A novel, cost-effective analytical method for measuring high-resolution vertical profiles of stratospheric trace gases using a GC-ECD

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Abstract. The radiative balance of the upper atmosphere is dependent on the magnitude and distribution of greenhouse gases and aerosols in that region. Climate models predict that with increasing surface temperature, the primary mechanism for transporting tropospheric air into the stratosphere (known as the Brewer-Dobson Circulation) will strengthen, leading to changes in the distribution of atmospheric water vapor, other greenhouse gases, and aerosols. Stratospheric relationships between greenhouse gases and other long-lived trace gases with various photochemical properties (such as N_2O, SF_6, and chlorofluorocarbons) provide a strong constraint for tracking changes in the stratospheric circulation. Therefore, a cost-effective approach is needed to monitor these trace gases in the stratosphere. In the past decade, the balloon-borne AirCore sampler developed at NOAA's Global Monitoring Laboratory has been routinely used to monitor the mole fractions of CO_2, CH_4, and CO from ground to approximately 25 km above mean sea level. Our recent development work adapted a gas chromatograph coupled with an electron capture detector (GC-ECD) to measure a suite of trace gases (N_2O, SF_6, CFC-11, CFC-12, H-1211, and CFC-113) in the stratospheric portion of AirCores. This instrument, called the StratoCore-GC-ECD, allows us to retrieve vertical profiles of these molecules at high resolution (5-7 hPa per measurement). We launched four AirCore flights and analyzed the stratospheric air samples for these trace gases. The results showed consistent and expected tracer-tracer relationships and good agreement with recent aircraft campaign measurements. Our work demonstrates that the StratoCore-GC-ECD system provides a low-cost and robust approach to measuring key stratospheric trace gases in AirCore samples and for evaluating changes in the stratospheric circulation.

1. Introduction

Monitoring the dry air mole fractions of a suite of trace gases in the stratosphere will significantly improve our understanding of the stratospheric mean meridional circulation's (Brewer-Dobson Circulation, or BDC) response to changing climate. The BDC is characterized by upwelling in the tropics, with upper and lower poleward branches and descent in the extra-tropics (Holton et al., 1995; Garcia and Randel, 2008; Butchart, 2014). Coupled chemistry-climate models (CCMs) predict an acceleration of the BDC in response to increasing greenhouse gas abundances and surface temperatures (Butchart

et al., 2006; Garcia and Randel, 2008; McLandress and Shepherd, 2009; Butchart, 2014), with far-reaching implications for surface weather, earth's radiation budget, and the climate (Forster and Shine, 2002; Randel et al., 2006; Gerber et al., 2012), recovery of the stratospheric ozone layer (Butchart and Scaife, 2001; Butchart et al., 2010) and potential impacts to surface air quality due to changes in stratosphere-to-troposphere ozone flux and the tropospheric oxidation capacity (Hegglin and

35 Shepherd, 2009). Furthermore, evaluating the impact of potential future climate intervention techniques also requires accurate modelling of the BDC in CCMs. However, directly measuring the strength and variation of the BDC is difficult.

The mean age of air (AoA) in the stratosphere (Ray et al., 1999; Andrews et al., 2001; Waugh and Hall, 2002) has been suggested to be an indicator of the BDC strength (Engel et al., 2009; Stiller et al., 2012, 2017). The measurement-derived AoA, traditionally using mole fractions of carbon dioxide (CO_2) and sulfur hexafluoride (SF_6), can be compared with modelled

40 AoA to investigate the model's performance in simulating the BDC. However, later studies showed that the BDC is not the only factor controlling the mean AoA, as it is also affected by the mixing of air from the extra-tropics back into the tropics, i.e., recirculation (Ray et al., 2014; Garny et al., 2014; Ploeger et al., 2015; Dietmüller et al., 2017). Also, recent work has shown that SF_6 has a non-negligible chemical sink in the mesosphere (Ray et al., 2017; Leedham Elvidge et al., 2018; Loeffel et al., 2022), biasing AoA calculations that rely on SF_6 mole fractions alone. Since the mesospheric SF_6 loss rate is proportional

45 to its mole fraction, which has been increasing rapidly in the past few decades, the measured SF_6 mole fraction in the midlatitudes now contains measurable information not only about AoA but also the mass exchange between the stratosphere and the mesosphere, which was only obtainable in polar vortex profiles before. Additional work has shown that the stratospheric dry mole fractions of some long-lived trace gases, such as nitrous oxide (N_2O), and several chlorofluorocarbons (CFCs), including dichlorodifluoromethane (CFC-12), trichlorofluoromethane (CFC-11), 1,1,2-trichloro-1,2,2-trifluoroethane

50 (CFC-113), and bromochlorodifluoromethane (halon 1211, or H-1211), could provide further constraints to help better understand stratospheric circulation and transport pathways of air into the stratosphere (Volk et al., 1996; Strahan et al., 1999; Schoeberl et al., 2000; Moore et al., 2014). This is because 1) photolytic destruction is the sole sink for these gases, 2) their photolytic destruction rates increase exponentially with altitude, and 3) the altitude-photolytic lifetime profiles for these trace gases are different (Moore et al., 2014). Therefore, observations of a suite of tracers are needed to carefully monitor, examine,

55 and verify simulated stratospheric transport.

The lightweight balloon-borne observation system known as the AirCore provides a low-cost approach to observing the composition of the stratosphere (Tans, 2009; Karion et al., 2010). High-quality *in situ* measurements of stratospheric air are rare since the cost of such field campaigns prohibits routine measurements. As a result, data collected from occasional high-altitude large-balloon ($>10^6$ m^3) and aircraft-based field campaigns since the 1980s is still relevant today for diagnosing

60 stratospheric composition and dynamical change (Hall et al., 1999; Andrews et al., 2001; Pan et al., 2010; Laube et al., 2020). The AirCore was developed at the NOAA Global Monitoring Laboratory (NOAA/GML) and has been widely used to measure CO_2, CH_4, and CO profiles in samples collected from the surface to the stratosphere. The AirCore consists of a long (approximately 100 meters), thin, coated stainless steel tube with one open end. The gas in the AirCore flows out as it ascends on a balloon. After the balloon is cut away at 30-32 km above mean sea level (AMSL), the AirCore descends and passively

65   collects ambient air. Due to the relatively low volumetric flow rate and small cross-section area of the AirCore, the mixing of air captured in the AirCore is limited to Taylor dispersion (Aris, 1956) and molecular diffusion, largely preserving the vertical structure of the atmosphere in the AirCore (Tans, 2009; Karion et al., 2010). After landing, the AirCore automatically closes, preserving the collected air sample until laboratory analysis. As AirCores are usually analyzed shortly (less than 4 hours) after landing, the composition of air from the ground to the mid-stratosphere can be measured with only a small amount (less than

70   0.7 m in both directions, Karion et al., 2010) of diffusion and dispersion mixing of the sample, allowing for the retrieval of vertical gradients of trace gases in air from the ground to the mid-stratosphere with significant fidelity at a relatively low cost (~$5000/profile).

The most common analytical approach for analyzing AirCore samples employs continuous-flow gas analyzers to derive the vertical profiles of CO_2, CO, CH_4, and N_2O (Karion et al., 2010; Membrive et al., 2017; Engel et al., 2017). In this

75   approach, the AirCore sample is pushed through one, or a series of, continuous-flow gas analyzer(s), during which the analyzer(s) measure the dry mole fractions of several gases (CO, CO_2, N_2O, and CH_4) at a relatively slow flow rate (approximately 30 mL min^{-1}) and data rate of approximately 0.45 Hz. These measurements are then combined with flight data (such as altitude, pressure, and temperature) to derive vertical profiles of measured trace gases with altitude, using estimates of flow impedance due to flow resistance in a laminar regime as sample air moves along the length of tubing (Tans, 2022).

80   Although this method provides fast, high-resolution measurements of several essential trace gases, the continuous analyzers cannot directly measure other trace gases of interest for evaluating changes in the BDC (such as CFCs or SF_6), limiting the species measured from an AirCore sample. Mrozek et al. (2016) and Laube et al. (2020) have designed sub-sampling systems that separate the AirCore samples into 20-30 mL aliquots, allowing for more detailed chemical and isotopic measurements using non-continuous flow analytical instruments. This method was then applied to measure the dry mole fractions of CFC-11

85   and other trace gases in each sub-sample and to investigate mass-independent fractionation in CO_2 (Mrozek et al., 2016, Laube et al., 2020). However, subsampling from AirCores only allows for a limited number of stratospheric measurements per small AirCore sample volume. In the case of Mrozek et al. (2016), ten stratospheric measurements from a 2 L volume AirCore can be measured in each flight. With the weight of NOAA unmanned free balloon payloads limited to 5.4 kg by the Federal Aviation Administration (FAA) in the United States, the subsampling method would provide lower vertical resolution and thus

90   limited utility in resolving critical stratospheric gradients of these gases. Additionally, NOAA's AirCore sampling program routinely deploys two samplers simultaneously, which currently restricts the total volume of each AirCore to less than 1 L. Therefore, an alternative approach is needed to measure the mole fractions of several critical trace gases in AirCores using a smaller sample volume per measurement.

Here, we present a novel analytical method using a modified gas chromatograph coupled with an electron capture

95   detector (GC-ECD) system to analyze the dry mole fractions of CFC-11, CFC-12, CFC-113, H-1211, N_2O, and SF_6 from the stratospheric portion of AirCore samples (approximately the first 20%-30% of the sampler tubing). We name this system the StratoCore-GC-ECD. The StratoCore-GC-ECD is designed to accomplish high-precision measurements of these six species using only ~4-5 mL of air sample per measurement from the stratospheric portion of AirCores while carefully registering the

altitude information of each data point, allowing us to acquire high-resolution measurements of the vertical gradient of these trace gases from the tropopause to the mid-stratosphere. This analytical method offers the potential for long-term monitoring of these gases using a balloon-borne sampling package that is regulated under the same flight rules as those that apply to weather balloons. This methodology, coupled with the AirCore, will provide us the flexibility to measure important stratospheric tracers at an enhanced spatial and temporal resolution over current analytical methods. Such observations will provide us with valuable information to monitor a suite of trace gases in the stratosphere long-term at low cost, define baseline stratospheric conditions for any perturbations in stratospheric composition due to future climate intervention techniques, and provide observational evidence to detect and monitor changes in the BDC.

