

# Enhancing characterization of organic nitrogen components in aerosols and droplets using high-resolution aerosol mass spectrometry

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**Abstract.** This study aims to enhance the understanding and application of the Aerodyne high-resolution aerosol mass spectrometer (HR-AMS) for the comprehensive characterization of organic nitrogen (ON) compounds in aerosol particles and atmospheric droplets. To achieve this goal, we analyzed/seventy-five N-containing organic compounds, representing a diverse range of ambient ~~ON~~non-organonitrate ON (NOON) types, including amines, amides, amino acids, N-heterocycles, protein, and humic acids. Our results show that ~~ON~~NOON compounds can produce significant levels of  $\text{NH}_x^+$  and  $\text{NO}_x^+$  ion fragments, which are typically recognized as ions representative of inorganic nitrogen species. We also ~~discovered~~identified the presence of  $\text{CH}_2\text{N}^+$  at  $m/z = 28.0187$ , an ion fragment that is rarely quantified in ambient datasets due to substantial interference from ~~air-related~~ $\text{N}_2^+$ . As a result, ~~we determined that the utilization of~~ an updated calibration factor of 0.79 is necessary ~~to accurately quantify ON content~~for accurate NOON quantification using via aerosol mass spectrometry. AMS. We also assessed the relative ionization efficiencies (RIEs) for ~~different ON~~various NOON species and found that the average RIE ~~of ON~~for NOON compounds ( $1.52 \pm 0.58$ ) aligns with the commonly used default value of 1.40 for organic aerosol ~~(OA)~~. Moreover, through a careful examination of the HR-AMS mass spectral features of various ~~ON~~NOON types, we propose fingerprint ion series that can aid the ~~ON~~NOON speciation analysis. ~~The~~For instance, the presence of  $\text{C}_n\text{H}_{2n+2}\text{N}^+$  ions is closely linked with amines, with  $\text{CH}_4\text{N}^+$  indicating primary amines,  $\text{C}_2\text{H}_6\text{N}^+$  suggesting secondary amines, and  $\text{C}_3\text{H}_8\text{N}^+$  representing tertiary amines.  $\text{C}_n\text{H}_{2n}\text{NO}^+$  ions (especially for  $n$  values of 1-4) are very likely derived from amides. The co-existence of three ions,  $\text{C}_2\text{H}_4\text{NO}_2^+$ ,  $\text{C}_2\text{H}_3\text{NO}^+$ , and  $\text{CH}_4\text{NO}^+$ , serves as an indicator for the presence of amino acids. Additionally, the presence of  $\text{C}_x\text{H}_y\text{N}_2^+$  ions indicates the occurrence of 2N-heterocyclic compounds. Notably, an elevated abundance of  $\text{NH}_4^+$  is a distinct signature for amines and amino acids, as inorganic ammonium salts produce only negligible amounts of  $\text{NH}_4^+$  in HR-AMS.

Finally, we quantified the ~~ON~~NOON contents in submicron particles ( $\text{PM}_{10}$ ) and fog ~~waters~~water in Fresno, California and  $\text{PM}_{10}$  in New York City (NYC). Our results revealed the substantial presence of amino compounds in both Fresno and NYC aerosols, whereas concurrently collected fog ~~waters~~water contained a broader range of ~~ON~~NOON species, including N-

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35 containing aromatic heterocycle (e.g., imidazoles) and amides. These findings highlight the significant potential of employing  
the widespread HR-AMS measurements of ambient aerosols and droplets to enhance our understanding of the sources,  
transformation processes, and environmental impacts associated with ~~ON~~NOON compounds in the atmosphere.

## 1. Introduction

40 Nitrogen-containing organic species (ON) have been ubiquitously observed in atmospheric particles and aqueous phases  
(e.g., fog/cloud droplets, and rainwater). These compounds can constitute a significant portion (~ 30%) of the total airborne  
nitrogen (Cape et al., 2011) and contribute to up to 25% of the total nitrogen (N) deposition flux (Jickells et al., 2013;  
Kanakidou et al., 2016). However, our understanding of atmospheric ON is generally limited compared to inorganic N species  
such as ammonium, nitrate and nitrite. This disparity arises from the diverse nature of ON, which encompasses a wide range  
of compounds with varying carbon numbers, functional groups, and physicochemical properties.

45 ON compounds, originating from multiple emission sources and formed through diverse chemical processes, exhibit  
notable spatial and temporal variations in the atmosphere. The presence of airborne ON carries significant implications for  
regional air quality, earth's climate, terrestrial and aquatic ecosystem, and human health (Cape et al., 2011; Cornell et al., 1995;  
Liu et al., 2017). Among these compounds, aliphatic amines have been found to play a significant role in new particle formation  
and growth (Smith et al., 2010; Almeida et al., 2013; Yao et al., 2018), and contribute to secondary organic aerosol formation  
(Murphy et al., 2007; Yu et al., 2017; Song et al., 2017). Furthermore, ON species have been demonstrated to significantly  
50 influence the optical and hygroscopic properties of aerosols (Powelson et al., 2014; Lee et al., 2013; Rovelli et al., 2017),  
thereby influencing cloud formation and atmospheric water cycles. These observations highlight the importance of studying  
ON compounds and their behaviorsbehaviours in the atmosphere.

55 Bulk ON, which represents the nitrogen content within organic compounds, is typically quantified by subtracting the  
inorganic nitrogen (IN) content from that of the total nitrogen (TN):  $ON = TN - IN$ . IN is the combined nitrogen amount in  
ammonium, nitrate and nitrite (Zhang and Anastasio, 2001). Measurement of TN involves converting N-containing species  
into inorganic ions (e.g.,  $NO_3^-$ ,  $NO_2^-$ , and  $NH_4^+$ ) or nitrogen gases (e.g.,  $NO_x$  and  $NH_3$ ) using chemical, photochemical, or  
high-temperature oxidation approaches (Cornell et al., 2003; Zhang and Anastasio, 2003a; Jickells et al., 2013). However,  
quantifying ON using the differing methods can introduce significant uncertainties, due to factors such as incomplete  
transformation of ON compounds and aggregation of measurement errors (Cornell et al., 2003; Cape et al., 2011). The latter  
60 is especially important when ON is minor compared to IN in the sample. Moreover, ON measurements have traditionally  
focused on the water-soluble fraction (WSON). However there are observations that water-insoluble ON (WION) can  
constitute a large fraction of the total ON and, sometimes may be more important than WSON (Miyazaki et al., 2011; Russell  
et al., 2003). Furthermore, most studies investigating atmospheric ON have relied on samples collected over periods of hours  
to days. The limited time resolution of these measurements hampers the ability to capture the rapid evolution processes of ON  
65 species in the atmosphere.

The speciation of atmospheric ON has been ~~analyzed~~[analysed](#) using various techniques, including ion chromatography (Vandenboer et al., 2011; Verrielle et al., 2012; Parworth et al., 2017; Liu et al., 2021; Place et al., 2017), gas chromatography coupled with different detectors (Özel et al., 2011; zel et al., 2009; Akyüz, 2007), infrared (IR) and nuclear magnetic resonance (NMR) spectroscopy (Herckes et al., 2007), liquid chromatography with mass spectrometry (LC-MS) (Ruiz-Jimenez et al., 2012; Samy and Hays, 2013; Samy et al., 2011; Ye et al., 2017), high-resolution electrospray ionization mass spectrometry (ESI-MS) (Laskin et al., 2009; Altieri et al., 2012; Rincon et al., 2012; Shi et al., 2020), aerosol mass spectrometry (Junninen et al., 2010; Huang et al., 2012; Setyan et al., 2014; Zhou et al., 2016; Ge et al., 2014; Xu et al., 2017; Huang et al., 2021; Yu et al., 2019), chemical ionization mass spectrometry (Yao et al., 2018; Zheng et al., 2015; Yao et al., 2016; Yu and Lee, 2012), and nano-secondary ion mass spectrometry (Li et al., 2016). It is important to note that no single technique can comprehensively capture the full spectrum of ON species. However, certain mass spectrometric techniques may provide broader coverage (Cape et al., 2011). A variety of ON species have been detected in the atmosphere, including amines, amino acids, urea, amides, nitriles, ~~organic nitrates~~[organonitrates](#), nitro-compounds, and N-heterocyclic compounds, and more. However, when considered individually, these identified ON classes often contribute only a small fraction to the overall ON pool. For example, amino compounds, including amines, amino acids, peptides and proteins, were found to constitute less than 20% of the total ON in the Central Valley of California – a region known for intense agricultural emissions of these compounds (Zhang and Anastasio, 2001; Zhang et al., 2002). Moreover, in some instances, ON compounds were only qualitatively characterized without quantitative information. These limitations restrict a comprehensive understanding of the chemistry of atmospheric ON compounds and their environmental impacts.

