



## 1 Technical note: Surface fields for global environmental modelling

Margarita Choulga<sup>1</sup>, Francesca Moschini<sup>1</sup>, Cinzia Mazzetti<sup>1</sup>, Stefania Grimaldi<sup>2</sup>, Juliana
 Disperati<sup>3</sup>, Hylke Beck<sup>4</sup>, Peter Salamon<sup>2</sup>, Christel Prudhomme<sup>1</sup>

4 <sup>1</sup>European Centre for Medium-Range Weather Forecasts (ECMWF), Reading, RG2 9AX, United Kingdom

<sup>2</sup>Joint Research Centre (JRC), European Commission, Ispra, 21027, Italy

<sup>3</sup>Fincons Group, Vimercate, 20871, Italy
 <sup>4</sup>King Abdullah University of Science an

<sup>4</sup>King Abdullah University of Science and Technology (KAUST), Thuwal, Saudi Arabia

8 *Correspondence to*: Margarita Choulga (margarita.choulga@ecmwf.int) and Christel Prudhomme 9 (christel.prudhomme@ecmwf.int)

10 Abstract. Climate change has resulted in more frequent occurrences of extreme events, such as flooding and 11 heavy snowfall, which can have a significant impact on densely populated or industrialised areas. Numerical 12 models are used to simulate and predict these extreme events, enabling informed decision-making and planning 13 to minimise human casualties and protect costly infrastructure. LISFLOOD is an integrated hydrological model 14 underpinning the European and Global Flood Awareness Systems (EFAS and GloFAS, respectively) developed 15 by the Copernicus Emergency Management Service (CEMS). The CEMS\_SurfaceFields\_2022 dataset is a new 16 set of high-resolution surface fields at 1 and 3 arc min (approximately 2 and 6 km at the Equator respectively) 17 covering Europe and the global land surface (excluding Antarctica) respectively, based on a wide variety of high-18 resolution and up-to-date data sources. The dataset has been created together with upgrades to the open source 19 LISFLOOD code. The set encompasses (i) catchment morphology and river network, (ii) land use, (iii) vegetation 20 21 cover type and properties, (iv) soil properties, (v) lake information, and (vi) water demand. This manuscript details the complete workflow to generate CEMS\_SurfaceFields\_2022 fields, including data sources and methodology. 22 23 24 The use of these fields is expected to significantly improve accuracy, detail, and realism of LISFLOOD simulations. CEMS\_SurfaceFields\_2022 can also be used as input for other Earth system models or for carrying out general statistical analyses across various spatial scales, ranging from global and regional to local levels.

#### 25 1 Introduction

26 27 Current numerical Earth system models are highly complex. Thanks to the availability of High Performance Computers, cloud computing, and a wide range of high-resolution environmental data derived from the use of 28 29 30 ground, unconventional and satellite measurement sensors, numerical global models are even able to reach kilometre-scale horizontal resolution. But increase in spatial resolution also means that the Earth system and environmental models have to represent more surface and atmospheric processes and their interactions, which can 31 become challenging, for example in complex orographic areas. Model accuracy heavily depends on the quality of 32 the input surface fields (i.e. how realistic and up-to-date they are), and it is essential to minimise errors in surface 33 34 fields. New high-resolution (i.e. 10-100 m) surface datasets based on daily satellite observations are now frequently released and continuously supported by e.g. the Copernicus program (e.g. Global Land Cover: 35 Buchhorn et al., 2021; GHSL-BUILT-S: Pesaresi and Politis, 2022; Schiavina et al., 2022), which helps in 36 37 achieving the goal of minimising surface field errors. It was shown, e.g. in Kimpson et al. (2023, in review), that the use of accurate and up-to-date underlying information to generate model's input surface fields can substantially 38 39 reduce skin temperature errors even at 30 km horizontal resolution (Kimpson et al., 2023 in review).

Following the digital revolution of cloud archiving and computing, where data, software and IT infrastructure can 40 be accessed by anyone from everywhere, the Earth systems and environmental modelling community has also 41 moved from codes developed by a single organisation and few contributors, to so-called 'community models' 42 where a reference code is open for free use and/ or development according to sharing principles. Such models 43 44 include Joint UK Environmental Simulator JULES, a land surface model whose development is coordinated by the UK Met Office and UKCEH (Best et al., 2011; Clark et al., 2011; Marthews et al., 2022), OpenIFS, a 45 46 Numerical Weather Forecast model available to external users for research and training (Sparrow et al., 2021; Carver, 2022; Huijnen et al., 2022; Köhler et al., 2023), the Community Land Model CLM, an Earth System 47 48 Model with strong climate component maintained by the National Centre for Atmospheric Research but available for use by the wider research community (Lawrence et al., 2019), or LISFLOOD-OS, a spatially distributed water 49 resources model developed by the Joint Research Centre (JRC; Van Der Knijff and De Roo, 2008) and available 50 for use and development through a share code repository (https://ec-jrc.github.io/lisflood/#lisflood; https://ec-51 jrc.github.io/lisflood-code/).

To promote the seamless development of science, and facilitate research community efforts in working with the same code and input data, providing feedback, and improving the code and the data itself, powerful web-based





platforms can be used. One of them is the Google Earth Engine (GEE; Gorelick et al., 2017), a free-of-charge platform that provides easy, web-based access to an extensive catalogue of satellite imagery and other geospatial data in an analysis-ready format. The data catalogue is embedded into Google computing platform that lets you easily implement all personal workflows, which facilitates global-scale analysis and visualization (GEE: FAQ, 2023). GEE was chosen for the generation of a new vast surface field set due to its high resolution data catalogue and powerful computation capabilities.

60 This manuscript presents the methodology used to prepare the CEMS\_SurfaceFields\_2022 dataset containing all 61 surface fields necessary to run the LISFLOOD-OS model at 1 arc min (over Europe) and 3 arc min (globally). 62 CEMS\_SurfaceFields\_2022 can also be used in the set-up of the Early Warning Systems of the Copernicus 63 Emergency Management Service of the European Union for the European (European Flood Awareness System 64 EFAS version 5; Smith et al., 2016; https://www.efas.eu/) and global (Global Flood Awareness System GloFAS 65 version 4; Hirpa et al., 2018; Harrigan et al., 2023; https://www.globalfloods.eu/) domains expected to become 66 operational during 2023. The detailed explanation, encompassing raw data collection, scientific protocol, and 67 technical details, will allow the adequate understanding and interpretation of the surface field datasets (openly 68 available from the data catalogue of the JRC - for EFAS https://data.jrc.ec.europa.eu/dataset/f572c443-7466-69 4adf-87aa-c0847a169f23, for GloFAS https://data.jrc.ec.europa.eu/dataset/68050d73-9c06-499c-a441-70 dc5053cb0c86), with clear methodological protocols that can be replicated or adapted easily to prepare alternative 71 72 fields over a different geographical domain, spatial resolution or different content as relevant for downstream application. Finally, the resulting surface fields are expected to be a useful resource not only for hydrological 73 modelling but also for weather prediction, Earth system modelling, environmental modelling, or statistical 74 analysis in general, with a spatial scale allowing for global, regional and even national applications.

#### 75 2 Surface fields for distributed environmental modelling

76 Environmental models, especially land surface and hydrological models, simulate how water moves across 77 canopy, surface, subsurface, ground and eventually river channels using mechanistic equations that describe the 78 79 physics of these processes. Each model represents processes with more or less complexity, depending on the model purpose and expected output (Rosbjerg and Madsen, 2005). With most represented terrestrial processes 80 depending on the landscape, information describing the spatial variation in the geophysical and vegetation 81 characteristics is needed. Such characteristics include morphological features (e.g. channel geometry, orography 82 or slope), soil hydraulic property, land and vegetation features (e.g. ecosystem cover type, leaf area index (LAI), 83 evaporation rates, crop type, planting and harvesting dates), and if relevant, human intervention information such 84 as population density or type of water usage.

85 LISFLOOD is a semi-distributed, physically based hydrological model which has been designed for the modelling 86 of rainfall-runoff processes in large and transnational catchments (Bates and De Roo, 2000; De Roo et al., 2000; 87 De Roo et al., 2001; Van Der Knijff and De Roo, 2008; Van Der Knijff et al., 2010; Burek et al., 2013). In its 88 most prominent application, LISFLOOD is used by the Copernicus Emergency Management Services' EFAS and 89 GloFAS to provide medium range and seasonal riverine flow forecasts. LISFLOOD is also widely used for a 90 variety of applications, including water resources assessment (drought forecast); analysis of the impacts of land 91 use changes, river regulation measures, water management plans; climate change analysis (e.g. Vanham et al., 92 2021).

70 To facilitate users' uptake and enable the seamless development of science, LISFLOOD has been released as open source in 2019. The open-source suite includes the LISFLOOD hydrological model and a set of auxiliary tools for model setup, calibration, and post-processing of the results. For instance, the pre-processor LISFLOOD-LISVAP can be used to compute evapotranspiration, which is one of the three meteorological variables, along with total precipitation and average temperature, strictly required as input to the hydrological model.

98 The modelling of runoff processes in different climates and socio-economic contexts then requires a set of raster 99 fields to provide information of terrain morphology, surface water bodies, soil properties, land cover and land use 100 features, water demand. The total number of fields range between 66, when only the essential rainfall-runoff 101 processes are modelled, to a total 108 for a more comprehensive model set-up in which, for instance, lakes, 102 reservoirs, water demand for anthropogenic use are included (https://ec-jrc.github.io/lisflood-model/).

In this section, we introduce the main characteristics of environmental fields dataset produced, grouped according to their role in process representation (name in brackets next to each field correspond to the name in the data repository). The main model's technical field is 'mask' – a Boolean field that defines model boundaries, i.e. gridcells over which the model performs calculations and grid-cells which are skipped (e.g. ocean grid-cells). Whilst the fields described in this manuscript follow some specific requirements of the LISFLOOD model, they can be

108 used for any environmental modelling application, either directly, or following a transformation, as relevant.



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#### 109 2.1 Catchment morphology and river network

110 Morphology and channel shape information are essential for the computation of snow melting, temperature 111 scaling, and river routing. Land morphology is derived from elevation and its variability within a single cell can 112 be represented through slope, standard deviation, aspect, etc. River drainage information, derived from elevation, 113 is used to connect the model cells according to the direction of the surface runoff, with channel geometry 114 information used for routing processes.

115 The dataset contains 14 morphology and river network variables:
116 • Morphologic information: local drainage direction (i.e.

- <u>Morphologic information</u>: local drainage direction (i.e. flow direction from one cell to another; *LDD*, dimensionless), upstream area (*upArea*, m<sup>2</sup>), grid-cell area (*pixarea*, m<sup>2</sup>), grid-cell length (*pixleng*, m), standard deviation of elevation (*elvstd*, m), gradient (i.e. elevation gradient; *gradient*, m/m);
- <u>Kinematic wave equation for routing</u>: channel bottom width (*chanbw*, m), channel length (*chanlenght*, m), channel gradient (*changrad*, m/m), Manning's roughness coefficient for channels (*chanman*, s/m<sup>1/3</sup>);
- m), channel gradient (*changrad*, m/m), Manning's roughness coefficient for channels (*chanman*, s/m<sup>1/3</sup>);
   <u>River network information</u>: channel mask (i.e. presence of river channel; *chan*, dimensionless), channel
- side slope (i.e. channel's horizontal distance divided by vertical distance; *chans*, m/m);
  Open water evaporation: bankfull channel depth (*chanbnkf*, m), channel flood plain (i.e. width
  - <u>Open water evaporation</u>: bankfull channel depth (*chanbnkf*, m), channel flood plain (i.e. width of the area where the surplus of water is distributed when the water level in the channel exceed the channel depth; *chanflpn*, m).

#### 126 2.2 Land use fields

Land use is an essential component of environmental models. Many models use a sub-grid-cell approach where a
single grid-cell can include several different land uses with each land use being subject to different prominent
physical processes. This approach allows to keep a high level of accuracy when representing how different types
of land cover affect e.g. the hydrological cycle (e.g. evaporation is different in urban areas compared to forests)
while limiting the increase in computational time.

