

1 **Photoenhanced sulfates formation by the heterogeneous uptake of SO<sub>2</sub> on non-**  
2 **photoactive mineral dust**

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7 **Short summary.** We provide direct evidences that light prominently enhances the conversion  
8 of SO<sub>2</sub> to sulfates on non-photoactive mineral dust, where triplet states of SO<sub>2</sub> (<sup>3</sup>SO<sub>2</sub>) can act  
9 as a pivotal trigger to generate sulfates. Photochemical sulfate formation depends on H<sub>2</sub>O, O<sub>2</sub>,  
10 and basicity of mineral dust. It is suggested that the SO<sub>2</sub> photochemistry on non-photoactive  
11 mineral dust significantly contributes to sulfates, highlighting previously unknown pathway to  
12 better explain the missing sources of atmospheric sulfates.

13

14 **Abstract.** Heterogeneous uptake of SO<sub>2</sub> on mineral dust is a predominant formation pathway  
15 of sulfates, whereas the contribution of photo-induced SO<sub>2</sub> oxidation to sulfates on the dust  
16 interfaces still remains unclear. Here, we investigated heterogeneous photochemical reactions  
17 of SO<sub>2</sub> on five mineral oxides (SiO<sub>2</sub>, kaolinite, Al<sub>2</sub>O<sub>3</sub>, MgO, and CaO) without photocatalytic  
18 activity. Light significantly enhanced the uptake of SO<sub>2</sub>, and its enhancement effects negatively  
19 depended on the basicity of mineral oxides. The initial uptake coefficient ( $\gamma_{0, \text{BET}}$ ) and the  
20 steady-state uptake coefficient ( $\gamma_{s, \text{BET}}$ ) of SO<sub>2</sub> positively relied on light intensity, relative  
21 humidity (RH) and O<sub>2</sub> content, while they exhibited a negative relationship with the initial SO<sub>2</sub>  
22 concentration. Rapid sulfate formation during photo-induced heterogeneous reactions of SO<sub>2</sub>  
23 with all mineral oxides was confirmed to be ubiquitous, and H<sub>2</sub>O and O<sub>2</sub> played the key roles  
24 in the conversion of SO<sub>2</sub> to sulfates. Specially, triplet states of SO<sub>2</sub> (<sup>3</sup>SO<sub>2</sub>) was suggested to be  
25 the trigger for photochemical sulfate formation. Atmospheric implications supported a  
26 potential contribution of interfacial SO<sub>2</sub> photochemistry on non-photoactive mineral dust to  
27 atmospheric sulfate sources.

28

29 **Keywords:** SO<sub>2</sub>; Sulfates; Non-photoactive mineral dust; Heterogeneous photochemistry

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31

## 32 **1 Introduction**

33 As an important trace gas in the atmosphere, SO<sub>2</sub> is mainly emitted by volcanic eruption and  
34 fuel combustion. There is an uneven distribution of atmospheric SO<sub>2</sub> concentrations that show  
35 a distinctive seasonal and regional differentiation. Typical ratios of SO<sub>2</sub> in the troposphere are  
36 below 0.5 ppb for a clean weather in remote areas, rising to around several hundred ppb during  
37 the polluted days in urban regions (Ma et al., 2020). About half of SO<sub>2</sub> is oxidized to sulfates  
38 (He et al., 2012), which is one of the most significant compositions in fine particles. **The mass**  
39 **fraction of sulfates in PM<sub>2.5</sub> is high up to 30%** (Shao et al., 2019), especially in polluted regions  
40 where high-sulfur fuels are usually used (Olson et al., 2021). They significantly alter  
41 physicochemical properties of aerosols in terms of hygroscopicity, acidity and light absorption  
42 property (Chan and Chan, 2003; Cao et al., 2013; Lim et al., 2018). Sulfates also pose a human  
43 health risk through causing respiratory illness and cardiovascular (Shiraiwa et al., 2017). In  
44 addition, the deposition of sulfates leads to adverse effects on ecosystems via the acidification  
45 of soils and lakes (Golobokova et al., 2020). Therefore, the oxidation of SO<sub>2</sub> to form sulfates  
46 has attracted widespread attention in the past decades.

47 The conversion of SO<sub>2</sub> to sulfates in the atmosphere usually occurs in different phases: gas-  
48 phase oxidation of SO<sub>2</sub> by hydroxyl radicals ( $\bullet$ OH) or Criegee intermediate radicals (Mauldin  
49 et al., 2012; Davis et al., 1979); aqueous-phase reaction of SO<sub>2</sub> with O<sub>3</sub>, peroxides or transition  
50 metal ions dissolved in cloud and fog droplets (Alexander et al., 2009; Herrmann et al., 2000;  
51 Harris et al., 2013; Liu et al., 2020a; Li et al., 2020); and heterogeneous SO<sub>2</sub> uptake on aerosols  
52 including authentic mineral dust, soot, inorganic ion and organic compounds (Adams et al.,  
53 2005; He et al., 2018a; **Ye et al., 2018; Wang et al., 2019; Yao et al., 2019; Zhang et al., 2020a;**  
54 **Liu et al., 2020; Liu et al., 2021**). However, the oxidation of SO<sub>2</sub> in gas and aqueous phases  
55 fails to explain high sulfate concentrations under polluted conditions. Model simulation

56 suggests that the rapid sulfate formation can be attributed to the heterogeneous SO<sub>2</sub> uptake (Li  
57 et al., 2017). A positive relationship between the fraction of sulfates and mineral dust in haze  
58 days has been reported, implying that mineral dust may account for the formation of sulfates  
59 (Wang et al., 2020a). Moreover, a large amount of sulfates was observed to be formed on the  
60 surface of mineral dust after long-distance transport (Prospero, 1999). Thus, investigating the  
61 heterogeneous oxidation of SO<sub>2</sub> on mineral dust can provide basic data for the model  
62 calculation to evaluate atmospheric sulfates.

63 Mineral dust, regarded as the dominant component of particulate matters in the atmosphere,  
64 accounts for about 30%–60% mass fractions of global aerosols (Dentener et al., 1996; Peng et  
65 al., 2012). It primarily contains SiO<sub>2</sub> (40%–80%), followed by Al<sub>2</sub>O<sub>3</sub> (10%–15%), Fe<sub>2</sub>O<sub>3</sub>  
66 (6%–13%), CaO (3%–10%), MgO (1%–7%) and TiO<sub>2</sub> (0.1%–5%) (Urupina et al., 2021;  
67 Urupina et al., 2019; Usher et al., 2003). Mineral dust can provide active sites for adsorption  
68 and reaction of gases. Up to now, the heterogeneous SO<sub>2</sub> uptake on authentic mineral aerosols  
69 and model mineral oxides has been widely reported (Ma et al., 2019; Goodman et al., 2001;  
70 Wang et al., 2018; Wang et al., 2020b), with various uptake coefficients ( $\gamma$ ) of SO<sub>2</sub> varying  
71 from 10<sup>-9</sup> to 10<sup>-4</sup> (Urupina et al., 2019; Usher et al., 2002).

72 It was recognized that light could significantly enhance heterogeneous conversion of SO<sub>2</sub> to  
73 sulfates on the surface of photocatalytic mineral dust (Chen et al., 2021; Li et al., 2019; Wang  
74 et al., 2020b). Electron-hole pairs are produced via photo-induced electrons from the valence  
75 band to the conduction band of photocatalytic metal oxides, and then react with H<sub>2</sub>O and O<sub>2</sub> to  
76 generate reactive oxygen species (ROS), such as •OH and •O<sub>2</sub><sup>-</sup> (Chu et al., 2019). Sulfates are  
77 produced by the heterogeneous reactions of SO<sub>2</sub> with ROS (Park and Jang, 2016; Park et al.,  
78 2017; Langhammer et al., 2020; Bounechada et al., 2017). In particular, due to the large  
79 abundance of non-photoactive mineral dust (more than 85% mass of total mineral dust in the  
80 atmosphere) (Usher et al., 2003; Liu et al., 2022), revealing the photooxidation processes of  
81 SO<sub>2</sub> on these mineral dust is of great importance to better reevaluate the sulfate formation on  
82 aerosols in the global scale.

