\Gamma_P^2}{r^4}$ ) surpasses the tendency towards geometrical contraction (quantified by the factor  
 302  $\frac{\Gamma_P^2}{r^4}$ ). The [area of the patch](#) thus expands as rigid particles are drawn towards the center of the  
 303 vortex. [Note, however, that a patch surrounding the center of the vortex can only shrink.](#) This  
 304 behavior is made possible by the singularity at the center, and although this feature is artificial,  
 305 point vortices are often used in idealized models of fluid flow and will act as sinks or “black  
 306 holes” for buoyant particles even though  $2Q_r < 0$ .  
 307 The sign of  $Q_a$  is clearly not the whole story and does not encompass the effects of boundaries.  
 308 For example, consider the fate of heavy ( $\rho_f < \rho_p$ ) particles in the eddy show in Fig. 1a. The  
 309 particles will migrate outward in each case, and no interior attraction will occur unless the eddy  
 310 is surrounded by a boundary, which would then act as an attractor.

311 In the next section, we will consider a more general, 3D, eddy-like circulation: one that has both  
312 vertical and horizontal components of vorticity, time dependence, and a variety of vortical  
313 structures that act as candidates for attraction. Our model is based on the incompressible flow in  
314 a rotating cylinder (Greenspan, 1986), which has been studied in many configurations by  
315 numerous authors as a model of ocean circulation ([Hart and Kittelman, 1996](#); [Pedlosky & Spall,](#)  
316 [2005](#)), ocean eddies ([Pratt et al., 2014](#); [Rypina et al., 2015](#)), or industrial processes and  
317 engineering applications ([Lopez & Marques, 2010 and references therein](#)), and can be easily set  
318 up in the laboratory setting ([Fountain et al., 2000](#); [Lackey and Sotiropoulos \(2006\)](#)). In its  
319 original configuration the cylinder rotates about a vertical axis at a constant (positive) angular  
320 velocity ( $\vec{\Omega} = \Omega \vec{k}$ ), and the lid, which is in contact with the fluid, rotates with a slightly greater  
321 angular speed. The differential rotation sets up an azimuthal circulation in the horizontal and an  
322 overturning circulation in the vertical. (Overturning is observed in ocean eddies as well and  
323 [Ledwell et al. \(2008\)](#) present an example.) The steady, axially symmetric state of the rotating  
324 cylinder flow that is established will be our first object of investigation. A steady but  
325 asymmetrically-perturbed variant can be established by moving the axis of rotation of the lid  
326 away from the axis of rotation of the cylinder, and this offset can also be varied in order to  
327 induce time dependence. [Fountain et al. \(2000\)](#) set a similar situation up in a laboratory cylinder  
328 using a submerged impeller that can be tilted, rather than the differentially rotating lid that can be  
329 shifted, to establish an asymmetric disturbance flow. The authors discussed the Lagrangian  
330 characteristics of the undisturbed flow and demonstrated the existence of secondary vortical  
331 structures generated when the flow is perturbed. [Pratt et al. \(2014\)](#) reproduced similar structures  
332 using a primitive equation simulation and explored the rich assembly of chaotic regions and non-  
333 chaotic vortical structures as functions of the Ekman and Rossby numbers of the flow. The time-



334 dependent version of the rotating cylinder flow and a theory describing the resulting vortical  
 335 structures were discussed by Rypina et al. (2015), who based their examples on a  
 336 phenomenological model that reproduced many of the qualitative features of the numerically-  
 337 obtained velocity field. In dimensionless Cartesian coordinates, the model velocity field is given  
 338 by

$$339 \quad u^{(x)} = -bx(1 - 2z) \frac{r_o - r}{3} - ay(c + z^2) + \varepsilon \left[ y(y - y_o + \gamma \cos(\sigma t)) - \frac{r_o^2 - r^2}{2} \right] (1 - \beta z), \quad (9a)$$

$$340 \quad u^{(y)} = -by(1 - 2z) \frac{r_o - r}{3} + ax(c + z^2) - \varepsilon x(y - y_o + \gamma \cos(\sigma t))(1 - \beta z), \quad (9b)$$

$$341 \quad u^{(z)} = bz(1 - z) \frac{2r_o - 3r}{3}, \quad (9c)$$

342 in which  $r = (x^2 + y^2)^{1/2}$  and  $r_o$  is the cylinder radius. The velocity field consists of a steady,  
 343 axially symmetric flow of strength  $a$  with an overturning circulation of strength  $b$ . To this  
 344 symmetric state one can add an asymmetric, possibly unsteady and depth dependent, perturbation  
 345 of amplitude  $\varepsilon$  (not to be confused with the Stokes number  $\tilde{\varepsilon}$ ). The perturbation is quantified by  
 346 an offset parameter  $y_o$  that introduces axial asymmetry in the velocity field, a frequency  $\sigma$ , and  
 347 an amplitude  $\beta$  for linear depth dependence and an amplitude  $\gamma$  for the time dependence. For the  
 348 case of axially symmetric, steady flow ( $\varepsilon = 0$ ) the horizontal velocity field, in cylindrical  
 349 coordinates, becomes

$$350 \quad u^{(r)} = -br(1 - 2z) \frac{r_o - r}{3} \quad (10a)$$

351 and

$$352 \quad u^{(\theta)} = ar(c + z^2), \quad (10b)$$

353 where  $\theta$  is the azimuthal angle. Table 1 lists the parameter values used for each numerical  
354 experiment.

355 [We now review the main features of the Lagrangian circulation in the rotating cylinder flow.](#) In  
356 the steady, symmetric configuration, each fluid trajectory is confined to the surface of a torus as  
357 it winds around the cylinder. The typical torus is associated with quasi-periodic trajectories and  
358 any such trajectory, followed for a sufficient length of time [so that it completes many](#)  
359 [overturning and azimuthal rotations around the cylinder](#), will sketch out the torus in 3D. Fig. 2b  
360 contains several examples of such tori and Fig. 2a shows the corresponding Poincare map, made  
361 by marking the crossing points of trajectories through a vertical slice through the cylinder. After  
362 a large number of crossings each quasi-periodic trajectory traces out the cross section of the torus  
363 on which it lives. The tori are nested within each other, with a single, horizontal, periodic  
364 trajectory located at the center of the nest. Certain tori contain periodic trajectories, and these  
365 will show up as a finite number of dots on the Poincare map. Because of this geometry, the  
366 motion of fluid parcels is most naturally described in terms of action-angle-angle variables,  
367 where the action,  $I$ , acts a label for a particular torus and is constant following each trajectory,  
368 and the two angle variables,  $\tilde{\theta}$  and  $\phi$ , define the location of a parcel on the torus. Here  $\tilde{\theta}$  is an  
369 azimuthal angle that differs from the above cylindrical coordinate  $\theta$  in how its origin is defined,  
370 while the ‘poloidal’ angle  $\phi$  wraps around the cross-section of each torus. The coordinates are  
371 non-orthogonal but are defined in such a way that the angular velocities,  $\Omega_{\tilde{\theta}}$  and  $\Omega_{\phi}$ , are also  
372 constant following a trajectory. The explicit transformations to the action-angle-angle variables  
373 are given in Mezic and Wiggins (1994).

374 When the symmetric RC flow is perturbed by a small, steady, symmetry-breaking perturbation,  
 375 as controlled by the parameters  $\varepsilon$  and  $y_o$  in [Eq. \(9\)](#), the tori that are populated by periodic orbits  
 376 potentially become resonant and break up, resulting in chaotic motion of fluid parcels in the  
 377 vicinity (Fig. 2d-i). Tori with quasiperiodic orbits deform but stay intact. Examples are discussed  
 378 by Fountain et al. (2000) and Pratt et al. (2013), and the latter found that chaos generally  
 379 dominates in a large region that includes the central axis of the cylinder and extends around the  
 380 boundaries of the cylinder. Away from this region the space is occupied by tori that have  
 381 survived the perturbation, and these are sandwiched between tori that have broken up and created  
 382 braided regions of chaos. The breakup of a torus also gives rise to new tori that appear as islands  
 383 in the Poincare maps (Fig. 3d and 3g) and these contain non-chaotic trajectories. The number of  
 384 islands can be predicted by a theory that decomposes the symmetry-breaking perturbation into  
 385 Fourier modes, written in the  $(I, \tilde{\theta}, \phi)$  coordinates, with wave numbers  $n$  and  $m$  in the  $\tilde{\theta}$  and  $\phi$   
 386 direction. If the angular velocities  $\Omega_{\tilde{\theta}}$  and  $\Omega_{\phi}$  characterizing the trajectories on a particular torus  
 387 satisfy the resonance condition  $n\Omega_{\tilde{\theta}} + m\Omega_{\phi} = 0$  for some  $n$  and  $m$ , equivalent to the trajectories  
 388 on that torus being periodic, then that torus will break up and a new set of invariant tori (islands)  
 389 will form. Running through the center of the islands will be a periodic trajectory that will execute  
 390  $n$  azimuthal cycles to every  $m$  poloidal (overturning) cycles. In the case shown in Fig. 3a,  
 391  $n = m = 1$ , so the periodic trajectory circles the cylinder horizontally once for each overturning  
 392 cycle: a so-called 1:1 resonance.

393 If the symmetry breaking perturbation is quasi-periodic in time, with underlying frequencies  $\sigma_i$ ,  
 394 the resonance condition for the breakup of a torus becomes  $n\Omega_{\tilde{\theta}} + m\Omega_{\phi} + l_i\sigma_i = 0$ , where  $l_i$ 's  
 395 are integers (Rypina, et al. 2015). Unlike the resonance condition for the steady perturbation,  
 396 which is only satisfied on tori foliated by periodic trajectories, this new resonant condition may

397 be satisfied on tori that have quasi-periodic orbits, and the resonant islands that form will have a  
398 shape and location that vary in time. An example (Fig. 2g,h) of the case of a resonance with a  
399 single-frequency (i.e., time-periodic) perturbation shows a number of resonant islands. These  
400 features vary in time, recovering their shape and location periodically, and the snapshots shown  
401 are obtained by strobing the trajectories in 3D and at the forcing frequency. The green and blue  
402 islands in Fig. 2h have resulted from the breakup of tori with quasiperiodic trajectories, and  
403 center of the island corresponds to a closed material curve that is populated with quasiperiodic  
404 trajectories.