132 The dataset differentiates between six different land uses:
133 • Forest: areas where the main hydrological procession of the second s

- <u>Forest</u>: areas where the main hydrological processes are canopy interception, evapotranspiration from canopies, canopies drainage and evapotranspiration, root uptake and evaporation from the soil (fraction of forest; *fracforest*, dimensionless fraction);
- Sealed surface: impervious areas where there is no water infiltration into the soil, i.e. water is accumulated in the surface depression, yet evaporates, but once the depression is full, water is transported by a surface runoff (fraction of sealed surface; *fracsealed*, dimensionless fraction);
  - <u>Inland water</u>: open water bodies where the most prominent hydrological process is evaporation (fraction of inland water; *fracwater*, dimensionless fraction);
- Irrigated crops: areas used by agriculture water is abstracted from ground water and surface water bodies to irrigate the fields. The main hydrological processes connected with the irrigated crops are canopy interception, evapotranspiration from canopies, canopies drainage and evapotranspiration, root uptake and evaporation from the soil (fraction of all irrigated crops, excluding rice; *fracirrigated*, dimensionless fraction);
- Irrigated rice: areas used to grow rice with flooded irrigation agricultural technique, when water is abstracted from the inland water bodies and delivered to the rice fields. The main hydrological processes connected with rice fields are soil saturation, flooding, rice growing phase, soil drainage phase (fraction of irrigated rice; *fracrice*, dimensionless fraction);
- Other land cover: used in canopy interception, evaporation from the canopies, canopy drainage, plant
   evapotranspiration, evaporation from the soil hydrological processes. The relative importance of these
   processes depends on the LAI (fraction of other cover types; *fracother*, dimensionless fraction).

### 153 2.3 Vegetation properties

154 Vegetation-related information contributes to the computation of precipitation interception, evaporation, 155 transpiration, and root water uptake. Depending on the model, vegetation dynamics can be represented with 156 different degrees of complexity including in hydrology processes, vegetation growth and feedback on climate 157 (Bonan et al., 2002). Rice being the world's most important food crop and having specific water demands, its 158 water cycle is often considered explicitly, with planting and harvesting dates being critical information to represent 159 the inter-annual variability in its water demand, provided the maximum three growing seasons. The variables 160 allow to model how vegetation affects the hydrology, with a particular focus on root water uptake and transpiration 161 depending on vegetation type and vegetation state (e.g. water stress conditions). For example, the crop group



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162 number depends on the critical amount of soil moisture below which water uptake from plants is reduced as they 163 start closing their stomata.

- 164 The dataset describes vegetation properties through four variables (note that LAI consists in total of 36 10-day 165 average fields) for each of forest (\_f), irrigated crops (\_i) and other land cover types (\_o), and another six (two 166 types times three seasons) variables for rice: 167
  - <u>Transpiration rate</u>: crop coefficient (*cropcoef\_f*, *cropcoef\_i*, *cropcoef\_o*, dimensionless);
  - <u>Water uptake</u>: crop group number (*cropgrpn\_f*, *cropgrpn\_i*, *cropgrpn\_o*, dimensionless);
- 169 Surface runoff generation and water routing: Manning's surface roughness coefficient (mannings\_f, 170 mannings\_o, s/m1/3), rice planting and harvesting days (riceplantingday1, riceplantingday2, 171 riceplantingday3, calendar day number; riceharvestday1, riceharvestday2, riceharvestday3, calendar 172 day number);
- 173 Water interception and evaporation: leaf area index (laif, laii, laio, m<sup>2</sup>/m<sup>2</sup>).

#### 174 2.4 Soil properties

175 In land surface and distributed hydrological models, the water movement, storage and plants' water-uptake from 176 the soil are often described by the soil/ water retention curve (SWRC). The SWRC is derived empirically by 177 measuring how water is retained and released by different soil types. Throughout time different SWRC have been 178 developed and integrated into models, the most widely applied are Van Brooks and Corey (Brooks and Corey, 179 1964), Fredlund and Xing (Fredlund and Xing, 1994), van Genuchten (van Genuchten, 1980), and Gardner 180(Gardner, 1956) SWRCs. Different SWRC equations require different parameters, some shared between different 181 SWRC concepts, e.g. referring physical soil characteristics such as water saturated and unsaturated content, 182 hydraulic conductivity and pore size, others uniquely describing the SWRC function shape, not directly related to 183 soil properties. Often, for computational reasons, the soil profile from ground level to bedrock depth is sliced into 184 layers, at the modeller's choice, and the SWRC function is applied to each soil layer.

185 The dataset includes variables required to apply the Van Genuchten SWRC equations (van Genuchten, 1980) to 186 describe the water dynamics through a vertical soil profile composed of three layers (1, 2, 3), each variable is 187 required for each soil layer and for forest (\_f) or non-forest (\_o) land use, with different soil depth in forest (\_f) 188 and non-forest (\_o) areas following root depth values from Allen at al. (1998), further referred as FAO56, (total 189 of 29 variables; see Section 4.4 for detailed definition and calculation):

- 190 Soil profile: surface layer depth (soildepth1\_f, soildepth1\_o, mm), middle layer depth (soildepth2\_f, 191 soildepth2\_o, mm), subsoil depth (soildepth3\_f, soildepth3\_o, mm);
- 192 Soil hydraulic properties: saturated (thetas1\_f, thetas1\_o, thetas2\_f, thetas2\_o, thetas3, m3/m3) and 193 residual (*thetar1*, *thetar2*, *thetar3*, m<sup>3</sup>/m<sup>3</sup>) volumetric soil moisture content, pore size index (*lambda1\_f*, 194 lambda1\_o, lambda2\_f, lambda2\_o, lambda3, dimensionless), Van Genuchten equation parameter 195 (genual\_f, genual\_o, genua2\_f, genua2\_o, genua3, cm<sup>-1</sup>), saturated soil conductivity (ksat1\_f, ksat1\_o, 196 ksat2\_f, ksat2\_o, ksat3, mm/day).

#### 197 2.5 Lakes

198 Lakes (and reservoirs) are important as they influence the atmosphere regionally and globally as well as the river 199 discharge. The area covered by lakes is used for computing evaporation from open water surfaces. In LISFLOOD 200 the volume of evaporated water is not subtracted from the storage volume of lakes. Here the dataset only includes 201 data on lake extent and not reservoirs (generally smaller): lake mask (i.e. presence of lakes consistent with fraction 202 of inland water; lakemask, dimensionless).

#### 203 2.6 Water demand

204 Some environmental models explicitly represent a number of the human interventions impacting on the water 205 cycle. One of the most common is water demand, which represents the withdrawal of water from natural water 206 sources (e.g. rivers, reservoirs, groundwater) to satisfy the water demand for anthropogenic use. The segregation 207 of the total water demand for anthropogenic use into four main sectors, namely domestic, energy, industrial, and 208 livestock water withdrawal, enables a more accurate representation of the processes. Following the Food and 209 Agriculture Organisation of the United Nations (FAO) terminology (Kohli et al., 2012), domestic water 210withdrawal represents indoor and outdoor household water use as well as other uses (e.g. industrial and urban 211 agriculture) connected to the municipal system (e.g., water use by shops, schools, and public buildings). Electricity 212 (energy) water withdrawal is the water use for the cooling of thermoelectric and nuclear power plants. Water 213 withdrawal for industry is the water used for fabricating, processing, washing, cooling or transporting products, 214 also includes water within the final products and water used for sanitation within the manufacturing facility. 215 Livestock withdrawal is the demand for drinking and cleaning purposes of livestock.





Higher accuracy in environmental modelling is achieved by differentiating water demand sources and by allocating different levels of priority to different usages. Within LISFLOOD, for instance, water demand for the energy sector and flooded irrigation (rice crops) is supplied by surface water bodies only, while non-flooded irrigation, domestic, industrial, livestock water demand can be supplied by both groundwater and surface water bodies. Moreover, domestic water demand has the highest priority in case of water scarcity conditions.

221 It must be noted that the fields of water demand for agriculture are not included in this dataset because LISFLOOD 222 computes crops water demand internally by accounting for climatic conditions, information on land cover (see  $\bar{2}\bar{2}\bar{3}$ Section 2.2), crops properties (see Section 2.4), and soil properties (see Section 2.5). Conversely, fields 224 225 representing the volume of water to satisfy the domestic, energy, industrial, and livestock demand must be provided as input. Domestic, industrial, energy, and livestock water demand volumes have seasonal (e.g. due to 226 227 temperature differences) and inter-annual variations (e.g. due to population changes and different economic conditions). In order to account for this variability, in LISFLOOD the four sectoral water demand fields provide 228 daily water demand data with monthly or annual variability from 01.01.1979 to 31.12.2019: the water demand 229 values are provided in mm/day, one field per month (the first day of each month is used as representative 230 timestamp for the entire month) for domestic and energy demand, one value per year (the monthly fields are 231 repeated twelve times per each year) for industrial and livestock demand.

The dataset includes water demand for four main sectors (note that each sector consists in total of 12 daily water demand fields per 41 (1979-2019) years, so 492 fields per sector) for: livestock (*liv*, mm/day), industry (*ind*, mm/day), energy production, (*ene*, mm/day) and domestic use (*dom*, mm/day). The temporal extension of the water demand fields presented in this manuscript includes the most recent information of water demand at the time of the dataset's preparation. Readers that are interested in using more recent water demand data are invited to follow the protocol presented in Section 4.6 to further extend in time the provided fields.

### 238 2.7 Specific requirements for the dataset

239 The dataset produced follows the specific requirements of LISFLOOD for EFAS (European domain, 1 arc min

240 resolution at mid-latitude of the domain (47.50 N) is ~1.25 km) and GloFAS (Global domain) implementation,

- summarised in Table 1.
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#### Table 1. Dataset files technical specifications.

Туре	Specification
Format	NetCDF
Projection	EPSG:4326 - WGS84: World Geodetic System
Horizontal	Europe: 1 arc min (~1.86 km at the Equator) [file size 4530x2970 grid-cells]
resolution	Globe: 3 arc min (~5.57 km at the Equator) [file size 7200x3600 grid-cells]
Domain hound	Europe: [North = 72.25 N; South = 22.75 N; West = 25.25 W; East = 50.25 E]
Domain bound	Globe: [North = 90.00 N; South = 90.00 S; West = 180.00 W; East = 180.00 E]
Missing value (i.e.	Over land: none
NoData) location	Over ocean: all ocean grid-cells have missing value (i.e. ocean is masked based on 'mask' field)
Missing value (i.e.	For Integer variable type: 0
NoData) number	For Real variable type: -999999.0
Variable type	Integer: Int8
variable type	Real: Float32

#### 244 3 Reference data and overall methodology

245 This section describes all data sources used to produce dataset's surface fields introduced in Section 2. All data 246 considered were open source, freely available, updated as recently as possible, with recognised reference on their 247 quality. Note that whilst the majority of surface fields contain no time element, vegetation and water demand 248 fields explicitly describe the annual cycle (vegetation, rice) or annual time evolution (water demand) and therefore 249 have more stringent requirements regarding the data source. Global single-source datasets (e.g. Te Chow, 1959; 250 Supit et al., 1994; Allen et al., 1998; Buchhorn et al., 2021) were favoured to regional and/ or multiple data sources 251 that needed to be combined in order to produce the required data unless sub-set information was of much better 252 quality (e.g. Moiret-Guigand, 2021).





#### 253 **3.1 Catchment morphology and river network**

#### 254 **3.1.1 Digital Elevation Model**

255 The MERIT DEM: Multi-Error-Removed Improved-Terrain Digital Elevation Model v.1.0.3 [15 October, 256 2018] (further referred as MERIT DEM) is a high accuracy global DEM at 3 arc second resolution (~90 m at the 257 Equator) covering land area from 90 N to 60 S, selected for its ability to clearly represent landscapes such as river 258 networks and hill-valley structures even in flat areas where height errors could be larger than topography 259 variability (Yamazaki et al., 2017; Bhardwaj, 2021; Chai et al., 2022). It is derived from seven different open-260 source datasets, delivered as 57 GeoTiff files 30° by 30° region each, at ~90 m resolution (in total 90.0 GB), 261 representative of the year 2018. More detail on method, data content and access can be found in Yamazaki et al. 262 (2017) and MERIT DEM web-page http://hydro.iis.u-tokyo.ac.jp/~yamadai/MERIT\_DEM.

