



### Identifying and correcting interferences to PTR-ToF-1

### MS measurements of isoprene and other urban volatile 2

### organic compounds 3

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35 Abstract: Proton-transfer-reaction time-of-flight mass spectrometry (PTR-ToF-MS) is a technique commonly used to measure ambient volatile organic compounds (VOCs) in urban, rural, 36 37 and remote environments. PTR-ToF-MS is known to produce artifacts from ion fragmentation, 38 which complicates the interpretation and quantification of key atmospheric VOCs. This study 39 evaluates the extent to which fragmentation and other ionization processes impacts urban 40 measurements of the PTR-ToF-MS ions typically assigned to isoprene (m/z 69,  $C_5H_8H^+$ ), 41 acetaldehyde (m/z 45, CH<sub>3</sub>CHO<sup>+</sup>), and benzene (m/z 79, C<sub>6</sub>H<sub>6</sub>H<sup>+</sup>). Interferences from 42 fragmentation are identified using gas-chromatography (GC) pre-separation and the impact of 43 these interferences are quantified using ground-based and airborne measurements in a number of

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44 US cities, including Las Vegas, Los Angeles, New York City, and Detroit. In urban regions with 45 low biogenic isoprene emissions (e.g., Las Vegas), fragmentation from higher carbon aldehydes 46 and cycloalkanes emitted from anthropogenic sources may contribute to m/z 69 by as much as 50% during the day, while the majority of the signal at m/z 69 is attributed to fragmentation during 47 the night. Interferences are a higher fraction of m/z 69 during airborne studies, which likely results 48 49 from differences in the reactivity between isoprene and the interfering species along with the subsequent changes to the VOC mixture at higher altitudes. For other PTR masses, including m/z 50 51 45 and m/z 79, interferences are observed due to the fragmentation and secondary ionization of 52 VOCs typically used in solvents, which are becoming a more important source of anthropogenic 53 VOCs in urban areas. We present methods to correct these interferences, which provide better 54 agreement with GC measurements of isomer specific molecules. These observations show the 55 utility of deploying GC pre-separation for the interpretation PTR-ToF-MS spectra. 56

### 57 1. Introduction

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Volatile organic compounds (VOCs) are an important contributor to urban air pollution. Once
emitted to the atmosphere, VOCs undergo chemical reactions that contribute to the formation of
hazardous pollutants such as ozone and secondary organic aerosol. It is important to quantify VOC
mixing ratios in urban areas to determine strategies that may reduce air pollution.

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64 Proton-transfer-reaction time-of-flight mass spectrometry (PTR-ToF-MS) is a technique used to measure a wide spectrum of VOCs, including oxygenates, aromatics, furanoids, nitriles, and 65 biogenic species such as isoprene and monoterpene isomers (Yuan et al., 2017). PTR-ToF-MS 66 67 measurements in urban regions enable the determination of VOC mixing ratios from an extensive range of emission sources, including fossil fuels, solvent evaporation from volatile chemical 68 69 products (VCPs), residential wood burning, cooking, and urban foliage (Yuan et al., 2017). The 70 fast-time resolution and broad selectivity of PTR-ToF-MS measurements enables source 71 apportionment, flux estimates, and spatial mapping on mobile platforms that yield important information about urban VOC source strengths (e.g., Gkatzelis et al., 2021a;Karl et al., 72 73 2018;Pfannerstill et al., 2023)

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76 77 78 VOC detection by PTR-ToF-MS relies on analyte reactions with protonated water (Reaction 1).

$$VOC + H_3O^+ \rightarrow VOC \cdot H^+ + H_2O \tag{R1}$$

79 Proton transfer is exothermic and spontaneous for VOCs with a proton affinity that is greater than 80 water. For many VOCs, including ketones, aromatics, and nitriles, the protonated product 81  $(VOC \cdot H^+)$  is the primary signal detected by PTR-ToF-MS. For other VOCs, secondary reactions 82 including dehydration, fragmentation, and water clustering results in additional product ions that 83 can complicate the mass spectra. Pagonis et al. (2019) summarizes studies that have reported fragmentation for a wide spectrum of species. Fragmentation is most prevalent in alcohols, 84 85 aldehydes, and other species with long-chain alkane functionality. Small alcohols and aldehydes 86 (C < 3) primarily react to form protonated products following R1, while at higher carbon numbers, 87 a large fraction of the reactions follow dehydration and/or fragmentation (R2-R3). 88

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 $(R - 0H, RH = 0) \cdot H^+ \rightarrow R^+ + H_2 0$  (R2)





90  $R^+ \rightarrow R_1^+ + R_2$ 91 (R3) 92 93 R is the carbon backbone of an alcohol (R-OH) or aldehyde (RH=O), R<sup>+</sup> is the dehydration product, R<sub>1</sub><sup>+</sup> is a fragment, and R<sub>2</sub> is a neutral product. Fragmentation may also result from protonation of 94 95 cycloalkanes or alkyl aromatics. PTR-ToF-MS is not sensitive to small alkanes (C < 5), but larger 96 alkanes and cycloalkanes are detected at low sensitivity and upon ionization subsequently 97 fragment to produce ions that often overlap with the dehydration and fragmentation products of 98 alcohols and aldehydes (Arnold et al., 1998; Gueneron et al., 2015; Jobson et al., 2005). The degree 99 of dehydration and fragmentation is partially dependent on the strength of the drift field (E) and 100 density (N) (characterized by the E/N ratio), which impacts ion kinetic energy (Arnold et al., 101 1998;Krechmer et al., 2018;Yuan et al., 2017;Holzinger et al., 2019). Lower E/N results in lower 102 fragmentation, but higher clustering with neutral water, which may further complicate the mass 103 spectra (Holzinger et al., 2019). Additional products may also be formed by reactions of analytes 104 with  $O_2^+$  and  $NO^+$  ions, which are present due to the ionization of small amounts of air in the 105 discharge region.

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107 In the atmosphere, complex mixtures of emissions may result in PTR-ToF-MS mass spectra where 108 dehydration and fragmentation products interfere with the quantification of important atmospheric 109 VOCs. For example, PTR-ToF-MS measurements in regions with significant oil and natural gas 110 development show that substituted cycloalkanes fragment to produce significant signal at m/z 69 111 (C5H9<sup>+</sup>, Koss et al., 2017; Warneke et al., 2014; Pfannerstill et al., 2019). Likewise, interferences 112 have been observed downwind of urban and industrial environments (e.g., Inomata et al., 2010; Choi et al., 2022). These fragments overlap with protonated isoprene and these previous 113 114 studies have shown that interferences make isoprene quantification challenging in these regions. 115 In forested areas, isoprene is largely emitted from biogenic sources and previous studies that have 116 compared PTR-ToF-MS measurements to those from gas-chromatography show good agreement 117 (e.g., Kaser et al., 2013). The impact of an interference to specific molecules, such as isoprene, 118 depends on atmospheric composition which changes spatially (e.g., urban vs. rural regions) and 119 temporally (e.g., summer vs. winter).

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121 Assessments of interferences on PTR-MS measurements in urban atmospheres have been 122 conducted previously (e.g., Warneke et al., 2003), but the sources that contribute to urban VOCs change on decadal timescales as fossil fuel emissions steadily decline (Kim et al., 2022; Warneke 123 124 et al., 2012). The urban atmospheric composition, once dominated by motor vehicle emissions, is 125 now composed of a higher proportion of oxygenates from solvents, other VOCs emitted from 126 sources such as VCPs, and cooking (Gkatzelis et al., 2021b;McDonald et al., 2018;Peng et al., 127 2022; Wernis et al., 2022). Significant advances in PTR-ToF-MS detectors (quadrupole vs time-128 of-flight mass spectrometers) and drift tube designs have enhanced instrument capabilities to 129 acquire mass spectra with greater resolution and sensitivity (Breitenlechner et al., 2017;Krechmer 130 et al., 2018; Yuan et al., 2016; Holzinger et al., 2019). These technological advances enable better 131 identification and quantification and also an improved understanding of the interferences that 132 impact PTR-ToF-MS spectra.

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With the changes in atmospheric composition and technological advances, it is necessary to revisit
 potential interferences to commonly observed and reported VOCs by PTR-ToF-MS. In this study,





136 we investigate the interferences that impact PTR-ToF-MS spectra measured across several US 137 urban areas including Los Angeles, CA, Las Vegas, NV, Detroit, MI, and New York City, NY. 138 Interferences are identified using GC pre-separation, similar to previous measurements that 139 quantified PTR-ToF-MS fragmentation and interferences observed in complex mixtures, including 140 wildfire and urban emissions (e.g., Koss et al., 2018;Warneke et al., 2003). We show that commonly measured species, such as acetaldehyde, benzene, and isoprene, exhibit interferences 141 142 from larger molecules associated with solvent use and cooking. The extent of these interferences 143 depends on the temporal and spatial variability of VOC emission sources. We present methods to 144 correct for these interferences based on the measurement capabilities of modern PTR-ToF-MS 145 instruments.

#### 147 2. Methods

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149 Table 1 summarizes the key field campaigns and instrumentation used to quantify VOCs and PTR-150 ToF-MS interferences. Multiple PTR-ToF-MS instruments are used in this study and Table 1 outlines the PTR-ToF-MS reactor designs and drift tube operating parameters that play an 151 152 important role in determining ion distributions. In this study, all instruments were operated with 153 E/N 120 - 140 Td. The instruments described in this study use ion-molecular reactors and ion 154 optics devolved by Ionicon Analytik and Tofwerk, AG as described by Müller et al. (2014) and Krechmer et al. (2018), respectively. The following sections describe each campaign and provide 155 additional details of instrument operation 156

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158 Table 1: Summary of campaigns, instrumentation, drift tube operating parameters, and interferences reported for each 159 campaign.

Campaign	Location	Dates	Instrumentation	Reactor Design <sup>1</sup>	E/N (Td)	Studied Interferences
SUNVEx	Las Vegas, NV	July 1 – 30, 2021	NOAA PTR-ToF-MS <sup>2</sup> NOAA GC-PTR-ToF-MS NOAA iWAS	Vocus	140	Isoprene, aromatics, oxygenates
RECAP-CA	Los Angeles, CA; Central Valley, CA	June 1 – Aug 30, 2021	NOAA PTR-ToF-MS <sup>2</sup> UC Berkeley PTR-ToF-MS <sup>2</sup> NOAA GC-MS	Vocus (NOAA) Vocus (Berkeley)	140 130	Isoprene, aromatics, oxygenates
FIREX-AQ	Los Angeles, CA	September 5, 2019	NOAA PTR-ToF-MS <sup>2</sup> Oslo PTR-ToF-MS <sup>3</sup> NOAA iWAS, UCI Irvine WAS NCAR TOGA-TOF	HC / DT (NOAA) HC/ DT (Oslo)	125 120	Isoprene
MOOSE	Detroit, MI	May 21 – June 30, 2021	Aerodyne PTR-ToF-MS <sup>2</sup> Aerodyne GC-EI-ToF-MS	Vocus	125	Isoprene, aromatics
LISTOS	New York City, NY	January 2020 – April 2021.	Stony Brook PTR-ToF-MS <sup>3</sup>	HC / DT	130	Isoprene

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<sup>1</sup> HC / DT refers to the hollow cathode / drift tube design used in traditional PTR-MS instruments. This reactor is used in this study with both Tofwerk and Ionicon systems. The Vocus reactor is used with Tofwerk instruments.

162 163 <sup>2</sup> Tofwerk design PTR-ToF-MS using quadrupole ion optics as described by Krechmer et al. (2018)

164 <sup>3</sup> Ionicon design PTR-ToF-MS with ion optics consisting of two einzel lens as described by Müller et al. (2014)

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2.1. Campaign Descriptions 166

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168 2.1.1. SUNVEx / RECAP-CA

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170 PTR-ToF-MS measurements were performed as part of the 2021 Southwest NOx and VOC

171 Experiment (SUNVEx, https://csl.noaa.gov/projects/sunvex/) and Re-Evaluating the Chemistry of





- 172 Air Pollutants in California (RECAP-CA). SUNVEx was a ground campaign conducted to study
- air quality in the Las Vegas Valley during the summer ozone season using both mobile and ground based sampling. The RECAP-CA campaign was conducted in Los Angeles and included mobile,
- 174 based sampling. The RECRIPTION campaign was conducted 175 ground-based, and airborne sampling.
- 176
- 177 Measurements in Las Vegas were conducted between 30 June-27 July 2021 at an air quality
- monitoring station located near the Jerome Mack Middle School (Fig. 1). Jerome Mack is an urban
  background site located ~ 8km east of downtown Las Vegas. The site was chosen based on its
- suite of trace gas and PM<sub>2.5</sub> monitors and classification as a US Environmental Protection Agency
- 181 (EPA) Photochemical Assessment Monitoring Station (Annual Monitoring Network Plan, 2022).

Measurements in the Los Angeles Basin were conducted between 2 August and 5 September 2021 at the California Institute of Technology in Pasadena, CA (Caltech, Fig. 1). The ground site was located within 0.5 km of the site used during the California Research at the Nexus of Air Quality and Climate Change (CalNex) field study in order to directly compare with air quality measurements conducted in 2010 (Ryerson et al., 2013). During this portion of the campaign, instruments were situated in a trailer and sampled air from the top of a 10-m tower.

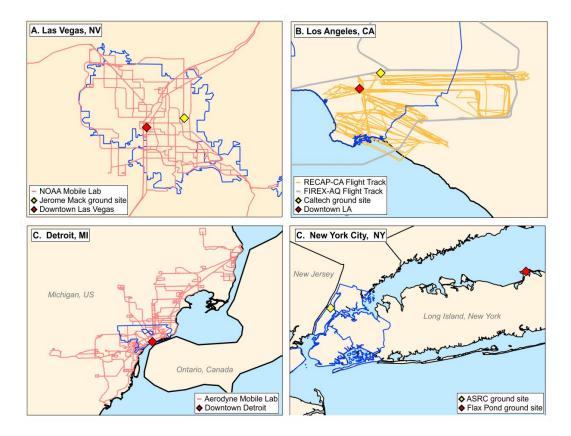
During both SUNVEx and RECAP-CA, the NOAA mobile laboratory was deployed to sample the spatial distribution of VOCs and NO<sub>x</sub> in regions of varying population density. A similar sampling strategy was employed previously to study urban VOC enhancements in New York City and is useful for identifying VOC signatures emitted from major sources, such as fossil fuels, VCPs, and

useful for identifying voc signatures enfitted from high sources, such as fossil fuels, ver s, and
 cooking activities (Coggon et al., 2021;Gkatzelis et al., 2021a;Gkatzelis et al., 2021b;Stockwell et

- al., 2021). Drive tracks from the mobile laboratory are shown on the map in Fig. 1, along with the
- 194 locations of the ground sites and major population centers.







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Figure 1. Mobile laboratory drive tracks, flight tracks, ground site locations, and locations of interest for the field
 campaigns outlined in Table 1. The blue lines highlight the statistical metropolitan areas for Las Vegas, Los Angeles,
 and Detroit, and the five boroughs of New York City.

The airborne component of RECAP-CA was conducted onboard the Naval Postgraduate School UV-18A Twin Otter aircraft and was based out of an airport located in Burbank, CA. Measurements of VOCs, NO<sub>x</sub>, and greenhouse gases took place on nine days between 1 June and 23 June 2021. The Twin Otter typically flew at ~ 300 m above ground level at air speeds of  $50 - 60 \text{ m s}^{-1}$ . Each flight covered approximately 5000 km of distance across the Los Angeles area, including downtown, the coast, the Santa Ana area, and the San Bernardino Valley. The flight track of the Twin Otter is shown in Fig. 1b.