405 [Note that the resonance condition above and our results in general are applicable to quasi-](#)  
406 [periodic disturbances with finite number of frequencies, rather than only periodic disturbances.](#)  
407 [\(We only show numerical simulations for the time-periodic case for simplicity.\) Because any](#)  
408 [broad-spectrum function can be arbitrary closely represented by a quasi-periodic function with a](#)  
409 [finite number of frequencies, this could be applicable to some oceanic flows, especially those](#)  
410 [with pronounced peaks in the spectrum. However, for flows with truly broadband spectrum, this](#)  
411 [approach is probably poorly applicable and/or at least impractical because of the very large](#)  
412 [number of discrete frequencies needed. This is similar in its utility/applicability to other](#)  
413 [Kolmogorov-Arnold-Moser—based and resonance—based arguments used in prior papers by](#)  
414 [many authors \(including both us and the reviewer\), see, for example, Rypina et al., 2007 and](#)  
415 [Beron et al., 2008; 2010.](#)

### 416 **III. Results**

417 Aggregation of rigid particles will occur in presence of an attractor, an object with a dimension  
418  $< 3$  to which particles tend asymptotically in time. We are most interested in attractors that

419 occur in the interior of the rotating cylinder, and are set up by the background circulation, as  
 420 opposed to the physical boundaries of cylinder. We will see that a closed material contour  
 421 consisting of periodic orbits near the core of the nested tori in the steady symmetric case act as  
 422 an attractor for slightly buoyant particles, and that similar material contours consisting of  
 423 periodic or quasiperiodic orbits near the centers of the resonant islands in the asymmetric cases  
 424 can play the same role. We will explore three cases in increasing complexity, beginning with  
 425 steady flows with axial symmetry, and proceeding to steady, asymmetric flows and finally  
 426 unsteady asymmetric flows.

427 The search for attractors is motivated by the hypothesis that for cases of strong drag, where the  
 428 rigid particle velocity lies close to the fluid velocity, a periodic orbit for the rigid particle motion  
 429 will exist in the vicinity of a periodic trajectory for the fluid parcel motion, and that if  $Q_a > 0$  in  
 430 a region surrounding the latter, that it should attract buoyant particles. For the time-dependent  
 431 case, we extend the search to included closed loops that contain recirculating rigid particles and  
 432 that vary periodically in time.

433 (a) steady, axially-symmetric 3D flows

434 The fluid velocity field for this case is given by Eqs. (9c) and (10), and these indicate that the  
 435 location of the horizontal, periodic fluid parcel trajectory living at the center of the nested tori, is  
 436 given by  $r = 2r_o/3$  and  $z = \frac{1}{2}$ . It is natural to ask whether a periodic trajectory for rigid particles  
 437 also exists nearby. In the slow-manifold approximation, the steady radial, azimuthal and vertical  
 438 particle velocities are obtained by writing Eq. (4) in cylindrical coordinates, leading to

$$439 \quad \mathbf{v}^{(r)} = \mathbf{u}^{(r)} + \tilde{\varepsilon} \left( \frac{3R}{2} - 1 \right) \left[ \left( u^{(r)} \frac{\partial}{\partial r} + u^{(z)} \frac{\partial}{\partial z} \right) \mathbf{u}^{(r)} - u^{(\theta)} \left( 2\Omega + \frac{u^{(\theta)}}{r} \right) - \Omega^2 \mathbf{r} \right] \quad (11a)$$

440 
$$v^{(\theta)} = u^{(\theta)} + \tilde{\varepsilon} \left( \frac{3R}{2} - 1 \right) \left[ \left( u^{(r)} \frac{\partial}{\partial r} + u^{(z)} \frac{\partial}{\partial z} \right) u^{(\theta)} + u^{(r)} \left( 2\Omega + \frac{u^{(\theta)}}{r} \right) \right] \quad (11b)$$

441 
$$v^{(z)} = u^{(z)} + \tilde{\varepsilon} \left( \frac{3R}{2} - 1 \right) \left[ \left( u^{(r)} \frac{\partial}{\partial r} + u^{(z)} \frac{\partial}{\partial z} \right) u^{(z)} + g \right] \quad (11c)$$

442 *Position of attracting periodic orbit; approximate analytical expression on a slow manifold*

443 Searching for points  $r = r_c$  and  $z = z_c$  for which  $v^{(r)} = v^{(z)} = 0$ , and that lie in the proximity of

444 the horizontal trajectory of the flow, we introduce

445 
$$r_c = \frac{2r_o}{3} + \tilde{\varepsilon} \left( \frac{3R}{2} - 1 \right) \tilde{r} \text{ and } z_c = \frac{1}{2} + \tilde{\varepsilon} \left( \frac{3R}{2} - 1 \right) \tilde{z}.$$

446 Substituting into the right-hand sides of (11a,c) and setting both to zero results, after neglect of

447  $O(\tilde{\varepsilon}^2)$  terms, in

448 
$$r_c = \frac{2r_o}{3} + \tilde{\varepsilon} \left( \frac{3R}{2} - 1 \right) \frac{g}{b} r_o \quad (12a)$$

449 and

450 
$$z_c = \frac{1}{2} + \frac{9}{2br_o} \tilde{\varepsilon} \left( \frac{3R}{2} - 1 \right) \left[ \Omega^2 + a \left( c + \frac{1}{4} \right) \left( 2\Omega + a \left( c + \frac{1}{4} \right) \right) \right]. \quad (12b)$$

451 For the parameters  $a > 0$  and  $b > 0$ , circulation is cyclonic with upwelling in the center of the

452 cylinder, and  $(3R/2) - 1 > 0$  for buoyant particles, so the  $O(\tilde{\varepsilon})$  corrections are positive and the

453 periodic particle orbit lies at larger radius and elevation than the periodic fluid orbit. Note also

454 from Eq. (11b) that the azimuthal velocity component of the rigid particle on the periodic orbit is

455 equal to that of the fluid.

456 An explanatory sketch (Fig. 3) shows the position of the periodic orbit of the rigid particle

457 relative to that of the periodic orbit of the fluid. Since the rigid particle is buoyant, it can

458 maintain its level  $z$  only if it is situated in a region where the vertical fluid velocity is  $< 0$ , here  
 459 to the right of the fluid periodic orbit. Also, the horizontal pressure gradients associated with the  
 460 centripetal acceleration associated with the frame rotation (term  $\Omega^2 r$ ), the Coriolis acceleration  
 461 (term  $2\Omega u^{(\theta)}$ ), and the centripetal acceleration due to the azimuthal velocity  $u^{(\theta)^2}/2r$  are all  
 462 positive for this flow, so that low pressure exists at  $r=0$  and the rigid particle is forced  
 463 horizontally inward. To remain stationary the particle must sit in a region where the radial  
 464 velocity of the fluid is outward. In this manner, the periodic trajectory exists at a location where  
 465 the forces of inertia, buoyancy and added mass can be countered by the drag due to the  
 466 background flow. If we fix all other parameters and increase  $\Omega$  through positive values, the term  
 467 multiplying  $\tilde{\epsilon}$  in [Eq. \(12b\)](#) will become dominated by the  $\Omega^2$  term and will grow without bound  
 468 and the periodic trajectory may cease to exist. At the same time, a periodic orbit for the rigid  
 469 particle can always be found close to that of the fluid, regardless of the magnitudes of the  
 470 parameters  $\Omega$ ,  $a$ ,  $b$  etc., provided that the relative particle size  $d/L$  (and thus  $\tilde{\epsilon}$ ), and/or the  
 471 relative density difference  $\frac{(\rho_f - \rho_p)}{\rho_f}$  (and thus  $\frac{3R}{2} - 1$ ) are made sufficiently small.

472 *Position of attracting periodic orbit; conditions for the loss of periodic orbit*

473 We have suggested that periodic orbits for rigid particles are encouraged when the  $\tilde{\epsilon} \left( \frac{3R}{2} - \right.$   
 474  $\left. 1 \right) \ll 1$ , and in the case of Run 1 the value is 0.0066. A cross-sectional plot of the radial and  
 475 vertical components of the slow manifold particle velocity in a vertical section through the  
 476 cylinder (Fig. 4a) shows that the periodic orbit lies at  $r = 0.369$  and  $z = 0.504$  (as compared to  
 477 the values  $r_c = 0.338$  and  $z_c = 0.502$  predicted by [Eq. \(12\)](#). (The convergence of the  
 478 surrounding velocity field is too weak to be seen in the graphic.) If  $\tilde{\epsilon} \left( \frac{3R}{2} - 1 \right)$  is raised to the  
 479 moderately small value 0.02, the position of periodic trajectory migrates towards larger radius

480 (Fig. 4b), the reason being that the greater buoyancy (larger value of  $\frac{3R}{2} - 1$ ) or smaller drag  
 481 (larger  $\tilde{\epsilon}$ ) requires a larger downward fluid velocity for equilibrium. Since the maximum  
 482 downward fluid velocity occurs at the outer cylinder wall (see Eq. (9c)) the position of the  
 483 periodic orbit continues to migrate outward and is lost (Fig. 4c) when  $\tilde{\epsilon} \left( \frac{3R}{2} - 1 \right)$  exceeds a  
 484 value close to 0.3.

485 *Position of periodic orbit in numerical simulations:*

486 The slow-manifold reduction yields to the prediction (Eq. (12)) of the position of the attracting  
 487 material contour, or loop, for slightly buoyant particles. We can compare this prediction to what  
 488 is observed in numerical simulations using the Maxey-Riley [Eqs. \(1\) and \(2\)](#) over a range of  
 489 particle size  $d$  (and thus  $\tilde{\epsilon}$ ) and frame rotation  $\Omega$ . As shown in Fig. 5, qualitative agreement with  
 490 the slow-manifold prediction, and the sketch in Fig. 3, holds for a very small  $d$  (when  $\tilde{\epsilon}$  is small).  
 491 Here the attractor in Fig. 5 is located close to the central periodic fluid parcel trajectory that lives  
 492 at mid-depth,  $z = 0.5$  and  $r = \frac{2R}{3} \approx 0.33$ . As  $d$  (and  $\tilde{\epsilon}$ ) increases, the attractor moves  
 493 increasingly up and outward, and although the theory captures the trends, quantitative agreement  
 494 with the numerical results worsens. Also, when frame rotation  $\Omega$  is increased (panel c), the  
 495 attractor responds by shifting up from mid-depth, again in qualitative but not quantitative  
 496 agreement with the slow-manifold prediction in [Eq. \(12b\)](#).