263 The MERIT DEM was used to compute standard deviation of elevation, gradient and channel geometry fields.

#### 264 3.1.2 Hydromorphology

265 The Catchment-based Macro-scale Floodplain (CaMa-Flood) Global River Hydrodynamics Model v4.0 266 maps (further referred as CaMa-Flood) are used for the basic maps describing all physical properties of the river 267 network. It is derived from MERIT Hydro (MERIT Hydro is a global hydrography dataset, created by using 268 elevation (i.e. MERIT DEM) and several inland water maps; more detail can be found in Yamazaki et al. (2019) 269 and MERIT Hydro web-page http://hydro.iis.u-tokyo.ac.jp/~yamadai/MERIT\_Hydro) for high resolution river 270 routing applications using the FLOW algorithm (Yamazaki et al., 2009; Yamazaki et al., 2011). The maps include 271 information on channel length, river topography parameters, floodplain elevation profile, channel width and 272 channel depth. The maps exist at 15, 6, 5, 3 and 1 arc min resolutions covering land area from 90 N to 60 S, 273 representative of the year 2017, and for each resolution, they are available as one single file with all variables in 274 NetCDF format (for 1 arc min 737.0 MB). More detail on method, data content and access can be found in 275 Yamazaki et al. (2011) and CaMa-Flood web-page http://hydro.iis.u-tokyo.ac.jp/~yamadai/cama-276 flood/index.html. Note that whilst the CaMaFlood maps where originally generated for the specific use of the 277 CaMa-Flood model, they can also serve as basic to derive alternative maps for other environmental models, as 278 done here.

The CaMa-Flood maps were used to create the local drainage direction (LDD), upstream area, channel geometryand land masks fields.

#### 281 3.2 Land use fields

#### 282 3.2.1 Land cover

283 The Copernicus Global Land Service (CGLS) Land Cover (LC) 100m map (further referred as CGLS-LC100) 284 is a global land cover map of the year 2015 (Buchhorn et al., 2020). It is derived from the PROBA-V 100 m 285 satellite image collection, a database of high quality land cover training sites and ancillary datasets, reaching an 286 accuracy of 80 % at Level1 (Buchhorn et al., 2021). It contains 23 classes for discrete classification and 10 classes 287 for continuous cover fractions; and it is delivered as 15 files in GeoTiff format (in total 39.3 GB) at 100 m 288 resolution covering land area from 90 N to 60 S and representative of the year 2015. More detail on method, data 289 content and access can be found in Buchhorn et al. (2021) and Copernicus web-site 290 https://land.copernicus.eu/global/products/lc.

The CGLS-LC100 was used to generate crop parameters and Manning's surface roughness coefficient for forest and other land cover types, to generate forest, inland water, and sealed surface fraction fields, following a basic quality check on large water bodies (i.e. correcting Fox Basin and Caspian Sea).

295 The Coordination of Information on the Environment (CORINE) Land Cover (CLC) inventory for 2018 296 (further referred as CLC2018) is a set of maps describing the land cover/ land use status of 2018 covering 297 39 countries in Europe with a total area of over 5.8 Mkm<sup>2</sup>. The dataset is derived from satellite imagery (mainly 298 Sentinel-2, based on a constellation of two satellites orbiting Earth at altitude of 786 km 180° apart revisiting 299 equator every 5 days, and for gap filling Landsat-8, making a constellation together with Landsat-9 satellite 300 orbiting Earth at altitude of 705 km each revisiting equator every 16 days) and in-situ data and contains 44 classes, 301 delivered as one GeoTiff raster file (125.0 MB) at 100m resolution covering land area over Europe, representative 302 of the time period 2017-2018. The overall accuracy for CLC2018 is 92 % for the blind analysis (i.e. validation 303 team had no knowledge of the CLC2018 thematic classes) but there are regional variations: the Black Sea 304 geographical region has the lowest accuracy of 84 %; country-wise overall accuracy vary from 86 % for Portugal 305 to 99 % for Iceland, lowest accuracy being linked to the landscape complexity (Moiret-Guigand, 2021). More

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- detail on method, data content and access can be found in Büttner and Kosztra (2017) and Moiret-Guigand (2021),
- 307 and Copernicus web-site <u>https://land.copernicus.eu/pan-european/corine-land-cover/clc2018</u>.
- 308 The CLC2018 was used to generate the irrigated crop fraction and rice fraction fields.
- 309 3.2.2 Crop cover

310 The Spatial Production Allocation Model (SPAM) – Global Spatially-Disaggregated Crop Production 311 Statistics Data for 2010 v2.0 (further referred as SPAM2010) is a global dataset generated in 2020, which 312 redistributes crop production information from country and sub-national provinces level to a finer grid-cell level 313 (IFPRI, 2019). It is derived from numerous data sources, including crop production statistics, cropland data, 314 biophysical crop "suitability" assessments, spatial distribution of specific crops or crop systems, and population 315 density. SPAM2010 contains estimates of crop distributions within disaggregated units (based on a cross-entropy 316 approach) for 42 crops and two production systems (irrigated and rainfed), and is delivered as 84 files in shapefile 317 format at 10 km (5 arc min) resolution covering land area from 90 N to 60 S and representative of the year 2010 318 (in total 2.2 GB). Based on crop expert judgement from international (i.e. International Rice Research Institute, 319 International Maize and Wheat Improvement Center) and national organisations (i.e. The Chinese Academy of 320 Agricultural Sciences) SPAM2010 over Europe and America is more accurate than over Africa and South East 321 Asia, with best performance in allocating rice; grid-by-grid comparison of crop areas with independent Cropland 322 Data Layer (produced by using satellite images and vast amount of ground truth) over continental United States 323 shows coefficient of determination ( $R^2$ ) 0.7-0.9 and root mean square error (RMSE) 231-307 ha indicating a 324 relatively high reliability, with highest R<sup>2</sup> and lowest RMSE values are for maize and soybean (Yu et al., 2020). 325 More detail on method, data content and access can be found in Yu et al. (2020) and MapSPAM web-site 326 https://mapspam.info.

SPAM2010 was used to compute the irrigated crop and rice fractions, crop parameters and Manning's surface
 roughness coefficient for irrigated crop fields.

#### 329 **3.3 Vegetation properties**

#### 330 **3.3.1 Crop properties**

The Food and Agriculture Organisation (FAO) of the United Nations Irrigation and Drainage Paper No. 56 (further referred as FAO56) is a publication covering geographically referenced statistics for crop development stages, crop coefficients, crop height, rooting depth, and soil water depletion fraction for common crops found across the world; it also covers procedures for information aggregation, e.g. on the grid. It is delivered as an article with a set of tables and equations and can be considered as the most complete source of information on crop properties. More detail on method and data content can be found in Allen et al. (1998) and FAO online crop information web-page http://www.fao.org/land-water/databases-and-software/crop-information/tobacco/en/.

FAO56 was used to compute the crop coefficients for forest, irrigated crops and other land cover types (online
 crop information was specifically used for tobacco); and for intermediate computations such as depletion fraction
 for different crop and surface types (table), crop height and root depth fields.

341 342

**Intara** et al. (2018) is a publication covering oil palm roots architecture.

- 343 Intara et al. (2018) was used for oil palm root depth information in addition to FAO56.
- 344

Burek et al. (2014) is a publication covering summarised information for crop coefficients, rooting depth, crop
 group number and Manning's surface roughness coefficient for different surface types.

Burek et al. (2014) was used for built-up, bare/ sparse vegetation, snow & ice, permanent inland water, ocean &
seas, herbaceous wetland, moss & lichen surface types crop coefficients, rooting depth, crop group number and
Manning's surface roughness coefficient information in addition to FAO56 and other sources.

350

The Wofost 6.0 crop simulation model description (further referred as SUPIT) is a publication on developing, validating, and testing new or already existing agrometeorological models (Supit et al., 1994). It contains crop group information for several crops as examples, and relation of a crop group from water depletion fraction. The publication is delivered as a book with a set of tables and equations. Information on crop group is still considered up-to-date. More detail on method and data content can be found in Supit et al. (1994).

356 SUPIT was used to compute the crop group fields for forest, irrigated crops and other land cover types.





#### 357 3.3.2 River hydraulics properties

358 The Open-Channel Hydraulics manual (further referred as CHOW) is a publication on open-channel hydraulics, including basic principles and different types of flows, i.e. uniform, gradually varied, rapidly varied, and unsteady (Te Chow, 1959). It contains information on roughness coefficient over different surfaces. The

gublication is delivered as a book with a set of tables and equations. More detail on method and data content can
 be found in Te Chow (1959).

363 CHOW was used to compute the Manning's surface roughness coefficient fields for forest, irrigated crops and 364 other land cover types.

### 365 3.3.3 Vegetation time evolution

366 The Copernicus Global Land Service (CGLS) Leaf Area Index (LAI) 1km Version 2 collection (further 367 referred as CGLS-LAI) is a set of global maps without missing data describing vegetation dynamics - the annual 368 evolution of LAI at 10-day intervals over the period of 1999-2020. The dataset is derived from 369 SPOT/VEGETATION and PROBA-V data. The dataset's root mean square deviations over 20 GBOV sites over 370 the period 2014-2018 is 0.92, compared to 1.19 for MODIS C6 LAI product (Martinez-Sanchez, 2020). The 371 dataset is delivered as one multi-band file per year in NetCDF (netCDF4 CF-1.6) format (14.7 GB per year) at 1 372 km resolution covering land area from 90 N to 60 S and representative of the 10-year period of 2010-2019. More 373 detail on method, data content and access can be found in Smets (2019) and Martinez-Sanchez (2020), and 374 Copernicus web-site https://land.copernicus.eu/global/products/lai

375 CGLS-LAI was used to compute the LAI fields for forest, irrigated crops and other land cover types.

#### 376 **3.3.4 Crop time evolution**

377 The RiceAtlas v3 (further referred as RiceAtlas) is a spatial database of global rice calendars and production. It 378 contains information on start, peak and end dates of sowing, transporting and harvesting rice, derived from global 379 and regional databases, national publications, online reports, and expert knowledge. It is delivered as 7 files in 380 shapefile format (in total 195.8 MB) for administrative units (in total 2725 spatial units) at 1 km resolution for the 381 national production totals to match the years 2010-2012 (Laborte et al., 2017a). RiceAtlas is ~10 times more 382 spatially detailed, and has ~7 times more special units comparing with other global datasets (Laborte et al., 2017b). 383 More detail on method, data content and access can be found in Laborte et al. (2017a) and Laborte et al. (2017b). 384 RiceAtlas was used to compute rice planting and rice harvesting days for three different seasons.

#### 385 3.4 Soil properties

386 The International Soil Reference and Information Centre (ISRIC) SoilGrids250m global gridded soil 387 information release 2017 (further referred as SoilGrids250m) is an output of special predictions produced by the 388 SoilGrids system, as a set of global soil property and class maps at 250 m resolution. It is derived from soil profile 389 data (from ~150,000 sites globally) with the use of machine learning, and contains information on soil 390 characteristics at six standard depths, including soil textures (clay, silt, sand), depth to bedrock, bulk density, 391 organic carbon, pH and cation exchange capacity. It is delivered as 43 files in GeoTiff format (in total 111.8 GB) 392 at 250 meters resolution covering land area with no permanent ice and representative for the year 2010 (according 393 to land cover) (Hengl et al., 2017). SoilGrids250m pH comparison with SSURGO data over California (depth 0-394 200 cm) and Soil and Landscape Grid of Australia data over Tasmania (depth 0-5 cm) show high correlation, 0.79 395 and 0.71 respectively (Hengl et al., 2017). Despite its limited accuracy (i.e. between 30 and 70 %, according to 396 the SoilGrids web-site) due to the scarcity of soil profile observations (especially in Central Asia, Artic regions 397 costal area and desert), low resolution of covariates data and algorithms, it was selected as the most recent source 398 of information. More detail on method, data content and access can be found in Hengl et al. (2017) and 399 SoilGrids250m web-site https://www.isric.org/explore/soilgrids/faq-soilgrids-2017.

400 SoilGrids250m was used to compute the soil depth and soil hydraulic properties for forest and non-forest.

### 401 3.5 Lakes

402 The Global Lakes and Wetlands Database (further referred as GLWD) is a global database of water bodies. It 403 is derived from a combination of global and regional lake data sets, registers and inventories (i.e. point information 404 with descriptive attributes), and digital maps (i.e. polygons, rasterised global land cover and land use maps). The 405 database consists of two global files in shapefile format at spatial resolutions of up to 1:1 million – GLWD-1 with 406 3067 largest lake and 654 largest reservoir polygons (6.4 MB), and GLWD-2 with ~250000 smaller lake and 407 reservoir polygons (32.0 MB); and of one global file in ADF raster format at 30 arc sec resolution – GLWD-3





408 combines GLWD-1, GLWD-2 and additional information (8.9 MB). Validation against documented data shows 409 that GLWD represents good wetland maximum extent, and describes comprehensively lakes with surface area 410 greater or equal 1 km<sup>2</sup> (Lehner and Döll, 2004). More detail on method, data content and access can be found in

411 Lehner and Döll (2004) and GLWD web-site https://www.worldwildlife.org/pages/global-lakes-and-wetlands-

412 database

413 GLWD (i.e. only GLWD-1 and GLWD-2) was used to compute the discrete lake mask field.

#### 414 3.6 Water demand

415 AQUASTAT is the FAO's global information system on water resources and agricultural water management. 416 AQUASTAT collects information on water use via the network of AQUASTAT National Correspondents who 417 are required to fill the annual questionnaire and collaborate with AQUASTAT team in the data validation process. 418 Five types of manual checks are followed by automatic implementation of almost 200 validation rules. The dataset 419 includes data for 180 countries worldwide, yearly data from 1979 to 2019 were used to produce the maps presented 420 by this manuscript. Float, lumped values for each country for the variables "Gross Domestic Product (GDP)", 421 "Industry, value added to GDP", "Agricultural water withdrawal", "Industrial water withdrawal", "Municipal 422 water withdrawal", "Total water withdrawal", and "Irrigation water withdrawal" were obtained in CSV format (2 423 files. in total 2.0 MB) from the AQUASTAT data acquisition dashboard 424 (https://tableau.apps.fao.org/views/ReviewDashboard-v1/country\_dashboard). More detail on method, data 425 content and access can be found in AQUASTAT web-site 426 https://www.fao.org/aquastat/en/overview/methodology/.