83 Hence, photochemical SO<sub>2</sub> uptake and sulfate formation on non-photoactive mineral oxides

84 were firstly investigated using a flow reactor and an *in situ* diffuse reflectance infrared Fourier  
85 transform spectroscopy (DRIFTS). The SO<sub>2</sub> conversion to sulfates was examined under various  
86 conditions, and the roles of light intensity, SO<sub>2</sub> concentration, H<sub>2</sub>O, O<sub>2</sub> and basicity of mineral  
87 oxides were determined. Reaction mechanisms and atmospheric implications were proposed,  
88 highlighting a new and important pathway accounting for photochemical uptake of SO<sub>2</sub> to form  
89 sulfates on the non-photoactive surfaces.

90

## 91 **2 Experimental methods**

### 92 **2.1 Materials**

93 Analytical grade SiO<sub>2</sub> (Sinopharm Chemical Reagent Co., Ltd.), kaolinite (Sinopharm  
94 Chemical Reagent Co., Ltd.), Al<sub>2</sub>O<sub>3</sub> (Alfa Aesar), MgO (Sigma-Aldrich), and CaO (Sigma-  
95 Aldrich) were used in the experiments. Through the nitrogen Brunauer-Emmett-Teller (BET)  
96 physisorption analysis, their specific surface areas were detected to be 0.419, 6.407, 8.137,  
97 10.948 and 6.944 m<sup>2</sup> g<sup>-1</sup>, respectively. **With BaSO<sub>4</sub> used as the reference, the ultraviolet-visible**  
98 **(UV-vis) light absorption spectra of samples (Figure S1) in the wavelength range of 300–800**  
99 **nm were obtained by the Shimadzu UV-2550 spectrophotometer, which was equipped with**  
100 **diffuse reflection integrating sphere attachment.** The solid powder (0.2–5 g) was uniformly  
101 dispersed into 10.0 mL ethanol solution. The mixed liquid was poured into a rectangle quartz  
102 sample dish (14.0 cm × 7.0 cm) and dried to form a solid coating in an oven at 353 K for 10 h.  
103 SO<sub>2</sub> standard gas (50 ppm in N<sub>2</sub>, Shenyang Air Liquide Co., LTD) and high purity N<sub>2</sub> and O<sub>2</sub>  
104 (99.999 vol.%, Shenyang Air Liquide Co., LTD) were used as received. The solid sample  
105 powder (0.2 g) was immersed into 10 mL deionized water (20 mg mL<sup>-1</sup>), and then the  
106 suspension was vigorously stirred for 10 min. The pH of SiO<sub>2</sub>, kaolinite, Al<sub>2</sub>O<sub>3</sub>, MgO and CaO  
107 suspension was measured to be 6.27, 6.58, 9.33, 10.61 and 12.72 using a pH meter, respectively,  
108 which was employed to characterize the basicity of mineral oxides.

### 109 **2.2 Rectangular flow reactor**

110 The uptake experiments of SO<sub>2</sub> on mineral dust were performed in a horizontal rectangular  
111 flow reactor (26.0 cm length × 7.5 cm width × 2.0 cm height), which was depicted in Figure

112 S2. In a previous study, a similar rectangular flow reactor was designed and the feasibility of  
113 the reactor has been verified (Knopf et al., 2007). The reactor was made of quartz to allow the  
114 transmission of light. The temperature was maintained at 298 K by circulating temperature-  
115 controlled water through the outer jacket of the reactor. Synthetic air with a N<sub>2</sub>/O<sub>2</sub> volume ratio  
116 of 4:1 was introduced into the flow reactor, and its total flow rate was 1000 mL min<sup>-1</sup>. The  
117 Reynolds number (*Re*) was calculated to be 28.2 (*Re* < 200), as described in the Supporting  
118 Information, indicating a laminar flow state. SO<sub>2</sub> with high purity N<sub>2</sub> (100 mL min<sup>-1</sup>) as carrier  
119 gas were introduced into the reactor through a movable T-shaped injector equipped with six  
120 exit holes of 0.2 mm diameter, so that the gas could be uniformly distributed over the width of  
121 the reactor. The SO<sub>2</sub> concentration was 40–200 ppb and measured with a SO<sub>2</sub> analyzer (Thermo  
122 43i). Wet N<sub>2</sub> generated with a bubbler containing deionized water was introduced by two  
123 parallel inlets on the side of T-shaped injector. Relative humidity (RH, 10%–75%) was  
124 controlled by regulating the ratio of dry N<sub>2</sub> to wet N<sub>2</sub> and measured via a hygrometer (Center  
125 314). The equivalent layer numbers of water on surface was 0.9–4.0 according to the Brunauer-  
126 Emmett-Teller (BET) model (Sumner et al., 2004), and the thickness of the film of adsorbed  
127 water varied between 2.7–12 nm at RH=10%–75%. There were three equally spaced exhaust  
128 ports to mitigate the outlet turbulence. A xenon lamp (CEL-LAX500, China Education Au-light  
129 Co., Ltd) was used to simulate sunlight and vertically located above the reactor. A filter was  
130 placed on the reactor to remove the light with wavelengths shorter than 300 nm. The spectrum  
131 irradiance of the xenon lamp was displayed in Figure S3 and measured using a calibrated  
132 spectroradiometer (ULS2048CL-EVO, Avantes). The spectral irradiance was measured inside  
133 the reactor, after passing the water cooling and in the absence of a sample. The total irradiance  
134 ( $0\text{--}7.93 \times 10^{16}$  photons cm<sup>-2</sup> s<sup>-1</sup>) on the coating can be adjusted by varying the distance of the  
135 xenon lamp to the reactor.

### 136 2.3 Uptake coefficient of SO<sub>2</sub>

137 The heterogeneous reaction kinetics of SO<sub>2</sub> with mineral dust can be described by a pseudo-  
138 first-order reaction. SiO<sub>2</sub> was taken as an example, and Figure S4 showed a linear relationship  
139 between the natural logarithm of the SO<sub>2</sub> concentration and the reaction time. The apparent rate

140 constant ( $k_{\text{obs, SiO}_2}$ ) of SO<sub>2</sub> with SiO<sub>2</sub> can be calculated using the equation 1,

141 
$$\frac{\ln(C_0/C_t)}{t} = k_{\text{obs, SiO}_2} \quad (1)$$

142 where  $C_0$  and  $C_t$  (ppb) are the initial SO<sub>2</sub> concentration and the SO<sub>2</sub> concentration, respectively;

143  $t$  was calculated by dividing the length of the reactive surface by the average flow velocity. **The**

144 **loss of SO<sub>2</sub> on the internal wall of the reactor in blank experiments was carried out under**

145 **various conditions (Figure S5 as an example), and it has been deducted for the  $\gamma$  calculation.**

146 Assuming that the wall loss was constant in the experiments with and without samples, the

147 geometric uptake coefficient ( $\gamma_{\text{geo}}$ ) was determined by the equation 2 (Knopf et al., 2007),

148 
$$\gamma_{\text{geo}} = \frac{4Vk}{S\omega} \quad (2)$$

149 where  $k$  (s<sup>-1</sup>),  $V$  ( $4 \times 10^{-4}$  m<sup>3</sup>),  $S$  ( $9.8 \times 10^{-3}$  m<sup>2</sup>) and  $\omega$  ( $314.05$  m s<sup>-1</sup>) are the reaction rate

150 constant, the volume of the rectangular reactor, the surface area of the sample dish, and the

151 mean molecular speed of SO<sub>2</sub>, respectively.

