1	Aggregation of Slightly Buoyant Microplastics in Three-Dimensional Vortex Flows
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19 Abstract

20 Although the movement and aggregation of microplastics at the ocean surface has been well studied, less is known about the subsurface. Within the Maxey-Riley framework governing the 21 22 movement of small rigid spheres with high drag in fluid, aggregation of buoyant particles is 23 encouraged in vorticity-dominated regions. We explore this process in an idealized model that is qualitatively reminiscent of a three-dimensional eddy with an azimuthal and overturning 24 circulation. In the axially symmetric state, buoyant spherical particles that do not accumulate at 25 the top boundary are attracted to a loop consisting of periodic orbits. Such a loop exists when 26 drag on the particle is sufficiently strong. For small slightly-buoyant particles, this loop is located 27 close to the periodic fluid parcel trajectory. If the symmetric flow is perturbed by a symmetry-28 29 breaking disturbance, additional attractors for small rigid slightly-buoyant particles may arise near periodic orbits of fluid parcels within the resonance zones created by the disturbance. 30 31 Disturbances with periodic or quasi-periodic time dependence may produce even more attractors, with a shape and location that recurs periodically. However, not all such loops attract, and rigid 32 particles released in the vicinity of one loop may instead be attracted to a nearby attractor. 33 Examples are presented along with mappings of the respective basins of attraction. 34

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40 Significance statement

This paper investigates the phenomenon of aggregation of small, spherical, slightly buoyant, finite size, rigid particles in a simple analytically-prescribed three-dimensional vortex flow. Our goal was to gain insights into the behaviour of slightly buoyant marine microplastic particles in a flow that qualitatively resembles ocean eddies. Attractors are mapped out for the steady axisymmetric, steady asymmetric, and non-steady asymmetric vortices over a range of flow and particle parameters. Simple theoretical arguments are used to interpret the results.

I. Introduction

59 Marine microplastic pollution has been a rising concern for the ocean environmental and for human health. Microplastics (scales < 5mm) and nanoplastics (scales $< 1 \mu$ m) have been found 60 in the tissues of marine animals, some of which are consumed by humans (Landrigan, et al. 61 2023). This comes at a time when global production of plastics is projected to increase. 62 63 Observations of marine microplastics have been conventionally carried out using net tows and mostly occurred at or near the sea surface (van Sebille et al., 2015). However, the density of 64 many types of microplastic particles, including high-density polyethylene, is sufficiently close to 65 that of sea water that suspension within the water column for long periods of time is feasible. For 66 the near-surface microplastics, Kukulka et al. (2010) and Kooi et al. (2016) present observational 67 68 evidence for the fast decay in concentrations with depth over the top 5 - 20 m of the water column, with the vertical penetration of plastic particles dependent on the wind speed. 69 Pabortsava and Lampitt (2020), on the other hand, show observational evidence for much deeper, 70 71 below-the-mixed-layer subsurface peaks for three common types of microplastics in the Atlantic 72 Ocean. Processes such as biofouling and bio-geo-chemical or photo degradation might increase 73 the density of the plastic particles and eventually lead to the sinking of microplastics from the 74 surface into the deeper part of the water column (Kaiser et al., 2017; Kreczak et al., 2021; Kvale et al., 2020). Consumption by biomass with the subsequent downward vertical transport is 75 76 another vehicle for redistributing microplastics from the surface down. For example, Choy et al. 77 (2019) suggest that this mechanism, specifically, consumption by pelagic red crabs and giant larvaceans, was responsible for the subsurface peaks in plastic particles concentrations observed 78 at depths near 250 m in Monterey Bay. Thus, microplastics have been found well beneath the 79

80 <u>ocean surface, but less is known regarding their spatio-temporal and size/density distributions</u>
81 (Shamskhany et al., 2021).

A potentially important aspect of the movement of plastics and microplastics is aggregation, a 82 process that occurs at the surface over large scales near the centers of the five major subtropical 83 gyres and has been attributed to Ekman drift, windage and inertia (Beron-Vera, 2021). Many 84 85 early models concentrated on the ocean surface, but Wichmann et al. (2019) has highlighted the importance of resolving the full three dimensional circulation. If aggregation also occurs below 86 the surface, well beneath the direct influence of Ekman layers, the dynamics is likely to be 87 88 different. Indeed, modeling results by Wichmann et al. (2019), based on a framework created by Lange and van Sebille (2017) and Delandmeter and van Sebille (2019), suggests that the large 89 scale accumulation associated with the garbage patches disappears below 60m depth. 90

91 To avoid confusion, we will refer to infinitesimal fluid elements as "fluid parcels", and to rigid 92 plastic particles of finite size as "rigid particles". Typically the position $x_p(t)$ of a rigid particle 93 is tracked according to

94
$$x_p(t + \Delta t) = x_p(t) + \int_t^{t+\Delta t} u dt + dx_b$$

1

where *u* is the fluid velocity and dx_b is an extra displacement due the non-fluid nature of the rigid particle. The user can introduce custom schemes for calculating contributions to dx_b due to factors such as windage and inertia (e.g. Beron-Vera et al., 2016), turbulent diffusion (e.g. Kulkulka, 2012), wave induced Stokes drift (Onink et al., 2019), etc. Eulerian schemes in which plastic particles are treated as concentrations, are rare, but Mountford and Morales Maqueda (2019) developed an Eulerian model in which concentrations are advected by the fluid and are subject to parameterized turbulence as well as sinking or rising according to a simple law involving buoyancy and friction. <u>In a similar fashion, Kvale et al. (2020) propose an Eulerian</u>
 model for the biological uptake and the resulting re-distribution of microplastics.

An alternative approach would be to use the Maxey-Riley equation (discussed below) to solve 104 for the rigid particle velocity, v, and then use the latter to compute the trajectory of that rigid 105 particle, i.e., $x_p(t + \Delta t) = x_p(t) + \int_t^{t+\Delta t} v dt$. This equation would account for the non-fluid-106 following effects in a deductive way, however the resulting 6th-order system (for the three 107 components of velocity and position) would be computationally challenging. 108 To better understand the implications of the use of this approach while avoiding the computational burden 109 and complexity, we have elected to analyze the movement and aggregation of individual rigid 110 111 particles using a Maxey-Riley framework in connection with an idealized, analyticallyprescribed, 3D vortex flow that qualitatively resembles the geometry of the circulation in an 112 ocean eddy but is not a solution to any dynamical oceanographic equations of motion. As shown 113 by Pratt (2014) and Rypina et al. (2015), kinematic models that reproduce the correct geometry 114 are able to also reproduce the important Lagrangian features of the flow. Even in our simple 115 flow, aggregation is non-trivial, often with multiple attractors present and lack of attraction in 116 some circumstances. Thus, we wanted to thoroughly explore this simple example before 117 investigating more realistic oceanic flows. We note that other idealized studies have been carried 118 119 out in connection with 2D wave fields and vortex flows (e.g. DiBenedetto 2018a,b and Kelly et al., 2021). 120

Aggregation can be attributed to the presence of an attractor: here, an object with a dimension less than three that is somehow set up by the fluid circulation patterns and towards which rigid particle trajectories attract. As long as the fluid is incompressible, fluid parcels will not experience attraction and will not aggregate, but plastic particles with inertia, added mass, and drag may do so. Note also that because each attractor is generally associated with its
 corresponding basin of attraction, if rigid particles are introduced outside of the basin of
 attraction, they will not be attracted and will not aggregate towards this attractor.

In order to reach a better understanding of what leads to attraction and attractors in 3D flows, we 128 129 explore a simple canonical example in geophysical fluid dynamics, namely the flow in a rotating 130 cylinder. This flow resembles some of the characteristics of ocean eddies, including a horizontal swirl and an overturning component in the vertical, but is much less complex than any realistic 131 oceanic eddy. Specifically, we use a simple analytically-prescribed phenomenological velocity 132 133 introduced by Rypina et al. 2015. The Lagrangian properties of this circulation have been previously studied (Fountain, et al. 2000; Pratt et al. 2014; Rypina et al. 2015) allowing us to 134 begin to investigate inertial rigid particles from an established base of knowledge. A prior theory 135 (Haller and Sapsis, 2008) governing the movement of rigid particles with high drag indicates that 136 accumulation is favored for slightly buoyant particles in flows dominated by vorticity, and this 137 138 also motivates our choice of background flow. Identification of the attractors that can arise in this 139 flow field, evaluating their reach and domains of attraction, and clarifying the circumstances that lead to their formation are the primary objectives of this work. Although motivated by the 140 141 problem of marine microplastics, this study is, for now, mainly a curiosity-driven research aiming to develop a basic understanding of the mechanisms that might lead to aggregation of 142 143 rigid particles in 3D flows. The hope is that with such basic understanding in hand, one could 144 later start investigating aggregation phenomena in more complex and more realistic ocean mesoscale and submesoscale eddying flows. 145

146 II. Methods

The physics of the motion of a small, rigid sphere that moves with velocity $\vec{v}(t)$ through a fluid with pre-existing velocity distribution $\vec{u}(\vec{x}, t)$ has been the subject of investigation by Stokes (1851), Basset (1888), Boussinesq (1903), Faxen (1922), Oseen (1927), Tchen (1947) and many others, and was put in a unifying framework by Maxey and Riley (1983). More recent theoretical extensions include Beron-Vera et al. (2019) and Beron-Vera (2021). We will use a form of the Maxey-Riley equation that has been extended to include constant frame rotation with angular velocity $\vec{\Omega}^*$:

154
$$\frac{d\vec{v}}{dt} = \frac{\rho_f}{\rho_p} \frac{D\vec{u}}{Dt} + \frac{\rho_f}{2\rho_p} \left(\frac{D\vec{u}}{Dt} - \frac{d\vec{v}}{dt}\right) - \frac{9\nu\rho_f}{2\rho_p d^2} (\vec{v} - \vec{u}) + \left(1 - \frac{\rho_f}{\rho_p}\right) \vec{g} + \frac{\rho_f}{\rho_p} \vec{\Omega}^* \times (\vec{u} - \vec{v})$$

155
$$+ \frac{\rho_f}{\rho_p} 2\vec{\Omega}^* \times \vec{u} - 2\vec{\Omega}^* \times \vec{v} + \left(\frac{\rho_f}{\rho_p} - 1\right) \vec{\Omega}^* \times \vec{\Omega}^* \times \vec{x} .$$
(1)

156 The frame rotation was introduced into the non-rotating Maxey-Riley equation by replacing

157
$$\overrightarrow{v_s} = \overrightarrow{v_r} + \overrightarrow{\Omega} \times \overrightarrow{x_r}, \ \overrightarrow{u_s} = \overrightarrow{u_r} + \overrightarrow{\Omega} \times \overrightarrow{x_r},$$