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### 207 2.1.2. FIREX-AQ

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The 2019 Fire Influence on Regional to Global Environments and Air Quality (FIREX-AQ) campaign was a large field study designed to investigate the emissions and atmospheric chemistry of biomass burning emissions. A detailed description of the campaign, instrumentation, and science goals is provided by Warneke et al. (2023). As part of the measurements, urban flights were performed through the Los Angeles Basin on 5 September 2019. VOC measurements were conducted onboard the NASA DC-8 by the NOAA PTR-ToF-MS and University of Oslo PTR-





- ToF-MS. VOC measurements were also conducted by three GC instruments: the NOAA improved Whole Air Sampler (iWAS), the University of California, Irvine Whole Air Sampler (WAS) and
- 217 the NCAR Trace Organic Gas Analyzer with a Time-of-Flight mass spectrometer (TOGA-TOF).
- 218 The flight tracks conducted in Los Angeles are shown in Fig. 1b.
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# 220 2.1.3. MOOSE Campaign221

222 The Michigan-Ontario Ozone Source Experiment (MOOSE) campaign was a multi-institutional 223 ground-based and mobile sampling effort conducted in 2021 to study ozone, meteorology, and 224 pollution in and around Michigan and Ontario. This region is currently designated as non-225 attainment of the US federal ozone standard. Aerodyne Research, Inc. scientists deployed the 226 Aerodyne Mobile Laboratory (Herndon et al., 2005; Yacovitch et al., 2015) as part of the 227 CHEmical Source Signatures (CHESS) sub-experiment in order to measure emission plumes from 228 point sources and gain insight to the drivers of local ozone pollution during MOOSE. Other goals 229 of CHESS included developing emission source fingerprints for significant industrial source sites 230 in the area.

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232 Ambient VOC measurements were conducted onboard the mobile laboratory using a PTR-TOF-MS which was co-located with an *in situ* GC-EI-TOF-MS equipped with thermal desorption 233 preconcentration (Claflin et al., 2021). The AML sampled air around the Detroit metropolitan 234 235 region between 21 May and 30 June 2021 (Fig 1c). During mobile sampling, the mobile laboratory 236 transited through major population centers and targeted industrial point sources. Overnight and 237 when not driving, the mobile laboratory was stationed at the Salina Elementary/Intermediate 238 Schools in Dearborn, MI, parked at the Michigan Department of Environment, Great Lakes, and 239 Energy air monitoring station [AQS ID 26-163-0033]. 240

## 241 **2.1.4.** LISTOS

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243 The Stony Brook PTR-ToF-MS was deployed on the rooftop observatory at the Advanced 244 Sciences Research Center (ASRC) of the City University of New York to make continuous, high 245 time-resolution measurements of VOCs during the COVID lockdown from January 2020 to April 246 2021, (Fig. 1c; Cao et al., 2023). This campaign was a part of the broader Long Island Sound 247 Tropospheric Ozone Study (LISTOS). ASRC is located in the Manhattan Borough of New York 248 City, which is a highly urbanized region. Air was continuously sampled from a rooftop observatory 249 that is situated ~90 m above sea level on one of the tallest buildings in the vicinity of the site. In 250 June 2022, the Stonybrook PTR-ToF-MS was moved to the Flax Pond Marine Laboratory 251 (40°57'36"/N, 73°8'24" W) near Stony Brook, New York, which is about 60 miles east of ASRC 252 and located on the north side of Long Island in a forested suburban area. The Flax Pond Marine 253 Laboratory is a 0.6 km<sup>2</sup> preserve that encompasses a tidal wetland area and is operated for research purposes by the School of Marine and Atmospheric Sciences of Stony Brook University. At Flax 254 255 Pond, air was continuously sampled from a ~10 m tower.

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- 257 2.2.Instrument Descriptions PTR-ToF-MS
- 258 259 **2.2.1. NOAA PTR-ToF-MS**
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The NOAA PTR-ToF-MS was deployed during SUNVEx, RECAP-CA, and FIREX-AQ. During
FIREX-AQ, the NOAA PTR-ToF-MS used a traditional ion source and drift tube as described by
Yuan et al. (2016). A full description of the operating parameters, VOC measurements, and
calibration methods are provided by Gkatzelis et al. (2022).

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During SUNVEx and RECAP-CA, the instrument was modified to use the Vocus focusing ion 266 267 molecule reactor (TOFWERK, AG) and was operated following the recommendations by Krechmer et al. (2018). The Vocus provides greater sensitivity to VOCs compared to the 268 269 traditional drift tube design due to the use of quadrupole ion guides that increase ion transmission. 270 Here, the Vocus was operated at 2.5 mbar and with an axial electric field gradient of 65 V cm<sup>-1</sup> 271  $(E/N \sim 140 \text{ Td})$ . The water flow to the ion source was maintained at 23 mL min<sup>-1</sup> and the drift tube 272 was heated to 110°C. Typically, the quadrupole ion guide in the Vocus PTR-ToF-MS is operated 273 at voltages > 275 V to reduce the transmission of reagent ions that would otherwise limit the 274 lifetime of the detectors (Krechmer et al., 2018). Here, the quadrupole ion guide was tuned to 250 275 V to increase the transmission of ions produced from important VOCs with low molecular weights, 276 such as ethanol (m/z 47), acetonitrile (m/z 42), and methanol (m/z 33). Figure S1 compares the 277 product distribution of VOCs measured by the Vocus against those measured with the traditional 278 drift tube. In general, the ion product distributions are comparable, though small differences in 279 water clusters and fragmentation in the Vocus reflect the higher amount of water in the drift tube 280 and a higher operating E/N. The degree of fragmentation in the NOAA Vocus PTR-ToF-MS is 281 comparable to other PTR-MS systems with E/N > 120 Td (e.g., Buhr et al., 2002;Pagonis et al., 282 2019). Other Vocus PTR-ToF-MS instruments used in this study observed higher fragmentation 283 owing to differences in operating conditions of the big-segmented quadrupole (BSQ). The 284 implications of fragmentation from the BSQ are discussed further in Section 3.1.

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286 To guide the identification of the proton-transfer-reaction products, a GC was used to trap and pre-287 separate ambient VOCs during SUNVEx. The GC deployed here is the same instrument used by 288 Stockwell et al. (2021) to identify molecular isomers measured from coating headspaces. Briefly, 289 the GC consists of a liquid nitrogen cryotrap coupled to a DB-624 column (Restek MXT-624; 30-290 m length  $\times$  0.25-mm inner diameter (I.D.), 1.4-µm film thickness). Samples were collected onto 291 the cryotrap at predetermined volumes (typically 80 cm<sup>3</sup>), then injected onto the column via rapid 292 heating to 100°C. Nitrogen gas carried the sample through the column at 8 sccm while the column 293 was heated from 40°C to 150°C at a rate of 12°C min<sup>-1</sup>. The effluent from the column was injected 294 into the PTR-ToF-MS inlet. In this study, we use this setup (termed GC-PTR-ToF-MS) to 295 qualitatively assess isomer distributions and fragmentation patterns for VOCs detected during in 296 situ sampling.

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The GC-PTR-ToF-MS was primarily deployed during the ground-based sampling phase in Las Vegas, while PTR-ToF-MS only was used in Los Angeles. In GC-mode, samples were collected every 2 hours and automatically analyzed by PTR-ToF-MS. In between GC measurements, the PTR-ToF-MS sampled ambient air through a 10-m Teflon inlet at 2 L min<sup>-1</sup>. During ambient sampling, instrument backgrounds were determined hourly by passing ambient air through platinum catalyst heated to 350°C. The PTR-ToF-MS was calibrated using gravimetricallyprepared gas standards, or by liquid calibration (Coggon et al., 2018).





When installed on the mobile laboratory, the PTR-ToF-MS sampled air through a 1-m Teflon inlet at 2 L min<sup>-1</sup>, and instrument backgrounds were determined every 15 minutes. During an evening drive on 31 July 2021, the GC-PTR-ToF-MS was deployed to speciate VOCs on the Las Vegas
Strip, where large crowds of people were present and anthropogenic emissions from personal care products, cooking, and other human activities were expected to be highest.

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## 312 2.2.2. University of Oslo PTR-ToF-MS

The University of Oslo PTR-ToF-MS was deployed during FIREX-AQ to target NH<sub>3</sub>, but also measured the same VOCs as the NOAA PTR-ToF-MS. The instrument was operated as described by Müller et al. (2014) with modifications to reduce the formation of NH<sub>4</sub><sup>+</sup> in the ion source as described by Tomsche et al. (2023). Briefly, the instrument sampled air through a heated inlet at a flowrate of 10–60 L min<sup>-1</sup> in order to reduce losses of NH<sub>3</sub> to inlet surfaces. The drift tube was operated at 2.1 mbar and 120°C with corresponding E/N ratio of 120 Td. VOC sensitivities were determined via calibrated using gravimetrically-prepared standards.

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### 322 2.2.3. University of California Berkeley PTR-ToF-MS

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The Berkeley Vocus PTR-ToF-MS (Aerodyne Research, Inc., Billerica, USA) was deployed 324 325 during the RECAP-CA aircraft campaign on the US Navy Twin Otter. The PTR-ToF-MS was 326 operated with a Vocus reactor set to 60°C, 2.0 mbar, and an E/N ratio of ~ 130 Td. The potential gradient along the drift tube was 590 V. The gradient between BSQ skimmer 1 and skimmer 2 was 327 328 changed once during the campaign from 6 to 9.1 V, which resulted in an improved sensitivity for 329 some VOCs, but significantly stronger fragmentation for other compounds such as nonanal (Fig. 330 S2). Both operating conditions were calibrated. The reagent water flow was 20 mL min<sup>-1</sup>. Similar 331 to the NOAA-PTR-ToF-MS, the voltage of the quadrupole ion guide was operated at 200 V to improve the transmission for low-mass VOCs like methanol. 332

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334 Ambient air was sampled via a 90 cm long heated (40°C) <sup>1</sup>/<sub>4</sub>-inch Teflon line through a Teflon 335 filter from an isokinetic inlet (flow rate ~ 6 m s<sup>-1</sup> for 5 m length) with a mass flow controller at 1.5 L min<sup>-1</sup>. Mass spectra were recorded at 10 Hz time resolution for a mass range of 10-500 Da. Zero-336 337 air blank measurements were conducted several times in each flight for 1-5 minutes during aircraft 338 turns ~ 2-4 times per flight, each followed by a pulse of calibration gas ~ 1-5 minutes in duration. 339 These in-flight calibrations were used to validate the sensitivities calculated from ground 340 calibrations. Ground calibrations were conducted every 1-3 days (in total, 19 times) during the 341 campaign using one of three gravimetrically prepared multicomponent VOC standards (Apel-342 Riemer Environmental Inc., Miami, FL, USA). More details on the instrument operation and 343 calibration can be found in Pfannerstill et al. (2023).

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### 345 2.2.4. Aerodyne PTR-ToF-MS

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Aerodyne Research, Inc. deployed a Vocus PTR-ToF-MS during MOOSE 2021 (Krechmer et al., 2018;Riva et al., 2019). The Vocus was operated at a pressure of 2.2 mbar and axial voltage
gradient of 600 V, corresponding to an E/N ratio of 125 Td. Data were recorded and processed at
1 Hz time resolution using the Tofware software (Aerodyne Research Inc. and TOFWERK) in
Igor Pro (WaveMetrics) (Stark et al., 2015). Background measurements were conducted every 16





minutes by overflowing the Vocus inlet with air from a zero-air generator (ZAG) equipped with a Pt/Pd catalyst at 400°C. Calibrations were performed every 4 hours with a multicomponent VOC mixture (Apel-Riemer Environmental Inc., Miami, FL, USA; nominal 1 ppm in N<sub>2</sub>) diluted with ZAG air. The sensitivities of species in the calibration mixture were correlated to their protoncapture-rate coefficient (Sekimoto et al., 2017). To calculate sensitivities for compounds not present in the calibration mixture, the slope of the linear fit was multiplied by the proton-capturerate coefficient of the species of interest (Holzinger et al., 2019;Krechmer et al., 2018).

- 360 2.2.5. Stony Brook PTR-ToF-MS
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362 The Stony Brook Ionicon high-resolution PTR-ToF-MS (Ionicon 8000, Analytik GmbH, Austria) 363 was deployed in New York City during the COVID shutdown, and subsequently at Flax Pond on 364 Long Island. In this study, the Stony Brook PTR-ToF-MS was operated with a drift field of ~130 365 Td, drift voltage of 600 V, reactor temperature of 60°C, and drift tube pressure of 2.3 mbar. The 366 instrument sampled VOCs through a heated (60°C) 1/16-inch outer diameter (O.D.) capillary 367 PEEK inlet (~1 m length) with bypass flow line teed off the 1/2-inch PTFE inlet line fitted with a 368 blower on the back end (residence time of the gas was  $\sim 10$  s). Data were collected at 1 Hz and 369 integrated to 5-minute averages.

Calibrations were performed using a dynamic dilution system. VOC-free air was produced by
pumping ambient air through a Pt-based catalytic converter at 400°C, then mixed with
multicomponent gas calibration mixture (Apel-Riemer Environmental Inc., Miami, FL, USA) that
included isoprene. Calibration was performed spanning a concentration range of observed values
(0, 5, 10, 15, 20 ppbv). At ASRC, the calibration gas was typically analyzed twice a week prior to
the COVID-19 lockdown, and typically every 1-2 weeks during the lockdown given limited access
to the observatory. At Flax Pond, the calibration gas was analyzed once a week.

- 377
- 378 2.3.Instrument Descriptions Gas Chromatography
- 379

380 2.3.1. NOAA GC-MS

381

The NOAA GC-MS provided speciated VOC measurements during SUNVEx, RECAP-CA, and 382 383 FIREX-AQ. During RECAP-CA, the NOAA GC-MS was deployed to the ground site to sample 384 ambient air on a 20 minute duty cycle. The GC-MS collects two separate 240 mL samples and 385 analyzes each on two channels. Channel 1 consists of a CO<sub>2</sub> trap (Ascarite II, Thomas Scientific), 386 a water trap operated at -55°C, and a sample trap operated at -165°C. The series of traps are linked to an Al<sub>2</sub>O<sub>3</sub>-KCl porous layer open tubular column (Restek RT-Alumina bonded porous 387 388 polymer/KCl; 30-m length  $\times$  0.25-mm I.D., and 4-µm film thickness) designed to separate light hydrocarbons. Channel 2 consists of the water and sample traps, but is coupled to a DB-624 column 389 390 identical to the column used in the GC-PTR-ToF-MS. This column separates hydrocarbons up to 391  $C_{12}$ , as well as select oxygen-, halogen-, and nitrogen-containing VOCs. The effluent of each 392 column is analyzed using a quadrupole mass spectrometer (Agilent 5975C) operated in selected 393 ion monitoring/scan mode. The GC-MS was calibrated using a gravimetrically-prepared gas 394 mixture containing 50 VOC components.

395

VOCs analyzed by NOAA GC-MS in Las Vegas and during FIREX-AQ were first sampled using
 a whole air sampling canister system (iWAS, Lerner et al., 2017). The iWAS system consists of a





398 stainless steel compressor and 24 2.7 L electropolished stainless steel canisters. During SUNVEx, 399 canisters were filled every 2 hours during stationary ground-based sampling, while targeted 400 samples were taken during mobile drives. During FIREX-AQ, targeted samples were filled on 401 demand up to a total number of 72 per flight. The canisters were shipped to either Boulder, CO, or 402 Pasadena, CA, and analyzed by GC-MS within four days of collection to minimize sampling 403 artifacts.

404 405

### 406 2.3.2. UCI WAS

407

408 The University of California, Irvine Whole Air Sampler (UCI WAS) was deployed on the DC-8 409 during FIREX-AQ to sample VOCs and halocarbons. The UCI WAS operated similarly to the 410 NOAA canister system, where samples were collected into electropolished stainless steel canisters, 411 then analyzed offline within seven days using a series of laboratory GC systems. The operation of 412 the airborne WAS and laboratory GCs is fully described elsewhere (Colman et al., 2001;Simpson 413 et al., 2020; Simpson et al., 2010). All samples were analyzed on a multi-column GC system 414 coupled to flame ionization, electron capture, and mass selective detectors. The whole system is 415 calibrated using a suite of VOC standards.

416

### 417 **2.3.3. TOGA-TOF**

418

419 The NCAR Total Organic Gas Analyzer with a TOFWERK electron ionization high-resolution 420 time-of-flight mass spectrometer (TOGA-TOF) was deployed on the DC-8 during FIREX-AQ to 421 provide in situ GC measurements of a large suite of VOCs including hydrocarbons, oxygenated 422 VOCs (OVOCs), and halogen-, nitrogen- and sulfur-containing VOCs. A full description of the 423 TOGA system is provided by Apel et al. (2015). Briefly, during FIREX-AQ the TOGA-TOF 424 continuously sampled 13-mL aliquots of ambient air for approximately 35 seconds every 105 425 seconds, concentrating the VOCs in two cryogenic preconcentration steps prior to injection onto a 426 Restek MTX-624 column (I.D. =  $0.18 \,\mu$ m, length = 8 m). Helium gas carried the samples through 427 the column, which was heated from 25-120°C at a rate of 100°C min<sup>-1</sup>, and the effluent was 428 analyzed by the electron ionization time-of-flight mass spectrometer. The TOF-MS was operated 429 at 70 eV and nominal mass resolution 3000  $\Delta m$  m<sup>-1</sup>. The system was calibrated several times per 430 flight and in the laboratory before and after the campaign using a series of multicomponent VOC 431 standards (Apel-Riemer Environmental Inc., Miami, FL, USA).