152 The uptake process of SO<sub>2</sub> on SiO<sub>2</sub> depended on the reaction of SO<sub>2</sub> with SiO<sub>2</sub> and the mass

153 transport of SO<sub>2</sub> to the surface. It can be expressed with the equation 3,

154 
$$k'_{\text{r, SiO}_2} = \left[ \frac{1}{k_{\text{obs, SiO}_2} - k_{\text{obs, wall}}} - \frac{a}{N_u D} \right]^{-1} \quad (3)$$

155 where  $k_{\text{obs, SiO}_2}$  and  $k_{\text{obs, wall}}$  (s<sup>-1</sup>) are the apparent rate constants measured with and without

156 SiO<sub>2</sub> samples, respectively.  $k'_{\text{r, SiO}_2}$  is the reaction rate constant of SO<sub>2</sub> accounting for the

157 diffusion effect;  $D$  ( $0.1337$  cm<sup>2</sup> s<sup>-1</sup>) is the diffusion coefficient of SO<sub>2</sub> in air;  $a$  (1 cm) is one

158 half height of the flow reactor;  $N_u$  is the Nusselt numbers obtained with a calculation method

159 from Solbrig and Gidaspow (1967), which represents the mass transport. Then, the corrected  $\gamma$

160 can be calculated by the equation 2 where  $k$  was replaced by  $k'_{\text{r, SiO}_2}$ . In our experiments, the

161 correction for  $\gamma$  was estimated to be approximate 10%. Initial uptake coefficients ( $\gamma_0$ ) and

162 steady-state uptake coefficients ( $\gamma_s$ ) were calculated by averaging the signals within the 1.0 and

163 40–60 min reaction time, respectively.

164 To understand the diffusion depth of SO<sub>2</sub> and determine the interaction of SO<sub>2</sub> with the

165 underlying layers of SiO<sub>2</sub>, the uptake of SO<sub>2</sub> as a function of the SiO<sub>2</sub> mass under irradiation

166 was shown in Figure S6. The  $\gamma$  exhibited a linear increase in the SiO<sub>2</sub> mass range of 0.05–2.0  
167 g, while it remained unchanged at the SiO<sub>2</sub> mass > 3.0 g. Therefore, the uptake coefficient of  
168 SO<sub>2</sub> in the linear regions was normalized using the BET surface area of SiO<sub>2</sub> by the equation 4  
169 (Brunauer et al., 1938),

$$170 \gamma_{\text{BET}} = \frac{S_{\text{geo}} \times \gamma_{\text{geo}}}{S_{\text{BET}} \times m_{\text{SiO}_2}} \quad (4)$$

171 where  $\gamma_{\text{BET}}$  is the SO<sub>2</sub> uptake coefficient normalized to the BET surface area;  $S_{\text{geo}}$  ( $9.8 \times 10^{-3}$   
172  $\text{m}^2$ ) is the geometric area of the sample dish;  $S_{\text{BET}}$  ( $0.419 \text{ m}^2 \text{ g}^{-1}$ ) is the BET surface area of  
173 SiO<sub>2</sub>;  $m_{\text{SiO}_2}$  ( $0.05\text{--}2.0 \text{ g}$ ) is the SiO<sub>2</sub> mass. The same method was also used to calculate the  
174 uptake coefficients of SO<sub>2</sub> on other mineral oxides.

## 175 2.4 In Situ DRIFTS analysis

176 The changes in the chemical compositions on mineral oxides in the SO<sub>2</sub> uptake process were  
177 investigated by the Fourier transform infrared (FTIR) spectrometer (Thermo Nicolet iS50)  
178 equipped with an *in situ* diffuse reflectance accessory and a mercury cadmium telluride (MCT)  
179 detector. About 14 mg mineral oxides was placed into a stainless-steel cup inside the reaction  
180 cell. To remove adsorbed impurities, SiO<sub>2</sub> was purged with a 150 mL min<sup>-1</sup> airflow (N<sub>2</sub>/O<sub>2</sub>  
181 volume ratio = 4:1) at RH=40% for 1 h. Then, a background spectrum of unreacted samples  
182 was collected. SO<sub>2</sub> (2 ppm) was introduced into the reaction cell, and the IR spectra was  
183 recorded as a function of time at a resolution of 4 cm<sup>-1</sup> by averaging 100 scans. The light from  
184 the xenon lamp (500 W) was transmitted into the DRIFTS reaction cell via a fiber. To verify  
185 the role of intermediate, Ru(bpy)<sub>3</sub>(Cl)<sub>2</sub> and NaHCO<sub>3</sub>, acting as <sup>3</sup>SO<sub>2</sub> and •OH scavengers  
186 (Bulgakov and Safonova, 1996; Gen et al., 2019a), respectively, were mixed with SiO<sub>2</sub> powder  
187 in an agate mortar, and the mixture was put in the reaction cell of DRIFTS.

188

## 189 3 Results and discussion

### 190 3.1 Photo-enhanced uptake of SO<sub>2</sub>

191 Acting as the most abundant mineral oxides, SiO<sub>2</sub> was firstly used to investigate the uptake  
192 behaviors of SO<sub>2</sub>. As shown in Figure 1A, when SO<sub>2</sub> (80 ppb) was exposed to SiO<sub>2</sub> in the dark,

193 the SO<sub>2</sub> concentration decreased to 70 ppb, and then it quickly increased and reached the steady  
194 state after 20 min. Upon exposure to SiO<sub>2</sub> under irradiation, the SO<sub>2</sub> concentration exhibited a  
195 greater drop than that in the dark. The deactivation of SO<sub>2</sub> uptake on SiO<sub>2</sub> was very slight after  
196 20 mins under irradiation. These suggest that light can significantly promote the heterogeneous  
197 reaction of SO<sub>2</sub> on SiO<sub>2</sub>. When SO<sub>2</sub> **did not** contact with SiO<sub>2</sub>, its concentration recovered  
198 rapidly. The desorption of SO<sub>2</sub> was observed when SO<sub>2</sub> was isolated from SiO<sub>2</sub> in the dark and  
199 under irradiation, indicating that the physical adsorption partially contributed to the SO<sub>2</sub> loss  
200 during the photochemical process. **The proportion of the desorbed SO<sub>2</sub> during the uptake**  
201 **process can be quantified by dividing the integral of reversible desorption of SO<sub>2</sub> ( $t = 80\text{--}100$**   
202 **min) into the integral of the SO<sub>2</sub> uptake ( $t = 20\text{--}80$  min), which was calculated to be 95% and**  
203 **12% in the dark and under irradiation, respectively. This implies that SO<sub>2</sub> uptake in the dark**  
204 **was primarily ascribed to the physical adsorption of SO<sub>2</sub>, while SO<sub>2</sub> uptake under irradiation**  
205 **was mainly attributed to chemical processes or irreversible adsorption.**

206 The uptake coefficients of SO<sub>2</sub> on SiO<sub>2</sub> as a function of irradiation intensity were shown in  
207 Figure 1B. The errors in all figures are the standard deviations of three repetitive experiments.  
208 Both  $\gamma_{0, \text{BET}}$  and  $\gamma_{s, \text{BET}}$  displayed a well linear relationship with the irradiation intensity, further  
209 confirming the photochemical nature for the reactions of SO<sub>2</sub> on SiO<sub>2</sub>. In particular,  $\gamma_{0, \text{BET}}$  and  
210  $\gamma_{s, \text{BET}}$  on SiO<sub>2</sub> under simulated solar irradiation was comparable with those ( $10^{-7}\text{--}10^{-6}$ ) on  
211 Gobi Desert dust (GDD) and Arizona Test Dust (ATD) under UV irradiation, which contained  
212 photocatalytic metal oxides (Park et al., 2017). As for the SO<sub>2</sub> uptake on TiO<sub>2</sub>,  $\gamma_{0, \text{BET}}$  and  $\gamma_{s, \text{BET}}$   
213 were measured to be  $10^{-6}$  and  $10^{-7}$ , respectively, by using the flow tube (Ma et al., 2019), which  
214 were similar to our results. **It should be pointed out that the similar uptake coefficient did not**  
215 **mean the comparable ability of photoactive and non-photoactive mineral oxides to SO<sub>2</sub> uptake,**  
216 **since the uptake coefficient was highly dependent on environmental conditions (SO<sub>2</sub>**  
217 **concentration, relative humidity, mineral oxides mass, light source and pressure) and reactor**  
218 **type (chamber and flow tube reactor), and the uptake coefficients mentioned here were not**  
219 **obtained under the exact same reaction conditions used in our study. The purities of different**



220 mineral substances are 95%–98%. If photoactive impurities mainly contributed to the SO<sub>2</sub>  
 221 uptake in the experiment, the SO<sub>2</sub> uptake coefficient on impurities should be 20–50 times  
 222 higher than the current SO<sub>2</sub> uptake coefficient and range from 10<sup>-5</sup> to 10<sup>-3</sup>. The SO<sub>2</sub> uptake  
 223 coefficient on photoactive substances was reported to be 10<sup>-7</sup>–10<sup>-6</sup> in previous studies (Ma et  
 224 al., 2019; Park et al., 2017). Thus, the impurities in minerals were less likely responsible for  
 225 the SO<sub>2</sub> uptake.