2. The StratoCore-GC-ECD setup

2.1 Gas Chromatography

The sample analysis portion of the StratoCore-GC-ECD system is adopted from previous GC systems designed and built for rapid, high-frequency *in situ* analysis on aircraft and large balloon platforms (Elkins et al., 1996; Romashkin et al., 2001; Moore et al., 2003; Hintsa et al., 2021). Figure 1 displays a diagram of the StratoCore-GC-ECD system. The analysis component consists of a two-channel GC-ECD that mimics the design of the UAS Chromatograph for Atmospheric Trace Species (UCATS, Hintsa et al., 2021) and the *in situ* GC system used during the Lightweight Airborne Chromatograph Experiment campaign (LACE, Moore et al., 2003). The GC system uses ultra-high purity nitrogen (N_2) as the carrier gas. In each GC channel, a Valco 12-port 2-position valve (VICI, TX, USA) is used to switch between sample loading (into two 1 mL sample loops) and injecting modes. Each analysis takes 120 seconds in this setup. Channel 1 uses a 10% dimethylsilicone (OV-101) packed column as the pre-column to separate CFC-12, H-1211, CFC-11, and CFC-113 in the sample, which subsequently passes through the main column and is analyzed by the ECD detector. A temperature controller (model CNI16-AL, Omega, CT, USA) is used to control column temperature at 38 °C. The flow rate of carrier gas in this channel is 70 mL per minute, and the pre-column is backflushed for 85 seconds at 100 mL per minute in each analysis to remove the residual sample. Similarly, Channel 2 uses HayeSep® D porous polymer packed columns followed by Molecular Sieve 5A to separate SF_6 and N_2O (controlled at 110 °C), which are then analyzed by the second ECD detector. The carrier gas flow rate in Channel 2 is 70 mL per minute. Backflushing in channel 2 occurs after 55 seconds of each 120-second analysis at 100 mL per minute to remove any residual sample. In addition, a small flow of pure CO_2 (0.2 mL per minute) is mixed into the ECD detector in Channel 2 as a dopant to minimize the matrix effect and improve the ECD response to N_2O (Fehsenfeld et al., 1981).

The StratoCore-GC-ECD system displays adequate analytical precisions suitable for measuring the dry air mole fractions of CFC-11, CFC-12, CFC-113, H-1211, N_2O, and SF_6 (hereafter referred to as target molecules) in the stratosphere. Typical chromatographs are shown in Figure 2. The analytical repeatability of the GC-ECD for the target molecules is evaluated by measuring gas cylinders with well-determined dry mole fractions of target molecules multiple times, and the

130 uncertainties are shown in Table 1. Considering the dry mole fractions of these species in the stratosphere display wide ranges (50%-100% overall variations for CFCs and N_2O, 20% for SF_6), such analytical precisions (<0.7%) of the GC-ECD should be sufficient to understand the stratospheric variability of these species. A set of five gas mixtures in Aculife-treated aluminium cylinders, spanning the range of expected stratospheric dry air mole fractions (20% to 100% of tropospheric values) were prepared and used to calibrate the GC-ECD. Examples of the most recent calibration curves are shown in Figure 3.

135 Furthermore, as new trace gas species emerge and grow in the atmosphere, identifying possible interferences caused by the GC co-elution of target molecules and potential new trace gases is important. We therefore tie the StratoCore-GC-ECD measurements to the surface network program at NOAA/GML, where atmospheric samples are analyzed on both GC-ECD and GC-MS. The intercomparisons between the different analytical techniques could be used to detect potential interferences if they emerge in the future.

2.2 Sample handling

Airborne *in situ* GC-ECD systems typically use a high sample flow (~100 mL per minute) to flush the sample loading system prior to analysis (Elkins et al., 1996; Romashkin et al., 2001; Moore et al., 2003; Hintsa et al., 2021). However, the limited amount of AirCore sample (here, <250 mL of air) requires an alternative approach to load air into the GC-ECD. To achieve this goal, an in-house sample handling system was specially designed and built to capture and inject sample gas from 145 the AirCore into the GC-ECD system (Figure 1). The flow path in the sample handling system is controlled by a 6-port 2-position valve, and AirCore samples are carefully pushed at low pressure (approximately 300 hPa above ambient pressure) by an in-house standard gas cylinder with well-determined dry mole fractions of target molecules. The 6-port valve setup allows this "push gas" to also act as a calibration standard that can be directly injected into the GC-ECD periodically (through the bypass position, Figure 1). For each sample loading, the sample flow rate is controlled by a mass flow controller (Mykrolis, 150 MA, USA) at 4-5 mL per minute to maintain a stable pressure profile and constant flow. The flow rates during the sample loading process are then measured by a mass flow meter (Omega, CT, USA) at 1 Hz at the outlet of the StratoCore-GC-ECD system with a precision better than 0.6% (see section 2.3). The flow measurements associated with each sample loading process are integrated to calculate the total volume of air coming out of the AirCore for each measurement. The total sample volume data are used for registering the location of each measurement along the length of the AirCore (given a known total volume of 155 the AirCore) which is a crucial step for registering the GC-ECD measurements with altitude (Tans, 2022).

2.3 In-lab testing

To examine the potential contamination of target molecules from the AirCore, evaluate the mixing of air samples along the direction of flow during analysis, and assess the accuracy of AirCore volume registration, a series of tests were conducted with the StratoCore-GC-ECD system. The AirCore used in the tests shares the same material and surface coating 160 as the AirCores currently used by NOAA/GML for measuring atmospheric vertical profiles of CO_2, CH_4, and CO (Karion et

al., 2010; Baier et al., 2018). It consists of a 25 m long 304-grade stainless steel tube, with an outer diameter of 0.32 cm and an inner diameter of 0.29 cm. The inside wall of the tubing was treated with SilcoNert coating (SilcoTek Corp., PA, USA).

In the first test, we conduct storage tests of air inside the AirCore to determine whether the AirCore tubing surfaces contaminate the air sample. The AirCore was first flushed with zero-grade air, then stored overnight (~14 hours), and subsequently analyzed using the StratoCore-GC-ECD system, during which a standard gas of typical tropospheric composition was used as the push gas following the stored sample. The analytical results are shown in Figure 4A: the dry mole fractions of all target molecules were below the StratoCore-GC-ECD detection limit for these species in the entire AirCore. A similar test was then conducted: the AirCore was first flushed with the standard gas of tropospheric composition, stored overnight (~14 hours), then zero air was used as a push gas to analyze the AirCore using the StratoCore-GC-ECD system. The results (Figure 4B) show that none of the target molecules measured demonstrated any significant change in value after 14 hours of storage. Considering the storage time of actual AirCore samples (the time from AirCore landing to analysis) is usually within 4 hours, these results show that the AirCore sampler surfaces do not contaminate air samples during regular AirCore flights.

The two tests also demonstrated limited mixing of the sample during the analysis. The push gas used in both tests differed from the gas in the AirCore, defining the transition between the AirCore sample and the push gas after each analysis. The measurements display sharp transitions from the sample to the push gas within 1-2 injections (<10 mL of air). These abrupt transitions indicate that due to the carefully controlled flow during the sample loading process and minimized pressure drop during valve switching, the sample mixing during the analysis is minimal. Here, we used a simple mixing model to estimate the molecular diffusion and Taylor dispersion between the push gas and the sample gas in the AirCore:

$$X_{rms} = (2 * D_{eff} * t)^{0.5}, \qquad \text{Equation 1}$$

in which X_{rms} is the root-mean-squared diffusion distance, t is time, and D_{eff} is an effective diffusivity incorporating both the molecular diffusion and the Taylor dispersion:

$$D_{eff} = D + \frac{a^2 v^2}{48D}. \qquad \text{Equation 2}$$

In Equation 2, D is the molecular diffusion coefficient, a is the tube radius, and v is the average air velocity. Using the equations, we modelled the diffusion between the push gas and the sample gas during the analysis period (~1 hour) for the test AirCore, shown in Figure 4C. The modelled X_{rms} on the back end of the AirCore was 37.4 cm (equivalent to 2.5 mL of air), in line with the observed sharp transition between the push gas and the sample gas. Therefore, we suggest that the flow in the StratoCore-GC-ECD system during sample analysis remains a rigid "slug flow" moving through the system.

Additionally, the results from the two tests demonstrated high accuracy of volume registration by the StratoCore-GC-ECD system. Using the mass flow meter, we registered each data point to the volume of the AirCore. The measured volume of the AirCore by the StratoCore-GC-ECD system is defined as the midpoint of the transition between the AirCore gas and the push gas. In the meantime, we carefully measured the true volume of the AirCore multiple times by weighing the amount of water needed to fill the entire AirCore. The total volume of the AirCore measured by the StratoCore-GC-ECD system agrees with the actual volume of the AirCore within ± 1 mL, suggesting the volume measurements are accurate to within ±0.6%.

A final set of tests was also conducted to further verify that using a simple 1-D diffusion model, we can quantify the diffusion of the sample in the AirCore over a known storage time, and minimal mixing occurred during the sample loading process. The AirCore was filled with two alternating slugs using two calibrated dry standard gas cylinders. The two standards have different dry mole fractions of all the target molecules and CO_2, CO, and CH_4, therefore the transition between the two slugs can be observed using both a continuous analyzer that measures the CO_2 mole fraction, and the StratoCore-GC-ECD. During the filling process, the transitions between slugs in the AirCore (Karion et al., 2010) were directly measured by a continuous gas analyzer (G2401-m, Picarro, CA, USA), shown as the varying CO_2 mole fractions in Figure 5A. Subsequently, the AirCore was closed, quickly connected to the StratoCore-GC-ECD system then analyzed immediately. The N_2O mole fractions show the transitions between slugs measured by the StratoCore-GC-ECD (Figure 5A). The transitions observed by both the direct continuous measurements and the StratoCore-GC-ECD display good agreement, again suggesting mixing induced by StratoCore-GC-ECD's sample loading process is negligible. In addition, the AirCore was filled with two alternating slugs (the same as the previous test), then stored for 26 hours before being analyzed by the StratoCore-GC-ECD. After the longer, 26-hr storage, the StratoCore-GC-ECD measurements suggest that the mixing between slugs was significantly enhanced (Figure 5B). Assuming a 1-D diffusion model along the length of the AirCore, we used Equation 1 to model the molecular diffusion inside the AirCore during storage. The modelled diffusive mixing between slugs can capture the observed mixing well (Figure 5B), suggesting that the horizontal mixing in the tubing coil during the storage time between AirCore landing and sample analysis can be calculated and accounted for in the calculation of the uncertainty of the AirCore altitude registration.

3. Balloon-borne AirCore flights

Four balloon-borne AirCore test flights were conducted to retrieve the vertical profiles of CFC-11, CFC-12, CFC-113, H-1211, N_2O, and SF_6 in the stratosphere. The balloons were launched in North-eastern Colorado, USA on September 8, 2021 (flight 1), November 16, 2021 (flight 2), March 31, 2022 (flight 3), and August 9, 2022 (flight 4), respectively. A 3000 g weather balloon filled with helium carried the flight train on each flight to ~ 29.5 km AMSL. The payload package in the test flights is identical to routine AirCore flights of the ongoing AirCore Program at NOAA/GML (Figure 6). Each package contains a parachute; a cutter that can be remotely controlled to release the balloon and open the parachute; a GPS transmitter for real-time tracking of the flight, an Automatic Dependent Surveillance-Broadcast (ADS-B) transponder; a radiosonde (InterMet systems iMet-1 RSB, MI, USA) for recording latitude, longitude, altitude, temperature, atmospheric pressure, and relative humidity; and two AirCores. The two AirCores used in each flight are identical to ensure the sampling processes of both AirCores are the same and confirm that there is no contamination to the AirCore samplers. In flights 1-3, the AirCores consisted of 91-meter-long tubing, with an outer diameter of 0.32 cm and inner diameter of 0.29 cm (total volume: 600 mL), with a valve placed on each end (top and bottom valves). The AirCores used in flight 4 (August 9, 2022) were optimized for stratospheric air sampling (see discussion below) and consisted of two tubing segments: the tropospheric portion (open end

when sampling) was a 21-meter tube with an inner diameter of 0.58 cm, and the stratospheric portion was a 28-meter tube with an inner diameter of 0.29 cm (740 mL total volume with approximately 25% in the thinner tubing). Before the flight, each AirCore is insulated using a polymer foam package, wrapped with plastic wrap to minimize damage in the field, and covered by a custom-made bag using high-strength, lightweight Dyneema composite fabric. A data logger (Arduino, MA, USA) placed next to the AirCore inside the polymer foam recorded the coil temperature at multiple locations, latitude, longitude, altitude, and atmospheric pressure at 1 Hz.