432 433

## 434 2.3.4. Aerodyne GC-MS

435

436 The ARI GC-MS system consists of three main components: (1) a thermal desorption 437 preconcentrator (TDPC) (Aerodyne Research, Inc.) for sample collection, (2) a GC (Aerodyne 438 Research, Inc.) for sample separation, and (3) an electron ionization time-of-flight mass 439 spectrometer (EI-ToF-MS) (TOFWERK AG, model EI-HTOF) for sample detection (Gilman et 440 al., 2013; Obersteiner et al., 2016; Claflin et al., 2021). For the MOOSE campaign, the TOF-MS 441 was operated at 70 eV and nominal mass resolution 3000  $\Delta m m^{-1}$ . The GC is a 2-channel system, 442 where both separation channels use identical preconcentration steps. The TDPC employed for this 443 campaign relied upon two-stage adsorbent trapping for preconcentration of analytes. The sample





444 is initially collected onto multibed (Tenax TA/Graphitized Carbon/Carboxen 1000), 445 preconditioned glass sorbent tubes that are optimized for C<sub>2/3</sub>-C<sub>30</sub> species. The first stage of 446 trapping allows sampling rates up to 100 sccm, followed by forward-flushing with UHP nitrogen 447 to remove water. For the higher volatility channel (Channel 1), additional H<sub>2</sub>O was removed from 448 the ambient sample stream via passing through a cooled PFA tube (1/8-inch O.D., 1/16-inch I.D., 449 3.375-inch length,  $15^{\circ}$ C) prior to trapping of VOCs to avoid water condensed in the sample tube. 450 After the initial collection and water purge, the sample was then transferred to the focusing stage, 451 which is a multibed (Tenax, Carbopack X, Carboxen 1003) glass cold trap. After preconcentration, 452 the samples are transferred to different separation columns (Restek Rt-Q-Bond and Rxi-624) for 453 Channels 1 and 2, respectively. Channel 1 is optimized for separation of  $C_3$ - $C_4$  alkanes,  $C_1$ - $C_2$ 454 halocarbons and C1-C3 oxygenates; Channel 2 is optimized for C5-C12 alkanes, C6-C10 aromatics, 455 C<sub>3</sub>-C<sub>6</sub> oxygenates.

456

457 The GC inlet consisted of approximately 8 m of 0.25-inch O.D., 0.15625-inch I.D. PFA tubing 458 connected to a sample pump, with an inlet flow of approximately 1 slpm. The GC pulled a sub-459 flow from the main GC inlet via 1-m length, 0.125-inch O.D., 0.0625-inch I.D. PFA tubing. The 460 sample flow was 80 sccm to each GC channel for 10 minutes during each 30-minute analytical 461 cycle. Before preconcentration, the ambient sample is passed through a bed of pre-cleaned sodium 462 sulfite (nominal 1 g) to scrub ozone and thereby reduce sampling artifacts [Helmig, 1997].

463

464 The TDPC employed for this campaign relied upon two-stage adsorbent trapping for preconcentration of analytes. The sample is initially collected onto multibed (Tenax 465 466 TA/Graphitized Carbon/Carboxen 1000), preconditioned glass sorbent tubes that are optimized for 467  $C_{2/3}$ - $C_{30}$  species. The first stage of trapping allows sampling rates up to 100 sccm, followed by forward-flushing with UHP nitrogen to remove water. For the higher volatility channel (Channel 468 469 1), additional H<sub>2</sub>O was removed from the ambient sample stream via passing through a cooled 470 PFA tube (1/8-inch O.D., 1/16-inch I.D., 3.375-inch length, 15°C) prior to trapping of VOCs to 471 avoid water condensed in the sample tube. After the initial collection and water purge, the sample 472 was then transferred to the focusing stage, which is a multibed (Tenax, Carbopack X, Carboxen 473 1003) glass cold trap.

474

Calibration was performed by multicomponent VOC calibrant gas (Apel-Riemer Environmental
Inc., Miami, FL, USA) diluted in UHP N2 at sufficient total flow to overflow the GC sub-inlet.
Instrument calibrations were performed in-field each night throughout the campaign via automated
valve switching. Note that this calibration mixture was the same as described in Section 2.2.3.

### 480 **3. Results**

481

482 The following sections outline PTR-ToF-MS interferences observed for ions typically assigned to isoprene, oxygenated VOCs, and aromatic VOCs. The primary data used for this analysis are from 483 484 the NOAA PTR-ToF-MS, which provided direct evidence of interferences via GC pre-separation. 485 Each section begins with a description of GC-PTR-ToF-MS samples collected along the Las Vegas 486 Strip during SUNVEx. This region comprises many hotels and entertainment establishments and 487 is impacted by emissions from fossil fuels, VCPs, and restaurant cooking. This region had the 488 highest observed mixing ratios compared to New York, Detroit, and Los Angeles. These 489 interferences are then compared to ground site data collected by the NOAA PTR-ToF-MS during





SUNVEx/RECAP-CA. Finally, other observations are used to show the ubiquity of these
interferences across instruments and urban environments. A key focus of this discussion is the
impact of fragmentation on PTR-ToF-MS observations of m/z 69 (exact mass 69.070), which is
typically assigned to isoprene. We also focus on methods to determine urban isoprene mixing
ratios when fragmentation from other higher molecular weight molecules is important.

- 496 **3.1.** Isoprene
- 497

495

### 498 **3.1.1.** Known interferences to isoprene (m/z 69)

499

500 Biogenic VOCs are commonly reported by PTR-ToF-MS, including isoprene and the sum of 501 monoterpene isomers. Isoprene is the dominant biogenic VOC emitted by urban foliage and is a 502 major contributor to urban OH reactivity (Calfapietra et al., 2013). Interferences to isoprene in 503 PTR-ToF-MS spectra results from the production of the C<sub>5</sub>H<sub>8</sub>H<sup>+</sup> ion, which is a common fragment 504 for higher-carbon aldehydes (> C<sub>5</sub>), alkenes, and cycloalkanes (Buhr et al., 2002;Gueneron et al., 505 2015; Pagonis et al., 2019; Romano and Hanna, 2018). Previous studies have characterized ambient 506 isoprene interferences from 2-methyl-3-buten-2-olalkenes emitted from biogenic sources (e.g., 507 Karl et al., 2012) and cycloalkanes emitted from fossil fuels use and oil and natural gas production 508 (e.g., Gueneron et al., 2015; Warneke et al., 2014; Pfannerstill et al., 2023). For example, Gueneron et al. (2015) showed that substituted cyclohexanes and cyclohexenes produce fragmentation 509 510 patterns that consist largely of m/z 111, m/z 125, m/z 69, m/z 83, m/z 57, and other lesser-abundant 511 hydrocarbon fragments. In regions with significant oil and natural gas development, these 512 compounds may produce interferences at m/z 69 which can interfere with the signal resulting from 513 biogenic sources of isoprene (e.g., Warneke et al., 2014). Similarly, Kilgour et al. (2021) show 514 that aldehydes emitted due to ozone deposition to surface ocean waters may interfere with the 515 quantification of isoprene. The key aldehydes observed to produce an interference were nonanal 516 and octanal. The same aldehydes may be produced on inlet surfaces exposed to high ozone concentrations and result in an isoprene artifact (Ernle et al., 2023; Vermeuel et al., 2022). 517

518

# 519 3.1.2. Characterizing aldehyde interferences to m/z 69 using GC-PTR-ToF-MS 520

Figure 2 shows a GC-PTR-ToF-MS chromatogram of the ion typically assigned to isoprene (m/z 521 522 69, C<sub>5</sub>H<sub>8</sub>H<sup>+</sup>). This sample was collected on the Las Vegas Strip in the evening (~ 22:15 local time, 523 LT) when biogenic emissions of isoprene are expected to be low. The chromatogram shows that 524 isoprene (retention time, RT = 65 s) is only a small contributor to the signal at m/z 69 measured in 525 this region. Additional peaks are observed at RT = 210 s, 250 s, 600 s, and 710s. These peaks are 526 not cycloalkanes, as might be expected from mobile source emissions (Gueneron et al., 2015); 527 rather, these are dehydration and fragmentation products of saturated aldehydes, including 528 methylbutanal, pentanal, octanal, and nonanal. Chromatograms of the parent ions attributed to 529 octanal and octanone (m/z 129,  $C_8H_{16}OH^+$ ), together with nonanal and nonanone (m/z 143, 530  $C_9H_{18}OH^+$ ) are shown in Fig. S3. The parent ion and the dehydration products (m/z 111  $C_8H_{15}^+$ 531 and m/z 125 C<sub>9</sub>H<sub>17</sub><sup>+</sup>, respectively) are observed, but at different ratios between the aldehydes and 532 ketones. Pentanal and methylbutanal almost entirely dehydrate and do not exhibit significant signal 533 at the parent ion mass. Comparisons of ambient observations with GC-PTR-ToF-MS 534 chromatograms of standard mixtures show that only aldehydes (and not the ketones) are observed 535 in significant quantities on the Las Vegas Strip (Fig. S3).





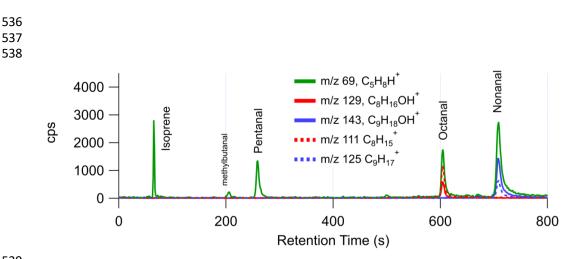


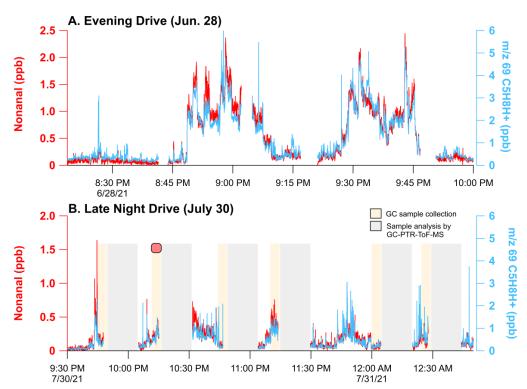
Figure 2. GC-PTR-ToF-MS chromatogram from downtown Las Vegas at 22:15 on 30 July 2021, showing the contributions of isomers and fragments to the ion typically assigned to isoprene (m/z 69, C<sub>5</sub>H<sub>8</sub>H<sup>+</sup>).

543 These aldehydes emissions likely result from cooking (Arata et al., 2021;Klein et al., 2016;Schauer 544 et al., 1999;Karl et al., 2018;Wernis et al., 2022) and their significant presence on the Las Vegas 545 Strip possibly reflects the high density of restaurants along Las Vegas Boulevard. Figure 3 shows 546 mobile laboratory measurements of nonanal and m/z 69 during evening drives on 28 June and 30 547 July 2021. GC-PTR-ToF-MS sampling was only conducted on 30 July and the location of the sample (Las Vegas Strip) is shown in Fig. 1a. During both drives, m/z 69 was enhanced along Las 548 549 Vegas Boulevard and mixing ratios reached a maximum of 6 ppb (assuming a sensitivity 550 equivalent to isoprene). On 28 June, m/z 69 and nonanal detected at m/z 143 ( $C_9H_{18}OH^+$ ) are highly correlated ( $r^2 > 0.93$ ), suggesting that these ions share a common source. A similar 551 correlation was observed between m/z 69 and octanal ( $r^2 = 0.90$ ). 552

553







555 556 Figure 3. Mobile laboratory data showing PTR-ToF-MS measurements of nonanal (m/z 143, C<sub>9</sub>H<sub>18</sub>OH<sup>+</sup>, red) and m/z 557  $69 (C_5H_8H^+, blue)$  on the Las Vegas Strip during nighttime hours on (a) 28 June and (b) 30 July 2021. GC samples 558 were only collected on 30 July, and the shaded regions in (b) show periods of sample collection (beige) and sample 559 analysis (grey). The red marker in (b) indicates the time of the GC-PTR-ToF-MS sample shown in Fig. 2.

560 561

562 Long-chain aldehydes are not routinely reported in urban datasets and the isoprene interference 563 due to aldehyde fragmentation is underappreciated in ambient PTR-ToF-MS datasets. Studies have described how aldehydes produced on the surfaces of inlet tubing interfere with isoprene measured 564 565 by PTR-ToF-MS in remote forests (Vermeuel et al., 2022) and in the stratosphere (Ernle et al., 2023), and aldehydes emitted from ocean surface waters also interfere with isoprene measurements 566 in laboratory studies and ambient measurements near coastal regions (Kilgour et al., 2021). Long-567 chain aldehydes are likely ubiquitous in cities, and cooking activities are likely a major source of 568 octanal and nonanal resulting in an isoprene interference (Wernis et al., 2022;Peng et al., 2022). 569

570

571 The interference from aldehydes is likely common across PTR designs, even though differences could exist due to operating conditions (e.g., the E/N ratio). Figure S1 compares the fragmentation 572 573 patterns of pentanal, octanal, and nonanal observed in the NOAA PTR-ToF-MS (E/N ~ 140 Td), 574 which utilizes the Vocus ion source, to those reported by Buhr et al. (2002) (E/N  $\sim 120-130$  Td), 575 which employ a traditional drift tube and quadrupole. In both reactor designs, C5-aldehydes 576 dehydrate to produce  $C_5H_8H^+$  (m/z 69) directly, while larger aldehydes such as octanal and nonanal 577 dehydrate to produce  $C_8H_{15}^+$  (m/z 111, exact mass: 111.117) and  $C_9H_{17}^+$  (m/z 125, exact mass: 578 125.123), then further fragment to produce the  $C_5H_8H^+$  ion (Buhr et al., 2002; Pagonis et al., 2019).





579 These dehydration products are unique to aldehydes since ketone isomers do not undergo 580 significant fragmentation (Fig. S3).

581

582 We note that even though the fragmentation was similar between two instruments with different 583 reactor designs, fragmentation may result from the operation of other instrument components. For example, the intensity of aldehyde fragmentation was found to vary strongly with voltage gradients 584 within the BSQ of the Berkeley Vocus-PTR-ToF-MS (Fig. S2). These results show that PTR-ToF-585 586 MS systems employing quadrupole ion guides may exhibit fragmentation outside of the drift tube 587 region; consequently, care may be needed when tuning instruments to minimize unwanted 588 secondary reactions.

589

### 590 3.1.3. Corrections to ground site observations of m/z 69 in Los Angeles and Las Vegas

591

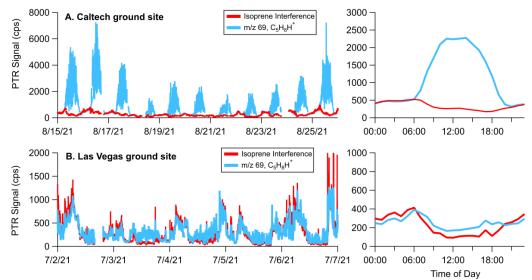
592 GC-PTR-ToF-MS and mobile laboratory measurements described in Section 3.1.2 indicate that 593 aldehydes significantly contribute to the signal at m/z 69 in urban areas. Chromatograms show that 594 the dehydration products from nonanal (m/z 125) and octanal (m/z 111) are useful markers that 595 distinguish aldehydes from ketone isomers. Coincidentally, the dehydration products from nonanal 596 and octanal are identical to the fragments produced from substituted cyclohexanes, which interfere 597 with isoprene in hydrocarbon-rich environments (see Gueneron et al., 2015; Warneke et al., 598 2014;Pfannerstill et al., 2023). Here, it is proposed that the signals at m/z 111 and m/z 125 can be used as proxies to calculate the contribution from aldehyde and cycloalkane fragmentation on the 599 600 signal at m/z 69 in urban areas.

