



1 **Exploring the joint probability of precipitation and soil moisture** 2 **over Europe using copulas**

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7 **Abstract.** The joint probability of precipitation and soil moisture is here investigated over Europe
8 with the goal to extrapolated meaningful insights on the potential joint use of these variables for
9 the detection of agricultural droughts within a probabilistic modeling framework. The use of
10 copulas is explored as a parametric approach often used in hydrological studies for the analysis of
11 bivariate distributions. The analysis is performed for the period 1996-2020 on the ERA5
12 precipitation and LISFLOOD soil moisture datasets, both available as part of the Copernicus
13 European Drought Observatory. The results show an overall good correlation between the
14 empirical frequency series derived from the two datasets (Kendall's $\tau = 0.42 \pm 0.1$), but also clear
15 spatial patterns in the tail-dependence derived with both non-parametric and parametric
16 approaches. About half of the domain shows symmetric tail-dependences, well reproduced by the
17 Student-t copula, whereas the rest of the domain is almost equally split between low and high tail-
18 dependences (modeled with the Gumbel family of copulas). These spatial patterns are reasonably
19 reproduced by a random forest classifier, suggesting that this outcome is not driven by chance.
20 This study stresses how a joint use of precipitation and soil moisture for agriculture drought
21 characterization may be more beneficial in areas with strong low tail-dependence, such as southern
22 France, northern UK, northern Germany, and Denmark in this study, and how this behavior should
23 be carefully considered in drought studies.



24 **1. Introduction**

25 Agricultural drought, defined as a condition of unusually high precipitation shortages and/or soil
26 water deficits causing adverse effects on crop yields and production (Panu and Sharma, 2002), is
27 probably the most recognized of the four main drought types (Wilhite and Glantz, 1985) due to
28 the more direct and easier to understand impacts (Mishra and Singh, 2010). The scientific literature
29 on agricultural drought has produced a very large number of indices (WMO and GWP, 2016), with
30 the aim of reproducing the temporal dynamic and the effects of crop water deficit through a
31 combination of climatic observations, hydrological modeling, and remote sensing data (Zargar et
32 al., 2011).

33 The difficulty in capturing the multi-facet nature of agricultural drought events across the
34 world with a single index (Sivakumar et al., 2011) is confirmed by the absence of consensus in the
35 scientific literature on the most reliable agricultural drought index. However, despite the large
36 range of available indices, some common characteristics can be identified, such as the focus on
37 some proxy variables of plant water availability – through soil moisture (Dutra et al., 2008), actual
38 evapotranspiration (Anderson et al., 2011) or basic meteorological information (Vicente-Serrano
39 et al., 2010) – and the need to account for deviations from long-term conditions.

40 Meteorological drought indicators computed on appropriate aggregation time scales
41 (McKee et al., 1993; Vicente-Serrano et al., 2010) have demonstrated a good capability to
42 represent agricultural drought conditions in several case studies (e.g., Bachmair et al., 2018;
43 Mohammed et al., 2022; Tian et al., 2018). They have been successfully integrated in a number of
44 operational drought monitoring systems, thanks to their minimal input data requirements and ease
45 of use. Among those indices, the Standardized Precipitation Index (SPI, McKee et al., 1993)
46 computed on short- to medium-aggregation periods (i.e., SPI -3 and -6) is often adopted as a
47 suitable proxy variable for agricultural droughts (WMO, 2012).

48 As highlighted by Sheffield and Wood (2007), simplified indices for drought monitoring,
49 such as the Palmer Drought Severity index (PDSI; Palmer, 1965) or the previously mentioned
50 meteorological indicators, have been slowly integrated with indices directly based on modeled soil
51 moisture data, thanks to the increasing availability worldwide of physically-based hydrological
52 models. Soil moisture percentile, or similarly standardized quantities, are often used for this scope
53 (Mo and Lettenmeier, 2013; Xia et al., 2014). The ever-growing records of remote sensing-based



54 estimates of soil moisture are becoming an additional data source to support the development of
55 dedicated soil moisture drought indices (Cammalleri et al., 2017; Carrão et al., 2016).

56 In the context of agricultural drought, an overall good agreement between SPI and soil
57 moisture indices has been demonstrated over a large range of agricultural practices, crop types and
58 climatic conditions. Halwatura et al. (2017) showed how SPI-3 represents a good approximation
59 of modeled soil moisture over three different climatic regions in eastern Australia. Sims et al.
60 (2002) found high correlation between short-term precipitation deficit and soil moisture variations
61 in North Carolina, while Ji and Peters (2003) highlighted the high correlation between SPI-3 and
62 vegetation growth over croplands and grasslands in the U.S. Great Plains. Wang et al. (2015)
63 observed a good matching between soil moisture dynamics and SPI at the scale of 1-3 months
64 when testing various indices over China. In Europe, Manning et al. (2018) highlighted how
65 precipitation is the main driver of soil moisture droughts for a set of both dry and wet sites.

66 In spite of the above-mentioned consistencies, the index selected to characterize drought
67 conditions over a certain study region will inevitably affect the outcome of the drought analysis,
68 as highlighted by Quiring and Papakryiakou (2003) in testing different indices over the Canadian
69 prairies. These Authors suggest that a variety of drought indices should always be tested to
70 determine the most appropriate one for each specific application. It follows that the synergy
71 between multiple indices can be exploited by the use of multivariate indicators (Hao and Singh,
72 2015), a family of approaches that encompasses a variety of merging strategies, including
73 combined cascading indices (Cammalleri et al., 2021a; Rembold et al., 2019), composite and
74 integrated approaches (Brown et al., 2008; Svoboda et al., 2002), and joint probability functions
75 (Bateni et al., 2018; Hao and AghaKouchak, 2013; Kanthavel et al., 2022).

76 The latter class of approaches, in particular, aims at capturing the complex statistical
77 interdependence among different drought-related variables (Hao and Singh, 2015), and it has seen
78 a growing relevance in many hydrological applications thanks to the introduction of copula
79 functions and their ability to model a wide range of dependence structures (Nelsen, 2006; Salvadori
80 et al., 2007; Joe, 2015). In the field of drought indices, the approach proposed by Kao and
81 Govindaraju (2010) for the computation of the Joint Deficit Index (JDI) has been applied to a
82 variety of drought-related quantities over different regions, often including precipitation and soil
83 moisture (i.e., Dash et al., 2019; Kwon et al., 2019).