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Hours before each flight, AirCores are flushed with special gases to distinguish between the residual fill gas not evacuated from the AirCore during flight and atmospheric samples during analysis. The AirCores used for CO_2, CH_4, and CO measurements (by continuous analyzers) are flushed with an air mixture with a high CO dry mole fraction (1765 ppb) and known CO_2 and CH_4 dry mole fractions. The AirCores analyzed on the StratoCore-GC-ECD system are flushed with zero-air containing elevated H-1211 (6.6 ppt). We selected H-1211 as the tracer for the remaining fill gas because of the rapid photochemical destruction of H-1211 in the lower stratosphere: atmospheric models and *in situ* aircraft observations (Portmann et al., 1999; Papanastasiou et al., 2013; Moore et al., 2014; Elkins et al., 2020) showed H-1211 is entirely destroyed above 25 hPa (~20-22 km AMSL). Since the topmost sampling altitude of our AirCore is higher than 25 km AMSL, H-1211 is an ideal tracer for distinguishing the residual fill gas in the AirCore from air samples collected above 24-25 km AMSL. After the AirCores are thoroughly flushed, the bottom valves are kept closed until minutes before the flight to minimize potential contamination from ambient air.

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The flight trajectories of all four test flights are shown in Figure 7. The balloon setup is designed such that the payload for all the flights should have similar ascent and descent processes: the balloon carries the payload to 28.9 km AMSL (10 hPa) at ~6 m/s, then the cutter is activated to release the payload from the balloon. After the balloon cutaway, the parachute is deployed and the payload descends, during which the AirCores passively collect ambient air. One exception was flight 3: after the balloon cutaway, the parachute did not fully open, resulting in a faster descent rate. After landing, the bottom valve on the AirCore closes automatically after ~30 seconds to minimize sample loss and potential contamination and loss of sample due to warming. The AirCores were quickly transferred back to the lab for analysis. In each flight, one of the AirCores was analyzed by a Picarro-2401-m continuous gas analyzer for CO_2, CH_4, and CO dry mole fractions, and the other AirCore was analyzed by the StratoCore-GC-ECD (here, only the stratospheric portion (approximately the first 20-30% of the AirCore) was analyzed). After the analysis, the filling process of the AirCore during descent is modelled using the meteorological data and a fluid dynamic program (Tans, 2022). The modelled results are then used to register the sample measurement time series with the altitude at which each sample was collected to derive the vertical profiles of all the trace gas measured by both instruments.

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The AirCore dimensions have a significant impact on the sampling efficiency in the stratosphere (e.g., Membrive et al., 2017; Baier et al., 2023). The AirCores used in flight 4 had wider tubing (0.58 cm inside diameter) at the bottom (open end) and thinner tubing on top (0.29 cm inside diameter), while the AirCores used in flights 1-3 consisted of one piece of thin tubing (0.29 cm inside diameter). As a result, their sampling efficiencies in the stratosphere are drastically different. Using the fluid dynamic model described in Tans (2022), we modelled the outflow and inflow of AirCores during each flight, shown in

Figure 8. The model suggests that after the balloon cutaway, the volumetric inflow during flight 4 (Figure 8A) increased much more rapidly compared to flights 1-3. The higher inflow in flight 4 is due to the larger diameter on the bottom portion of the AirCores used in flight 4, which produces a much smaller pressure gradient along the length of the tubing, making it easier for air to enter the sampler. Therefore, the stratospheric sampling efficiency, i.e., the ability of the AirCore to collect stratospheric air (Figure 8B), of flight 4 was significantly higher than those of flights 1-3. The increased flow rates in flight 3 (compared to flights 1 and 2) were caused mainly by its fast descent, creating a large pressure gradient at the inlet of the AirCores. However, the combination of high flow rates and short descent time (60% faster than other flights) in flight 3 still resulted in higher flow resistance in the AirCore which reduced sampling efficiency. Indeed, the model shows that compared to the AirCores in flights 1 and 2, the AirCores in flight 4 are much closer to pressure equilibrium between the closed and open ends of the AirCore (Figure 8C), while the AirCores in flight 3 displayed the most significant imbalance between open and closed ends during descent. This is demonstrated by the observed pressure differential between the open and closed ends of AirCore: the pressure differential of the entire AirCore in every flight was measured at 1 Hz by a pressure transducer mounted on the closed end of the AirCore. Using the model output, we also calculated the overall pressure differential and compared it with the measurements. The modelled pressure differential (Figure 8D) time series agrees with measurements in all the flights (data not shown in the figure): during the AirCore descent, the root-mean-square error (RMSE) between model output and measurements is less than 0.6 hPa (corresponding to approximately 240 m at 28 km AMSL and 20 m at 12 km AMSL), demonstrating that the model successfully reproduced the air sampling process. The AirCores used in flight 4 showed a much smaller pressure differential compared to the AirCores used in flights 1-3, highlighting the higher stratospheric sampling efficiency of the AirCore using modified tubing coil.

The StratoCore-GC-ECD analysis of AirCores from the four test flights yielded high vertical resolution profiles of target molecules, agreeing well with their predicted stratospheric photochemical loss processes. For all the target molecules measured by the StratoCore-GC-ECD system (Figure 9A-F), analyzing the 600 mL AirCores (in flights 1-3) produces 31 to 38 stratospheric measurements from each AirCore, equivalent to one measurement every 5-7 hPa. The larger dual-diameter AirCore used in flight 4 yielded 50 measurements in the stratosphere with a resolution of 4.5 hPa per measurement (corresponding to approximately 1.6 km per measurement at 28 km AMSL and 0.14 km per measurement at 12 km AMSL). The decrease in dry mole fractions of the photolytic tracers with altitude can be explained by their stratospheric photochemical properties (Portmann et al., 1999; Moore et al., 2014): compared to mean tropospheric values, the average loss (in %) of each photolytic tracer at the 650 K isentrope (approximately 25 km AMSL) for the four test flights are 58±5%, 61±4%, 72±3%, 93±3%, and 100% for N_2O, CFC-12, CFC-113, CFC-11, and H-1211, respectively. These values qualitatively agree with the relative stratospheric loss of the tracers via photolysis (e.g., Moore et al., 2013): at any given altitude in the stratosphere, the photolysis lifetime of each molecule, from longest to shortest, are N_2O > CFC-12 > CFC-113 > CFC-11 > H-1211. In addition, the high-resolution analysis from the StratoCore-GC-ECD systems captured temporal stratospheric variability, such as the variable dry mole fractions of all molecules at 10-17 km in Flight 3 (Figure 9). Similar structures were also observed in the CO_2 and CH_4 profiles obtained from the continuous analyzers using the other AirCore on the same flight string (Figure 9G,

H). Therefore, the observed variability in the AirCore profiles is unlikely to originate from artifacts during the sampling or measurement processes but reflects short-term atmospheric conditions that may have developed from episodic stratospheric dynamic events. The observed variability, which is significantly larger than detection limits and calculated mixing in the AirCore, suggests that the StratoCore-GC-ECD system allows us to obtain high vertical resolution observations of the stratosphere.

The tracer-tracer relationships in the profiles collected to date show agreement with *in situ* observations from previous flight campaigns within analytical uncertainties. The relationships between different trace gases are shown in Figure 10 and compared with *in situ* aircraft measurements using the UAS Chromatograph for Atmospheric Trace Species (UCATS, Hintsa et al., 2021) during the NASA Dynamics and Chemistry of the Summer Stratosphere (DCOTSS) campaign in Kansas, United States in July-August 2021. The AirCores collected samples from a higher altitude (25 – 28 km AMSL), where there is more aged air and more pronounced photolytic loss of trace gases compared to the ER-2 research aircraft (up to 21 km AMSL). The relationships between CFC-11, CFC-12, and H-1211 measured from AirCores agree with those of UCATS (Figure 10A, B). For CFC-113, the StratoCore-GC-ECD measurements generally agree with UCATS measurements with a small but consistent (1-2 ppt) discrepancy (Figure 10C). We speculate this discrepancy originates from different analytical methods used to calibrate working standards for the two measurements: the standards used in the UCATS measurement were calibrated using a GC-ECD. In contrast, the standards used in StratoCore-GC-ECD were calibrated using a GC-MS. A previous study showed a 1.3 ppt offset in CFC-113 measurements between a GC-ECD and a GC-MS (Rhoderick et al., 2015). It is possible that the observed discrepancy between StratoCore-GC-ECD and UCATS measurements reflected a similar offset. Further work is needed to understand the origin of this offset. SF_6 measurements from StratoCore-GC-ECD have been corrected to account for their growth in the troposphere using the global average growth rate of tropospheric SF_6 in 2021, and the corrected N_2O - SF_6 relationship also shows general agreement with UCATS data (Figure 10D) with a small offset. We suggest that this small offset might originate from the uncertainty in estimating the tropospheric growth rate of SF_6. One exception is the upper-most AirCore samples in flight 3, which displays a slightly different N_2O-SF_6 relationship from the other three flights. We speculate that this might be due to the short-term stratospheric transport variability, which is most likely driven by a combination of seasonal changes in wave activity, quasi-biennial oscillation (QBO), El Niño–Southern Oscillation (ENSO), and other short-term, episodic events. Mapping this variation between AoA and photolytic loss-dominated tracers allows us to investigate these drivers of stratospheric dynamics further but is outside the scope of this analysis. As we accumulate additional data in further flights, we can likely distinguish between these short-term variations and long-term changes driven by climate change.

4. Conclusions

The StratoCore-GC-ECD system, with a specially designed AirCore sample handling system (capable of injection of 4-5 mL of air for each analysis), can measure a suite of long-lived trace gases (N_2O, SF_6, CFC-11, CFC-12, H-1211, and CFC-113) from AirCore samplers with analytical precisions below 0.7% for all gases. AirCore samplers designed with dimensions

specially optimized for stratospheric sampling can obtain high-resolution vertical profiles of these trace gases from the tropopause to 28 km AMSL. Four test AirCore flights were conducted in eastern Colorado from Fall 2021 to Summer 2022, with AirCores analyzed by the StratoCore-GC-ECD system. The results showed good agreement with model predictions and aircraft in situ measurements, suggesting that the StratoCore-GC-ECD system provides a robust, low-cost approach for

330 observing the chemical composition of the stratosphere. In the future, this system will be applied for regular monitoring of the changes of these trace gases in the stratosphere, providing additional observational constraints on global climate models in a changing climate. We suggest that the sample handling system of the StratoCore-GC-ECD can be adapted to other analytical techniques to allow even more measurements (such as isotopic measurements) from AirCore samples in the future.

