Understanding Changes in Cloud Simulations from E3SM Version 1 to Version 2

Yuying Zhang[1], Shaocheng Xie[1], Yi Qin[2], Wuyin Lin[3], Jean-Christophe Golaz[1], Xue Zheng[1], Po-Lun Ma[2], Yun Qian[2], Qi Tang[1], Christopher R. Terai[1], and Meng Zhang[1]

[1]Lawrence Livermore National Laboratory, Livermore, CA, USA
[2]Pacific Northwest National Laboratory, Richland, WA, USA
[3]Brookhaven National Laboratory, Upton, NY, USA

Correspondence to: Yuying Zhang (zhang24@llnl.gov)

Abstract. This study documents clouds simulated by the Energy Exascale Earth System Model (E3SM) version 2 (E3SMv2) and attempts to understand what causes the model behavior change in clouds relative to E3SMv1. This is done by analyzing the last 30-year (1985-2014) data from the 165-year historical simulations using E3SMv1 and v2 and four sensitivity tests to isolate the impact of changes in model parameter choices in its turbulence, shallow convection, and cloud macrophysics parameterization (CLUBB), microphysical parameterization (MG2), and deep convection scheme (ZM), as well as model physics changes in convective triggering. It is shown that E3SMv2 significantly improves the simulation of subtropical coastal stratocumulus clouds (Sc) and clouds with optical depth larger than 3.6 over the stratocumulus to cumulus transition regimes, where Shortwave Cloud Radiative Effect (SWCRE) is also improved, and the Southern Ocean (SO) while seeing an overall slight degradation in low clouds over other tropical and subtropical oceans. The better performance in E3SMv1 over those regions is partially due to error compensation between its simulated optically thin and intermediate low clouds for which E3SMv2 actually improves simulation of optically intermediate low clouds. Sensitivity tests indicate that the changes in low clouds are primarily due to the tuning made in CLUBB. The impact of the ZM tuning is mainly on optically intermediate and thick high clouds, contributing to improved SWCRE and Longwave Cloud Radiative Effect (LWCRE). The impact of the MG2 tuning and the new convective trigger is primarily on the high latitudes and the Southern Ocean (SO). They have a relatively smaller impact on clouds than do the CLUBB and ZM tunings. This study offers additional insights about clouds simulated in E3SMv2 by utilizing multiple data sets and the COSP diagnostic tool as well as through sensitivity tests. The improved understanding will benefit the future E3SM developments.

1 Introduction

Given the importance of clouds in global radiative balance and hydrological cycle, continuously improving the representation of clouds has been a key focus for global climate model (GCM) development. For instance, major efforts have been devoted to reducing outstanding errors in clouds simulated in the U.S. Department of Energy (DOE)'s Energy Exascale Earth System Model (E3SM) version 1 (E3SMv1) during development of its version 2 (E3SMv2) (Golaz et al. 2022). These efforts include

both improvements in representing its atmospheric physics and significant re-turning of atmospheric parameters in its deep convection, microphysics, and turbulence and macrophysics. As a result, the lack of stratocumulus (Sc) along the subtropical coasts, one outstanding error shown in E3SMv1 (Golaz et al. 2019, Rasch et al. 2019, Xie et al. 2018, Zhang et al. 2019), has been significantly reduced in E3SMv2, along with improvements in the stratocumulus to cumulus transition regimes. The updated physics and tuning parameters in E3SMv2 have also resulted in changes to other types of clouds over different cloud regimes such as the mixed-phase clouds in high latitudes (Zhang et al. 2022) and tropical high clouds (Golaz et al. 2022). Understanding model cloud behavior changes from E3SMv1 to E3SMv2 will provide necessary insights into clouds simulated by E3SMv2 and guide future E3SM development.

In this study, we perform a comprehensive evaluation of clouds simulated in E3SMv2. Our goal is to document the overall features in clouds simulated from this newly released model and understand what processes are primarily responsible for the changes from E3SMv1 to E3SMv2. To achieve our goal, we conduct and investigate a series of sensitivity tests using its atmospheric model (EAM) to isolate the impact of changes made in atmospheric physics and parameter choices in E3SMv2. This study provides a more comprehensive picture on E3SMv2 performance on cloud simulations, which was beyond the scope of the E3SMv2 overview paper by Golaz et al. (2022).

The cloud evaluation presented in this study is performed using the community satellite simulator package - the Cloud Feedback Model Intercomparison Project (CFMIP) Observation Simulator Package (COSP; Bodas-Salcedo et al. 2011, Swales et al. 2018) to improve the consistency between model clouds and satellite observations. COSP contains several independent satellite simulators for better comparing model clouds with satellite measurements collected by the International Satellite Cloud Climatology Project (ISCCP; Rossow & Schiffer, 1999), the Moderate Resolution Imaging Spectroradiometer (MODIS), the Multiangle Imaging Spectro-Radiometer (MISR), Cloud-Aerosol Lidar and Infrared Pathfinder Satellite Observation (CALIPSO), and CloudSat. The use of satellite simulators will not only make a fairer comparison between model clouds and satellite data, but also allow a more in-depth analysis of clouds. For example, clouds can be assessed in terms of their optical properties and vertical location, which dictate their radiative effects.

The manuscript is organized as follows. Section 2 provides a brief description of E3SMv2 with emphasis on changes in its atmospheric physical parameterizations and parameters settings compared to E3SMv1. In addition, satellite data and sensitivity tests analyzed in the study are also discussed. Section 3 documents the overall features in clouds simulated in E3SMv2. The processes that primarily attribute to the changes in simulated clouds from E3SMv1 to E3SMv2 are discussed in Section 4. A summary is provided in Section 5.

2 Model, Data, and Sensitivity Tests

2.1 E3SMv2

E3SMv2 is the version 2 of E3SM that was released in 2021 for public use. Compared to its precedent version (E3SMv1), E3SMv2 shows improved computational efficiency (approximately twice as fast) and simulated climate specifically related to

clouds and precipitation (Golaz et al., 2022). The improved computational efficiency mainly results from the implementation of high-order, property-preserving semi-Lagrangian tracer transport (Bradley et al., 2021) and high-order, property-preserving dynamics-physics-grid remap ("Physgrid") (Hannah et al., 2021). The improved simulation of clouds and precipitation is mainly attributed to the updates of atmospheric physics and parameter retuning. Here we emphasize the changes made in cloud and convection parameterizations in E3SMv2 since they are relevant to the improved simulation of clouds that are discussed in this study.

E3SMv2 uses the same set of atmospheric physics as E3SMv1 as described by Rasch et al. (2019) and Xie et al. (2018). Its cloud and convection parameterizations include a third-order turbulence closure parameterization (CLUBB; Cloud Layers Unified By Binormals) (Golaz et al., 2002; Larson, 2017; Larson & Golaz, 2005) for representing processes related to planetary boundary layer turbulence, shallow convection, and cloud macrophysics. The deep convection scheme is based on Zhang & McFarlane (1995) (ZM hereafter) with a dilute convective available potential energy (CAPE) modification by Neale et al. (2008). An updated version of the Morrison and Gettelman (2008) scheme (MG2, Gettelman et al., 2015) is used for representing cloud microphysics of stratiform and shallow convective clouds. The MG2 is combined with a classical nucleation theory-based ice nucleation (IN) parameterization for the heterogeneous ice formation in mixed phase clouds (Wang et al. 2014).

A notable update related to clouds and precipitation is the use of a new convective trigger function described by Xie et al. (2019) and Wang et al. (2020) in ZM to improve the simulation of precipitation and its diurnal cycle. The new convective trigger named as dCAPE-ULL uses the dynamic CAPE (dCAPE) trigger developed by Xie and Zhang (2000) with an unrestricted air parcel launch level (ULL) approach used by Wang et al. (2015). It was designed to address the unrealistically strong coupling of convection to the surface heating in ZM that often results in unrealistically too active model convection during the day in summer season over lands and improve the model capability to capture mid-level convection for nocturnal precipitation. Other updates include the implementation of a convective gustiness scheme following Redelsperger et al. (2000) to account for subgrid-scale surface wind gustiness and improve the representation of tropical clouds and precipitation (Harrop et al., 2018; Ma et al., 2022).

Significant model retunings have also been made to reduce errors in clouds and precipitation during the E3SMv2 development. Following Ma et al. (2022), several tuning parameters are recalibrated in CLUBB, ZM deep convection, and microphysics schemes to improve the simulation of cloud and precipitation. Information learned from the short Parameter Perturbation Ensemble simulations (Qian et al. 2018) is also used to guide EAMv2 model retuning effort. See Table A1 in Golaz et al. (2022) for details about the tuning parameters used in E3SMv2 model and the difference from E3SMv1.

2.2 Satellite data

Satellite measurements from ISCCP, MODIS, MISR, and CALIPSO are used to evaluate clouds simulated by E3SMv2. CERES-EBAF Edition 4.1 data are used to evaluate model simulated TOA shortwave cloud radiative effect (SWCRE) and longwave cloud radiative effect (LWCRE). More detailed information and references about these data are given in Table 1.

Zhang et al. (2019) discussed additional details about these measurements. Here we highlight some of the key points from Zhang et al. (2019) to facilitate interpreting the results of this study.

Measurements from the ISCCP, MODIS and MISR are from passive instruments. They contain the information about the area coverage of clouds stratified by cloud-top pressure (*ctp*) or cloud-top height (*cth*) of the highest cloud in a column and by the column integrated cloud optical thickness (*Tau*), which can be summarized in joint histograms of *ctp-tau* or *cth-tau*. As discussed in earlier studies (Marchand et al., 2010; Pincus et al., 2012), notable differences are found in the joint histograms among the three data sets, which are largely due to instrument limitations and different algorithms used for detecting clouds and retrieving the cloud heights and optical depths. For example, the ISCCP often has difficulties to detect small cumulus clouds and mistakenly put optically thin cirrus as a mid-topped cloud when there are low clouds underneath (Mace et al. 2006). In contrast, the retrievals used in MODIS do not work well for low-level clouds under temperature inversions or broken low-level clouds (Pincus et al. 2012) although MODIS cloud data are usually considered as the most accurate for high-topped cloud among the three passive instruments. In addition, partly cloudy pixels are excluded in MODIS while they are treated as homogeneous and included in ISCCP detected cloud. Pincus et al. (2012) found that this could lead to 15% difference in optically thinnest clouds estimated by ISCCP and MODIS. Compared to ISCCP and MODIS, MISR gives the most accurate estimate of cloud top height for low-level clouds and better detection of cumulus clouds. Different from ISCCP, MODIS, and MISR, CALIPSO uses active instruments to measure cloud height directly and therefore can provide information of cloud vertical structure. The CALIPSO data used in this study include high, middle, and low cloud fraction derived from the attenuated backscattered profile at 532 nm (Chepfer et al., 2010). Like Zhang et al. (2019), this study utilizes all these available observations to provide more complete information on model simulated clouds.

2.3 Sensitivity Tests

Table 2 lists the simulations analyzed in this study. In addition to the last 30-year (1985-2014) data from the 165-year historical simulations with E3SMv1 and E3SMv2 (Golaz et al. 2019 and 2022) used to document the overall features of clouds simulated in these models, we also analyze the sensitivity tests with the atmosphere model of E3SMv2 (EAMv2) to investigate the impact of each major change on the simulated clouds. Four major changes made in E3SMv2 were tested in Qin et al. (2023) and analyzed in this study. They include the new dCAPE_ULL convective trigger (Trigonly) and the tunings made in CLUBB (clubbonly), MG2 (MGonly), and ZM (ZMonly). As demonstrated in Xie et al. (2019), the dCAPE_ULL trigger effectively suppresses daytime suspicious deep convection particularly over land and captures elevated nocturnal convection above the boundary layer. It also considerably reduces convective precipitation over subtropical regions and the occurrence frequency of light-to-moderate precipitation. The changes in model convection can have large impact on the simulated clouds given the strong connection between clouds and convection. The tunings made in CLUBB are mainly to improve both stratocumulus and cumulus clouds and the transitions from stratocumulus to cumulus (Ma et al. 2022). As described in Table A1 of Golaz et al. (2022), the goal is achieved by 1) weakening the turbulence mixing in the planetary boundary layer (PBL) through adjusting a set of relevant parameters (e.g., clubb_c1, c14, c6, etc.), 2) promoting cloud formation by reducing the width of the w′

Probability Distribution Function (PDF) via reducing gamma_coef and gamma_coefb, 3) increasing stratocumulus clouds and
130 reducing cumulus clouds by reducing skewness (skw) via increasing C8, and 4) allowing larger horizontal variation in subgrid characteristics by enlarging the difference in parameter values between high- and low-skewness regimes. The major recalibrations of cloud microphysics (MG2) include re-setting the overly suppressed tuneable scaling factor (0.1) for the Wegener-Bergeron-Findeisen (WBF) process in EAMv1 to 0.7 to improve the simulation of mixed-phase clouds, adding a minimum cloud droplet number concentration (CDNC) to reduce aerosol forcing, and adjusting the exponent coefficient and
135 the liquid cloud accretion enhancement factor to better represent clouds and precipitation in subtropical low cloud regimes. In the deep convection scheme (ZM), the parcel buoyancy considers the subgrid temperature perturbation from the CLUBB scheme in addition to a constant value of 0.8 K used in EAMv1. A new tunable parameter with a default value of 2.0, zmconv_tp_fac (see Table A1), is introduced to scale the square root of the CLUBB subgrid temperature variance to be the subgrid temperature perturbation. Additionally, the parameters related to the autoconversion rate, detrained ice cloud effective
140 radius, and cloud fraction in deep convective clouds are reduced, while the parameters related to the downdraft mass flux fraction and the impact of the surface temperature change are enhanced compared to EAMv1. More details about what parameters were tuned in v2 are given in the Table A1 of Golaz et al. (2022). All the sensitivity tests are 6-year AMIP-type runs with present-day (2010) forcing from the Intergovernmental Panel on Climate Change (IPCC) AR5 emission data set (Lamarque et al., 2010) along with climatological sea surface temperature and sea ice prescribed from the observations
145 (repeating seasonal cycle without interannual variability). To accurately measure their impacts, the same 6-year AMIP runs were conducted for the default E3SMv2. The last 5 years from each run are analyzed. By comparing the AMIP runs with the 30-year coupled simulations, the short-term AMIP runs can well reproduce the major model errors shown in the long-term coupled simulations (not shown). This suggests that the E3SM cloud behaviors should be mostly controlled by its atmosphere model and most systematic errors in clouds are apparent in a few year simulations since clouds are associated with fast physics
150 (e.g., Xie et al. 2012, Ma et al. 2014, and Qian et al., 2018).