The Aerodyne high-resolution time-of-flight aerosol mass spectrometer (HR-AMS) is a highly promising tool for characterizing ON in atmospheric condensed phases. It utilizes thermal vaporization and electron impact (EI) ionization techniques to quantify aerosol components and has been widely used for real-time measurements of non-refractory components in submicron aerosols (PM<sub>1</sub>) with fast time resolution (Zhang et al., 2020; Decarlo et al., 2006). An important feature of the HR-AMS is its ability to differentiate fragment ions with the same integer mass-to-charge ratio ( $m/z$ ) but slightly different exact masses. This capability allows for the determination of ~~the~~ elemental compositions of the detected ions (Decarlo et al., 2006) and the average atomic ratios, such as oxygen-to-carbon (O/C), hydrogen-to-carbon (H/C), and nitrogen-to-carbon (N/C), for the bulk organic aerosol (OA) (~~Aiken et al., 2008; Ma et al., 2021~~)([Aiken et al., 2008; Ma et al., 2021](#)). These ratios are particularly valuable for estimating the total ON content present in ~~the~~ aerosol samples.

While the speciation analysis of ON using the HR-AMS mass spectra (MS) can be challenging due to the extensive fragmentation of parent molecules (~~Drewnick, 2012~~)([Drewnick, 2012](#)), it is still possible to identify different types of ON by leveraging unique spectral fingerprints, such as those observed for aliphatic amines (~~Murphy et al., 2007; Sun et al., 2011~~)([Murphy et al., 2007; Sun et al., 2011](#)), N-heterocycles (~~Hawkins et al., 2018; Kim et al., 2019~~)([Hawkins et al., 2018; Kim et al., 2019](#)), and organonitrates (~~Farmer et al., 2010~~)([Farmer et al., 2010](#)). Furthermore, the HR-AMS offers a new perspective for source identification of ON in ambient air. Specifically, the highly time-resolved HR-AMS data, typically collected at intervals of 2-5 minutes during field studies, offer valuable insights into the co-variation of ON species with other

100 known components and permit the application of multivariate factor analysis techniques to differentiate components that exhibit similar patterns of behavior. A recent study has successfully applied positive matrix factorization (PMF) on the N-containing ions in an HR-AMS field dataset, and determined ON factors that can facilitate the inference of specific sourced or formation processes (Qi et al., 2022).

105 To achieve a comprehensive characterization of atmospheric ON using HR-AMS, it is necessary to optimize and validate the technique. While the interpretation of OA behaviors has often relied on the O/C and H/C ratios (Aiken et al., 2008; Canagaratna et al., 2015), the N/C ratio has received limited attention—, and previous AMS measurements on ON standards are scarce. The work by Aiken et al. (2008) includes 27 ON compounds; however, ON spectra were not reported. These spectra are reported in this study. Ge et al. (2014) introduced a method using HR-AMS to characterize amines and their degradation products in postcombustion CO<sub>2</sub> capture (PCCC) processes. In that study, we performed analysis of the AMS spectra of 12 amino compounds and the NIST spectra for 37 ON compounds that were identified as the degradation products from PCCC amines. In a separate study, Price et al. (2023) tested the relative ionization efficiencies (RIEs) of three ON compounds. Additionally, Farmer et al. (2010) and Day et al. (2022) proposed AMS-based methods for quantifying organonitrates (aka organic nitrates), which have been broadly adopted by various AMS studies (e.g., Huang et al., 2021; Yu et al., 2019; Xian et al., 2023; Lin et al., 2021; Zhu et al., 2016).

110  
115 To address ~~this~~the knowledge gap, ~~our~~this study conducted analyses on a large set of ON standards, encompassing diverse chemical types likely present in the atmosphere. The HR-AMS spectra obtained were meticulously analyzed to characterize N-containing ion fragments and establish the relationship between the N/C ratios measured by HR-AMS and the nominal N/C values of the compounds. In addition, we carefully investigated mass spectral features of the ON standards and proposed speciation analysis protocols. Furthermore, we evaluated the effectiveness and applicability of the proposed method by examining three ambient HR-AMS datasets. Our results demonstrate that the optimized HR-AMS technique and the proposed analysis protocols carry great promise at enhancing our ability to investigate the sources, transformation processes, and environmental impacts of atmospheric ON.

## 2. Experiments and data analysis

125 A comprehensive set of seventy-five ~~pure~~ ON standards were analyzed by the HR-AMS, and detailed information regarding these standards can be found in Table S1 of the Supplementary Material. These standards, which include 27 amines (No. 1-27), 6 amides (No. 28-33), 27 amino acids (No. 34-59 and 71), 4 nitro-compounds (No. 60-63), 7 N-containing heterocycles (No. 64-70), 1 protein (No. 72), and 3 humic-like substances (No. 73-75), were mostly procured from Sigma-Aldrich with purities exceeding 98%. ~~In addition, to probe the relative ionization efficiencies (RIEs) of different ON~~Noted that organonitrates are not included, because the fragmentation of organonitrates in the AMS predominantly generates NO<sub>x</sub><sup>±</sup> ions (generally >50%, mostly NO<sup>+</sup> at *m/z* 30), and organic moieties that lack nitrogen. These ions cannot be used as fingerprint ions for quantifying organonitrates. Instead, the quantification of organonitrates is commonly achieved based on the differences

of  $\text{NO}^+/\text{NO}_2^+$  ratios between organic nitrates and ammonium nitrate (Farmer et al., 2010; Day et al. 2022). For clarity, we denote the compounds studied here non-organonitrate ON (NOON). In addition, to probe the RIEs of different NOON species, we analyzed 18 mixtures containing sulfate (in the form of ammonium sulfate) (Fisher Chemicals, > 99%) and individual ON/NOON standards in a 1:1 mass ratio (Table S2). All solutions were prepared using purified water (>18.2 M $\Omega$ ·cm) obtained from a Milli-Q system (Millipore, USA).

The procedures for HR-AMS analysis of liquid samples were detailed have been comprehensively documented in previous work prior works (Sun et al., 2010; Ge et al., 2014; Ge et al., 2017; O'brien et al., 2019; Niedek et al., 2023). Briefly, Each NOON standard was carefully weighted, dissolved in Milli-Q water, and diluted to a solute concentration of ~20 ppm. A portion of the aqueous solutions were solution (20-40 mL) was dispensed into a sample vial and then nebulized using a Collison-type atomizer with high purity argon as the carrier gas. Aerosol particles were dehydrated in a diffusion dryer was employed to dehumidify the aerosol-filled with silica-gel, reducing the relative humidity (RH) to <5%, before it was introduced into the HR-AMS. The HR-AMS was operated at a vaporizer temperature of ~ 600 °C and was switched alternated between the highly sensitive V-mode and the high mass resolution W-mode ( $m/\Delta m \sim 5000$ ). For this study,

W-mode data with an extended  $m/z$  range extended up to 500 amu was used, aligning with our focus on characterizing the chemical composition. To eliminate the interference of the  $\text{N}_2^+$  signal on  $\text{CH}_2\text{N}^+$  at  $m/z = 28.033$  and  $\text{CO}^+$  at 28.01, the aerosol generation system was initially purged by atomizing Milli-Q water under argon until the  $\text{N}_2^+$  signal measured by the HR-AMS reached a sufficiently low level. HR-AMS data for each sample were recorded under stable particle flow conditions, and the high-resolution mass spectrum (HRMS) of each NOON standard was derived from the average of at least three stable runs, each lasting 150 seconds. To account for any potential contamination or background signals, a blank measurement was conducted by aerosolizing Milli-Q water between every two samples, following the same procedure. Typically, the measured signals of Milli-Q water were minimal and had negligible influence on the samples.

The HR-AMS MS data were analyzed using the standard AMS data analysis software, namely SQUIRREL v1.46/v1.51H and PIKA v1.06/v1.10H (<http://cires.colorado.edu/jimenezgroup/ToFAMSResources/ToFSoftware/index.html>) written in Igor Pro (Wavemetrics, Portland, USA). During data post-processing, measures were taken to eliminate interferences caused by air components on organic fragments (e.g.,  $\text{N}_2^+$  at  $m/z$  28,  $\text{O}_2^+$  at  $m/z$  32). Given that the relative humidity RH values recorded during the measurements remained consistently very low (<5%), the influence of particle-bound water was deemed negligible. Thus, the MS mass spectral analyses were conducted using the directly measured signals of  $\text{O}^+$ ,  $\text{OH}^+$ ,  $\text{H}_2\text{O}^+$  and  $\text{CO}^+$ .