601

602 Figure 4 shows how the sum of m/z 111 and m/z 125 (termed the "isoprene interference") varies relative to the signal at m/z 69 measured at the ground sites in Los Angeles and Las Vegas. In Los 603 604 Angeles, high daytime emissions of isoprene dominate and comprise most of the signal of m/z 69 605 from 6:00–19:00 LT. The high variability in the signal at m/z 69 is caused by very localized 606 emissions from trees upwind of the measurement site. In Las Vegas, isoprene emissions are much 607 lower and the diel pattern of m/z 69 closely follows the behavior of the isoprene interference with 608 only small additional signal during the daytime.







610 611

**Figure 4.** Time series and diurnal pattern of the signal at m/z 69 ( $C_5H_8H^+$ ) and the isoprene interference (m/z 111 + m/z 125) measured at (a) the Caltech ground site and (b) the Las Vegas ground site. The time series data are shown for select periods to illustrate correlations between the isoprene interference and m/z 69. The diel patterns on the right are campaign averages.

615 616

617 Biogenic isoprene is predominantly emitted during daytime hours (e.g., Guenther et al., 2012), 618 while the isoprene interference in both cities is more prevalent at night. These differences in diurnal 619 patterns can be leveraged to subtract interferences from aldehydes and cycloalkanes from PTR-620 ToF-MS measurements of m/z 69. Here, the signals at m/z 69, m/z 111, and m/z 125 are analyzed 621 between 00:00-04:00 LT when daytime isoprene from biogenic sources is expected to be low. The 622 instrument response to aldehyde and cycloalkane fragmentation is calculated by determining the 623 ratio of m/z 69 to the sum of m/z 111 + m/z 125. This ratio is then applied to the full dataset 624 following Eq. 1.

625 626

627

$$m/z 69_{Corrected} = S_{69} - S_{111+125} \cdot f_{69/(111+125)}$$
 (Eq. 1)

628 S<sub>69</sub> is the signal measured at m/z 69, S<sub>111+125</sub> is the signal of the isoprene interference (sum of m/z 629 111 + m/z 125), and f<sub>69/(111+125</sub>) is the interference ratio determined at night. The nighttime 630 interference ratio is 3.0 in Las Vegas and 3.5 in Los Angeles (Fig. S4). The differences between 631 the cities may reflect variations in the distribution of aldehydes and cycloalkanes.

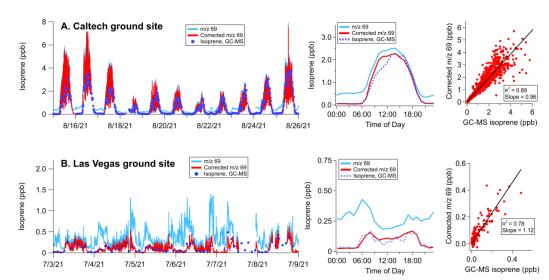
632

633 Figure 5 shows how measurements of m/z 69 change as a result of this correction and compares the corrected/calculated isoprene mixing ratios to the GC-MS measurements co-located with the 634 635 PTR-ToF-MS. In Los Angeles, the correction largely impacts m/z 69 signals at night. The diurnal 636 pattern shows that average mixing ratios approach zero in the evenings, though increases in 637 nighttime isoprene mixing ratios are observed during some periods (e.g., 22–24 August, Fig. S5). Corrections to m/z 69 during the daytime lead to a ~ 10% decrease in reported mixing ratios. This 638 639 shows that even when isoprene emissions are high, VOC fragmentation can have a significant 640 impact on the signal at m/z 69.





641



642 643

Figure 5. Uncorrected and corrected m/z 69 as time series, diel averages, and correlation plots for (a) the Caltech and 644 (b) Las Vegas ground sites. GC-MS measurements are shown for comparison against the corrected m/z 69 isoprene 645 mixing ratios. A detailed comparison of nighttime isoprene corrections in Los Angeles is shown in Fig. S5. 646

647 The corrected m/z 69 measurements are well correlated with GC-MS isoprene measurements ( $r^2$ 648 = 0.89) and agree to within 4%. At high isoprene mixing ratios, the measurements exhibit a greater 649 degree of scatter. This variability likely results from the differences in sampling timescales (1 s for 650 PTR-ToF-MS, ~ 120 s for GC-MS) along with the high variability of isoprene emissions from trees at the measurement site. When averaged to a diel profile, the daytime mixing ratios also agree 651 652 to within 4%. Both instruments show that average isoprene decreases to low mixing ratios at night (< 0.05 ppb). The GC-MS observed a number of periods of enhanced nighttime isoprene, likely 653 654 from non-biogenic sources. Remarkably, after accounting for the isoprene interference, the corrected m/z 69 mixing ratios from the PTR-ToF-MS captures the variability in nighttime 655 isoprene observed by GC-MS in Los Angeles (Fig. S5). On average, the isoprene interference 656 657 represents  $\sim 90\%$  of the nighttime signal of m/z 69.

658

659 The isoprene correction is most impactful on the Las Vegas measurements where isoprene 660 emissions are low and aldehydes + cycloalkane fragments constitute a larger fraction of the signal at m/z 69. Without correction, the variability in m/z 69 across all daytime hours is driven by the 661 isoprene interference (Fig. 4). After the interference contributions are subtracted, corrected 662 663 isoprene mixing ratios approach zero at night and decrease by nearly 50% to 0.1–0.15 ppb during 664 the day (Fig. 5b). The resulting diel pattern changes substantially and exhibits a daytime peak that 665 is consistent with the expected pattern for isoprene. GC-MS measurements show that isoprene 666 mixing ratios were typically < 0.2 ppb and the corrected m/z 69 diel pattern generally matches the 667 average diel pattern of isoprene reported by GC-MS. Though the number of canister samples in Las Vegas were limited (total 275, sampled every 2-4 h), a comparison between the corrected m/z 668 669 69 and GC-MS isoprene shows that the measurements agree to within 15%.

670

#### 671 3.1.4. Corrections to aircraft measurements of m/z 69 over Los Angeles





672

673 The isoprene interferences observed during SUNVEx and RECAP-CA show that PTR-ToF-MS 674 measurements of m/z 69 are significantly impacted by aldehydes and cycloalkanes. To assess the impact of isoprene interferences at higher altitudes, we analyze the FIREX-AQ and RECAP-CA 675 measurements of flights in the Los Angeles Basin and determine corrected m/z 69 signals 676 following Eq. (1). One challenge to this approach is that the DC-8 and Twin Otter aircraft did not 677 678 sample the Los Angeles Basin at night, and therefore the interference ratio (f69/(111+125)) is not easily 679 determined in the absence of isoprene. To overcome this limitation for FIREX-AQ, we vary the 680 interference ratio until the corrected m/z 69 signals reported by the NOAA PTR-ToF-MS matches 681 with the isoprene mixing ratios reported by GC instrumentation on the DC-8. The resulting ratio 682 determined by iteration (4.4) is ~ 20% higher than the ratio determined at ground level during 683 SUNVEx (3.5), which likely reflects differences between the operating conditions and drift tube designs used on the NOAA PTR-ToF-MS during FIREX-AQ and SUNVEx (traditional vs. 684 685 Vocus).

686

687 Figure 6 illustrates the spatial and temporal variability in (a) the corrected m/z 69 mixing ratios 688 and (b) the calculated interference. Transits to the north show that the interference is highest along 689 the San Gabriel Mountains where anthropogenic pollution typically builds in the Los Angeles 690 Basin (Angevine et al., 2013) and reached mixing ratios as high as 500 ppt. The interference 691 correlates well with both methylcyclohexane measured by the UCI WAS and methylpropanal measured by TOGA-TOF, which are proxies for the species known to fragment to produce the 692 693 isoprene interference (i.e., cycloalkanes and aldehydes). Corrected m/z 69 mixing ratios only 694 exhibit significant enhancements in regions where the DC-8 sampled air close to vegetation. Short 695 bursts of isoprene were observed above the San Gabriel Mountains, but mixing ratios were typically lower than 500 ppt. Over the entire flight, the isoprene interference constituted > 50% of 696 697 the signal observed at m/z 69.

698

699 Figure 6c compares the PTR-ToF-MS measurements to GC-based samples for uncorrected m/z 69 700 (top) and corrected m/z 69 mixing ratios (bottom). These comparisons show that the isoprene 701 interference resulted in an overestimation of PTR-ToF-MS measurements of isoprene by at least a 702 factor of 2. At times, the NOAA PTR-ToF-MS measured mixing ratios of m/z 69 that were 5 times 703 larger than the isoprene mixing ratios reported by GC-based methods. The Oslo PTR-ToF-MS also 704 sampled onboard the DC8 during FIREX-AQ and presented an opportunity to compare to the 705 fragmentation observed by the NOAA PTR-TOF-MS. Following a similar correction procedure as 706 described above, Fig. S6 shows that the Oslo PTR-ToF-MS measured the same degree of 707 interferences as the NOAA PTR-ToF-MS (i.e., fragmentation biased isoprene measurements high 708 by at least a factor of 2). The consistency between both instruments demonstrates that isoprene 709 interferences are common across PTR-ToF-MS designs (i.e., Tofwerk vs. Ionicon).





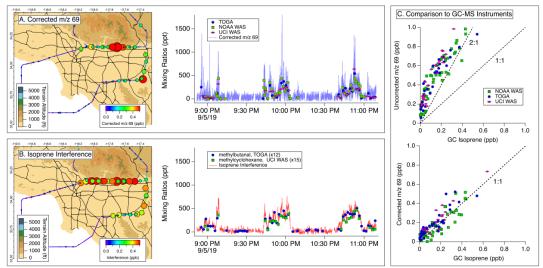




Figure 6. Impact of isoprene interference correction on m/z 69 measurements from the NOAA PTR-ToF-MS during
FIREX-AQ. (a) Map of corrected m/z 69 distribution (left) and time series with corresponding measurements of
isoprene from GC-MS samples (right). (b) Map of isoprene interference (left) and time series with GC-MS
measurements of methylcyclohexane and methylpropanal, which are proxies for cycloalkanes and aldehydes known
to contribute to the signal at m/z 69. (c) Comparisons of PTR-ToF-MS measurements of m/z 69 and GC-based isoprene
mixing ratios for uncorrected m/z 69 (top) and corrected m/z 69 (bottom) using Eq. (1) with an interference ratio =
10.

720 The Berkeley Vocus PTR-ToF-MS also observed interferences to m/z 69 during the RECAP-CA 721 flights. Unlike FIREX-AQ, GC-MS measurements were not available onboard the Twin Otter to compare against PTR-ToF-MS measurements. To evaluate the interference contributions to m/z 722 723 69 here, we determine the interference ratio from nonanal calibrations and compare it with data 724 collected from the Central Valley and Los Angeles Basin. In these regions, the signal at m/z 69 725 and the sum of  $m/z \ 111 + 125$  are well-correlated with a slope that closely matches the measured 726 fragmentation pattern for nonanal (Fig. S7). The interference ratio observed in the Central Valley 727 where oil and natural gas emissions are significant is similar to the ratio observed in Los Angeles 728 where aldehydes are more important. In the Central Valley, periods of high biogenic influence are clearly separated from periods of high interference from anthropogenic emissions. Building on 729 730 these responses, we use the calibrated data for nonanal to derive an m/z 69 correction in Los 731 Angeles. This method is limited in accounting for molecules that have fragmentation ratios 732 differing from nonanal, but since the interference ratio observed from oil and gas regions and from 733 the Los Angeles Basin is similar to the nonanal fragmentation ratio, we expect the uncertainty to 734 be relatively small despite there being no GC comparison.

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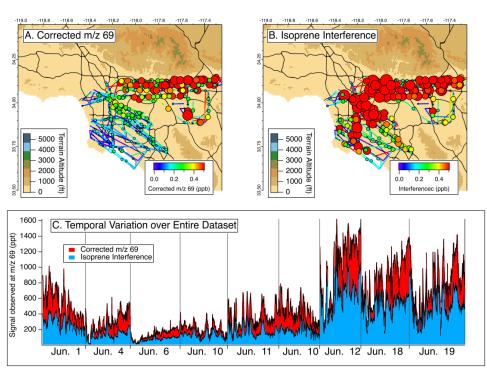
Figure 7 shows the impact of the isoprene interference on the Berkeley PTR-ToF-MS data. The
Twin Otter flew nine flights and the total signal of m/z 69 varied between 200–1200 ppt. Similar
to the observations by the NOAA PTR-ToF-MS during FIREX-AQ, the isoprene interference
during RECAP-CA was at least 50% of the signal observed at m/z 69 (Fig. 7c). The Twin Otter
sampled a larger swath of area than the DC-8, and Fig. 7 shows that the interference is persistent
across the Los Angeles Basin (Fig. 7b) at mixing ratios as high as 600 ppt. Similar to the DC-8
flight, corrected m/z 69 mixing ratios are highest along the San Gabriel Mountains. The Twin Otter





- 743 is capable of sampling at lower altitudes than the DC-8, and therefore larger mixing ratios of
- 744 isoprene were observed.
- 745





## 747 748

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Figure 7. Impact of isoprene interference correction on isoprene measurements during RECAP-CA as (a) the corrected isoprene distribution, (b) the isoprene interference, and (c) a pseudo time series of the total m/z 69 signal colored by the contributions of corrected m/z 69 and isoprene interference for all flights.

750 751

752 The contribution of the isoprene interference observed from the aircraft with the Berkeley PTR-753 ToF-MS differs from the observations at Caltech during RECAP-CA. On the ground, the isoprene 754 interference was  $\sim 10\%$  of the signal at m/z 69 during daytime hours (Fig. 5a), while at altitude it was > 50%. This difference can be explained by (1) the abundance of isoprene emitters close to 755 756 the ground site, (2) the differences in reactivities between isoprene, aldehydes, and cycloalkanes, and (3) the different instrument setting of the Berkeley PTR-ToF-MS (Fig. S2). Isoprene is highly 757 reactive towards atmospheric oxidants such as the OH radical ( $k_{OH} \sim 1 \times 10^{-10} \text{ cm}^3$  molecule<sup>-1</sup> s<sup>-1</sup> 758 ), whereas saturated aldehydes and cycloalkanes are expected to be 5-10 times less reactive 759 (Burkholder et al., 2019). This difference in reactivity may alter the distribution of VOCs that 760 contribute to m/z 69 and result in higher interferences aloft. The DC-8 and Twin Otter aircraft did 761 762 not specifically target altitude profiling while sampling in the Los Angeles Basin, but future work may help to characterize the impact of the isoprene interference at other altitudes. 763

764

#### 3.1.5. Corrections to m/z 69 measured in Detroit, MI 765

766

767 The SUNVEx/RECAP-CA/FIREX-AQ data reflect the behavior of the NOAA, Oslo, and Berkeley 768 instruments during summertime measurements in Los Angeles. Isoprene interferences likely





769 impact ground and airborne measurements in other cities and at other times of year. Figure 8 shows 770 the impact of interferences to the signal at m/z 69 reported by the Aerodyne PTR-ToF-MS 771 measurements during MOOSE. This campaign targeted emissions in Detroit, MI, where the Aerodyne Mobile Laboratory conducted a mix of mobile and stationary sampling in select 772 locations around the metropolitan area (Fig. 1). Figure 8a shows a subset of PTR-ToF-MS 773 measurements of m/z 69 and corrected m/z 69, along with isoprene measurements by the Aerodyne 774 GC-MS. Corrected m/z 69 is calculated by determining the interference ratio at night (= 2.54), 775 776 similar to the approach used to calculate interferences during SUNVEx/RECAP-CA. Figure 8b 777 shows normalized histograms for each measurement over the entire deployment.

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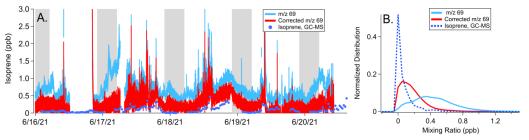


Figure 8. Impact of isoprene interference correction on the PTR-ToF-MS data collected from the Aerodyne Mobile
Laboratory during the MOOSE campaign as a (a) time series of GC-MS measurements, uncorrected and corrected
m/69 isoprene mixing ratios. The shaded regions show nighttime measurements (22:00–05:00 LT). (b) Histograms
showing the distribution of m/z 69 measured by the Aerodyne PTR-ToF-MS, corrected m/z 69 isoprene mixing ratios, and isoprene mixing ratios measured by the GC-MS.