497 *Geometry of rigid particle trajectories and evidence of attraction in numerical simulations:*

498 If in the neighborhood of the periodic rigid particle trajectory  $Q_a > 0$ , the phase space for  
 499 buoyant particles will contract and the periodic trajectory becomes a candidate for an attractor of  
 500 such particles. An example of the attraction towards the periodic orbit is shown in Figure 2c,  
 501 where a set of slightly buoyant particles ( $\frac{\rho_p}{\rho_f} = 0.97$ ) has been initialized over the volume of the



502 | cylinder, and Eqs. (1) and (2) have been integrated forward in time to determine their subsequent  
503 | trajectories. Each trajectory is shown using a unique color. It can be seen that the particles  
504 | aggregate within a ring-like structure of decreasing thickness in the general vicinity of the  
505 | periodic orbit of the fluid flow.

506 | Basin of attraction – relationship to  $Q_a$ :

507 | To map out the basin of attraction for the particle periodic orbit, we first consider the region over  
508 | which phase space contraction for the buoyant particles (i.e.  $Q_a > 0$ ) occurs. This region is  
509 | shown in Fig. 6a for the current example, along with the streamlines of the fluid overturning  
510 | stream function. Much of the fluid flow recirculates entirely within the region of positive  $Q_a$ ,  
511 | whereas some of the outer streamlines cross the boundary (thick contour) between positive and  
512 | negative  $Q_a$ . If it were the case that rigid particles exactly followed streamlines of the fluid  
513 | overturning circulation, then net contraction or expansion of phase space along a rigid particle  
514 | trajectory would depend on the sign of the time-integrated value of  $Q_a$  along streamlines. The  
515 |  $Q_a = 0$  contour, shown by a bold contour in each frame of Fig. 6, might then approximately  
516 | delineate the basin of attraction for buoyant rigid particles. In the slow-manifold approximation,  
517 | where rigid particle velocities lie close to the fluid velocities, the  $Q_a = 0$  contour might continue  
518 | to do so.

519 | To test this conjecture, we locate the basin of attraction in the numerical simulations by releasing  
520 | buoyant particles at various locations in the cross-section  $0 < x < r_o$  and  $0 < z < 1$ , integrating  
521 | the subsequent trajectories over many overturning cycles, and recording the position ( $x_{final}$  and  
522 |  $z_{final}$ ) of each particle where it crosses the same plane the final time (i.e., recording final  
523 | crossing with the Poincare section). [We use the variable-step 4-th order Runge-Kutta integration](#)

524 scheme, which we implemented in Matlab via the built-in function “ode45”. In our simulations, the  
525 relative and absolute tolerances are set to the value of  $10^{-9}$  to integrate particle trajectories (Eqs.  
526 (2) and (3)) (our results were not sensitive to the further decrease in tolerance values). Since the  
527 flow (Eqs. (9a,b,c)) is prescribed analytically and has no normal flow component at the perimeter  
528 and top and bottom of the cylinder, no interpolation scheme is needed and no extra boundary  
529 conditions are enforced during the integration. Integration of a trajectory is stopped when a  
530 particle got within one particle radius from the cylinder walls or top/bottom. The values of  $z_{final}$   
531 as a function of initial particle position are mapped in Fig. 7a, where the large green area  
532 corresponding to  $z_{final} \cong 0.5$  indicates the region from which particles are attracted. Only  
533 particles initiated near the central axis of the cylinder, and close to the cylinder boundaries lie  
534 outside this region, and these rise to the surface of the cylinder, contact the upper lid, and are no  
535 longer followed. It can be seen that the green area in Fig. 7a has an oval shape that somewhat  
536 resembles the overturning streamlines at small  $x$  in the central part of the cylinder, but extends to  
537 near the top, bottom and outer cylinder boundaries at larger  $x$ . Thus the  $Q_a = 0$  contour provides  
538 a rough indication of the size and shape of the basin of attraction, but misses some important  
539 details.

#### 540 *Basin of attraction – dependence on $\Omega$*

541 We have seen that the location of the periodic orbit that acts as an attractor for buoyant particles  
542 shifts up and out in response to increasing frame rotation  $\Omega$  (Fig. 5c). In Fig. 8 we indicate the  
543 corresponding changes in the extent of the basin of attraction with respect to changing  $\Omega$  by re-  
544 computing Fig. 8a with  $\Omega = 0.3, 1, \text{ and } 10$ . The two smaller  $\Omega$  values (0.3 and 1) correspond  
545 roughly to Rossby numbers  $a/2\Omega$  of about 1 and 0.2, i.e., are representative of the ocean  
546 submesoscale and mesoscale flows. The  $Q_a$ -functions for these cases are plotted in Fig. 6b-c.

547 Most submesoscale eddies are going to tend to have  $u^{(\theta)}/r$  about the same magnitude as  $\Omega$   
548 (except on the equator) and mesoscale eddies will have  $u^{(\theta)}/r \ll \Omega$ . The results in Fig. 8  
549 suggest that, while the basin of attraction does shrink slightly with increasing  $\Omega$ , this dependence  
550 is weak. The main difference between the three numerical runs in Fig. 8 is in the associated  
551 attraction time, which gets significantly shorter for larger values of  $\Omega$ . This is explored in more  
552 detail below.

553 Attraction time:

554 It follows from Eq. (6) that the attraction time towards the periodic orbit should scale as  $T_a =$   
555  $\left[ 2\tilde{\epsilon} \left( \frac{3R}{2} - 1 \right) Q_a \right]^{-1}$  where  $Q_a = \frac{1}{2} \left( \frac{1}{2} |\vec{\zeta}_a|^2 - |S|^2 \right)$  with  $\vec{\zeta}_a = \vec{\zeta}_r + \overline{2\Omega}$ . Thus, for  $\vec{\zeta}_r \geq 0$ , as in  
556 most of our numerical runs (except Experiment 1e), attraction time decreases with increasing  $\Omega$   
557 for positive  $\Omega \geq 0$ . For negative  $\vec{\zeta}_r$ , which corresponds to the reversed direction of the flow in  
558 our simulations (Experiment 1e), an increase in  $\Omega$  will initially slow the attraction by decreasing  
559 the magnitude of  $\vec{\zeta}_a$  all the way to 0, at which point the periodic orbit will lose its attraction  
560 properties, but then will speed up the attraction as  $\Omega$  is further increased. This trend is confirmed  
561 numerically in Fig. 9, where for the flow parameters corresponding to the “reversed flow” run in  
562 Table 1 (Experiment 1e, with  $\vec{\zeta}_r < 0$ ), we release a sample trajectory within the basin of  
563 attraction and plot its  $z$ -coordinate as it winds around the can and eventually approaches the  
564 attracting periodic orbit. As anticipated, the attraction time initially increases as  $\Omega$  is increased  
565 from 0 to 0.6, but then decreases as  $\Omega$  is further increased to 2.

566 Disappearance of the subsurface attractor when  $\tilde{\epsilon}$  becomes too large:

567 Finally, to illustrate the disappearance of the subsurface attractor when  $\tilde{\epsilon}$  becomes too large, in  
568 Fig. 10, we contrast 2 numerical simulations with the same flow parameters (corresponding to

569 the “slow overturn” run 1c in Table 1) but different particle diameters,  $d = 10^{-3}$  vs  $d = 5 \times$   
570  $10^{-4}$ . For larger  $d$ , the subsurface periodic orbit for rigid particles is no longer present within the  
571 can, leading to all particles rising up to the surface (Fig. 10b). For smaller  $d$ , the periodic orbit is  
572 still present and acts as an attractor for buoyant rigid particles over a significant portion of the  
573 | can (green region in Fig. 10a). We note that this run [would be more qualitatively similar to the](#)  
574 | oceanic mesoscale or submesoscale eddies, where the overturning component of circulation is  
575 | weak in comparison to the horizontal swirl.

576 (b) steady non-symmetrically perturbed case

577 We now consider a case in which the axial symmetry of the steady flow has been broken, here  
578 through a change in the perturbation amplitude parameter  $\varepsilon$  from zero to 0.25, and in the offset  
579 | parameter  $y_0$  from 0 to  $-0.2$  in the Eqs. (9a,b). The fluid velocity field now contains something  
580 | like a stationary, “mode-1” azimuthal wave in the horizontal velocity field.

581 The resulting Lagrangian structure (Fig. 2d and e) has a sea of chaos that covers the near-axial  
582 and outer regions of the cylinder, where no unbroken tori survive. Within this chaotic sea is a  
583 region containing a nest of unbroken tori that surround a central periodic orbit. This orbit has  
584 evolved from the central periodic orbit of the symmetry case and is now tilted. Within the nest of  
585 unbroken tori there exist resonant layers, in which new tori have arisen, and the most prominent  
586 is the “island” that is centered near  $x = 0.4$  and  $z = 0.2$  in the right-half (and near  $x = 0.4$  and  
587  $z = 0.2$  in the right half) of Fig. (2d). We further note that this center lies within the region of  
588 | positive  $Q_a$  (Fig. 6b). The island corresponds to the yellow tori in Fig. 3e and is produced by a  
589 | 1: 1 resonance, so that the periodic trajectory running through its center executes one complete  
590 | azimuthal cycle and one overturning cycle before connecting back onto itself. Thus, in this  
591 | steady asymmetric configuration, we now have 2 periodic orbits of the fluid flow – the central

592 slightly-tilted periodic orbit near mid-depth (that evolved from the central horizontal periodic  
593 orbit of the axisymmetric flow) and a new periodic orbit running through the center of the  
594 resonant island (resulting from the break-up of the resonant torus satisfying  $\Omega_{\bar{\theta}} + \Omega_{\phi} = 0$ ).

595 We speculate that for sufficiently small  $\tilde{\epsilon}$  a periodic orbit for the rigid particle motion exists in  
596 the vicinity of each of the 2 periodic orbits of the fluid flow. This conjecture is difficult to prove  
597 due to a complex geometry, leading to centrifugal forces that act in different directions at  
598 different locations along the particle path. For now we simply search for the supposed attractors  
599 by releasing particles and following their trajectories.