Data availability

335 Data used in this work is available at: https://doi.org/10.15138/VA4C-CY20 (Li et al., 2023).

Author contribution

Conceptualization: JL, FM, and BCB; Data curation: JL, BCB, FM, TN, SW, JH, GD, EH, and BH; Formal analysis: JL, FM, and BCB; Funding acquisition: FM and BCB; Investigation: JL, BCB, FM, TN, SW, JH, GD, and EH; Methodology: JL, BCB, FM, TN, SW, JH, GD, and CS; Project administration: BCB, CS, and BH; Resources: BCB, FM, GD, CS, and BH;

340 Supervision: BCB, CS, and BH; Validation: BCB, FM, and EH; Visualization: JL, and FM; Writing – original draft preparation: JL; Writing – review & editing: JL, BCB, FM, TN, SW, JH, GD, EH, CS, and BH.

Competing interests

The authors declare that they have no conflict of interest.

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References

Andrews, A. E., Boering, K. A., Daube, B. C., Wofsy, S. C., Loewenstein, M., Jost, H., Podolske, J. R., Webster, C. R., Herman, R. L., and Scott, D. C.: Mean ages of stratospheric air derived from in situ observations of CO_2, CH_4, and N_2O, Journal of Geophysical Research: Atmospheres, 106, 32295–32314, 2001.

Aris, R.: On the dispersion of a solute in a fluid flowing through a tube, Proc R Soc Lond A Math Phys Sci, 235, 67–77, 1956.

Baier, B., Sweeney, C., Andrews, A. E., Fleming, E. L., Higgs, J., Hedelius, J., Kiel, M., Parker, H. A., Newberger, T., and Tans, P. P.: AirCore profiles of greenhouse and trace gas species for sites within the Total Carbon Column Observing Network (TCCON) and comparisons to models and ground-based Fourier Transform Spectrometer (FTS) retrievals, AGU Fall Meeting Abstracts, A34C-02, 2018.

Butchart, N.: The Brewer-Dobson circulation, Reviews of geophysics, 52, 157–184, 2014.

Butchart, N. and Scaife, A. A.: Removal of chlorofluorocarbons by increased mass exchange between the stratosphere and troposphere in a changing climate, Nature, 410, 799–802, 2001.

Butchart, N., Scaife, A. A., Bourqui, M., de Grandpré, J., Hare, S. H. E., Kettleborough, J., Langematz, U., Manzini, E., Sassi, F., and Shibata, K.: Simulations of anthropogenic change in the strength of the Brewer–Dobson circulation, Clim Dyn, 27, 727–741, 2006.

Butchart, N., Cionni, I., Eyring, V., Shepherd, T. G., Waugh, D. W., Akiyoshi, H., Austin, J., Brühl, C., Chipperfield, M. P., and Cordero, E.: Chemistry–climate model simulations of twenty-first century stratospheric climate and circulation changes, J Clim, 23, 5349–5374, 2010.

Dietmüller, S., Garny, H., Plöger, F., Jöckel, P., and Cai, D.: Effects of mixing on resolved and unresolved scales on stratospheric age of air, Atmos Chem Phys, 17, 7703–7719, 2017.

Elkins, J. W., Fahey, D. W., Gilligan, J. M., Dutton, G. S., Baring, T. J., Volk, C. M., Dunn, R. E., Myers, R. C., Montzka, S. A., and Wamsley, P. R.: Airborne gas chromatograph for in situ measurements of long-lived species in the upper troposphere and lower stratosphere, Geophys Res Lett, 23, 347–350, 1996.

Elkins, J. W., Hintsa, E. J., and Moore, F. L.: ATom: Measurements from the UAS Chromatograph for Atmospheric Trace Species (UCATS), ORNL DAAC, 2020.

Engel, A., Möbius, T., Bönisch, H., Schmidt, U., Heinz, R., Levin, I., Atlas, E., Aoki, S., Nakazawa, T., Sugawara, S., Moore, F., Hurst, D., Elkins, J., Schauffler, S., Andrews, A., and Boering, K.: Age of stratospheric air unchanged within uncertainties over the past 30 years, Nat Geosci, 2, 28–31, https://doi.org/10.1038/ngeo388, 2009.

Engel, A., Bönisch, H., Ullrich, M., Sitals, R., Membrive, O., Danis, F., and Crevoisier, C.: Mean age of stratospheric air derived from AirCore observations, Atmos Chem Phys, 17, 6825–6838, 2017.

Fehsenfeld, F. C., Goldan, P. D., Phillips, M. P., and Sievers, R. E.: Selective electron-capture sensitization, in: Journal of Chromatography Library, vol. 20, Elsevier, 69–90, 1981.

380   Forster, P. M. de F. and Shine, K. P.: Assessing the climate impact of trends in stratospheric water vapor, Geophys Res Lett, 29, 10–11, 2002.

Garcia, R. R. and Randel, W. J.: Acceleration of the Brewer–Dobson circulation due to increases in greenhouse gases, J Atmos Sci, 65, 2731–2739, 2008.

Garny, H., Birner, T., Bönisch, H., and Bunzel, F.: The effects of mixing on age of air, Journal of Geophysical Research:
385   Atmospheres, 119, 7015–7034, 2014.

Gerber, E. P., Butler, A., Calvo, N., Charlton-Perez, A., Giorgetta, M., Manzini, E., Perlwitz, J., Polvani, L. M., Sassi, F., and Scaife, A. A.: Assessing and understanding the impact of stratospheric dynamics and variability on the Earth system, Bull Am Meteorol Soc, 93, 845–859, 2012.

Hall, T. M., Waugh, D. W., Boering, K. A., and Plumb, R. A.: Evaluation of transport in stratospheric models, Journal of
390   Geophysical Research: Atmospheres, 104, 18815–18839, 1999.

Hegglin, M. I. and Shepherd, T. G.: Large climate-induced changes in ultraviolet index and stratosphere-to-troposphere ozone flux, Nat Geosci, 2, 687–691, 2009.

Hintsa, E. J., Moore, F. L., Hurst, D. F., Dutton, G. S., Hall, B. D., Nance, J. D., Miller, B. R., Montzka, S. A., Wolton, L. P., and McClure-Begley, A. et al: UAS Chromatograph for Atmospheric Trace Species (UCATS)–a versatile instrument for trace
395   gas measurements on airborne platforms, Atmos Meas Tech, 14, 6795–6819, 2021.

Holton, J. R., Haynes, P. H., McIntyre, M. E., Douglass, A. R., Rood, R. B., and Pfister, L.: Stratosphere-troposphere exchange, Reviews of geophysics, 33, 403–439, 1995.

Karion, A., Sweeney, C., Tans, P., and Newberger, T.: AirCore: An innovative atmospheric sampling system, J Atmos Ocean Technol, 27, 1839–1853, 2010.

400   Laube, J. C., Elvidge, E. C. L., Adcock, K. E., Baier, B., Brenninkmeijer, C. A. M., Chen, H., Droste, E. S., Grooß, J.-U., Heikkinen, P., and Hind, A. J.: Investigating stratospheric changes between 2009 and 2018 with halogenated trace gas data from aircraft, AirCores, and a global model focusing on CFC-11, Atmos Chem Phys, 20, 9771–9782, 2020.

Leedham Elvidge, E., Bönisch, H., Brenninkmeijer, C. A. M., Engel, A., Fraser, P. J., Gallacher, E., Langenfelds, R., Mühle, J., Oram, D. E., and Ray, E. A.: Evaluation of stratospheric age of air from CF_4, C_2F_6, C_3F_8, CHF_3, HFC-125, HFC-227ea and
405   SF_6; implications for the calculations of halocarbon lifetimes, fractional release factors and ozone depletion potentials, Atmos Chem Phys, 18, 3369–3385, 2018.

Li, J., Baier, B., Dutton, G., Hall, B., Higgs, J., Hintsa, E., Moore, F., Newberger, T., Sweeney, C., Wolter, S., & NOAA Global Monitoring Laboratory. (2023). Vertical profiles of stratospheric CFC-11, CFC-12, CFC-113, H-1211, N_2O, SF_6 mole fractions acquired from four AirCore flights in Eastern Colorado using StratoCore-GC-ECD [Data set]. NOAA GML.
410   https://doi.org/10.15138/VA4C-CY20

Loeffel, S., Eichinger, R., Garny, H., Reddmann, T., Fritsch, F., Versick, S., Stiller, G., and Haenel, F.: The impact of sulfur hexafluoride (SF_6) sinks on age of air climatologies and trends, Atmos Chem Phys, 22, 1175–1193, 2022.

McLandress, C. and Shepherd, T. G.: Simulated anthropogenic changes in the Brewer–Dobson circulation, including its extension to high latitudes, J Clim, 22, 1516–1540, 2009.

415 Membrive, O., Crevoisier, C., Sweeney, C., Danis, F., Hertzog, A., Engel, A., Bönisch, H., and Picon, L.: AirCore-HR: a high-resolution column sampling to enhance the vertical description of CH_4 and CO_2, Atmos Meas Tech, 10, 2163–2181, 2017.

Moore, F. L., Elkins, J. W., Ray, E. A., Dutton, G. S., Dunn, R. E., Fahey, D. W., McLaughlin, R. J., Thompson, T. L., Romashkin, P. A., and Hurst, D. F.: Balloonborne in situ gas chromatograph for measurements in the troposphere and stratosphere, Journal of Geophysical Research: Atmospheres, 108, 2003.

420 Moore, F. L., Ray, E. A., Rosenlof, K. H., Elkins, J. W., Tans, P., Karion, A., and Sweeney, C.: A cost-effective trace gas measurement program for long-term monitoring of the stratospheric circulation, Bull Am Meteorol Soc, 95, 147–155, https://doi.org/10.1175/BAMS-D-12-00153.1, 2014.

Mrozek, D. J., van der Veen, C., Hofmann, M. E. G., Chen, H., Kivi, R., Heikkinen, P., and Röckmann, T.: Stratospheric Air Sub-sampler (SAS) and its application to analysis of $\Delta^{17}O$ (CO_2) from small air samples collected with an AirCore, Atmos

425 Meas Tech, 9, 5607–5620, 2016.

Pan, L. L., Bowman, K. P., Atlas, E. L., Wofsy, S. C., Zhang, F., Bresch, J. F., Ridley, B. A., Pittman, J. v, Homeyer, C. R., and Romashkin, P.: The stratosphere–troposphere analyses of regional transport 2008 experiment, Bull Am Meteorol Soc, 91, 327–342, 2010.

Papanastasiou, D. K., Carlon, N. R., Neuman, J. A., Fleming, E. L., Jackman, C. H., and Burkholder, J. B.: Revised UV

430 absorption spectra, ozone depletion potentials, and global warming potentials for the ozone-depleting substances CF_2Br_2, CF_2ClBr, and CF_2BrCF_2Br, Geophys Res Lett, 40, 464–469, 2013.