3 General Cloud Features Simulated in E3SMv1 and E3SMv2

Results shown in this section are based on the last 30-year (1985-2014) data from the 165-year historical simulations, which are part of the E3SMv1 and E3SMv2 CMIP6 DECK plus historical simulation campaigns (Eyring et al., 2016). The exception is for MODIS simulator output. A bug was found in the MODIS simulator output shortly after the 165-year's historical
155 simulations were completed. Therefore, the results from MODIS analyzed here are from the default E3SMv1 and E3SMv2 simulations conducted in the sensitivity tests, that is, the 6-year AMIP runs. As we will discuss later, the 6-year AMIP runs can well represent the general features in cloud simulations as we see in the 30-year coupled runs. So, we do not expect this issue will affect what we will learn from this study. In this section, we will emphasize the overall features in clouds simulated in E3SMv2 by utilizing COSP. We focus on annual mean climatology of clouds simulated by both models and changes from
160 E3SMv1 to E3SMv2.

3.1 Evaluation of model cloud with COSP

3.1.1 Total cloud fraction

The annual mean total cloud fraction between the models and the ISCCP, MODIS, and CALIPSO observations is shown in Figure 1. Note that clouds with *Tau* < 1.3 in ISCCP and MODIS are neglected due to the large uncertainty of cloud detection from passive instruments as discussed earlier (Marchand et al. 2010; Pincus et al. 2012). The impact of the exclusion of the optically thin clouds in ISCCP and MODIS was discussed by Zhang et al. (2019). In general, the omission of the optically thin clouds leads to a noticeable reduction of total cloud fraction between 60S and 60N, however, the exclusion of the optically thin clouds in MODIS has minor effect on its total cloud fraction (not shown). Beyond 60S and 60N, measurements from the passive instruments are less reliable due to their difficulties in detecting clouds over surfaces with ice and snow (Klein et al. 2013; Zhang et al. 2019) and therefore caution needs to be taken for model results discussed in these regions.

Overall, both models produce comparable results with slightly fewer clouds simulated in E3SMv2 compared to its previous version. E3SMv2 generally shows slightly larger errors than E3SMv1 compared to ISCCP and CALIPSO. Relative to ISCCP, E3SMv2 underpredicts clouds in the tropical and extratropical regions and has more clouds over the Arctic than observations. As we will show later, this is mainly related to its simulated low clouds. E3SMv1 shows a similar error pattern in most regions except for the Pacific Ocean where clouds are generally overestimated. Relative to CALIPSO, both models underpredict clouds globally except for the Arctic. Compared to MODIS, however, E3SMv2 shows a slightly better result than E3SMv1 with reduced biases over most regions. The discrepancy in model performance against different satellite data is likely not due to the use of 6-year AMIP simulations for MODIS since similar discrepancy is seen when 6-year simulation data are used for ISCCP and CALIPSO (not shown).

A robust improvement made in E3SMv2 is the considerable increase of stratocumulus cloud over the eastern ocean basins along the coasts, such as the west coasts of South Africa and North and South America. The comparison with all three satellite observations shows this improvement. This is significant since the lack of stratocumulus along the coasts in E3SMv1 (common in most climate models) is the one of the main outstanding issues that the development of E3SMv2 tackled. Cloud biases over N. Hemisphere storm tracks, the stratocumulus to cumulus transition regimes, and the Southern Ocean (SO) are also noticeably reduced in E3SMv2 as seen in the comparisons particularly with ISCCP and MODIS. It is worth mentioning that the marine stratocumulus cloud biases are substantially reduced in the E3SMv2 North American regionally refined configuration (Tang et al., 2023), i.e., the high-resolution version of the E3SMv2 release. This is due to the finer resolutions in both atmosphere and ocean and highlights the benefits of refining resolutions in multiple components than in a single component.

The model simulated clouds are further examined by analyzing the column-integrated cloud optical depth distributions for total cloud fraction averaged in the domain between 60S and 60N and over the SO region between 45S and 60S, respectively (Fig. 2). The domain between 60S and 60N is selected because measurements from ISCCP and MODIS are more reliable while the SO region is selected since most climate models have difficulties in accurately capturing clouds over this region where E3SMv2 shows clearly improvements in cloud simulations compared to E3SMv1. ISCCP and MODIS agree relatively better

for clouds with *Tau* > 3.6, while for clouds with *Tau* < 3.6, the total cloud fraction in ISCCP is significantly larger than that in MODIS, primarily due to the different assumptions on partly cloudy pixels in their retrievals (Zhang et al. 2019). For 3.6 < *Tau* < 23, MODIS shows larger cloud fraction than ISCCP. In contrast, model results from the two simulators are very close to each other. This is because ISCCP and MODIS simulators use nearly the same method to determine the values of *Tau* (Pincus et al. 2012).

Both models agree with MODIS much better than with ISCCP in both examined regions. Relative to MODIS, E3SMv1 generally well reproduces optically thin clouds (1.3 < *Tau* < 3.6) and notably overestimates optically intermediate to thick clouds (*Tau* >= 9.4). In contrast, E3SMv2 underpredicts the optically thin clouds and shows slightly larger overestimation of thick clouds (*Tau* >23) compared to E3SMv1. The largest discrepancy between E3SMv1 and v2 is seen for *Tau* between 3.6 and 9.4 where E3SMv2 performs worse than v1 and considerably underestimates the MODIS observed clouds. This suggests that some features seen in the total cloud fraction in Figure 1 are a result from compensating errors in clouds with different optical properties. For example, the reduction of positive cloud fraction bias over the SO in E3SMv2 relative to MODIS, as shown in Figure 1, appears to stem more from the deficiencies in optically thin and intermediate clouds than the slightly more excessive optically thick clouds with respect to E3SMv1.

3.1.2 Cloud vertical structure

The annual mean high, middle, low cloud fraction from CALIPSO and model biases from the observation are shown in Figure 3. The high, middle, and low clouds in both models and CALIPSO are defined as cloud top pressure lower than 440mb, between 440mb and 680mb, and higher than 680mb, respectively. E3SMv1 and E3SMv2 show similar bias patterns for all types of clouds with clouds at all heights mostly underpredicted in the low latitudes and low clouds overpredicted in the northern high latitudes and the SO. Over the Tropics, E3SMv2 slightly improves high and middle clouds with noticeable reduction of errors in South Pacific (high clouds) and tropical and subtropical Pacific (middle clouds), when compared with E3SMv1. On the other hand, it degrades the representation of low clouds over tropical and subtropical oceans except along the subtropical coasts in both Hemispheres where E3SMv2 produces much more stratocumulus (better). Over the SO, the overpredicted clouds in E3SMv1 are reduced in E3SMv2. The degradation of low clouds over tropical and subtropical oceans is partially due to error compensation in low clouds with different optical properties in E3SMv1 as will be shown in a comparison with the MISR low cloud data, which give a more accurate estimate of low clouds over oceans compared to CALIPSO (Zhang et al. 2019). Also over the SO E3SMv2 shows larger overestimation of middle-level clouds and smaller overestimation of low clouds. This breakdown of biases in terms of clouds vertical structure from CALIPSO provides additional information to understand the overall model performance in simulating total cloud fraction. For instance, the degradation of total cloud fraction simulation in E3SMv2 over the tropical and subtropical regions as shown in Figure 1i is primarily due to errors in low clouds since middle and high clouds are generally improved over these regions (Figures 3c, 3f, and 3i).

3.1.3 High clouds

The simulated high clouds in three optical thickness ranges (thin, intermediate, and thick) are further examined with MODIS (Fig. 4) since MODIS provides more accurate information about high clouds than other datasets. Despite some regional differences, E3SMv2 shows a very similar error pattern to E3SMv1. In general, both models overestimate the MODIS cloud fraction (particularly over land) for optically thin clouds and underestimate it (mainly over ocean) for optically intermediate and thick clouds. E3SMv2 shows slightly larger error for optically intermediate clouds and slightly smaller error for optically thick clouds. There are some improvements seen in E3SMv2 along the Antarctic coasts for optically intermediate clouds and in the Arctic for optically thick clouds where E3SMv1 slightly overestimates the observed clouds.

3.1.4 Low clouds

As MISR can detect low clouds well, it is used to provide additional insights into model deficiencies in simulating low clouds. The optically thin, intermediate, and thick low cloud fractions (cth below 3 km) from MISR (only available over oceans) and the difference between the models and MISR are displayed in Figure 5. For optically thin low clouds ($0.3 < Tau < 3.6$), there is little change in model errors from E3SMv1 to E3SMv2. Both models largely underestimate the optically thin low clouds in both tropical and midlatitude oceans. For optically intermediate low clouds ($3.6 < Tau < 23$), E3SMv2 dramatically improve the simulation. The large overestimation over the N.H. storm tracks, the SO, and the stratocumulus to cumulus transition regimes and underestimation in the Sc regions in E3SMv1 are significantly reduced in E3SMv2, consistent with the earlier discussions and Golaz et al. (2022). However, the smaller positive biases in optically intermediate low clouds in E3SMv2 leads to a bigger negative bias in total low-cloud fraction than E3SMv1 as shown in Fig. 3 (compared to CALIPSO low clouds). The better performance shown in E3SMv1 is clearly due to error compensation between its simulated optically thin (underestimated) and intermediate (overestimate) low clouds. The optically thick low clouds are not the dominant cloud type. Only a few optically thick low clouds are detected by MISR, for which the models show a slightly positive bias in the N. Hemisphere storm tracks and the SO, especially for E3SMv2.

3.2 Cloud radiative effect

Clouds have a large impact on radiation. The biases in model simulated shortwave and longwave cloud radiative effect (SWCRE and LWCRE) from the CERES-EBAF Ed4.1 dataset (Fig. 6) are closely related to those in model clouds. SWCRE is largely influenced by low clouds. The most noticeable improvement from E3SMv1 to E3SMv2 is over the stratocumulus regimes along the west coast of continents where the severely underestimated SWCRF in E3SMv1 is significantly reduced due to the improvement of Sc as discussed earlier. In contrast, the biases in LWCRE are more related to high clouds. The major error in E3SMv1 is the much weaker LWCRE due to large underestimation of (optically intermediate) high clouds over the tropical Indo-West Pacific region (Figure 4e and 4f). This problem is reduced over the West Pacific region in E3SMv2 while slightly exaggerated over eastern Indian Ocean. For other regions, the change is little from E3SMv1 to E3SMv2.

4 Impact of major changes in v2 on cloud simulations

The previous section indicates that E3SMv2 largely improves the stratocumulus clouds over the eastern ocean basins in both hemispheres along the coast of the continents of South Africa and North and South America along with the improvement seen in the stratocumulus to cumulus transition regimes, and the Southern Ocean (SO) while it degrades the simulations of high clouds and total cloud fraction compared to E3SMv1. The improvements in low clouds are mainly from optically intermediate clouds with *Tau* in a range of 3.6 and 23. For optically thin (*Tau* < 0.3) and thick (*Tau* > 23) low clouds, E3SMv2 actually produces slightly worse result. The degradation of high clouds is primarily from the intermediate clouds.