HR-AMS data acquired from the online measurements of  $\text{PM}_{10}$  during July-August 2009 in NYC (Sun et al., 2011), and during January 2010 in Fresno, CA (Ge et al., 2012b; Ge et al., 2012a) were used in this study. Additionally, a set of 11 fog water samples collected during January 2010 in Fresno (Kim et al., 2019) were measured and analyzed following the ON standard analysis procedure described above. were used in this study. Additionally, we examined a set of 11 fog water samples collected during January 9-16, 2010, in Fresno. These fog waters were collected using two Caltech Active Strand Cloud water Collectors (CASCC). Immediately after collection, the fog water samples were filtrated with 0.45  $\mu\text{m}$  Syringe Filters (Pall Laboratory) and then stored in pre-cleaned high-density polyethylene (HDPE) bottles within a freezer at -20°C until analysis.

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The analysis of fog water samples involved the utilization of the same HR-AMS instrument and a similar procedure employed for the NOON standards. More details on fog water analysis can be found in Kim et al. (2019). The analytical technique developed in this work was applied on these datasets to assess its effectiveness and applicability.

### 3. Results and discussion

#### 170 3.1. N/C calibration factor

Accurate determination of elemental ratios using HR-AMS spectra requires calibrations due to the loss of neutral fragments and the inability to detect very small  $m/z$  ions such as  $H^+$ . The calibration factor for N/C reported in Aiken et al. (2008) was established based on 27 ON standards, mainly consisting of amino acids. ~~The~~ the N/C ratios derived from HR-AMS spectra for these standards agree well with the nominal values (slope = 0.96;  $r^2 = 0.95$ ) and the estimated uncertainty associated with this calibration was 22%. In this study, we re-evaluated the N/C calibration factor using a considerably larger dataset comprising 75 ~~ON~~ standards that represent a broad range of N-containing functional groups (see Table S1). The new calibration plot (Figure 1a) shows a similar slope ( $0.99 \pm 0.03$ ,  $r^2 = 0.84$ ), albeit with a higher uncertainty of 32% compared to the previous calibration. This increased uncertainty is partially attributed to the large positive biases observed for a few primary amines (further details provided in Section 3.4). In addition, the calibration factors for O/C (0.75) and H/C (0.93) (Figures 1c and 1d), determined based on the ~~ON~~NOON standards in this study, exhibit good agreements with those reported by Aiken et al. (2008) (0.75 for O/C and 0.91 for H/C).

However, it should be noted that the N/C calibration factors reported here (0.99) and in Aiken et al. (2008) (0.96) include contributions from  $NH_x^+$  and  $NO_x^+$  ions. Yet, these ions are typically assigned to inorganic ammonium and nitrate, respectively, during the analysis of ambient AMS data. Consequently, adjustments should be made to the N/C correction factors to ensure accurate estimation of ~~ON~~NOON content in ambient aerosols. Table 1 shows that  $NH_x^+$  ions are consistently observed in the MS of various types of ~~ON~~NOON standards, contributing approximately 10% to the total N content on average.  $NO_x^+$  ions are particularly significant in nitro-compounds, contributing on average 16.2 ( $\pm 8.3$ )% to the total N. ~~A previous study conducted on organic nitrates also demonstrated the predominant occurrence of  $NO_x^+$  ions (mostly  $NO^+$  at  $m/z$  30) compared to the much less abundant  $C_xH_yO_zN_p^+$  fragments observed in the HR-AMS mass spectra (Farmer et al., 2010).~~ Based on the standards analyzed in this study, excluding the contributions from  $NH_x^+$  and  $NO_x^+$  would lead to an overall reduction of  $\sim 13\%$  in the N/C calibration ratio. In addition, the  $CH_2N^+$  ion ( $m/z = 28.0187$ ) is an important fragment in most ~~ON~~NOON compounds, particularly amines. However, due to interference from the adjacent  $N_2^+$  ion, it is necessary to exclude the  $CH_2N^+$  ion in the analysis of ambient HR-AMS data. By excluding this ion, the N/C calibration ratio would further decrease by  $\sim 7\%$ .

Considering the exclusion of  $NH_x^+$ ,  $NO_x^+$ , and  $CH_2N^+$  ions, we have derived a new N/C calibration factor of 0.79 (Figure 1b). We highly recommend using this value to determine the N/C ratio for ambient samples. ~~However, it is important to note~~ It should be noted again that even with this new ratio is for NOON compounds, there may still be underestimations of the ON content in cases where aerosols/droplets contain a significant amount of urea or organic nitrates. This is organonitrates. The

underestimation also occurs in case of significant presence of urea because a large fraction of the N mass (generally > 50%) is present in  $\text{NH}_4^+$  for urea and in  $\text{NO}_x^+$  for organic nitrates.

### 200 3.2 Relative ionization efficiencies (RIEs) of different $\text{ON}_{\text{NOON}}$ species

205 TheFor the AMS measurements, relative ionization efficiency (RIE) defined as the ratio of the ionization efficiency of a species to that of nitrate, quantified in terms of mass (Canagaratna et al., 2007); it is crucial for accurately quantifying the mass of the specific species in question. Here, the RIEs of  $\text{ON}_{\text{NOON}}$  species were determined by analyzing the HR-AMS spectra of 18 mixtures containing  $\text{ON}_{\text{NOON}}$  species and sulfate in 1:1 mass ratio. The measured ratios between the total organic signal and the sulfate signal varied from approximately 0.3 to 2, with an average of 1.33 (Figure S1). Assuming that the aerosol composition measured by HR-AMS corresponds to the solute composition in the original mixtures, these ratios represent the RIEs of different  $\text{ON}_{\text{NOON}}$  compounds relative to sulfate. In this study, since the RIE of sulfate relative to nitrate (based on mainly  $\text{NO}^+$  and  $\text{NO}_2^+$ ) was found to be 1.14, the average RIE of  $\text{ON}$  species relative to nitrate was determined to be  $1.52 \pm 0.58$ . It should be noted that the obtained result from the analysis of 18 mixtures aligns well with the average RIE of  $1.6 \pm 0.5$  reported by Price et al. (2023).  $\text{NOON}$  species relative to nitrate was determined to be  $1.52 \pm 0.58$ . Varying mass ratios of  $\text{NOON}$  to sulfate in the solutions should have negligible influence on the determined RIE value. It is important to note that the obtained result from the analysis of 18 mixtures aligns very well with the average RIE of  $1.6 \pm 0.5$  reported by Price et al. (2023), which was based on a study involving only 3 compounds (4-nitrocatechol, isosorbide mononitrate and triammonium citrate). This RIE value (1.52) is only ~8.6% higher than the widely used default value of 1.4 for the quantification of total OA (Canagaratna et al., 2007). However, considering the inherent measurement uncertainties, such as variations in volatility leading to deviations in HR-AMS-determined particle compositions from the original mixture ratios, and the potential influences of impurities in the solutions on the measured mass of  $\text{ON}_{\text{NOON}}$  and thus the resulting RIE, we recommend maintaining the default RIE of 1.4 for the quantification of  $\text{ON}_{\text{NOON}}$  in ambient aerosols and droplets.

### 220 3.3. Quantification of $\text{ON}_{\text{NOON}}$ in ambient aerosols/droplets

The mass concentration of  $\text{ON}_{\text{NOON}}$  can be determined based on HR-AMS measurements using the following equations:

$$\text{OC} = \text{OM}_{\text{mass}} / (\text{OM}/\text{OC}) \quad (1)$$

$$\text{ON}_{\text{NOON}} = \text{OC} \times (\text{N}/\text{C}) \times (14/12) \quad (2)$$

where OC represents the mass concentration of carbon present in OA,  $\text{OM}_{\text{mass}}$  denotes the measured total OA mass,  $\text{ON}_{\text{NOON}}$  denotes the mass of nitrogen in  $\text{NOON}$  species. The OM/OC (organic mass to organic carbon ratio) and N/C (calculated using the new correction factor of 0.79) can be derived from HR-AMS elemental analysis.

For liquid samples, such as aqueous extracts of collected filter samples, fog, cloud, and rain water, the  $\text{ON}_{\text{NOON}}$  concentrations can be determined in ppm or  $\text{mg} \cdot \text{L}^{-1}$  using the OC content measured by a total organic carbon (TOC) analyzer, typically representing the water-soluble OC (WSOC) in ppm or  $\text{mg} \cdot \text{L}^{-1}$ . The corresponding  $\text{WSOC}_{\text{NOON}}$  concentration can be calculated as follows:

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$$\frac{W_{\text{SON}}}{W_{\text{SONNOON}}} = \text{WSOC} \times (\text{N/C}) \times (14/12) \quad (3)$$

where the factor (14/12) is the ratio of the molar masses of nitrogen and carbon, respectively.

In the above calculations, accurate quantification of  $\frac{\text{ONNOON}}{\text{NOON}}$  relies on the accuracy of N/C determination, which can be influenced by the need to estimate  $\text{H}_x\text{O}^+$  (mainly  $\text{H}_2\text{O}^+$ ,  $\text{OH}^+$ ,  $\text{O}^+$ ) and  $\text{CO}^+$  signals for ambient datasets (Aiken et al., 2008).