Ploeger, F., Riese, M., Haenel, F., Konopka, P., Müller, R., and Stiller, G.: Variability of stratospheric mean age of air and of the local effects of residual circulation and eddy mixing, Journal of Geophysical Research: Atmospheres, 120, 716–733, 2015.

Portmann, R. W., Brown, S. S., Gierczak, T., Talukdar, R. K., Burkholder, J. B., and Ravishankara, A. R.: Role of nitrogen

435 oxides in the stratosphere: A reevaluation based on laboratory studies, Geophys Res Lett, 26, 2387–2390, 1999.

Randel, W. J., Wu, F., Voemel, H., Nedoluha, G. E., and Forster, P.: Decreases in stratospheric water vapor after 2001: Links to changes in the tropical tropopause and the Brewer-Dobson circulation, Journal of Geophysical Research: Atmospheres, 111, 2006.

Ray, E. A., Moore, F. L., Elkins, J. W., Dutton, G. S., Fahey, D. W., Vömel, H., Oltmans, S. J., and Rosenlof, K. H.: Transport

440 into the Northern Hemisphere lowermost stratosphere revealed by in situ tracer measurements, Journal of Geophysical Research: Atmospheres, 104, 26565–26580, 1999.

Ray, E. A., Moore, F. L., Elkins, J. W., Rosenlof, K. H., Laube, J. C., Röckmann, T., Marsh, D. R., and Andrews, A. E.: Quantification of the SF_6 lifetime based on mesospheric loss measured in the stratospheric polar vortex, Journal of Geophysical Research: Atmospheres, 122, 4626–4638, 2017.

445   Ray, E. A., Moore, F. L., Rosenlof, K. H., Davis, S. M., Sweeney, C., Tans, P., Wang, T., Elkins, J. W., Bönisch, H., and Engel, A.: Improving stratospheric transport trend analysis based on SF_6 and CO_2 measurements, Journal of Geophysical Research: Atmospheres, 119, 14–110, 2014.

Rhoderick, G. C., Hall, B. D., Harth, C. M., Kim, J. S., Lee, J., Montzka, S. A., Mühle, J., Reimann, S., Vollmer, M. K., and Weiss, R. F.: Comparison of halocarbon measurements in an atmospheric dry whole air sampleInter-comparison of

450   halocarbons in an atmospheric dry whole air sample, Elementa: Science of the Anthropocene, 3, 2015.

Romashkin, P. A., Hurst, D. F., Elkins, J. W., Dutton, G. S., Fahey, D. W., Dunn, R. E., Moore, F. L., Myers, R. C., and Hall, B. D.: In situ measurements of long-lived trace gases in the lower stratosphere by gas chromatography, J Atmos Ocean Technol, 18, 1195–1204, 2001.

Schoeberl, M. R., Sparling, L. C., Jackman, C. H., and Fleming, E. L.: A Lagrangian view of stratospheric trace gas

455   distributions, Journal of Geophysical Research: Atmospheres, 105, 1537–1552, 2000.

Stiller, G. P., Clarmann, T. von, Haenel, F., Funke, B., Glatthor, N., Grabowski, U., Kellmann, S., Kiefer, M., Linden, A., and Lossow, S.: Observed temporal evolution of global mean age of stratospheric air for the 2002 to 2010 period, Atmos Chem Phys, 12, 3311–3331, 2012.

Stiller, G. P., Fierli, F., Ploeger, F., Cagnazzo, C., Funke, B., Haenel, F. J., Reddmann, T., Riese, M., and Clarmann, T. von:

460   Shift of subtropical transport barriers explains observed hemispheric asymmetry of decadal trends of age of air, Atmos Chem Phys, 17, 11177–11192, 2017.

Strahan, S. E., Loewenstein, M., and Podolske, J. R.: Climatology and small-scale structure of lower stratospheric N2O based on in situ observations, Journal of Geophysical Research: Atmospheres, 104, 2195–2208, 1999.

Tans, P. P. (2009). System and method for providing vertical profile measurements of atmospheric gases. Google Patents.

465   Tans, P.: Fill dynamics and sample mixing in the AirCore, Atmos Meas Tech, 15, 1903–1916, 2022.

Volk, C. M., Elkins, J. W., Fahey, D. W., Salawitch, R. J., Dutton, G. S., Gilligan, J. M., Proffitt, M. H., Loewenstein, M., Podolske, J. R., and Minschwaner, K.: Quantifying transport between the tropical and mid-latitude lower stratosphere, Science (1979), 272, 1763–1768, 1996.

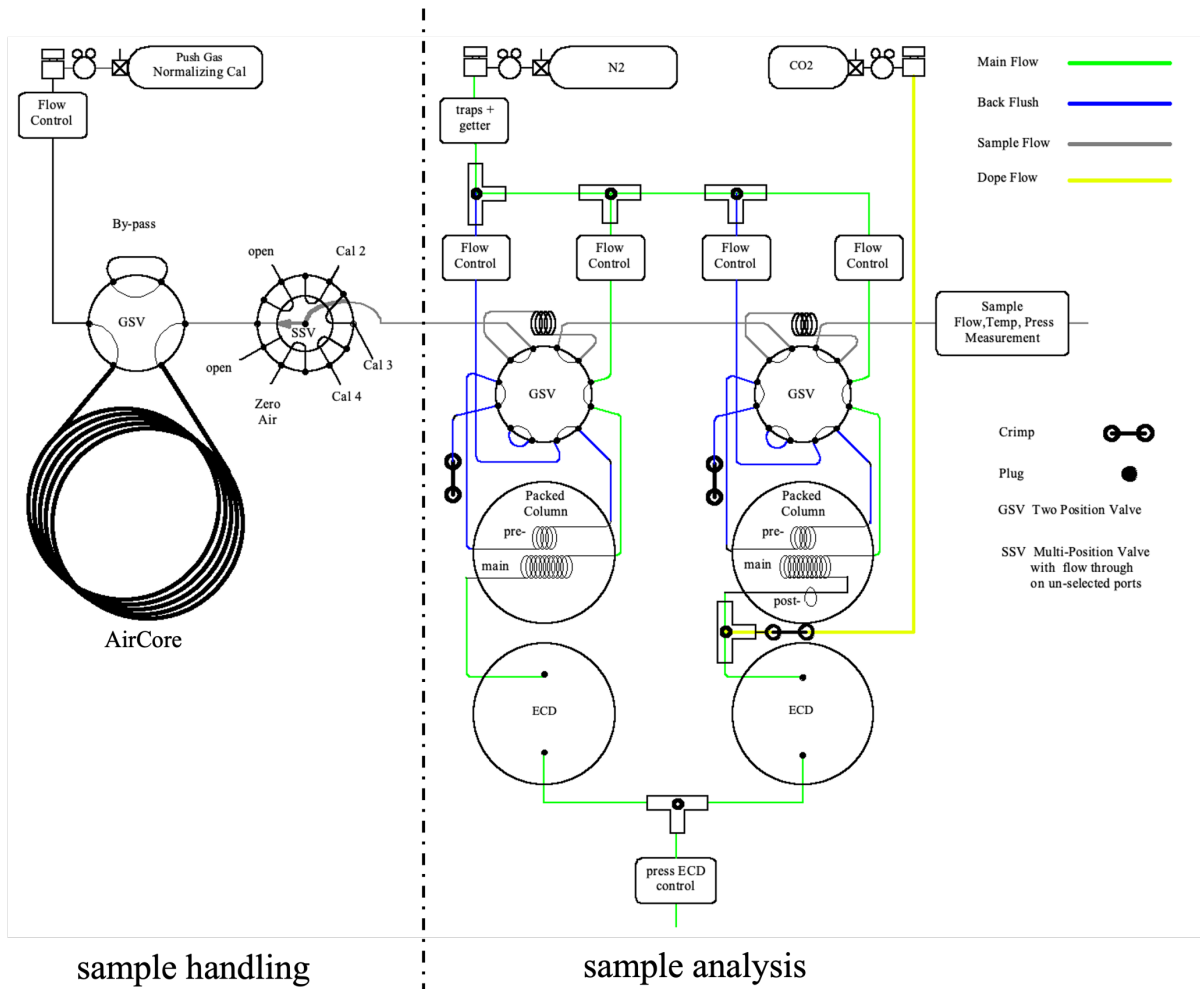
Waugh, D. and Hall, T.: Age of stratospheric air: Theory, observations, and models, Reviews of Geophysics, 40, 1, 2002.

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Figure 1: Simplified sketch of the StratoCore-GC-ECD system. The dashed line marks the boundary between the sample handling system and the sample analysis system. Left of the dashed line is the sample handling system, which carefully injects the sample from AirCores (4-5 mL per analysis) into the GC-ECD. Right of the dashed line is the sample analysis system, which measures the mole fractions of six trace gases in each injection.

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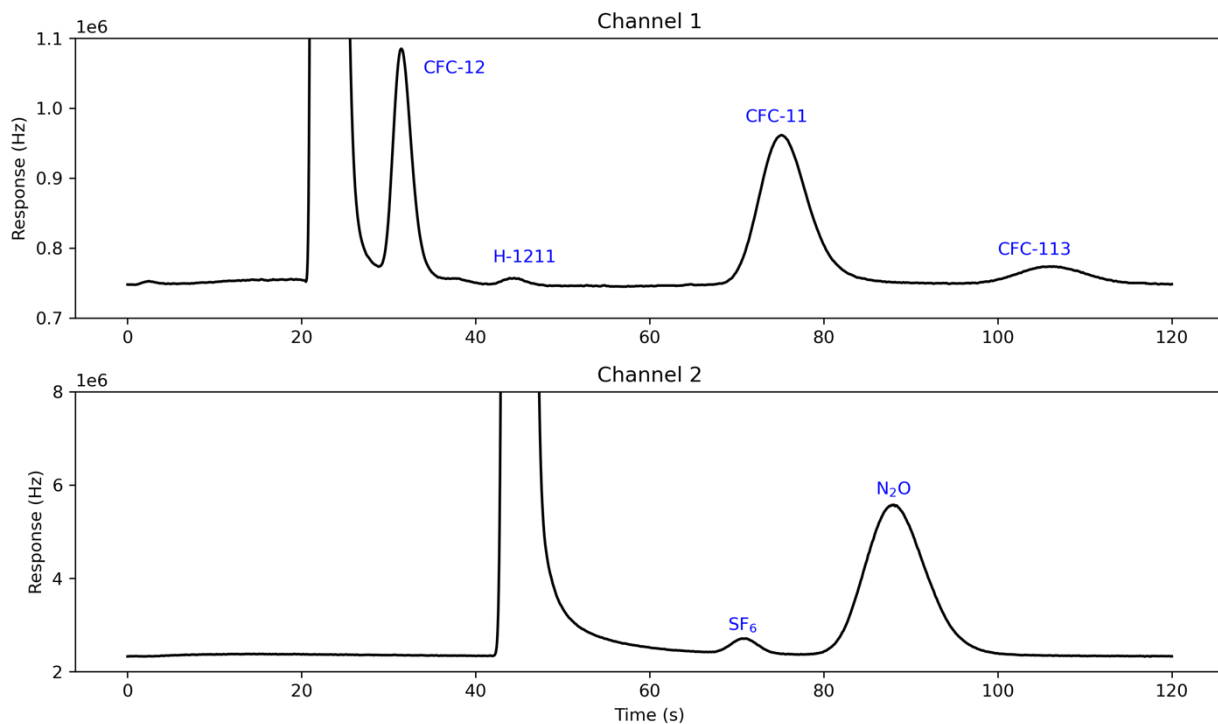
Figure 2: A typical chromatogram from StratoCore-GC-ECD analysis. The x-axis is the retention time of each analysis, the y-axis is the response of the ECD. The top panel is the response of channel 1 (analyzing the mole fractions of CFC-12, H-1211, CFC-11, and CFC-113) and the bottom panel is the response of channel 2 (analyzing the mole fractions of N_2O and SF_6).