158
$$\frac{D_{S}\overline{u_{S}}}{Dt} = \frac{D_{r}\overline{u}_{r}}{Dt} + 2\,\overrightarrow{\Omega}\times\overrightarrow{u}_{r} + \overrightarrow{\Omega}\times\overrightarrow{\Omega}\times\overrightarrow{x}_{r}, \ \frac{d_{S}\overline{v_{S}}}{Dt} = \frac{d_{r}\overline{v}_{r}}{Dt} + 2\,\overrightarrow{\Omega}\times\overrightarrow{v}_{r} + \overrightarrow{\Omega}\times\overrightarrow{\Omega}\times\overrightarrow{x}_{r}.$$

- 166 <u>complex tangled-filament-like shapes which are poorly represented by an ellipsoid, and no</u>
- 167 <u>corrections for tangled filaments are currently available.</u>

In Eq. (1), which is a statement of Newton's second law for the rigid particle, the right-hand side 168 represents, in order, the effects of inertia, added mass, drag, buoyancy, Coriolis acceleration 169 170 associated with the added mass, the Coriolis acceleration associated with the particle mass, 171 Coriolis acceleration associated with the fluid motion, and centrifugal acceleration. A similar equation has been previously derived by Beron-Vera et al. (2019), though the centrifugal 172 acceleration does not appear there explicitly, having been combined with the acceleration due to 173 gravity in order to define an effective gravity and corresponding geopotential. Coordinates are 174 then imagined to be aligned with geopotential surfaces, though standard spherical or Cartesian 175 176 coordinates are usually used in practice (Vallis, 2006). Our explicit retention of the centrifugal 177 acceleration will later allow absolute vorticity to arise naturally as a quantity of central 178 importance. We have omitted the lift force, the Basset history force, and the Faxen corrections (Gatignol, 1983). Faxen corrections account for the variation of the flow across the rigid particle 179 and are proportional to $a^2\Delta u$. For a particle size that is much smaller than the typical length scale 180 of the flow, these corrections are small and typically neglected (Haller and Sapsis, 2008; Beron-181 182 Vera et al., 2019). The history term, which is an integral along a particle path, accounts for the boundary layer effects that a particle leaves behind. It is typically ignored under the assumption 183 184 that the chances of other particles crossing that localized boundary layer before it decays are small (Beron-Vera et al., 2019; see also Langlois et al., 2015 and Daitche and Tel., 2011 for 185 more info on the influence of the history term on the behavior of rigid particles). Finally, the lift 186 force arises when a particle rotates in a horizontally sheared flow. As shown in Beron-Vera 187 2019, the inclusion of the lift force leads to the next-order, $O(\tilde{\epsilon}^2)$ correction in the slow-188

189 manifold approximation, and thus can also be neglected for small $\tilde{\epsilon}$. In Eq. (1), ρ_p and ρ_f are

- 190 densities of the rigid particle and the fluid, d is the particle radius, v is viscosity of the fluid, \vec{g} is
- 191 the gravity vector, and $\frac{D\vec{u}}{Dt} = \frac{\partial\vec{u}}{\partial t} + \vec{u} \cdot \nabla\vec{u}$ is the fluid material derivative, evaluated for

192 undisturbed fluid velocity at the position of the center of the rigid particle. The position x(t) of a 193 particle is determined by

194
$$\frac{d\vec{x}}{dt} = \vec{v}(\vec{x}, t), \tag{2}$$

and together <u>Eqs.</u> (1) and (2) compose a coupled, 6th-order system for computation of the particle
position and velocity as functions of time.

197 If the velocities and lengths are nondimensionalized using characteristic scales U and L for the 198 background fluid flow, and L/U is used as a time scale, then Eq. (2) remains formally unchanged 199 while the nondimensional form of Eq. (1) is

$$200 \qquad \frac{d\vec{v}}{dt} = \frac{3R}{2} \frac{D\vec{u}}{Dt} + \tilde{\varepsilon}^{-1} (\vec{v} - \vec{u}) + \left(1 - \frac{3R}{2}\right) \vec{g}_r + 3R\vec{\Omega} \times (\vec{u} - \vec{v}) + 2\left(\frac{3R}{2} - 1\right) \vec{\Omega} \times \vec{v}, \tag{3}$$

201 where
$$=\frac{2\rho_f}{\rho_f+2\rho_p}$$
, $\vec{g}_r = (\vec{g} - \vec{\Omega}^* \times \vec{\Omega}^* \times \vec{x})/(\frac{U^2}{L})$, $\vec{\Omega} = \frac{\vec{\Omega}^*L}{U}$ and $\vec{\varepsilon} = \frac{2}{9} \left(\frac{d}{L}\right)^2 \frac{UL}{v} \frac{1}{R}$ is the Stokes

number, the ratio of the adjustment time scale of a particle (due to drag) to the time scale of the background flow. For $\tilde{\varepsilon} \ll 1$, viscous drag is the dominant force acting on the particle, implying that a particle with an initial velocity differing by an amount > $O(\tilde{\varepsilon})$ from the local fluid velocity will be rapidly accelerated over a time scale $\tilde{\varepsilon}$ to a velocity proximal to that of the fluid. Thereafter the particle will undergo a slow evolution in which the weaker forces due to inertia, added mass, and buoyancy cause slight departures from the movement of the fluid itself. The limit $\tilde{\varepsilon} \to 0$ constitutes a singular perturbation of Eq. (3), a problem that can be addressed using an approach due to Fenichel (1979) that was originally formally developed for a steady background flow, but that has been extended by Haller and Sapsis (2008) to include a timevarying background flow. In either case, it can be shown that following the initial viscous adjustment, the particle position and velocity tend toward a subspace or "slow manifold" on which the particle velocity is determined directly by the fluid velocity through an "inertial" equation, here extended to include frame rotation:

215
$$\vec{v} = \vec{u} + \tilde{\varepsilon} \left(\frac{3R}{2} - 1\right) \left[\frac{D\vec{u}}{Dt} + 2\vec{\Omega} \times \vec{u} - \vec{g}_r\right] + O(\tilde{\varepsilon}^2).$$
 (4)

216 This result is the same as obtained by Beron-Vera et al. 2019, provided that their gravity vector 217 is interpreted as our \vec{g}_r . The same authors also present more general cases, including those with 218 the lift force and on the sphere. In Supplementary Material we present a simple derivation of Eq. 219 (4) based on a multiple-scale expansion. It provides a quick, though less rigorous, alternative to 220 the Fenichel approach.

- A chief advantage of the slow manifold reduction is that the 6^{th} order system given by Eqs. (2)
- 222 and (3), in which particle velocity needs to be solved for, is reduced to a 3rd order system given
- 223 by Eqs. (2) and (4), where the particle velocity is explicitly written as a function of fluid velocity
- 224 and flow and particle parameters (and thus is known). The bracketed expression in Eq. (4), which
- determines the velocity of the rigid particle relative to the fluid, is nothing more than $\frac{\partial}{\partial x_i} \tau_{ij}$,
- where τ_{ij} is the stress tensor for the fluid. Thus the relative velocity of a rigid particle on the
- slow manifold is in the same direction as the net force that would act on a fluid parcel occupying
- the same space. Ordinarily, for a fluid parcel, that force would equate with an acceleration, but
- on the slow time scale, the relative particle velocity points in the same direction as the net fluid

force and its magnitude is proportional to $\tilde{\varepsilon}\left(\frac{3R}{2}-1\right) = \frac{2}{9}\frac{d^2}{L^2}\frac{UL}{v}\frac{(\rho_f-\rho_p)}{\rho_f}$. Since the aggregation of rigid particles requires departures of the particle velocity from the (divergence free) velocity field of the fluid, one can expect that aggregation will occur more slowly if *d* and $(\rho_f - \rho_p)/\rho_f$ are small, or if *v* is large. At the same time, the existence of attractors internal to the fluid may depend on $(\rho_f - \rho_p)/\rho_f$ being small: for example, a large density difference may mean that rigid particles simply sink to the bottom or rise to the surface (and are thus attracted to attractors external to the fluid interior).

As pointed out by Haller and Sapsis (2008) (also see Beron-Vera et al. 2019), we can consider a
continuous concentration of rigid particles with <u>similar</u> properties, and with smoothly varying
velocity <u>given by Eq. (4)</u>. The aggregation of such a concentration would appear to require that
the divergence of that velocity be negative (though see an apparent counterexample <u>in Fig 1c</u>,
presented later). Following Haller and Sapsis (2008), consider the evolution of a material
volume of rigid particles. The time rate of change of this volume is

243
$$\frac{dV}{dt} = \oiint \vec{v} \cdot \vec{n} \, dA_V = \iiint (\nabla \cdot \vec{v}) \, \mathrm{dV} = \iiint \nabla \cdot \left[\vec{u} + \tilde{\varepsilon} \left(\frac{3R}{2} - 1 \right) \left(\frac{D\vec{u}}{Dt} + 2\vec{\Omega} \times \vec{u} - \vec{g}_r \right) \right] \mathrm{d}V \tag{5}$$

where $\nabla \cdot \vec{u} = 0$ for an incompressible fluid. Shrinking *V* to an infinitesimal size allows the righthand side to be approximated by *V* times the local value in the integrand, and the result may be integrated in time, yielding

247
$$V(t) = V_0 \exp\left(\tilde{\varepsilon} \left(\frac{3R}{2} - 1\right) \int_{t_0}^t \nabla \cdot \left(\frac{D\vec{u}}{Dt} + 2\vec{\Omega} \times \vec{u} - \vec{g}_r\right) ds\right)$$

248
$$= V_0 exp\left(-2\tilde{\varepsilon} \left(\frac{3R}{2} - 1\right) \int_{t_0}^t [Q_r(x(s), s) + \vec{\Omega} \cdot \vec{\zeta}_r + \left|\vec{\Omega}\right|^2] ds\right)$$

249
$$= V_0 exp\left(-2\tilde{\varepsilon} \left(\frac{3R}{2} - 1\right) \int_{t_0}^t Q_a(x(s), s) ds\right).$$
(6)

Here $Q_r = \frac{1}{2} \left(\frac{1}{2} |\vec{\zeta_r}|^2 - |S|^2 \right)$ is the three-dimensional Okubo-Weiss parameter (Okubo, 1970; Weiss, 1991), $\vec{\zeta_r}$ represents the relative vorticity vector for the fluid, $S = 1/2(\nabla \vec{u} + (\nabla \vec{u})^T)$ is the strain tensor, and |S| is its Frobenius norm. The final step in Eq. (6) follows from introduction of the absolute vorticity vector

$$254 \quad \vec{\zeta}_a = \vec{\zeta}_r + \overline{2\Omega} \tag{7}$$

and the corresponding function $Q_a = \frac{1}{2} \left(\frac{1}{2} |\vec{\zeta_a}|^2 - |S|^2 \right)$. We note that for a volume *V* of any size:

257
$$\frac{dV}{dt} = 2\tilde{\varepsilon} \left(\frac{3R}{2} - 1\right) \iiint Q_a \, \mathrm{dV} = \tilde{\varepsilon} \left(\frac{3R}{2} - 1\right) \iiint \frac{\partial^2}{\partial x_i \partial x_j} \tau_{ij} \, \mathrm{dV} = \frac{2}{9} \frac{d^2}{L^2} \frac{UL}{\nu} \frac{(\rho_f - \rho_p)}{\rho_f} \oiint \frac{\partial}{\partial x_j} \tau_{ij} n_i \, dA_V,$$
258 (8)

where
$$n_j$$
 denote the components of the outward unit vector normal to the bounding surface A_V .
The first equality in Eq. (8) is a modest modification of Eq. (31) from Haller and Sapsis (2008),
and one could probably have guessed that our more general result could be obtained by replacing
 $Q_{\rm with} Q_a$. The remainder of the equation expresses volume changes in terms of the fluid
stresses. Thus for buoyant particles, a volume $V(t)$ of any size will contract if the force normal
to A_V due to the fluid stresses, integrated around A_V , is inward. In many cases, including
quasigeostrophic eddies and gyres, internal waves, and the surface gravity waves considered by
DiBenedetto et al. (2018a,b) and all inviscid flows, the stress tensor is dominated by pressure,
i.e., $\frac{\partial}{\partial x_j} \tau_{ij} \simeq -\frac{1}{\rho_f} \nabla p$, so the tendency to aggregate is determined entirely by the pressure field.