The  $\text{H}_x\text{O}^+$  signals do not directly affect the N/C ratio, and hence have no impact on the  $\frac{W_{\text{SON}}}{W_{\text{SONNOON}}}$  calculation in Eq.

3. However, the current estimation of  $\text{H}_x\text{O}^+$  ( $\text{H}_2\text{O}^+ = 0.225 \text{CO}_2^+$ ,  $\text{OH}^+ = 0.25 \text{H}_2\text{O}^+$ ,  $\text{O}^+ = 0.05 \text{H}_2\text{O}^+$ ) for ambient OA (Aiken et al., 2008) can introduce uncertainties in the O/C and H/C ratios, thereby influencing the accuracy of the OM/OC ratio. If Eq. 1 is used to calculate the OC content, an overestimation of the OM/OC ratio would lead to a lower OC and an underestimation of  $\frac{\text{ONNOON}}{\text{NOON}}$  in Eq. 2. The estimation of  $\text{CO}^+$  ( $= \text{CO}_2^+$ ; (Aiken et al., 2008)) can also affect the elemental ratios and  $\frac{\text{ONNOON}}{\text{NOON}}$  quantification, albeit to a lesser extent compared to the  $\text{H}_x\text{O}^+$  ions. These uncertainties can be minimized,

though not entirely eliminated, by conducting measurements on HEPA-filtered ambient air during field campaigns.

Additionally, in the HR-AMS spectra of ambient samples, ON ions are typically observed at substantially lower intensities compared to adjacent ions, and are often located at the edges or between  $\text{C}_x\text{H}_y^+$  and  $\text{C}_x\text{H}_y\text{O}_z^+$  ions. For instance,  $\text{C}_x\text{H}_y\text{N}_1^+$  ions are often found between  $\text{C}_x\text{H}_{y-2}\text{O}_1^+$  and  $\text{C}_{x+1}\text{H}_{y+2}^+$  ions, while  $\text{C}_x\text{H}_y\text{O}_1\text{N}_1^+$  ions are located between  $\text{C}_x\text{H}_{y-2}\text{O}_2^+$  and  $\text{C}_{x+1}\text{H}_{y+2}\text{O}_1^+$  ions (refer to Figure S2 for two examples at  $m/z$  42 and 44). Therefore, accurate determination of ON ions requires careful optimization of the peak shape and peak width parameters prior to performing peak fittings. This optimization helps reduce uncertainties associated with fitting of ON ions, particularly those with low signal intensities. Furthermore, caution should be given when fitting ON ions with high  $m/z$  values (e.g.,  $>100$  amu) since the mass resolution of HR-AMS may not be sufficient to unambiguously distinguish ON ions from neighboring high  $m/z$  isobaric ions.

### 3.4. HR-AMS mass spectral features of $\frac{\text{ONNOON}}{\text{NOON}}$ compounds

Detailed analyses of the HR-AMS spectra of all tested  $\frac{\text{ONNOON}}{\text{NOON}}$  standards (Figure S3) reveal unique fragmentation patterns associated with specific functional groups. Aliphatic amines, characterized by the strong electron-donating ability of the nitrogen atom, predominantly undergo  $\alpha$ -cleavage, resulting in the formation of ion series of  $\text{C}_n\text{H}_{2n+2}\text{N}^+$ , e.g.,  $\text{CH}_4\text{N}^+$  ( $m/z$  30.0344),  $\text{C}_2\text{H}_6\text{N}^+$  ( $m/z$  44.0500),  $\text{C}_3\text{H}_8\text{N}^+$  ( $m/z$  58.0657),  $\text{C}_4\text{H}_{10}\text{N}^+$  ( $m/z$  72.0813),  $\text{C}_5\text{H}_{12}\text{N}^+$  ( $m/z$  86.0970), depending on the molecular structures (McLafferty and Turecek, 1993). Notably, primary amino compounds with the  $-\text{CH}_2-\text{NH}_2$  functional group, such as ethylamine, ethanolamine, and glycine (Figure 2a), exhibit the most abundant peak of  $\text{CH}_4\text{N}^+$  ( $m/z$  30.0344). In the case of butylamine, iso-butylamine and pentylamine, the  $\text{CH}_4\text{N}^+$  ion can constitute  $>60\%$  of the total signals in their respective MS, leading to significant positive biases in the N/C ratios determined by HR-AMS, as shown in Figures 1a and 1b. Additionally, secondary elimination from initial  $\alpha$ -cleavage may contribute significantly to the formation of  $\text{C}_n\text{H}_{2n+2}\text{N}^+$  ions for secondary and tertiary amines (e.g., diethylamine and triethylamine; Figure 2a).

$\alpha$ -cleavage can be an important fragmentation pathway for amides as well, causing the loss of alkyl groups and the formation of ion series of  $\text{C}_n\text{H}_{2n}\text{NO}^+$ , such as  $\text{CH}_2\text{NO}^+$  ( $m/z$  44.0136),  $\text{C}_2\text{H}_4\text{NO}^+$  ( $m/z$  58.0293),  $\text{C}_3\text{H}_6\text{NO}^+$  ( $m/z$  72.0449), and  $\text{C}_4\text{H}_8\text{NO}^+$  ( $m/z$  86.0606) (McLafferty and Turecek, 1993). Particularly, the  $\text{CH}_2\text{NO}^+$  ion (sometimes also  $\text{CHNO}^+$ ) can be



generated from the amide group (-CO-NH<sub>2</sub>), as observed for urea, pyrazinecarboxamide, and L-asparagine (Figure 2b).

Additional, cleavage of the C(O)-N bond can be significant for certain amides, resulting in the formation of

265 C<sub>3</sub>H<sub>3</sub>O<sup>+</sup> from acetaminophenol and C<sub>7</sub>H<sub>5</sub>O<sup>+</sup> from benzamide (Figure 2b).

Amino acids containing the functional group -CH(NH<sub>2</sub>)-COOH commonly produce significant signals of C<sub>2</sub>H<sub>4</sub>NO<sub>2</sub><sup>+</sup> (*m/z* 74.0242) resulting from α-cleavage. Subsequent losses of -OH and -CO from the C<sub>2</sub>H<sub>4</sub>NO<sub>2</sub><sup>+</sup> ion lead to the formation of C<sub>2</sub>H<sub>3</sub>NO<sup>+</sup> (*m/z* 57.0215) and CH<sub>4</sub>NO<sup>+</sup> (*m/z* 46.0293), respectively. This pattern is observed in amino acids such as serine, valine, cysteine, leucine, and tyrosine as depicted in Figure 2c.

270 2N-heterocyclic compounds show characteristic C<sub>x</sub>H<sub>y</sub>N<sub>2</sub><sup>+</sup> peaks in their spectra (Figure 2d). For example, imidazole and nitro-imidazole produce C<sub>x</sub>H<sub>y</sub>N<sub>2</sub><sup>+</sup> ions at *m/z* 67.0296/68.0374 (C<sub>3</sub>H<sub>3</sub>N<sub>2</sub><sup>+</sup>/C<sub>3</sub>H<sub>4</sub>N<sub>2</sub><sup>+</sup>) while histine generates C<sub>4</sub>H<sub>5</sub>N<sub>2</sub><sup>+</sup>/C<sub>4</sub>H<sub>6</sub>N<sub>2</sub><sup>+</sup> at *m/z* 81.0452/82.0531. The imidazolyl compounds also produce ion series of C<sub>3+n</sub>H<sub>5+n</sub>N<sub>2</sub><sup>+</sup> (*n*>1). On the other hand, compounds containing the pyrazinyl group show notable C<sub>4+n</sub>H<sub>4+n</sub>N<sub>2</sub><sup>+</sup> series, with C<sub>4</sub>H<sub>4</sub>N<sub>2</sub><sup>+</sup> at *m/z* 80.0374 representing pyrazinecarboxamide.

275 Nitro-compounds containing the -NO<sub>2</sub> group (e.g., nitrophenol, 2-nitrobenzaldehyde and 2-nitrobenzyl alcohol) display much higher NO<sup>+</sup>/NO<sub>2</sub><sup>+</sup> ratio (averaging 8.64 for the 4 compounds tested in this study). This ratio is similar to those observed for organic nitrates (i.e., RONO<sub>2</sub>) (Farmer et al., 2010). In comparison, the NO<sup>+</sup>/NO<sub>2</sub><sup>+</sup> ratio in pure ammonium nitrate is much lower, which is 2.69 in this work, ~2.7 in Fry et al. (2009) and 2.4 in Bruns et al. (2010). In addition, although not measured in this study, nitriles have been shown to generate principal ion series of C<sub>n</sub>H<sub>2n-1</sub>N<sup>+</sup> and C<sub>n</sub>H<sub>2n-2</sub>N<sup>+</sup> (McLafferty and Turecek, 1993).