427 AQUASTAT variables were used accordingly to compute water demand fields for domestic, industrial, energy, 428 livestock use.

429

430 United States Geological Survey National Water Information System (further referred as USGS NWIS) is a 431 national database on water use data for the United States (US) with annual statistics provided every 5 years since 432 1950. The water use data are best estimates produced by the USGS in cooperation with local, state, and federal 433 agencies as well as academic and private organisations. The water use data are lumped values (float numbers) for 434 each state, delivered in plain text format (52 files, in total 56.0 MB). Following variables were used: "Domestic total self-supplied withdrawals, fresh, in Mgal/d", "Public Supply total self-supplied withdrawals, fresh, in 435 436 Mgal/d", "Industrial total self-supplied withdrawals, fresh, in Mgal/d", "Total Thermoelectric Power total self-437 supplied withdrawals, fresh, in Mgal/d", "Total Thermoelectric Power power generated, in gigawatt-hours", and 438 "Livestock total self-supplied withdrawals, fresh, in Mgal/d". More detail on method, data content and access can 439 be found in USGS NWIS web-site https://waterdata.usgs.gov/nv/nwis/wu. For this study, data from 1985 to 2015 440 were used.

441 USGS NWIS variables were used accordingly to refine the global water demand fields for the domestic, industrial, 442 energy, livestock use sectors for the US.

443

444 Global Change Analysis Model (further referred as GCAM) is an integrated, multi-sector model developed by 445 the Joint Global Change Research Institute (JGCRI) to explore the overall behaviour of human and physical 446 systems dynamics and interactions. GCAM includes five main systems. One of these systems, the water module, 447 provides information about water withdrawals for energy, agriculture, and municipal uses as lumped values of 448 235 hydrologic basins; a detailed explanation can be found in Calvin et al. (2019). Estimates of industrial, 449 thermoelectric water withdrawals (energy sector) and electricity consumption were computed by running the 450 GCAM model, the output used are two files in CSV format (in total 4.0 MB). Data from the following sectors was 451 used: "biomass", "electricity", "nuclearFuelGenIII", "nuclearFuelGenIII", "regional coal", "regional natural gas", 452 "regional oil", "SheepGoat", "Beef", "Dairy", "Pork", and "Poultry". More detail on method, data content and 453 access can be found in the documentation of the open source package https://github.com/JGCRI/gcam-454 core/tree/gcam-v6.0.

455

GCAM variables were used accordingly to estimate water withdrawals for industrial, energy, livestock use.

456 457 Global-scale gridded estimates of thermoelectric power and manufacturing water use (further referred as 458 Vassolo and Doll, 2005) is a global-scale gridded estimate of water withdrawal for cooling of thermal power 459 stations and for manufacturing. Estimates of values for the year 1995 are provided with a spatial resolution of  $0.5^{\circ}$ 460 by 0.5°. Thermoelectric power water use is based on the geographical location of 63590 thermal power stations. 461 Manufacturing water use is computed by estimating country-specific water withdrawal values, and spatial 462 downscaling using city night-time lights. Dataset verification of Vassolo and Doll (2005) showed satisfactory 463 representation of thermoelectric power water use but high uncertainty in the representation of manufacturing water 464 use. The data are delivered as one shapefile (2.5 MB). More details on method, data content and validation, and 465 data access can be found in Vassolo and Doll (2005).





466 Vassolo and Doll (2005) dataset was used for the computation of energy demand fields. 467

The Gridded Livestock of the World (GLW) version3 (further referred as GLW3) is a spatial gridded dataset
of the global distribution of eight livestock species for 2010. It is delivered as 8 GeoTiff files at 0.083333° (~10
km at the Equator) resolution (in total 208.0 MB). The species abundance was converted to total livestock mass.

471 More detail on method, data content and access can be found in Gilbert et al. (2018).

472 GLW3 was used to spatially disaggregate the water demand for livestock use.

473

World Bank manufacturing value added and gross domestic product (further referred as World Bank) data
provide "Manufacturing, value added (constant 2015 US\$)" values (further referred as MVA) and "Gross
Domestic Product GDP (constant 2015 US\$)" values. The data provided as a table, downloaded in CSV format
(6 files, in total 6.0 MB) from <a href="https://data.worldbank.org">https://data.worldbank.org</a>.

478 World Bank dataset was used to temporally downscale the values of water demand fields for the industrial and 479 energy sectors.

480

481 The Global Human Settlement Population Grid multitemporal version R2019A (further referred as GHS-482 POP) is a spatial raster dataset that depicts the distribution of population, expressed as the number of people per 483 grid-cell (Freire et al., 2016; Florczyk et al., 2019; Schiavina et al., 2019). GHS-POP residential population 484 estimates for target years provided by CIESIN GPWv4.10 were disaggregated from census or administrative units 485 to grid-cells, informed by the distribution and density of built-up as mapped in the Global Human Settlement 486 Layer. The dataset has a spatial resolution of 9 arc sec (~300 m at the Equator) resolution and is delivered as 487 individual files in GeoTiff format for 1975, 1990, 2000 and 2015 (4 files, in total 6.5 GB; available online: 488 https://ghsl.jrc.ec.europa.eu/ghs\_pop2019.php, last accessed: 21.02.2023).

GHS-POP was used to spatially disaggregate the country, state, basin-level information of domestic, industrial,
 energy water withdrawal.

491

492 Thematic Mapping Country Borders shapefile (further referred as TM 'country borders') was derived from 493 Thematic Mapping <sup>TM</sup>, which is a tool enabling web browsers to create thematic maps and associated world 494 datasets. For this work, the TM World Borders Dataset was downloaded as one shapefile (10.0 MB). The United 495 States Census Bureau Cartographic Boundary Files - Shapefile (further referred as US CB) provides the State 496 boundaries for the USA. For this work, the 2018 version was retrieved as one shapefile (3.2 MB; available online: 497 https://www.census.gov/geographies/mapping-files/time-series/geo/carto-boundary-file.html, last accessed: 498 21.02.2023). More detail on method, data found content and access can be in 499 http://thematicmapping.org/downloads/.

500 TM 'country borders' and US CB were used to spatially disaggregate the information of water withdrawal for 501 domestic, industrial, energy use.

502

Multi-Source Weather (further referred as MSWX) is a high-resolution (3-hourly, 0.1°), bias-corrected
 meteorological product with global coverage from 1979 to 7 months into the future. The data for 42 years
 (~316700 files in NetCDF format, in total 128.0 GB) were retrieved via www.gloh2o.org/mswx/. For more
 detailed information, see Beck et al. (2022).

MSWX 2-meter daily and monthly maximum and minimum air temperature were used to account for the climate induced intra- and inter- annual fluctuations of domestic, livestock, and energetic water demand.

509

Huang et al. (2018) is a publication presenting 0.5° resolution global monthly gridded sectoral water withdrawal
 dataset for the period 1971–2010.

512 Huang at al. (2018) Table 3 (calibrated R coefficient values) and Eq. (2) to (6) for temporal downscaling of 513 domestic and energy water demands were used in this study, respectively.

#### 514 **3.7 Surface field creation overview**

515 Considering the high resolution (i.e. hundreds of meters) and volume of data (i.e. GB) of most input datasets used 516 to generate the surface fields, a high performing data manipulation platform was needed. GEE (Gorelick et al., 517 2017) was selected as it provides (embedded) a vast high resolution data catalogue (e.g. ready available MERIT 518 DEM, CGLS-LC100, CLC2018) and powerful computation capabilities. It also allows to upload any raster and 519 vector data (e.g. GeoTiff or shapefiles) and to conduct each surface field tailored computations. All GEE scripts 520 were written in JavaScript to produce GeoTiff files, converted to the final file format (NetCDF) locally after 521 transfer from GEE platform.

522 To ensure a consistent representation of physical processes at all scales, surface fields should be as coherent as 523 possible among each other – between variables and across scales. Coherency can be achieved by using, where





524 possible, the same input datasets to derive different field types (e.g. unique forest information input to create all 525 forest-related surface fields), and making sure spatial aggregation/disaggregation across scales results in expected 526 values. Figure 1 shows a simplified scheme that relates input datasets (e.g. CGLS-LC100) with the resulting 527 surface fields (e.g. surface cover fractions – forest, inland water, and sealed surface fraction fields), also 528 highlighting fields requiring intermediary and sequential steps (e.g. forest fraction is needed to create soil 529 parameter fields over forested and non-forested areas).

530





534 For processes with horizontal dependency such as river routing, the relationship between grid-cells (e.g. how the 535 grid-cells are connected) must be defined first so that all dependent fields can be generated on the same grid 536 coordinates, spatial resolution and using consistent input data. For example, LDD defines how water moves across 537 the model grid-cells as a river drainage network (see Figure ) and strongly depends on elevation data (see Section 538 3.1.2 for more details). Because of the complex spatial dependency of a river drainage network, LDD must be 539 created directly from elevation data at the required grid and resolution and cannot be resampled from a previous 540 LDD field of a different grid and/ or resolution. It is then used to define information on the river network, including 541 upstream area and gradient. Note, Figure 1 misses an arrow from MERIT DEM to LDD only because this step 542 was mainly done by CaMa-Flood developers (see Section 3.1.2 for more details).

543 Four steps are involved in generating a particular surface field (see Table 2), with step 3 being the most complex 544 and varied (see Figure 2 for an example), and step 4 being necessary only for some model specifications (here as 545 required by LISFLOOD). Further details on specific manipulations associated with each field category are given 546 in Section 4 as relevant.





548 Figure 2. Examples of data manipulation for (a) transformation of elevation data into LDD (done within CaMa-Flood),





#### 550 4 Generation of surface fields

551 This section details the complex data manipulations required to generate the surface fields introduced in Section 2, 552 with examples of resulting fields. The techniques are reproduceable to different input data and/ or for different 553 output data specifications. Full technical descriptions for all fields needed by the LISFLOOD model are available 554 in the LISFLOOD user guide (available online: <u>https://ec-jrc.github.io/lisflood-code/4\_Static-Maps-introduction/</u>, 555 last accessed: 21.02.2023).

555 556

### 557 Table 2. The four steps of a particular surface field generation and associated data manipulations.

		-	
Order	Description	Purpose	Function
1	Raw file	Vector gridding, region merging	
1	preparation	Upscaling (spatial/ temporal aggregation)	Arithmetic mean, mode, sum, standard deviation (weighted) resampling from auxiliary data
2	Unit conversion         Converting values from native to fraction per grid-cell         Surface area, percentage or categorical to fractions per (see Annex 1 for more details)		Surface area, percentage or categorical to fractions per grid-cell (see Annex 1 for more details)
		Transforming	Mathematical equation/ function needed to generate the output variable
		Reprojecting	Interpolation (changing grid, preserving resolution in meters)
3	Value computation	Upscaling (spatial [default]/ temporal aggregation)	Arithmetic mean, mode, sum, standard deviation (weighted) resampling from auxiliary data (changing resolution, preserving grid)
		Downscaling (spatial [default]/ temporal disaggregation)	Nearest neighbour (changing resolution, preserving grid)
		Limiting	Force a minimum/ maximum value to satisfy e.g. calculation precision, physical meaning and/ or model requirement
4	Zero/	Replace zero/ NoData by the most appropriate values	LIGHT. Constant value, unweighted global mean, unweighted global mode
7	filling		DEEP. Values from next coarser resolution (up to an agreed maximum resolution); if still missing, method LIGHT

#### 558 4.1 Catchment morphology and river network

559 Environmental models require an accurate description of terrain and hydro-morphology to represent the 560 hydrodynamics at the spatial resolution of the model. Here all catchment morphology and river network fields are 561 derived from CaMa-Flood and MERIT DEM (see Table 3). They follow a complex sequential workflow (see 562 Figure ). Note that whilst some river network fields were already directly available from the CaMa-Flood 563 catalogue (e.g. LDD, channel length), they had to be adapted to the specific requirements of LISFLOOD, 564 specifically consistent with an interconnected river network described by the D8 algorithm (O'Callaghan and 565 Mark, 1984; Figure 2a) different to that used by the CaMa-Flood algorithm.