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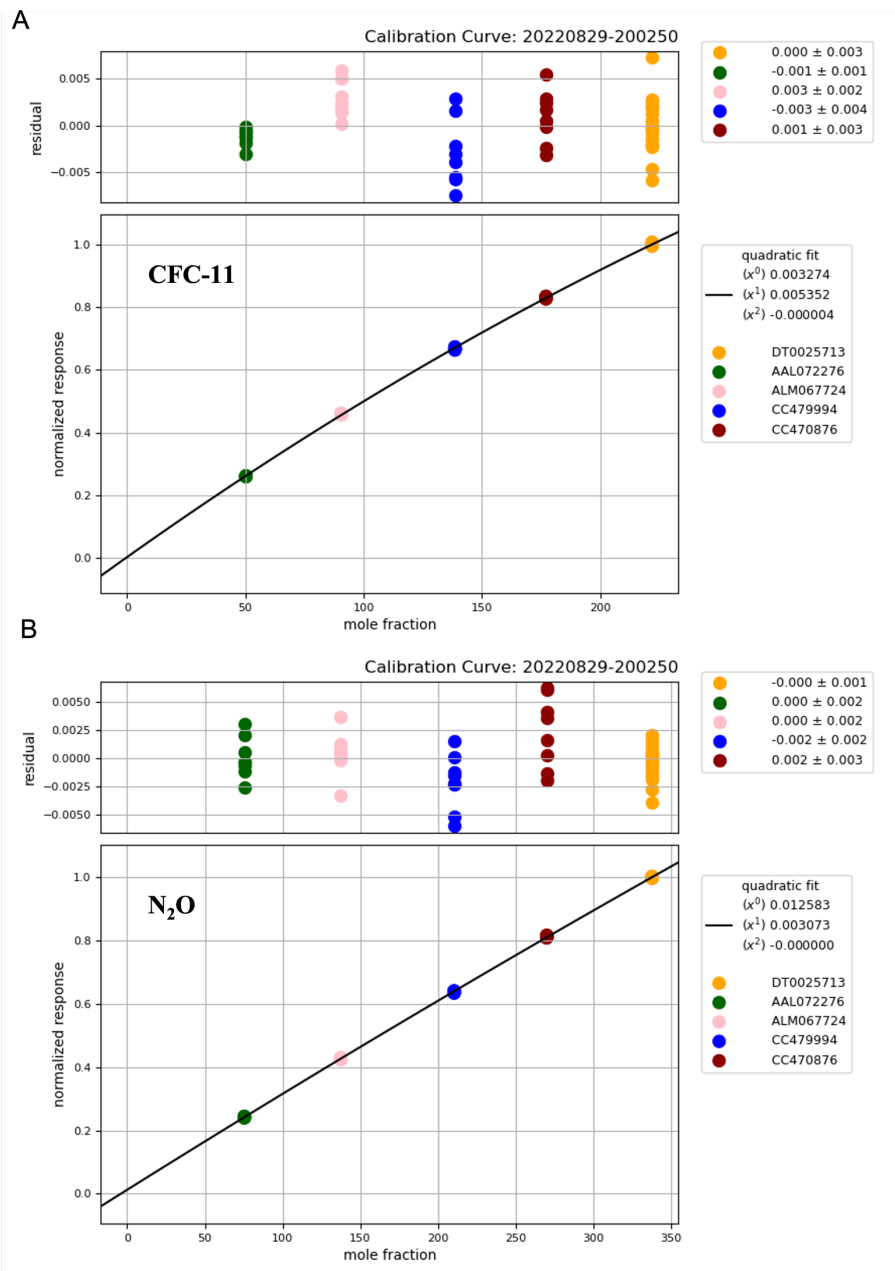
Figure 3: Examples of calibration curves (A: CFC-11; B: N₂O) generated by analyzing five standard tanks using the StratoCore-GC-ECD system. Each color represents a different calibration tank, and each tank was measured a total of seven times. In each panel, the upper figure shows the relative residual mole fraction (unitless) between the measured value and the true curve, and the legend shows the mean residual of each tank; the lower figure shows the calibration curve and the parameters of the quadratic fit function; the legend shows the standards used.

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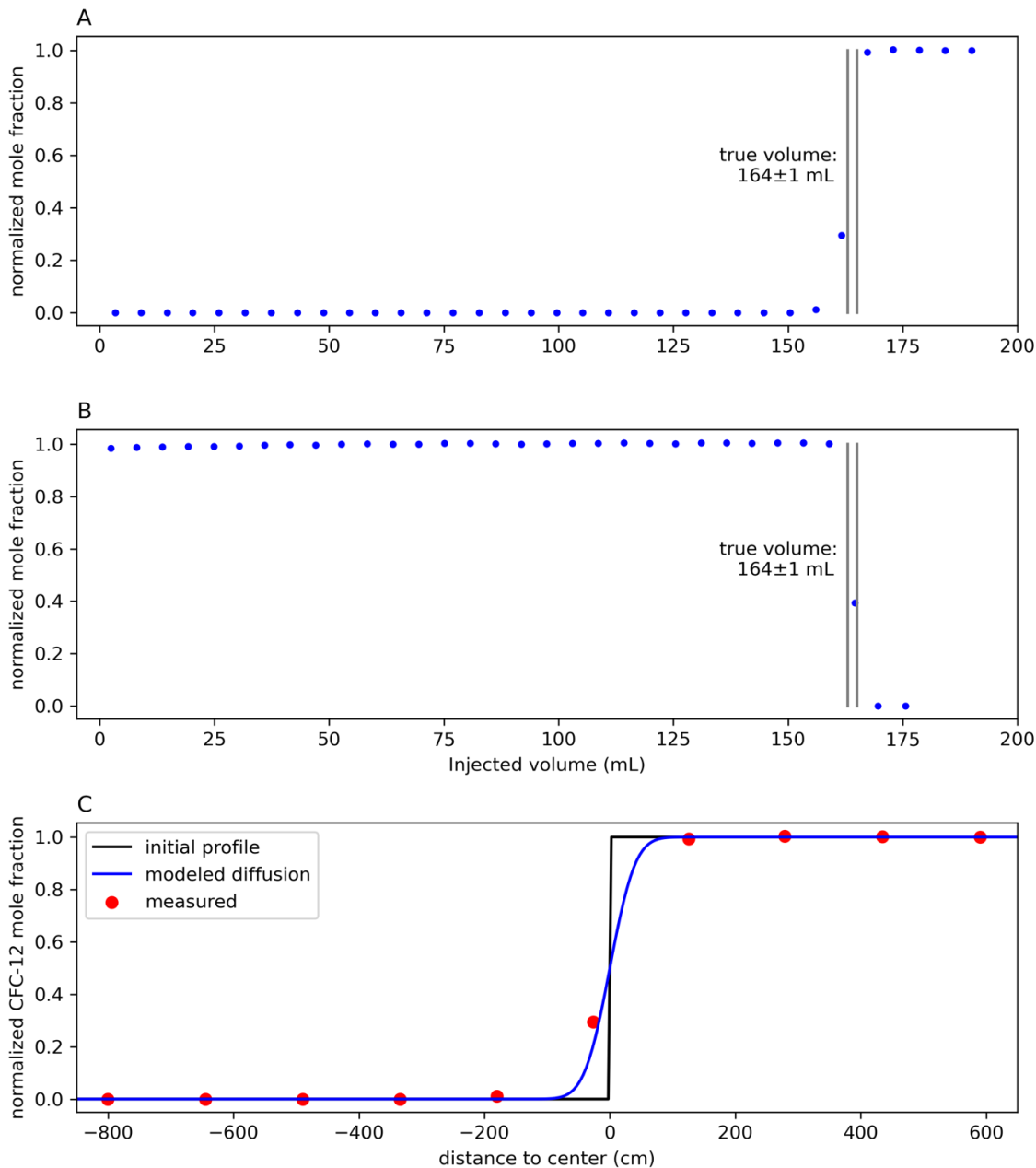
Figure 4: Results of StratoCore-GC-ECD flow-through tests from AirCore samples (showing CFC-12 as an example). A: filling the AirCore with zero air, then analyzing using StratoCore-GC-ECD with air with tropospheric mole fractions as the push gas. B: filling the AirCore with tropospheric air, then analyzing the AirCore using StratoCore-GC-ECD with zero air as push gas. The grey lines in A and B represent the true volume of the test AirCore. C: modelled diffusion at the AirCore gas–push gas boundary during AirCore analysis after 12 hours.

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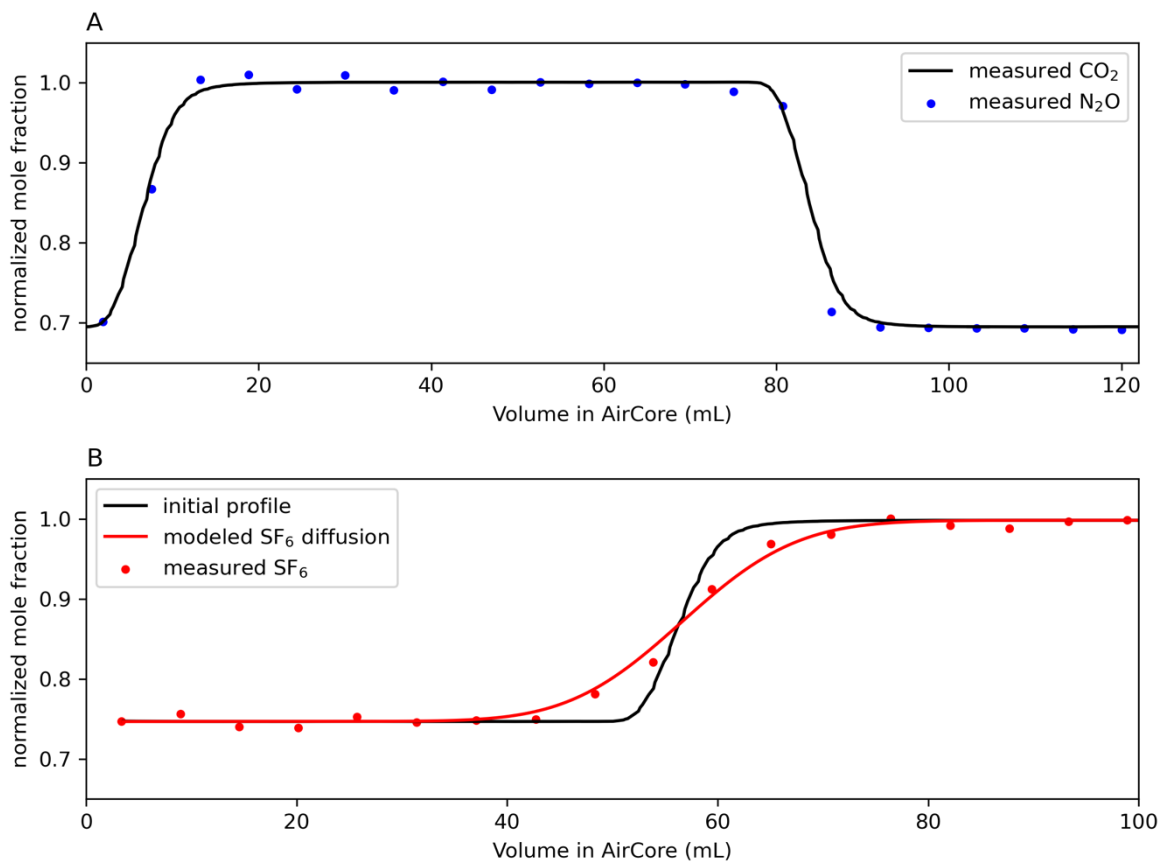
Figure 5: A: results from the alternating-slug test. The AirCore was filled by two alternating slugs (with normalized CFC-11 dry mole fractions of 0.7 and 1, respectively), then immediately analyzed by StratoCore-GC-ECD. The black line represents the transition between two slugs measured by the continuous flow analyzer; blue points are StratoCore-GC-ECD measurements. B: Results of the storage test using the same alternating slugs. The black line represents the transition between two slugs when the AirCore is being filled; the red line represents the modelled 1-D diffusion after 26-hour storage; the red dots are the measurements from the StratoCore-GC-ECD.