280 In the HR-AMS spectra of ONNOON compounds, the contributions of molecular ions are generally low, averaging around 3.3% of the total signals. (see average contributions of different NOON categories in Table 1). The relationship between the compounds' molecular weights and the abundance of molecular ions is not straightforward (Figure S4). However, for compounds with stable structures including benzene or heterocyclic rings, a high abundance of molecular ion peak is typically observed. For example, pyrazole and imidazole exhibit a prominent C<sub>3</sub>H<sub>4</sub>N<sub>2</sub><sup>+</sup> peak at *m/z* 68.0374, (21.2 for pyrazole and 28.2% for imidazole), 1,3-phenylenediamine shows a C<sub>6</sub>H<sub>8</sub>N<sub>2</sub><sup>+</sup> peak at *m/z* 108.0687, (20.6%), 4-aminophenol displays a C<sub>6</sub>H<sub>7</sub>NO<sup>+</sup> peak at *m/z* 109.0528, (17.1%), and nicotinamide exhibits a C<sub>6</sub>H<sub>6</sub>N<sub>2</sub>O<sup>+</sup> peak at *m/z* 122.0480, (5.8%). It is worth noting that a few simple and low molecular weight aliphatic amines, such as methylamine, (17.0%), dimethylamine, (15.8%), and trimethylamine, (12.4%), and urea (a simple amide), (8.1%), also show relatively high contributions from molecular ions in their HR-AMS spectra as well.

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### 3.5. Speciation analysis of ONNOON in ambient aerosols and droplets

The analysis of ONNOON compounds or classes in AMS spectra of ambient samples presents challenges due to their low concentrations, the complexities and uncertainties involved in assigning ion fragments to potential parent molecules. However, by examining the spectral features as discussed earlier, we can effectively address these challenges. To provide a

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295 comprehensive overview, we have compiled the mass contributions of different N ion categories to the total ON mass for  
different types of  $\text{ONNOON}$  compounds and summarized them in Table 1.

300 With the exception of amides, the majority of ON mass originates from  $\text{C}_x\text{H}_y\text{N}_1^+$  ions, accounting for over 60% of the  
total mass. This is particularly true for amines and amino acids. Considering the atmospheric abundances of amines (Ge et al.,  
2014b, 2011a, b), significant presence of  $\text{C}_x\text{H}_y\text{N}_1^+$  ions strongly suggest the presence of amines. Moreover,  $\text{C}_n\text{H}_{2n+2}\text{N}^+$  ions  
305 may indicate the existence of aliphatic amines. To evaluate this further, we calculated the fractional contributions of three  
specific ions –  $\text{CH}_4\text{N}^+$ ,  $\text{C}_2\text{H}_6\text{N}^+$  and  $\text{C}_3\text{H}_8\text{N}^+$  (left panel of Figure 3) – in the HR-AMS MS of different  $\text{ONNOON}$  classes. It is  
evident that these three ions are **only clearly** abundant in amines rather than in other types of N-containing species. Furthermore,  
a close examination of the distribution of these ions among different types of amines (right panel of Figure 3) reveals that  
 $\text{CH}_4\text{N}^+$  is mainly associated with primary amines (e.g., methylamine, ethylamine, butylamine, pentylamine and ethanolamine),  
305  $\text{C}_2\text{H}_6\text{N}^+$  is more likely linked with secondary amines (e.g., dimethylamine and diethylamine), and  $\text{C}_3\text{H}_8\text{N}^+$  is mainly produced  
by tertiary amines (e.g., trimethylamine and triethylamine).

Another intriguing observation is that only amines and amino acids can generate significant  $\text{NH}_4^+$  ( $m/z$  18.0344) signal  
compared to other  $\text{ONNOON}$  standards (Figure 4). Conversely, the HR-AMS spectra of pure ammonium nitrate, ammonium  
sulfate and ammonium chloride (Figure S6) do not produce the  $\text{NH}_4^+$  ion. This disparity is likely due to the cleavage of C-N  
310 bond, accompanied by hydrogen migration to the  $\text{NH}_x^+$  ( $x=0, 1, 2$ ) fragments. Importantly, this finding strongly indicates that  
 $\text{NH}_4^+$  can serve as an HR-AMS marker for the identification of amino compounds, including both amines and amino acids.

In addition, the contribution of  $\text{C}_x\text{H}_y\text{N}_1\text{O}_1^+$  ions to the ON mass is generally small (<6%), except for amides, where it can  
reach ~26% (Table 1). This observation suggests that the abundance of  $\text{C}_x\text{H}_y\text{N}_1\text{O}_1^+$  ions, particularly  $\text{CH}_2\text{NO}^+$ , can serve as an  
indicator of the presence of amides. **Furthermore in addition**, we propose that the co-occurrence of  $\text{C}_2\text{H}_4\text{NO}_2^+$  ( $m/z$  74.0242),  
315  $\text{C}_2\text{H}_3\text{NO}^+$  ( $m/z$  57.0215), and  $\text{CH}_4\text{NO}^+$  ( $m/z$  46.0293) can act as tracer ions for amino acids. Furthermore, an enrichment of  
 $\text{C}_x\text{H}_y\text{N}_2^+$  ions suggests the presence of certain 2N-heterocyclic compounds, such as imidazoles and pyrazines (Kim et al., 2019;  
Hawkins et al., 2018). Finally, a high  $\text{NO}^+/\text{NO}_2^+$  ratio is likely a signature indicating the abundance of nitro-compounds/organic  
nitrates (Farmer et al., 2010; Fry et al., 2009).

320 Since AMS uses standard 70-eV electron ionization, the obtained spectra are expected to be comparable with the EI spectra  
available in the NIST database (National Institute of Standards and Technology Standard Reference Database 1A, NIST 11,  
Software version 2.0g). Figure S5 shows the correlation coefficients (Pearson's  $r^2$ ) between the AMS-measured spectra  
(summed in unit mass resolution) and the NIST spectra for the 75 ~~pure~~  $\text{ONNOON}$  standards. The level of spectral similarity  
varies among different types of  $\text{ONNOON}$  compounds. Higher correlations are observed for amines (average  $r^2=0.70$ ) and  
amides (average  $r^2=0.63$ ), while amino acids exhibit low correlations (average  $r^2=0.23$ ). The  $r^2$  values for N-heterocycles and  
325 nitrocompounds are 0.57 and 0.50, respectively. Although an increase in molecular weight generally tends to decrease the  
spectral similarities, the influence of molecular weight is not particularly evident (Figure S5).

A more detailed analysis reveals that the AMS spectra of aromatic compounds with ring structures exhibit close  
agreements with the corresponding NIST spectra (e.g., 1,3-phenylenediamine, 4-aminophenol, and 2-picolinic acid in Figure

S7). Conversely, spectra of aliphatic molecules (long-chain or branched) show less similarity and contain more signals towards low  $m/z$  range compared to the NIST spectra (e.g., tri-n-amyamine and lysine in Figure S7). Additionally, AMS spectra of oxygenated compounds tend to deviate more significantly from their NIST counterparts than other compound types. This divergence can be attributed to additional thermal energy provided by the 600°C AMS oven, resulting in more extensive fragmentation compared to the NIST spectra. Indeed, previous investigations by our group have demonstrated that employing a lower AMS vaporizer temperature ( $T_{\text{vap}}$ ) (e.g., 250°C) can reduce fragmentation, enhance the signals of parent molecular ions and/or fingerprint fragments, and overall increase the resemblance to the corresponding NIST spectra (Ge et al., 2014)(Ge et al., 2014). Docherty et al. (2015) also showed that the OA MS changes substantially with variations in  $T_{\text{vap}}$ , and utilizing a lower  $T_{\text{vap}}$  can improve the analysis of some reduced OA species (with more  $\text{C}_x\text{H}_y^+$  ions). Docherty et al. (2015) also showed that the OA MS changes substantially with variations in  $T_{\text{vap}}$ , and utilizing a lower  $T_{\text{vap}}$  can improve the analysis of some reduced OA species (with more  $\text{C}_x\text{H}_y^+$  ions). Therefore, during field deployments (or offline laboratory studies), it would be valuable to periodically acquire AMS spectra at lower  $T_{\text{vap}}$  in addition to the standard 600°C spectra, as they may provide richer chemical information for compound speciation.

### 3.6. Characterization of ~~ON~~NOON in ambient aerosols and fog ~~waters~~water

Using the new N/C calibration factor (0.79), we determined the N/C ratios of NYC PM<sub>1</sub>, Fresno PM<sub>1</sub> and fog ~~waters~~water and summarized them in Figure 5. On average, Fresno PM<sub>1</sub> exhibited a higher N/C ratio compared to NYC (0.019 vs. 0.015), although both ratios were significantly lower than that of fog organics (0.078). The mean ~~ON~~NOON mass concentrations were estimated to be 70 ng·m<sup>-3</sup>, 120 ng·m<sup>-3</sup> and 2.2 mg·L<sup>-1</sup> for the three sample sets, respectively. The ON levels observed in NYC and Fresno PM<sub>1</sub> are relatively low compared to those observed in a forest site in southern China (0.2~1.1 μg m<sup>-3</sup>) (Yu et al., 2020) and in Hong Kong (0.24 ± 0.09 μg m<sup>-3</sup>) (Li et al., 2022; Yu et al., 2021), but are similar to ~~the ON level that~~ observed in a recent study conducted in summertime Nanjing, China (0.08~0.14 μg m<sup>-3</sup>) (Xian et al., 2023).