84 Studies on the marginal distribution of either precipitation or soil moisture have somewhat
85 converged on adopting the Gamma distribution for precipitation and the Beta distribution for soil
86 moisture. The use of the Gamma family for the implementation of the SPI at different accumulation
87 periods has been a standard practice in many applications (e.g., Mo and Lyon, 2015; Yuan and
88 Wood, 2013). While other distributions have also proven to be reliable, such as the exponentiated
89 Weibull (Pieper et al., 2020) and the Person Type III (Ribeiro and Pires, 2016), the Gamma is still
90 the most adopted one. Over Europe, Stagge et al. (2015) demonstrated how the Gamma
91 outperformed the other tested distributions across all accumulation periods and regions.

92 A more limited number of applications based on soil moisture data are available in the
93 literature compared to SPI. The use of the Beta distribution for soil moisture data has been
94 introduced as early as the late 70s, with the pioneer study of Ravelo and Decker (1979), following
95 the consideration that soil moisture is a double-bounded quantity, ranging between residual and
96 saturation. Sheffield et al. (2004) successfully applied this standardization for drought analyses
97 over the US, while the same distribution has been adopted by Cammalleri et al. (2016) on modeled
98 data over Europe. Most recently, the Beta distribution was also used to characterize the frequency
99 of global satellite soil moisture data (Sadri et al., 2020).

100 Conversely, no standard approaches have been identified for the application of copulas to
101 model the bivariate joint distribution of precipitation and soil moisture, mainly due to the large
102 variety of probabilistic structures than may be observed between these two quantities. Common
103 fitting strategies rely on the application of various copula families to identify the best fitting for
104 each specific site (e.g., Hao and AghaKouchak, 2013), or are based on an a-priori selection of a
105 copula family following empirical evidence (e.g., Dixit and Jayakumar, 2021).

106 A comprehensive study on the joint probabilistic dynamic of these two quantities, and on
107 their bivariate distribution, is currently lacking in the scientific literature of multivariate drought
108 modeling. Hence, the main goal of this study is to fill this gap, by investigating the mutual
109 relationship between precipitation (cumulated over 3 months, as for SPI-3) and soil moisture
110 datasets as available over Europe as part of the European Drought Observatory of the Copernicus
111 Emergency Management Service (EDO, <https://edo.jrc.ec.europa.eu>).

112 A large set of copulas is tested for this purpose across the entire European domain, to
113 identify an optimal modeling of the dependence especially in proximity of the tails (given its major
114 role in extreme detection). The spatial distribution of the results is analyzed to infer evidence of



115 common patterns and behavior, which may support future operational applications based on
116 similar parametric approaches.

117

118 **2. Materials and Methods**

119 **2.1 Precipitation and soil moisture datasets**

120 The study focuses on Europe and makes use of the dataset of indicators available over the region
121 as part of EDO. Precipitation data accumulated over consecutive 3-month periods are used here,
122 as the quantity at the base of the SPI-3 index. Hourly total precipitation maps from the ECMWF
123 ERA5 global atmospheric reanalysis model ([https://www.ecmwf.int/en/forecasts/dataset/ecmwf-](https://www.ecmwf.int/en/forecasts/dataset/ecmwf-reanalysis-v5)
124 [reanalysis-v5](https://www.ecmwf.int/en/forecasts/dataset/ecmwf-reanalysis-v5)) are collected through the Copernicus Climate Change Service (C3S,
125 <https://climate.copernicus.eu/>) and cumulated at monthly updates (no missing values are present
126 in the reanalysis dataset). This dataset has proven to be quite reliable over Europe for drought
127 analyses (e.g., Cammalleri et al., 2021b; van der Wiel et al., 2022), as it is currently employed in
128 near-real time as part of the operational tools of EDO.

129 Soil moisture records over the entire European domain are derived from the simulations of
130 the LISFLOOD distributed hydrological rainfall–runoff model (de Roo et al., 2000). LISFLOOD
131 runs in near-real time as part of the European Flood Awareness System (Thielen et al., 2009), and
132 it provides daily soil moisture maps for the root zone at a spatial resolution of 5-km. Daily modeled
133 data are averaged at monthly scale and converted into a Soil Moisture Index (SMI) as in
134 Seneviratne et al. (2010). The model is calibrated and validated over an extensive network of river
135 discharge stations following the procedure described in Arnal et al. (2019), and it has been
136 successfully tested for drought analyses over Europe as part of EDO for the computation of the
137 Soil Moisture Anomaly (SMA) index (Cammalleri et al., 2015).

138 In this study, data collected on the most recent 25 years (1996-2020) are used as a common
139 period. This period is chosen to minimize the effects of non-stationarity in precipitation records
140 and to avoid the inclusion of early LISFLOOD records that are affected by a lower number of
141 ground meteorological stations in the forcing (Thieming et al., 2022). The 300 maps (12 months ×
142 25 years) for the two datasets are then spatially interpolated on a common Lambert azimuthal
143 equal-area (LAEA) projection on a regular grid of 5-km using the nearest neighbor algorithm. This
144 is done to preserve the high-resolution information of the soil moisture and by considering the
145 smooth spatial dynamics of precipitation accumulated over 3 months.



146 **2.2 Copula families**

147 The introduction of copulas in multivariate probability modeling has provided to hydrologists a
148 flexible tool to reproduce the joint probability of multiple dependent variables characterized by a
149 variety of marginal distributions (De Michele and Salvadori, 2003; Salvadori and De Michele,
150 2004).

151 Limiting the focus on bivariate variables, the joint probability distribution, F , of two
152 random variables (X_1 and X_2) can be expressed, thanks to the Sklar's theorem, as:

153
$$F(x_1, x_2) = C(F_1(x_1), F_2(x_2)) \tag{1}$$

154 where F_1 and F_2 are the marginal distribution of X_1 and X_2 , respectively, and C is the copula
155 function (Salvadori et al., 2007).