785

786 Without correction, m/z 69 reported by PTR-ToF-MS exhibits a broad distribution with a peak mixing ratio of ~ 0.4 ppb (Fig. 8b). After applying the corrections described by Eq. (1), m/z 69 787 788 signals decrease by nearly a factor of 2 and show better agreement with isoprene reported by GC-789 MS. Corrected m/z 69 mixing ratios are still a factor of 2 higher than the isoprene mixing ratios 790 reported by the GC-MS. One possibility is that other VOCs in the Detroit region may also contribute to the signal at m/z 69. GC pre-separation measurements directly using PTR-ToF-MS 791 792 were not conducted during MOOSE; consequently, it is difficult to determine what other species 793 might contribute to m/z 69 in this region. Broad deployment of GC-PTR-ToF-MS measurements 794 in urban areas may help to better quantify the contributions of fragmenting species to PTR-ToF-795 MS measurements of m/z 69.

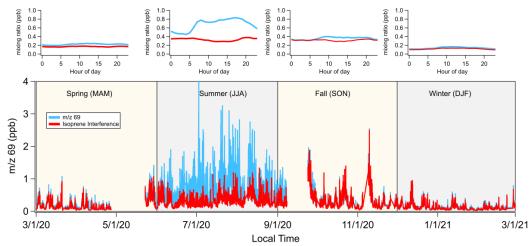
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- 797 798

### 3.1.6. Seasonal changes to the m/z 69 interferences observed in New York City

Figure 9 shows the impact of the isoprene interference on data reported by the Stony Brook PTR-ToF-MS during a year-long sampling effort to characterize emissions in New York City during the COVID-19 lockdown. Shown here are mixing ratios of m/z 69 along with the estimated contribution to m/z 69 from the isoprene interference. We calculate the isoprene interference for each season, and present the diurnal patterns in the top row. The interference ratio (f<sub>69/(111+125)</sub>, Eq. (1)) is similar in spring, summer, and winter (2.2–2.5), but lower during fall (1.9).







Local Time
 Figure 9. (bottom) Time series of m/z 69 and isoprene interference measured by the Stony Brook PTR-ToF-MS at the urban ASRC ground site in New York City. (top) Diel patterns of m/z 69 and isoprene interference mixing ratios for each season.

810

811 The signal at m/z 69 is variable across seasons and the highest mixing ratios are observed during 812 summer. The isoprene interference is a major contributor to m/z 69 in fall, winter, and spring (77-813 88% of total signal) and strongly influences the day-to-day variability. During summertime 814 isoprene emissions from urban foliage increases the variability in m/z 69 and results in higher 815 mixing ratios of m/z 69 during the day. The isoprene interference increases the background mixing 816 ratios of m/z 69 and dominates the total signal at night.

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The ASRC site is located in a heavily urbanized region and the PTR-ToF-MS sampled air at the 818 819 top of the building where mixing ratios of isoprene are likely lower. The persistent, high contribution from the isoprene interference to m/z 69 during all seasons likely reflects the high 820 821 emissions of aldehydes and cycloalkanes from anthropogenic sources in this region. Figure S8 822 contrasts the measurements at ASRC with those reported from the Flax Pond site. Flax Pond is 823 located in a less-densely populated region of Long Island where biogenic sources of isoprene are 824 more abundant. There, interferences are a much smaller fraction of the signal at m/z 69 ( < 10%) 825 and the variability is largely driven by isoprene during the summer months. Mixing ratios at Flax 826 Pond are lower during the winter, but comparable to those observed at ASRC during the same 827 season (~ 100-150 ppt). Furthermore, the variability is predominantly driven by the isoprene interference. Figures 9 and S8 demonstrate that interferences will vary spatially between heavily 828 829 urban and biogenic-dominated regions. In addition, outside of the summer months, isoprene is 830 unlikely to be a major contributor to m/z 69 in both regions.

831

The contribution of the isoprene interference to m/z 69 in New York City is comparable to the ground level measurements in Las Vegas, Los Angeles, and Detroit (0.25–0.5 ppb), which demonstrates that isoprene measurements by PTR-MS are likely to be significantly impacted across most, if not all, urban regions. Identifying anthropogenic or nighttime sources of isoprene by PTR-MS will be difficult if not confirmed unambiguously by GC-PTR-ToF-MS or separately by GC-MS.





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### 839 3.2. Oxygenated VOCs

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# 841 **3.2.1.** Characterizing interferences to oxygenated VOCs using GC-PTR-ToF-MS 842

Small oxygenated VOCs are an important contributor to the reactivity and ozone produced in urban 843 844 areas. Alcohols, ketones, and small aldehydes ( $< C_4$ ) may be emitted to the atmosphere from 845 mobile sources, VCPs, cooking activities, and other sources (Klein et al., 2016;McDonald et al., 846 2018), but are also formed as secondary products of atmospheric chemistry. Some studies have 847 reported that certain alcohols, such as ethanol, may ionize to product products that overlap with 848 proton-transfer products of other important oxygenates, such as acetaldehyde (Buhr et al., 849 2002; Pagonis et al., 2019). Previous intercomparisons have shown that acetaldehyde is one 850 example of oxygenated VOCs where there may be large disagreements between PTR-ToF-MS and 851 GC-MS (Yuan et al., 2016).

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Figure 10 shows the GC-PTR-ToF-MS chromatogram collected on the Las Vegas Strip for protontransfer products typically assigned to oxygenated VOCs. Here, we present small oxygenates
typically reported in ambient data sets that are subject to fragmentation or interferences, including
methanol, acetaldehyde, ethanol, and C4-carbonyls, which represent the sum of methacrolein
(MACR), methyl vinyl ketone (MVK), and crotonaldehyde.

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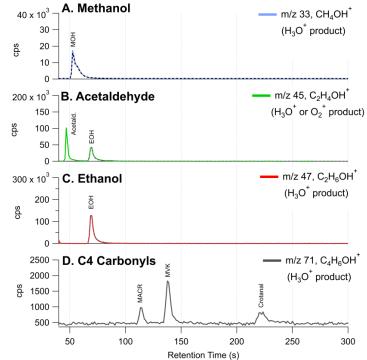
First, GC-PTR-ToF-MS data show that no other species elute through the GC column to yield a significant interference to methanol and ethanol. The signal at m/z 59 (C<sub>3</sub>H<sub>6</sub>OH<sup>+</sup>) is also observed to result entirely from acetone + propanal (not shown). This is consistent with previous studies that show good agreement between GC-MS and PTR-ToF-MS (e.g., Warneke et al., 2003).

863

864 Crotonaldehyde is a major fraction of the C4-carbonyls observed on the Las Vegas Strip. Typically, 865 MVK and MACR are treated as the dominant isomers to the C<sub>4</sub>-carbonyl product (m/z 71, 866  $C_{4}H_{6}OH^{+}$ ), since these are secondary products from isoprene oxidation and are expected to be 867 present at high mixing ratios (Yuan et al., 2017). Crotonaldehyde is observed to be a major 868 contributor to m/z 71 in biomass burning emissions (Koss et al., 2018), but its presence on the Las 869 Vegas Strip likely points to other important aldehyde sources, such as cooking. The higher fraction 870 of crotonaldehyde reflects that isoprene mixing ratios are lower in Las Vegas than other cities (Fig. 871 5) and that cooking is an important source of VOCs along the Las Vegas Strip (Fig. 2). Xu et al. 872 (2022) showed that measurements of C4-carbonyls by the NOAA PTR-ToF-MS, ammonium-873 adduct chemical ionization mass spectrometer (NH4-CIMS), and NOAA GC-MS agreed in the 874 daytime during RECAP-CA when MVK and MACR were high, but disagreed at night when 875 isoprene products were low and crotonaldehyde mixing ratios were likely elevated. Additional 876 interferences at m/z 71 could result from decomposition of ISOPOOH on inlet surfaces (Rivera-877 Rios et al., 2014).







Retention Time (s)
 Figure 10. GC-PTR-ToF-MS chromatograms from the Las Vegas Strip showing the contributions of isomers and fragments to ions typically assigned to small oxygenates. The labels highlight the traditionally assigned isomers for (a) methanol, (a) acetaldehyde, (c) ethanol, and (d) C<sub>4</sub>-carbonyls including methacrolein (MACR), methyl vinyl ketone (MVK), and crotonaldehyde.

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887 The most significant interference for the small oxygenated VOCs observed by GC-PTR-ToF-MS 888 is associated with the ionization of ethanol to produce signal at the mass typically assigned to 889 acetaldehyde (m/z 45,  $C_2H_5O^+$ ). Ethanol has been shown by Inomata and Tanimoto (2009) to produce fragments at m/z 19 (H<sub>3</sub>O<sup>+</sup>), m/z 31 (CH<sub>3</sub>O<sup>+</sup>), and m/z 29 (C<sub>2</sub>H<sub>5</sub><sup>+</sup>). Buhr et al. (2002) 890 identified m/z 45 as product and found that it correlated to the ethanol proton transfer product (m/z 891 892 47,  $C_{2}H_{6}OH^{+}$ ) with a ratio of 0.22. The likely pathway for the formation of m/z 45 is by ethanol reactions with  $O_{2^+}$ , which has been identified by Spanel and Smith (1997) as the dominant  $O_{2^+}$ 893 894 product using selective ion flow tube (SIFT) mass spectrometry. The NOAA Vocus PTR-ToF-MS 895 observes a ratio that is higher than that determined by Buhr et al. (2002) (~0.38), although the 896 distribution of total fragmentation (i.e., the sum of all ethanol fragments relative to m/z 47) appears 897 similar (Fig. S1).

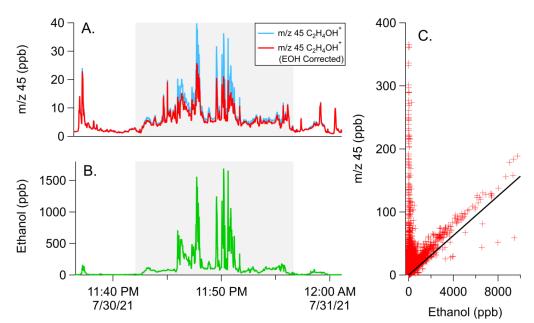
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Figure 11 shows the temporal behavior of m/z 45 and m/z 47 (ethanol) during the nighttime drive
on 30 July. Figure 11a shows the mixing ratio of m/z 45 assuming the sensitivity of acetaldehyde
and Fig. 11b shows the mixing ratio of ethanol. Figure 11c shows a scatter plot of the signal at m/z
45 vs. that of m/z 47 for the entire mobile laboratory dataset. First, ethanol and m/z 45 are
correlated when ethanol mixing ratios are high (Fig. 11a, b). Ethanol on the Las Vegas Strip





904 reached mixing ratios of 1.5 ppm and corresponding increases in m/z 45 were observed that point
905 towards a contribution from ethanol. Figure 11c shows that a subset of the m/z 45 signal measured
906 throughout the Las Vegas dataset exhibit a ratio to ethanol that agrees with the fragmentation ratio
907 observed from GC-PTR-ToF-MS measurements. These observations point towards a broader
908 impact of ethanol on m/z 45 throughout the Las Vegas region.



...

910 911 Figure 11. Demonstration of impacts of ethanol on mobile drive data in downtown Las Vegas during the evening 912 drive on 30 July. (a) Time series of the signal at m/z 45 (C<sub>2</sub>H<sub>4</sub>OH<sup>+</sup>) with and without the subtraction of the ethanol 913 interference. (b) Time series of ethanol (m/z 47, C<sub>2</sub>H<sub>6</sub>OH<sup>+</sup>). The shaded regions show when the mobile laboratory was 914 sampling along the Las Vegas strip. (c) Correlation plot of mobile drive data for the entire Las Vegas dataset. The 915 solid line shows the fragmentation ratio of m/z 45 to m/z 47 for ethanol, as derived from the GC-PTR-ToF-MS data 916 (Fig. 10).

### 918 3.2.3. Corrections to m/z 45 measured in Las Vegas

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920 The extent to which ethanol contributed to the signal at m/z 45 can be determined by correction
921 techniques. Figure 11a shows the m/z 45 signal with the contribution from ethanol subtracted
922 following:

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- 924 925

$$m/z 45_{Corrected} = S_{45} - S_{47} \cdot f_{45/47}$$
 (Equation 2)

S45 is the signal from m/z 45, S47 is the signal of ethanol, and f45/47 is the ratio determined by GCPTR-ToF-MS. Generally, the ethanol-corrected data on m/z 45 show that ethanol contributed
~40% to the signal on the Las Vegas Strip. Outside of this region, ethanol ionization has a modest
impact on m/z 45. Over the average mobile laboratory dataset, ethanol may have contributed as
much as 5% to the total signal at m/z 45. Similar contributions are estimated for the ground site
data collected during ground sampling at Caltech. Consequently, ethanol reactions with O2<sup>+</sup> may





only be an important contributor to m/z 45 in highly-concentrated ethanol plumes, which may be encountered during mobile sampling or upon aircraft encounters with point sources. This ratio may also be affected by humidity, which changes the distribution of  $O_2^+/H_3O^+$  in drift tubes operated at low water mixing ratio.

936

937 The GC-PTR-ToF-MS provides some insights into the interferences of oxygenates, but there are 938 limits to the extent to which oxygenates elute through a DB-624 or other GC columns with similar 939 polarity. Interferences towards oxygenated masses may be an important focus for future work, as 940 recent studies have pointed towards the increasing fraction of oxygenated VOCs observed in urban 941 air (Karl et al., 2018;Xu et al., 2022;Khare et al., 2022) and instrumentation capable of measuring unfragmented oxygenates are becoming more common (e.g., Khare et al., 2022;Xu et al., 942 943 2022; Riva et al., 2019). Intercomparisons with GC-MS measurements employing polar columns, 944 or with mass spectrometers employing softer ionization chemistry (e.g., iodide or NH4<sup>+</sup> adduct 945 mass spectrometers) may help to better characterize the response and selectivity of PTR-ToF-MS 946 to oxygenates.

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# 948 **3.3.** Aromatic VOCs 949

### 950 **3.3.1.** Known interferences to aromatic masses

952 PTR-ToF-MS is well suited to measure ambient mixing ratios of C<sub>6</sub>–C<sub>9</sub> aromatics; however, it is 953 known that alkyl aromatics (e.g., ethylbenzene and ethyltoluene isomers) and aromatic 954 monoterpenes and monoterpenoids (e.g., cymene and fenchone, Kari et al., 2018; Tani, 2013) 955 fragment and contribute to the signals typically attributed to benzene (m/z 79,  $C_6H_6H^+$ ) and toluene 956  $(m/z 93, C_7H_8H^+)$  (Pagonis et al., 2019; Yuan et al., 2017). The abundance and distribution of 957 aromatics depends on the relative mix of VOC emissions from petrochemical sources, including 958 fossil fuels, solvents emitted from VCPs (e.g. paints and architectural coatings), and asphalt paving 959 (Gkatzelis et al., 2021a;Gkatzelis et al., 2021b;Khare et al., 2020;Stockwell et al., 2021). Higher 960 aromatics, such as xylenes and ethylbenzene, are prevalent in both fossil fuel and VCP emissions, 961 whereas benzene is restricted in consumer products and is therefore almost entirely associated with 962 fossil fuels (McDonald et al., 2018). Consequently, PTR-ToF-MS measurements in urban regions 963 with significant solvent emissions may exhibit a greater degree of interference on benzene and 964 toluene than regions with greater fossil fuel usage.

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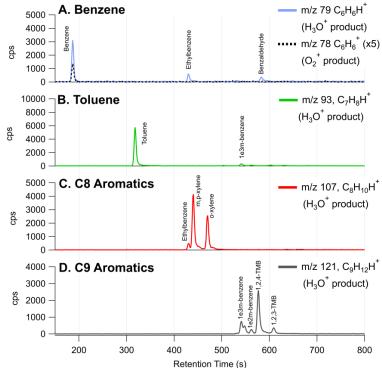
### 966 **3.3.2.** Characterizing interferences to benzene (m/z 79) using GC-PTR-ToF-MS

967

968 Figure 12 shows GC-PTR-ToF-MS measurements of key aromatic compounds measured in 969 downtown Las Vegas where both fossil fuels and VCP emissions were prevalent. Each panel is 970 labeled by the typical compound assignment and shows chromatograms of the corresponding 971 proton-transfer product. In general, chromatograms show that C<sub>9</sub>- and C<sub>8</sub>-aromatics are the 972 expected key contributors to the signals at m/z 121 ( $C_9H_{12}H^+$ ) and m/z 107 ( $C_8H_{10}H^+$ ), 973 respectively. This is consistent with previously observed PTR-ToF-MS behavior (Yuan et al., 974 2017), and shows that urban measurements at these masses continue to be reliably assigned to 975 simple alkyl aromatics.