In this section, we discuss what changes made in E3SMv2 have led to these changes in clouds from E3SMv1 to E3SMv2 through carefully designed sensitivity experiments, which are shown in Table 2 and discussed in Section 2.3. Since model responses to these changes are similar regardless of which simulator output is used, for simplicity, we use MODIS for total cloud fraction and CALIPSO for vertical cloud structure. Results from both MODIS and MISR simulators are also examined and discussed in this section for high and low clouds, respectively, given their relatively more accurate measurements in these two types of clouds. The inclusion of MODIS and MISR clouds are particularly meaningful to aid in the breakdown of the model responses in terms of cloud opacity. The relevant figures for MODIS high clouds and MISR low clouds are included in Appendix.

4.1 MODIS total cloud fraction

Figure 7 shows the impact of these changes on the simulation of total clouds. The improved stratocumulus along the west coast of continents (i.e., South Africa, North and South America) is clearly due to the retuning of CLUBB, which promote stratocumulus-like symmetric mixing by increasing the damping coefficients and allowing larger horizontal variation in subgrid vertical velocity as described in Ma et al. (2022). However, the retuning of CLUBB also leads to a considerable reduction of clouds over tropical and subtropical oceans, which contributes to the fewer clouds produced in E3SMv2 compared to v1. This is mainly from tuning those parameters that can decrease boundary layer mixing and decoupling between boundary layer and free troposphere. The cloud lateral entrainment is also decreased in v2 due to the CLUBB-retuning, which could lead to a reduced cloudiness in shallow cumulus regime. As indicated by CALIPSO (Figures 9-11) and MISR (Figures A1 – A3), the CLUBB-induced cloud changes are mainly in the optically intermediate low clouds. The reduction of total clouds is partially offset by the tuning made in ZM, which extends clouds originated from deep convections to almost everywhere. This is largely due to the tuning of the autoconversion for convective clouds, which is significantly tuned down from 0.007 in E3SMv1 to a nominal value of 0.002 in E3SMv2. This increases cloud condensate amount detrained from deep convections and increases overall cloudiness. An analysis of MODIS high clouds (Figures A4 – A6) supports that the major impact on high clouds is from the ZM tuning, which largely increases the optically intermediate and thick clouds. The impact of MG2-tuning and the new convective trigger on total cloud fraction is relatively small compared to the tuning made in CLUBB and ZM. The MG2 tuning mainly influences high latitudes, particularly along the Antarctic coastline while the new trigger produces

290   more clouds over most of the regions except for the tropical western Pacific and the SO between 30°S and 60°S where a reduction of clouds is noticeable. The large impact of MG2-tuning on high latitude clouds is likely due to the change in the WBF process rate, which has a large impact on mixed-phase clouds that are common in high latitudes.

The impact on total cloud fraction with different cloud optical properties is further examined in Figure 8. Over 60S – 60N, consistent with Figure 7, the CLUBB tuning has led to a reduction of clouds regardless of their optical properties, which results

295   in a better simulation of optically thick clouds ($Tau > 9.4$) and a worse simulation of optically thin clouds ($Tau <= 9.4$). The ZM tuning response is to increase clouds with overall minor impact on optically thin clouds ($Tau < 3.6$) and considerably large impact on clouds with cloud optical depth between 3.6 and 23 that are more relevant to deep convection. The MG tuning mainly affect simulation of intermediate clouds ($3.6 < Tau < 23$) where the overestimation of clouds is reduced. The impact of the new trigger on clouds is more clearly demonstrated when examining clouds with different optical properties. It acts to

300   reduce optically thin clouds and increase optically intermediate and thick clouds. This is consistent with Xie et al. (2019), which showed that the new trigger helped suppress light precipitation and enhance intermediate and heavy precipitation.

Similar impacts of these changes are found over the SO. The tunings made in CLUBB help make optically intermediate and thick clouds ($Tau > 9.4$) closer to the observations and the MG tuning dramatically improves optically intermediate clouds ($9.4 < Tau < 23$). The new trigger acts to reduce optically thin clouds and increase optically intermediate and thick clouds like

305   over 60S to 60N. The reduction of optically thin clouds from the new trigger is mainly from low clouds as indicated in Figure A1d. The exception is for the ZM tuning, which has minor impact over the SO.

4.2 CALIPSO vertical cloud structure

Figures 9-11 display the impact of these changes on the simulation of high, middle, and low clouds, respectively. For high clouds (Figure 9), it is not surprising to see that the deep convection ZM tuning has the largest impact, which considerably

310   increases high clouds globally, especially in the tropics. The new trigger leads to a considerable reduction of high clouds over the western Pacific, which is consistent with the reduction in precipitation over this region as shown in Xie et al. (2019). The CLUBB tuning typically increases high clouds over land and reduces them over ocean. Overall, the impact of MG2 on high clouds is minor with slight increase of clouds in the Arctic and reduction of clouds along the Antarctic coastlines. By comparing with MODIS high clouds, the changes in high clouds are mainly in intermediate and thick clouds due to the changes in ZM

315   (Figures A5c and A6c). All the changes have minor impact on optically thin high clouds.

For middle clouds (Figure 10), the CLUBB tuning leads to a slight reduction of cloud amount globally. The MG2 tuning leads to a considerable reduction of clouds in both Arctic and Antarctic and a slight increase of clouds over the SO. The impact of the ZM tuning and the new trigger is mainly to increase clouds in the tropics (the ZM tuning) and the SO and the Arctic regions (the new trigger).

320   For low clouds (Figure 11), the CLUBB tuning plays the largest role in the E3SMv2 cloud changes. It dramatically reduces the low clouds in the tropical and subtropical regions except for the stratocumulus regime where low clouds along the subtropical coasts largely increase. These changes are primarily from the optically intermediate clouds as indicated earlier.

The tuning made in ZM as well as the use of the new trigger help offset the reduction of clouds made by the CLUBB tuning over the tropical and midlatitude regions. The exception is over the SO where the new trigger acts to reduce low clouds and
325   MG2 tuning to substantially increase low clouds around the Antarctica. The reduced low clouds by the new trigger are mainly from optically thin clouds while the increased low clouds by MG2 are from optically thin and thick clouds suggested by MISR. The analysis of the vertical cloud structure clearly indicates compensation effect of cloud changes in the vertical that has impact on the total cloud fraction shown in Figure 7. For instance, the opposite sign of changes in middle and low clouds over the SO with the new trigger leads to overall small changes over that region in Figure 7.

330   **4.3 SWCRE and LWCRE**

The impacts of these model changes on SWCRE and LWCRE are shown in Figures 12 and 13, respectively. For SWCRE, consistent with the large reduction of low clouds, the CLUBB tuning largely reduces SWCRE over the tropical and subtropical oceans. In contrast, the SWCRE in the tropical deep convection regions becomes much stronger mainly due to the increase of optically intermediate and thick high clouds due to the ZM tuning. The reduction of high clouds over the Western Pacific with
335   the new trigger leads to reduction of SWCRE. Overall, the impact of the MG2 tuning on SWCRE is minor.
For LWCRE, only the ZM tuning has noticeable impact, leading to an increase in LWCRE due to the increase of high clouds. The minor impact of the CLUBB tuning in LWCRE indicates the changes in low clouds have minor impact on LWCRE.

5 Summary

We performed systematic evaluation of clouds simulated in the newly developed Energy Exascale Earth System Model
340   (E3SM) version 2 (E3SMv2) with satellite observations by utilizing the satellite simulator package (COSP) to mitigate sampling and algorithmic differences between modeled and observed clouds. Multiple cloud observations measured by various instruments were used to address potential data uncertainty and instrument and retrieval limitations. Our focus is to document E3SMv2 performance on clouds and understand what updates in E3SMv2 have caused the changes in clouds from E3SMv1 to E3SMv2. For the second purpose, results from 4 sensitivity tests conducted in Qin et al. (2023) were used to isolate the
345   impact of tuning made in CLUBB, MG2 and ZM, as well as the use of the new dCAPE_ULL trigger in E3SMv2 on clouds.
In general, E3SMv2 shows a similar error pattern in clouds as exhibited in its previous version. One robust improvement is seen in the stratocumulus regime where stratocumulus clouds along the subtropical west coast of continents in both Hemispheres are largely increased. This is true by comparing all the satellite datasets used in this study and consistent with a much stronger (better) SWCRE over these stratocumulus regions. The MISR data further indicates that the improvement is
350   mainly from optically intermediate low clouds ($3.6 < Tau < 23$) along the subtropical coasts. Note that the lack of stratocumulus clouds is one of the most outstanding problems in E3SMv1. This improvement represents a big achievement made in E3SMv2 (Golaz et al. 2022). Relative to CALIPSO, however, E3SMv2 shows a larger negative bias in total low clouds than E3SMv1 in other regions in tropical and subtropical oceans. A comparison with MISR suggests that the smaller

biases in E3SMv1 partially result from error compensation in its simulated low clouds with different optical properties, for which E3SMv1 shows an underestimation of optically thin low clouds (similar to E3SMv2) while it largely overestimates optically intermediate low clouds in these regions. Relative to the large changes in low clouds, E3SMv1 and E3SMv2 show very similar simulation of CALIPSO high clouds although some regional changes are seen in their simulated optical properties as demonstrated in comparison of MODIS.

The sensitivity tests indicate that the improved stratocumulus along the coast is primarily from the retuning of parameters related to CLUBB. Other changes only have minor contributions. However, the CLUBB tuning also resulted in a reduction of low clouds, due to the reduction of optically intermediate clouds, in the tropical and subtropical regions. This change reduced the overpredicted cumulus clouds in the subtropical cumulus regions but exaggerated the issue with underpredicted cloud in other regions in tropical and subtropical oceans. This led to an overall slight degradation of cloud simulation in E3SMv2 compared to E3SMv1 although the better performance in the latter is partially due to error compensation in clouds with different optical properties as discussed earlier.

The sensitivity tests indicate that the impact of the MG tuning on clouds is mainly in the high latitudes over both Hemispheres where it increased low clouds and decreased middle clouds. Over the SO, its overall effect is minor. However, a close look at the cloud optical properties over the SO shows a significant improvement in optically intermediate clouds compared to MODIS. Overall, its impact on clouds is smaller than the other updates examined in this study. The ZM tuning primarily increased optically intermediate and thick high clouds over the tropical deep convection regions. The increase in the high clouds partially offset the decrease of low clouds by the CLUBB tuning in the total cloud fraction and had a positive impact on simulation of both SWCRE and LWCRE in E3SMv2 over the tropical deep convection regions. Similar to the MG2 tuning, the impact of the new trigger on clouds is also smaller than the tuning made in CLUBB and ZM. The most noticeable change using the new trigger is the large increase in optically thick low clouds near the Antarctic coast and in the northern high latitudes, whereas changes of opposite sign are seen in its produced optically thin clouds, leading to a much smaller overall changes over these areas in Figure 11d.

This study offered additional insights about clouds simulated in E3SMv2 by utilizing multiple data sets and the COSP diagnostic tool and through sensitivity tests. The improved understanding will benefit future E3SM developments and application of E3SM in various science applications.

Code and data availability

The E3SM source code is available on GitHub: https://github.com/E3SM- Project/E3SM (E3SM Project, 2023) under the 3-Clause BSD Open Source license (https://e3sm.org/resources/policies/open-source-license/). The simulations of version 1 and version 2 are reproduced using maintenance branches maint-1.0 at https://github.com/E3SM-Project/E3SM/tree/maint-1.0 (last access: April 2023; Rasch et al., 2019) and maint-2.0 at https://github.com/E3SM-Project/E3SM/tree/maint-2.0 (last access: April 2023; Golaz et al., 2022), respectively. These model codes and the simulator output from the model simulations

are archived at https://doi.org/10.5281/zenodo.8021851. The simulator output is also accessible at https://portal.nersc.gov/project/e3sm/yuying/E3SMv2/ModelOutput.

The original cloud observations for model evaluation and CERES are available at https://climserv.ipsl.polytechnique.fr/cfmip-obs and https://asdc.larc.nasa.gov/project/CERES/CERES_EBAF_Edition4.1, respectively. The observational climo data are also available at https://portal.nersc.gov/project/e3sm/yuying/E3SMv2/observation.

Author contribution

YZ and SX initiated this study and led the analysis and manuscript writing. YQ performed the sensitivity tests. QT and XZ performed the E3SM v1 and v2 historical simulations. YZ coordinated the manuscript writing with input from all the co-authors.

Competing interest

Po-Lun Ma is a Topical Editor of Geoscientific Model Development. Other authors declare that they have no conflict of interest.

Acknowledgement

This research was primarily supported as part of the Energy Exascale Earth System Model (E3SM) project and partially supported by the PCMDI cloud feedback project of the Regional and Global Model Analysis (RGMA) Program, funded by the U.S. Department of Energy, Office of Science, Office of Biological and Environmental Research. The authors thank all E3SM team members for their efforts in developing and supporting the E3SM model version 2. Work at LLNL was performed under the auspices of the US DOE by Lawrence Livermore National Laboratory under contract No. DE-AC52-07NA27344. The Pacific Northwest National Laboratory (PNNL) is operated for DOE by Battelle Memorial Institute under contract DEAC06-76RLO 1830. This research used high-performance computing resources of the National Energy Research Scientific Computing Center, a DOE Office of Science User Facility supported by the Office of Science of the U.S. Department of Energy under Contract No. DE-AC02-05CH11231.