156 A large variety of parametric formulations has been introduced in the literature to explicitly
157 link the marginal to the joint distributions, with some of the most common copula families used in
158 hydrology belonging to the Elliptical and Archimedean copulas (Chen and Guo, 2019). Two non-
159 parametric measures of dependence play a major role in parametric copula inference. The Kendall
160 rank correlation coefficient (τ) is commonly used as a measure of overall ordinal association, while
161 the so-called Tail-Dependence (TD, Salvadori et al., 2007) is used to estimate the asymptotical
162 degree of dependence in the upper and lower extremes (upper tail-dependence, λ_U , and lower tail-
163 dependence, λ_L , respectively). The non-parametric values of both TDs can be evaluated following
164 the method proposed by Schmidt and Stadtmueller (2006).

165 In this study, the parametric bivariate probability of precipitation and soil moisture is
166 assessed using the R package "VineCopula" (Aas et al., 2009; Dissman et al., 2013). The Akaike
167 Information Criterion (AIC, Stoica and Selen, 2004) is used to select, for each spatial cell, the best
168 fitting copula among the wide range of families available. The main properties of some relevant
169 copulas are reported in Table 1, as they will be useful to interpret the successive results.

170 In particular, from the data in Table 1 it is important to highlight how the BB7 copula is a
171 combination of the Joe and Clayton, of which it inherits the tail-dependences, and how the TD
172 behavior of a copula can be inverted (i.e., the upper tail-dependence can become the lower and
173 *vice versa*) by simply considering the reciprocal marginals (commonly known as rotated forms,
174 identified by the suffix 180).

175

176



177 **Table 1.** Main copulas analyzed in this study and their coefficients for the upper and lower tail-
 178 dependences (λ_L and λ_U , respectively).

Copula	λ_L	λ_U
Gaussian	0	0
Student-t	$2t_{\nu+1}\left(-\sqrt{\nu+1}\sqrt{\frac{1-\rho}{1+\rho}}\right)$	$2t_{\nu+1}\left(-\sqrt{\nu+1}\sqrt{\frac{1-\rho}{1+\rho}}\right)$
Gumbel	0	$2-2^{\frac{1}{\theta}}$
Clayton	$2^{-\frac{1}{\theta}}$	0
Joe	0	$2-2^{\frac{1}{\delta}}$
BB7	$2^{-\frac{1}{\theta}}$	$2-2^{\frac{1}{\delta}}$

179

180 Even if a copula is selected as the optimal based on the AIC, this does not necessarily
 181 exclude that other copulas may perform similarly. For this reason, we introduced a further test

182 based on the relative likelihood criterion (Burnham and Anderson, 2002), $\exp\left(\frac{AIC_{\min} - AIC_i}{2}\right)$,
 183 to establish the likelihood that an AIC value of a given copula (AIC_i) is statistically significantly
 184 different that the minimum value (AIC_{\min}) obtained for the optimal solution.

185 2.3 Random forest classification of selected copulas

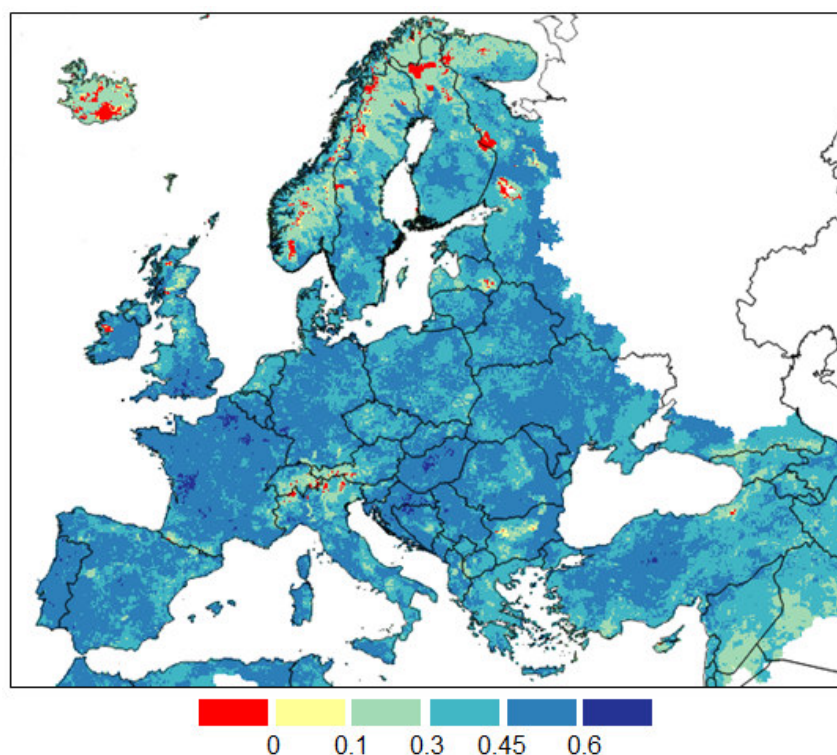
186 The interpretation of the selected copula functions may help highlighting the transferability of the
 187 observed results over different contexts. For this reason, the observed spatial distribution is
 188 analyzed through a random forest classifier (Breiman, 2001), in order to find evidences of
 189 reproducible patterns beyond simple chance.

190 As input features we consider a set of commonly available variables, such as: ground
 191 elevation, annual average temperature, annual total precipitation, precipitation seasonality (ratio
 192 between total precipitation in warm and cold months), annual average Normalized Difference
 193 Vegetation Index (NDVI), annual average soil moisture, and soil type. As hyperparameters for the
 194 random forest, we tuned the number of trees (ntree) and the number of features randomly sampled
 195 at each split (mtry) using the “randomForest” R package (Breiman, 2001).



196 **3. Results**

197 A preliminary analysis of the degree of correlation between the monthly empirical frequencies of
198 3-month precipitation and soil moisture is tested on the full timeseries of each grid cell using the
199 non-parametric Kendall rank correlation coefficient (also known as Kendall's τ), as depicted in
200 Fig. 1 for the entire European domain.