In general, Q_a can change sign along a particle trajectory, making it hard to predict whether the surrounding volume shrinks or expands with time. If a buoyant particle is trapped in a region in which Q_a is predominatly positive, then this region is a good candidate for aggregation.

271 Persistent ocean eddies and other vortical structures are possibilities, not only because vorticity

tends to dominate over strain, but also because such features have the ability to trap fluid for long

273 periods of time. For dense particles, contraction occurs in areas dominated by strain, and it has

274 been shown that aggregation of heavy particles can occur in strain-dominated filaments that arise

275 in particle-laden turbulent flows, though the considered particle-to-fluid density differences tend

to be quite large (see Brandt and Coletti, 2022 for a review). In our study, we will focus on

277 vortex flows reminiscent of ocean eddies, and on lower dimension objects within such flows that

278 <u>can act as attractors for buoyant particles.</u>

279 A simple example of aggregation is given by Haller and Sapsis (2006), who argue that the elliptical center of a steady, non-divergent 2d eddy, with $\vec{g} = |\vec{\Omega}| = 0$, acts as an attractor for 280 buoyant particles. Here Q_a (now = Q_r), is ostensibly positive near the elliptical center of the 281 eddy, corresponding to contraction of the phase space (which in our case coincides with the 282 physical space) of the rigid particle motion. Since the central fixed point of the velocity field of 283 284 the eddy is also a fixed point of the slow manifold particle velocity (Eq. (4)), buoyant particles initiated about the center should migrate towards the center. If the eddy is inviscid and its 285 streamlines are circular, then the pressure and azimuthal velocity are related by the cyclostrophic 286 balance $\frac{1}{\rho_f} \frac{\partial p}{\partial r} = \frac{u_{\theta}^2}{r}$ so that $2Q_r = \frac{1}{\rho_f} \left(\frac{1}{r} \frac{\partial p}{\partial r} + \frac{\partial^2 p}{\partial r^2} \right)$, and for an eddy in solid body rotation ($u_{\theta} =$ 287 $\Gamma_s r$), $2Q_r = \frac{1}{\rho_f} \left(\frac{1}{r} \frac{\partial p}{\partial r} + \frac{\partial^2 p}{\partial r^2} \right) = 2\Gamma_s^2$. As suggested in Figure 1a, a small concentration of rigid 288 particles indicated by the cross hatched area shrinks as it moves towards the center of the eddy. 289

290 The contraction is partially due to the geometric effect of movement towards smaller radius $\left(\operatorname{term} \frac{1}{r} \frac{\partial p}{\partial r}\right)$ but also due to the fact that the pressure gradient decreases to zero as the center is 291 approached and thus the inner edge of the path moves more slowly inward than the outer part 292 $(\text{term} \frac{\partial^2 p}{\partial r^2})$. In the case of solid body rotation the two terms contribute equally. A second 293 example (Fig. 1b) is of an eddy with an azimuthal velocity given by $u_{\theta} = \Gamma_C r^{1/2}$. Here $\frac{\partial^2 p}{\partial r^2} = 0$ 294 and $2Q_r = \frac{1}{\rho_f} \left(\frac{1}{r} \frac{\partial p}{\partial r} \right) = \Gamma_c^2 / r > 0$, so the contraction of the patch is entirely due to the geometric 295 effect of its movement towards smaller radius. The most curious case is that of a point vortex: 296 $u_{\theta} = \Gamma_P r^{-1}$, for which $2Q_r = \frac{1}{\rho_f} \left(\frac{1}{r} \frac{\partial p}{\partial r} + \frac{\partial^2 p}{\partial r^2} \right) = \frac{\Gamma_P^2}{r^4} - \frac{3\Gamma_P^2}{r^4} < 0$. Here the vorticity is zero away from 297 the eddy center and the velocity field is dominated by strain. The pressure gradient *increases* as 298 the center of the vortex is approached, meaning that the inner part of the patch moves towards 299 300 the center more rapidly than the outer portion (Fig. 1c) and this tendency (quantified by the factor $-\frac{3\Gamma_P^2}{r^4}$) surpasses the tendency towards geometrical contraction (quantified by the factor 301 $\frac{\Gamma_P^2}{r^4}$). The <u>area of the patch</u> thus expands as rigid particles are drawn towards the center of the 302 vortex. Note, however, that a patch surrounding the center of the vortex can only shrink. This 303 304 behavior is made possible by the singularity at the center, and although this feature is artificial, 305 point vortices are often used in idealized models of fluid flow and will act as sinks or "black holes" for buoyant particles even though $2Q_r < 0$. 306

The sign of Q_a is clearly not the whole story and does not encompass the effects of boundaries. For example, consider the fate of heavy ($\rho_f < \rho_p$) particles in the eddy show in Fig. 1a. The particles will migrate outward in each case, and no interior attraction will occur unless the eddy is surrounded by a boundary, which would then act as an attractor.

311	In the next section, we will consider a more general, 3D, eddy-like circulation: one that has both
312	vertical and horizontal components of vorticity, time dependence, and a variety of vortical
313	structures that act as candidates for attraction. Our model is based <u>on</u> the incompressible flow in
314	a rotating cylinder (Greenspan, 1986), which has been studied in many configurations by
315	numerous authors as a model of ocean circulation (Hart and Kittelman, 1996; Pedlosky & Spall,
316	2005), ocean eddies (Pratt et al., 2014; Rypina et all., 2015), or industrial processes and
317	engeneering applications (Lopez & Marques, 2010 and references therein), and can be easily set
318	up in the laboratory setting (Fountain et al., 2000; Lackey and Sotiropoulos (2006)). It its
319	original configuration the cylinder rotates about a vertical axis at a constant (positive) angular
320	velocity $(\vec{\Omega} = \Omega \vec{k})$, and the lid, which is in contact with the fluid, rotates with a slightly greater
321	angular speed. The differential rotation sets up an azimuthal circulation in the horizontal and an
322	overturning circulation in the vertical. (Overturning is observed in ocean eddies as well and
323	Ledwell et al. (2008) present an example.) The steady, axially symmetric state of the rotating
324	cylinder flow that is established will be our first object of investigation. A steady but
325	asymmetrically-perturbed variant can be established by moving the axis of rotation of the lid
326	away from the axis of rotation of the cylinder, and this offset can also be varied in order to
327	induce time dependence. Fountain et al. (2000) set a similar situation up in a laboratory cylinder
328	using a submerged impeller that can be tilted, rather than the differentially rotating lid that can be
329	shifted, to establish an asymmetric disturbance flow. The authors discussed the Lagrangian
330	characteristics of the undisturbed flow and demonstrated the existence of secondary vortical
331	structures generated when the flow is perturbed. Pratt et al. (2014) reproduced similar structures
332	using a primitive equation simulation and explored the rich assembly of chaotic regions and non-
333	chaotic vortical structures as functions of the Ekman and Rossby numbers of the flow. The time-

dependent version of the rotating cylinder flow and a theory describing the resulting vortical

335 structures were discussed by Rypina et al. (2015), who based their examples on a

phenomenological model that reproduced many of the qualitative features of the numerically-

obtained velocity field. In dimensionless Cartesian coordinates, the model velocity field is givenby

339
$$u^{(x)} = -bx(1-2z)\frac{r_o - r}{3} - ay(c+z^2) + \varepsilon \left[y \left(y - y_o + \gamma cos(\sigma t) \right) - \frac{r_o^2 - r^2}{2} \right] (1 - \beta z) , \quad (9a)$$

340
$$u^{(y)} = -by(1-2z)\frac{r_o - r}{3} + ax(c+z^2) - \varepsilon x (y - y_o + \gamma \cos(\sigma t))(1 - \beta z),$$
(9b)

341
$$u^{(z)} = bz(1-z)\frac{2r_o-3r}{3}$$
, (9c)

in which $r = (x^2 + y^2)^{1/2}$ and r_0 is the cylinder radius. The velocity field consists of a steady, 342 343 axially symmetric flow of strength a with an overturning circulation of strength b. To this symmetric state one can add an asymmetric, possibly unsteady and depth dependent, perturbation 344 345 of amplitude ε (not to be confused with the Stokes number $\tilde{\varepsilon}$). The perturbation is quantified by an offset parameter y_o that introduces axial asymmetry in the velocity field, a frequency σ , and 346 an amplitude β for linear depth dependence and an amplitude γ for the time dependence. For the 347 case of axially symmetric, steady flow ($\varepsilon = 0$) the horizontal velocity field, in cylindrical 348 coordinates, becomes 349

350 $u^{(r)} = -br(1-2z)\frac{r_o-r}{3}$ (10a)

351 and

352
$$u^{(\theta)} = ar(c+z^2),$$
 (10b)

where θ is the azimuthal angle. Table 1 lists the parameter values used for each numerical experiment.