References

Bodas-Salcedo, A., Webb, M. J., Bony, S., Chepfer, H., Dufresne, J. L., Klein, S. A.: COSP: Satellite simulation software for model assessment, Bull. Amer Meteor. So., 92, 1023–1043, https://doi.org/10.1175/2011BAMS2856.1, 2011.

Bradley, A. M., Bosler, P. A., and Guba, O.: Islet: Interpolation semi-lagrangian element-based transport.Geosci. Model Dev. Discuss, 1–48. doi: 10.5194/gmd-2021-296, 2021.

CERES_EBAF_Ed4.1 Data Quality Summary. Available online: https://ceres.larc.nasa.gov/documents/DQ_summaries/ CERES_EBAF_Ed4.1_DQS.pdf.

415 Cesana, G. and H. Chepfer: Evaluation of the cloud thermodynamic phase in a climate model using CALIPSO-GOCCP, J. Geophys. Res., DOI:10.1002/jgrd.50376, 2013.

Chepfer, H., S. Bony, D. Winker, G. Cesana, J. L. Dufresne, P. Minnis, C. J. Stubenrauch, and S. Zeng: The GCM-Oriented CALIPSO Cloud Product (CALIPSO-GOCCP), J. Geophys. Res., 115, D00H16, doi:10.1029/2009JD012251, 2010.

Eyring, V., Bony, S., Meehl, G. A., Senior, C. A., Stevens, B., Stouffer, R. J., & Taylor, K. E.: Overview of the Coupled

420 Model Intercomparison Project Phase 6 (CMIP6) experimental design and organization. Geoscientific Model Development,9(5), 1937–1958. doi:10.5194/gmd-9-1937-2016, 2016.

Gettelman, A., Morrison, H., Santos, S., Bogenschutz, P., & Caldwell, P. M.: Advanced Two-Moment Bulk Microphysics for Global Models. Part II: Global Model Solutions and Aerosol–Cloud Interactions. Journal of Climate, 28(3), 1288–1307. https://doi.org/10.1175/ JCLI-D-14-00103.1, 2015.

425 Golaz, C., et al.: The DOE E3SM coupled model version 1: Overview and evaluation at standard resolution. Journal of Advances in Modeling Earth Systems,11(7), 2089-2129. doi:10. 1029/2018MS001603, 2019.

Golaz, C., et al.: The DOE E3SM Model Version 2: Overview of the physical model and initial model evaluation, Journal of Advances in Modeling Earth Systems, https://doi.org/10.1029/2022MS003156, 2022.

Golaz, J.-C., Larson, V. E., and Cotton, W. R.: A PDF-Based Model for Boundary Layer Clouds. Part I: Method and Model

430 Description. Journal of the Atmospheric Sciences, 59(24), 3540–3551, https://doi.org/10.1175/1520-0469(2002)059<3540:APBMFB>2.0.CO;2, 2002.

Hannah, W. M., Bradley, A. M., Guba, O., Tang, Q., Golaz, J.-C., and Wolfe, W.: Separating physics and dynamics grids for improved computational efficiency in spectral element Earth System Models, J. Adv. Model Earth Sy., 13(7), e2020MS002419, doi: 10.1029/2020ms002419, 2021.

435 Harrop, B. E., Ma, P.-L., Rasch, P. J., Neale, R. B., and Hannay, C.: The role of convective gustiness in reducing seasonal precipitation biases in the tropical west pacific. Journal of Advances in Modeling Earth Systems, 10(4), 961-970. doi:10.1002/2017MS001157, 2018.

Klein, S. A., Zhang, Y., Zelinka, M. D., Pincus, R., Boyle, J., & Gleckler, P. J.: Are climate model simulations of clouds improving? An evaluation using the ISCCP simulator. Journal of Geophysical Research: Atmospheres, 118, 1329–1342.

440 https://doi.org/10.1002/jgrd.50141, 2013.

Larson, V. E., and Golaz, J.-C.: Using Probability Density Functions to Derive Consistent Closure Relationships among Higher-Order Moments. Monthly Weather Review, 133(4), 1023–1042. https://doi.org/10.1175/MWR2902.1, 2005.

Larson, V. E.: CLUBB-SILHS: A parameterization of subgrid variability in the atmosphere.arXiv. doi: 10.48550/ARXIV.1711.03675, 2017.

445 Loeb, N. G., and D. R. Doelling: Clouds and the Earth's Radiant Energy System (CERES) Energy Balanced and Filled (EBAF) Top-of-Atmosphere (TOA) Edition=4.0 Data Product, J. Climate, https://doi.org/10.1175/JCLI-D-17-0208.1, 2018.

Ma, H.-Y., Xie, S., Klein, S. A., Williams, K. D., Boyle, J. S., Bony, S., Douville, H., Fermepin, S., Medeiros, B., Tyteca, S., Watanabe, M., and Williamson, D.: On the correspondence between mean forecast errors and climate errors in CMIP5 models.Journal of Climate,27(4), 1781–1798. https://doi.org/10.1175/JCLI-D-13-00474.1, 2014.

450 Ma, P.-L., Harrop, B. E., Larson, V. E., Neale, R. B., Gettelman, A., Morrison, H., et al.: Better calibration of cloud parameterizations and subgrid effects increases the fidelity of the E3SM Atmosphere Model version 1. Geoscientific Model Development, 15(7), 2881–2916.doi:10.5194/gmd-15-2881-2022, 2022.

Mace, G. G., Benson, S., Sonntag, K. L., Kato, S., Min, Q., Minniss, P., Twohy, C. H., Poellot, M., Dong, X., Long, C., Zhang, Q., and Doelling, D. R.: Cloud radiative forcing at the ARM Program Climate Rsearch Facility: 1. Technique, validation, and

455 comparison to satellite-derived diagnostic quantities. J. Geophys. Res., 111, D11S90, doi:10.1029/2005JD005921, 2006.

Marchand, R., T. Ackerman, M. Smyth, and W. B. Rossow: A review of cloud top height and optical depth histograms from MISR, ISCCP, and MODIS. J. Geophys. Res., doi: 10.1029/2009JD013422, 2010.

Morrison, H., and Gettelman, A.: A New Two-Moment Bulk Stratiform Cloud Microphysics Scheme in the Community Atmosphere Model, Version 3 (CAM3). Part I: Description and Numerical Tests. Journal of Climate, 21(15), 3642–3659,

460 https://doi.org/10.1175/ 2008JCLI2105.1, 2008.

Neale, R. B., Richter, J. H., and Jochum, M.: The impact of convection on ENSO: From a delayed oscillator to a series of events. Journal of Climate, 21(22), 5904–5924. https://doi.org/ 10.1175/2008JCLI2244.1, 2008.

Pincus, R., Platnick, S., Ackerman, S. A., Hemler, R. S., & Patrick Hofmann, R. J.: Reconciling Simulated and Observed Views of Clouds: MODIS, ISCCP, and the Limits of Instrument Simulators. J. Climate, https://doi.org/10.1175/JCLI-D-11-

465 00267.1, 2012.

Qian Y, H. Wan, B. Yang, J.C. Golaz, B. Harrop, Z. Hou, V.E. Larson, L. R Leung, G. Lin, W. Lin, P. Ma, H. Ma, P. Rasch, B. Singh, H. Wang, S. Xie, and K. Zhang, Parametric sensitivity and uncertainty quantification in the version 1 of E3SM Atmosphere Model based on short Perturbed Parameters Ensemble simulations, J. Geophysical Research, 123, no. 23:13046-13073, doi:10.1029/2018JD028927, 2018.

470 Qin, Y., Zheng, X., Klein, S. A., Zelinka, M. D., Ma, P.-L., Golaz, J.-C. and Xie, S.: Causes of Reduced Climate Sensitivity in E3SM from Version 1 to Version 2. Submitted to JAMES, 2023.

Rasch, P. J., Xie, S., Ma, P.-L., Lin, W., Wang, H., Tang, Q., Bur- rows, S. M., Caldwell, P., Zhang, K., Easter, R. C., Cameron-Smith, P., Singh, B., Wan, H., Golaz, J.-C., Harrop, B. E., Roesler, E., Bacmeister, J., Larson, V. E., Evans, K. J., Qian, Y., Taylor, M., Leung, L. R., Zhang, Y., Brent, L., Branstet- ter, M., Hannay, C., Mahajan, S., Mametjanov, A., Neale, R., Richter,

475 J. H., Yoon, J.-H., Zender, C. S., Bader, D., Flan- ner, M., Foucar, J. G., Jacob, R., Keen, N., Klein, S. A., Liu, X., Salinger, A. G., Shrivastava, M., and Yang, Y.: An overview of the atmospheric component of the Energy Exascale Earth System Model, Journal of Advances in Modeling Earth Systems,11(8), 2377–2411, doi: 10.1029/2019ms001629, 2019.

Redelsperger, J.-L., Guichard, F., and Mondon, S.: A parameterization of mesoscale enhancement of surface fluxes for large-scale models, Journal of Climate,13(2), 402 - 421. doi: 10.1175/1520-0442(2000)013 〈0402:APOMEO〉 2.0.CO;2, 2000.

480 Rossow, W. B., and R. A. Schiffer: Advances in understanding clouds from ISCCP, Bull. Amer. Meteor. Soc., 80, 2261–2288. https://doi.org/10.1175/1520-0477(1999)080<2261: AIUCFI>2.0.co;2, 1999.

Rossow, W.B.; Walker, A.; Golea, V.; Knapp, K. R.; Young, A.; Inamdar A.; Hankins, B.; and NOAA's Climate Data Record Program (2016): International Satellite Cloud Climatology Project Climate Data Record, H-Series NOAA National Centers for Environmental Information. doi:10.7289/V5QZ281S, 2016.

485 Swales, D., R. Pincus, and A. Bodas-Salcedo: The Cloud Feedback Model Intercomparison Project Observational Simulator Package: Version 2, Geosci. Model Dev., 11, 77-81, https://doi.org/10.5194/gmd-11-77-2018, 2018.

Tang, Q., Golaz, J.-C., Van Roekel, L. P., Taylor, M. A., Lin, W., Hillman, B. R., Ullrich, P. A., Bradley, A. M., Guba, O., Wolfe, J. D., Zhou, T., Zhang, K., Zheng, X., Zhang, Y., Zhang, M., Wu, M., Wang, H., Tao, C., Singh, B., Rhoades, A. M., Qin, Y., Li, H.-Y., Feng, Y., Zhang, Y., Zhang, C., Zender, C. S., Xie, S., Roesler, E. L., Roberts, A. F., Mametjanov, A.,

490 Maltrud, M. E., Keen, N. D., Jacob, R. L., Jablonowski, C., Hughes, O. K., Forsyth, R. M., Di Vittorio, A. V., Caldwell, P. M., Bisht, G., McCoy, R. B., Leung, L. R., and Bader, D. C.: The fully coupled regionally refined model of E3SM version 2: overview of the atmosphere, land, and river results, Geosci. Model Dev., 16, 3953–3995, https://doi.org/10.5194/gmd-16-3953-2023, 2023.

Wang, Y., Liu, X., Hoose, C., and Wang, B.: Different contact angle distributions for heterogeneous ice nucleation in the

495 Community Atmospheric Model version 5. Atmospheric Chemistry and Physics, 14(19), 10411–10430. https://doi.org/10.5194/acp-14-10411-2014, 2014.

Wang, Y.-C., Pan, H.-L., & Hsu, H.-H.: Impacts of the triggering function of cumulus parameterization on warm-season diurnal rainfall cycles at the atmospheric radiation measurement southern great plains site, J. Geophys. Res. - Atmosphere, 120(20), 10681-10702, doi:10.1002/2015JD023337, 2015.

500 Wang, Y.-C., Xie, S., Tang, S., and Lin, W.: Evaluation of an improved convective triggering function: Observational evidence and SCM tests, J. Geophys. Res - Atmosphere, 125, 2019JD031651. https://doi.org/10.1029/2019JD031651, 2020.

Xie, S., and Zhang, M.: Impact of the convection triggering function on single-2156column model simulations.Journal of Geophysical Research - Atmosphere,2157105(D11), 14983-14996. doi: 10.1029/2000JD900170, 2000.

Xie, S., Ma, H.-Y., Boyle, J. S., Klein, S. A., and Zhang, Y.: On the correspondence between short- and long-time-scale

505 systematic errors inCAM4/CAM5 for the year of tropical convection.Journal of Climate,25(22), 7937–7955. https://doi.org/10.1175/JCLI-D-12-00134.1, 2012.