Figure 3. Workflow of complex manipulations to create some of the morphology and river network fields; solid arrows
 indicate a function transformation, dashed – modification of existing input data to LISFLOOD specifications.

#### 570 Table 3. Morphology and river network fields, their description, data source and applied transformation; \* denotes 571 transformation following Burek et al. (2014).

Field type	Description	Data source (variable)	Transformation
Local drainage direction ( <i>LDD</i> )	Connects every grid-cell forming a river network from springs to mouth	CaMa-Flood (flwd)	Direction coding, ensuring grid-cell connectivity
Grid-cell area ( <i>pixarea</i> )	Area of every grid-cells	CaMa-Flood (flwd)	Grid-cell area based on a given coordinate reference system and resolution





Grid-cell length	Length of every grid-cell	pixarea	$pixlength = \frac{pixarea}{resolution}$ , where
(pixlength)			resolution – 1.86 km and 5.57 km for 1
Unstream area	Accumulated area of all	LDD nixarea	PCRaster Acculux function
(upArea)	connected grid-cells of the	LDD, pixureu	(Karssenberg et al., 2010)
	LDD from springs (start;		
	lowest values) to mouth		
0. 1 11	(end; highest values)		
of elevation ( <i>elvstd</i> )	Amount of elevation variation within a grid-cell	MERII DEM	deviation deviation
Gradient (gradient)	Elevation gradient between	MERIT DEM; LDD	$aradient = \frac{abs(elv_{uc} - elv_{dc})}{abs(elv_{uc} - elv_{dc})}$ where $elv_{uc}$
	two connected grid-cells		$D_{uc,dc}$ , where $civ$
			downstream cell, $D_{ucdc}$ – distance
			between upstream and downstream
			cells
Channel bottom	Width of the bottom of the	CaMa-Flood (width);	Recomputing zero and negative values
width ( <i>chanbw</i> )	channel	upArea	based on equation* chambu = un Area + 0.0022
Channel length	Length of river channel in	CaMa-Flood (rivlen)	No transformation was carried out
(chanlength)	each grid-cell (can exceed	cultur Plood (IPPloi)	
	grid-size to account for		
	meandering river)		
Channel gradient	Gradient (slope) of river	MERIT DEM; LDD,	$changrad = \frac{abs(elv_{uc} - elv_{dc})}{chanlenath_{uc}}$ , where
(changraa)	channel mside a grid-cen	chantengin	elv – elevation, $uc$ and $dc$ – upstream
			and downstream cell;
Manainala	Maurinala analara	MEDIT DEM	Note: LDD is used to define <i>uc</i> and <i>dc</i>
roughness	coefficient of river channel	MERII DEM; upArea	$chanman = 0.25 \pm 0.015$
coefficient for	for each grid-cell		$\min\left(\frac{50}{1}\right) \pm 0.030$
channels (chanman)	0		$\lim_{upArea} \left( \frac{1}{upArea} + \frac{1}{upArea} \right) + 0.050$
			$\min(\frac{evm}{2000}, 1)$ , where $elv - elevation$ ,
Channel	Channel presence in the	'mask' (main model's	km and $m$ – values in km and m Channel mask is equal to 1 everywhere
mask (chan)	grid-cell indicator. Note	technical field)	channel mask is equal to 1 everywhere
	LISFLOOD specific	,	
	requirement to have		
	channels in every 'mask'		
Side slope (chans)	Slope of river banks (i.e.		Side slope of all channels is 45°, hence
2000 000 F C (00000)	horizontal distance divided		side slope is equal to 1 everywhere
	by vertical distance)		
Bankfull channel	Channel depth (i.e. river	upArea	Transformation based on equation*
depth (chanonkf)	beu depin)		$chanbnkf = 0.27 \cdot upArea_{km^2}^{,,,s}$ where $km^2 - values$ in $km^2$

#### 572 4.2 Land use fields

573 574 575

In models explicitly accounting for sub-grid variability, the fraction of each land use in every cell must be provided so that process representation for each land use can be weighted accordingly. Here, the fractions of the five land use classes used in LISFLOOD (and additional ocean fraction for consistency check) are derived from super-high 576 577 resolution datasets each following specific steps summarised in Table 4. Note that LISFLOOD requires all 'mask' (main model's technical field) grid-cells to have at least one non-zero fraction type, hence the extra step in the 578 579 generation of the inland water fraction field was to set empty grid-cells (i.e. grid-cells that based on the data source are fully covered with ocean) as fully covered with inland water.

580 581 582 Table 4. Fraction of land use fields, their description, data source and applied transformations; 'sum' refers to the sum of all fractions except 'other land cover fraction'; grey cells show required intermediate fields.

Field type	Description	Data source (variable)	Transformation (in order)
Forest fraction	Evergreen and deciduous	CGLS-LC100 (tree-	Unit conversion % to fraction;
(fracforest)	needle leaf and broad leaf	coverfraction)	Reprojecting and upscaling to final
	tree areas		grid and resolution with mean;
			Consistency check with other fractions





Sealed surface fraction ( <i>fracsealed</i> )	Urban areas, characterizing the human impact on the environment	CGLS-LC100 (urban- coverfraction)	Unit conversion % to fraction, scaled by 0.75; Reprojecting and upscaling to final grid and resolution with mean; Consistency check with other fractions
Inland water fraction (fracwater)	Rivers, freshwater and saline lakes, ponds and other permanent water bodies over the continents	CGLS-LC100 (water- permanent-coverfraction)	Force Fox Basin and Caspian Sea to be fully covered with water; Unit conversion % to fraction; Reprojecting and upscaling to final grid and resolution with mean; Consistency check with other fractions; Cross-checking with 'mask' and forcing empty grid-cells as inland water
Irrigated crops fraction (fracirrigated)	Irrigated areas of all possible crops excluding rice	SPAM (spam2010v1r0_global_physi cal-area_CROP_i, 41 crops rice excluding) CLC2018 (landcover = '212')	Shapefile gridding to its native resolution (~10 km); Unit conversion ha to fractions; Reprojecting and downscaling to CLC2018 grid and resolution (~100 m) with nearest neighbour Unit conversion class to fraction Merging SPAM- and CLC2018- derived fractions, priority to CLC2018; Reprojecting and upscaling to final grid and resolution with mean; Convictors up shock with other fractions
Irrigated rice fraction (fracrice)	Irrigated areas of rice	SPAM (spam2010v1r0_global_physi cal-area_RICE_i) CLC2018 (landcover = '213')	Shapefile gridding to its native resolution (~10 km); Unit conversion ha to fractions; Reprojecting and downscaling to CLC2018 grid and resolution (~100 m) with nearest neighbour Unit conversion class to fraction Merging SPAM- and CLC2018- derived fractions, priority to CLC2018; Reprojecting and upscaling to final grid and resolution with mean; Consistency check with other fractions
Other land cover fraction ( <i>fracother</i> )	Agricultural areas, non- forested natural area, pervious surface of urban areas	Non-negative residual from 1 subtracting 'sum' of all other fractions	$fracother = \max((1 - sum), 0)$
Ocean fraction ( <i>fracocean</i> )	Oceans	CGLS-LC100 (discrete_classification = '200')	Unit conversion class to fraction; Forcing NoData to zero over 'mask' grid-cells, otherwise – fully covered; Reprojecting and upscaling to final grid and resolution with mean; Consistency check with other fractions

583

584 For the sealed surface fraction, it is assumed that water can infiltrate in roughly 25 % of urban areas at kilometre 585 scale through e.g. trees along the road, bushes along the fence, grass or moss between concrete tiles or cobble 586 stones.

587 To ensure consistency between fractions, the sum of all fraction fields must be 1 at any resolution. When sum is 588 greater than 1, the inland water fraction value is assumed correct (input data corrected prior computation over Fox 589 Basin and Caspian Sea) and all other fractions are corrected ( $fr\_corr$ ) following Eq. (1):

$$590 \quad fr\_corr = fr\left(1 - \frac{fr_{inlandWater+fr_{ocean+fr_{forest+fr_{sealed+fr_{irigated+fr_{rice}-1}}}{fr_{forest+fr_{sealed+fr_{irigated+fr_{rice}-1}}}\right),$$

(1)

591 where fr refers to the original (i.e. before consistency check) fraction of the forest, irrigated crops, rice and sealed 592 surfaces.

593 The generated fraction fields, e.g. forest (see Figure 4a) and other land cover (see Figure 4b), have generally good

594 consistency with other up-to-date products like ESA CCI Land Cover time-series v2.0.7 (ESA CCI map viewer 595 https://maps.elie.ucl.ac.be/CCI/viewer/; Defourny et al., 2017).

596







597 598 Figure 4. Fraction fields for forest (a) and other land cover (b) at 3 arc min (~5.6 km at the equator) resolution.

### 599 4.3 Vegetation properties

600 In complement to the land use fraction, the distribution of vegetation type and characteristics is required to capture 601 the difference in environmental processes such as water intake of evaporation to be represented accurately (see 602 Section 2). Here the vegetation properties are derived from many data sources using maps to account for the 603 species spatial distribution and tables to obtain associated hydro-dynamics properties. This requires assumptions 604 to be made in case different sources did not contain the same information, and transformations to be applied 605 depending on the vegetation type. The main data sources and general transformation steps to derive the 606 18 vegetation properties fields are summarised in Table and following text. Note that 'crop group number' 607 variable corresponds to a water depletion value and can be averaged across different crop types.

608

Field type	Description	Data source	Transformation (in order)
Crop coefficient	Ratio between	CGLS-LC100	Force Fox Basin and Caspian Sea to be
for forest,	the potential	(discrete_classification = '111',	fully covered with water;
irrigated crops	(reference)	'112', '113', '114', '115', '116',	Unit conversion class to fraction (in total
and other land	evapotranspirati	'121', '122', '123', '124', '125',	12 forest related and 7 other land cover
cover type	on rate, in	'126' [forest types], '20', '30',	related fraction fields);
(cropcoef_f,	mm/day, and	'40', '60', '70', '90', '100' [other	Reprojecting and upscaling to final grid
cropcoef_i,	the potential	land cover types])	and resolution with mean
cropcoef_o)	evaporation rate	SPAM	Shapefile gridding to its native resolution
	of a specific	(spam2010v1r0_global_physical-	(~10 km);
	crop (averaged	area_CROP_i/r, 42 crops, 'i' -	Unit conversion ha to fractions (in total 42
	by time and	irrigated, 'r' - rainfed)	irrigated crop related and 42 rainfed crop
	ecosystem type)		related fraction fields);
			Reprojecting and downscaling to final grid
			and resolution with nearest neighbour;
			Limiting values to 0.0-1.0 interval
		FAO56 (Table 11, 12 -	Average crop coefficient value across
		information on crop coefficient	climate zones for each crop growing stage
		and crop height); Intara et al.	and crop/ land cover type;
		(2018), Burek et al. (2014)	Weighted average of crop coefficient per
			different crop growth stages (weighted by
			stage duration in days if available,
			otherwise mean);
			Average crop height value across climate
			zones for each crop/ land cover type
			Weighted average of relevant crop
			coefficient for forest, irrigated crops and
			other land cover type (weighted by crop
			height and fraction) following Eq. (2);
			Note: for other land cover type
			computation of crop coefficient of all
			rainfed crops is used for CGLS-LC100
			(discrete_classification = '40');
			Zero/ NoData filling with global mean
Crop group	Represents a	CGLS-LC100	Same steps as for crop coefficient
number for forest,	vegetation type	(discrete_classification = '111',	

609	Fable 5. Vegetation property fields, their description, data source and applied transformations; grey cells	s show
610	required intermediate fields.	