201

202 **Fig. 1.** Spatial distribution of the Kendall's τ between monthly empirical frequencies of 3-month
203 precipitation and soil moisture. Roughly, values lower than 0.1 are not statistically significant at p
204 = 0.05 (two-tails).
205

206 The results reported in Fig. 1 confirms the expected direct relation between the two
207 quantities, with a relatively homogeneous distribution of medium/high correlation τ values
208 between 0.3 and 0.5 ($\tau = 0.42 \pm 0.1$). Limited regions with low (and sometime even slightly
209 negative) τ values are sporadically observed, mostly concentrated over the Alps, Iceland and the
210 coldest regions of the Scandinavia peninsula. Correlations over these regions are likely affected

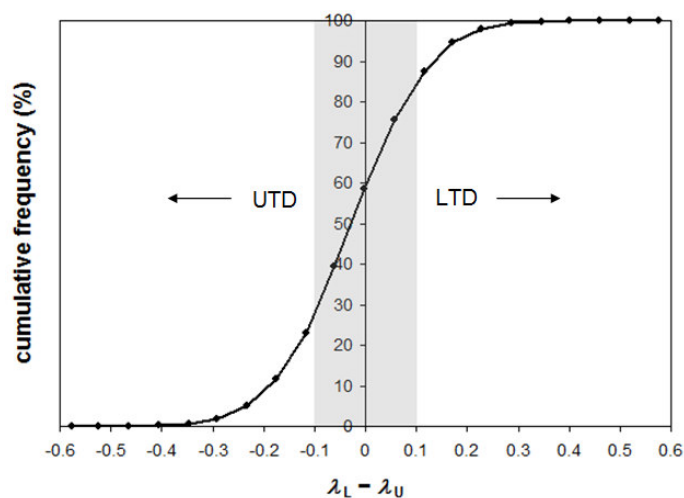


211 by the presence of snow coverage during extended periods of the year. Overall, the observed
 212 τ values cannot be considered statistically significant (at $p = 0.05$) only for less than 2% of the
 213 domain.

214 The analysis of the non-parametric tail-dependence values is summarized in the plot
 215 depicted in Fig. 2, where the cumulative frequency of the difference between the empirical λ_L and
 216 λ_U values is reported. Symmetric behaviors in Fig. 2 can be identified by setting a maximum value
 217 for $|\lambda_L - \lambda_U|$. To identify this threshold, non-parametric TDs were re-computed on shuffled time
 218 series (to artificially reconstruct conditions of null tail-dependencies), and the $|\lambda_L - \lambda_U|$ value
 219 corresponding to a cumulative frequency of 90% of the grid cells after the shuffling was detected
 220 as threshold, corresponding to 0.1.

221 The plot in Fig. 2 highlights how the majority (about 50%) of the grid cells can be
 222 considered characterized by a symmetric behavior in the tail-dependence ($|\lambda_L - \lambda_U| < 0.1$), whereas
 223 the rest of the grid cells are almost equally split between a predominance of the Upper Tail-
 224 Dependence (UTD, corresponding to negative differences) or a predominance of Lower Tail-
 225 Dependence (LTD, positive differences).

226



227

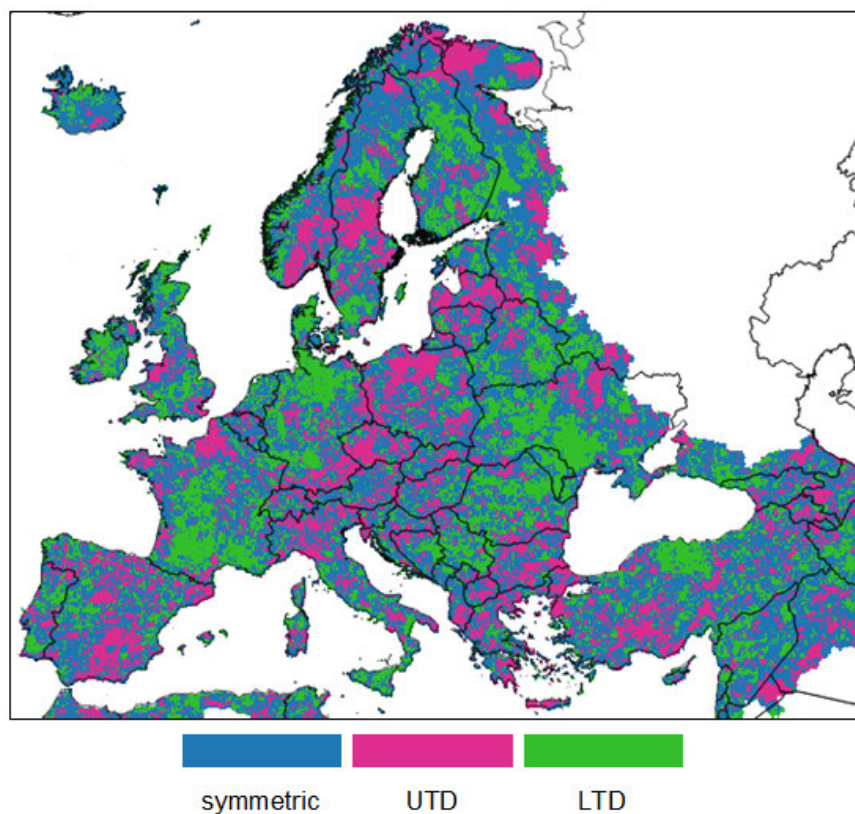
228 **Fig. 2.** Analysis of the frequency of the empirical tail-dependencies. The plot shows the cumulative
 229 frequency distribution of the differences between the empirical λ_L and λ_U values computed
 230 according to Schmidt and Stadtmueller (2006). The domain with a roughly symmetric behavior
 231 ($|\lambda_L - \lambda_U| < 0.1$) is highlighted by the grey box area.



232

233 The results reported in Fig. 2 were used to divide the entire domain in three categories
234 (symmetric, LTD and UTD) as depicted in Fig. 3. This map shows evidence of some coherent
235 spatial patterns, such as the predominance of LTD in southern France, southern Italy, northern
236 Germany and Denmark, and western Ukraine (among others), and a clustering of UTD in Poland,
237 Czech Republic, southern Scandinavia, and Greece. The symmetric condition seems overall more
238 spread across the entire domain, also thanks to the higher frequency, with a slightly predominance
239 over northern Europe (i.e., northern Scandinavian peninsula and Iceland).

240



241

242 **Fig. 3.** Spatial distribution of the three categories derived from the differences in the empirical tail-
243 dependencies.