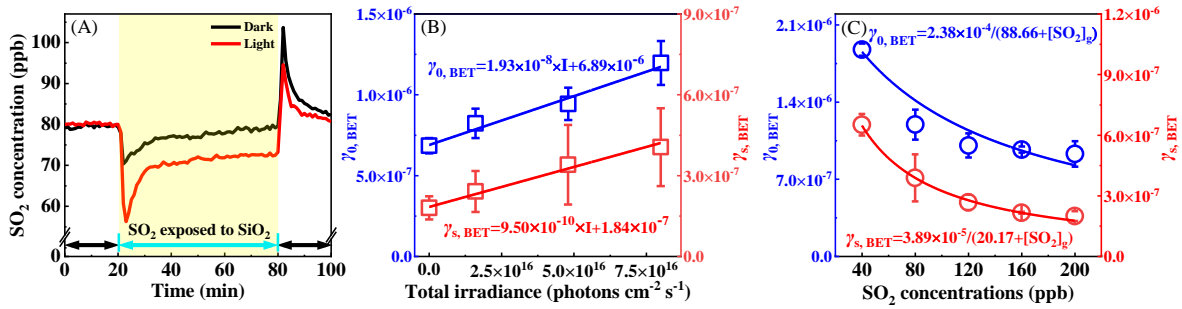
226 Figure 1C shows the evolution of  $\gamma_{0, \text{BET}}$  and  $\gamma_{s, \text{BET}}$  at different SO<sub>2</sub> concentrations under  
 227 irradiation. An inverse dependence of  $\gamma_{0, \text{BET}}$  and  $\gamma_{s, \text{BET}}$  on the SO<sub>2</sub> concentration was observed,  
 228 meaning that both initial and steady-state uptake reactions were lower efficient at higher SO<sub>2</sub>  
 229 concentrations. The uptake of gases on the solid surfaces usually follows the Langmuir-  
 230 Hinshelwood (L-H) mechanism (Ammann et al., 2003; Zhang et al., 2020b), suggesting that  
 231 gaseous molecules are quickly absorbed on the surfaces, and then the reactions occur among  
 232 the absorbed species. Assuming that the adsorption of SO<sub>2</sub> on SiO<sub>2</sub> is in accord with the  
 233 Langmuir isotherm, the dependence of  $\gamma$  on the SO<sub>2</sub> concentration can be described by the  
 234 equation 5 (Zhang et al., 2020b),

$$235 \gamma = \frac{(4V/S\omega)k[\text{SiO}_2]_{\text{T}}K_{\text{SO}_2}}{1 + K_{\text{SO}_2}[\text{SO}_2]_{\text{g}}} \quad (5)$$

236 where [SO<sub>2</sub>]<sub>g</sub> is the concentration of gaseous SO<sub>2</sub>; [SiO<sub>2</sub>]<sub>T</sub> is the total number of active sites  
 237 on SiO<sub>2</sub>;  $k$  is the reaction rate constant of SO<sub>2</sub> absorbed on SiO<sub>2</sub>;  $K_{\text{SO}_2}$  represents the Langmuir  
 238 adsorption constant of SO<sub>2</sub>. Because the SiO<sub>2</sub> mass remained constant during the reaction, the  
 239 equation 5 can be written as the equation 6,

$$240 \gamma = \frac{a}{1 + K_{\text{SO}_2}[\text{SO}_2]_{\text{g}}} \quad (6)$$

241 where  $a = (4V/S\omega)k[\text{SiO}_2]_{\text{T}}K_{\text{SO}_2}$ . As shown in Figure 1C, the equation 6 can well describe the  
 242 correlation of the SO<sub>2</sub> uptake coefficient with the SO<sub>2</sub> concentration, suggesting that the L-H  
 243 mechanism can explain the influence of the SO<sub>2</sub> concentration on  $\gamma_{0, \text{BET}}$  and  $\gamma_{s, \text{BET}}$ .



244

245 **Figure 1.** (A) The temporal variation of the SO<sub>2</sub> concentration on SiO<sub>2</sub> in the dark and under  
 246 irradiation ( $7.93 \times 10^{16}$  photons cm<sup>-2</sup> s<sup>-1</sup>); The background changes of the SO<sub>2</sub> concentration

247 in the blank reactor have been deducted. (B) The  $\gamma_{0, \text{BET}}$  and  $\gamma_{s, \text{BET}}$  of SO<sub>2</sub> on SiO<sub>2</sub> as a

248 function of the light intensity. (C) The  $\gamma_{0, \text{BET}}$  and  $\gamma_{s, \text{BET}}$  of SO<sub>2</sub> on SiO<sub>2</sub> at different SO<sub>2</sub>

249 concentrations under irradiation ( $7.93 \times 10^{16}$  photons cm<sup>-2</sup> s<sup>-1</sup>); The fitting lines for  $\gamma_{0, \text{BET}}$

250 and  $\gamma_{s, \text{BET}}$  were based on the Langmuir-Hinshelwood mechanism using equation 6. Reaction

251 conditions: SiO<sub>2</sub> mass of 0.2 g, temperature of 298 K, RH of 40% and O<sub>2</sub> content of 20%.

### 252 3.2 Photo-induced formation of sulfates by the SO<sub>2</sub> uptake

253 To investigate the products formed on SiO<sub>2</sub>, *in situ* DRIFTS spectra were recorded, as shown

254 in Figure 2. The band at 1359 cm<sup>-1</sup> was assigned to physically-adsorbed SO<sub>2</sub> on SiO<sub>2</sub> (Urupina

255 et al., 2019). The bidentate sulfate and bisulfate contributed to the bands at 1260 and 1229/1074

256 cm<sup>-1</sup> (Urupina et al., 2019; Yang et al., 2020), respectively. The bands at 1038 cm<sup>-1</sup> may be

257 related to the monodentate sulfite (Yang et al., 2019; Wang et al., 2019). It was noted that the

258 intensity of physically-adsorbed SO<sub>2</sub> (1359 cm<sup>-1</sup>) under irradiation was lower than that in the

259 dark (Figure S7), which may be ascribed to further conversion of SO<sub>2</sub> absorbed on SiO<sub>2</sub> under

260 irradiation. Especially, the sulfate bands (1260, 1229 and 1074 cm<sup>-1</sup>) only appeared under