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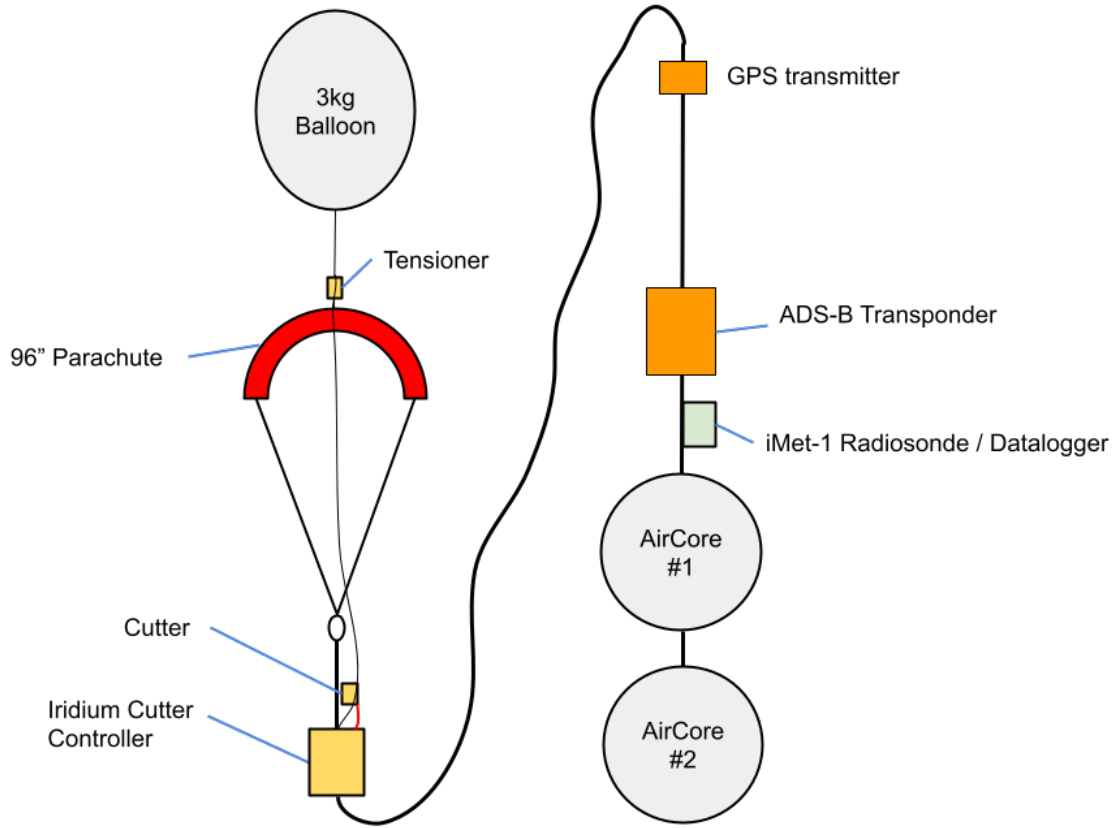
Figure 6: Components of a typical NOAA AirCore flight train.

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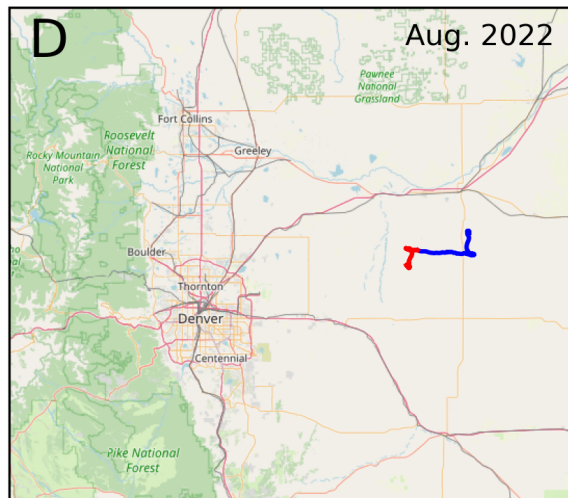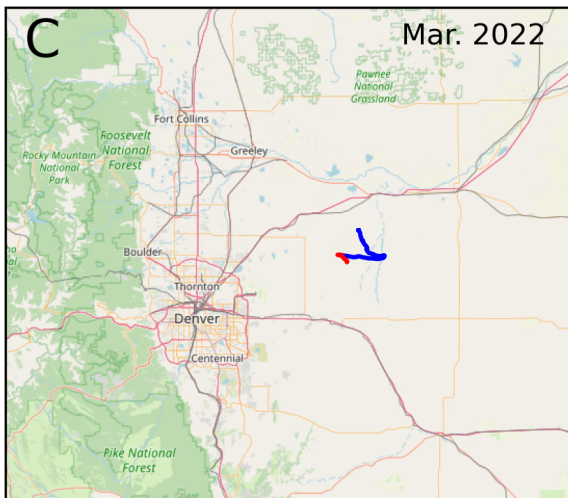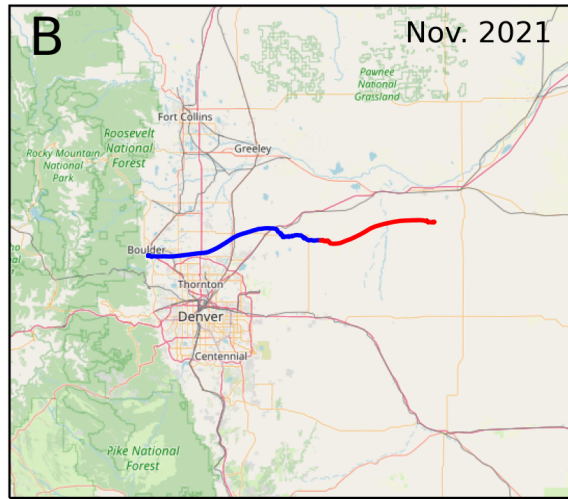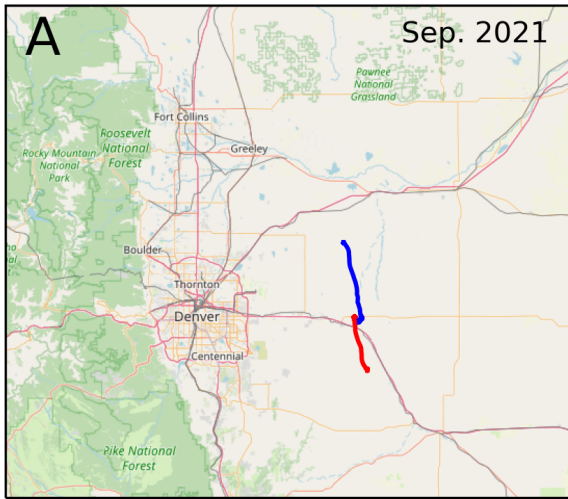
Figure 7: Flight trajectories of the four test flights in northeastern Colorado, USA. Panels A-D represent flights 1-4. The blue lines represent the payload ascent, and the red lines represent the descent. The base map is provided by © OpenStreetMap contributors 2023. Distributed under the Open Data Commons Open Database License (ODbL) v1.0.