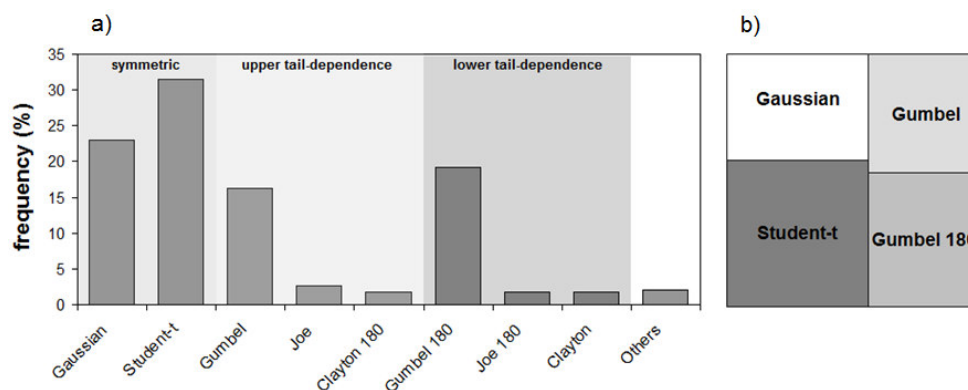
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245 Given the results observed in terms of tail-dependence, is it useful to focus on the capability
 246 to reproduce such patterns instead of finding the single copula that can perform reasonably well
 247 over the entire domain. Indeed, the search for the optimal copula based on the minimum AIC
 248 returns the BB7 as the optimal solution in about 80% of the domain (not shown). This result is
 249 thanks to the flexibility of its formulation (derived as a combination of two purely asymmetric
 250 functions), which allows reproducing both symmetric and asymmetric tail-dependencies
 251 depending on the values assumed by the two parameters. However, the fact that a single flexible
 252 copula works well over a large range of conditions may hide the key spatial patterns observed in
 253 TD, which can be highlighted instead by adopting a limited number of more common copulas
 254 specialized in reproducing specific behaviors.

255 By limiting the search to a subset of copula functions, comprising only purely symmetric
 256 or purely asymmetric tail behaviors, more interesting results are obtained, as summarized by the
 257 frequency plot in Fig. 4. The grid cells where symmetric tail behavior copulas are selected as
 258 optimal are about 55% of the domain (see Fig. 4b), with a predominance of Student-t copula but
 259 also with a non-negligible fraction of cells (23%) where the Gaussian (symmetric and without tail-
 260 dependences) is chosen (see Fig. 4a). The remaining grid cells are almost equally split between
 261 upper and lower tail-dependences, with Gumbel (and its rotated counterpart, Gumbel 180) as the
 262 most selected among the asymmetric options.

263



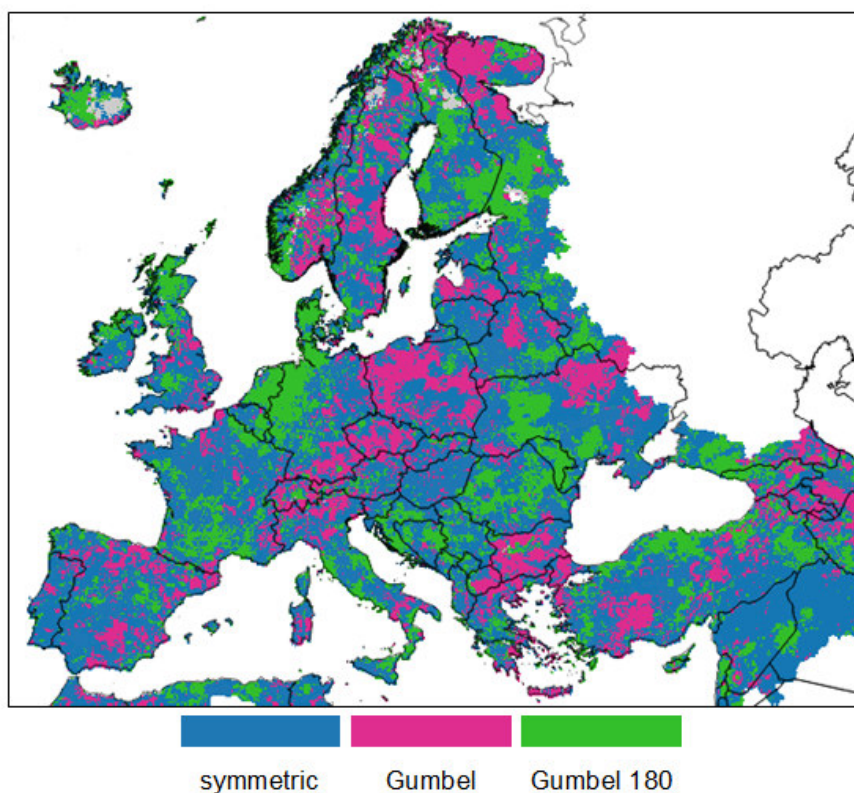
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265 **Fig. 4.** Frequency of the optimal copulas based on the minimum AIC. The barplot in panel a)
 266 shows the frequency for each copula, whereas the box in panel b) reports a synthetic description
 267 of the subdivision of the entire domain among the 4 most frequent copulas.

268



269 The spatial distribution of these optimal copulas (Fig. 5) confirms most of the patterns
270 observed in Fig. 3, as a further confirmation that a rather limited range of simple copula functions
271 is able to capture the overall dynamics of dependence between precipitation and soil moisture over
272 the entire European domain. Despite the observed spatial clusters in the obtained optimal copulas,
273 the overall patterns observed in Fig. 5 are still rather noisy and may be difficult to interpret. This
274 erratic behavior can be partially explained by the fact that different copulas may perform quite
275 similarly over some grid cells, hence the AIC of the optimal copula (AIC_{min}) may not differ
276 significantly from the AIC of other functions.