We now review the main features of the Lagrangian circulation in the rotating cylinder flow. In 355 356 the steady, symmetric configuration, each fluid trajectory is confined to the surface of a torus as it winds around the cylinder. The typical torus is associated with quasi-periodic trajectories and 357 any such trajectory, followed for a sufficient length of time so that it completes many 358 overturning and azimuthal rotations around the cylinder, will sketch out the torus in 3D. Fig. 2b 359 contains several examples of such tori and Fig. 2a shows the corresponding Poincare map, made 360 by marking the crossing points of trajectories through a vertical slice through the cylinder. After 361 a large number of crossings each quasi-periodic trajectory traces out the cross section of the torus 362 on which it lives. The tori are nested within each another, with a single, horizontal, periodic 363 364 trajectory located at the center of the nest. Certain tori contain periodic trajectories, and these will show up as a finite number of dots on the Poincare map. Because of this geometry, the 365 motion of fluid parcels is most naturally described in terms of action-angle-angle variables, 366 where the action, *I*, acts a label for a particular torus and is constant following each trajectory, 367 and the two angle variables, $\tilde{\theta}$ and ϕ , define the location of a parcel on the torus. Here $\tilde{\theta}$ is an 368 azimuthal angle that differs from the above cylindrical coordinate θ in how its origin is defined, 369 while the 'poloidal' angle ϕ wraps around the cross-section of each torus. The coordinates are 370 non-orthogonal but are defined in such a way that the angular velocities, $\Omega_{\tilde{\theta}}$ and Ω_{ϕ} , are also 371 constant following a trajectory. The explicit transformations to the action-angle-angle variables 372 are given in Mezic and Wiggins (1994). 373

374 When the symmetric RC flow is perturbed by a small, steady, symmetry-breaking perturbation, 375 as controlled by the parameters ε and y_0 in Eq. (9), the tori that are populated by periodic orbits potentially become resonant and break up, resulting in chaotic motion of fluid parcels in the 376 377 vicinity (Fig. 2d-i). Tori with quasiperiodic orbits deform but stay intact. Examples are discussed 378 by Fountain et al. (2000) and Pratt et al. (2013), and the latter found that chaos generally dominates in a large region that includes the central axis of the cylinder and extends around the 379 380 boundaries of the cylinder. Away from this region the space is occupied by tori that have 381 survived the perturbation, and these are sandwiched between tori that have broken up and created 382 braided regions of chaos. The breakup of a torus also gives rise to new tori that appear as islands 383 in the Poincare maps (Fig. 3d and 3g) and these contain non-chaotic trajectories. The number of 384 islands can be predicted by a theory that decomposes the symmetry-breaking perturbation into Fourier modes, written in the $(I, \tilde{\theta}, \phi)$ coordinates, with wave numbers *n* and *m* in the $\tilde{\theta}$ and ϕ 385 direction. If the angular velocities $\Omega_{\tilde{\theta}}$ and Ω_{ϕ} characterizing the trajectories on a particular torus 386 satisfy the resonance condition $n\Omega_{\tilde{\theta}} + m\Omega_{\phi} = 0$ for some *n* and *m*, equivalent to the trajectories 387 388 on that torus being periodic, then that torus will break up and a new set of invariant tori (islands) will form. Running through the center of the islands will be a periodic trajectory that will execute 389 *n* azimuthal cycles to every *m* poloidal (overturning) cycles. In the case shown in Fig. 3a, 390 n = m = 1, so the periodic trajectory circles the cylinder horizontally once for each overturning 391 cycle: a so-called 1:1 resonance. 392

393 If the symmetry breaking perturbation is quasi-periodic in time, with underlying frequencies σ_i ,

the resonance condition for the breakup of a torus becomes $n\Omega_{\tilde{\theta}} + m\Omega_{\phi} + l_i\sigma_i = 0$, where l_i 's

are integers (Rypina, et al. 2015). Unlike the resonance condition for the steady perturbation,

396 which is only satisfied on tori foliated by periodic trajectories, this new resonant condition may

397	be satisfied on tori that have quasi-periodic orbits, and the resonant islands that form will have a
398	shape and location that vary in time. An example (Fig. 2g,h) of the case of a resonance with a
399	single-frequency (i.e., time-periodic) perturbation shows a number of resonant islands. These
400	features vary in time, recovering their shape and location periodically, and the snapshots shown
401	are obtained by strobing the trajectories in 3D and at the forcing frequency. The green and blue
402	islands in Fig. 2h have resulted from the breakup of tori with quasiperiodic trajectories, and
403	center of the island corresponds to a closed material curve that is populated with quasiperiodic
404	trajectories.
405	Note that the resonance condition above and our results in general are applicable to quasi-
406	periodic disturbances with finite number of frequencies, rather than only periodic disturbances.
407	(We only show numerical simulations for the time-periodic case for simplicity.) Because any
408	broad-spectrum function can be arbitrary closely represented by a quasi-periodic function with a
409	finite number of frequencies, this could be applicable to some oceanic flows, especially those
410	with pronounced peaks in the spectrum. However, for flows with truly broadband spectrum, this
411	approach is probably poorly applicable and/or at least impractical because of the very large
412	number of discrete frequencies needed. This is similar in its utility/applicability to other
413	Kolmogorov-Arnold-Moser—based and resonance—based arguments used in prior papers by
414	many authors (including both us and the reviewer), see, for example, Rypina et al., 2007 and
415	<u>Beron et al., 2008; 2010.</u>

416 III. Results

417 Aggregation of rigid particles will occur in presence of an attractor, an object with a dimension
418 < 3 to which particles tend asymptotically in time. We are most interested in attractors that

419 occur in the interior of the rotating cylinder, and are set up by the background circulation, as 420 opposed to the physical boundaries of cylinder. We will see that a closed material contour consisting of periodic orbits near the core of the nested tori in the steady symmetric case act as 421 422 an attractor for slightly buoyant particles, and that similar material contours consisting of periodic or quasiperiodic orbits near the centers of the resonant islands in the asymmetric cases 423 can play the same role. We will explore three cases in increasing complexity, beginning with 424 steady flows with axial symmetry, and proceeding to steady, asymmetric flows and finally 425 unsteady asymmetric flows. 426

The search for attractors is motivated by the hypothesis that for cases of strong drag, where the rigid particle velocity lies close to the fluid velocity, a periodic orbit for the rigid particle motion will exist in the vicinity of a periodic trajectory for the fluid parcel motion, and that if $Q_a > 0$ in a region surrounding the latter, that it should attract buoyant particles. For the time-dependent case, we extend the search to included closed loops that contain recirculating rigid particles and that vary periodically in time.

433 (a) steady, axially-symmetric 3D flows

The fluid velocity field for this case is given by Eqs. (9c) and (10), and these indicate that the location of the horizontal, periodic fluid parcel trajectory living at the center of the nested tori, is given by $r = 2r_o/3$ and $z = \frac{1}{2}$. It is natural to ask whether a periodic trajectory for rigid particles also exists nearby. In the slow-manifold approximation, the steady radial, azimuthal and vertical particle velocities are obtained by writing Eq. (4) in cylindrical coordinates, leading to

439
$$v^{(r)} = u^{(r)} + \tilde{\varepsilon} \left(\frac{3R}{2} - 1\right) \left[\left(u^{(r)} \frac{\partial}{\partial r} + u^{(z)} \frac{\partial}{\partial z} \right) u^{(r)} - u^{(\theta)} \left(2\Omega + \frac{u^{(\theta)}}{r} \right) - \Omega^2 r \right]$$
(11a)

440
$$v^{(\theta)} = u^{(\theta)} + \tilde{\varepsilon} \left(\frac{3R}{2} - 1\right) \left[\left(u^{(r)} \frac{\partial}{\partial r} + u^{(z)} \frac{\partial}{\partial z} \right) u^{(\theta)} + u^{(r)} \left(2\Omega + \frac{u^{(\theta)}}{r} \right) \right]$$
(11b)

441
$$v^{(z)} = u^{(z)} + \tilde{\varepsilon} \left(\frac{3R}{2} - 1\right) \left[\left(u^{(r)} \frac{\partial}{\partial r} + u^{(z)} \frac{\partial}{\partial z} \right) u^{(z)} + g \right]$$
(11c)

442 *Position of attracting periodic orbit; approximate analytical expression on a slow manifold*

443 Searching for points $r = r_c$ and $z = z_c$ for which $v^{(r)} = v^{(z)} = 0$, and that lie in the proximity of 444 the horizontal trajectory of the flow, we introduce

445
$$r_c = \frac{2r_o}{3} + \tilde{\varepsilon}\left(\frac{3R}{2} - 1\right)\tilde{r}$$
 and $z_c = \frac{1}{2} + \tilde{\varepsilon}\left(\frac{3R}{2} - 1\right)\tilde{z}$.

Substituting into the right-hand sides of (11a,c) and setting both to zero results, after neglect of $O(\tilde{\varepsilon}^2)$ terms, in

448
$$r_c = \frac{2r_o}{3} + \tilde{\varepsilon} \left(\frac{3R}{2} - 1\right) \frac{g}{b} r_o$$
 (12a)

449 and

450
$$Z_{c} = \frac{1}{2} + \frac{9}{2br_{o}} \tilde{\varepsilon} \left(\frac{3R}{2} - 1\right) \left[\Omega^{2} + a\left(c + \frac{1}{4}\right) \left(2\Omega + a\left(c + \frac{1}{4}\right)\right)\right].$$
(12b)

For the parameters a > 0 and b > 0, circulation is cyclonic with upwelling in the center of the cylinder, and (3R/2) - 1 > 0 for buoyant particles, so the $O(\tilde{\varepsilon})$ corrections are positive and the periodic particle orbit lies at larger radius and elevation than the periodic fluid orbit. Note also from Eq. (11b) that the azimuthal velocity component of the rigid particle on the periodic orbit is equal to that of the fluid.

456 An explanatory sketch (Fig. 3) shows the position of the periodic orbit of the rigid particle

457 relative to that of the periodic orbit of the fluid. Since the rigid particle is buoyant, it can

maintain its level z only if it is situated in a region where the vertical fluid velocity is < 0, here 458 to the right of the fluid periodic orbit. Also, the horizontal pressure gradients associated with the 459 centripetal acceleration associated with the frame rotation (term $\Omega^2 r$), the Coriolis acceleration 460 (term $2\Omega u^{(\theta)}$), and the centripetal acceleration due to the azimuthal velocity $u^{(\theta)^2}/2r$ are all 461 positive for this flow, so that low pressure exists at r=0 and the rigid particle is forced 462 horizontally inward. To remain stationary the particle must sit in a region where the radial 463 velocity of the fluid is outward. In this manner, the periodic trajectory exists at a location where 464 the forces of inertia, buoyancy and added mass can be countered by the drag due to the 465 background flow. If we fix all other parameters and increase Ω through positive values, the term 466 multiplying $\tilde{\varepsilon}$ in Eq. (12b) will become dominated by the Ω^2 term and will grow without bound 467 and the periodic trajectory may cease to exist. At the same time, a periodic orbit for the rigid 468 particle can always be found close to that of the fluid, regardless of the magnitudes of the 469 470 parameters Ω , a, b etc., provided that the relative particle size d/L (and thus $\tilde{\epsilon}$), and/or the relative density difference $\frac{(\rho_f - \rho_p)}{\rho_f}$ (and thus $\frac{3R}{2} - 1$) are made sufficiently small. 471 Position of attracting periodic orbit; conditions for the loss of periodic orbit 472 We have suggested that periodic orbits for rigid particles are encouraged when the $\tilde{\varepsilon} \left(\frac{3R}{2}\right)$ 473 1) << 1, and in the case of Run 1 the value is 0.0066. A cross-sectional plot of the radial and 474 vertical components of the slow manifold particle velocity in a vertical section through the 475 476 cylinder (Fig. 4a) shows that the periodic orbit lies at r = 0.369 and z = 0.504 (as compared to the values $r_c = 0.338$ and $z_c = 0.502$ predicted by Eq. (12). (The convergence of the 477 surrounding velocity field is too weak to be seen in the graphic.) If $\tilde{\varepsilon}\left(\frac{3R}{2}-1\right)$ is raised to the 478 479 moderately small value 0.02, the position of periodic trajectory migrates towards larger radius

(Fig. 4b), the reason being that the greater buoyancy (larger value of $\frac{3R}{2} - 1$) or smaller drag

- 481 (larger $\tilde{\varepsilon}$) requires a larger downward fluid velocity for equilibrium. Since the maximum
- downward fluid velocity occurs at the outer cylinder wall (see Eq. (9c)) the position of the
- 483 periodic orbit continues to migrate outward and is lost (Fig. 4c) when $\tilde{\varepsilon}\left(\frac{3R}{2}-1\right)$ exceeds a
- 484 value close to 0.3.