976Retention Time (s)977Figure 12. GC-PTR-ToF-MS chromatogram from downtown Las Vegas showing the contributions of isomers and978fragments to ions typically assigned to  $C_6-C_9$  aromatics. The labels highlight the traditionally assigned isomers for (a)979benzene, (b) toluene, (c)  $C_8$ -aromatics including o,m,p-xylene + ethylbenzene, and (d)  $C_9$ -aromatics including980ethyltoluene isomers + trimethylbenzene isomers.

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In contrast, the masses typically assigned to benzene  $(m/z 79, C_6H_6H^+)$  and toluene  $(m/z 93, C_6H_6H^+)$ 982 983  $C_7H_8H^+$ ) show greater contributions from the fragmentation of alkyl aromatics. At m/z 93, most 984 of the signal is attributed to toluene and a small fraction (< 5%) results from the fragmentation of 985 1-ethyl-3-methylbenzene. At m/z 79,  $\sim 80\%$  of the signal results from the proton-transfer product of benzene and the remainder from the fragmentation of ethylbenzene and benzaldehyde. Previous 986 work has shown contributions of ethylbenzene to m/z 79 in urban air (Inomata et al., 2010), 987 988 whereas contributions from benzaldehyde are not well studied. Benzaldehyde may result from 989 VCPs, cooking, motor vehicle emissions, biomass burning, or secondary production (Gkatzelis et 990 al., 2021a;Koss et al., 2018;McDonald et al., 2018).

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# 3.3.3. Corrections to ground site observations of m/z 79 in Los Angeles and Las Vegas 993

994 The interferences at m/z 79 are significant and present a challenge for reliably quantifying benzene 995 in Las Vegas and other urban regions. To quantify this interference, Fig. 13 highlights benzene 996 measurements from the Jerome Mack and Caltech ground sites. Figure 13a, b show corrected and 997 uncorrected benzene at m/z 79 can be attributed to benzene calculated from two methods:

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999  $m/z 79_{Corrected} = S_{C6H6H+} - S_{C7H6OH+} \cdot f_{79/Benzald} - S_{C8H10H+} \cdot f_{79/Ethylbenzene}$  (Method 1)





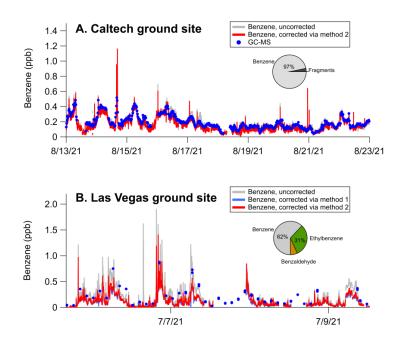
1000 1001 1002  $m/z 79_{Corrected} = S_{C6H6+}$  (Method 2) 1003 1004 1005 Where  $S_{C6H6H+}$  is the signal of  $C_6H_6H^+$ ,  $S_{C7H6OH+}$  is the signal attributable to benzaldehyde,  $S_{C8H10H+}$ is the signal attributable to ethylbenzene, and  $S_{C6H6+}$  is the signal attributed to the benzene charge-1006 1007 transfer product at m/z 78. f79/Benzald and f79/Ethylbenzene are the fragmentation patterns of 1008 benzaldehvde ( $f_{79/\text{Benzald}} = C_6H_6H^+/C_7H_6OH^+$ ) and ethylbenzene ( $f_{79/\text{Ethylbenzene}} = C_6H_6H^+/C_8H_{10}H^+$ ) 1009 as determined by GC-PTR-ToF-MS chromatograms. Method 1 corrects for benzene by subtracting 1010 the contributions of benzaldehyde and ethylbenzene to the signal at  $C_6H_6H^+$ . GC-PTR-ToF-MS 1011 measurements show that benzaldehyde is the primary contributor to Sc7H6OH+, whereas 1012 ethylbenzene is one of four isomers that contributes to S<sub>C8H10H+</sub>. We use the GC-PTR-ToF-MS 1013 measurements at the Jerome Mack ground site and find that ethylbenzene contributes ~12.5% of 1014 the total C<sub>8</sub>-aromatic signal. Method 2 simply estimates benzene mixing ratios based on 1015 calibrations applied to the charge-transfer product at m/z 78 (C<sub>6</sub>H<sub>6</sub><sup>+</sup>). This mass has no discernible 1016 interference from other VOCs in the GC-PTR-ToF-MS data (Fig. 12) and is detected with 1017 sufficient sensitivity to reliably quantify benzene (~ $180 \text{ cps ppb}^{-1}$ ). We note that Method 1 requires regular quantification of C8-aromatic distributions by GC in order to account for ethylbenzene 1018 1019 fragmentation, whereas Method 2 relies only on measurements of the  $O_2^+$  charge-transfer product. 1020 We note that Method 2 may present limitations if other species are present that fragment to produce 1021 the  $O_2^+$  product. Deployment of GC-PTR-ToF-MS in other cities may help to determine whether 1022 the charge-transfer product is unambiguously linked to benzene. 1023 1024 On average, the interferences from ethylbenzene and benzaldehyde constitute ~ 38% of the signal

1025 at m/z 79 detected at Jerome Mack, and ~ 3% of signal detected at Caltech (pie charts, Fig 13). 1026 We speciate the interference at Jerome Mack using GC-PTR-ToF-MS and find that the majority 1027 of the interference is associated with fragmentation of ethylbenzene (31% of total signal) with a 1028 small contribution from benzaldehyde (7% of total signal). Ethylbenzene was likely emitted from 1029 a local source due to a cabinet-making shop upwind of the ground site. The two methods for 1030 correcting benzene agree well for Jerome Mack data, which confirms that ethylbenzene and 1031 benzaldehyde are the primary contributors to the benzene interferences. We note that we only use 1032 Method 2 for data collected at Caltech since GC-PTR-ToF-MS measurements were unavailable 1033 during this period of the deployment.

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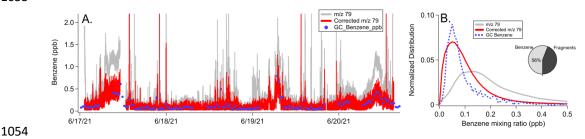


1036<br/>1037Figure 13. Impact of fragmentation on the signal at m/z 79 ( $C_6H_6H^+$ ) and corresponding benzene mixing ratios<br/>measured at (a) Caltech and (b) the Jerome Mack ground site in Las Vegas. The corrections using the two methods<br/>are shown compared to uncorrected data. The pie charts show the average contribution of benzene, ethylbenzene, and<br/>benzaldehyde to the signal at m/z 79 ( $C_6H_6H^+$ ).1041

### 1042 3.3.4. Corrections to observations of m/z 79 in Detroit

1044 The significant interferences to m/z 79 observed in Las Vegas are also observed in PTR-ToF-MS 1045 data collected downwind of Detroit during the MOOSE campaign. Figure 14 shows four days of mobile laboratory sampling with the Aerodyne PTR-ToF-MS and GC-MS instrumentation. 1046 1047 Similar to Fig. 13, corrected via Method 2 and uncorrected m/z 79 measurements are shown (Fig. 1048 14a). Figure 14b shows the histograms of PTR-ToF-MS measurements alongside those from the 1049 GC-MS during the entire campaign. The pie chart shows the average fraction of m/z 79 attributed to benzene vs. the fraction associated with fragments. Benzaldehyde and ethylbenzene 1050 1051 contributions are not separated since GC-PTR-ToF-MS measurements were unavailable. 1052









1055Figure 14. (a) Time series of GC-MS samples of benzene and PTR-ToF-MS mixing ratios of m/z 79 ( $C_6H_6H^+$ ) and1056corrected m/z 79 reported during sampling by the Aerodyne Mobile Laboratory downwind of Detroit, MI. (b)1057Histograms showing the distribution of m/z 79 measured by the Aerodyne PTR-ToF-MS, corrected m/z 79 mixing1058ratios calculated using Method 2, and benzene mixing ratios measured by the GC-MS. The pie chart highlights the1059fraction of the m/z 79 associated with benzene vs. the fraction associated with fragments.

1060

1061 Similar to Las Vegas, fragmentation of higher aromatic species plays an important role in 1062 determining the benzene signal in m/z 79. The distribution of uncorrected m/z 79 shows a peak around 0.12 ppb and a broad tail biased towards higher mixing ratios. The GC-MS measures a 1063 1064 distribution of benzene with a maximum at 0.05 ppb, and a much lower frequency of higher mixing 1065 ratios. When the PTR-ToF-MS data are calibrated using the benzene charge-transfer product 1066 (Method 2), corrected m/z 79 mixing ratios show a better agreement with GC-MS measurements. 1067 The distribution of corrected m/z 79 is wider than that reported by GC-MS, which may reflect the 1068 faster sampling of the PTR-ToF-MS and more frequent observations of concentrated aromatic 1069 plumes (Fig. 14a). Over the entire sampling period, the average distribution of m/z 79 shows that 1070 benzene accounts for ~ 56% of the total signal. This is consistent with the observations in Las 1071 Vegas (interference  $\sim 62\%$  of the signal), indicating that m/z 79 in both datasets were influenced 1072 by solvent emissions to a significant extent.

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### 1074 4. Conclusions

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1076 Urban VOCs significantly contribute to the degradation of air quality, and PTR-ToF-MS provides 1077 important constraints on the emissions and chemical transformation of many gas-phase organics. 1078 Advances in PTR-ToF-MS sensitivity and detection provide opportunities to identify, characterize, 1079 and revisit measurement interferences to commonly reported VOCs (Yuan et al., 2017). Here, we 1080 find that long-chain aldehydes, along with previously identified cycloalkanes, are important 1081 contributors to the signal of m/z 69 typically associated with isoprene in many urban areas. The 1082 fragmentation of these molecules can be larger than the signal associated with the proton-transfer 1083 product from isoprene, depending on the mixture of anthropogenic and biogenic VOCs, time of 1084 day, and season. We find that interferences at ground level in Los Angeles (large isoprene 1085 emissions, large anthropogenic emissions) are highest at night and constitute ~10% of the signal 1086 observed during the day. Interferences are a higher fraction of m/z 69 at altitude (> 50%) and are 1087 observed to be widespread throughout the Los Angeles Basin. In Las Vegas (low isoprene 1088 emissions, large anthropogenic emissions), interferences dominate the signal at m/z 69 throughout 1089 the day and night. These interferences are observed in other cities (e.g., Detroit and New York 1090 City), depend on season, and are common among drift tube designs operated at similar E/N.

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1092 Other PTR-ToF-MS masses also exhibit interferences, including those typically assigned to 1093 oxygenates and aromatic VOCs. Fragmentation from ethanol impacts measurements of 1094 acetaldehyde on m/z 45, though this interference is only significant in regions with large ethanol 1095 emissions. PTR-ToF-MS measurements of benzene using m/z 79 exhibit significant interferences 1096 from the fragmentation of ethylbenzene and benzaldehyde. In Las Vegas and Detroit, 1097 fragmentation impacts m/z 79 mixing ratios by as much as 40%. The growing contribution of 1098 interferences to aromatics and oxygenates may reflect the changing mix of urban VOC emissions 1099 from one dominated by mobile sources to one dominated by solvents (Gkatzelis et al., 1100 2021b;McDonald et al., 2018). In the case of benzene, other ions such as the charge-transfer 1101 product (m/z 78,  $C_6H_6^+$ ) can be used to quantify benzene without significant influence from





1102 fragmentation from higher carbon VOCs. As instrument sensitivity improves, there may be other 1103 ions that can be used to improve the quantification of additional VOCs.

1104

1105 Corrections to these interferences are feasible, though it is unlikely that a universal correction 1106 factor is sufficient to resolve instrument discrepancies across datasets. Instrument responses, as 1107 well as changes to the VOC mixture in different regions, require that detailed characterization be 1108 performed on a dataset-by-dataset basis. GC-MS measurements provide an opportunity to compare 1109 against PTR-ToF-MS measurements for a wide-variety of key VOCs, including isoprene, small 1110 oxygenates, and aromatics. Likewise, information about fragmentation and instrument-specific 1111 responses to reactive hydrocarbons can be determined using GC-PTR-ToF-MS. For species such 1112 as oxygenates, intercomparisons against other mass spectrometers using softer ionization (e.g., 1113 iodide or ammonium-adduct CIMS) or use of GC pre-separation using polar columns may yield 1114 valuable information about instrument artifacts.

- 1116 Data Availability
- 1117

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1118 Data for SUNVEx and RE-CAP are available at the NOAA CSL data repository 1119 (<u>https://csl.noaa.gov/projects/sunvex/</u>). Data for FIREX-AQ, MOOSE, and LISTOS are available 1120 at the NASA data repository (<u>https://www-air.larc.nasa.gov/missions.htm</u>).

## 1122 Author Contribution

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MMC, CES, XL, JBG, AL, CW, EYP, CA, EFK, and AHG conducted measurements during
SUNVEx and RE-CAP. MMC, GIG, CW, JBG, AL, AW, FP, DB, RH, and ECA conducted
measurements during FIREX-AQ. MSC, BML, FM, and MC conducted measurements during
MOOSE. JM, CC, and JEM conducted measurements during LISTOS. MMC and CW wrote the
paper with contributions from all authors.

1129

### 1130 Competing Interests

1131

EAC is a co-editor of Atmospheric Measurement Techniques. The peer-review process was guided
by an independent editor and the authors also have no other competing interests to declare.

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- 1149 **References**
- 1150
- 1151 Angevine, W. M., Brioude, J., McKeen, S., Holloway, J. S., Lerner, B. M., Goldstein, A. H., Guha,
- 1152 A., Andrews, A., Nowak, J. B., Evan, S., Fischer, M. L., Gilman, J. B., and Bon, D.: Pollutant
- 1153 transport among California regions, Journal of Geophysical Research: Atmospheres, 118, 6750-
- 1154 6763, <u>https://doi.org/10.1002/jgrd.50490</u>, 2013.
- Annual Monitoring Network Plan, Clark County Department of Environment and Sustainability,1-77, 2022.
- 1157 Apel, E. C., Hornbrook, R. S., Hills, A. J., Blake, N. J., Barth, M. C., Weinheimer, A., Cantrell, C.,
- 1158 Rutledge, S. A., Basarab, B., Crawford, J., Diskin, G., Homeyer, C. R., Campos, T., Flocke, F., Fried,
- 1159 A., Blake, D. R., Brune, W., Pollack, I., Peischl, J., Ryerson, T., Wennberg, P. O., Crounse, J. D.,
- 1160 Wisthaler, A., Mikoviny, T., Huey, G., Heikes, B., O'Sullivan, D., and Riemer, D. D.: Upper
- 1161 tropospheric ozone production from lightning NOx-impacted convection: Smoke ingestion case
- study from the DC3 campaign, Journal of Geophysical Research: Atmospheres, 120, 2505-2523,
- 1163 <u>https://doi.org/10.1002/2014JD022121</u>, 2015.
- 1164 Arata, C., Misztal, P. K., Tian, Y., Lunderberg, D. M., Kristensen, K., Novoselac, A., Vance, M. E.,
- 1165 Farmer, D. K., Nazaroff, W. W., and Goldstein, A. H.: Volatile organic compound emissions
- 1166 during HOMEChem, Indoor Air, 31, 2099-2117, <u>https://doi.org/10.1111/ina.12906</u>, 2021.
- Arnold, S. T., Viggiano, A. A., and Morris, R. A.: Rate Constants and Product Branching Fractions
  for the Reactions of H3O+ and NO+ with C2–C12 Alkanes, The Journal of Physical Chemistry A,
  102, 8881-8887, 10.1021/jp9815457, 1998.
- Breitenlechner, M., Fischer, L., Hainer, M., Heinritzi, M., Curtius, J., and Hansel, A.: PTR3: An
  Instrument for Studying the Lifecycle of Reactive Organic Carbon in the Atmosphere, Analytical
- 1172 Chemistry, 89, 5824-5831, 10.1021/acs.analchem.6b05110, 2017.
- 1173 Buhr, K., van Ruth, S., and Delahunty, C.: Analysis of volatile flavour compounds by Proton
- 1174 Transfer Reaction-Mass Spectrometry: fragmentation patterns and discrimination between
- isobaric and isomeric compounds, International Journal of Mass Spectrometry, 221, 1-7,
- 1176 <u>https://doi.org/10.1016/S1387-3806(02)00896-5</u>, 2002.
- 1177 Burkholder, J. B., Sander, S. P., Abbatt, J., Barker, J. R., Cappa, C., Crounse, J. D., Dibble, T. S.,
- 1178 Huie, R. E., Kolb, C. E., Kurylo, M. J., Orkin, V. L., Percival, C. J., Wilmouth, D. M., and Wind, P. H.:
- 1179 Chemical Kinetics and Photochemical Data for Use in Atmospheric Studies, Evaluation No. 19,
- 1180 Jet Propulsion Laboratory, Pasadena, CA, 2019.
- 1181 Calfapietra, C., Fares, S., Manes, F., Morani, A., Sgrigna, G., and Loreto, F.: Role of Biogenic
- 1182 Volatile Organic Compounds (BVOC) emitted by urban trees on ozone concentration in cities: A
- review, Environmental Pollution, 183, 71-80, <u>https://doi.org/10.1016/j.envpol.2013.03.012</u>,
- 1184 2013.