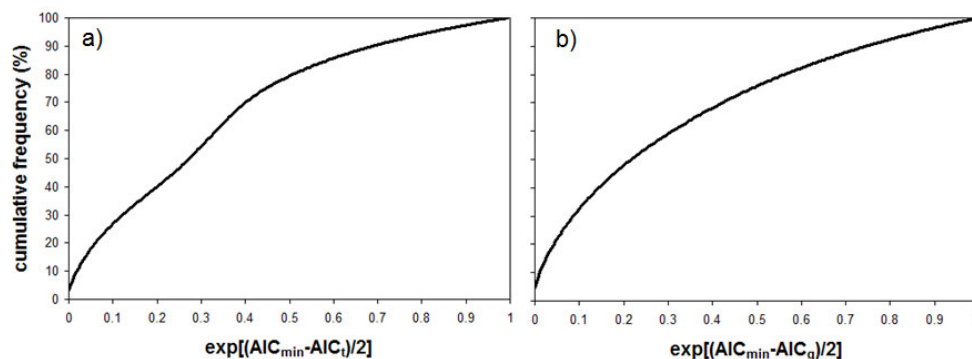


278 **Fig. 5.** Spatial distribution of the optimal copulas obtained by minimizing the AIC. The symmetric
279 tail behavior class includes both Gaussian and Student-t copulas.
280

281 To further investigate this hypothesis, we evaluated the possibility to replace the optimal
282 copulas with either a Student-t or a Gumbel (direct and rotated) over the entire domain. The



283 Gaussian copula was excluded from this analysis under the assumption that the no tail-dependence
 284 of the Gaussian can be adequately reproduced by the Student-t with a small enough tail-
 285 dependence. The plots in Fig. 6 reports the relative likelihood for the Student-t (panel a) and
 286 Gumbel families (panel b) compared to the locally selected optimal copulas. Low values of this
 287 metric correspond to conditions where the optimal copula cannot be replaced by the alternative
 288 function (being either the Student-t or the Gumbel).
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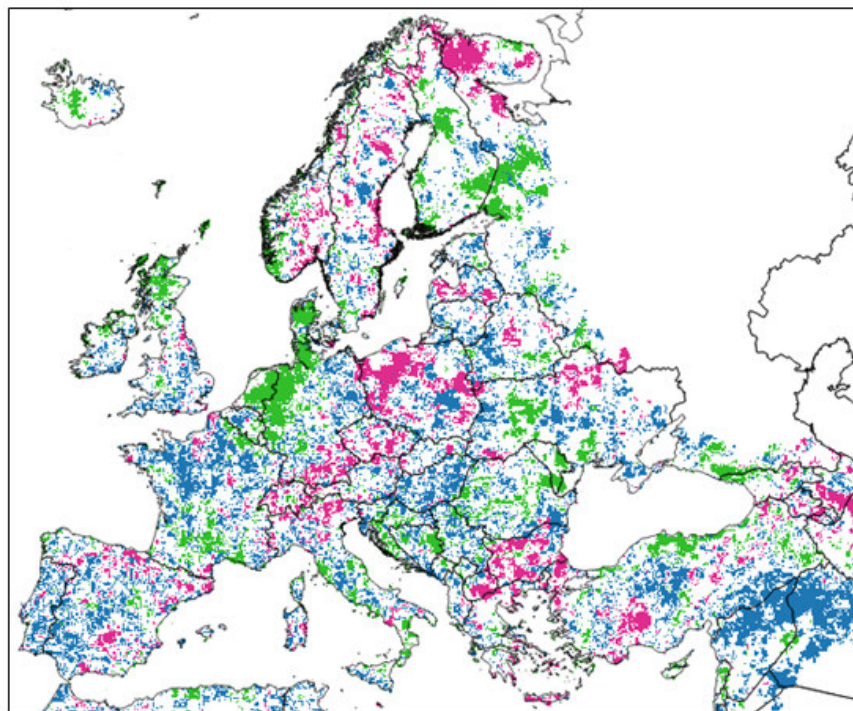
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291 **Fig. 6.** Frequency analysis of the relative likelihood computed between the optimal AIC (AIC_{min})
 292 and: a) Student-t (AIC_t), or b) Gumbel (AIC_g) families. The grid cells where either the Student-t
 293 or the Gumbel was already the optimal solution were excluded from the respective frequency
 294 analysis.
 295

296 The results in Fig. 6 show that, if we assume a relative likelihood of 0.1 as a threshold to
 297 detect a statistically significant difference, the Student-t cannot reasonably replace the local
 298 optimal copula in about 18% of the entire domain, whereas this fraction is about 17% for the
 299 Gumbel family. It is also possible to observe how the Gumbel family is the optimal copula in
 300 almost the totality (about 99%) of the grid cells where the Student-t is not a suitable replacement
 301 of the local optimal, whereas almost only symmetric copulas (63% Student-t and 34% Gaussian)
 302 are the optimal functions where the Gumbel family is not a suitable replacement. Overall, these
 303 results suggest that the selection of the optimal copula is “univocal” (i.e. cannot be reasonably
 304 replaced by another function) in about 35% (18+17) of the domain, whereas either the Student-t
 305 or the Gumbel families can be adopted in the remain fraction of the domain with similar
 306 performances in terms of AIC. It is worth mentioning how this analysis confirms the assumption



307 that all the areas where the Gaussian was chosen as optimal copula can be satisfactory modeled
308 also by using the Student-t (i.e. without a statistically significant increase in AIC).
309



310

311 **Fig. 7.** Spatial distribution of the grid cells where the selection of the optimal copula is “univocal”
312 according to the relative likelihood criterion.

313

314 The “univocal” areas derived from the previous analysis are mapped in Fig. 7, highlighting
315 some of the more consistent spatial clusters already observed in both Figs. 3 and 5, as well as a
316 large fraction of cells in northern Europe where a “univocal” optimal copula cannot be selected.
317 These grid cells with “univocal” copula are used as a starting point for the random forest
318 classification, given the robustness in their signal.