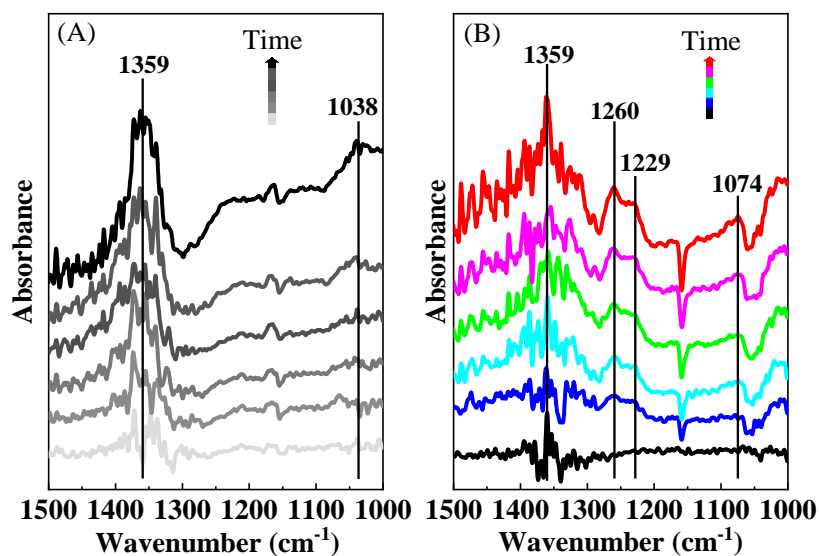
261 irradiation, while the sulfites (1038 cm<sup>-1</sup>) were only detected in the dark. This suggests that

262 light changed the SO<sub>2</sub> conversion pathways on SiO<sub>2</sub>. As shown in Figure S7, the bands at

263 1157/1055 cm<sup>-1</sup> were assigned to the asymmetric stretching of Si-O (Hu et al., 2003). Sulfate

264 generated on the surface may interact with SiO<sub>2</sub>, leading to a decrease in the intensity of peaks

265 (1157/1055 cm<sup>-1</sup>).



266

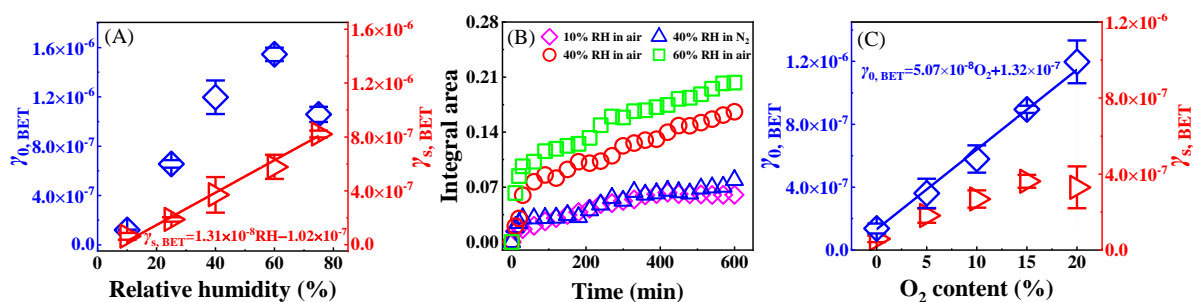
267 **Figure 2.** *In situ* DRIFTS spectra of SiO<sub>2</sub> during the uptake process of SO<sub>2</sub> (2 ppm) in the  
 268 dark (A) and under irradiation (B). Reaction conditions: RH of 40%, temperature of 298 K  
 269 and O<sub>2</sub> content of 20%.

### 270 3.3 Key roles of H<sub>2</sub>O and O<sub>2</sub> in photochemical conversion of SO<sub>2</sub> to sulfates

271 Figure S8A shows temporal variations of the SO<sub>2</sub> concentration in the reaction with SiO<sub>2</sub> at  
 272 RH=10% and 60% under irradiation. The uptake of SO<sub>2</sub> was very weak at RH=10%, whereas  
 273 it was obvious at RH=60%. Moreover, H<sub>2</sub>O markedly prolonged the time to reach the steady-  
 274 state uptake of SO<sub>2</sub>. This definitely determines that H<sub>2</sub>O plays a distinct enhancement role in  
 275 the photochemical uptake of SO<sub>2</sub>. As shown in Figure 3A,  $\gamma_{0, \text{BET}}$  had a continuous increase  
 276 from  $(1.20 \pm 0.04) \times 10^{-7}$  to  $(1.54 \pm 0.07) \times 10^{-6}$  with increasing the RH in the 10%–60% range,  
 277 but it decreased to  $(1.05 \pm 0.09) \times 10^{-6}$  at RH=75%. The  $\gamma_{s, \text{BET}}$  linearly depended on the RH,  
 278 and linear fitting to  $\gamma_{s, \text{BET}}$  versus RH yielded the equation  $\gamma_{s, \text{BET}} = 1.31 \times 10^{-8} \times \text{RH} - 1.02 \times 10^{-7}$ .  
 279 Adsorbed H<sub>2</sub>O promoted the hydration and dissociation of SO<sub>2</sub> (Huang et al., 2015), and it may  
 280 generate reactive oxygen species (ROS) such as •OH or HO<sub>2</sub> radicals to oxidize SO<sub>2</sub> under  
 281 irradiation (Li et al., 2020; Ma et al., 2019), which would lead to positive effects of RH on the  
 282 SO<sub>2</sub> uptake. Adsorbed H<sub>2</sub>O also occupied adsorptive and active sites on the surface, and  
 283 produced the competition with SO<sub>2</sub>. When this competitive role was dominated, the uptake of

284 SO<sub>2</sub> would be hindered.

285 The DRIFTS spectra of SiO<sub>2</sub> during the SO<sub>2</sub> uptake at different RHs are shown in Figure  
286 S9A. The band intensities of sulfates (1260 and 1229 cm<sup>-1</sup>) at RH=60% were greatly stronger  
287 than those at RH=10%, suggesting that H<sub>2</sub>O significantly promotes the sulfate formation. To  
288 further investigate the influence of H<sub>2</sub>O on the sulfate formation, the integrated area of sulfates  
289 in the DRIFTS spectra (1289–1202 cm<sup>-1</sup>) as a function of the time at different RHs is illustrated  
290 in Figure 3B. Sulfates exhibited a fast formation in the initial 30 min at any RH, and then they  
291 were continuously generated at a relatively slow rate. Absorptive sites for SO<sub>2</sub> can be blocked  
292 because of the accumulation of H<sub>2</sub>O and products (sulfites and sulfates), resulting in the gradual  
293 deactivation of the surface. It was noted that sulfates had a more distinct formation trend with  
294 increasing the RH, revealing that H<sub>2</sub>O can act as an important participator in the production of  
295 sulfates by the photochemical uptake of SO<sub>2</sub> on SiO<sub>2</sub>.



296

297 **Figure 3.** (A) The dependence of  $\gamma_{0, \text{BET}}$  and  $\gamma_{s, \text{BET}}$  on RH. (B) Integrated area of sulfates in  
298 DRIFTS spectra (1289–1202 cm<sup>-1</sup>) as a function of time. (C) The dependence of  $\gamma_{0, \text{BET}}$  and  
299  $\gamma_{s, \text{BET}}$  on O<sub>2</sub>. Reaction conditions: SiO<sub>2</sub> mass of 0.2 g, irradiation intensity of  $7.93 \times 10^{16}$   
300 photons cm<sup>-2</sup> s<sup>-1</sup>, temperature of 298 K, O<sub>2</sub> content of 20% for (A) and RH of 40% for (B).