- 1185 Cao, C., Gentner, D. R., Commane, R., Toledo-Crow, R., Schiferl, L. D., and Mak, J. E.: Policy-
- 1186 Related Gains in Urban Air Quality May Be Offset by Increased Emissions in a Warming Climate,
- 1187 Environmental Science & Technology, 10.1021/acs.est.2c05904, 2023.
- Choi, J., Henze, D. K., Cao, H., Nowlan, C. R., González Abad, G., Kwon, H.-A., Lee, H.-M., Oak, Y.
  J., Park, R. J., Bates, K. H., Maasakkers, J. D., Wisthaler, A., and Weinheimer, A. J.: An Inversion
- 1190 Framework for Optimizing Non-Methane VOC Emissions Using Remote Sensing and Airborne
- 1191 Observations in Northeast Asia During the KORUS-AQ Field Campaign, Journal of Geophysical
- 1192 Research: Atmospheres, 127, e2021JD035844, <u>https://doi.org/10.1029/2021JD035844</u>, 2022.
- 1193 Claflin, M. S., Pagonis, D., Finewax, Z., Handschy, A. V., Day, D. A., Brown, W. L., Jayne, J. T.,
- 1194 Worsnop, D. R., Jimenez, J. L., Ziemann, P. J., de Gouw, J., and Lerner, B. M.: An in situ gas
- 1195 chromatograph with automatic detector switching between PTR- and EI-TOF-MS: isomer-
- 1196 resolved measurements of indoor air, Atmos. Meas. Tech., 14, 133-152, 10.5194/amt-14-133-1197 2021, 2021.
- 1198 Coggon, M. M., McDonald, B. C., Vlasenko, A., Veres, P. R., Bernard, F., Koss, A. R., Yuan, B.,
- 1199 Gilman, J. B., Peischl, J., Aikin, K. C., DuRant, J., Warneke, C., Li, S.-M., and de Gouw, J. A.:
- 1200 Diurnal Variability and Emission Pattern of Decamethylcyclopentasiloxane (D5) from the
- 1201 Application of Personal Care Products in Two North American Cities, Environmental Science &
- 1202 Technology, 52, 5610-5618, 10.1021/acs.est.8b00506, 2018.
- 1203 Coggon, M. M., Gkatzelis, G. I., McDonald, B. C., Gilman, J. B., Schwantes, R. H., Abuhassan, N.,
- 1204 Aikin, K. C., Arend, M. F., Berkoff, T. A., Brown, S. S., Campos, T. L., Dickerson, R. R., Gronoff, G.,
- 1205 Hurley, J. F., Isaacman-VanWertz, G., Koss, A. R., Li, M., McKeen, S. A., Moshary, F., Peischl, J.,
- 1206 Pospisilova, V., Ren, X., Wilson, A., Wu, Y., Trainer, M., and Warneke, C.: Volatile chemical
- 1207 product emissions enhance ozone and modulate urban chemistry, Proceedings of the National
- 1208 Academy of Sciences, 118, e2026653118, doi:10.1073/pnas.2026653118, 2021.
- 1209 Colman, J. J., Swanson, A. L., Meinardi, S., Sive, B. C., Blake, D. R., and Rowland, F. S.:
- 1210 Description of the Analysis of a Wide Range of Volatile Organic Compounds in Whole Air
- 1211 Samples Collected during PEM-Tropics A and B, Analytical Chemistry, 73, 3723-3731,
- 1212 10.1021/ac010027g, 2001.
- Ernle, L., Ringsdorf, M. A., and Williams, J.: Influence of ozone and humidity on PTR-MS and GCMS VOC measurements with and without a Na2S2O3 ozone scrubber, Atmos. Meas. Tech., 16,
  1179-1194, 10.5194/amt-16-1179-2023, 2023.
- 1216 Gkatzelis, G. I., Coggon, M. M., McDonald, B. C., Peischl, J., Aikin, K. C., Gilman, J. B., Trainer, M.,
- 1217 and Warneke, C.: Identifying Volatile Chemical Product Tracer Compounds in U.S. Cities,
- 1218 Environmental Science & Technology, 55, 188-199, 10.1021/acs.est.0c05467, 2021a.
- 1219 Gkatzelis, G. I., Coggon, M. M., McDonald, B. C., Peischl, J., Gilman, J. B., Aikin, K. C., Robinson,
- 1220 M. A., Canonaco, F., Prevot, A. S. H., Trainer, M., and Warneke, C.: Observations Confirm that





- Volatile Chemical Products Are a Major Source of Petrochemical Emissions in U.S. Cities,
- 1222 Environmental Science & Technology, 55, 4332-4343, 10.1021/acs.est.0c05471, 2021b.
- 1223 Gkatzelis, G. I., Coggon, M. M., Stockwell, C. E., Hornbrook, R. S., Allen, H., Apel, E. C., Bela, M.
- 1224 M., Blake, D. R., Bourgeois, I., Brown, S. S., Campuzano-Jost, P., St. Clair, J. M., Crawford, J. H.,
- 1225 Crounse, J., Day, D. A., DiGangi, J. P., Diskin, G. S., Fried, A., Gilman, J. B., Guo, H., Hair, J. W.,
- 1226 Halliday, H. S., Hanisco, T. F., Hannun, R. A., Hills, A., Huey, L. G., Jimenez, J. L., Katich, J. M.,
- 1227 Lamplugh, A., Lee, Y. R., Liao, J., Neuman, J. A., Nowak, J. B., Pagonis, D., Peischl, J., Perring, A.
- 1228 E., Piel, F., Rickly, P. S., Robinson, M. A., Rollins, A. W., Ryerson, T. B., Schueneman, M. K.,
- 1229 Schwantes, R. H., Schwarz, J. P., Sekimoto, K., Selimovic, V., Shingler, T., Tanner, D. J., Tomsche,
- 1230 L., Vasquez, K. T., Veres, P., Washenfelder, R. A., Weibring, P., Wennberg, P. O., Wisthaler, A.,
- 1231 Wolfe, G. M., Womack, C. C., Xu, L., Yokelson, R., and Warneke, C.: Parameterizations of US
- 1232 wildfire and prescribed fire emission ratios and emission factors based on FIREX-AQ aircraft
- 1233 measurements, Atmos. Chem. Phys., In prep, 2022.
- Gueneron, M., Erickson, M. H., VanderSchelden, G. S., and Jobson, B. T.: PTR-MS fragmentation
   patterns of gasoline hydrocarbons, International Journal of Mass Spectrometry, 379, 97-109,
   <u>https://doi.org/10.1016/j.ijms.2015.01.001</u>, 2015.
- Guenther, A. B., Jiang, X., Heald, C. L., Sakulyanontvittaya, T., Duhl, T., Emmons, L. K., and Wang,
  X.: The Model of Emissions of Gases and Aerosols from Nature version 2.1 (MEGAN2.1): an
  extended and updated framework for modeling biogenic emissions, Geosci. Model Dev., 5,
  1471-1492, 10.5194/gmd-5-1471-2012, 2012.
- Herndon, S. C., Jayne, J. T., Zahniser, M. S., Worsnop, D. R., Knighton, B., Alwine, E., Lamb, B. K.,
  Zavala, M., Nelson, D. D., McManus, J. B., Shorter, J. H., Canagaratna, M. R., Onasch, T. B., and
  Kolb, C. E.: Characterization of urban pollutant emission fluxes and ambient concentration
  distributions using a mobile laboratory with rapid response instrumentation, Faraday
  Discussions, 130, 327-339, 10.1039/B500411J, 2005.
- Holzinger, R., Acton, W. J. F., Bloss, W. J., Breitenlechner, M., Crilley, L. R., Dusanter, S., Gonin,
  M., Gros, V., Keutsch, F. N., Kiendler-Scharr, A., Kramer, L. J., Krechmer, J. E., Languille, B.,
  Locoge, N., Lopez-Hilfiker, F., Materić, D., Moreno, S., Nemitz, E., Quéléver, L. L. J., Sarda Esteve,
  R., Sauvage, S., Schallhart, S., Sommariva, R., Tillmann, R., Wedel, S., Worton, D. R., Xu, K., and
  Zaytsev, A.: Validity and limitations of simple reaction kinetics to calculate concentrations of
  organic compounds from ion counts in PTR-MS, Atmos. Meas. Tech., 12, 6193-6208,
  10.5194/amt-12-6193-2019, 2019.
- Inomata, S., and Tanimoto, H.: A deuterium-labeling study on the reproduction of hydronium
  ions in the PTR-MS detection of ethanol, International Journal of Mass Spectrometry, 285, 9599, <u>https://doi.org/10.1016/j.ijms.2009.05.001</u>, 2009.
- 1256 Inomata, S., Tanimoto, H., Kato, S., Suthawaree, J., Kanaya, Y., Pochanart, P., Liu, Y., and Wang,
- 1257 Z.: PTR-MS measurements of non-methane volatile organic compounds during an intensive field





- 1258 campaign at the summit of Mount Tai, China, in June 2006, Atmos. Chem. Phys., 10, 7085-7099,
  1259 10.5194/acp-10-7085-2010, 2010.
- 1260 Jobson, B. T., Alexander, M. L., Maupin, G. D., and Muntean, G. G.: On-line analysis of organic
- 1261 compounds in diesel exhaust using a proton transfer reaction mass spectrometer (PTR-MS),
- 1262 International Journal of Mass Spectrometry, 245, 78-89,
- 1263 <u>https://doi.org/10.1016/j.ijms.2005.05.009</u>, 2005.
- 1264 Kari, E., Miettinen, P., Yli-Pirilä, P., Virtanen, A., and Faiola, C. L.: PTR-ToF-MS product ion
- 1265 distributions and humidity-dependence of biogenic volatile organic compounds, International
- 1266 Journal of Mass Spectrometry, 430, 87-97, <u>https://doi.org/10.1016/j.ijms.2018.05.003</u>, 2018.
- 1267 Karl, T., Hansel, A., Cappellin, L., Kaser, L., Herdlinger-Blatt, I., and Jud, W.: Selective
- 1268 measurements of isoprene and 2-methyl-3-buten-2-ol based on NO<sup>+</sup> ionization
- 1269 mass spectrometry, Atmos. Chem. Phys., 12, 11877-11884, 10.5194/acp-12-11877-2012, 2012.

Karl, T., Striednig, M., Graus, M., Hammerle, A., and Wohlfahrt, G.: Urban flux measurements
reveal a large pool of oxygenated volatile organic compound emissions, Proc. Natl. Acad. Sci.
U.S.A., 115, 1186-1191, 10.1073/pnas.1714715115, 2018.

Kaser, L., Karl, T., Schnitzhofer, R., Graus, M., Herdlinger-Blatt, I. S., DiGangi, J. P., Sive, B.,
Turnipseed, A., Hornbrook, R. S., Zheng, W., Flocke, F. M., Guenther, A., Keutsch, F. N., Apel, E.,
and Hansel, A.: Comparison of different real time VOC measurement techniques in a ponderosa
pine forest, Atmos. Chem. Phys., 13, 2893-2906, 10.5194/acp-13-2893-2013, 2013.

1277 Khare, P., Machesky, J., Soto, R., He, M., Presto, A. A., and Gentner, D. R.: Asphalt-related

- emissions are a major missing nontraditional source of secondary organic aerosol precursors,
  Science Advances, 6, eabb9785, doi:10.1126/sciadv.abb9785, 2020.
- 1280 Khare, P., Krechmer, J. E., Machesky, J. E., Hass-Mitchell, T., Cao, C., Wang, J., Majluf, F., Lopez-
- 1281 Hilfiker, F., Malek, S., Wang, W., Seltzer, K., Pye, H. O. T., Commane, R., McDonald, B. C.,
- 1282 Toledo-Crow, R., Mak, J. E., and Gentner, D. R.: Ammonium adduct chemical ionization to
- 1283 investigate anthropogenic oxygenated gas-phase organic compounds in urban air, Atmos.
- 1284 Chem. Phys., 22, 14377-14399, 10.5194/acp-22-14377-2022, 2022.
- Kilgour, D., Novak, G., and Bertram, T.: Observations of Biotic and Abiotic Marine Volatile
  Organic Compounds Emitted from Coastal Seawater, December 01, 2021, 2021.

1287 Kim, S.-W., McDonald, B. C., Seo, S., Kim, K.-M., and Trainer, M.: Understanding the Paths of
1288 Surface Ozone Abatement in the Los Angeles Basin, Journal of Geophysical Research:
1289 Atmospheres, 127, e2021JD035606, https://doi.org/10.1029/2021JD035606, 2022.

- 1290 Klein, F., Platt, S. M., Farren, N. J., Detournay, A., Bruns, E. A., Bozzetti, C., Daellenbach, K. R.,
- 1291 Kilic, D., Kumar, N. K., Pieber, S. M., Slowik, J. G., Temime-Roussel, B., Marchand, N., Hamilton,
- 1292 J. F., Baltensperger, U., Prévôt, A. S. H., and El Haddad, I.: Characterization of Gas-Phase





- Organics Using Proton Transfer Reaction Time-of-Flight Mass Spectrometry: Cooking Emissions,
   Environmental Science & Technology, 50, 1243-1250, 10.1021/acs.est.5b04618, 2016.
- 1295 Koss, A., Yuan, B., Warneke, C., Gilman, J. B., Lerner, B. M., Veres, P. R., Peischl, J., Eilerman, S., 1296 Wild, R., Brown, S. S., Thompson, C. R., Ryerson, T., Hanisco, T., Wolfe, G. M., Clair, J. M. S.,
- 1297 Thayer, M., Keutsch, F. N., Murphy, S., and de Gouw, J.: Observations of VOC emissions and
- 1298 photochemical products over US oil- and gas-producing regions using high-resolution H3O+
- 1299 CIMS (PTR-ToF-MS), Atmos. Meas. Tech., 10, 2941-2968, 10.5194/amt-10-2941-2017, 2017.
- Koss, A. R., Sekimoto, K., Gilman, J. B., Selimovic, V., Coggon, M. M., Zarzana, K. J., Yuan, B.,
  Lerner, B. M., Brown, S. S., Jimenez, J. L., Krechmer, J., Roberts, J. M., Warneke, C., Yokelson, R.
- 1302 J., and de Gouw, J.: Non-methane organic gas emissions from biomass burning: identification,
- 1303 quantification, and emission factors from PTR-ToF during the FIREX 2016 laboratory
- 1304 experiment, Atmos. Chem. Phys., 18, 3299-3319, 10.5194/acp-18-3299-2018, 2018.
- Krechmer, J., Lopez-Hilfiker, F., Koss, A., Hutterli, M., Stoermer, C., Deming, B., Kimmel, J.,
  Warneke, C., Holzinger, R., Jayne, J., Worsnop, D., Fuhrer, K., Gonin, M., and de Gouw, J.:
  Evaluation of a New Reagent-Ion Source and Focusing Ion–Molecule Reactor for Use in ProtonTransfer-Reaction Mass Spectrometry, Analytical Chemistry, 90, 12011-12018,
  10.1021/acs.analchem.8b02641, 2018.
- 1310 Lerner, B. M., Gilman, J. B., Aikin, K. C., Atlas, E. L., Goldan, P. D., Graus, M., Hendershot, R.,
- 1311 Isaacman-VanWertz, G. A., Koss, A., Kuster, W. C., Lueb, R. A., McLaughlin, R. J., Peischl, J.,
- 1312 Sueper, D., Ryerson, T. B., Tokarek, T. W., Warneke, C., Yuan, B., and de Gouw, J. A.: An
- improved, automated whole air sampler and gas chromatography mass spectrometry analysis
  system for volatile organic compounds in the atmosphere, Atmos. Meas. Tech., 10, 291-313,
- 1315 10.5194/amt-10-291-2017, 2017.
- 1316 McDonald, B. C., de Gouw, J. A., Gilman, J. B., Jathar, S. H., Akherati, A., Cappa, C. D., Jimenez, J.
- 1317 L., Lee-Taylor, J., Hayes, P. L., McKeen, S. A., Cui, Y. Y., Kim, S. W., Gentner, D. R., Isaacman-
- 1318 VanWertz, G., Goldstein, A. H., Harley, R. A., Frost, G. J., Roberts, J. M., Ryerson, T. B., and
- 1319 Trainer, M.: Volatile chemical products emerging as largest petrochemical source of urban
- 1320 organic emissions, Science, 359, 760-764, 10.1126/science.aaq0524, 2018.
- 1321 Müller, M., Mikoviny, T., Feil, S., Haidacher, S., Hanel, G., Hartungen, E., Jordan, A., Märk, L.,
- 1322 Mutschlechner, P., Schottkowsky, R., Sulzer, P., Crawford, J. H., and Wisthaler, A.: A compact
- 1323 PTR-ToF-MS instrument for airborne measurements of volatile organic compounds at high
- 1324 spatiotemporal resolution, Atmos. Meas. Tech., 7, 3763-3772, 10.5194/amt-7-3763-2014, 2014.
- Pagonis, D., Sekimoto, K., and de Gouw, J.: A Library of Proton-Transfer Reactions of H3O+ Ions
  Used for Trace Gas Detection, Journal of The American Society for Mass Spectrometry, 30,
  1330-1335, 10.1007/s13361-019-02209-3, 2019.
- Peng, Y., Mouat, A. P., Hu, Y., Li, M., McDonald, B. C., and Kaiser, J.: Source appointment ofvolatile organic compounds and evaluation of anthropogenic monoterpene emission estimates