319 A sample of 25% of the “univocal” grid cells (corresponding to about 8% of the entire
320 domain) was used to train the random forest, adopting a number of trees (ntree) of 80 and a single



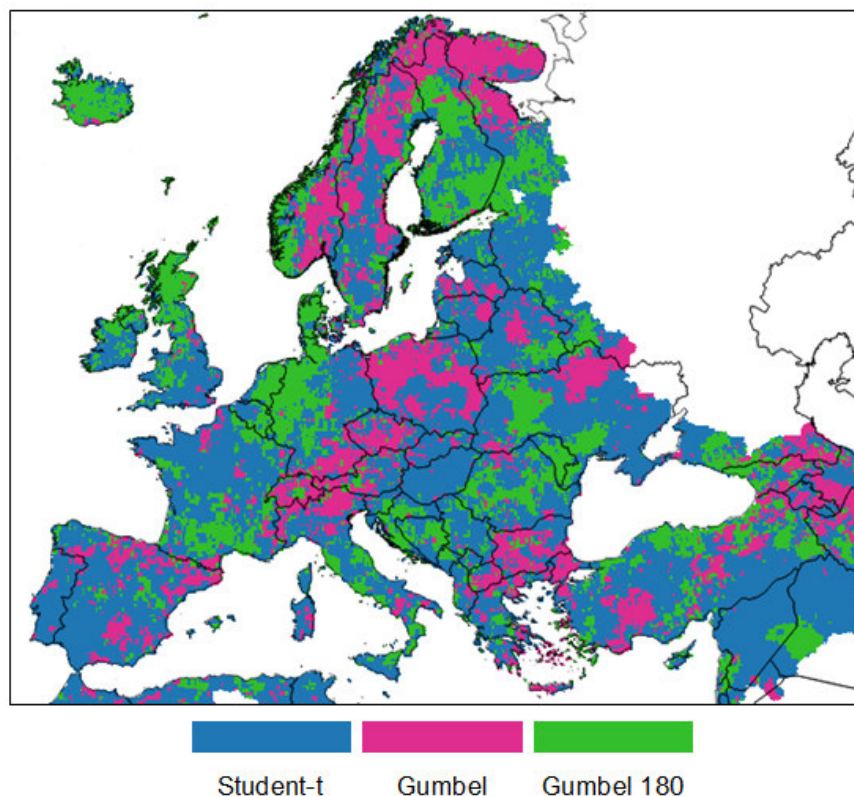
321 feature randomly sampled at each split ($m_{try} = 1$). The training size and minimum values of
322 hyperparameters were chosen to reduce the problem of overfitting. Among the possible features,
323 three variables were selected by analyzing the variable importance plots, as well as the ease of
324 access: annual average temperature, annual total precipitation, and precipitation seasonality. The
325 trained classifier was then applied to the testing subset (the remaining 75% of the “univocal” grid
326 cells) and the outcomes were analyzed by mean of a confusion matrix, which results are
327 summarized in Table 2. Overall, the obtained classification has a very satisfactory matching with
328 the test subset, with a general high accuracy ($ACC = 0.86$) and with all the metrics pointing toward
329 a significant improving in the performance compared to the reference No-Information-Rate (NIR)
330 (i.e., small p-values) and a high probability to have correct modeled values compared to simple
331 chance (i.e., high Cohen’s K).

332

333 **Table 2.** Summary of the confusion matrix analysis applied to the trained random forest on the
334 testing subset.

Accuracy (ACC)	0.86
No-Information-Rate (NIR)	0.50
p-value ($ACC > NIR$)	$< 2.2 \times 10^{-16}$
McNemar’s test p-value	3.44×10^{-5}
Cohen’s kappa statistic (K)	0.78

335



337 **Fig. 8.** Map of the optimal copula as modeled by the trained random forest classifier.
338

339 Finally, the trained classifier was extended to the entire domain to obtain a classification
340 of the entire European domain in term of the expected optimal copula and TD behavior. This map,
341 reported in Fig. 8, bears a strong resemblance to both the empirically-derived map in Fig. 3 and
342 the optimal AIC fitting in Fig. 5. Beside this overall agreement, some notable discrepancies can
343 be observed over northern Scandinavia and Iceland, two regions where low Kendall's τ and a small
344 fraction of "univocal" selected copulas were already identified.

345

346 **4. Discussion**

347 The overarching goal of the study is to investigate the joint probability of two variables
348 aiming at capturing agricultural drought conditions, hence the overall agreement between these
349 two quantities is a fundamental prerequisite. The expected direct relationship between 3-month



350 cumulated precipitation and soil moisture, as both SPI-3 and SMA are similarly-used agricultural
351 drought indices, can be seen as a first clue for the identification of the most suitable set of copula
352 families (Salvadori et al., 2007; Genest et al., 2007). This behavior is overall confirmed by the
353 positive Kendall's τ values observed over most of the domain ($\tau = 0.42 \pm 0.1$). Moderately high
354 correlation values were observed between precipitation and soil moisture also in other studies.
355 Kwon et al. (2018) reporting Pearson's r values between 0.4 and 0.6 for 55 stations in South Korea,
356 albeit with seasonal patterns; Gaona et al. (2022) reported similar values over the Ebro basin with
357 both land-surface modeled and satellite soil moisture, and Sepulcre-Cantó et al. (2012) obtained
358 an average value of r of about 0.6 over nine stations across Europe.

359 Sehler et al. (2019) studied the correlation between remote sensing-based precipitation and
360 soil moisture, finding moderate correlation over southern Europe, and a weak (often not
361 significant) correlation in central Europe. However, central Europe is close to the upper limit of
362 the analyzed remote sensing products, which can explain such low performance. Limited
363 correlation even among different soil moisture products has been observed in northern Europe in
364 other studies (Almenda-Martin et al., 2022), confirming the difficulty to model the soil moisture
365 dynamics over this region.

366 The obtained values for the Kendall's τ fall in a somewhat optimal range for the analysis
367 of the joint probability, since τ values are statistically significant almost everywhere (i.e.,
368 consistency in the produced outcomes) but not too high to make meaningless any joint use of the
369 two datasets (i.e., too similar products).

370 Even more interesting is the outcome of the tail-dependence analysis, given the role that
371 such quantity, and in particular the low-tail, plays in the detection of drought extreme events. The
372 TD investigation is sometime overlooked in multivariate drought analyses, where previous studies
373 often focused on optimizing the copula to the local data without analyzing the empirical TD and
374 the implications for the modeling of drought conditions. Indeed, TD is rarely the focus of extensive
375 analyses, such as the one reported in this study for the entire Europe, and previous references in
376 the scientific literature for precipitation and soil moisture are rather scarce.

377 As an example, Manning et al. (2018) performed a very detailed analysis over 11 FluxNet
378 sites in Europe on the role of precipitation and evapotranspiration on soil moisture drought, based
379 on pair copula constructions, but the authors did not provide any indication on which bivariate
380 copula was the optimal for each site. Kwon et al. (2018) reported that Frank copula was the most



381 frequent optimal choice in their study over South Korea, but some clear spatial patterns were also
382 observed in their outcomes, with Frank being the selected copula mostly in the central area, but
383 also Gumbel and Student-t performing the best in the southern and eastern coasts, respectively.