600 As shown in Fig. 2f, separate attractors arise in the vicinity of two periodic orbits. The first  
601 appears as a ring-like structure (purple core) lying near the center of the original nested tori and  
602 the second is a similar feature with a red core near the center of the resonant island. The two are  
603 chained together and each has its own basin of attraction (Fig. 7c): the first consisting of a  
604 roughly elliptical patch (inner green region) in the x-z-plane, which corresponds of a slice  
605 through a tube-like structure in 3D, and the second consisting on an annular (blue) region that  
606 surrounds the green region and that occupies a relatively larger volume.

607 In order to check that attraction of slightly-buoyant rigid particles towards periodic orbits located  
608 near the centers of the resonant islands in the perturbed flow is not limited to the case of the 1: 1  
609 resonance, in an additional simulation (Fig. 11, experiment 2c in Table 1), we adjusted the  
610 background flow parameter  $b$  in Eqs. (9), which is responsible for the overturning strength, to  
611 create a 2: 1 resonance instead of a 1: 1 resonance, as in the original run. In this case, the  
612 resonant torus breaks down giving rise to a 2-island chain on the corresponding Poincare section  
613 (Fig. 11a), and the fluid periodic orbit that goes through the centers of both islands completes 2

614 full cycles in azimuth and 1 complete cycle in vertical before connecting onto itself. Also, as in  
615 the original run, a second slightly-tilted periodic orbit still exists near mid-depth of the can.  
616 When buoyant particles are released into this flow, two attractors arise, corresponding to the 2  
617 periodic orbits of rigid particles – one near mid-depth (purple core in Fig. 11c) and another in red  
618 near the center of the 2: 1 resonant island.

619 Shift in position of the periodic orbit associated with a resonant island as a function of flow and  
620 particle parameters, and frame rotation

621 The position of the attracting periodic orbit for rigid particles that is located within the resonant  
622 islands (we will refer to it as the resonant periodic orbit) in the asymmetrically-perturbed flow  
623 depends both on the perturbation strength (via  $\varepsilon$ ), on the flow and particle parameters (via  $\tilde{\varepsilon}$ ), and  
624 on the frame rotation  $\Omega$ . Specifically, this resonant periodic orbit for the rigid particles will shift  
625 away from the corresponding periodic trajectory of the fluid flow as  $\tilde{\varepsilon}$  and  $\Omega$  are increased. The  
626 same is true for the slightly-tilted central attracting periodic orbit near mid-depth. This is  
627 qualitatively similar to the shifting of the central periodic orbit up and out from  $z = 0.5$ ,  
628  $r = 0.34$  in the axisymmetric flow in response to changing  $\tilde{\varepsilon}$  and  $\Omega$ , which we explored in detail  
629 the previous section both analytically (Eqs. (12)) and numerically (Fig. 3-5).

630 In order to numerically illustrate the shift in the position of the attracting periodic orbits, we  
631 present (Figs. 12 and 13) numerical simulations in the steady perturbed flow configuration for 3  
632 values of  $d$  (and thus  $\tilde{\varepsilon}$ ) and 3 values of  $\Omega$ . As both parameters increase, the attractors move  
633 away from the corresponding periodic orbits of the fluid flow. This shift is evident from the  
634 change in the color of the attraction basins in (a,d,g) and from the location of the yellow cloud of  
635 dots in (c,f,i) in Figs. 12-13. Increases in  $\tilde{\varepsilon}$  and  $\Omega$  also lead to the shrinkage of the attraction

636 basins for both attractors and to a faster convergence rate, as is evident from the tighter cloud of  
637 yellow dots in (c,f,i), as discussed in more detail below. The basin of attraction for the central  
638 attractor – the green region in Fig. 12 – seems to shrink faster than the basin of attraction for the  
639 resonant attractor (the blue-ish region) as  $d$  increases, so when  $d$  is increased from  $2 \times 10^{-3}$  to  
640  $3 \times 10^{-3}$ , the central attractor vanishes, whereas the resonant attractor is still present (Fig. 12g).  
641 On the other hand, the increase in  $\Omega$  (Fig. 13) causes a faster shrinkage of the basin of attraction  
642 for the resonant attractor than for the central attractor, so when  $\Omega$  is increased from 2 to 5 in Fig.  
643 13g, the resonant attractor disappears, whereas the central attractor is still present. Figs. 12g,h,i  
644 (and Fig. 13g,h,i) show cases where this threshold has been exceeded, and one of the attractors  
645 has been lost, whereas the other is still present.

646 Attraction time:

647 Similar to the unperturbed flow, the attraction time for attractors in the steady, perturbed flow  
648 may still scale as  $T_a = \left[ 2\tilde{\varepsilon} \left( \frac{3R}{2} - 1 \right) Q_a \right]^{-1}$ , provided that  $Q_a$  is regarded as a typical value  
649 within the corresponding basin of attraction. The predicted decrease in attraction time with  
650 increasing  $\tilde{\varepsilon}$  and  $Q_a$  is evident from the numerical simulations in Figs. 12-13, where in (c,f,i) we  
651 color-coded trajectory crossings with the x-z Poincare plain by time, with blue/yellow  
652 corresponding to initial/final time. For smaller values of  $\tilde{\varepsilon}$  and  $\Omega$ , we observe a wider and more  
653 diffuse cloud of dots (because trajectories wind around the can many times before approaching  
654 the attractor), whereas as  $\tilde{\varepsilon}$  and  $\Omega$  increase, the clouds at comparable times become denser and  
655 more compact around the attractors.

656 Basin of attraction

657 For the slightly-tilted central periodic orbit located within the central non-chaotic region near  
658 mid-depth in Fig. 2f, we observe that the basin of attraction – green region in Fig. 7b – extends  
659 roughly from the location of the periodic orbit to the edge of the central non-chaotic region (that  
660 is foliated by discretely sampled closed curves in Fig. 2d). Note that as  $\tilde{\epsilon}$  increases, the attracting  
661 periodic orbit moves away from the center of this non-chaotic region towards its edge, leading to  
662 the shrinkage and eventual disappearance of the corresponding basin of attraction, shown by the  
663 green regions in Fig. 12a,d,g).

664 Similarly, in all of our numerical simulations, we observe that for the resonant attracting periodic  
665 orbit running through the resonant islands, the basin of attraction seems to cover the region  
666 between the orbit and the edge of the corresponding resonant island. An analytical expression for  
667 the width of the (non-degenerate) resonant island in the fluid flow (Pratt et al., 2014) predicts

668 that  $\Delta I = \sqrt{\frac{\epsilon F_{nm}^0(I_0)}{\left(n \frac{d^j \Omega_\phi}{dI^j} + m \frac{d^j \Omega_\theta}{dI^j}\right)_{I_0}}}$ , where  $\Delta I$  is the deviation in the action coordinate away from  $I_0$ , the

669 value of action at the resonant torus (i.e., at the center of the island). This width depends on the  
670 strength of the perturbation  $\epsilon$ , the order of the resonance (via  $n$  and  $m$  in the resonance

671 condition), the background flow (via  $\frac{d^j \Omega_\phi / \theta_j}{dI^j}$ ), and the structure of the perturbation (via  $F_{nm}^0(I_0)$ ).

672 This expression could be used as an upper limit on the extent of the basin of attraction. However,  
673 because the attracting periodic orbit will move away from the center of the island towards its  
674 edge as  $\tilde{\epsilon}$  and  $\Omega$  increase, the basin of attraction for the resonant attractor (blue region in Figs.  
675 12a,d and 13a,d) becomes increasingly smaller than  $\Delta I$ . One might speculate, then, that the  
676 attractor will completely disappear when the attracting periodic orbit reaches the edge of the  
677 resonant island. This is the case in Figs. 13g where the resonant attractor is no longer present.



678 (c) non-steady, non-symmetrically perturbed case

679 The final case that we will consider is one in which the perturbation is asymmetric and varies  
680 periodically in time. The chosen perturbation frequency,  $\sigma = 2\pi/9.1$ , causes 2 strong additional  
681 resonances (compared to the steady perturbed case) – one with  $n = 0$ ,  $m = 1$ , and  $l = 1$  (i.e.,  
682 with a torus whose overturning frequency is equal to the perturbation frequency) that is shown in  
683 blue in Fig. 2g,h and is located near the outer edge of the central non-chaotic region, and another  
684 resonance, shown in green in Fig. 2g,h, with  $n = 1$ ,  $m = 1$ , and  $l = 1$ , which is located between  
685 the central non-chaotic region and the larger  $n = 1$ ,  $m = 1$  resonant island (that was present in  
686 the steady case as well). Both of these new resonant structures are time dependent, their shape  
687 and position recurring periodically. For example, the blue island, which looks like a crescent  
688 moon pointing upward on the Poincare section at  $t = 0$ , becomes a crescent moon pointing  
689 downward at time 4.55. The movement of the green island is more complex, as it turns both in  
690 azimuth and vertical, making one complete loop over 9.1 time units. Because of the time-  
691 dependence, trajectories must be strobed at the forcing frequency  $\sigma$  in order to capture  
692 ‘snapshots’ of their forms as they recur at a particular phase in the time cycle. At the center of  
693 each feature is a closed material curve that also varies periodically. Where the island has  
694 emerged from the breakup of a torus with quasiperiodic orbits, the individual trajectories that  
695 populate the material curves are themselves quasiperiodic.

696 Particle trajectory computations in this case confirm that the purple, red and green islands give  
697 rise to attractors (Fig. 3i), whereas the blue island does not. In fact, slightly-buoyant rigid  
698 particles that are released in the blue region converge towards the attractor that lies near the  
699 purple region. This is also indicated by the basin of attraction of the central attractor extending  
700 across the space occupied by the blue resonant island in Fig. 7c.