Xie, S., Lin, W., Rasch, P. J., Ma, P.-L., Neale, R., Larson, V. E., Qian, Y., Bogenschutz, P. A., Caldwell, P., Cameron-Smith, P., Golaz, J.-C., Mahajan, S., Singh, B., Tang, Q., Wang, H., Yoon, J.-H., Zhang, K., and Zhang Y.: Understanding cloud and convective characteristics in version 1 of the E3SM atmosphere model, J. Adv. Model. Earth Sy., 10, 2618–2644,

510 https://doi.org/10.1029/2018MS001350, 2018.

Xie, S., Wang, Y.-C., Lin, W., Ma, H.-Y., Tang, Q., Tang, S., Zheng, X., Golaz, J., Zhang, G. J., and Zhang, M.: Improved Diurnal Cycle of Precipitation in E3SM with a Revised Convective Triggering Function. Journal of Advances in Modeling Earth Systems, 11. https://doi.org/10.1029/2019MS001702, 2019.

Zhang, G. J., and McFarlane, N. A.: Sensitivity of climate simulations to the parameterization of cumulus convection in the Canadian climate centre general circulation model, Atmosphere-Ocean, 33(3), 407–446. https://doi.org/10.1080/07055900.1995.9649539, 1995.

Zhang, M., Xie, S., Liu, X., Lin, W., Zheng, X., Golaz, J.-C., and Zhang, Y.: Cloud phase simulation at high latitudes in EAMv2: Evaluation using CALIPSO observations and comparison with EAMv1, J. Geophys. Res. – Atmospheres, 127, e2022JD037100. https://doi.org/10.1029/2022JD037100, 2022.

Zhang, Y., Xie, S., Covey, C., Lucas, D. D., Gleckler, P., Klein, S. A., et al.: Regional assessment of the parameter-dependent performance of CAM4 in simulating tropical clouds. Geophy. Res. Let, 39(14), L14708. Doi: 10.1029/2012GL052184, 2012.

Zhang, Y., Xie, S., Lin, W., Klein, S. A., Zelinka, M., Ma, P.-L., Rasch, P. J., Qian, Y., Tang, Q., and Ma, H.-Y.: Evaluation of Clouds in Version 1 of the E3SM Atmosphere Model With Satellite Simulators. Journal of Advances in Modeling Earth Systems, 11(5), 1253-1268, doi: 10.1029/2018MS001562, 2019.

 Table 1. Satellite data used in the study

	Quantity	Time Period	Temporal/spatial resolution	Data Source	Reference
ISCCP#	Cloud cover - joint histogram	1983-2017	Monthly / 1 deg	ISCCP H	Rossow et al. 2016 Pincus et al. 2012 Zhang et al. 2012
MODIS#	Cloud cover - joint histogram	2002-2016	Monthly / 1 deg	Collection 5.1	Pincus et al. 2012
MISR#	Cloud cover -joint histogram	2000-2020	Monthly / 1 deg	CTH-OD	Marchand et al. 2010
CALIPSO#	High, middle, and low cloud fraction	2006-2020	Monthly / 2 deg	GOCCP v3.1.2	Chepfer et al. 2010
CERES	TOA shortwave and longwave cloud radiative effect	2000-2021	Monthly / 1 deg	CERES_EBAF_Ed4.1	Loeb et al. 2018

#: The joint histogram of cloud cover (as a function of cloud-top pressure (*ctp*) or cloud-top height (*cth*) and optical thickness (*Tau*)) and CALIPSO cloud data are from the CFMIP GCM Simulator-Oriented cloud products developed and can be
530 downloaded from http://climserv.ipsl.polytechnique.fr/cfmip-obs/.

Table 2. Description of simulations

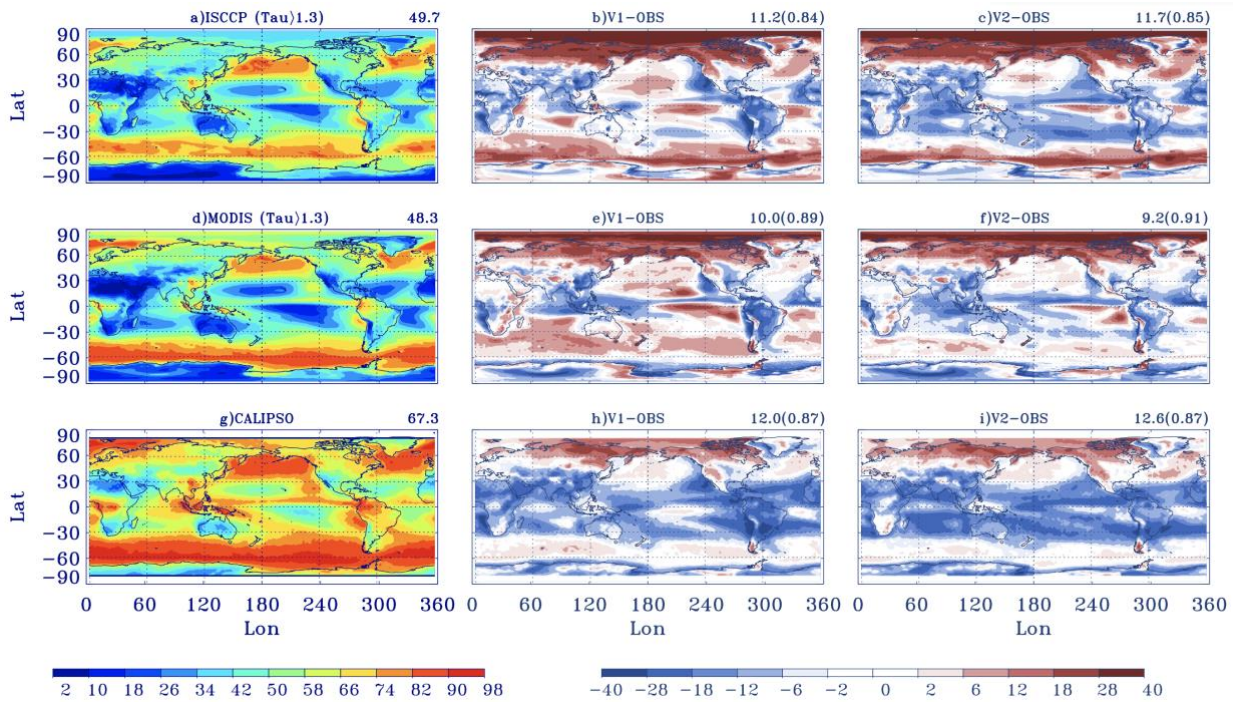
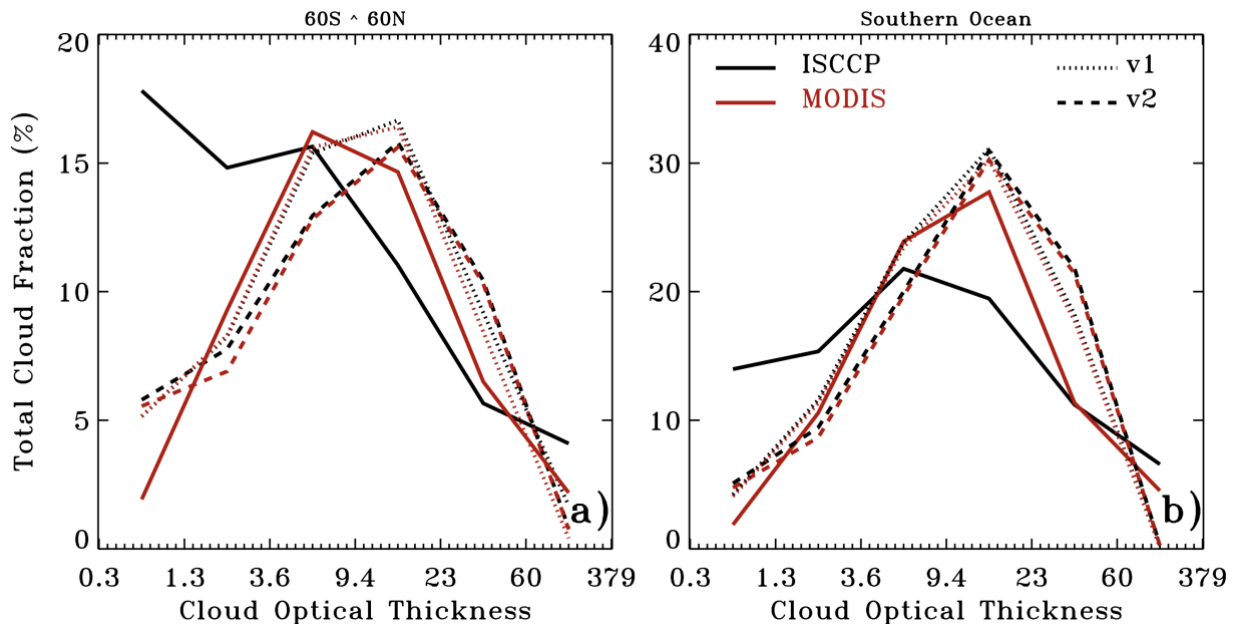
Name	Description	Purpose
v1_coupled	The last 30-year (1985-2014) data from the 165-year historical simulations with E3SMv1	E3SMv1 baseline
v2_coupled	The last 30-year (1985-2014) data from the 165-year historical simulations with E3SMv2	E3SMv2 baseline
v2	EAM v2 configuration: 6 yr control – AMIP	V2 baseline for sensitivity tests
Trigonly	V2 with ZM trigger reverted to v1, i.e., turn off the dCAPE_ULL trigger	Test the impact of the new trigger
clubbonly	V2 with CLUBB related parameters reverted to v1	Test the impact of CLUBB tuning
MGonly	V2 with MG2 related parameters reverted to v1 (Bergeron factor, minimum CDNC, accretion factor, autoconversion factor, …)	Test the impact of MG tuning
ZMonly	V2 with ZM related parameters reverted to v1, dCAPE_ULL trigger retained	Test the impact of ZM tuning

535 **Figure 1.** Annual mean total cloud fraction from a) ISCCP (Tau>1.3), d) MODIS (Tau>1.3), and g) CALIPSO, the differences between E3SMv1 and observations b), e), and h), and the differences between E3SMv2 and observations c), f), and i). The number after each panel name is the global annual mean of cloud fraction for observations or the global annual mean RMSE and correlation (in parentheses) for cloud fraction differences.

540

Figure 2. Column-integrated cloud optical depth distributions averaged a) over 60S-60N and b) over Southern Ocean (45S-60S) for ISCCP, MODIS, E3SMv1 and E3SMv2.

545

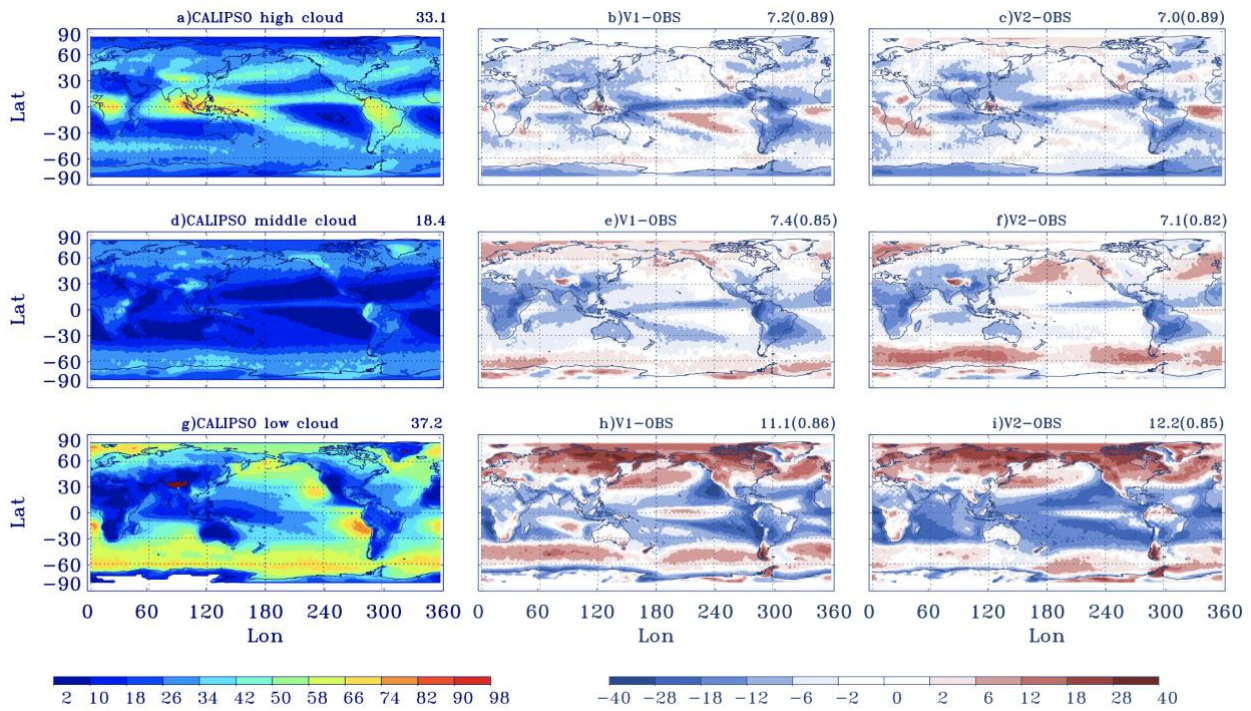



Figure 3. Annual mean CALIPSO-derived a) high, d) middle, g) low cloud fraction, the differences between E3SMv1 and observation for b), e), h), and the differences between E3SMv2 and observation for c), f), i). The number after each panel name is the global annual mean of cloud fraction for observations or the global annual mean RMSE and correlation (in parentheses) for cloud fraction differences.