irrigated crops and other land cover type (cropgrpn_f, cropgrpn_i,	and is an indicator of its adaptation to dry climate (averaged by	'112', '113', '114', '115', '116', '121', '122', '123', '124', '125', '126' [forest types], '20', '30', '40', '60', '70', '90', '100' [other land cover types])	
cropgrpn_o)	ecosystem type)	SPAM (spam2010v1r0_global_physical- area_CROP_i/r, 42 crops, 'i' – irrigated, 'r' – rainfed)	Same steps as for crop coefficient
		FAO56 (Table 22 – information on crop depletion fraction), SUPIT (Table 6.1, 6.2 – information on crop groups), Burek et al. (2014)	Applying function (SUPIT) to water depletion fraction (FAO56) for each crop/ land cover type $cropgrpn = 10 \cdot fr_{dep} - 1.5$ , where $fr_{dep}$ – water depletion fraction; Limiting values to 1.0-5.0 interval; Note: if $fr_{dep}$ missing – using precomputed crop group number (Burek et al., 2014)
			Same steps as for crop coefficient, but in Eq. (2) weighted by fraction only
Manning's surface roughness coefficient for forest, irrigated crops and other land cover type (mannings_f,	Roughness or friction applied to the flow by the surface on which water is flowing (averaged by	CGLS-LC100 (discrete_classification = '111', '112', '113', '114', '115', '116', '121', '122', '123', '124', '125', '126' [forest types], '20', '30', '40', '60', '70', '90', '100' [other land cover types])	Same steps as for crop coefficient
mannings_0)	ecosystem type)	SPAM (spam2010v1r0_global_physical- area_CROP_i/r, 42 crops, 'i' – irrigated, 'r' – rainfed)	Same steps as for crop coefficient
		CHOW (Table 5, 6 – information on roughness coefficient n, Burek et al. (2014)	Matching roughness coefficient for each crop/ land cover type
			Same steps as for crop coefficient, but in Eq. (2) weighted by fraction only
Leaf area index for forest, irrigated crops and other land cover type ( <i>laif</i> , <i>laii</i> , <i>laio</i> )	Defined as half the total area of green elements of the canopy per unit horizontal ground area m <sup>2</sup> /m <sup>2</sup> (10-day average; 36 fields in total)	CGLS-LAI 10-day average for 2010-2019; <i>fracforest</i> , <i>fracirrigated</i> , <i>fracother</i>	Upscaling to final temporal resolution (in total 36 LAI fields); Reprojecting and upscaling to final grid and spatial resolution with unweighted mean; Filtering sparce areas of relevant fractions fr < 0.7, where $fr$ – fraction; NoData filling DEEP (upscaling to 1, 3, 15 arc min, 1, 3, 15, 60 degrees spatial resolution with unweighted mean; replacing NoData at final resolution with first available precomputed less coarser resolution, if not – with zero)
Rice planting day (riceplantingday1, riceplantingday2, riceplantingday3)	Most probable day of the year when rice is planted for the first, second and third time	RiceAtlas (PLANT_PKn, 3 seasons)	Ordering planting seasons by increasing Julian day (in total 3 planting dates per spatial unit); Shapefile gridding to final grid and spatial resolution (in total 3 fields); Note: if less than 3 seasons – repeating last
Rice harvest day (riceharvestday1, riceharvestday2, riceharvestday3)	Most probable day of the year when rice is harvested after planting for the first, second and third time	RiceAtlas (HARV_PKn, 3 seasons)	available planting/ harvesting seasons date; NoData filling with global unweighted mode date of first planting/ harvesting season (i.e. 105 – 15 <sup>th</sup> April/ 227 – 15 <sup>th</sup> August)
Root depth for forest and non- forest (root_depth_f, root_depth_o)	Deepest soil depth reached by the crop roots	CGLS-LC100 (discrete_classification = '111', '112', '113', '114', '115', '116', '121', '122', '123', '124', '125', '126' [forest types], '20', '30',	Same steps as for crop coefficient





	'40', '60', '70', '90', '100' [other land cover types])	
	SPAM (spam2010v1r0_global_physical- area_CROP_i/r, 42 crops, 'i' – irrigated, 'r' – rainfed)	Same steps as for crop coefficient
	FAO56 (Table 22 – information on crop rooting depth), Burek et al. (2014)	Matching rooting depth for each crop/ land cover type
		Same steps as for crop coefficient, but in Eq. (2) weighted by fraction only; Downscaling to native SoilGrids250m resolution with nearest neighbour (for soil depth calculations)

611

612 The final step of the crop coefficient, crop group number, Manning's surface roughness coefficient, and additional 613 crop height (for crop coefficient calculation) and root depth (for soil depth calculation, see Section 4.4) for forest, 614 irrigated crops and other land cover type is to compute weighted average of their components (e.g. different forest 615 types) following Eq. (2):

 $K = \frac{A_1 \cdot fr_1 \cdot K_1 + A_2 \cdot fr_2 \cdot K_2 + \dots + A_N \cdot fr_1}{A_1 \cdot fr_1 + A_2 \cdot fr_2 + \dots + A_N \cdot fr_N}$  $+A_N \cdot fr_N \cdot K_N$ 616

(2)

where A is a scaling parameter (equals 1, except for crop coefficient where it equals to crop height), fr refers to 617 618 the fraction of crop or land cover type, K – default (i.e. source table based) variable in question values, 1..N – 619 number of crop or land cover types included in the field (i.e. for forest N=12, irrigated crops N=41, other land 620 cover type N=7 and for CGLS-LC100 type '40' (cropland) default values are based on 42 rainfed crops). 621



Crop coefficient other



622 623 Figure 5. Crop coefficient for forest (a) and other land cover type (b) at 3 arc min (~5.6 km at the equator) resolution.

#### 624 4.4 Soil properties

625 Soil proprieties are derived from SoilGrids250m (see Section 3.4) and are computed for both forested and non-626 forested (also known in literature as 'others') areas, expressed as fractions (see Section 4.2), where non-forested 627 area is the complementary fraction of forest. Soil depth layers are derived first and used as input to the soil hydraulic equations used to derive the properties, following a sequential workflow (see Error! Reference source 628 629 not found.). Equations used are from Toth et al. (2015). 630



Figure 6. Workflow to generate the soil related fields; solid arrows indicate a function transformation, dotted -633 upscaling; 'SoilGrids250m depths' - fields at the SoilGrids250m native grid and resolution with six default depths, 634 'final grid and resolution' - fields at the dataset's final grid and resolution, boxes with no explicit indication - fields at 635 SoilGrids250m native grid and resolution only.

636 Table 6. Soil property fields, their description, and applied transformations.





Field type	Description	Data Source	Transformation (in order)
Soil depth layers 1, 2, 3 for	Root depths	SoilGrids250m	Transforming at SoilGrids250m
forest and non-forest	assumed to divide	(absolute depth to bedrock);	native grid and resolution as
(soildepth1 f,	the total soil depth	root depth f, root depth o	described in Annex 2 'Soil
soildepth1 o, soildepth2 f,	between topsoil		Depth' (in total 3 forest and 3
soildepth2 o, soildepth3 f,	(surface [layer 1]		non-forest soil depth layer
soildepth3 o)	and middle [layer		fields);
1 = 7	21) and subsoil		Reprojecting and upscaling to
	(bottom [layer 3])		final grid and resolution with
			unweighted mean;
			NoData filling DEEP (upscaling
			to 1, 3, 15 arc min, 1, 3, 15, 60
			degrees spatial resolution with
			unweighted mean; replacing
			NoData at final resolution with
			first available precomputed less
			coarser resolution, if not - with
			zero)
Saturated volumetric soil	Saturated water	SoilGrids250m (clay_content,	Transforming at SoilGrids250m
moisture content for soil	content soil	silt content. bulk density):	native grid and resolution as
depth layers 1, 2, 3, and for	hydraulic property	soildepth1 f. soildepth1 o.	described in Annex 2 'Soil
forest and non-forest	representing the	soildepth2 f, soildepth2 o,	hydraulic parameters' (in total 5
(thetas1 f, thetas1 o,	maximum water	soildepth3 f, soildepth3 o;	fields per soil hydraulic
thetas2 f, thetas2 o,	content in the soil	fracforest	parameter, except thetar - only 3
thetas3)		5 5	as no forest/ non-forest
Residual volumetric soil	Residual water	SoilGrids250m (clay content,	separation);
moisture content for soil	content soil	silt_content); soildepth1_f,	Limiting values and weighting
depth layers 1, 2, 3	hydraulic property	soildepth1_o, soildepth2_f,	by forest/ non-forest fraction
(thetar1, thetar2, thetar3)	representing the	soildepth2_o, soildepth3_f,	(limits thetas $< 1.0$ , thetar $<$
	minimum water	soildepth3_o; fracforest	thetas, lambda $\leq 0.42$ ,
	content in the soil		$genua \le 0.055, ksat > 0.0);$
Pore size index for soil	Van Genuchten	SoilGrids250m (clay_content,	Upscaling to final grid and
depth layers 1, 2, 3, and for	parameter $\lambda$ (also	silt_content, bulk_density,	resolution with unweighted
forest and non-forest	referred as 'n-1' in	organic_carbon_content);	mean;
(lambda1_f, lambda1_o,	literature) soil	soildepth1_f, soildepth1_o,	NoData filling DEEP (upscaling
lambda2_f, lambda2_o,	hydraulic property	soildepth2_f, soildepth2_o,	to 1, 3, 15 arc min spatial
lambda3)	representing the	soildepth3_f, soildepth3_o;	resolution with unweighted
	pore size index of	fracforest	mean; replacing NoData at final
	the soil		resolution with first available
Van Genuchten equation	Van Genuchten	SoilGrids250m (clay_content,	precomputed less coarser
parameter for soil depth	parameter α soil	silt_content, bulk_density,	resolution, if not – with global
layers 1, 2, 3, and for forest	hydraulic property	organic_carbon_content);	unweighted mean)
and non-forest (genual_f,		soildepth1_f, soildepth1_o,	
genua1_o, genua2_f,		soildepth2_f, soildepth2_o,	
genua2_o, genua3)		soildepth3_f, soildepth3_o;	
		fracforest	
Saturated soil conductivity	Saturated hydraulic	SoilGrids250m (clay_content,	
for soil depth layers 1, 2, 3,	conductivity soil	silt_content, soil_pH,	
and for forest and non-	hydraulic property	cation_exchange_capacity);	
forest (ksat1_f, ksat1_o,	representing the	soildepth1_f, soildepth1_o,	
ksat2_f, ksat2_0, ksat3)	ease with which	soildepth2_f, soildepth2_o,	
	water moves	soildepth3_f, soildepth3_o;	
	through pore spaces	fracforest	
	of the soil		1

637

Two of the most common soil parameters of land surface and hydrological models, saturated hydraulic conductivity *ksat* and saturated water content, are shown in Figure 7.

640 Saturated hydraulic conductivity ksat (see Figure 7a) ranges from 2 to 7445 mm/day. The highest ksat values are 641 concentrated in desertic areas such as the Sahara, Arabian Peninsula, Gobi, Patagonian, Sonoran-Mojave and 642 Kalahari and Namib deserts. Low ksat between, 2 and 18 mm/day, are found in the Amazon river basin, the lower 643 Mississippi river basin and South East Asia. ksat was visually compared against 8 global datasets developed with 644 different input data and/ or PTFs (Zhang and Schaap, 2019; Gupta et al., 2021); a general agreement is noticeable 645 in areas that show low variability across all datasets. Northern Russia, Canada, South East Asia and Sonoran-646 Mojave Desert are the areas with high variability among datasets, with values ranging from very low to very high 647 ksat. Source of uncertainties in ksat values are primarily due to little availability of soil samples and measurements





648 carried out in those areas. Moreover, the climatic context plays a relevant role in clay mineralogy composition,
649 organic composition and soil pores structure (Hodnett and Tomasella, 2002), which influence how water flows
650 through the soil. Therefore, the PTF developed using soil samples collected in temperate areas (such as Europe)
651 are expected to have a different hydraulic behaviour compared to those collected in tropical climates (Gupta et
652 al., 2021), as also seen in Figure 7a.

- Saturated water content (see Figure 7b) ranges between 0.27 to 0.79, with 80% of values between 0.40 and 0.46.
   A comparison with other global datasets was not carried out, however uncertainties are expected to be of the same order of magnitude than those of *ksat* given the fact the saturated water content is calculated using bulk density and clay content data.
- 657



Figure 7. Saturated hydraulic conductivity (a) and saturated water content (b) for forested areas of soil depth layer 2.