#### 701 IV. Discussion

702 We have considered attraction phenomena for small, finite size, spherical, buoyant, rigid  
703 particles in a three-dimensional rotating cylinder flow with azimuthal rotation and overturning,  
704 and both with or without time dependence. The aim has been to gain insights into the behavior of  
705 slightly buoyant microplastic particles in 3D vortex flows that qualitatively resemble ocean  
706 eddies. The rigid particle motion is governed by a simplified version of the Maxey-Riley  
707 equations (accounting for inertia, buoyancy and simplified quantification of drag and added  
708 mass), and, approximately, by the slow-manifold reduction of these equations. We have  
709 illustrated the possibility of aggregation of slightly-buoyant rigid particles in 3D vortex flows  
710 towards closed loop attractors located subsurface within the interior of the flow. Even in our  
711 idealized flow and for spherical particles with fixed radius and buoyancy, aggregation is non-  
712 trivial, often with multiple attractors present and/or the lack of attraction in some circumstances.  
713 Our rotating cylinder model is much less complex than any real ocean eddy in many respects,  
714 including the assumed quasiperiodic time dependence and the absence of decay and interaction  
715 with the surroundings. Understanding aggregation in a simple periodic flow seems like a  
716 reasonable first step towards understanding aperiodic, interacting, and decaying oceanic eddies.  
717 This approach is common in applications of dynamical systems theory to oceanography and  
718 meteorology. For example, arguments relating to the increased stability of jets due to the strong  
719 Kolmogorov-Arnold-Moser stability near shearless trajectories have first been developed for  
720 spatially-periodic and time-quasiperiodic flows and tested using idealized toy models, before  
721 exploring these ideas in more realistic oceanic and atmospheric settings (see Rypina et al., 2007  
722 and Beron et al., 2008; 2010). Note also that our results are applicable to quasi-periodic  
723 disturbances with finite number of frequencies rather than just periodic disturbances (we only

724 [show numerical simulations for the time-periodic case for simplicity\), and a quasiperiodic](#)  
725 [function might potentially be useful for approximating temporal variability in some oceanic](#)  
726 [flows, especially those with pronounced peaks in the spectrum.](#)

727 We have explored a steady axisymmetric rotating cylinder flow and a steady flow with its axial  
728 symmetry broken. In all cases, we have observed emergence of subsurface attracting structures  
729 that lead to the aggregation of buoyant particles towards them. We have linked these attractors to  
730 the periodic orbits of rigid particles that exist in a region of net contraction of the phase space of  
731 the particle motion. The slow manifold equations suggest that periodic orbits for rigid particles  
732 exist near periodic orbits of the underlying fluid flow, provided the drag is sufficiently strong  
733 (Stokes number  $\ll 1$ ).

734 We have also explored one case of an axially asymmetric and time-periodic flow, with focus on  
735 the resonant “islands” that arise due to the time-dependence. At the center of such islands are  
736 closed material contours, or loops, composed of quasi-periodic orbits of the fluid flow. One such  
737 structure has a nearby attractor, also a closed loop of quasiperiodic orbits for rigid particles,  
738 while a second example does not. A detailed explanation awaits formulation of a quantitative  
739 theory, something that is beyond the scope of the present paper and that will be presented in a  
740 future work.

741 We have observed that the disappearance of an attractor, which can occur as the result of  
742 increasing rigid particle size or frame rotation, coincides roughly with the displacement of the  
743 position of the attractor to the outer edge of the resonant island from which it sprang. Whether  
744 this purely geometric observation forms the basis for a general criterion for the loss of attraction  
745 is unknown, as a dynamical justification is needed.

746 [Marine microplastics can have complex non-spherical tangled-filament shapes, change their](#)  
747 [physical and chemical properties in time due to aging and photo- or chemical-decay processes](#)  
748 [\(Andrady 2011\), are subject to biofouling \(see recent relevant work by Kreczak et al., 2021\), and](#)  
749 [may interact leading to the formation of clusters. None of these effects were considered in this](#)  
750 [paper, and all will need to be taken into account for the realistic prediction of marine](#)  
751 [microplastic evolution and re-distribution in the ocean. Real ocean eddies are also decaying in](#)  
752 [time and are usually moving \(translating\) rather than stationary. Translation with a constant](#)  
753 [velocity can be handled by considering the flow in a moving frame of reference, but decay and](#)  
754 [interactions will likely change the geometry of the circulation and make the flow truly aperiodic.](#)  
755 [Our simplified model cannot account for these effects, which will need to be explored separately](#)  
756 [later.](#)

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760 simulations using an analytical vortex model are provided in text.

761 **Author Contribution Statement:** IR led the overall effort and performed most of the numerical  
762 simulations, LP contributed towards the theoretical understanding and interpretation of the  
763 results, MD participated in the overall effort.

764 **Competing interests Statement:** no competing interests

765

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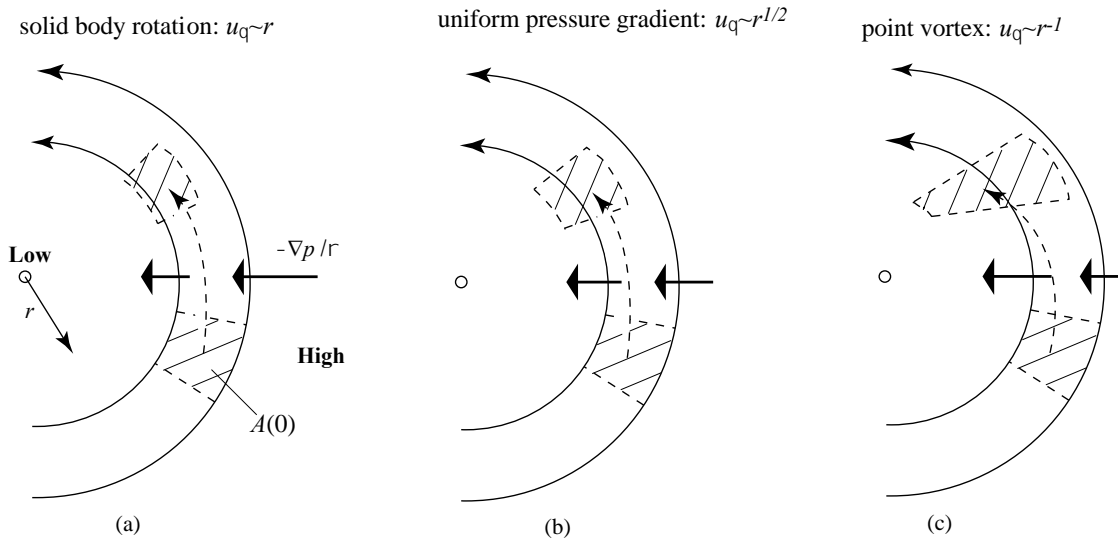
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Experiment	$a$	$b$	$\varepsilon$	$y_o$	$\sigma$	$\gamma$	$\beta$	$\Omega$	d
1 – steady symmetric	0.62	7.5	0	0	0	0	0	0	$10^{-3}$
1a (small $\Omega$ )	0.62	7.5	0	0	0	0	0	0.3	$10^{-3}$
1b (large $\Omega$ )	0.62	7.5	0	0	0	0	0	1	$10^{-3}$
1c (slow overturn)	0.62	0.25	0	0	0	0	0	1	$10^{-3}$ vs. $5 \times 10^{-4}$
1d ( $z_{attractor}$ vs $\Omega$ )	0.62	7.5	0	0	0	0	0	Sweep 0 to 10	$10^{-3}$
1e (reversed flow)	-0.62	-7.5	0	0	0	0	0	0, 0.6, 2	$10^{-3}$
2 – steady asymmetric	0.62	7.5	0.25	-0.2	0	0	0	0	$10^{-3}$
2a (small $\Omega$ )	0.62	7.5	0.25	-0.2	0	0	0	0.3	$10^{-3}$
2b (large $\Omega$ )	0.62	7.5	0.25	-0.2	0	0	0	1	$10^{-3}$
2c (2:1 resonance)	0.62	3.8	0.25	-0.2	0	0	0	0	$10^{-3}$
3 - non-steady asymmetric	0.62	7.5	0.25	-0.2	$\frac{2\pi}{9.1}$	0.2	1	0	$10^{-3}$

891 Table 1: Dimensionless parameter values for numerical experiments. Fixed parameters in the  
892 kinematic model (Eqs. 9a-c) are  $c = 0.69$ , and  $r_0 = 1/2$  in all cases. Parameters that appear in  
893 the nondimensional Maxey-Riley Eq. (3) are also nondimensional, with  $L$ ,  $U$ ,  $L/U$  as length,  
894 velocity and time scales. Fixed parameter values based on  $L = 1\text{m}$  and  $U = 1\text{m/s}$  include

895  $\frac{\rho_p}{\rho_f} = 0.97$ ,  $R = \frac{2\rho_f}{\rho_f + 2\rho_p} = 0.680$ ,  $\frac{3R}{2} - 1 = .020$   $\vec{g}_r = \frac{gL}{U^2} = 10.0$ ,  $\tilde{\varepsilon} = \frac{2}{9} \left(\frac{d}{L}\right)^2 \frac{UL}{\nu R} = 0.33$ , and

896  $\tilde{\varepsilon} \left(\frac{3R}{2} - 1\right) = 0.0067$ . Note that  $\vec{\Omega} = \Omega \vec{k} = \frac{\vec{\Omega}^* L}{U}$ .



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898 Figure 1. Three types of two-dimensional eddies with zero frame rotation and for which gravity  
 899 is imagined to be zero: solid body rotation (a), constant pressure gradient (b), and point vortex  
 900 (c). In each case, the cross hatched area represents a concentration of rigid particles with area  
 901  $A(t)$ .