301 Figure S8B displays effects of O<sub>2</sub> on the photochemical uptake of SO<sub>2</sub> on SiO<sub>2</sub>. Negligible  
302 SO<sub>2</sub> uptake occurred in N<sub>2</sub>, while there was a significant decrease in the SO<sub>2</sub> concentration in  
303 air. The  $\gamma_{0, \text{BET}}$  greatly increased from  $(1.37 \pm 0.45) \times 10^{-7}$  in N<sub>2</sub> to  $(1.19 \pm 0.13) \times 10^{-6}$  in 20%  
304 O<sub>2</sub> (Figure 3C), confirming that O<sub>2</sub> was involved in the reaction of SO<sub>2</sub> on SiO<sub>2</sub>. The  $\gamma_{s, \text{BET}}$

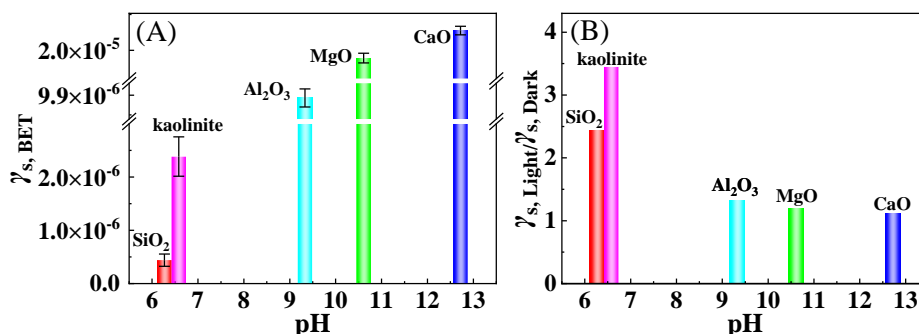
305 displayed different dependence behaviors on O<sub>2</sub>. It exhibited an increase from  $(7.10 \pm 2.85) \times$   
306  $10^{-8}$  in N<sub>2</sub> to  $(4.37 \pm 0.58) \times 10^{-7}$  in 15% O<sub>2</sub>, whereas it remained unchanged in 20% O<sub>2</sub>.

307 DRIFTS spectra of SiO<sub>2</sub> during the SO<sub>2</sub> uptake in N<sub>2</sub> and air was compared in Figure S9B.  
308 In both air and N<sub>2</sub>, the bands of absorbed SO<sub>2</sub> (1359 cm<sup>-1</sup>), sulfates (1260, 1229 and 1074  
309 cm<sup>-1</sup>). Nevertheless, their intensities in N<sub>2</sub> were weaker than those in air. According to the  
310 integrated area of sulfates in the DRIFTS spectra (1289–1202 cm<sup>-1</sup>) as a function of time, the  
311 formation trends of sulfates were similar in N<sub>2</sub> and air (Figure 3B), while the sulfate formation  
312 rate in N<sub>2</sub> was obviously lower than that in air, meaning that O<sub>2</sub> enhanced the sulfate production.  
313 It was reported that the production rate of sulfates from the SO<sub>2</sub> uptake on TiO<sub>2</sub> and by the  
314 photolysis of nitrates under UV irradiation in N<sub>2</sub> was also smaller than that in air (Ma et al.,  
315 2019; Gen et al., 2019b). In addition, it was noted that sulfates can be generated in N<sub>2</sub>, meaning  
316 that O<sub>2</sub> was not necessary and some pathways contributed to sulfates without O<sub>2</sub>.

### 317 **3.4 Ubiquitously photoenhanced conversion of SO<sub>2</sub> to sulfates**

318 To better assess the potential for photochemical conversion of SO<sub>2</sub> to sulfates, the SO<sub>2</sub> uptake  
319 experiments were further performed for typical mineral oxides without photocatalytic activity.  
320 As displayed in Figure S10, more obvious uptake behaviors of SO<sub>2</sub> on kaolinite, Al<sub>2</sub>O<sub>3</sub>, MgO  
321 and CaO were observed under irradiation when compared to those in the dark. Figure 4A shows  
322 that there was the largest  $\gamma_{s, \text{BET}}$  for CaO among five minerals, and  $\gamma_{s, \text{BET}}$  positively depended  
323 on the basicity (pH) of mineral oxides. Basic oxides generally contains more surface hydroxyls,  
324 which is in favor of sulfite and sulfate formation to enhance the heterogeneous uptake of SO<sub>2</sub>  
325 (Zhang et al., 2006). The ratios of steady-state uptake coefficients under irradiation to those in  
326 the dark ( $\gamma_{s, \text{Light}}/\gamma_{s, \text{Dark}}$ ) were larger than 1.0 for all mineral oxides (Figure 4B). **The**  
327 **experiments for the pH dependence on SiO<sub>2</sub> have been also performed (Figure S11). The pH**  
328 **of SiO<sub>2</sub> suspension was adjusted to pH = 9, and  $\gamma_{s, \text{BET}}$  and  $\gamma_{s, \text{Light}}/\gamma_{s, \text{Dark}}$  were determined to**  
329 **be  $(8.79 \pm 0.85) \times 10^{-6}$  and 1.31, respectively. These results suggest that light can generally**  
330 **enhance the SO<sub>2</sub> uptake on minerals at a wide pH range.** However, the  $\gamma_{\text{Light}}/\gamma_{\text{Dark}}$  had smaller  
331 values with an increase in the basicity, suggesting that the promotion effect of the light was less

332 remarkable for basic oxides.



333

334 **Figure 4.** (A) The dependence of  $\gamma_{s, \text{BET}}$  under irradiation on the basicity (pH) of mineral

335 oxides. (B) The ratios of steady-state uptake coefficients under irradiation to those in the dark

336 ( $\gamma_{s, \text{Light}}/\gamma_{s, \text{Dark}}$ ). Reaction conditions: mineral oxides mass of 0.2 g, irradiation intensity of

337  $7.93 \times 10^{16}$  photons  $\text{cm}^{-2} \text{s}^{-1}$ , temperature of 298 K, RH of 40% and O<sub>2</sub> content of 20%.

338 As shown in Figure 5A and B, the band at  $1300 \text{ cm}^{-1}$  should be ascribed to the sulfate. The

339 intensity of sulfate ( $1300$  and  $1220 \text{ cm}^{-1}$ ) under irradiation was larger than those in the dark.

340 Compared to weaker peaks of sulfates ( $1200$  and  $1260 \text{ cm}^{-1}$ ) for Al<sub>2</sub>O<sub>3</sub> in the dark (Figure 5C),

341 a stronger band of bisulfates appeared at  $1220 \text{ cm}^{-1}$  under irradiation (Figure 5D). By contrast

342 to the generation of sulfates for kaolinite and Al<sub>2</sub>O<sub>3</sub>, both sulfites and sulfates formations were

343 observed for MgO and CaO (Figure 5E–H). Sulfites were dominant in the dark, as shown by

344 the peaks at  $966$  and  $1020 \text{ cm}^{-1}$  for MgO and  $943 \text{ cm}^{-1}$  for CaO, whereas the sulfate formation

345 was significantly enhanced under irradiation according to peak intensities at  $1163 \text{ cm}^{-1}$  for

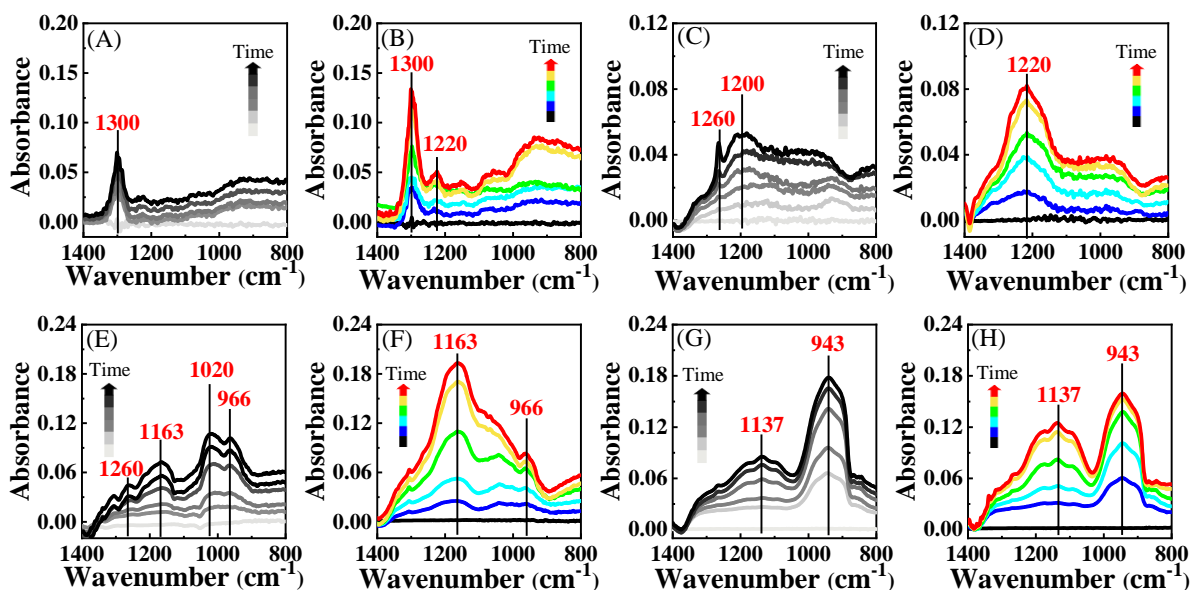
346 MgO and  $1137 \text{ cm}^{-1}$  for CaO. It should be noted that these mineral oxides were non-

347 photoactive because of their poor light absorption property (Figure S1). Nevertheless, it was

348 very surprised that the light can greatly promote the formation of sulfates via the SO<sub>2</sub> uptake

349 process on mineral oxides without photocatalytic activity, which was strongly suggested to be

350 a new and important finding for atmospheric sulfate sources.