485 *Position of periodic orbit in numerical simulations:*

486 The slow-manifold reduction yields to the prediction (Eq. (12)) of the position of the attracting 487 material contour, or loop, for slightly buoyant particles. We can compare this prediction to what is observed in numerical simulations using the Maxey-Riley Eqs. (1) and (2) over a range of 488 particle size d (and thus $\tilde{\varepsilon}$) and frame rotation Ω . As shown in Fig. 5, qualitative agreement with 489 the slow-manifold prediction, and the sketch in Fig. 3, holds for a very small d (when $\tilde{\varepsilon}$ is small). 490 Here the attractor in Fig. 5 is located close to the central periodic fluid parcel trajectory that lives 491 at mid-depth, z = 0.5 and $r = \frac{2R}{3} \approx 0.33$. As d (and $\tilde{\epsilon}$) increases, the attractor moves 492 493 increasingly up and outward, and although the theory captures the trends, quantitative agreement 494 with the numerical results worsens. Also, when frame rotation Ω is increased (panel c), the attractor responds by shifting up from mid-depth, again in qualitative but not quantitative 495 496 agreement with the slow-manifold prediction in Eq. (12b).

497 *Geometry of rigid particle trajectories and evidence of attraction in numerical simulations:*

- 498 If in the neighborhood of the periodic rigid particle trajectory $Q_a > 0$, the phase space for
- 499 buoyant particles will contract and the periodic trajectory becomes a candidate for an attractor of
- such particles. An example of the attraction towards the periodic orbit is shown in Figure 2c,
- 501 where a set of slightly buoyant particles ($\frac{\rho_p}{\rho_f} = 0.97$) has been initialized over the volume of the

cylinder, and Eqs. (1) and (2) have been integrated forward in time to determine their subsequent
trajectories. Each trajectory is shown using a unique color. It can be seen that the particles
aggregate within a ring-like structure of decreasing thickness in the general vicinity of the
periodic orbit of the fluid flow.

506 <u>Basin of attraction – relationship to Q_a :</u>

To map out the basin of attraction for the particle periodic orbit, we first consider the region over 507 which phase space contraction for the buoyant particles (i.e. $Q_a > 0$) occurs. This region is 508 shown in Fig. 6a for the current example, along with the streamlines of the fluid overturning 509 510 stream function. Much of the fluid flow recirculates entirely within the region of positive Q_a , whereas some of the outer streamlines cross the boundary (thick contour) between positive and 511 negative Q_a . If it were the case that rigid particles exactly followed streamlines of the fluid 512 513 overturning circulation, then net contraction or expansion of phase space along a rigid particle trajectory would depend on the sign of the time-integrated value of Q_a along streamlines. The 514 $Q_a = 0$ contour, shown by a bold contour in each frame of Fig. 6, might then approximately 515 delineate the basin of attraction for buoyant rigid particles. In the slow-manifold approximation, 516 where rigid particle velocities lie close to the fluid velocities, the $Q_a = 0$ contour might continue 517 to do so. 518

To test this conjecture, we locate the basin of attraction in the numerical simulations by releasing buoyant particles at various locations in the cross-section $0 < x < r_o$ and 0 < z < 1, integrating the subsequent trajectories over many overturning cycles, and recording the position (x_{final} and z_{final}) of each particle where it crosses the same plane the final time (i.e., recording final crossing with the Poincare section). We use the variable-step 4-th order Runge-Kutta integration 524 scheme, which we implemented in Matlab via the built-in function "ode45". In our simulations, the relative and absolute tolerances are set to the value of 10^{-9} to integrate particle trajectories (Eqs. 525 (2) and (3)) (our results were not sensitive to the further decrease in tolerance values). Since the 526 flow (Eqs. (9a,b,c)) is prescribed analytically and has no normal flow component at the perimeter 527 and top and bottom of the cylinder, no interpolation scheme is needed and no extra boundary 528 conditions are enforced during the integration. Integration of a trajectory is stopped when a 529 particle got within one particle radius from the cylinder walls or top/bottom. The values of z_{final} 530 as a function of initial particle position are mapped in Fig. 7a, where the large green area 531 corresponding to $z_{final} \approx 0.5$ indicates the region from which particles are attracted. Only 532 particles initiated near the central axis of the cylinder, and close to the cylinder boundaries lie 533 outside this region, and these rise to the surface of the cylinder, contact the upper lid, and are no 534 longer followed. It can be seen that the green area in Fig. 7a has an oval shape that somewhat 535 resembles the overturning streamlines at small x in the central part of the cylinder, but extends to 536 near the top, bottom and outer cylinder boundaries at larger x. Thus the $Q_a = 0$ contour provides 537 538 a rough indication of the size and shape of the basin of attraction, but misses some important 539 details.

540 <u>Basin of attraction – dependence on Ω </u>

We have seen that the location of the periodic orbit that acts as an attractor for buoyant particles shifts up and out in response to increasing frame rotation Ω (Fig. 5c). In Fig. 8 we indicate the corresponding changes in the extent of the basin of attraction with respect to changing Ω by recomputing Fig. 8a with $\Omega = 0.3$, 1, and 10. The two smaller Ω values (0.3 and 1) correspond roughly to Rossby numbers $a/2\Omega$ of about 1 and 0.2, i.e., are representative of the ocean submesoscale and mesoscale flows. The Q_a -functions for these cases are plotted in Fig. 6b-c. 547 Most submesoscale eddies are going to tend to have $u^{(\theta)}/r$ about the same magnitude as Ω 548 (except on the equator) and mesoscale eddies will have $u^{(\theta)}/r \ll \Omega$. The results in Fig. 8 549 suggest that, while the basin of attraction does shrink slightly with increasing Ω , this dependence 550 is weak. The main difference between the three numerical runs in Fig. 8 is in the associated 551 attraction time, which gets significantly shorter for larger values of Ω . This is explored in more 552 detail below.

553 <u>Attraction time:</u>

It follows from Eq. (6) that the attraction time towards the periodic orbit should scale as $T_a =$ 554 $\left[2\tilde{\varepsilon} \left(\frac{3R}{2}-1\right)Q_a\right]^{-1}$ where $Q_a = \frac{1}{2}\left(\frac{1}{2}|\vec{\zeta}_a|^2-|S|^2\right)$ with $\vec{\zeta}_a = \vec{\zeta}_r + \overline{2\Omega}$. Thus, for $\vec{\zeta}_r \ge 0$, as in 555 most of our numerical runs (except Experiment 1e), attraction time decreases with increasing Ω 556 for positive $\Omega \ge 0$. For negative $\overline{\zeta_r}$, which corresponds to the reversed direction of the flow in 557 our simulations (Experiment 1e), an increase in Ω will initially slow the attraction by decreasing 558 the magnitude of $\vec{\zeta}_a$ all the way to 0, at which point the periodic orbit will lose its attraction 559 properties, but then will speed up the attraction as Ω is further increased. This trend is confirmed 560 numerically in Fig. 9, where for the flow parameters corresponding to the "reversed flow" run in 561 Table 1 (Experiment 1e, with $\bar{\zeta}_r < 0$), we release a sample trajectory within the basin of 562 attraction and plot its z-coordinate as it winds around the can and eventually approaches the 563 attracting periodic orbit. As anticipated, the attraction time initially increases as Ω is increased 564 from 0 to 0.6, but then decreases as Ω is further increased to 2. 565

566 *Disappearance of the subsurface attractor when* $\tilde{\varepsilon}$ *becomes too large:*

567 Finally, to illustrate the disappearance of the subsurface attractor when $\tilde{\varepsilon}$ becomes too large, in

Fig. 10, we contrast 2 numerical simulations with the same flow parameters (corresponding to

the "slow overturn" run 1c in Table 1) but different particle diameters, $d = 10^{-3}$ vs $d = 5 \times 10^{-4}$. For larger *d*, the subsurface periodic orbit for rigid particles is no longer present within the can, leading to all particles rising up to the surface (Fig. 10b). For smaller *d*, the periodic orbit is still present and acts as an attractor for buoyant rigid particles over a significant portion of the can (green region in Fig. 10a). We note that this run would be more qualitatively similar to the oceanic mesoscale or submesoscale eddies, where the overturning component of circulation is weak in comparison to the horizontal swirl.

576 (b) steady non-symmetrically perturbed case

We now consider a case in which the axial symmetry of the steady flow has been broken, here through a change in the perturbation amplitude parameter ε from zero to 0.25, and in the offset parameter y_o from 0 to -0.2 in the Eqs. (9a,b). The fluid velocity field now contains something like a stationary, "mode-1" azimuthal wave in the horizontal velocity field.

The resulting Lagrangian structure (Fig. 2d and e) has a sea of chaos that covers the near-axial 581 582 and outer regions of the cylinder, where no unbroken tori survive. Within this chaotic sea is a 583 region containing a nest of unbroken tori that surround a central periodic orbit. This orbit has evolved from the central periodic orbit of the symmetry case and is now tilted. Within the nest of 584 unbroken tori there exist resonant layers, in which new tori have arisen, and the most prominent 585 586 is the "island" that is centered near x = 0.4 and z = 0.2 in the right-half (and near x = 0.4 and z = 0.2 in the right half) of Fig. (2d). We further note that this center lies within the region of 587 positive Q_a (Fig. 6b). The island corresponds to the yellow tori in Fig. 3e and is produced by a 588 1: 1 resonance, so that the periodic trajectory running through its center executes one complete 589 azimuthal cycle and one overturning cycle before connecting back onto itself. Thus, in this 590 steady asymmetric configuration, we now have 2 periodic orbits of the fluid flow – the central 591

slightly-tilted periodic orbit near mid-depth (that evolved from the central horizontal periodic orbit of the axisymmetric flow) and a new periodic orbit running through the center of the resonant island (resulting from the break-up of the resonant torus satisfying $\Omega_{\tilde{\theta}} + \Omega_{\phi} = 0$).

We speculate that for sufficiently small $\tilde{\varepsilon}$ a periodic orbit for the rigid particle motion exists in the vicinity of each of the 2 periodic orbits of the fluid flow. This conjecture is difficult to prove due to a complex geometry, leading to centrifugal forces that act in different directions at different locations along the particle path. For now we simply search for the supposed attractors by releasing particles and following their trajectories.