The distribution of ON ions in NYC PM<sub>1</sub>, Fresno PM<sub>1</sub> and fog water are further shown in Figure 6. Across all spectra,  $\text{C}_x\text{H}_y\text{N}_1^+$  ions dominated, accounting for 84%, 84% and 66% of the total ~~ON~~NOON mass respectively, indicating that amines are a major ON component in both cities. Notably, significant  $\text{C}_n\text{H}_{2n+2}\text{N}^+$  ions are observed, suggesting the presence of aliphatic amines, which aligns with previous findings (Zhang and Anastasio, 2003b, 2001). The presence of  $\text{C}_x\text{H}_y\text{NO}_1^+$  ions likely indicates the presence of amides.

Comparatively, the HR-AMS spectra of Fresno fog waters also exhibited prominent  $\text{C}_x\text{H}_y\text{N}_2^+$  ions, likely originating from imidazoles formed through aqueous-phase reactions (De Haan et al., 2009a; De Haan et al., 2009b; Kim et al., 2019). In addition, appreciable signals of  $\text{C}_2\text{H}_4\text{NO}_2^+$  ( $m/z$  74.0242),  $\text{C}_2\text{H}_3\text{NO}^+$  ( $m/z$  57.0215), and  $\text{CH}_4\text{NO}^+$  ( $m/z$  46.0293), indicative of amino acids, are observed in fog ~~waters~~water but are negligible in PM<sub>1</sub>, consistent with the findings reported in Zhang et al. (2002), where the quantification of amino compounds were achieved through derivatization HPLC analysis. The enrichment of amino acids in fog ~~waters~~water compared to PM<sub>1</sub> can be attributed to a significant fraction of amino acids originating from

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proteinaceous matter mainly associated with coarse mode particles, such as soil dust and pollen. While these amino compounds may evade effective detection by the AMS in real time, they can be scavenged by the fog droplets.

#### 4. Method limitations

In the preceding sections, we have explored the successful optimization of the HR-AMS methodology for the quantification and chemical characterization of ~~ON/NOON~~ species in atmospheric aerosols and droplets. Nonetheless, it is important to acknowledge certain limitations associated with this method. ~~Firstly, it~~

(1) It is important to note that previous studies have shown bimodal size distributions of ON in marine aerosols (Violaki and Mihalopoulos, 2010; Cornell et al., 2001; Miyazaki et al., 2010), with a significant contribution from supermicron particles ( $PM_{>1}$ ). For example, the average concentrations of WSON were found to be  $5.5 \pm 3.9$  and  $11.6 \pm 14.0$  nmol  $m^{-3}$  for coarse ( $PM_{1.3-10}$ ) and fine ( $PM_{1.3}$ ) particles, respectively, at a remote marine location in the Eastern Mediterranean (Violaki and Mihalopoulos, 2010). However, the HR-AMS has limitations in measuring ON in coarse particles in ambient air due to constraints in the transmission efficiency of the aerodynamic lens. ~~Secondly, as previously discussed in Section 3.6, the difference between Fresno  $PM_1$  and co-collected fog water could, in part, be attributed to this limitation.~~

(2) The AMS's thermal vaporization at  $\sim 600^\circ C$  ~~is not capable of~~ limits the detection of ~~vaporizing non-refractory materials/species~~, potentially ~~leading to resulting in the~~ underestimation of ~~ON/NOON~~ levels ~~by excluding if the amount of refractory ONN-containing species. Note is significant. It is worth noting that the contribution of refractory ON to TN total ON is still not yet well known. Thirdly, the~~ understood. Additionally, precise quantification of ~~ON/NOON~~ content faces challenges due to relatively significant uncertainties in the N/C correction factor and RIE values. Furthermore, the 70 eV ionization process leads to extensive fragmentation, which adds complexity to the identification of individual NOON compounds.

(3) The current method for analyzing liquid samples is applicable to only WSON. Future investigations may focus on using HR-AMS to analyze the water-insoluble ON (WION), for example, by measuring samples extracted with organic solvents (Jiang et al., 2022; Chen et al., 2017). ~~Lastly, while specific~~

(4) The current method provides insights into ion fragments ~~have been identified that are likely (fingerprint ions)~~ associated with ~~different specific~~ ON functional groups. However, the extensive fragmentation ~~caused/induced~~ by the ~~70 eV electron~~ 70 eV electron ionization in HR-AMS introduces inherent uncertainties in such ~~assignments. The identifications. For example, while a specific ion like  $CH_2N^+$  likely originates from amines, we cannot rule out the possibility (albeit low) of its association with other ON types. In real atmospheric samples containing multiple ON varieties, the overlay and interference of specific ion fragments can be significant, increasing the level of uncertainty. To mitigate this challenge, the incorporation of information/data from soft ionization mass spectrometry techniques (Wang et al., 2019; Song et al., 2022; Mao et al., 2022) in combination with alongside HR-AMS could greatly data can significantly enhance the ~~speciation~~ analyses of ~~ON/NOON~~ species, thereby reducing uncertainties and ~~further improving/advancing~~ our understanding of their composition.~~

## 5. Conclusions

This study focuses on the development of methodologies using HR-AMS for quantification and characterization of ~~ON/NOON~~ species in ambient aerosols and aqueous droplets. Through extensive analysis of HR-AMS spectra from 75 ~~ON/NOON~~ standards, we reach several important conclusions. Firstly, we show that ~~ON/NOON~~ compounds can generate significant  $\text{NH}_x^+$  and  $\text{NO}_x^+$  ion fragments, which are commonly assigned to inorganic nitrogen species in ambient AMS analysis. Additionally, although the presence of  $\text{CH}_2\text{N}^+$  ( $m/z = 28.0187$ ) is common in the AMS spectra of ~~ON/NOON~~ compounds, it is often overlooked due to interference from air-related  $\text{N}_2^+$  in ambient datasets. Secondly, the average RIE for ~~ON/NOON~~ was determined to be  $1.52 (\pm 0.58)$ , which aligns well with the default RIE value of 1.4 for total OA quantification. By considering these factors, we recommend a new N/C calibration factor of 0.79 for ~~ON/NOON~~ quantification.

Furthermore, we examined the ~~high-resolution mass spectra (HRMS)~~ of various ~~ON/NOON~~ compounds to identify chemical fingerprint ions that can aid in the ~~NOON~~ speciation of ~~ON-components~~ in ambient HR-AMS datasets. Our findings reveal distinct ion patterns associated with different ~~ON/NOON~~ types. Specifically,  $\text{C}_n\text{H}_{2n+2}\text{N}^+$  ions serve as tracer ions for amines (particularly  $\text{CH}_4\text{N}^+$ ,  $\text{C}_2\text{H}_6\text{N}^+$  and  $\text{C}_3\text{H}_8\text{N}^+$  are markers of primary, secondary and tertiary amines, respectively). The presence of  $\text{C}_n\text{H}_{2n}\text{NO}^+$  ions (especially for  $n=1-4$ ), is a strong signal for the presence of amides. Tracer ions for amino acids include  $\text{C}_2\text{H}_4\text{NO}_2^+$ ,  $\text{C}_2\text{H}_3\text{NO}^+$  and  $\text{CH}_4\text{NO}^+$ . The  $\text{C}_x\text{H}_y\text{N}_2^+$  ion series are characteristic of 2N-heterocyclic compounds. It is worth noting that  $\text{NH}_4^+$  ion can be produced from amines and amino acids, but not from inorganic nitrogen species. Additionally, we compared the HR-AMS and NIST database spectra for the ~~ON/NOON~~ species. The level of similarity varied significantly among different ~~ON-types~~. Pearson's correlation coefficients ( $r^2$ ) for amines, amides, N-heterocycles, nitro compounds and amino acids were found to be 0.70, 0.63, 0.57, 0.50 and 0.23, respectively.

Lastly, we applied the methodologies described above to analyze three HR-AMS datasets, which included online measurements of  $\text{PM}_1$  and offline analysis of fog ~~waterswater~~. The fog ~~waterswater~~ in Fresno exhibited significantly higher N/C ratios (average = 0.078) compared to the  $\text{PM}_1$  samples (0.019 in Fresno and 0.015 in NYC). Regarding ~~ON/NOON~~ constitution, both Fresno and NYC  $\text{PM}_1$  contained significant amounts of amino compounds, whereas Fresno fog ~~waterswater~~ exhibited a broader range of ~~ON/NOON~~ species, including N-containing aromatic heterocycles (e.g., imidazoles) and amides.