351  
 352 **Figure 5.** *In situ* DRIFTS spectra of kaolinite (A and B), Al<sub>2</sub>O<sub>3</sub> (C and D), MgO (E and F),  
 353 CaO (G and H) during the uptake process of SO<sub>2</sub> (2 ppm) for 600 min in the dark (black  
 354 lines) and under irradiation (colorful lines). Reaction conditions: RH of 40%, temperature of  
 355 298 K and O<sub>2</sub> content of 20%.

### 356 3.5 Conversion mechanisms of SO<sub>2</sub> to sulfates

357 Heterogeneous photochemical reaction mechanisms of SO<sub>2</sub> on non-photoactive mineral dust  
 358 were proposed in light of experimental observations (Figure 6). Gaseous SO<sub>2</sub> was adsorbed on  
 359 the surface (R1), and then reacted with H<sub>2</sub>O to form sulfites (R2). Under irradiation, adsorbed  
 360 SO<sub>2</sub> accepted photons to form its singlet states (<sup>1</sup>SO<sub>2</sub>) and <sup>3</sup>SO<sub>2</sub> (R3–5) (Sidebottom et al., 1972;  
 361 Martins-Costa et al., 2018). The reaction between <sup>3</sup>SO<sub>2</sub> and H<sub>2</sub>O resulted in the formation of  
 362 HOSO• and •OH (R6), which can combine with SO<sub>2</sub> to produce HOSO<sub>2</sub>• (R7). HOSO• and  
 363 HOSO<sub>2</sub>• can be transformed into SO<sub>3</sub>, which reacted with H<sub>2</sub>O to drive the sulfate formation  
 364 (R8 and R9). The interaction between <sup>3</sup>SO<sub>2</sub> and O<sub>2</sub> may also generate SO<sub>3</sub> directly, which  
 365 would be converted to sulfates subsequently (R10). Theoretical calculations suggested that the  
 366 multistep reactions between <sup>3</sup>SO<sub>2</sub> with H<sub>2</sub>O and O<sub>2</sub> had small energy barriers or were barrier-  
 367 free (Gong et al., 2022), which could enhance the generation of ROS and the transformation of  
 368 S(IV) to S(VI). As displayed by R11 and R12, SO<sub>2</sub> and H<sub>2</sub>SO<sub>3</sub> adsorbed on the surface may be  
 369 oxidized to form sulfates via the reactions with ROS including •O, •OH or HO<sub>2</sub>•, which were

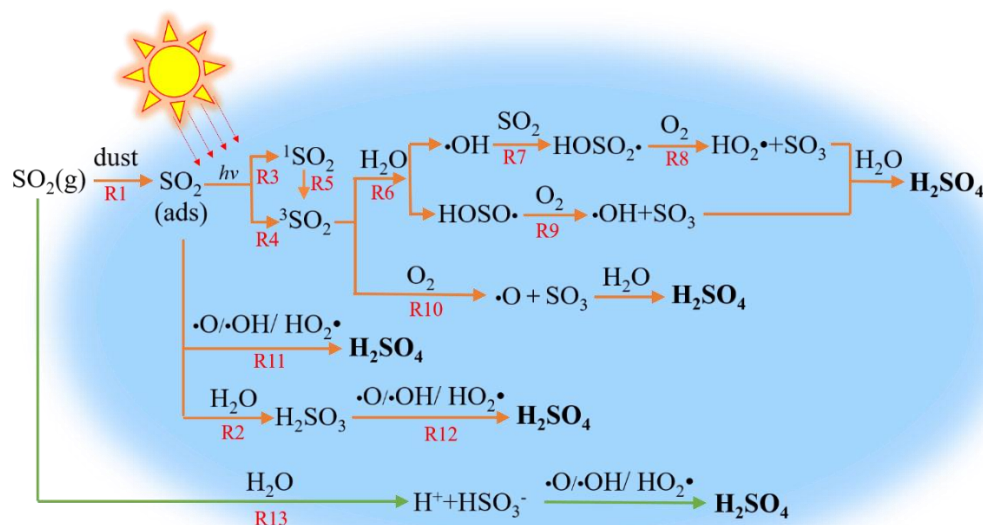
370 produced in R6 and R8–10. In addition, gaseous SO<sub>2</sub> could be dissolved into adsorbed H<sub>2</sub>O to  
371 generate bisulfites, which would be finally converted to sulfates by ROS (R13) (Urupina et al.,  
372 2019). As displayed in Figure S12A, the IR peaks of sulfates were not observed when tris (2,2'-  
373 bipyridine) ruthenium dihydrochloride (Ru(bpy)<sub>3</sub>(Cl)<sub>2</sub>) was employed as the quencher of <sup>3</sup>SO<sub>2</sub>.  
374 The peaks were assigned to the vibrations of excited Ru(bpy)<sub>3</sub>(Cl)<sub>2</sub> (Mukuta et al., 2014). This  
375 definitely proves that <sup>3</sup>SO<sub>2</sub> is the key trigger for the sulfate formation. Figure S12B shows that  
376 the peaks of sulfates were obviously weaker in the presence of NaHCO<sub>3</sub>, confirming the  
377 dominant contribution of •OH formed in R6 and R9 to the formation of sulfates.

378 Several photochemical mechanisms have been reported to explain the sulfate formation via  
379 the SO<sub>2</sub> uptake on various surfaces. Photoactive mineral oxides (such as TiO<sub>2</sub>, F<sub>2</sub>O<sub>3</sub> and ZnO)  
380 can accept photons to produce electron-hole pairs, which generated ROS for the conversion of  
381 SO<sub>2</sub> to sulfates (Ma et al., 2019; Li et al., 2019; Wang et al., 2020b). For example, •OH and  
382 HO<sub>2</sub>•, generated from the reaction of hole with H<sub>2</sub>O and electron with O<sub>2</sub>, respectively, can act  
383 as oxidizing agents for the reaction with SO<sub>2</sub> (Ma et al., 2019). Similarly, the reaction of SO<sub>2</sub>  
384 with photo-induced •OH obviously enhanced the formation of sulfate on diesel soot and actual  
385 PM<sub>2.5</sub> (Zhang et al., 2022; Zhang et al., 2020c). NO<sub>2</sub> and NO<sub>2</sub><sup>-</sup>/HNO<sub>2</sub> can be formed in the  
386 nitrates photolysis, and primarily contributed to the oxidation of SO<sub>2</sub> to sulfates on nitrates (Gen  
387 et al., 2019b; Gen et al., 2019a). Theoretically, the mechanism proposed in this study should  
388 also occur on photo-excited substrates. Taking TiO<sub>2</sub> as an example, SO<sub>2</sub> competed with TiO<sub>2</sub>  
389 for photons, and the production efficiency of <sup>3</sup>SO<sub>2</sub> and excited state of TiO<sub>2</sub> (TiO<sub>2</sub><sup>\*</sup>) depended  
390 on their light absorption properties. Meanwhile, <sup>3</sup>SO<sub>2</sub> had a competition electron-hole pairs  
391 generated from TiO<sub>2</sub><sup>\*</sup> for O<sub>2</sub> and H<sub>2</sub>O. Thus, the dominant mechanism for the SO<sub>2</sub> uptake on  
392 TiO<sub>2</sub> should be related to light absorption properties of precursors and the reactivity for <sup>3</sup>SO<sub>2</sub>  
393 and TiO<sub>2</sub><sup>\*</sup> to O<sub>2</sub> and H<sub>2</sub>O. By contrast, all mineral oxides used here cannot be excited under  
394 irradiation according to their light absorption spectra (Figure S1). Nevertheless, SO<sub>2</sub> adsorbed  
395 on mineral oxides can absorb the ultraviolet radiation (290–400 nm) to form the excited states  
396 of SO<sub>2</sub> (SO<sub>2</sub><sup>\*</sup>) (Kroll et al., 2018), which subsequently reacted with H<sub>2</sub>O and O<sub>2</sub>, finally  
397 converting SO<sub>2</sub> to sulfates. **The SO<sub>2</sub> uptake experiment in the dark and the visible light (>420**