#### 660 4.5 Lakes

661 The lake field is derived from the GLWD database.

662

#### 663 Table 7. Lake field, its description, data source and transformation.

Field type	Description	Data source	Transformation (in order)
Lake mask	Area covered by	GLWD (GLWD-1, GLWD-2,	Filtering non-lake spatial units;
(lakemask)	lakes only (binary	lake type only); fracwater	Shapefile gridding to final grid and resolution;
	representation)		If <i>fracwater</i> > 0 and GLWD is 'lake', then
			lakemask is 1, otherwise 0

#### 664 4.6 Water demand

665 Global gridded water demand fields with monthly variability were generated for the four sectors using the data 666 sources listed in Section 3.6 and following the transformations summarised in Table 8 (for additional information 667 and extra details see GitHub repository 'lisflood-utilities/water-demand-historic at feature/add\_h\_branches\_upd · 668 ec-jrc/lisflood-utilities · GitHub', last accessed: 21.02.2023). The water demand values are provided in mm/day, 669 one field per month from 01.01.1979 to 31.12.2019 (the first day of each month is used as the representative 670 timestamp for the entire month). The methodology applied largely follows Huang et al. (2018), with the key 671 differences being the use of freely available datasets and the higher resolution of the resulting fields. Spatial 672 downscaling was achieved following the approach of Hejazi et al. (2014); temporal downscaling was performed 673 following the approaches of Wada et al. (2011), Voisin et al. (2013) and Huang et al. (2018). It should be noted 674 that country-scale estimates (from AQUASTAT) were integrated with state-level water withdrawal estimates 675 (from USGS NWIS). The protocol for the integration of local information with global data sources was developed 676 for further use in the future, to enable the integration of other regional or national datasets as soon as they become 677 available.

678

# 679Table 8. Water demand fields, their description, data source and applied transformations; grey cells show required680intermediate fields.

Field type	Description	Data source	Transformations (in order)
Population density (pop)	Number of people per	GHS-POP R2019A (1975,	Reprojecting and upscaling from native (9 arc sec) to the final grid and intermediate resolution of 0.01°x0.01° with sum (in
	grid-cell	1990, 2000, 2015)	total four fields);





			Transforming from population number to density per grid-cell (i.e. dividing by grid-cell area) and upscaling from intermediate to final resolution with mean (in total four fields); NoData filling (year) with linear interpolation till 2015, and with years 2000 and 2015 trend extrapolation 2016 onwards ( <i>pop</i> <sup>grid</sup> <sub>year</sub> ; in total 41 fields)
		TM 'country borders', US CB 'state borders'	Shapefile (country, US State) gridding to final grid and intermediate resolution of 0.01°x0.01°, then to final resolution; Transforming from population density per grid-cell to population per country (i.e. multiplying by grid-cell area and summing grid- cells according to the country mask from step above; $pop_{vear}^{country}$ ; in total one table)
Water demand for domestic use ( <i>dom</i> )	Daily supply of water volume for indoor and outdoor household	AQUASTAT (per country), USGS NWIS (per US State), pop	Unit conversion from native to km <sup>3</sup> /year; NoData filling (year): for countries – with linear interpolation and forward/ backward extrapolation based on <i>pop<sub>vear</sub><sup>country</sup></i> , for US states – with linear interpolation and nearest neighbour extrapolation ( <i>demand<sub>vear</sub><sup>country</sup></i> , in total one table)
	purposes and for all the uses that are connected to the municipal system (e.g., water used by	<i>pop</i> , 1M 'country borders', US CB 'state borders'	Transforming water demand ( $demand_{year}^{our}$ ) to water demand per capita per country/ US State per year (in total one table): $perCapitaDemand_{year}^{country} = \frac{demand_{year}^{country}}{pop_{year}^{porty}}$ ; NoData filling (country) with nearest neighbour; Transforming from water demand per capita to water demand per grid-cell (i.e. weighting by $pop_{year}^{prid}$ ; in total one field per
	schools, and public buildings)	MSWX, Huang et al. (2018) [Table 3, Eq. (2)].	year): demand <sup>grid</sup> per CapitaDemand <sup>grid</sup> per pyear Temporal downscaling (month) to account for the withdrawal fluctuations between the warmest and coldest months based on Huang et al. (2018) Eq. (2) (in total 12 fields per year): demand <sup>grid</sup> <sub>month,year</sub> = $\frac{demandgrid}{month,year} \cdot \left(\frac{Terdit}{month,year} - \frac{aregrid}{rg} + \frac{Terdit}{rg} + R + 1\right)$ , where $avg\bar{r}_{year}^{grid}$ , $max\bar{r}_{year}^{grid}$ , $min\bar{r}_{year}^{grid}$ are the average, maximum, minimum monthly temperatures in a year; $\bar{T}_{month,year}^{grid}$ is the average temperature in a month of the year; $R$ is the amplitude of the monthly fluctuations from Huang et al. (2018) [Table 3]; $month_{year}^{number}$ is number of months in a year, i.e. 12; Temporal downscaling (day; in total 12 fields per year): $demand^{grid}_{day,month,year} = \frac{demand^{grid}_{month,year}}{day_{month}}$ , where $day_{month}^{number}$ is number of days in a month of a certain year
Water demand for industrial use ( <i>ind</i> )	Daily supply of water volume for fabricating, processing, washing and sanitation, cooling or transporting a product, incorporating water into a product	AQUASTAT (per country), USGS NWIS (per US State), GCAM (per region), Vassolo and Doll (2005), World Bank (MVA), <i>pop</i> , TM 'country borders'	<ul> <li>Unit conversion from native to km<sup>3</sup>/year; NoData filling (year; in total one table):</li> <li>regional data – downscaling (spatial) to country values (i.e. weighting by <i>pop<sub>year</sub><sup>v</sup></i>), then linear interpolation (between years) and nearest neighbour extrapolation in time, finally rescaling values according to Vassolo and Doll (2005);</li> <li>country data – with linear interpolation (between years) and forward/ backward extrapolation based on <i>MVA</i> or <i>pop<sub>year</sub><sup>v</sup></i>, value disaggregation from industrial water demand to manufacturing and thermoelectric water demands according to regional data results;</li> <li>for US States data – with linear interpolation;</li> <li>mosaicking results from US States and country data, from regional data, if not – with zero</li> </ul>
		<i>pop</i> , TM 'country borders', US CB 'state borders'	Transforming from water demand per country/ US State to per grid-cell (i.e. weighting by $pop_{year}^{grid}/pop_{year}^{country}$ ; in total one field per year): $demand_{year}^{grid} = \frac{demand_{year}^{goutry}}{pop_{year}^{country}} \cdot pop_{year}^{grid}$ ;





Water demand for thermoelectric use ( <i>ene</i> )	Daily supply of water volume for the cooling of thermoelectric and nuclear power plants	AQUASTAT (per country), USGS NWIS (per US State), GCAM (per region), Vassolo and Doll (2005), World Bank (MVA), pop, TM 'country borders'	Temporal downscaling (day; in total one field per year): $demand_{day,year}^{grid} = \frac{demand_{grid}}{day_{year}^{humber}}, \text{ where } day_{year}^{number} \text{ is number of}$ days in a year Same steps as for water demand for industrial use, but using the energy withdrawals as input data (in total one table) Same steps as for under domand for industrial use (in total one table)
		borders', US CB 'state borders'	Same steps as for water demand for industrial use (in total one field per year)
		GCAM (per region), MSWX, Huang et al. (2018) [Eq. (3)- (10)].	Temporal downscaling (month) to account for the withdrawal fluctuations between the warmest and coldest months based on Huang et al. (2018) Eq. (3)-(10) (in total 12 fields per year)
Water demand for livestock use ( <i>liv</i> )	Daily supply of water volume for domestic animal needs	AQUASTAT (per country), USGS NWIS (per US State), GCAM (per region), GLW3, TM 'country borders'	<ul> <li>Unit conversion from native to km<sup>3</sup>/year;</li> <li>NoData filling (year; in total one table):</li> <li>regional data – spatial downscaling from regional withdrawals to country values (i.e. weighting by total livestock mass estimates per country from GLW3, <i>livestock<sup>country</sup></i>):</li> <li><i>demand<sup>country</sup></i> = <sup>withdrawal<sup>region</sup></sup>/<sub>livestock<sup>region</sup></sub>. <i>livestock<sup>country</sup></i>, then value linear interpolation (between years) and nearest neighbour extrapolation, finally rescaled with country data (if available)</li> <li>for US States data – with linear interpolation (between years) and nearest neighbour extrapolation;</li> <li>mosaicking results from US States and regional data, if not – with zero</li> </ul>
		GLW3, TM 'country borders', US CB 'state borders'	Transforming from water demand per country/ US State to per grid-cell (i.e. weighting by $\frac{llvestockDensity_{year}^{yrid}}{llvestockDensity_{year}^{yrid}}$ ; in total one field per year): $demand_{year}^{grid} = \frac{demand_{year}^{country}}{llvestockDensity_{year}^{ord}}$ . $livestockDensity_{year}^{grid}$ ; Temporal downscaling (day; in total one field per year): $demand_{day,year}^{grid} = \frac{demand_{year}^{country}}{day_{year}^{number}}$ , where $day_{year}^{number}$ is number of days in a year

681 To the best of the authors' knowledge, no other publicly accessible temporally varying global water demand field 682 set exists (only static datasets). A rigorous validation of the temporally varying water demand fields is not 683 straightforward at the global scale, as the only comprehensive global data source, FAO AQUASTAT, was used 684 to create the fields.

#### 685 5 Data, access, licensing, documentation

686 The new CEMS\_SurfaceFields\_2022 is an open-source dataset of the Copernicus Emergency Management 687 Service describing key components of the Earth surface generally required in environmental and hydrological 688 modelling, including Earth system modelling and numerical weather prediction. The dataset includes static fields 689 (e.g. forest fraction), yearly cycle fields (e.g. 10-day average LAI, in total 36 fields), and yearly varying fields 690 (e.g. water demand). The surface fields are based on 25 different sources, including global and regional high 691 resolution (up to 100 m) gridded and vector datasets. They were processed into two set of fields (i) at 1 arc min 692 resolution (~1.86 km at the Equator) over Europe (72.25 N/ 22.75 N, 25.25 W/ 50.25 E; 4530x2970 grid-cells), 693 and (ii) at 3 arc min resolution (~5.57 km at the Equator) over the Globe (90.00 N/ 90.00 S, 180.00 W/ 180.00 E; 694 7200x3600 grid-cells), to provide an up-to-date surface state for six main field groups: (1) catchment morphology 695 and river network, (2) land use fields, (3) vegetation properties, (4) soil properties, (5) lakes, (6) water demand.





The CEMS\_SurfaceFields\_2022 dataset consist in total of 140 gridded fields at EPSG:4326 - WGS84: World 696 697 Geodetic System projection in NetCDF format with information on Earth's surface state (see Table 9 for the full 698 list of fields), which are grouped thematically in sub-folders. The 1 arc min European fields have a total volume 699 of 9.3 GB and the 3 arc min global fields have a total volume of 22.7 GB. The CEMS\_SurfaceFields\_2022 dataset 700 is freely available for download from the JRC Data Catalogue (https://data.jrc.ec.europa.eu/). The set of global 701 surface fields at 3 arc min resolution can be found here (JRC Data Catalogue - LISFLOOD static and parameter 702 maps for GloFAS - European Commission (europa.eu), https://data.jrc.ec.europa.eu/dataset/68050d73-9c06-703 499c-a441-dc5053cb0c86) and the set of surface fields for the European domain at 1 arc min resolution can be 704 found here (JRC Data Catalogue - LISFLOOD static and parameter maps for Europe - European Commission 705 (europa.eu), https://data.jrc.ec.europa.eu/dataset/f572c443-7466-4adf-87aa-c0847a169f23). The README.txt 706 file that can be found there contains the basic description of each surface fields including general information, 707 data description, file overview, methodological information and data access and sharing information (for detailed 708 technical description of how the surface fields were generated refer to the LISFLOOD User Guide, available 709 online: https://ec-jrc.github.io/lisflood-code/4\_Static-Maps-introduction/). The changelog.txt file - provides users 710 with information on updates to the datasets. The copyright.txt file - information about the data license (CC BY 711 4.0).