As shown in Fig. 2f, separate attractors arise in the vicinity of two periodic orbits. The first appears as a ring-like structure (purple core) lying near the center of the original nested tori and the second is a similar feature with a red core near the center of the resonant island. The two are chained together and each has its own basin of attraction (Fig. 7c): the first consisting of a roughly elliptical patch (inner green region) in the x-z-plane, which corresponds of a slice through a tube-like structure in 3D, and the second consisting on an annular (blue) region that surrounds the green region and that occupies a relatively larger volume.

In order to check that attraction of slightly-buoyant rigid particles towards periodic orbits located near the centers of the resonant islands in the perturbed flow is not limited to the case of the 1: 1 resonance, in an additional simulation (Fig. 11, experiment 2c in Table 1), we adjusted the background flow parameter *b* in Eqs. (9), which is responsible for the overturning strength, to create a 2: 1 resonance instead of a 1: 1 resonance, as in the original run. In this case, the resonant torus breaks down giving rise to a 2-island chain on the corresponding Poincare section (Fig. 11a), and the fluid periodic orbit that goes through the centers of both islands completes 2 full cycles in azimuth and 1 complete cycle in vertical before connecting onto itself. Also, as inthe original run, a second slightly-tilted periodic orbit still exists near mid-depth of the can.

616 When buoyant particles are released into this flow, two attractors arise, corresponding to the 2

617 periodic orbits of rigid particles – one near mid-depth (purple core in Fig. 11c) and another in red

618 near the center of the 2:1 resonant island.

619 <u>Shift in position of the periodic orbit associated with a resonant island as a function of flow and</u> 620 particle parameters, and frame rotation

621 The position of the attracting periodic orbit for rigid particles that is located within the resonant 622 islands (we will refer to it as the resonant periodic orbit) in the asymmetrically-perturbed flow depends both on the perturbation strength (via ε), on the flow and particle parameters (via $\tilde{\varepsilon}$), and 623 on the frame rotation Ω . Specifically, this resonant periodic orbit for the rigid particles will shift 624 away from the corresponding periodic trajectory of the fluid flow as $\tilde{\varepsilon}$ and Ω are increased. The 625 same is true for the slightly-tilted central attracting periodic orbit near mid-depth. This is 626 qualitatively similar to the shifting of the central periodic orbit up and out from z = 0.5, 627 628 r = 0.34 in the axisymmetric flow in response to changing $\tilde{\varepsilon}$ and Ω , which we explored in detail 629 the previous section both analytically (Eqs. (12)) and numerically (Fig. 3-5). In order to numerically illustrate the shift in the position of the attracting periodic orbits, we 630 present (Figs. 12 and 13) numerical simulations in the steady perturbed flow configuration for 3 631 632 values of d (and thus $\tilde{\varepsilon}$) and 3 values of Ω . As both parameters increase, the attractors move 633 away from the corresponding periodic orbits of the fluid flow. This shift is evident from the

634 change in the color of the attraction basins in (a,d,g) and from the location of the yellow cloud of

dots in (c,f,i) in Figs. 12-13. Increases in $\tilde{\varepsilon}$ and Ω also lead to the shrinkage of the attraction

636 basins for both attractors and to a faster convergence rate, as is evident from the tighter cloud of yellow dots in (c,f,i), as discussed in more detail below. The basin of attraction for the central 637 attractor – the green region in Fig. 12 – seems to shrink faster than the basin of attraction for the 638 resonant attractor (the blue-ish region) as d increases, so when d is increased from 2×10^{-3} to 639 3×10^{-3} , the central attractor vanishes, whereas the resonant attractor is still present (Fig. 12g). 640 On the other hand, the increase in Ω (Fig. 13) causes a faster shrinkage of the basin of attraction 641 642 for the resonant attractor than for the central attractor, so when Ω is increased from 2 to 5 in Fig. 13g, the resonant attractor disappears, whereas the central attractor is still present. Figs. 12g,h,i 643 (and Fig. 13g,h,i) show cases where this threshold has been exceeded, and one of the attractors 644 645 has been lost, whereas the other is still present.

646 <u>Attraction time:</u>

647 Similar to the unperturbed flow, the attraction time for attractors in the steady, perturbed flow may still scale as $T_a = \left[2\tilde{\varepsilon} \left(\frac{3R}{2} - 1\right)Q_a\right]^{-1}$, provided that Q_a is regarded as a typical value 648 within the corresponding basin of attraction. The predicted decrease in attraction time with 649 increasing $\tilde{\varepsilon}$ and Q_a is evident from the numerical simulations in Figs. 12-13, where in (c,f,i) we 650 color-coded trajectory crossings with the x-z Poincare plain by time, with blue/yellow 651 corresponding to initial/final time. For smaller values of $\tilde{\varepsilon}$ and Ω , we observe a wider and more 652 diffuse cloud of dots (because trajectories wind around the can many times before approaching 653 654 the attractor), whereas as $\tilde{\varepsilon}$ and Ω increase, the clouds at comparable times become denser and 655 more compact around the attractors.

656 <u>Basin of attraction</u>

For the slightly-tilted central periodic orbit located within the central non-chaotic region near mid-depth in Fig. 2f, we observe that the basin of attraction – green region in Fig. 7b – extends roughly from the location of the periodic orbit to the edge of the central non-chaotic region (that is foliated by discretely sampled closed curves in Fig. 2d). Note that as $\tilde{\varepsilon}$ increases, the attracting periodic orbit moves away from the center of this non-chaotic region towards its edge, leading to the shrinkage and eventual disappearance of the corresponding basin of attraction, shown by the green regions in Fig. 12a,d,g).

Similarly, in all of our numerical simulations, we observe that for the resonant attracting periodic
orbit running through the resonant islands, the basin of attraction seems to cover the region
between the orbit and the edge of the corresponding resonant island. An analytical expression for
the width of the (non-degenerate) resonant island in the fluid flow (Pratt et al., 2014) predicts

668 that
$$\Delta I = \sqrt{\frac{\epsilon F_{nm}^0(I_0)}{\left(n\frac{d^j \Omega_{\phi}}{dI^j} + m\frac{d^j \Omega_{\theta}}{dI^j}\right)_{I_0}}}$$
, where ΔI is the deviation in the action coordinate away from I_0 , the

value of action at the resonant torus (i.e., at the center of the island). This width depends on the strength of the perturbation ϵ , the order of the resonance (via *n* and *m* in the resonance

671 condition), the background flow (via $\frac{d^{j}\Omega_{\phi/\theta_{l}}}{dI^{j}}$), and the structure of the perturbation (via $F_{nm}^{0}(I_{0})$). 672 This expression could be used as an upper limit on the extent of the basin of attraction. However, 673 because the attracting periodic orbit will move away from the center of the island towards its 674 edge as $\tilde{\varepsilon}$ and Ω increase, the basin of attraction for the resonant attractor (blue region in Figs. 675 12a,d and 13a,d) becomes increasingly smaller than ΔI . One might speculate, then, that the 676 attractor will completely disappear when the attracting periodic orbit reaches the edge of the 677 resonant island. This is the case in Figs. 13g where the resonant attractor is no longer present. 678 (c) non-steady, non-symmetrically perturbed case

679 The final case that we will consider is one in which the perturbation is asymmetric and varies periodically in time. The chosen perturbation frequency, $\sigma = 2\pi/9.1$, causes 2 strong additional 680 681 resonances (compared to the steady perturbed case) – one with n = 0, m = 1, and l = 1 (i.e., with a torus whose overturning frequency is equal to the perturbation frequency) that is shown in 682 blue in Fig. 2g,h and is located near the outer edge of the central non-chaotic region, and another 683 resonance, shown in green in Fig. 2g,h, with n = 1, m = 1, and l = 1, which is located between 684 the central non-chaotic region and the larger n = 1, m = 1 resonant island (that was present in 685 the steady case as well). Both of these new resonant structures are time dependent, their shape 686 and position recurring periodically. For example, the blue island, which looks like a crescent 687 688 moon pointing upward on the Poincare section at t = 0, becomes a crescent moon pointing downward at time 4.55. The movement of the green island is more complex, as it turns both in 689 azimuth and vertical, making one complete loop over 9.1 time units. Because of the time-690 dependence, trajectories must be strobed at the forcing frequency σ in order to capture 691 'snapshots' of their forms as they recur at a particular phase in the time cycle. At the center of 692 each feature is a closed material curve that also varies periodically. Where the island has 693 694 emerged from the breakup of a torus with quasiperiodic orbits, the individual trajectories that populate the material curves are themselves quasiperiodic. 695

Particle trajectory computations in this case confirm that the purple, red and green islands give rise to attractors (Fig. 3i), whereas the blue island does not. In fact, slightly-buoyant rigid particles that are released in the blue region converge towards the attractor that lies near the purple region. This is also indicated by the basin of attraction of the central attractor extending across the space occupied by the blue resonant island in Fig. 7c.

We have considered attraction phenomena for small, finite size, spherical, buoyant, rigid 702 703 particles in a three-dimensional rotating cylinder flow with azimuthal rotation and overturning, 704 and both with or without time dependence. The aim has been to gain insights into the behavior of 705 slightly buoyant microplastic particles in 3D vortex flows that qualitatively resemble ocean 706 eddies. The rigid particle motion is governed by a simplified version of the Maxey-Riley 707 equations (accounting for inertia, buoyancy and simplified quantification of drag and added 708 mass), and, approximately, by the slow-manifold reduction of these equations. We have illustrated the possibility of aggregation of slightly-buoyant rigid particles in 3D vortex flows 709 towards closed loop attractors located subsurface within the interior of the flow. Even in our 710 711 idealized flow and for spherical particles with fixed radius and buoyancy, aggregation is nontrivial, often with multiple attractors present and/or the lack of attraction in some circumstances. 712 713 Our rotating cylinder model is much less complex than any real ocean eddy in many respects, 714 including the assumed quasiperiodic time dependence and the absence of decay and interaction 715 with the surroundings. Understanding aggregation in a simple periodic flow seems like a 716 reasonable first step towards understanding aperiodic, interacting, and decaying oceanic eddies. This approach is common in applications of dynamical systems theory to oceanography and 717 meteorology. For example, arguments relating to the increased stability of jets due to the strong 718 Kolmogorov-Arnold-Moser stability near shearless trajectories have first been developed for 719 720 spatially-periodic and time-quasiperiodic flows and tested using idealized toy models, before exploring these ideas in more realistic oceanic and atmospheric settings (see Rypina et al., 2007 721 722 and Beron et al., 2008; 2010). Note also that our results are applicable to quasi-periodic 723 disturbances with finite number of frequencies rather than just periodic disturbances (we only

724 show numerical simulations for the time-periodic case for simplicity), and a quasiperiodic

725 <u>function might potentially be useful for approximating temporal variability in some oceanic</u>

726 flows, especially those with pronounced peaks in the spectrum.