398 nm) was carried out (Figure S13). An ignorable difference was observed for the SO<sub>2</sub>  
 399 concentration with or without visible light, suggesting that visible light had a minor contribution  
 400 to the photoenhanced SO<sub>2</sub> uptake.

401 According to the experimental results, some surfaces, providing absorptive sites for SO<sub>2</sub>, can  
 402 enhance the photooxidation of SO<sub>2</sub> to sulfates. However, the promotion effect would vary with  
 403 different substances. For example, the current experiments on some basic minerals indicate that  
 404 light plays a minor enhancement role in the SO<sub>2</sub> uptake (Figure 4), but it could still enhance the  
 405 sulfate formation (Figure 5). The solubility and effective Henry's law constant of SO<sub>2</sub> were  
 406 positively dependent on pH. Thus, SO<sub>2</sub> was more liable to be dissolved to form HSO<sub>3</sub><sup>-</sup>/SO<sub>3</sub><sup>2-</sup>  
 407 on more alkaline surface, leading to a strong SO<sub>2</sub> uptake in the dark (Figure 4A and 4B), and  
 408 abundant sulfites on surfaces (Figure 5). Nevertheless, gaseous SO<sub>2</sub> tends to be adsorbed on  
 409 kaolinite and Al<sub>2</sub>O<sub>3</sub> due to less solubility of SO<sub>2</sub> on these surfaces, and then converted to sulfate  
 410 under irradiation (Figure 6). Accordingly, a strong promotion effect of light on SO<sub>2</sub> uptake was  
 411 observed on neutral and weakly alkaline surfaces (Figure 4B).



412  
 413 **Figure 6.** The proposed photochemical conversion mechanisms of SO<sub>2</sub> to sulfates on non-  
 414 photoactive mineral dust.

#### 415 4 Atmospheric implications

416 The lifetime ( $\tau$ ) for photochemical loss of SO<sub>2</sub> on mineral dust was given using the equation

417 7,

$$418 \quad \tau = \frac{4}{\gamma\omega A} \quad (7)$$

419 where  $\gamma$  and  $\omega$  are the uptake coefficient and the mean molecular speed of  $\text{SO}_2$ , respectively;  $A$   
420 is the surface area density of mineral dust, and it is estimated to be  $(1.4\text{--}4.8) \times 10^{-5} \text{ cm}^2 \text{ cm}^{-3}$   
421 (Zhang et al., 2019; He et al., 2018b). In this work,  $\gamma_{s, \text{BET}}$  of  $\text{SO}_2$  on several mineral oxides  
422 were measured to be from  $4.39 \times 10^{-7}$  to  $3.45 \times 10^{-5}$  under conditions with  $\text{SO}_2$  concentration  
423 of 40 ppb, irradiation intensity of  $7.93 \times 10^{16} \text{ photons cm}^{-2} \text{ s}^{-1}$  and RH of 40%. Thus, the  $\tau$  of  
424  $\text{SO}_2$  with respect to the photooxidation on mineral dust was calculated to be 0.9–240 days,  
425 which was shorter than that (54 years) for the photochemical uptake of  $\text{SO}_2$  on  $\text{TiO}_2$  and the  
426 corresponding one (346 days) for the heterogeneous oxidation of  $\text{SO}_2$  on ATD in the presence  
427 of nitrates (Ma et al., 2019; Zhang et al., 2019). **The reaction conditions in this study and those**  
428 **in the literatures are different in some respects, and the previously reported  $\text{SO}_2$  uptake**  
429 **coefficient ( $10^{-7}$ – $10^{-6}$ ) had a lower value (Ma et al., 2019). The huge difference in the  $\tau$  of  $\text{SO}_2$**   
430 **was also ascribed to the variation in the surface area density. The content of  $\text{TiO}_2$  in mineral**  
431 **dust was only about 1%, and thus the surface area density of  $\text{TiO}_2$  was about  $10^{-7} \text{ cm}^2 \text{ cm}^{-3}$ ,**  
432 **leading to a longer  $\tau$  (54 years) for  $\text{SO}_2$  on  $\text{TiO}_2$  (Ma et al., 2019).** It was comparable to the  
433 lifetime (3.6–20 days) of  $\text{SO}_2$  for the gas-phase reaction with  $\bullet\text{OH}$  at a concentration of  $\sim 10^{-6}$   
434 molecules  $\text{cm}^{-3}$  (Huang et al., 2015; Zhang et al., 2019). Therefore, the photochemical process  
435 with the excited state  $\text{SO}_2$  acting as a driver on mineral dust was an important pathway for the  
436  $\text{SO}_2$  sink in the atmosphere.

437 Sulfates show significant influences on the atmosphere, such as an important contributor to  
438 the haze formation, affecting the activity of aerosols acting as cloud condensation nuclei (CCN)  
439 and ice nuclei (IN), and modifying optical property and acidity of aerosols. A sulfate formation  
440 rate ( $R$ ) can be obtained using  $\gamma$  by the equation 8 (Cheng et al., 2016),

$$441 \quad R = \frac{d[\text{SO}_4^{2-}]}{dt} = \left[ \frac{R_p}{D} + \frac{4}{\gamma\omega} \right]^{-1} A[\text{SO}_2] \quad (8)$$

442 where  $R_p$  is the radius of mineral dust, which can be estimated using the equation 9 (Li et al.,

443 2020),

$$444 R_p = (0.254 \times [\text{PM}_{2.5}]/(\mu\text{g m}^{-3}) + 10.259) \times 10^{-9} \text{ m} \quad (9)$$

445 where  $[\text{PM}_{2.5}]$  was average  $\text{PM}_{2.5}$  mass concentration, and  $300 \mu\text{g m}^{-3}$  was used for the polluted  
446 periods in typical China cities (Li et al., 2020; Guo et al., 2014). It was assumed that mineral  
447 dust accounted for 50% mass of  $\text{PM}_{2.5}$  (Tohidi et al., 2022), and the mass fraction of  $\text{SiO}_2$ ,  
448  $\text{Al}_2\text{O}_3$ ,  $\text{MgO}$ , and  $\text{CaO}$  in mineral dust was 60%, 12.5%, 4% and 6.5%, respectively (Urupina  
449 et al., 2021; Urupina et al., 2019; Usher et al., 2003). Thus,  $R$  was determined to be  $2.15 \mu\text{g}$   
450  $\text{m}^{-3} \text{h}^{-1}$ . **This suggests that the  $\text{SO}_2$  on non-photoactive surfaces is a newly identified sulfate**  
451 **formation pathway in some dust-rich conditions.**

452

### 453 **Author contributions**

454 CH, WY and JM designed and conducted experiments; CH, WY and JM analyzed the data and  
455 prepared the paper with contributions from HY; FL conducted experiments; CH supervised the  
456 project.

457

### 458 **Competing interests**

459 The authors declare that they have no conflict of interest.

460

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466

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