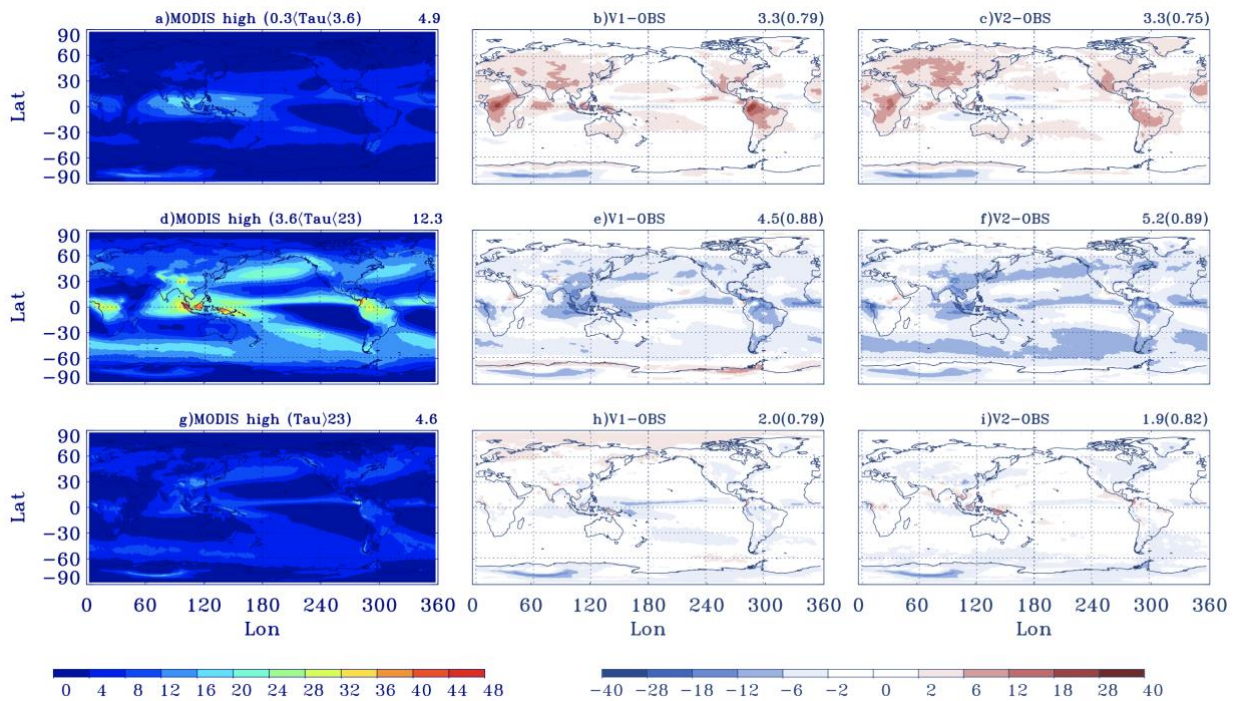



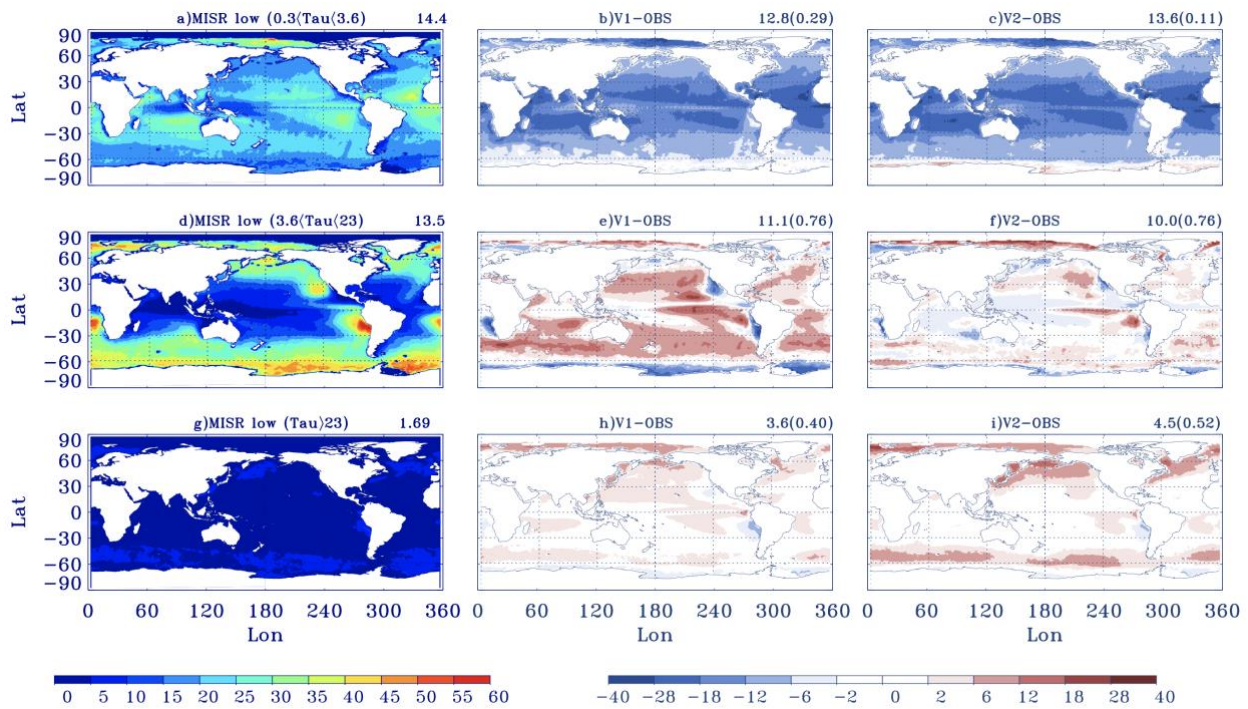
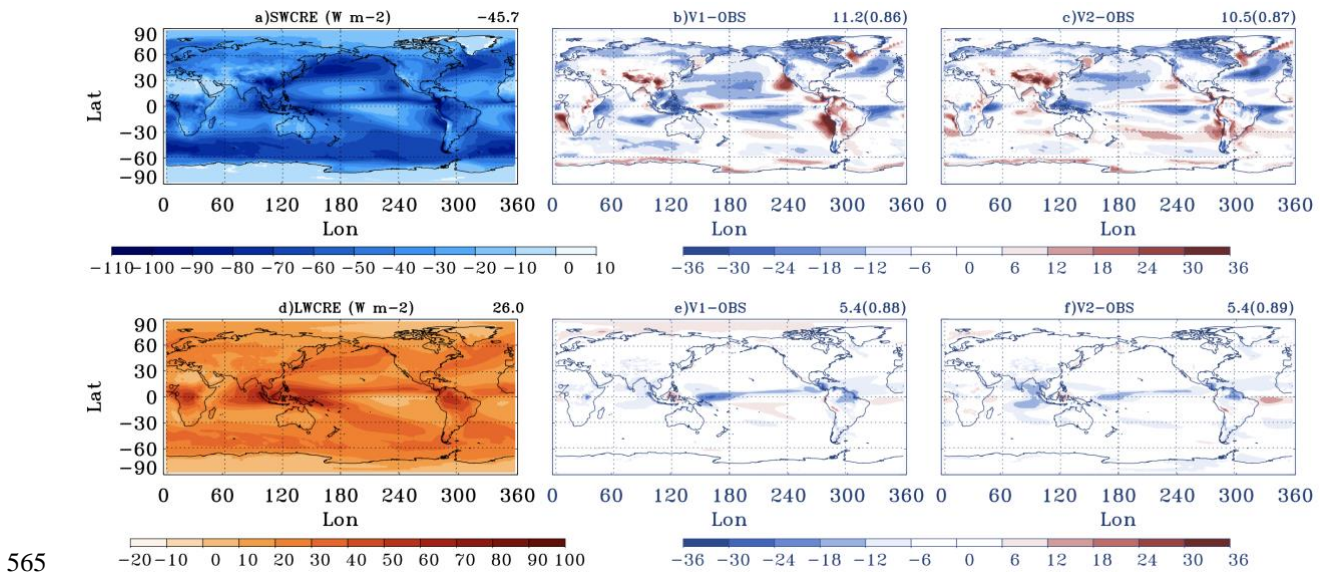
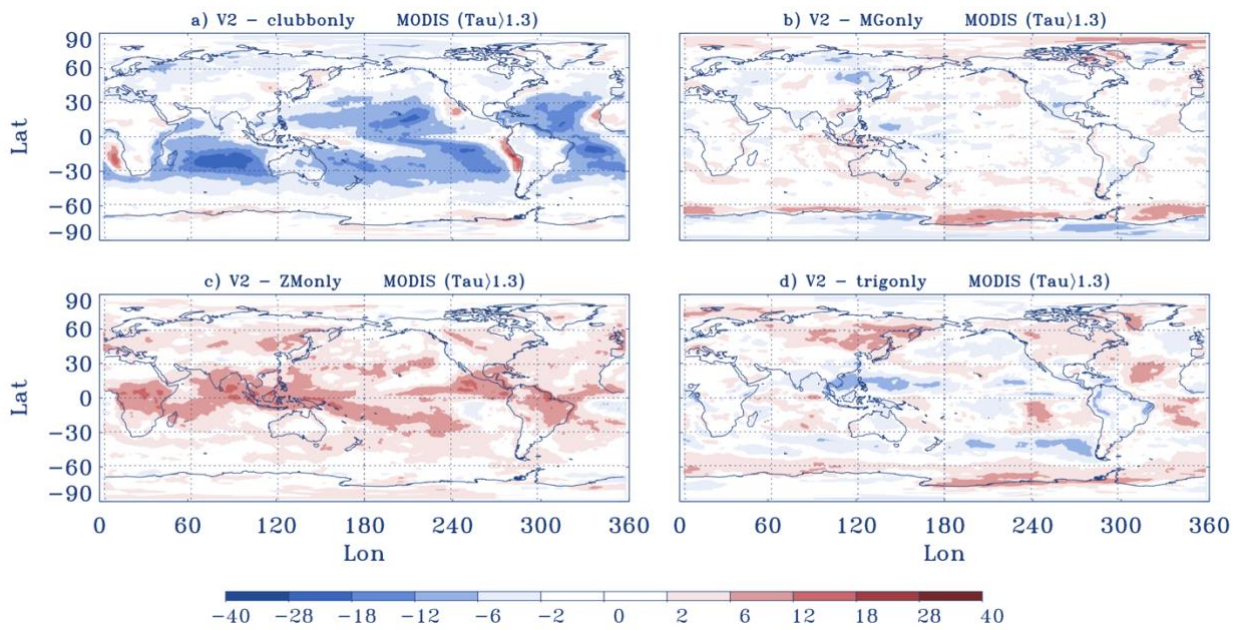
Figure 4. Annual mean MODIS high-topped thin (0.3<Tau<3.6), intermediate (3.6<Tau<23), and thick (Tau>23) clouds. a) & d) & g) are MODIS observations, b) & e) & h) are the differences between E3SMv1 and observation, and c) & f) & i) are the differences between E3SMv2 and observation. The number after each panel name is the global annual mean of cloud fraction for observations or the global annual mean RMSE and correlation (in parentheses) for cloud fraction differences.

560 **Figure 5.** Annual mean MISR low-topped optically thin (0.3<Tau<3.6), intermediate thickness (3.6<Tau<23) and thick (Tau > 23) clouds. a) & d) & g) are MISR observations, b) & e) & h) are the differences between E3SMv1 and observation, and c) & f) & i) are the differences between E3SMv2 and observation. The number after each panel name is the global annual mean of cloud fraction for observations or the global annual mean RMSE and correlation (in parentheses) for cloud fraction differences.

565

Figure 6. Annual mean a) SWCRE and d) LWCRE for CERES, and b) and e) differences between E3SMv1 and CERES, and c) and f) between E3SMv2 and CERES. The number after each panel name is the global annual mean of CRE for observations or the global annual mean RMSE and correlation (in parentheses) for cloud fraction differences.