We have explored a steady axisymmetric rotating cylinder flow and a steady flow with its axial symmetry broken. In all cases, we have observed emergence of subsurface attracting structures that lead to the aggregation of buoyant particles towards them. We have linked these attractors to the periodic orbits of rigid particles that exist in a region of net contraction of the phase space of the particle motion. The slow manifold equations suggest that periodic orbits for rigid particles exist near periodic orbits of the underlying fluid flow, provided the drag is sufficiently strong (Stokes number << 1).

We have also explored one case of an axially asymmetric and time-periodic flow, with focus on the resonant "islands" that arise due to the time-dependence. At the center of such islands are closed material contours, or loops, composed of quasi-periodic orbits of the fluid flow. One such structure has a nearby attractor, also a closed loop of quasiperiodic orbits for rigid particles, while a second example does not. A detailed explanation awaits formulation of a quantitative theory, something that is beyond the scope of the present paper and that will be presented in a future work.

We have observed that the disappearance of an attractor, which can occur as the result of
increasing rigid particle size or frame rotation, coincides roughly with the displacement of the
position of the attractor to the outer edge of the resonant island from which it sprang. Whether
this purely geometric observation forms the basis for a general criterion for the loss of attraction
is unknown, as a dynamical justification is needed.

746	Marine microplastics can	have complex non-spherical	tangled-filament shapes	, change their
		· ·	-	

- 747 physical and chemical properties in time due to aging and photo- or chemical-decay processes
- 748 (Andrady 2011), are subject to biofouling (see recent relevant work by Kreczak et al., 2021), and
- 749 <u>may interact leading to the formation of clusters. None of these effects were considered in this</u>
- 750 paper, and all will need to be taken into account for the realistic prediction of marine
- 751 <u>microplastic evolution and re-distribution in the ocean. Real ocean eddies are also decaying in</u>
- 752 time and are usually moving (translating) rather than stationary. Translation with a constant
- velocity can be handled by considering the flow in a moving frame of reference, but decay and
- 754 interactions will likely change the geometry of the circulation and make the flow truly aperiodic.
- 755 Our simplified model cannot account for these effects, which will need to be explored separately
 756 later.
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- 759 **Data Availability Statement:** No observational data was used. Details of the numerical
- simulations using an analytical vortex model are provided in text.
- 761 Author Contribution Statement: IR led the overall effort and performed most of the numerical
- simulations, LP contributed towards the theoretical understanding and interpretation of the
- results, MD participated in the overall effort.
- 764 Competing interests Statement: no competing interests
- 765

767 **References**

- Andrady, A.L., 2011. Microplastics in the marine environment. *Marine pollution bulletin*, 62(8),
 pp.1596-1605.
- Basset, A. B. (1988). Treatise on Hydrodynamics, Deighton Bell, London, Vol. 2, Chap. 22, pp.
 285-297.
- Beron-Vera, F. J., M. G. Brown, M. J. Olascoaga, I. I. Rypina, H. Kocak, and I. A.
 Udovydchenkov (2008). Zonal jets as transport barriers in planetary atmospheres. *Journal of Atmospheric Science*, 65, 3316-3326.
- Beron-Vera, F. J., M. J. Olascoaga, M. G. Brown, H. Kocak, and I. I. Rypina (2010). Invariant tori-like Lagrangian coherent structures in geophysical flows. *Chaos*, 20, 017514,
 doi:10.1063/1.3271342.
- Beron-Vera, F. J., M. J. Olascoaga and R. Lumpkin, 2016. Inertia-induced accumulation of
 flotsam in the subtropical gyres. Geophys. Res. Lett., 43, 12228-12233,
 <u>https://doi.org/10.1002/2016g1071443</u>.
- Beron-Vera, F. J., M. J. Olascoaga and P. Miron, 2019. Building a Maxey-Riley framework for
 surface ocean inertial particle dyamics. Phys. Fluids 31; doi: 10.1063/1.5110731.
- 783 Beron-Vera, F.J., 2021. Nonlinear dynamics of inertial particles in the ocean: From drifters and
- floats to marine debris and Sargassum. *Nonlinear dynamics*, *103*(1), pp.1-26.
- Boussinesq, J. (1903) Theorie Analytique de la Chaleur, L'Ecole Polytechnique, Paris, Vol.2, p.
 224.

- Brandt, L. and F. Coletti (2022) Particle-Laden Turbulence: Progress and Perspectives. Ann.
 Rev. Fluid Mich. 54, 159-189. <u>https://doi.org/10.1146/annurev-fluid-030121-021103</u>
- 789 Choy, C.A., Robison, B.H., Gagne, T.O., Erwin, B., Firl, E., Halden, R.U., Hamilton, J.A.,
- 790 Katija, K., Lisin, S.E., Rolsky, C. and S. Van Houtan, K., 2019. The vertical distribution and
- 791 <u>biological transport of marine microplastics across the epipelagic and mesopelagic water</u>
- 792 <u>column. Scientific reports, 9(1), p.7843.</u>
- Daitche, A. and Tél, T., 2011. Memory effects are relevant for chaotic advection of inertial
 particles. *Physical review letters*, 107(24), p.244501.
- 795 Delandmeter, P. and E. van Sebille (2019). The Parcels v2.0 Lagrangian framework: new field
- interpolation schemes. Geosci. Model Dev. Discuss., <u>https://doi.org/10.5194/gmd-2018-339</u>
- DiBenedetto, M. H., N. T. Ouellette, and J. R. Koseff, 2018a. Transport of anisotropic particles
 under waves. J. Fluid. Mech. 837, 320-340, doi:10.1017/jfm.2017.853.
- DiBenedetto, M.H. and Ouellette, N.T., 2018b. Preferential orientation of spheroidal particles in
 wavy flow. *Journal of Fluid Mechanics*, 856, pp.850-869.
- Faxén, H., 1922. Der Widerstand gegen die Bewegung einer starren Kugel in einer zähen
 Flüssigkeit, die zwischen zwei parallelen ebenen Wänden eingeschlossen ist. *Annalen der Physik*, 373(10), pp.89-119.
- 804 Fountain, G.O., Khakhar, D.V., Mezić, I. and Ottino, J.M., 2000. Chaotic mixing in a bounded
 805 three-dimensional flow. *Journal of Fluid Mechanics*, 417, pp.265-301.
- Froyland, G., Stuart, R. M., & van Sebille, E. (2014). How well-connected is the surface of the
- global ocean?. *Chaos*, 24(3), 033126. <u>https://doi.org/10.1063/1.4892530</u>

- Fenichel, N.: Geometric singular perturbation theory for ordinary differential equations. J. Differ.
 Equ. 31, 51–98 (1979)
- 810 Gatignol, R., 1983. The Faxén formulae for a rigid particle in an unsteady non-uniform Stokes811 flow.
- 812 <u>Greenspan, H. P., 1968. The theory of rotating fluids (Vol. 327). Cambridge: Cambridge</u>
 813 <u>University Press.</u>
- Haller, G., and T. Sapsis, 2008. Where do inertial particles go in fluid flows? Physica D, 237,
 573-583.
- Hart, J.E. and Kittelman, S., 1996. Instabilities of the sidewall boundary layer in a differentially
 driven rotating cylinder. *Physics of Fluids*, 8(3), pp.692-696.
- Kelly, R., D. B. Goldstein, S. Suryanarayanan, M. B. Tornielli and R. A. Handler, The nature of
- 819 bubble entrapment in a Lamb-Oseen vortex. Phys. Fluids 33, 061702;
 820 https://doi.org/10.1063/5.0053658.
- 821 Kooi, M., Reisser, J., Slat, B., Ferrari, F.F., Schmid, M.S., Cunsolo, S., Brambini, R., Noble, K.,
- 822 Sirks, L.A., Linders, T.E. and Schoeneich-Argent, R.I., 2016. The effect of particle properties
- 823 <u>on the depth profile of buoyant plastics in the ocean. Scientific reports, 6(1), p.33882.</u>
- <u>Kreczak, H., Willmott, A.J. and Baggaley, A.W., 2021. Subsurface dynamics of buoyant</u>
 <u>microplastics subject to algal biofouling</u>. *Limnology and Oceanography*, 66(9), pp.3287 <u>3299.</u>
- Kukulka, T., Proskurowski, G., Morét-Ferguson, S., Meyer, D. W., & Law, K. L.(2012). The
 effect of wind mixing on the vertical distribution of buoyant plastic debris. *Geophysical*

- 829 *Research Letters*, 39, L07601. <u>https://doi.org/10.1029/2012GL051116</u>
- 830 <u>Kvale, K., Prowe, A.F., Chien, C.T., Landolfi, A. and Oschlies, A., 2020. The global biological</u>
 831 microplastic particle sink. *Scientific reports*, *10*(1), p.16670.
- **Lackey, T.C. and Sotiropoulos, F., 2006. Relationship between stirring rate and Reynolds** number in the chaotically advected steady flow in a container with exactly counter-rotating
 lids. *Physics of Fluids*, 18(5).
- Landrigan PJ, Raps H, Cropper M, Bald C, Brunner M, Canonizado EM, Charles D, Chiles TC,
- B36 Donohue MJ, Enck J, Fenichel P, Fleming LE, Ferrier-Pages C, Fordham, R, Gozt A, Griffin
- 837 C, Hahn ME, Haryanto B, Hixson R, Ianelli H, James BD, Kumar P, Laborde A, Law KL,
- 838 Martin K, Mu J, Mulders Y, Mustapha A, Niu J, Pahl S, Park Y, Pedrotti M-L, Pitt JA,
- 839 Ruchirawat M, Seewoo BJ, Spring M, Stegeman JJ, Suk W, Symeonides C, Takada H,
- 840 Thompson RC, Vicini A, Wang Z, Whitman E, Wirth D, Wolff M, Yousuf AK, Dunlop S.
- 841 The Minderoo-Monaco Commission on Plastics and Human Health. Annals of Global
- 842 *Health.* 2023; 89(1): 23, 1–215. DOI: https://doi. org/10.5334/aogh.4056
- Lange, M. and E. van Sebille (2017) Parcels v0.9: prototyping a Lagrangian ocean analysis
 framework for the petascale age. Geosci. Model Dev., 10, 4175-4186.
 https://doi.org/10.5194/gmd-10-4175-2017
- 846 Langlois, G.P., Farazmand, M. and Haller, G., 2015. Asymptotic dynamics of inertial particles
- 847 with memory. *Journal of nonlinear science*, 25, pp.1225-1255.
- Ledwell, J. R., McGillicuddy, D. J., and Anderson, L. A., "Nutrient flux into an intense deep
 chlorophyll layer in a mode-water eddy," Deep Sea Res., Part II 55, 1139–1160 (2008).