- 1330 in Atlanta, Georgia, Atmospheric Environment, 288, 119324,
- 1331 <u>https://doi.org/10.1016/j.atmosenv.2022.119324</u>, 2022.
- 1332 Pfannerstill, E. Y., Wang, N., Edtbauer, A., Bourtsoukidis, E., Crowley, J. N., Dienhart, D., Eger, P.
- G., Ernle, L., Fischer, H., Hottmann, B., Paris, J. D., Stönner, C., Tadic, I., Walter, D., Lelieveld, J.,
  and Williams, J.: Shipborne measurements of total OH reactivity around the Arabian Peninsula
  and its role in ozone chemistry, Atmos. Chem. Phys., 19, 11501-11523, 10.5194/acp-19-11501-
- 1336 2019, 2019.
- 1337 Pfannerstill, E. Y., Arata, C., Zhu, Q., Schulze, B. C., Woods, R., Seinfeld, J. H., Bucholtz, A.,
- 1338 Cohen, R. C., and Goldstein, A. H.: Volatile organic compound fluxes in the San Joaquin Valley –
- 1339 spatial distribution, source attribution, and inventory comparison, EGUsphere, 2023, 1-42,
- 1340 10.5194/egusphere-2023-723, 2023.
- Riva, M., Rantala, P., Krechmer, J. E., Peräkylä, O., Zhang, Y., Heikkinen, L., Garmash, O., Yan, C.,
  Kulmala, M., Worsnop, D., and Ehn, M.: Evaluating the performance of five different chemical
  ionization techniques for detecting gaseous oxygenated organic species, Atmos. Meas. Tech.,
  12, 2403-2421, 10.5194/amt-12-2403-2019, 2019.
- 1345 Rivera-Rios, J. C., Nguyen, T. B., Crounse, J. D., Jud, W., St. Clair, J. M., Mikoviny, T., Gilman, J. B.,
- 1346 Lerner, B. M., Kaiser, J. B., de Gouw, J., Wisthaler, A., Hansel, A., Wennberg, P. O., Seinfeld, J. H.,
- and Keutsch, F. N.: Conversion of hydroperoxides to carbonyls in field and laboratory
- 1348 instrumentation: Observational bias in diagnosing pristine versus anthropogenically controlled
- 1349 atmospheric chemistry, Geophysical Research Letters, 41, 8645-8651,
- 1350 <u>https://doi.org/10.1002/2014GL061919</u>, 2014.
- 1351 Romano, A., and Hanna, G. B.: Identification and quantification of VOCs by proton transfer
- 1352 reaction time of flight mass spectrometry: An experimental workflow for the optimization of
- 1353 specificity, sensitivity, and accuracy, Journal of Mass Spectrometry, 53, 287-295,
- 1354 <u>https://doi.org/10.1002/jms.4063</u>, 2018.
- 1355 Ryerson, T. B., Andrews, A. E., Angevine, W. M., Bates, T. S., Brock, C. A., Cairns, B., Cohen, R. C.,
- 1356 Cooper, O. R., de Gouw, J. A., Fehsenfeld, F. C., Ferrare, R. A., Fischer, M. L., Flagan, R. C.,
- 1357 Goldstein, A. H., Hair, J. W., Hardesty, R. M., Hostetler, C. A., Jimenez, J. L., Langford, A. O.,
- 1358 McCauley, E., McKeen, S. A., Molina, L. T., Nenes, A., Oltmans, S. J., Parrish, D. D., Pederson, J.
- 1359 R., Pierce, R. B., Prather, K., Quinn, P. K., Seinfeld, J. H., Senff, C. J., Sorooshian, A., Stutz, J.,
- 1360 Surratt, J. D., Trainer, M., Volkamer, R., Williams, E. J., and Wofsy, S. C.: The 2010 California
- 1361 Research at the Nexus of Air Quality and Climate Change (CalNex) field study, Journal of
- 1362 Geophysical Research: Atmospheres, 118, 5830-5866, <u>https://doi.org/10.1002/jgrd.50331</u>,
  1363 2013.
- 1364 Schauer, J. J., Kleeman, M. J., Cass, G. R., and Simoneit, B. R. T.: Measurement of Emissions from
- Air Pollution Sources. 1. C1 through C29 Organic Compounds from Meat Charbroiling,
- 1366 Environmental Science & Technology, 33, 1566-1577, 10.1021/es980076j, 1999.





- 1367 Sekimoto, K., Li, S.-M., Yuan, B., Koss, A., Coggon, M., Warneke, C., and de Gouw, J.: Calculation
- 1368 of the sensitivity of proton-transfer-reaction mass spectrometry (PTR-MS) for organic trace
- 1369 gases using molecular properties, International Journal of Mass Spectrometry, 421, 71-94,
- 1370 <u>https://doi.org/10.1016/j.ijms.2017.04.006</u>, 2017.
- 1371 Simpson, I. J., Blake, N. J., Barletta, B., Diskin, G. S., Fuelberg, H. E., Gorham, K., Huey, L. G.,
- 1372 Meinardi, S., Rowland, F. S., Vay, S. A., Weinheimer, A. J., Yang, M., and Blake, D. R.:
- 1373 Characterization of trace gases measured over Alberta oil sands mining operations: 76
- 1374 speciated C<sub>2</sub>–C<sub>10</sub> volatile organic compounds (VOCs),
- 1375 CO<sub>2</sub>, CH<sub>4</sub>, CO, NO, NO<sub>2</sub>, NO<sub>y</sub>,
- 1376 O<sub>3</sub> and SO<sub>2</sub>, Atmos. Chem. Phys., 10, 11931-11954, 10.5194/acp-10-
- 1377 11931-2010, 2010.
- 1378 Simpson, I. J., Blake, D. R., Blake, N. J., Meinardi, S., Barletta, B., Hughes, S. C., Fleming, L. T.,
- 1379 Crawford, J. H., Diskin, G. S., Emmons, L. K., Fried, A., Guo, H., Peterson, D. A., Wisthaler, A.,
- 1380 Woo, J.-H., Barré, J., Gaubert, B., Kim, J., Kim, M. J., Kim, Y., Knote, C., Mikoviny, T., Pusede, S.
- 1381 E., Schroeder, J. R., Wang, Y., Wennberg, P. O., and Zeng, L.: Characterization, sources and
- 1382 reactivity of volatile organic compounds (VOCs) in Seoul and surrounding regions during
- 1383 KORUS-AQ, Elementa: Science of the Anthropocene, 8, 10.1525/elementa.434, 2020.

Spanel, P., and Smith, D.: SIFT studies of the reactions of H3O+, NO+ and O2+ with a series of
 alcohols, International Journal of Mass Spectrometry and Ion Processes, 167-168, 375-388,
 <u>https://doi.org/10.1016/S0168-1176(97)00085-2</u>, 1997.

- 1387 Stark, H., Yatavelli, R. L. N., Thompson, S. L., Kimmel, J. R., Cubison, M. J., Chhabra, P. S.,
- 1388 Canagaratna, M. R., Jayne, J. T., Worsnop, D. R., and Jimenez, J. L.: Methods to extract
- 1389 molecular and bulk chemical information from series of complex mass spectra with limited
- 1390 mass resolution, International Journal of Mass Spectrometry, 389, 26-38,
- 1391 <u>https://doi.org/10.1016/j.ijms.2015.08.011</u>, 2015.
- 1392 Stockwell, C. E., Coggon, M. M., Gkatzelis, G. I., Ortega, J., McDonald, B. C., Peischl, J., Aikin, K.,
- 1393 Gilman, J. B., Trainer, M., and Warneke, C.: Volatile organic compound emissions from solvent-
- and water-borne coatings compositional differences and tracer compound identifications,
- 1395 Atmos. Chem. Phys., 21, 6005-6022, 10.5194/acp-21-6005-2021, 2021.
- Tani, A.: Fragmentation and Reaction Rate Constants of Terpenoids Determined by Proton
   Transfer Reaction-mass Spectrometry, Environmental Control in Biology, 51, 23-29,
   10.2525/ecb.51.23, 2013.
- Tomsche, L., Piel, F., Mikoviny, T., Nielsen, C. J., Guo, H., Campuzano-Jost, P., Nault, B. A.,
  Schueneman, M. K., Jimenez, J. L., Halliday, H., Diskin, G., DiGangi, J. P., Nowak, J. B., Wiggins, E.
  B., Gargulinski, E., Soja, A. J., and Wisthaler, A.: Measurement report: Emission factors of NH3 and NHx for wildfires and agricultural fires in the United States, Atmos. Chem. Phys., 23, 2331-
- 1403 2343, 10.5194/acp-23-2331-2023, 2023.





- 1404 Vermeuel, M. P., Novak, G. A., Kilgour, D. B., Claflin, M. S., Lerner, B. M., Trowbridge, A. M.,
- 1405 Thom, J., Cleary, P. A., Desai, A. R., and Bertram, T. H.: Observations of biogenic volatile organic
- 1406 compounds over a mixed temperate forest during the summer to autumn transition,
- 1407 EGUsphere, 2022, 1-44, 10.5194/egusphere-2022-1015, 2022.
- 1408 Warneke, C., de Gouw, J. A., Kuster, W. C., Goldan, P. D., and Fall, R.: Validation of Atmospheric
- 1409 VOC Measurements by Proton-Transfer- Reaction Mass Spectrometry Using a Gas-
- 1410 Chromatographic Preseparation Method, Environmental Science & Technology, 37, 2494-2501,
- 1411 10.1021/es026266i, 2003.
- 1412 Warneke, C., de Gouw, J. A., Holloway, J. S., Peischl, J., Ryerson, T. B., Atlas, E., Blake, D.,
- 1413 Trainer, M., and Parrish, D. D.: Multiyear trends in volatile organic compounds in Los Angeles,
- 1414 California: Five decades of decreasing emissions, J. Geophs. Res., 117, 1-10,
- 1415 10.1029/2012jd017899, 2012.
- 1416 Warneke, C., Geiger, F., Edwards, P. M., Dube, W., Pétron, G., Kofler, J., Zahn, A., Brown, S. S.,
- 1417 Graus, M., Gilman, J. B., Lerner, B. M., Peischl, J., Ryerson, T. B., de Gouw, J. A., and Roberts, J.
- 1418 M.: Volatile organic compound emissions from the oil and natural gas industry in the Uintah
- 1419 Basin, Utah: oil and gas well pad emissions compared to ambient air composition, Atmos.
- 1420 Chem. Phys., 14, 10977-10988, 10.5194/acp-14-10977-2014, 2014.
- 1421 Warneke, C., Schwarz, J. P., Dibb, J., Kalashnikova, O., Frost, G., Al-Saad, J., Brown, S. S., Brewer,
- 1422 W. A., Soja, A., Seidel, F. C., Washenfelder, R. A., Wiggins, E. B., Moore, R. H., Anderson, B. E.,
- 1423 Jordan, C., Yacovitch, T. I., Herndon, S. C., Liu, S., Kuwayama, T., Jaffe, D., Johnston, N.,
- 1424 Selimovic, V., Yokelson, R., Giles, D. M., Holben, B. N., Goloub, P., Popovici, I., Trainer, M.,
- 1425 Kumar, A., Pierce, R. B., Fahey, D., Roberts, J., Gargulinski, E. M., Peterson, D. A., Ye, X., Thapa,
- 1426 L. H., Saide, P. E., Fite, C. H., Holmes, C. D., Wang, S., Coggon, M. M., Decker, Z. C. J., Stockwell,
- 1427 C. E., Xu, L., Gkatzelis, G., Aikin, K., Lefer, B., Kaspari, J., Griffin, D., Zeng, L., Weber, R., Hastings,
- 1428 M., Chai, J., Wolfe, G. M., Hanisco, T. F., Liao, J., Campuzano Jost, P., Guo, H., Jimenez, J. L.,
- 1429 Crawford, J., and Team, T. F.-A. S.: Fire Influence on Regional to Global Environments and Air
- 1430 Quality (FIREX-AQ), Journal of Geophysical Research: Atmospheres, 128, e2022JD037758,
- 1431 <u>https://doi.org/10.1029/2022JD037758</u>, 2023.
- 1432 Wernis, R. A., Kreisberg, N. M., Weber, R. J., Drozd, G. T., and Goldstein, A. H.: Source
- 1433 apportionment of VOCs, IVOCs and SVOCs by positive matrix factorization in suburban
- 1434 Livermore, California, Atmos. Chem. Phys., 22, 14987-15019, 10.5194/acp-22-14987-2022,
  1435 2022.
- 1436 Xu, L., Coggon, M. M., Stockwell, C. E., Gilman, J. B., Robinson, M. A., Breitenlechner, M.,
- 1437 Lamplugh, A., Crounse, J. D., Wennberg, P. O., Neuman, J. A., Novak, G. A., Veres, P. R., Brown,
- 1438 S. S., and Warneke, C.: Chemical ionization mass spectrometry utilizing ammonium ions (NH4+
- 1439 CIMS) for measurements of organic compounds in the atmosphere, Atmos. Meas. Tech., 15,
- 1440 7353-7373, 10.5194/amt-15-7353-2022, 2022.





- 1441 Yacovitch, T. I., Herndon, S. C., Pétron, G., Kofler, J., Lyon, D., Zahniser, M. S., and Kolb, C. E.:
- 1442 Mobile Laboratory Observations of Methane Emissions in the Barnett Shale Region,
- 1443 Environmental Science & Technology, 49, 7889-7895, 10.1021/es506352j, 2015.
- 1444 Yuan, B., Koss, A., Warneke, C., Gilman, J. B., Lerner, B. M., Stark, H., and de Gouw, J. A.: A high-
- 1445 resolution time-of-flight chemical ionization mass spectrometer utilizing hydronium ions (H3O+
- 1446 ToF-CIMS) for measurements of volatile organic compounds in the atmosphere, Atmos. Meas.
- 1447 Tech., 9, 2735-2752, 10.5194/amt-9-2735-2016, 2016.
- 1448 Yuan, B., Koss, A. R., Warneke, C., Coggon, M., Sekimoto, K., and de Gouw, J. A.: Proton-
- 1449 Transfer-Reaction Mass Spectrometry: Applications in Atmospheric Sciences, Chem. Rev., 117,
- 1450 13187-13229, 10.1021/acs.chemrev.7b00325, 2017.
- 1451