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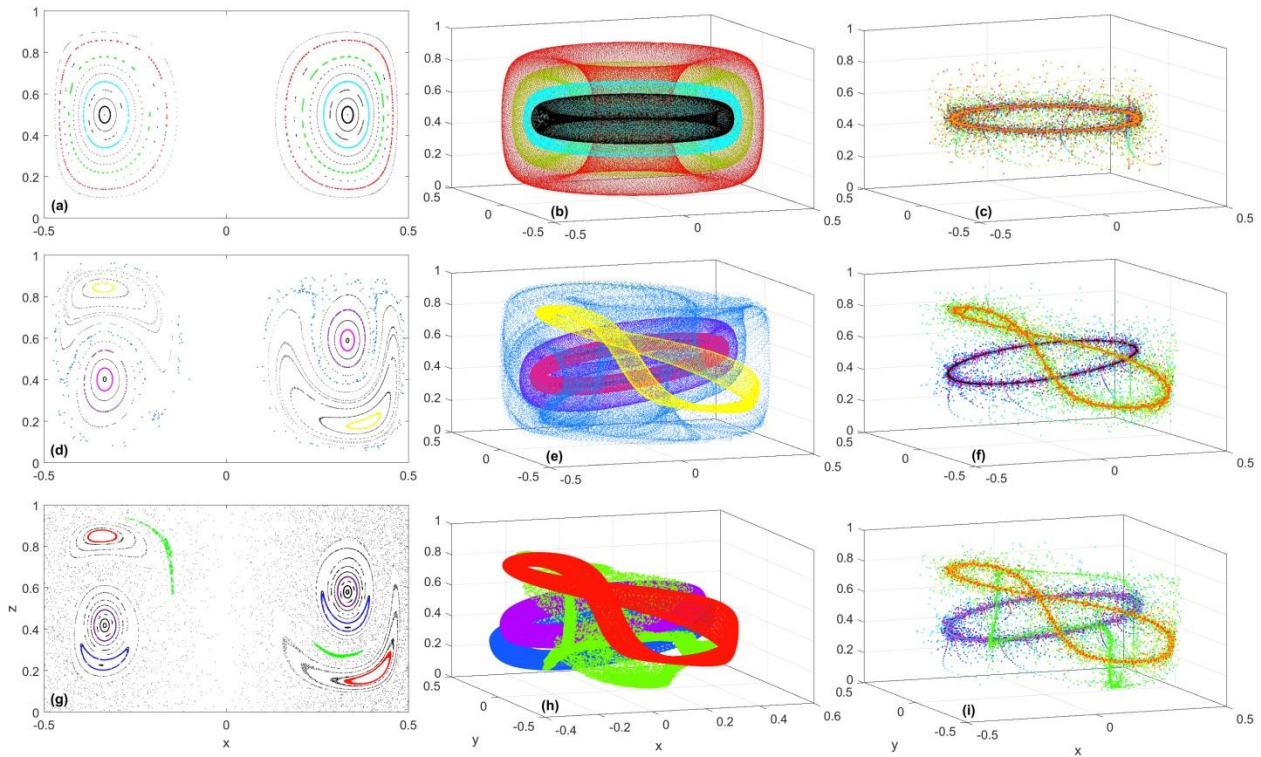
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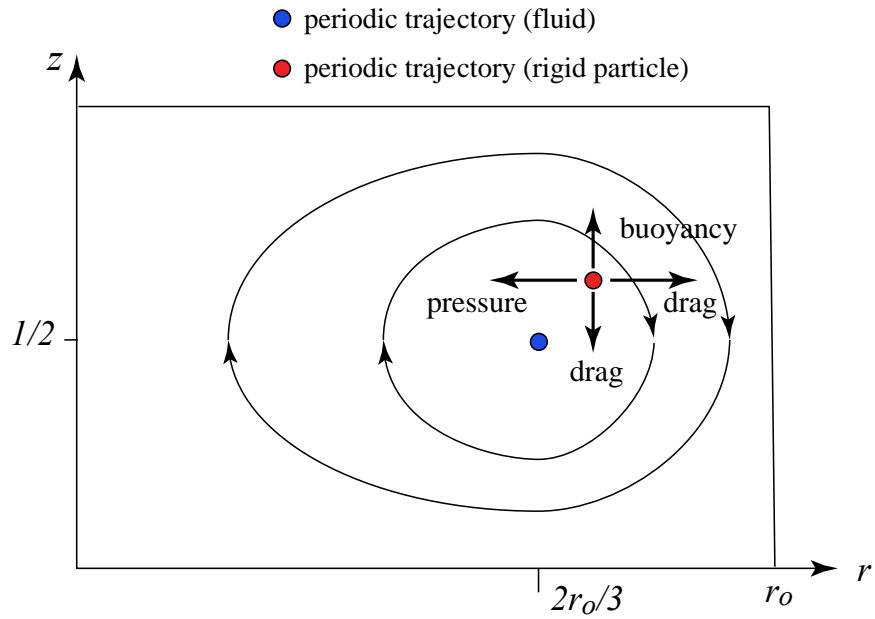
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911 Figure 2. (left) Poincare section, (middle) fluid parcels trajectories in 3D, (right) buoyant particle  
 912 trajectories in 3D for a steady symmetric fluid flow (top row), steady asymmetric flow (middle  
 913 row), and non-steady, asymmetric flow. Parameter setting are listed under Experiments 1, 2 and  
 914 3 in Table 1. Colors in the left column of panels match the corresponding panel in the middle  
 915 column, but the colors in the right column indicated time after release of the particles. Note the  
 916 attraction of buoyant particles to a single attractor at mid-depth in panel (c), to 2 attractors in  
 917 panel (f), and to 3 attractors in panel (i). Particles are released along a vertical line  $x = 0.334$ ,  
 918  $y = 0$ ,  $0 < z \leq 0.6$  with initial velocity equal to that of the co-located fluid parcels.

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923 Figure 3. Sketch showing the position in a vertical section of the periodic orbit (red dot) of the  
 924 rigid particle relative to the periodic orbit (blue dot) of the fluid flow. The viewer sees one half  
 925 of a vertical slide through the cylinder, with the azimuthal flow directed away from the viewer  
 926 and the cylinder center at the left edge.

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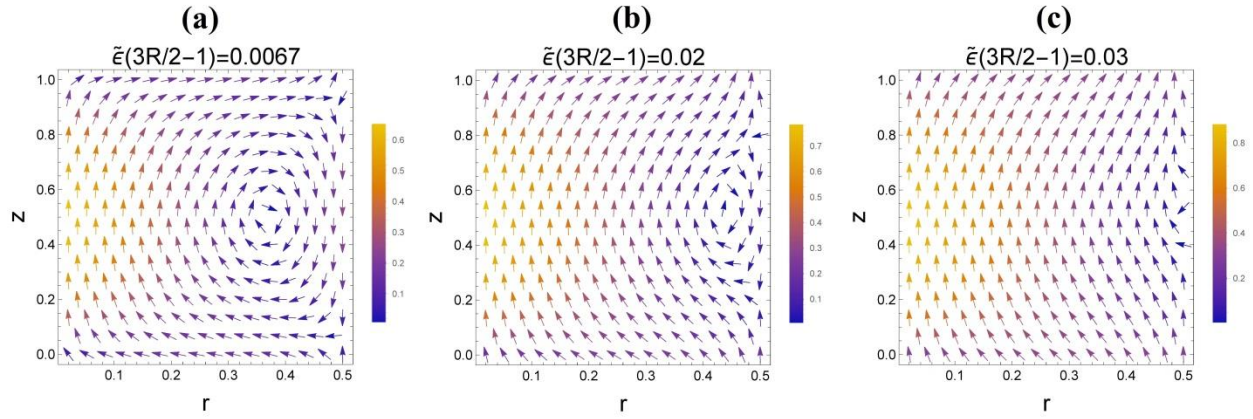
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935 Figure 4. The slow-manifold radial and vertical velocity components for the rigid particles,

936 plotted in the  $(r, z)$  plane for (a)  $\tilde{\epsilon} \left( \frac{3R}{2} - 1 \right) = 0.0067$ , (b) = 0.02, and (c) = 0.03. Other

937 parameters are as listed for Experiment 1a in Table 1.

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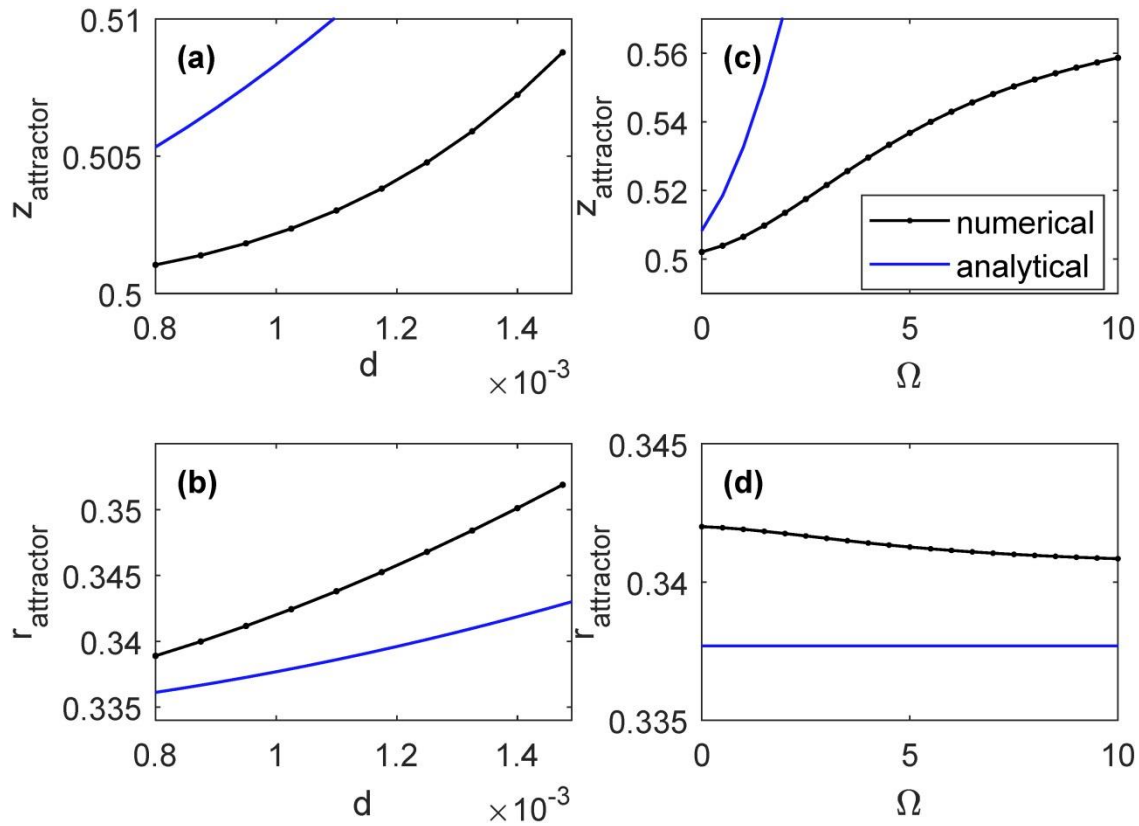
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949 Figure 5. For the steady symmetric rotating cylinder flow, the coordinates of the periodic orbit  
 950 that acts as an attractor for buoyant particles as a function of particle diameter (a-b) and frame  
 951 rotation (c). Flow parameters are listed in Table 1 and correspond to Experiment 1 for (a-b) and  
 952 Experiment 1d for (c-d).