712

Tuble STT an abe of Surface fields while blort description and anals medaded in Childs_Surface fields_
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Field group	Description	Name	Units
Main	model's technical field	mask	dimensionless
Catchment	local drainage direction (i.e. flow direction from	LDD	dimensionless
morphology	one cell to another)		
and river	grid-cell area	pixarea	m <sup>2</sup>
network	grid-cell length	pixlength	m
	upstream area	upArea	m <sup>2</sup>
	standard deviation of elevation	elvstd	m
	gradient	gradient	m/m
	channel bottom width	chanbw	m
	channel length	chanlenght	m
	channel gradient	changrad	m/m
	Manning's roughness coefficient for channels	chanman	s/m <sup>1/3</sup>
	channel mask (i.e. presence of river channel)	chan	dimensionless
	channel side slope (i.e. channel's horizontal	chans	m/m
	distance divided by vertical distance)		
	bankfull channel depth	chanbnkf	m
	channel floodplain (i.e. width of the area where	chanflpn	m
	the surplus of water is distributed when the water		
	level in the channel exceed the channel depth)		
Land use	fraction of forest	fracforest	dimensionless
fields	fraction of sealed surface	fracsealed	dimensionless
	fraction of inland water	fracwater	dimensionless
	fraction of irrigated crops	fracirrigated	dimensionless
	fraction of rice	fracrice	dimensionless
	fraction of other cover types	fracother	dimensionless
Vegetation	crop coefficient	<pre>cropcoef_f, cropcoef_i, cropcoef_o</pre>	dimensionless
properties	crop group number	cropgrpn_f, cropgrpn_i,	dimensionless
(for forest		cropgrpn_o	
[f], irrigated	Manning's surface roughness coefficient	mannings_f, mannings_o,	s/m <sup>1/3</sup>
crops [i],	rice planting days (3 seasons)	riceplantingday1, riceplantingday2,	calendar day
other land		riceplantingday3	number
cover types	rice harvesting days (3 seasons)	riceharvestday1, riceharvestday2,	calendar day
[o])		riceharvestday3	number
	leaf area index	laif, laii, laio	$m^2/m^2$
Soil	surface layer depth	soildepth1_f, soildepth1_o	mm
properties	middle layer depth	soildepth2_f, soildepth2_o,	mm
(for [1, 2,	subsoil depth	soildepth3_f, soildepth3_o	mm
<li>3] layers;</li>	saturated volumetric soil moisture content	thetas1_f, thetas1_o, thetas2_f,	m <sup>3</sup> /m <sup>3</sup>
for forest		thetas2_o, thetas3	
[f], non-	residual volumetric soil moisture content	thetar1, thetar2, thetar3	$m^3/m^3$
forest [o])	pore size index	lambda1_f, lambda1_o, lambda2_f,	dimensionless
		lambda2 o, lambda3	





	Van Genuchten equation parameter	genual_f, genual_o, genua2_f,	cm <sup>-1</sup>
	saturated soil conductivity	ksat1_f, ksat1_o, ksat2_f, ksat2_o, ksat3	mm/day
Lakes	lake mask (i.e. presence of lakes)	lakemask	dimensionless
Water	livestock	liv	mm/day
demand	industry	ind	mm/day
	thermoelectric production	ene	mm/day
	domestic use	dom	mm/day

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715 Whilst the CEMS\_SurfaceFields\_2022 dataset followed strict requirements of the LISFLOOD-OS model (e.g. 716 format, treatment of missing values, number of soil layers, etc...) it definitely can be used outside the LISFLOOD 717 context, using the full dataset or its parts, for applications such as modelling risk assessment. The workflow and 718 methodology used to generate the dataset and published in this manuscript can be used as reference and be easily 719 modified if further adaptation to the dataset is needed (e.g. using different set of equations to describe the soil 720 properties, or sourcing new/ more relevant local datasets).

### 721 6 Conclusion

The Earth's surface has a strong impact on the surface energy and water balance that drives lower atmosphere weather conditions and river discharge fluctuations. Depending on the surface type (e.g. land use, terrain or soil), weather in the region can be colder/ warmer, more/ less humid, drier/ rainier, and/ or calmer/ windier than its surroundings, and the terrestrial water cycle can differ, with water infiltrating more/ less in the soil, leaving as evaporation in a larger/ smaller rate, and reaching rivers faster/ slower. Surface information is provided by land use and ecosystem type (e.g., forest, rice paddy, bare ground, urban), river geometry (e.g., channel width, channel length), soil properties (e.g., depth, porosity, hydraulic properties), amongst others.

729 Information of underlying surface fields can be accounted for in Earth system and environmental models (e.g. 730 atmospheric, hydrological, etc.) to simulate the evolution in space and time of water, energy and carbon cycles. If 731 artificial influences and human intervention are included within the modelled processes (e.g. irrigation or water 732 management through reservoirs), the information required to describe the processes must also be integrated within 733 the modelling framework. Generally, this is achieved through a set of independent files used as input to the models. 734 Because of the temporal non-stationarity of some surface fields, typically associated with human intervention such 735 as land use and water use, but also due to climatic variation such as lake extent (new lakes forming or lakes 736 shrinking), input surface fields must be as representative as possible to the simulated period of interest. For 737 medium-range forecasting systems, this should be as close from present as possible, for example. When simulating 738 long periods, especially looking at past or future decades, caution must be given to results especially if some 739 surface fields which have substantially changed during the simulation period do not explicitly incorporate time 740 and instead are based on the most recent period, as they may not be representative to the full study period.

741 In addition, in recent years the horizontal resolution of global Earth system and environmental models has been 742 constantly increasing reaching the kilometre scale milestone, supported by the technological developments in the 743 field of High Performance Computers and the wealth of high resolution datasets freely available. This imposes 744 another condition to the input surface fields – it has to be of rather high horizontal resolution (i.e. ~2 and 6 km at 745 the Equator).

746 Thanks to the availability of a wide range of high resolution environmental data derived from the use of ground, 747 unconventional and satellite measurement sensors, new high resolution datasets describing the Earth's surface are 748 nowadays released regularly. Even though each dataset may have a very low absolute and root mean square errors 749 compared against available independent data, merging different datasets for modelling purposes (e.g. to model 750 hydrological surface parameters) might lead to questionable results and even model crash, due to possible 751 discontinuity or inconsistency in the combined datasets. In the specific case of hydrological modelling where river 752 flow is also represented, high horizontal resolution does not guarantee better modelling per se. Sources of 753 potentially large errors can be easily hidden in high resolution datasets. This is the case for instance of errors in 754 the Digital Elevation Models when they are used to obtain the rivers drainage network. Small errors in the 755 elevation of a grid cell can lead to a totally inaccurate representation of the location and the direction in which the 756 river is flowing in the model compared to reality. Mislocating a river or having a slightly inaccurate catchment 757 area can represent a trivial inaccuracy for most applications, but it can also lead to missed flood warning for 758 thousands of people within a flood awareness system. To benefit from different recent high resolution datasets 759 based on satellite and ground measurements, it is essential that a well-defined, thorough workflow is designed and 760 implemented so that the final products are consistent and compatible with each other, and can be used in 761 combination.





762 The work presented in this manuscript is focused not only on the final surface field set generation (i.e. 763 CEMS\_SurfaceFields\_2022), but also on deriving robust reproducible methodology that could be re-applied once 764 new versions of 25 or less input sources are released. Understanding of the methodology applied helps to interpret 765 values in the final surface fields and possibly even numerical model results that use these surface fields. The 766 collection of input sources and their preparation for actual use is a very important step as it includes going through 767 all technical documentation, comparison and verification of papers, and investigation of the actual data, as well 768 as data gridding, interpolation, and scaling. All input sources for CEMS\_SurfaceFields\_2022 are ranked according 769 to their quality and up-to-date in order to favour one value in ambiguous situations when several datasets provide 770 different information for the same location. Consistency check between all surface type fractions is carried out to 771 address that ambiguity during the merge of information of different origin (i.e. adjust fractions to sum to one in 772 each grid-cell). Some fields, like forest fraction, were rather straightforward to create from available source, yet 773 it was noted that prior correction of the source was needed to delete erroneous forest grid-cells from the Fox Basin 774 in Canada (the mismatch was only spotted during the investigation of the actual data, as it was absent from the 775 documentation). Other fields, like soil hydraulic properties, are created not only from the source information but 776 also from the forest fraction that had to be generated prior; the soil hydraulic property methodology also includes 777 several steps that have to be performed at the data native resolution (i.e. 250 m) using information from several 778 global fields simultaneously which becomes technically and computationally challenging. Surface fields with 779 clear multi-annual changes, like water demand maps, are created using temporal interpolation and extrapolation 780 from multiple data sources to create time series fields. A final and non-trivial task is to have all resulting fields on 781 the identical required grid without deterioration of the actual value precision, even after several file type 782 translations (e.g. local drainage direction field can be automatically checked and corrected if needed for required 783 boundaries only in PCRaster format, not NetCDF). Due to the number of data sources and surface fields required 784 to represent the main variables (i.e. 70) used in Earth system and environmental models, the overall effort to 785 generate the CEMS\_SurfaceFields\_2022 dataset (both human and computing resources) was substantial.

786 The CEMS\_SurfaceFields\_2022 dataset at 1 arc min (over Europe) and 3 arc min (globally) were tested and 787 indirectly validated using the LISFLOOD model through river discharge simulation (Grimaldi et al., 2024 in 788 preparation); they are the underlying surface fields of the EFAS version 5 and GloFAS version 4 operational 789 systems.

790 The CEMS\_SurfaceFields\_2022 dataset is a new data source open to all offering a kilometre-scale resolution of 791 high-quality data describing the Earth's surface, providing exceptional opportunity for the research and scientific 792 community to extend and multiply European and global applications in wide ranging fields of the water-energy-793 food nexus. 794

795 Data availability. The CEMS\_SurfaceFields\_2022 datasets are freely available for download from the JRC Data 796 Catalogue - global 3 arc min: https://data.jrc.ec.europa.eu/dataset/68050d73-9c06-499c-a441-dc5053cb0c86); 797 over Europe 1 arc min: https://data.jrc.ec.europa.eu/dataset/f572c443-7466-4adf-87aa-c0847a169f23.

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- 803

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- 1071 Annex
- 1072 Annex 1
- 1073 <u>Unit conversion to fraction</u>
- 1074 Hectare (ha):  $fraction = ha \cdot 10^4 / GridCellArea_{m^2}$ :
- 1075 Percentage (%):  $fraction = \frac{\%}{100};$
- 1076 Class (landcover type): *fraction* = 1, i.e. assumes full 100 % coverage of the grid-cell.
- 1077 Annex 2
- 1078 Soil depth

 $SD_1 = 50mm$ 

Soil depth layers are derived following Burek et al. (2014) in which the total soil depth is horizontally divided in three layers. The total soil depth is the 'absolute\_depth\_to\_bedrock' from SoilGrids250m, whereas root depths of forest and non-forest are derived from FAO56 and CGLS-LC100 dataset at SoilGrids250m native (~250 m) resolution (see Section 4.3 for more details). The methodology implemented for the creation of three soil layers is the following:

1084Soil depth layer 1 (surface)  $SD_1$  is assumed constant, equal to 50 mm all over the world for consistency with1085satellite-derived datasets (satellite signal penetration depth of 50 mm is a good approximation to take into account1086different meteorological conditions at different hour of the day globally based on Lv et al. (2018)), and follow Eq.1087(A1):

1089		
1090	Soil depth layer 2 (middle) $SD_2$ depends on the absolute depth to bedrock $adb$ – if it is equal or less than 300	) mm
1091	computation follow Eq. (A2), otherwise it is conditional of the root depths as per Eq. (A3), and must	meet
1092	requirement from Eq. (A4):	
1093		
1094	$SD_2 = (adb - SD_1)/2, adb \le 300mm$	(A2)
1095	$SD_2 = \min(root\_depth, (adb - 300mm - SD_1)), adb > 300m$	(A3)
1096	$SD_2 = 50$ mm, $SD_2 < 50$ mm	(A4)
1097		
1098	Soil depth layer 3 (bottom) $SD_3$ , is computed following Eq. (A5):	
1099		
1100	$SD3 = adb - (SD_1 + SD_2)$	(A5)
1101		
1102	This set of equations is used twice, once with the root depth of forest area and a second time with the root of	lepth
1103	of non-forested areas, resulting in a total of six soil depth layers computed at SoilGrids250m native resolution	on.

1104 1105 s

1088

Soil hydraulic parameters

(A1)





1106 Soil hydraulic parameters are derived by following three main steps (see Figure A1). 1107

1108 First, soil hydraulic properties are derived at native resolution by applying pedotransfer functions (PTFs) to each 1109 SoilGrids250m soil characteristics layer at each available depth. Pedotransfer functions translate field measured 1110 soil information (such as soil texture, pH and structure) into proprieties and parameters needed to describe soil 1111 processes. The PTFs implemented here are the ones proposed by Toth et al. (2015). Users can decide to derive 1112 soil proprieties from different PTFs, but the general principle presented here remains valid. 1113

1114 Second, the soil hydraulic parameters calculated at SoilGrids250m depths are vertically downscaled to the model 1115 soil depth (previously computed) by weighted average (Figure A1, Step 2 with theta saturated as an example) at 1116 the native SoilGrids250m resolution (~250 m).

1117

1118 Third, the soil hydraulic parameters at the final soil depths are upscaled from native to final resolution by average 1119 using forest and non-forest fraction layers as weights (Figure A1, Step 3).

1120





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- 1124