Moving forward, our aim is to extend this approach to multiple HR-AMS datasets collected from diverse locations worldwide. By doing so, we anticipate gaining a more comprehensive understanding of the chemical characteristics, sources, and processes of atmospheric ON. This broader application will contribute to advancing our knowledge regarding airborne organic nitrogen compounds and their role in atmospheric chemistry on a global scale.

**Code and data availability.** All datasets, including HR-AMS mass spectra, are available upon request from Qi Zhang ([dkwzhang@ucdavis.edu.cn](mailto:dkwzhang@ucdavis.edu.cn)) and Xinlei Ge ([caxinra@163.com](mailto:caxinra@163.com)).

**Supplement.** The supplementary material related to this article is available online.

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**Author contributions.** QZ conceived and designed the experiments, while XG and YS conducted the experiments. XG, YS, JT and QZ analyzed the HR-AMS spectra of the standards and the ambient data. XG and QZ wrote the paper with input from all authors.

430 **Competing interests.** The contact author has declared that none of the authors has any competing interests.

**Acknowledgements.** This work was funded by the U.S. Department of Energy Office of Science Atmospheric System Research Program (Grant #DE-SC0022140), the U.S. National Institute of Environmental Health Sciences Core Center (Grant P30 ES023513), and the Natural Science Foundation of China (91544220, 21976093, 42021004), and the Jiangsu Natural Science Foundation (BK20150042).

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## Tables

Table 1. Average values with  $\pm 1$  standard deviation for the % mass contributions (%) from major nitrogen-containing ion categories to the total organic nitrogen (ON) mass of different types of organic nitrogen standards. Note:  $M^+$  refers to the molecular ion of the compound, for example,  $M^+$  represents  $CH_5N^+$  for methylamine. Numbers in bold emphasize relatively large and significant values.

Categories	Amines <i>N</i> =27	Amides <i>N</i> =6	Amino Acids <i>N</i> =27	Nitro-compounds <i>N</i> =4	N-Heterocycles <i>N</i> =7	Protein <i>N</i> =1	Humic Acids <i>N</i> =3
$C_xH_yN_1^+$	<b>84.4±14.1</b>	48.1±20.8	<b>76.8±12.5</b>	60.8±13.8	69.8±19.3	75.5	<b>81.3±3.7</b>
$C_xH_yN_2^+$	2.9±7.7	11.0±11.5	1.9±5.3	3.7±5.5	<b>11.0±15.6</b>	3.7	1.1±0.2
$C_xH_yO_1N_1^+$	2.2±5.6	<b>25.6±11.9</b>	4.2±4.0	5.1±1.9	5.6±7.5	5.9	16.0±3.2
$C_xH_yO_2N_1^+$	0.26±0.97	0.64±1.44	3.0±4.4	1.5±0.9	2.8±3.5	0.3	1.2±0.6
$NO_x^+$	0.20±0.45	0.90±1.03	0.97±1.58	<b>16.2±8.3</b>	1.3±2.2	1.1	-
$NH_x^+$	<b>9.8±9.7</b>	<b>12.6±18.5</b>	<b>12.9±8.7</b>	<b>10.4±7.8</b>	<b>9.5±15.4</b>	<b>13.3</b>	-
$M^+$	<u>4.1±6.4</u>	<u>4.1±6.7</u>	<u>0.36±1.71</u>	<u>2.1±2.3</u>	<u>9.3±10.9</u>	=	=

## Figures

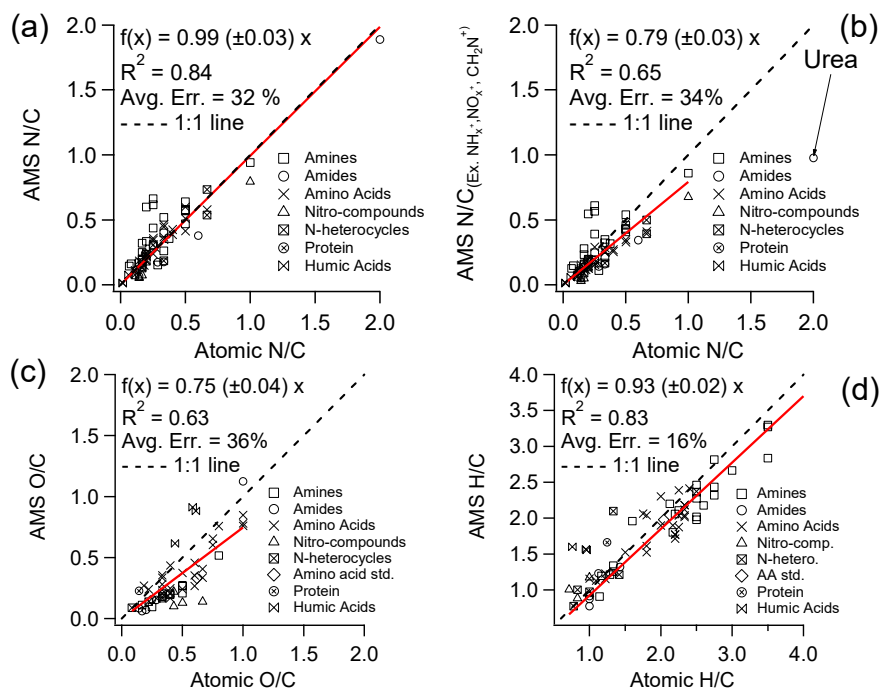


Figure 1. Atomic N/C ratios from the elemental analysis of 75 **ONNOON** standards versus the nominal values by (a) including and (b) excluding ion categories of  $\text{NH}_4^+$ ,  $\text{NO}_3^+$ , and  $\text{CH}_2\text{N}^+$ , and the atomic ratios of (c) O/C and (d) H/C from the same set of **ON** standards versus nominal values (the linear regressions are forced to zero. Urea. Table S1 provides details of all the standards used here. Note that urea is not included in the linear regression in (b); Table S1 details the standards used here.), and the linear regressions are orthogonal distance regression constrained to pass through the origin. The “average error” (Avg. Err.) represents the mean of the absolute relative deviations of all samples from the red fitted line in the corresponding figure.

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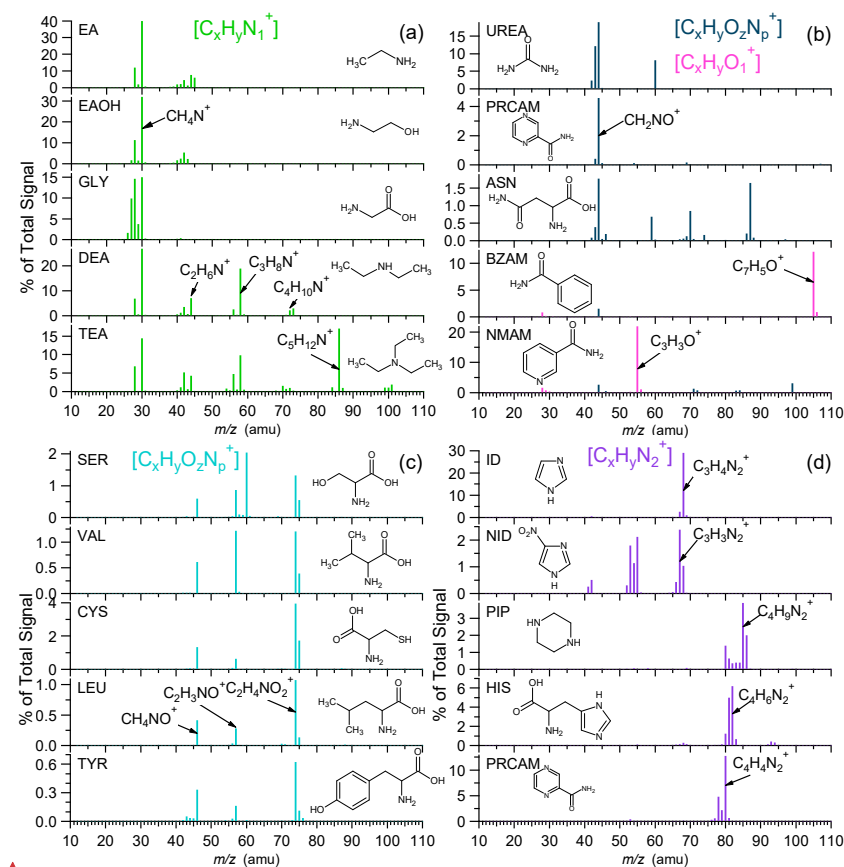
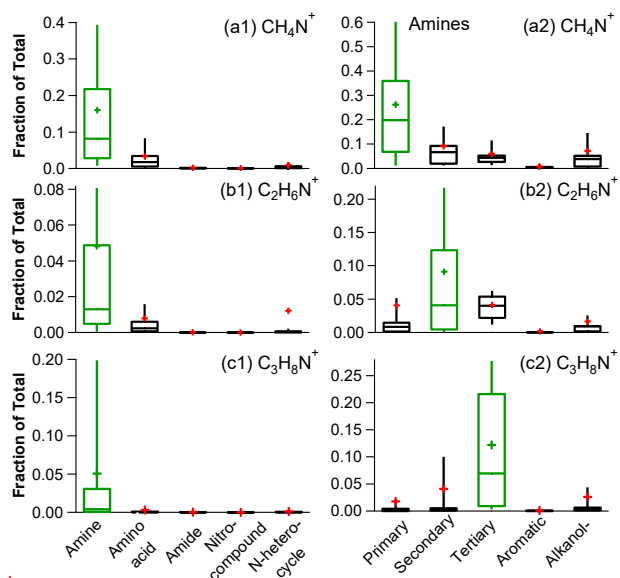