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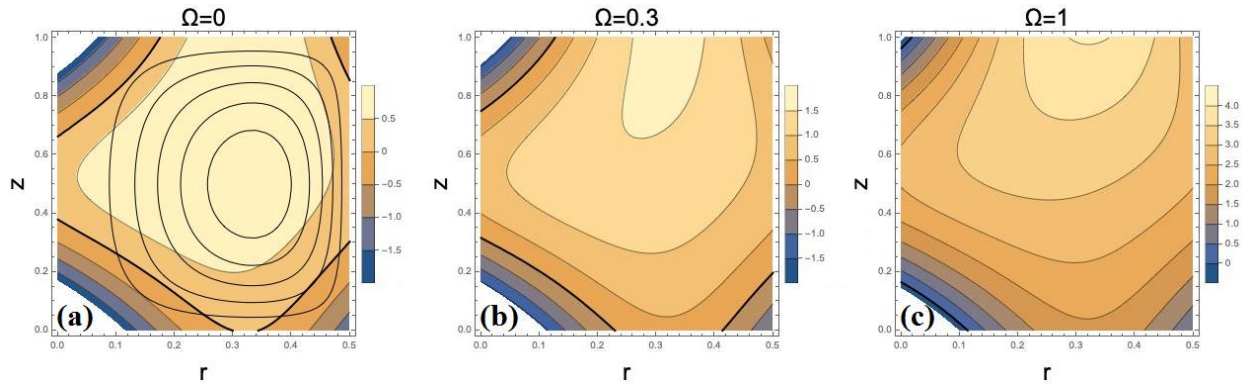
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959 Figure 6. (a): The  $Q_a$  function for the steady, axisymmetric, cylinder flow with the same  
 960 parameter setting (see Experiment 1a) as for Figure 3a-c, and plotted in  $(x,z)$  along with the  
 961 streamlines of the overturning circulation. The thick rigid curve corresponds to  $Q_a = 0$ . (b): The  
 962 same parameter settings, except  $\Omega$  has been raised from 0 to 0.3 (Rossby number  $\cong 1$ ) (c):  
 963  $\Omega = 1.0$ . (Rossby number  $\cong 0.2$ ).

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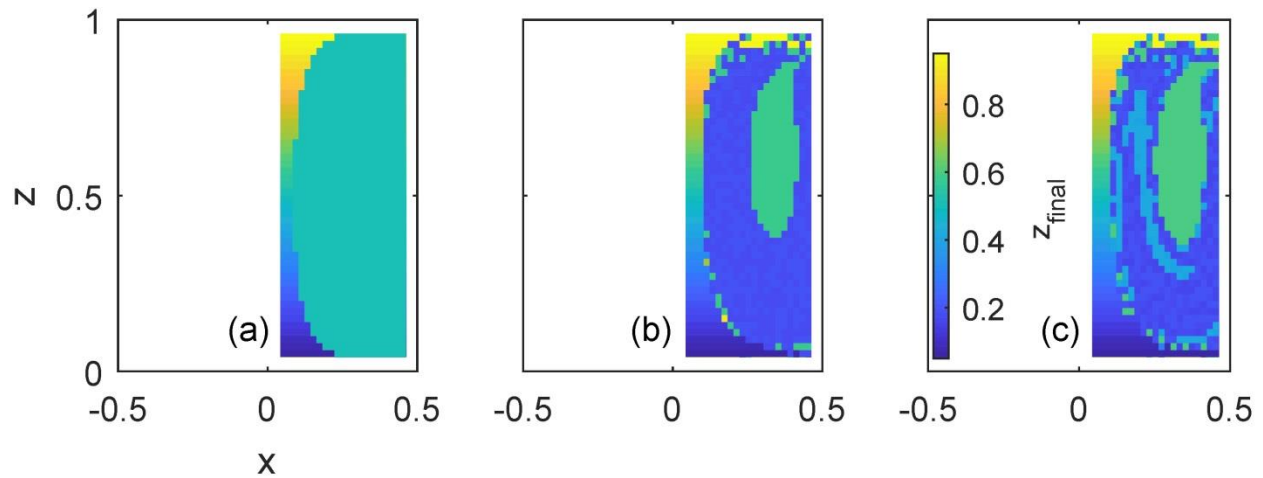
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974 Figure 7. Domain of attraction for the attractors in (a) steady symmetric (Experiment 1 in Table  
 975 1), (b) steady asymmetric (Experiment 2 in Table 1), and (c) time-periodic asymmetric rotating  
 976 cylinder flow (Experiment 3 in Table 1). (These are the same 3 experiments that were used to  
 977 produce Fig. 2.) The color indicates the height (i.e., value of  $z$ -coordinate) of the final crossing of  
 978 a trajectory with the Poincare section, as a function of particle's release location. Particles  
 979 attracted to the same attractor thus correspond to same color.

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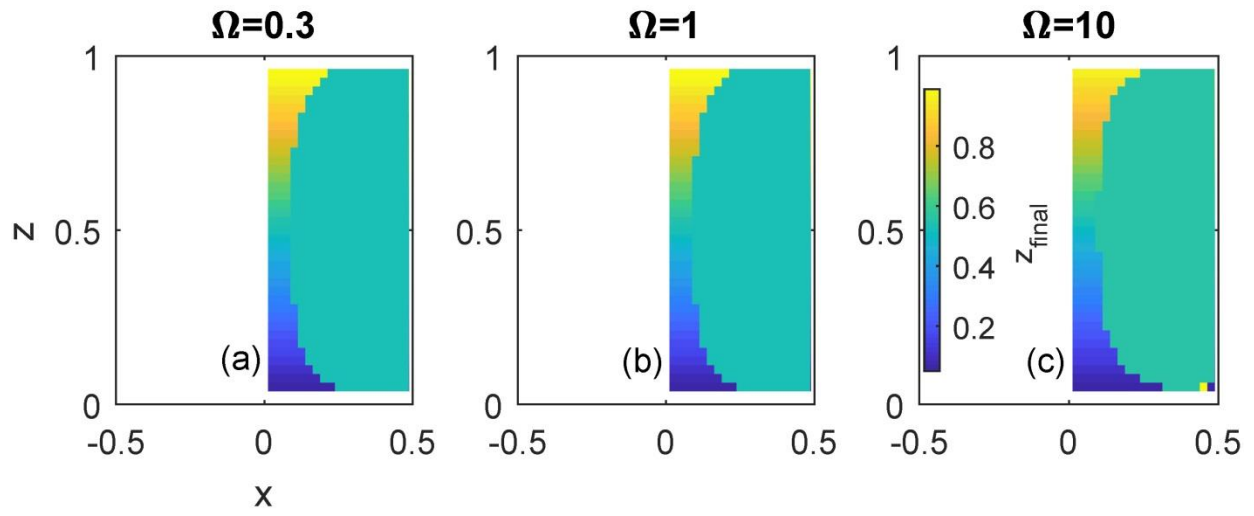
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988 Figure 8. Same as in Fig. 7a but with frame rotation.

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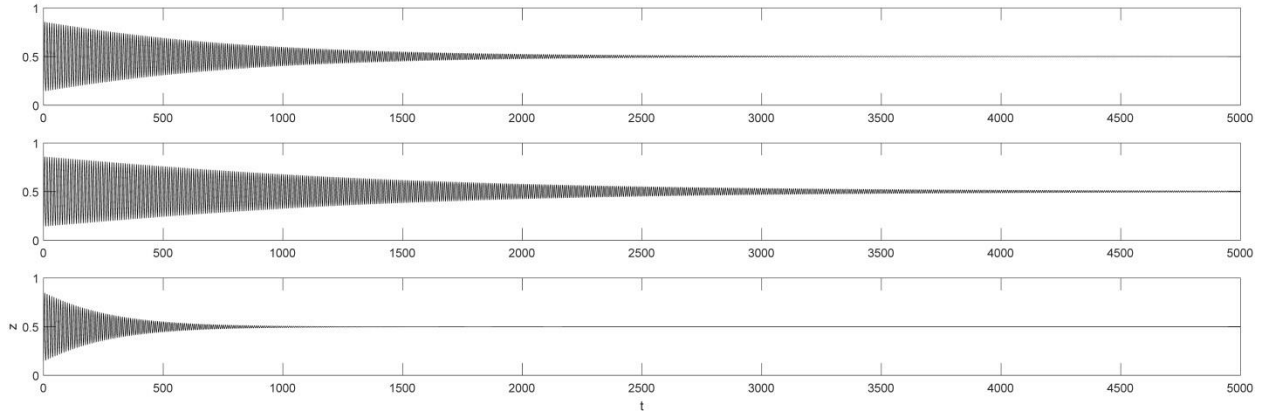
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1000 Figure 9. For the “reversed flow” experiment (Experiment 1e in Table 1), z-position of a sample  
 1001 particle trajectory as function of time for 3 values of  $\Omega$ : 0 (top), 0.6 (middle), and 2 (bottom).  
 1002 Time  $t$  is in dimensionless units (but since our scaling coefficient for time is equal to 1 sec, the  
 1003 numbers on the x-axis can also be read as dimensional time in sec.)