- Maxey, M.R. and J. J. Riley, 1983. Equation of motion for a small rigid sphere in a nonuniform
 flow. Phys. Fluids 26, 883.
- Mountford, A. S. and M. A. Morales Maqueda (2019) Eulerian Modeling of the
 Three=Dimensional Distribution of Seven Popular Microplastic Types in the Global Ocean. *J. Geophys. Res.*: Ocean, 124, 8558-8573. https://doi:10.1029/2019JC015050.
 Onink, V., Wichmann, D., Delandmeter, P., and van Sebille, E., 2019: The role of Ekman
- 856 currents, geostrophy and Stokes drift in the accumulation of floating microplastic, Journal of
- 857 Geophysical Research: Oceans, 124, 1474-1490, https://doi.org/10.1029/2018JC014547,
- 858 Okubo, A., 1970, June. Horizontal dispersion of floatable particles in the vicinity of velocity
- singularities such as convergences. In *Deep sea research and oceanographic abstracts* (Vol.
- 860 17, No. 3, pp. 445-454). Elsevier.
- 861 Pabortsava, K. and Lampitt, R.S., 2020. High concentrations of plastic hidden beneath the
- 862 <u>surface of the Atlantic Ocean. *Nature communications*, 11(1), p.4073.</u>
- Pedlosky, J. and Spall, M.A., 2005. Boundary intensification of vertical velocity in a β-plane
 basin. *Journal of physical oceanography*, *35*(12), pp.2487-2500.
- Pratt, L. J., I. I. Rypina, T. Özgökmen, H. Childs, and T. Bebieva, 2014. Chaotic Advection in a
 Steady, 3D, Ekman-Driven Circulation. *J. Fluid Mech*, 738, 143-183,
 DOI:10.1017/jfm.2013.583.
- 868 <u>Ripa, P., 1987. On the stability of elliptical vortex solutions of the shallow-water</u>
 869 equations. *Journal of Fluid Mechanics*, 183, pp.343-363.
- 870 Rypina, I. I., M. G. Brown, F. J. Beron-Vera, H. Kocak, M. J. Olascoaga, and I. A.

871	<u>Udovydchenkov</u>	(2007). Robust	transport bar	riers resulti	ng from s	strong Ko	olmogorov-
872	Arnold-Moser	stability.	Physical	Review	Letters,	98,	104102,
873	doi:10.1103/Phys	RevLett.98.1041	02.				

- 874 Rypina, I. I., L. J. Pratt, P. Wang, T. M. Ozgokmen, and I. Mezic, 2015. Resonance phenomena
- in a time-dependent, three-dimensional, Ekman-driven eddy. J. Chaos., 25, 087401,
- 876 <u>http://dx.doi.org/10.1063/1.4916086</u>.
- 877 Shamskhany, A., Z. Li, P. Patel and S. Karimpour, <u>2021</u>. Evidence of Microplastic Size Impact
- on Mobility and Transport in the Marine Environmet: A Review and Synthesis of Recent
- 879 Research. Front. Mar. Sci. 8:760649. Doi: 10.3389/fmars.2021.760649.
- Stokes, G. G. (1851) Trans. Camb. Phil. Soc. 9, 8. (Mathematical and Physical Papers 3,1.)
- 881 Tchen, C. M. (1947). Ph. D. thesis, Delft, Martinus Nijhoff, The Hague/
- van Sebille E., C. Wilcox, L. Lebreton, N. Maximenko, B.D. Hardesty, J.A. van Franeker, M.
- Eriksen, D. Siegel, F. Galgani, K.L. Law (2015) A global inventory of small floating plastic
- debris Environ. Res. Lett., 10 (2015), p. 124006, 10.1088/1748-9326/10/12/124006
- Wichmann, D., P. Delandmeter and E. van Sebille, 2019. Influence of Near-Surface Currents on
- the Global Dispersal of Marine Microplastics. J. Geophys. Res. (Oceans), 124(8),
- 887 https://doi.org/10.1029/2019JC015328
- 888 Weiss, J., 1991. The dynamics of enstrophy transfer in two-dimensional hydrodynamics. *Physica*
- 889 *D: Nonlinear Phenomena*, 48(2-3), pp.273-294.
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Experiment	а	b	Е	Уо	σ	γ	β	Ω	d
1 – steady symmetric	0.62	7.5	0	0	0	0	0	0	10 ⁻³
1a (small Ω)	0.62	7.5	0	0	0	0	0	0.3	10 ⁻³
1b (large Ω)	0.62	7.5	0	0	0	0	0	1	10 ⁻³
1c (slow overturn)	0.62	0.25	0	0	0	0	0	1	$10^{-3} vs.$
									5×10^{-4}
1d ($z_{attractor}$ vs Ω)	0.62	7.5	0	0	0	0	0	Sweep 0	10 ⁻³
								to 10	
le (reversed flow)	-0.62	-7.5	0	0	0	0	0	0, 0.6, 2	10 ⁻³
2 – steady asymmetric	0.62	7.5	0.25	-0.2	0	0	0	0	10 ⁻³
2a (small Ω)	0.62	7.5	0.25	-0.2	0	0	0	0.3	10 ⁻³
2b (large Ω)	0.62	7.5	0.25	-0.2	0	0	0	1	10 ⁻³
2c (2:1 resonance)	0.62	3.8	0.25	-0.2	0	0	0	0	10 ⁻³
3 - non-steady asymmetric	0.62	7.5	0.25	-0.2	$\frac{2\pi}{9.1}$	0.2	1	0	10 ⁻³

Table 1: Dimensionless parameter values for numerical experiments. Fixed parameters in the kinematic model (Eqs. 9a-c) are c = 0.69, and $r_0 = 1/2$ in all cases. Parameters that appear in the nondimensional Maxey-Riley Eq. (3) are also nondimensional, with *L*, *U*, *L/U* as length, velocity and time scales. Fixed parameter values based on L = 1m and U = 1m/s include

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$$\frac{\rho_p}{\rho_f} = 0.97, R = \frac{2\rho_f}{\rho_f + 2\rho_p} = 0.680, \frac{3R}{2} - 1 = .020 \ \vec{g}_r = \frac{gL}{U^2} = 10.0 \ , \ \tilde{\varepsilon} = \frac{2}{9} \left(\frac{d}{L}\right)^2 \frac{UL}{v} \frac{1}{R} = 0.33, \text{ and}$$

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$$\tilde{\varepsilon}\left(\frac{3R}{2}-1\right) = 0.0067$$
. Note that $\overline{\Omega} = \Omega \overline{k} = \frac{\overline{\Omega}^* L}{U}$.



Figure 1. Three types of two-dimensional eddies with zero frame rotation and for which gravity is imagined to be zero: solid body rotation (a), constant pressure gradient (b), and point vortex (c). In each case, the cross hatched area represents a concentration of rigid particles with area A(t).



Figure 2. (left) Poincare section, (middle) fluid parcels trajectories in 3D, (right) buoyant particle 911 trajectories in 3D for a steady symmetric fluid flow (top row), steady asymmetric flow (middle 912 row), and non-steady, asymmetric flow. Parameter setting are listed under Experiments 1, 2 and 913 3 in Table 1. Colors in the left column of panels match the corresponding panel in the middle 914 column, but the colors in the right column indicated time after release of the particles. Note the 915 attraction of buoyant particles to a single attractor at mid-depth in panel (c), to 2 attractors in 916 panel (f), and to 3 attractors in panel (i). Particles are released along a vertical line x = 0.334, 917 $y = 0, 0 < z \le 0.6$ with initial velocity equal to that of the co-located fluid parcels. 918

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Figure 3. Sketch showing the position in a vertical section of the periodic orbit (red dot) of the
rigid particle relative to the periodic orbit (blue dot) of the fluid flow. The viewer sees one half
of a vertical slide though the cylinder, with the azimuthal flow directed away from the viewer
and the cylinder center at the left edge.



Figure 4. The slow-manifold radial and vertical velocity components for the rigid particles,







Figure 5. For the steady symmetric rotating cylinder flow, the coordinates of the periodic orbit
that acts as an attractor for buoyant particles as a function of particle diameter (a-b) and frame
rotation (c). Flow parameters are listed in Table 1 and correspond to Experiment 1 for (a-b) and
Experiment 1d for (c-d).



Figure 6. (a): The Q_a function for the steady, axisymmetric, cylinder flow with the same parameter setting (see Experiment 1a) as for Figure 3a-c, and plotted in (x,z) along with the streamlines of the overturning circulation. The thick rigid curve corresponds to $Q_a = 0$. (b): The same parameter settings, except Ω has been raised from 0 to 0.3 (Rossby number \cong 1) (c): $\Omega = 1.0$. (Rossby number \cong 0.2).



Figure 7. Domain of attraction for the attractors in (a) steady symmetric (Experiment 1 in Table 1), (b) steady asymmetric (Experiment 2 in Table 1), and (c) time-periodic asymmetric rotating cylinder flow (Experiment 3 in Table 1). (These are the same 3 experiments that were used to produce Fig. 2.) The color indicates the height (i.e., value of z-coordinate) of the final crossing of a trajectory with the Poincare section, as a function of particle's release location. Particles attracted to the same attractor thus correspond to same color.







1000 Figure 9. For the "reversed flow" experiment (Experiment 1e in Table 1), z-position of a sample

particle trajectory as function of time for 3 values of Ω : 0 (top), 0.6 (middle), and 2 (bottom).

1002 Time t is in dimensionless units (but since our scaling coefficient for time is equal to 1 sec, the

1003 numbers on the x-axis can also be read as dimensional time in sec.)



Figure 10. For the "slow overturn" Experiment 1c from Table 1, color indicates the final zcoordinate of a particle's trajectory at the end of integration time as a function of particle's release location for 2 values of d: (a) 5×10^{-4} and (b) 10^{-3} . Yellow corresponds to particles rising up to the top, whereas green indicates the basin of attraction of the subsurface attracting periodic orbit. The insets at the left side of each frame show a sample trajectory whose release location is indicated by the black dot.

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1030 Figure 11. Same as Fig. 2(d-f) but with b = 3.8.



1044 Figure 12. For the steady perturbed system (Experiment 2 in Table 1), changes in the location of 1045 the attracting periodic orbits, basins of attractions, and time of attraction as a function of particle 1046 diameter d (and thus $\tilde{\epsilon}$). (a,d,g) show z-coordinate of the last crossing of trajectory with the x-z Poincare plane as a function of release location; flat regions are basins of attraction for the 2 1047 attactors. (b,e,h) show 20 trajectories in 3d released along a vertical line at y = 0, x = 0.334, 1048 0.05 < z < 0.95; denser cores indicate attractors. (c,f,i) show crossing of the same select 20 1049 trajectories with the x-z Poincare plane, color coded by time; blue corresponds to release 1050 location, yellow corresponds to final positions. 1051

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1056 Figure 13. For the steady perturbed system (Experiment 2 in Table 1), changes in the location of the attracting periodic orbits, basins of attractions, and time of attraction as a function of frame 1057 rotation Ω . (a,d,g) show z-coordinate of the last crossing of trajectory with the x-z Poincare plane 1058 as a function of release location; flat regions are basins of attraction for the 2 attactors. (b,e,h) 1059 show 20 select trajectories in 3d released along a vertical line at y = 0, x = 0.334, 0.05 < z < 0.051060 0.95; denser cores indicate attractors. (c,f,i) show crossing of the same 20 trajectories with the x-1061 z Poincare plane, color coded by time; blue corresponds to release location, yellow corresponds 1062 1063 to final positions.

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