384 Dash et al. (2019) found Frank (among the Archimedean copulas) working the best for 3-
385 month precipitation and soil moisture over an Indian basin, while Hao and AghaKouchak (2013)
386 highlighted the good performance of Frank and Gumbel in five regions in California, even if
387 neither Gaussian nor Student-t were considered. In all these applications, no specific
388 considerations on the observed TD behaviors were reported, even if a common trend seems to be
389 the good performance of Frank copula, which is in contrast with our results, where the Frank was
390 very rarely selected as optimal (less than 1% of the domain). A possible explanation of these results
391 may be our focus on empirical marginal frequencies rather than theoretical ones, given the well-
392 documented increasing uncertainty in parametric fitting in the tails (Farahmand and
393 AghaKouchak, 2015; Laimighofer and Laaha, 2022). As a possible confirmation of this
394 hypothesis, a good performance of Gumbel and Gaussian has been observed over Iran by Bateni
395 et al. (2018), similarly to our results, when a nonparametric form for SPI and SSI (Standardized
396 Soil Moisture Index) was used.

397 The absence of a strict standard procedure to investigate tail-dependence may be another
398 factor affecting the limited focus on the topic of drought studies. Non-parametric TD has the clear
399 advantage to avoid any alteration of the data due to the fitting procedure, but the outcomes in this
400 study also show a high degree of spatial noise likely due to the intrinsic nature of non-parametric
401 analyses, as well as to the limited sample size which affects the estimates of TD. For this last issue
402 see also the illustration 3.18 in Salvadori et al. (2007). The threshold used here to define a
403 symmetric behavior, based on a random shuffling of the data, seems to successfully overcome the
404 difficulty to define a self-consistent maximum difference in TDs.

405 The fitting of parametric copula functions returns more consistent spatial patters in our
406 study, but evidence on the absence of “univocal” fittings can be observed, as well as some
407 contrasting results compared to the non-parametric TD especially over northern Europe (areas with
408 low correlation). The grid cells where a given copula clearly outperforms the alternative options
409 is limited to roughly 1/3 of the domain, further stressing the evidence that clear-cut outcomes are
410 difficult to infer. In this regard, it seems reasonable to infer that only a critical concerted analysis



411 of both parametric and non-parametric TDs can return robust indications based on a converge of
412 evidence.

413 A clear outcome of our study is the predominance of regions with symmetric tail-
414 dependences, where the Student-t copula is suitable to reproduce the joint probability of
415 precipitation and soil moisture. An even split of the remaining domain between areas with either
416 lower or upper tail-dependence only is also observed, where the Gumbel copula (in either is direct
417 or 180 rotated forms) is proven to be a suitable option. These results are crucial in defining the
418 role of precipitation and soil moisture datasets in detecting drought events, and to which extent
419 they can work in synergy in a drought monitoring system. In fact, while the correlation between
420 the two datasets highlights the extent of their overall agreement, which in this study was somewhat
421 uniform across most of the domain, very different degrees of consistency can be obtained for
422 similar Kendall's τ if the TDs differ substantially. Regions with high LTD will have a high
423 agreement in the detection of drought extremes, hence a low number of false alarm and a higher
424 signal-to-noise ratio is expected.

425 In this regard, regions such as southern France, northern UK, northern Germany and
426 Denmark, where a strong LTD is observed, are appropriate candidates for a robust assessment of
427 agricultural drought conditions based on a joint precipitation-soil moisture index, whereas some
428 regions in central Europe (i.e., Poland, Czech Republic, Switzerland) may not equally benefit from
429 the use of a joint index due to the absence of LTD.

430 Overall, the fact that the copula fittings confirm most of the non-parametric TD patterns
431 suggests that a parametric approach is suitable for an operational implementation of a
432 precipitation-soil moisture joint drought index over most of Europe, as well as a tool to provide
433 meaningful insight on the potentiality of joint probability as detector of extreme droughts.

434 Even if, at first glance, it may seem difficult to assign a meaningful explanation for the
435 observed spatial patterns in LTD and UTD, the proven possibility to reasonably reconstruct these
436 spatial patterns with a random forest classifier, starting from only a small sample of robust training
437 data (less than 10% of the domain) and with common driving features, suggests that the observed
438 clusters are unlikely to be caused only by chance and that hidden structures may be present and
439 further explored. This result is encouraging for an extension of the derived considerations to other
440 spatial regions of the world.

441



442 **5. Summary and Conclusions**

443 The use of combined indices based on copula seems a promising development in the field of
444 drought detection and monitoring. In this study, we analyzed the joint probability of two quantities
445 commonly used to derive drought indices, 3-month cumulated precipitation and soil moisture, with
446 a special focus on the probabilistic characteristics that are key for their usage in agricultural
447 drought analyses.

448 The overall agreement in the marginal probability of the two variables suggests that they
449 are indeed valid candidates for the development of a joint drought index over the European domain.
450 However, an in-depth analysis of the tail-dependence, derived with both non-parametric and
451 parametric approaches, shows some clear spatial patterns, which have direct repercussion on the
452 capability of such data to provide robust estimates of the extremes. In this regard, regions such as
453 southern France, northern UK, northern Germany, and Denmark may benefit more from the joint
454 use of the two variables thanks to the observed strong low tail-dependence. The joint dependence
455 of precipitation and soil moisture is well reproduced using three commonly-used copulas (Student-
456 t, Gumbel and 180 rotated Gumbel), which spatial patterns were successfully reconstructed with a
457 random forest classification, suggesting the presence of a structure in the outcomes not related to
458 chance.

459

460 **Code availability:** The codes used for this analysis can be provided upon request via the
461 corresponding author.

462

463 **Data availability:** All the data used in this study can be access through the European Drought
464 Observatory (EDO) web portal (<https://edo.jrc.ec.europa.eu/gdo/php/index.php?id=2112>).

465

466 **Author contribution:** CC designed the experiments, with inputs from AT and CDM. CC
467 developed the codes and performed the analyses. CC prepared the manuscript, which was
468 expanded and revised by all co-authors.

469

470 **Competing interests:** At least one of the (co-)authors is a member of the editorial board of
471 Hydrology and Earth System Sciences.



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