Figure 2. Characteristic spectral peaks of selected ONNOON standards with different functional groups of (a) amines (R-NH<sub>2</sub>, R<sub>2</sub>-NH, R<sub>3</sub>-N), (b) amides (R-CO-NH<sub>2</sub>), (c) amino acids (R-CH(NH<sub>2</sub>)-COOH), and (d) 2N-heterocycles. (note full HRMS of these compounds can be found in Figure S3). EA: Ethylamine; EAOH: Ethanolamine; GLY: Glycine; DEA: Diethylamine; TEA: Triethylamine; UREA: urea; PRCAM: Pyrazinecarboxamide; ASN: Asparagine; BZAM: Benzamide; NMAM: N,N'-Methylenebisacrylamide; SER: Serine; VAL: Valine; CYS: Cysteine; LEU: Leucine; TYR: Tyrosine; ID: Imidazole; NID: Nitroimidazole; PIP: Piperazine; HIS: Histidine.

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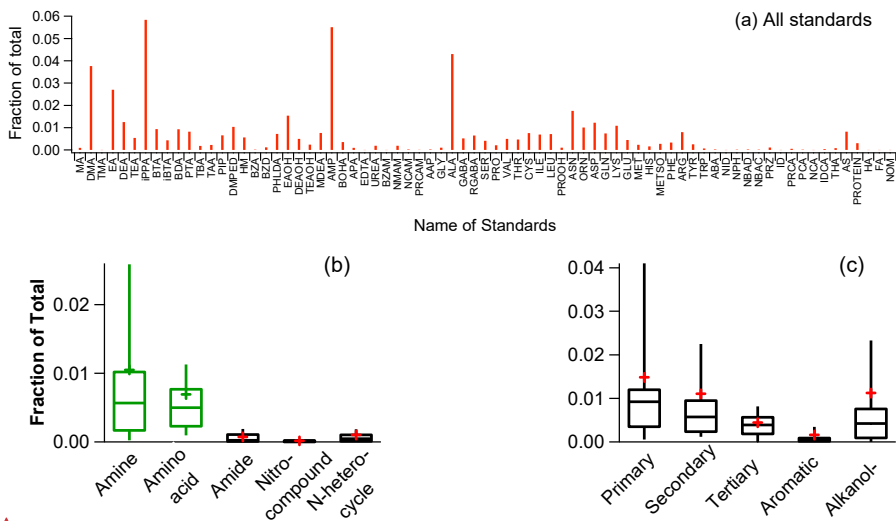
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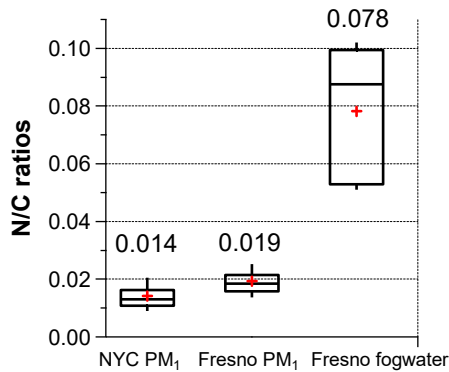


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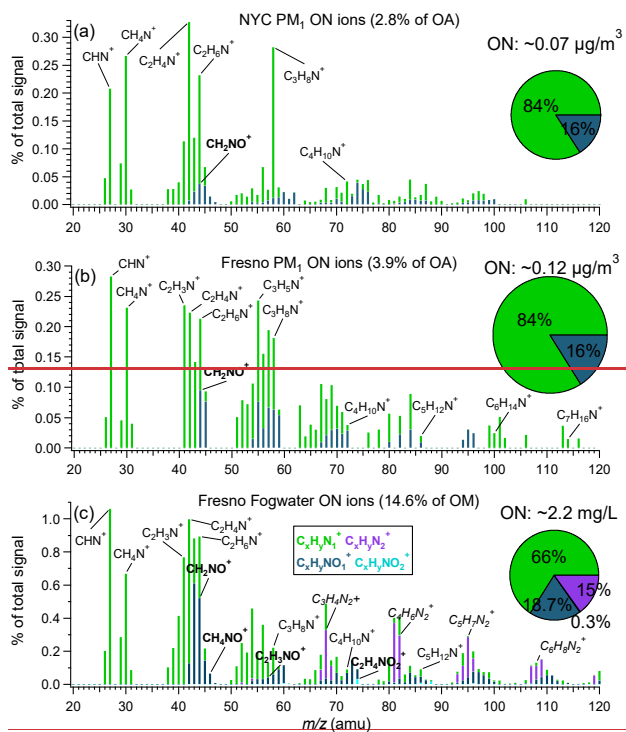
735 **Figure 3.** Fractional contributions of CH<sub>4</sub>N<sup>+</sup> (a), C<sub>2</sub>H<sub>6</sub>N<sup>+</sup> (b) and C<sub>3</sub>H<sub>8</sub>N<sup>+</sup> (c) to the mass spectra of the N-containing organic species with different functional groups (left panel) and different classes of amines (right panel) (the whiskers above and below the boxes indicate the 90<sup>th</sup> and 10<sup>th</sup> percentiles, the upper and lower boundaries of the boxes indicate the 75<sup>th</sup> and 25<sup>th</sup> percentiles, and the lines in the boxes indicate the median values and the cross symbols indicate the mean values).



740 Figure 4. Mass fractional contributions of  $\text{NH}_4^+$  in the HR-AMS spectra for all standards (a), different **FUNCTION** classes (b) and different amines (c).



745 Figure 5. N/C ratios of the organics in NYC  $\text{PM}_1$ , Fresno  $\text{PM}_1$  and fog **waterswater** (the whiskers above and below the boxes indicate the 90<sup>th</sup> and 10<sup>th</sup> percentiles, the upper and lower boundaries of the boxes indicate the 75<sup>th</sup> and 25<sup>th</sup> percentiles, and the lines in the boxes indicate the median values and the cross symbols indicate the mean values).



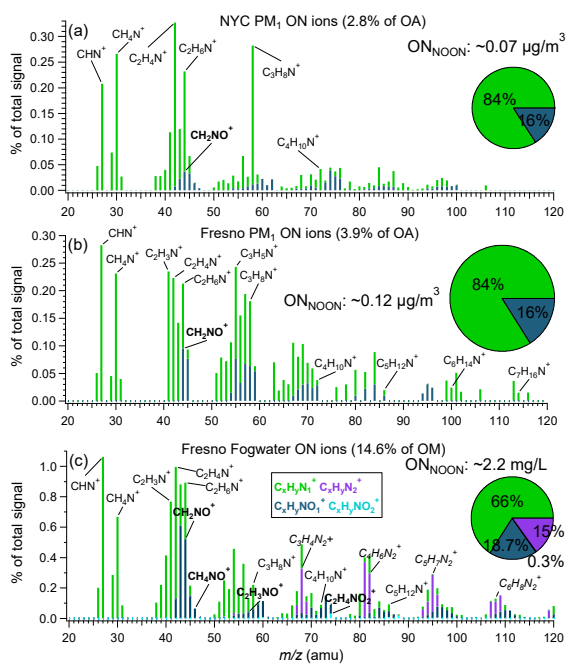


Figure 6. Average mass spectra of ON ion categories (C<sub>3</sub>H<sub>7</sub>N<sup>+</sup>, C<sub>3</sub>H<sub>9</sub>N<sup>+</sup>, C<sub>3</sub>H<sub>5</sub>N<sub>2</sub><sup>+</sup>, C<sub>3</sub>H<sub>7</sub>NO<sup>+</sup>, and C<sub>3</sub>H<sub>9</sub>NO<sub>2</sub><sup>+</sup>) for NYC PM<sub>1</sub> (a), Fresno PM<sub>1</sub> (b) and Fresno fog waterswater (c). The inset pies show the average  $\Theta_{\text{NONNOON}}$  concentrations and mass contributions of the four ion categories (legends in c) to the total  $\Theta_{\text{NONNOON}}$ , respectively.

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