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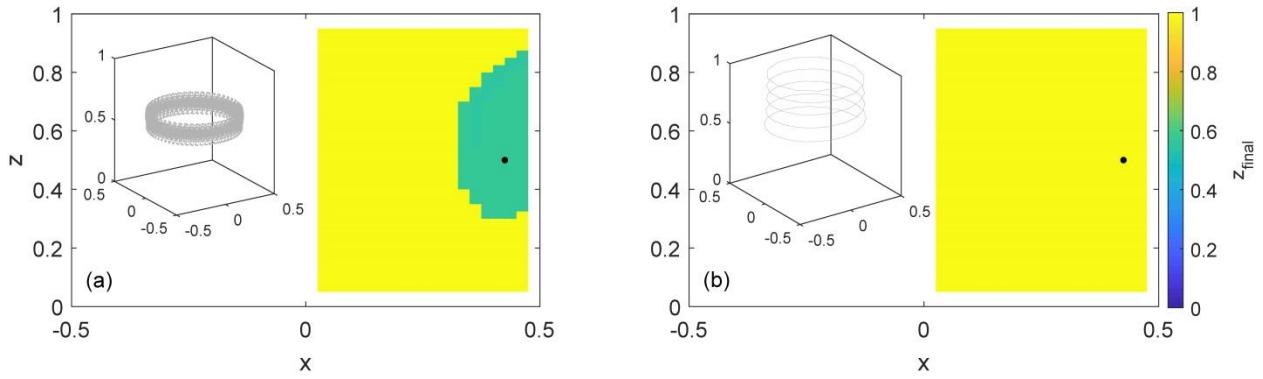
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1014 Figure 10. For the “slow overturn” Experiment 1c from Table 1, color indicates the final z-  
 1015 coordinate of a particle’s trajectory at the end of integration time as a function of particle’s  
 1016 release location for 2 values of  $d$ : (a)  $5 \times 10^{-4}$  and (b)  $10^{-3}$ . Yellow corresponds to particles  
 1017 rising up to the top, whereas green indicates the basin of attraction of the subsurface attracting  
 1018 periodic orbit. The insets at the left side of each frame show a sample trajectory whose release  
 1019 location is indicated by the black dot.

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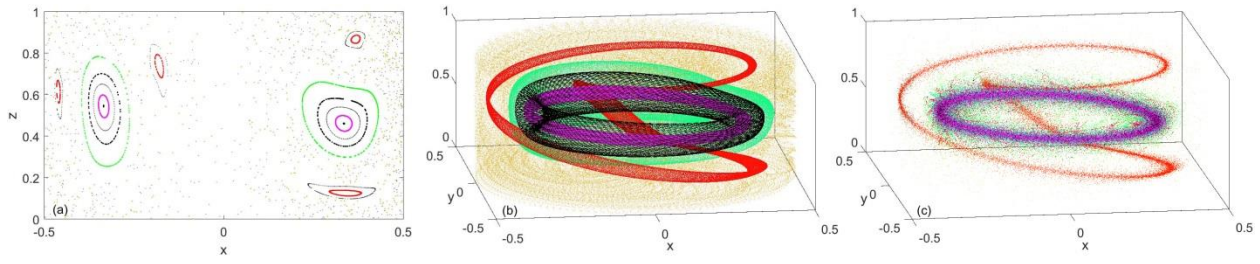
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1030 Figure 11. Same as Fig. 2(d-f) but with  $b = 3.8$ .

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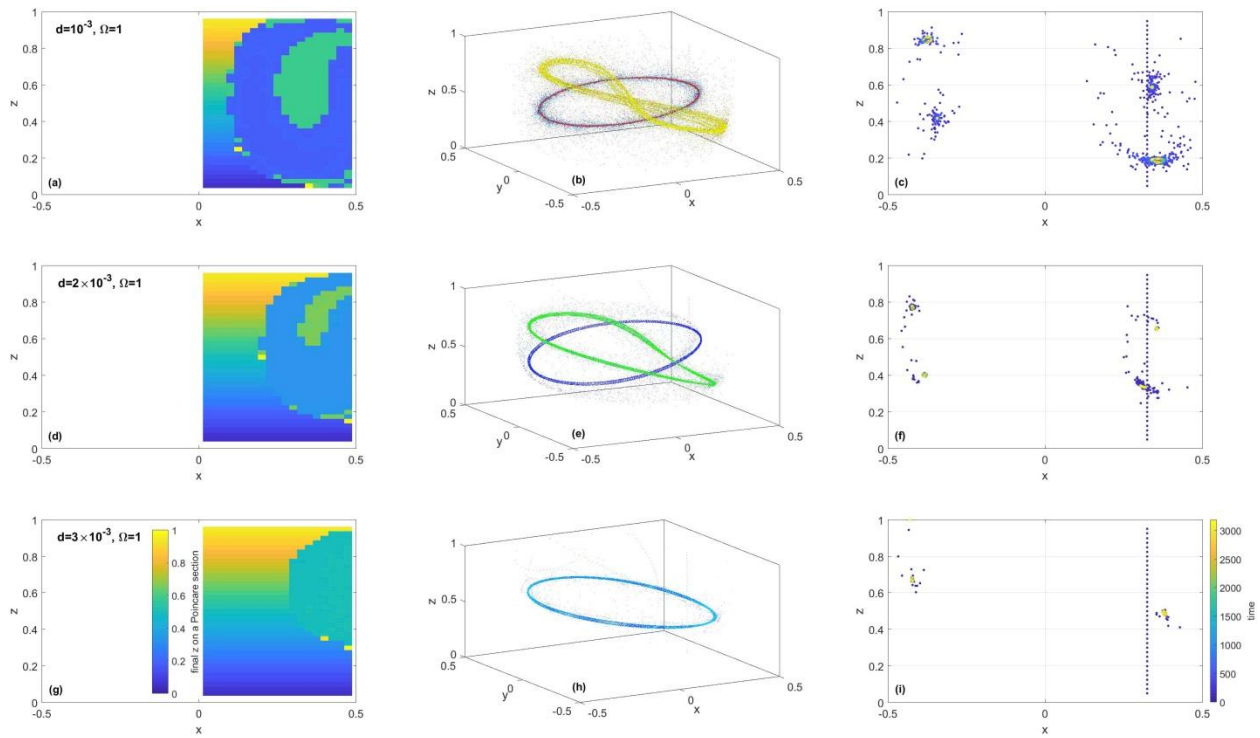
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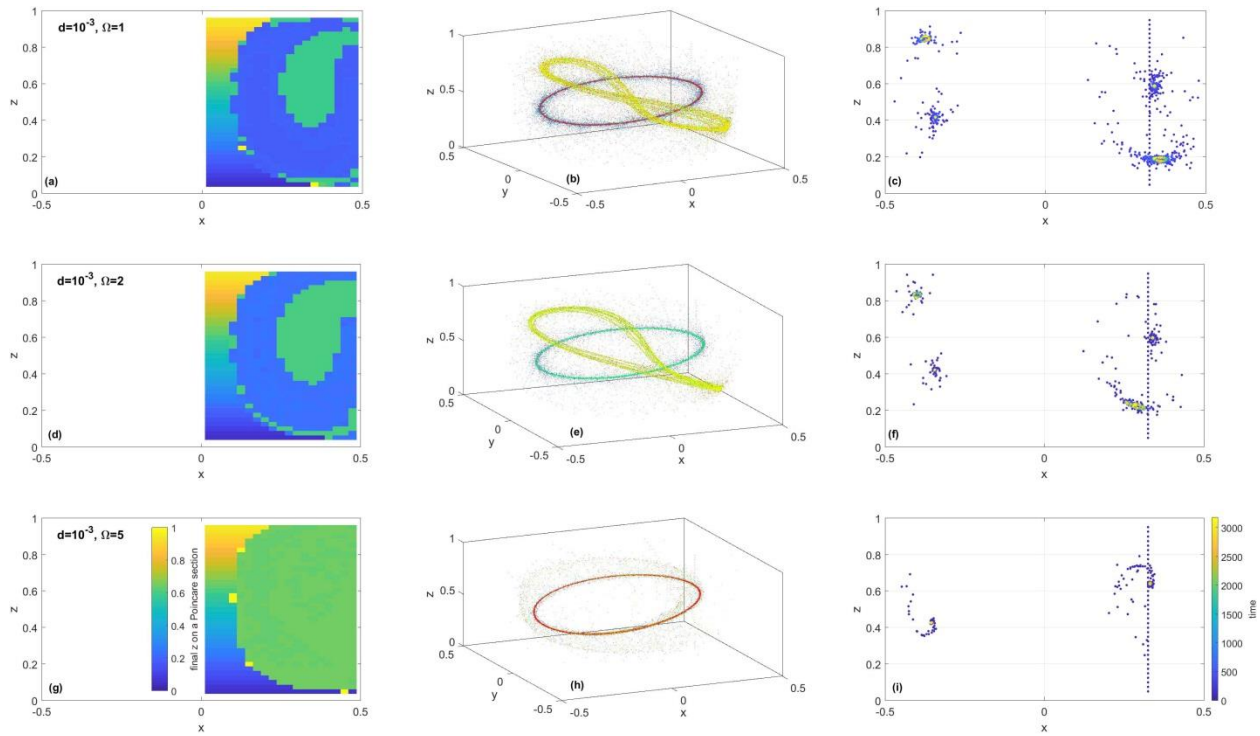
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1044 Figure 12. For the steady perturbed system (Experiment 2 in Table 1), changes in the location of  
 1045 the attracting periodic orbits, basins of attractions, and time of attraction as a function of particle  
 1046 diameter  $d$  (and thus  $\tilde{\epsilon}$ ). (a,d,g) show  $z$ -coordinate of the last crossing of trajectory with the  $x$ - $z$   
 1047 Poincare plane as a function of release location; flat regions are basins of attraction for the 2  
 1048 attractors. (b,e,h) show 20 trajectories in 3d released along a vertical line at  $y = 0, x = 0.334,$   
 1049  $0.05 < z < 0.95$ ; denser cores indicate attractors. (c,f,i) show crossing of the same select 20  
 1050 trajectories with the  $x$ - $z$  Poincare plane, color coded by time; blue corresponds to release  
 1051 location, yellow corresponds to final positions.

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1056 Figure 13. For the steady perturbed system (Experiment 2 in Table 1), changes in the location of  
 1057 the attracting periodic orbits, basins of attractions, and time of attraction as a function of frame  
 1058 rotation  $\Omega$ . (a,d,g) show z-coordinate of the last crossing of trajectory with the x-z Poincare plane  
 1059 as a function of release location; flat regions are basins of attraction for the 2 attractors. (b,e,h)  
 1060 show 20 select trajectories in 3d released along a vertical line at  $y = 0, x = 0.334, 0.05 < z <$   
 1061  $0.95$ ; denser cores indicate attractors. (c,f,i) show crossing of the same 20 trajectories with the x-  
 1062 z Poincare plane, color coded by time; blue corresponds to release location, yellow corresponds  
 1063 to final positions.

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