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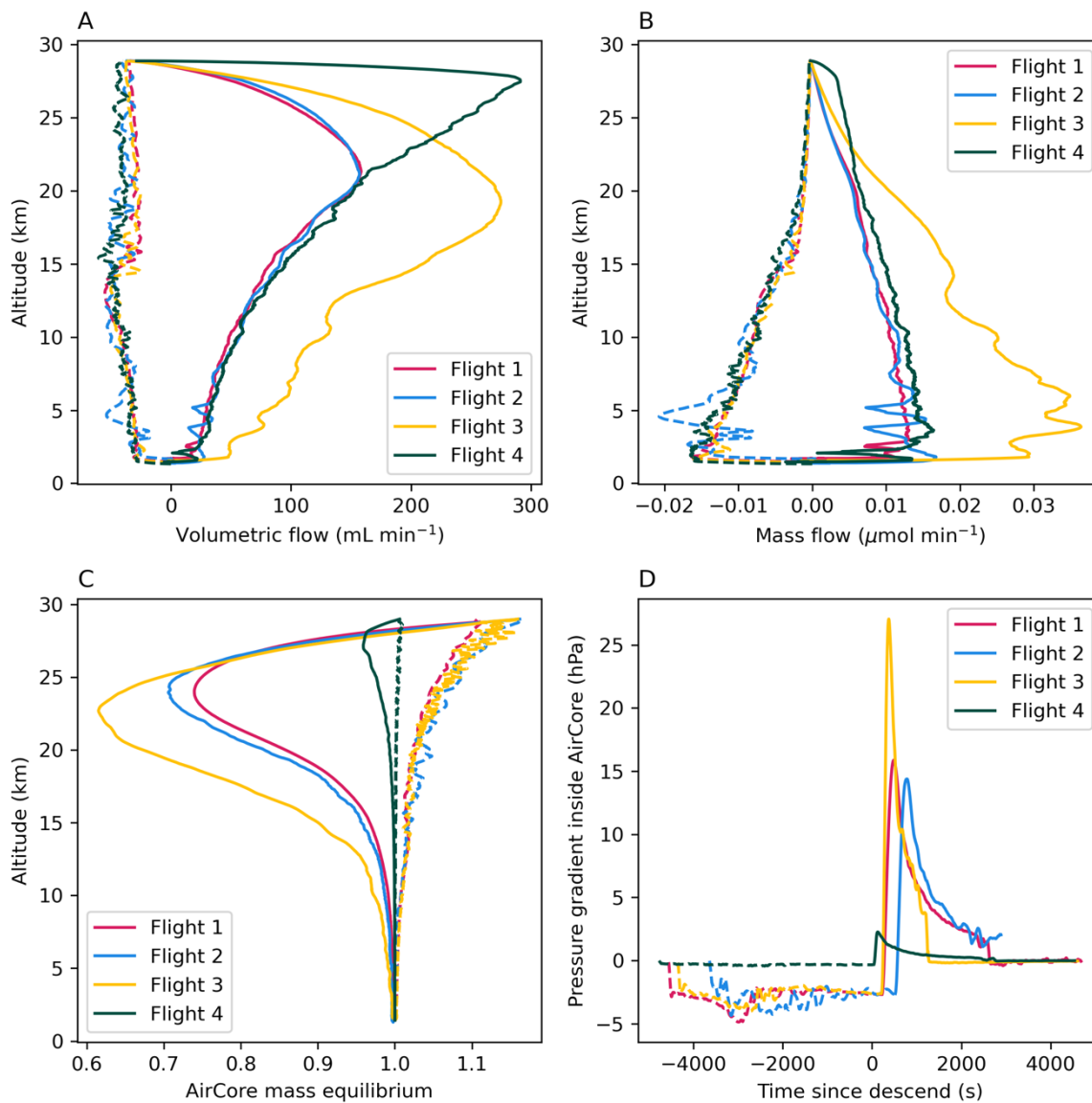
Figure 8: Modelled fluid dynamics of the AirCore filling process. Each line represents one flight, the dashed portion of each line represents the ascent, and the solid part represents the descent. A: the modelled volumetric flow (in mL min⁻¹) into the AirCores during the entire flight, plotted against altitude; B: the modelled mass flow (in µmol min⁻¹) into the AirCores; C: the AirCore mass equilibrium ratio (actual air mass divided by equilibrium air mass in the AirCore) during the entire flight: a mass equilibrium ratio equal to 1 means the air inside the AirCore reaches equilibrium with ambient air, a ratio lower than 1 means the air inside the AirCore is depleted and vice versa; D: time series of the modelled pressure gradient across the entire AirCore.

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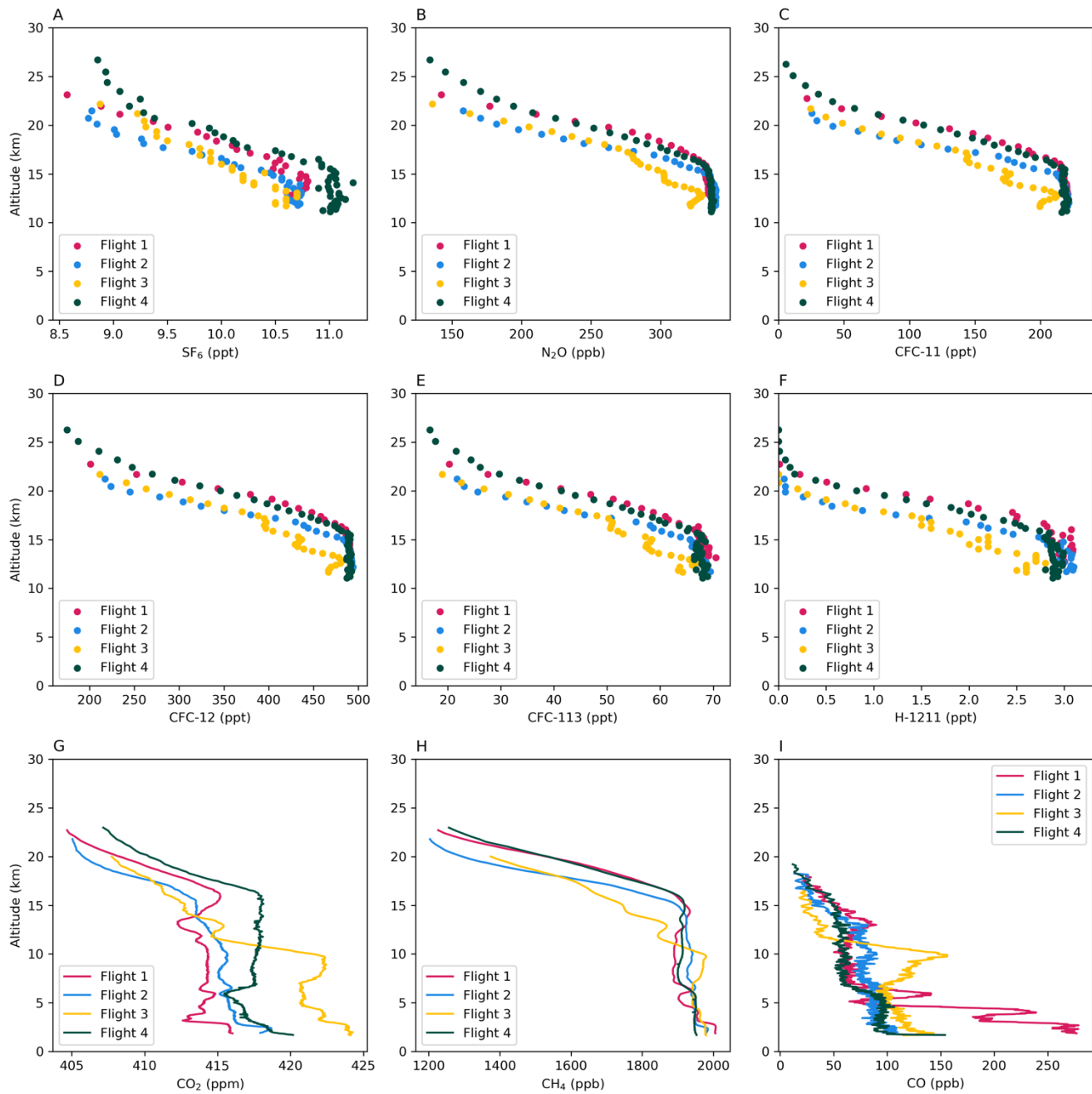
Figure 9: Vertical profiles of AirCore trace gas dry mole fractions measured from the 4 flights. Panels A-F are the results from the StratoCore-GC-ECD, from the tropopause to 25-28 km AMSL. Each color represents one of the flights. Panels G-I are the mole fractions of CO_2 (G), CH_4 (H), and CO (I) analyzed by a Picarro-2401 continuous gas analyzer.
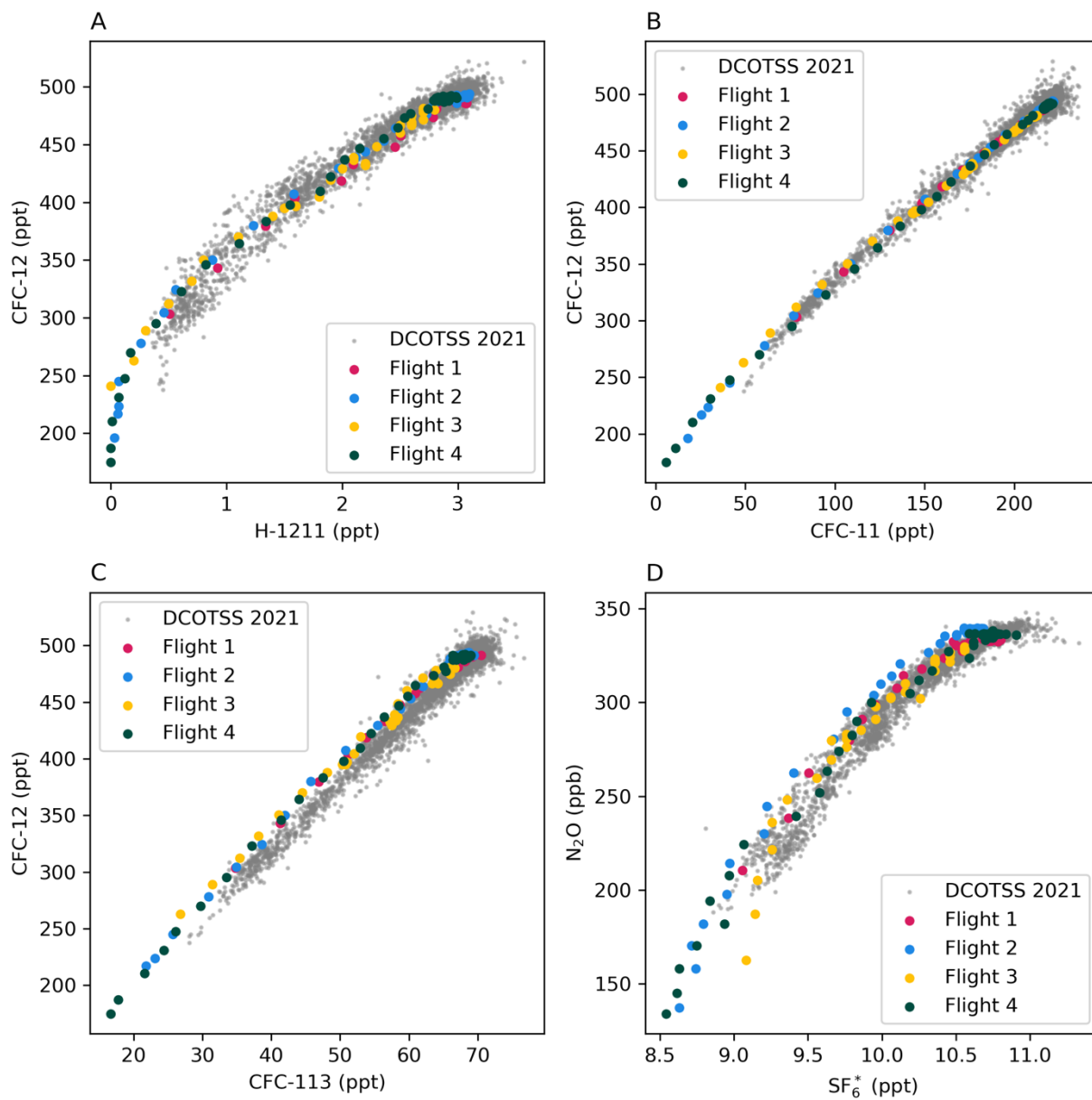
Figure 10: Relationships between different molecules measured by the StratoCore-GC-ECD for Flights 1-4. Grey points are aircraft measurements acquired from the 2021 NASA DCOTSS campaign over the central United States and Canada. *AirCore SF_6 data in panel D were corrected (based on the global average growth rate of SF_6 in 2020-2021) to account for the growth of SF_6 in the atmosphere.

530 **Table 1: Analytical repeatability of the StratoCore-GC-ECD, reported relative to the tropospheric mole fraction of each gas.**

Molecule	Analytical repeatability
CFC-11	±0.1 %
CFC-12	±0.1 %
CFC-113	±0.25 %
H-1211	±0.7 %
SF_6	±0.25 %
N_2O	±0.25 %