570

575 **Figure 7.** Difference in cloud fraction from MODIS simulator (Tau>1.3) between sensitivity tests and the default E3SMv2 run. a) E3SMv2 with CLUBB related parameters changed from v2 to v1; b) E3SMv2 with MG2 related parameters changed from v2 to v1; c) E3SMv2 with ZM related parameters changed from v2 to v1; and d) E3SMv2 with the dCAPE_ULL trigger turned off. All results are from 6-year AMIP-style climatology runs.

580

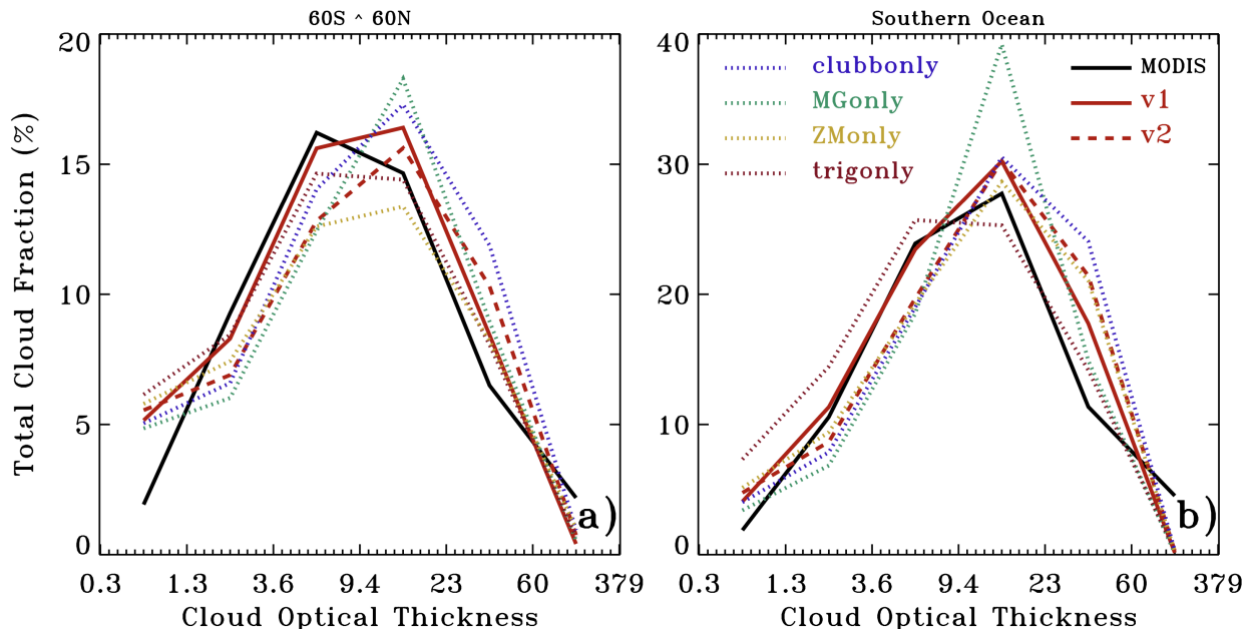



Figure 8. Column-integrated cloud optical depth distributions averaged a) over 60S-60N and b) over Southern Ocean (45S-60S) for MODIS, E3SMv1, E3SMv2, and sensitivity tests.

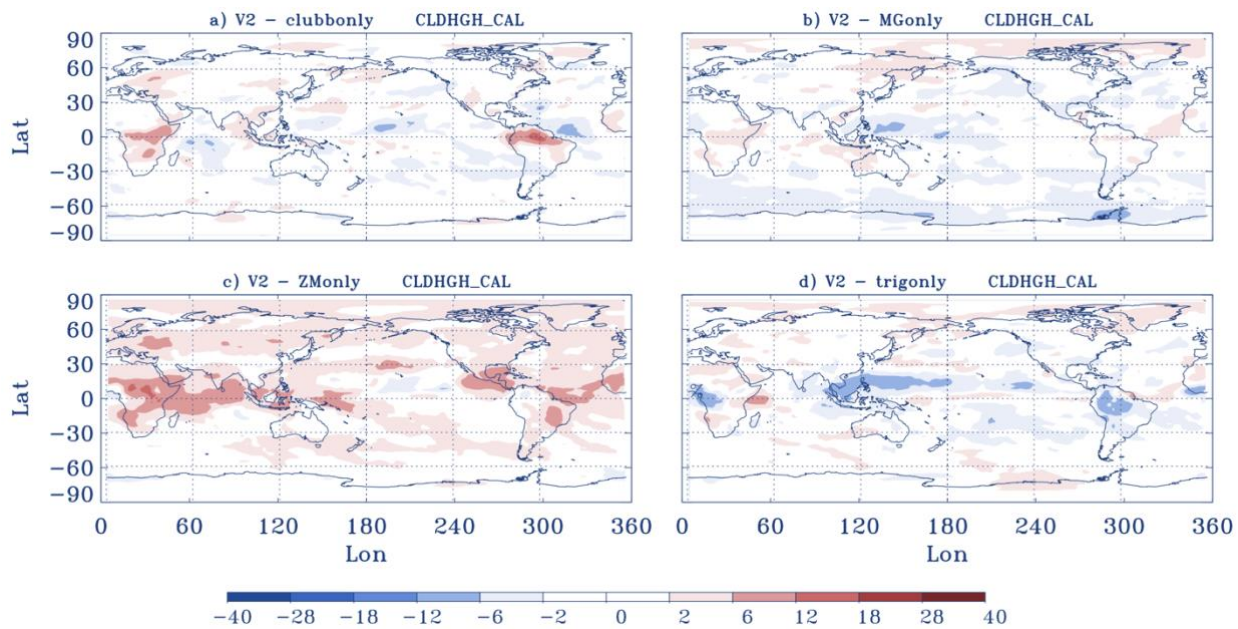
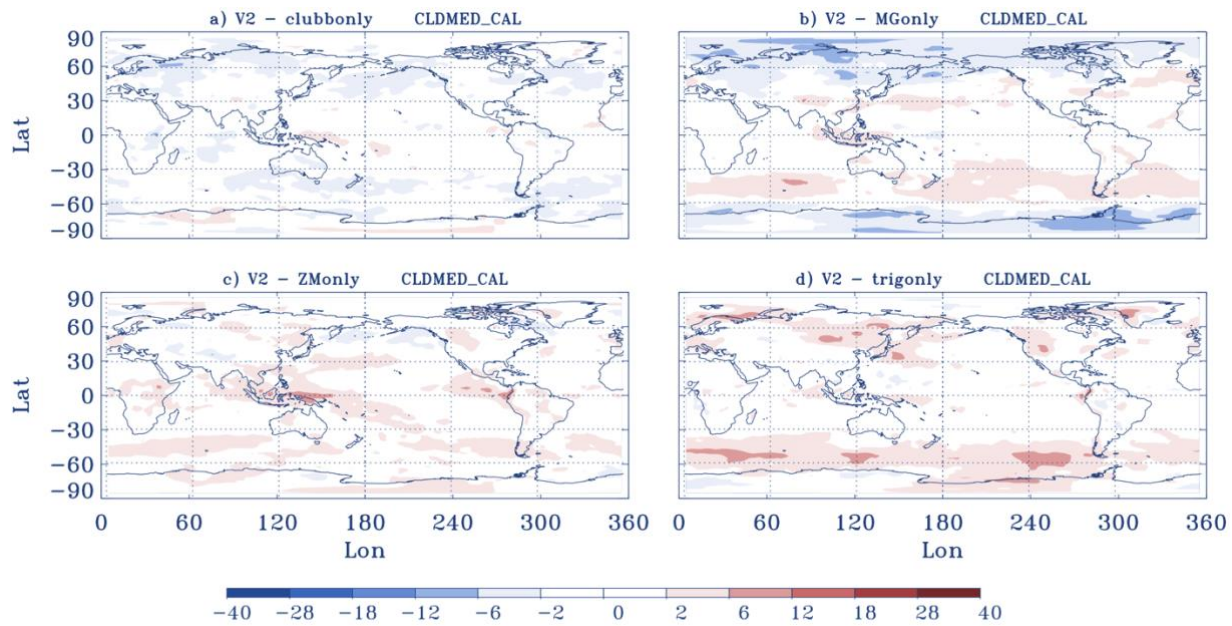
585

590 **Figure 9.** Difference in high cloud fraction from CALIPSO simulator between sensitivity tests and the default E3SMv2 run. a) E3SMv2 with CLUBB related parameters changed from v2 to v1; b) E3SMv2 with MG2 related parameters changed from v2 to v1; c) E3SMv2 with ZM related parameters changed from v2 to v1; and d) E3SMv2 with the dCAPE_ULL trigger turned off. All results are from 6-year AMIP-style climatology runs.

595

Figure 10. Difference in middle-level cloud fraction from CALIPSO simulator between sensitivity tests and the default E3SMv2 run. a) E3SMv2 with CLUBB related parameters changed from v2 to v1; b) E3SMv2 with MG2 related parameters changed from v2 to v1; c) E3SMv2 with ZM related parameters changed from v2 to v1; and d) E3SMv2 with the dCAPE_ULL trigger turned off. All results are from 6-year AMIP-style climatology runs.

600

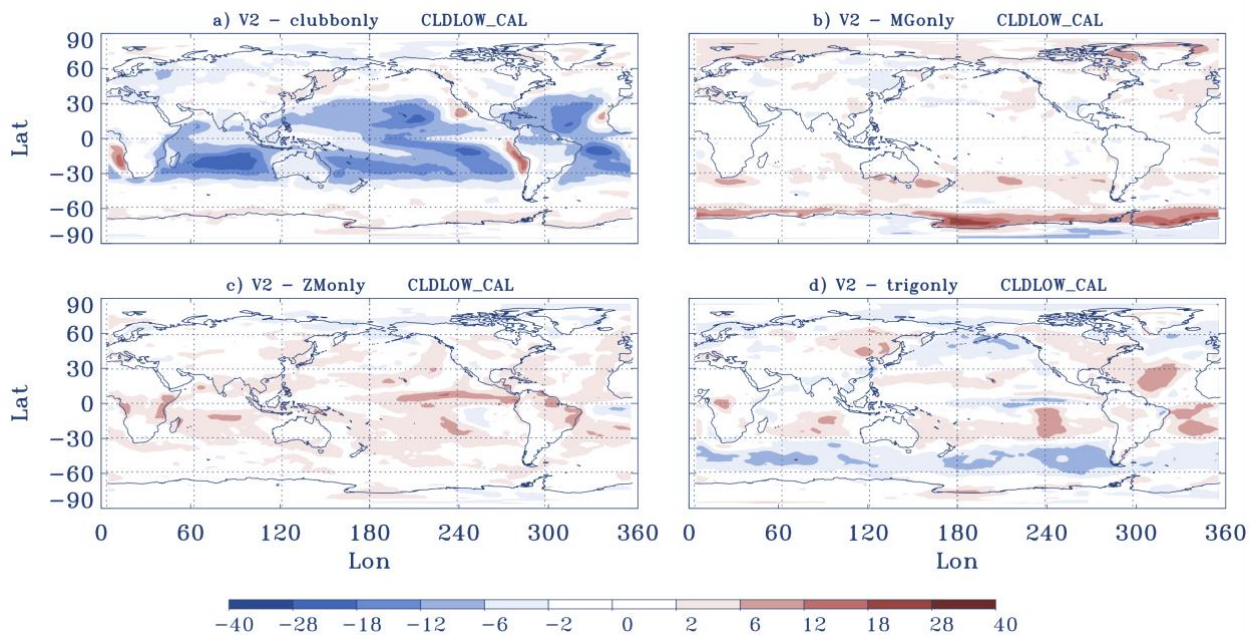



Figure 11. Difference in low cloud fraction from CALIPSO simulator between sensitivity tests and the default E3SMv2 run. a) E3SMv2 with CLUBB related parameters changed from v2 to v1; b) E3SMv2 with MG2 related parameters changed from v2 to v1; c) E3SMv2 with ZM related parameters changed from v2 to v1; and d) E3SMv2 with the dCAPE_ULL trigger turned off. All results are from 6-year AMIP-style climatology runs.

610

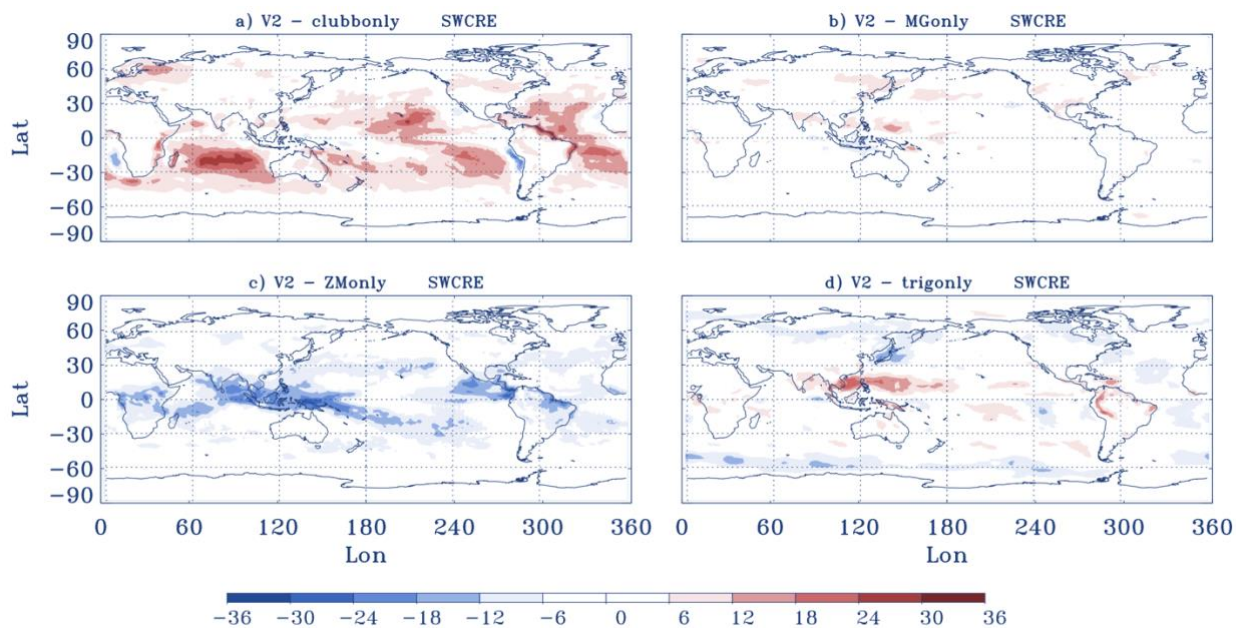



Figure 12. Difference in SWCRE between sensitivity tests and the default E3SMv2 run. a) E3SMv2 with CLUBB related parameters changed from v2 to v1; b) E3SMv2 with MG2 related parameters changed from v2 to v1; c) E3SMv2 with ZM related parameters changed from v2 to v1; and d) E3SMv2 with the dCAPE_ULL trigger turned off. All results are from 6-year AMIP-style climatology runs.

620

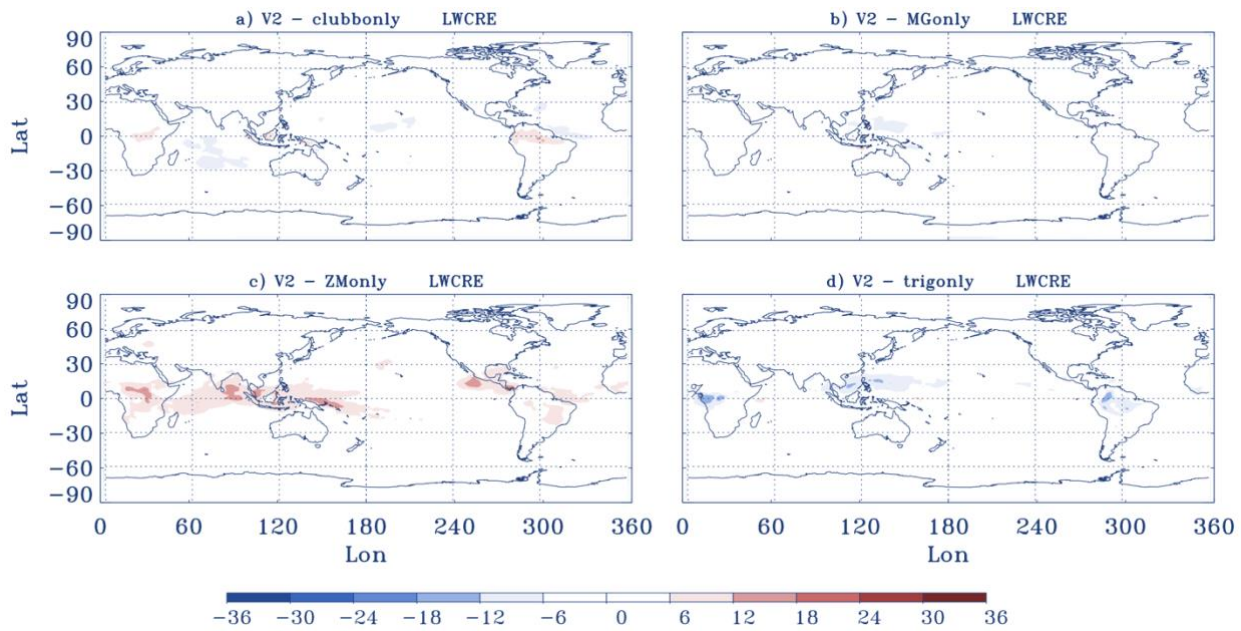



Figure 13. Difference in LWCRE between sensitivity tests and the default E3SMv2 run. a) E3SMv2 with CLUBB related parameters changed from v2 to v1; b) E3SMv2 with MG2 related parameters changed from v2 to v1; c) E3SMv2 with ZM related parameters changed from v2 to v1; and d) E3SMv2 with the dCAPE_ULL trigger turned off. All results are from 6-year AMIP-style climatology runs.

630

635

Appendix A

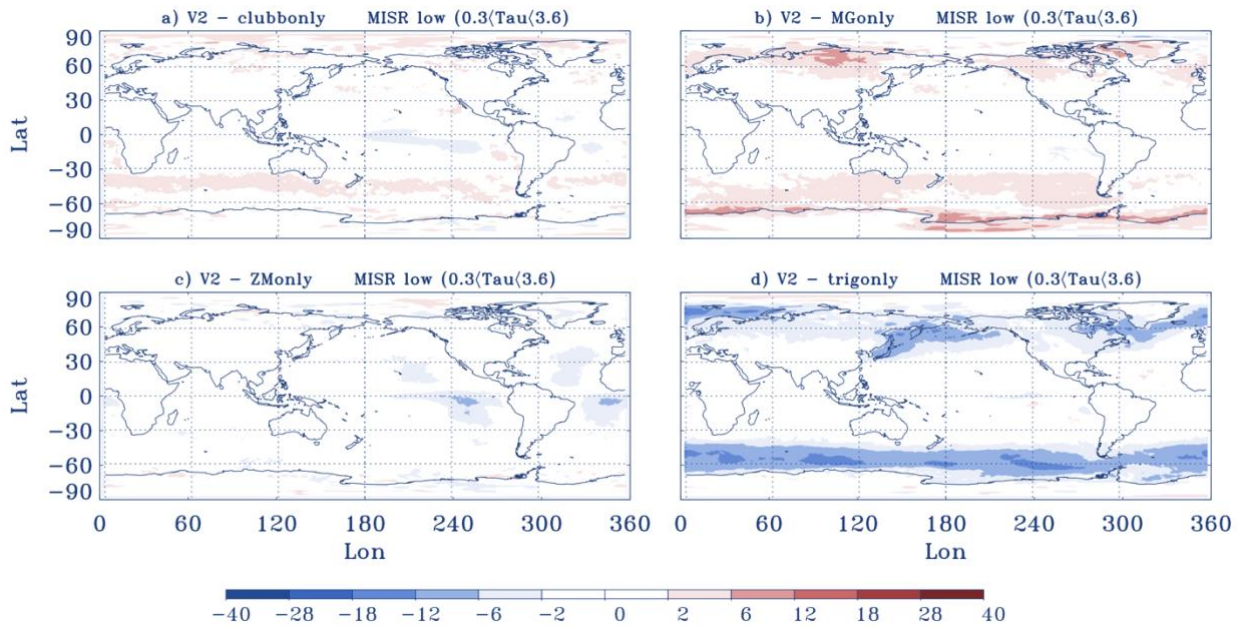



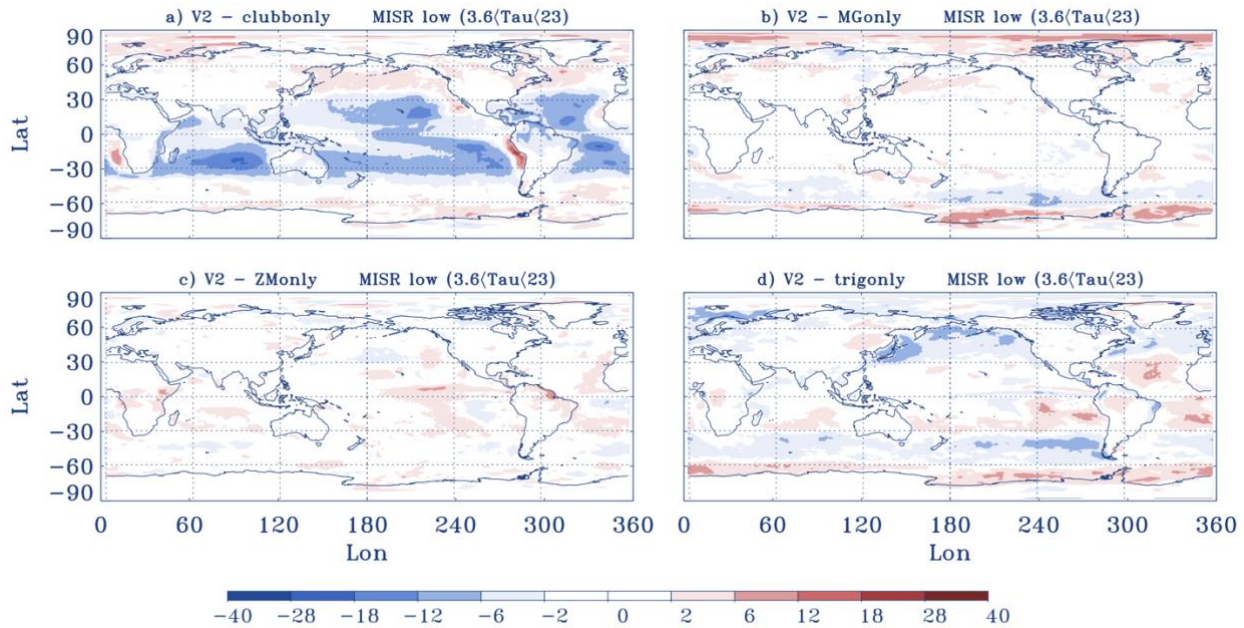
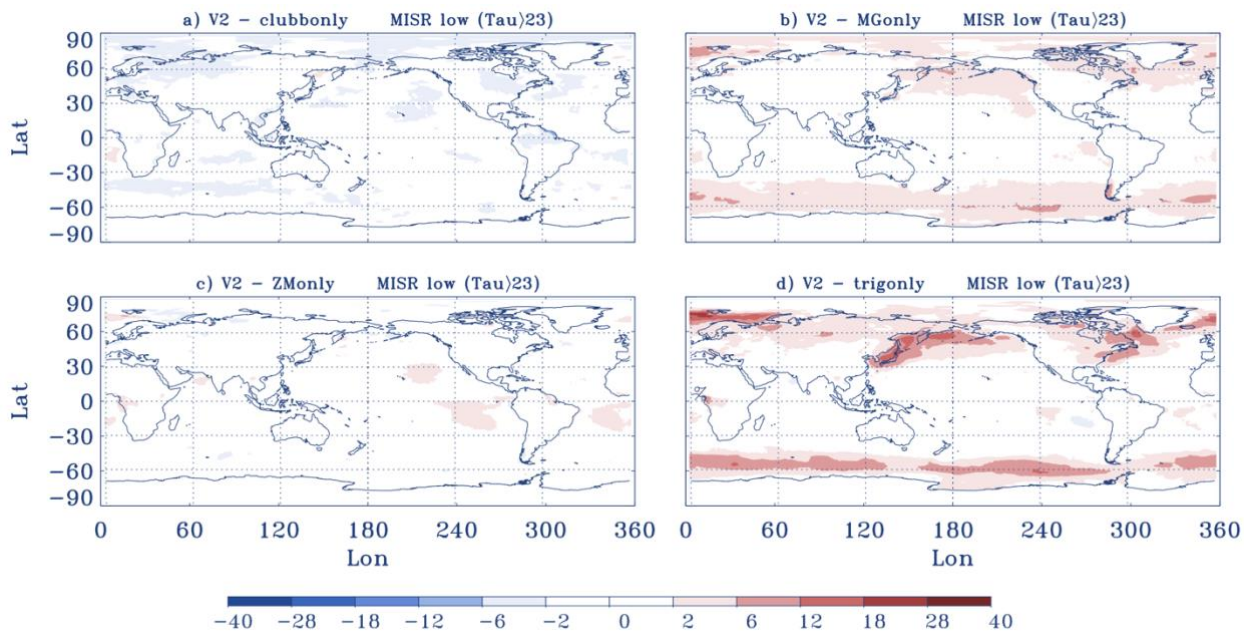
Figure A1. Difference in optically thin low cloud fraction from MISR simulator (0.3 < Tau < 3.6) between sensitivity tests and the default E3SMv2 run. a) E3SMv2 with CLUBB related parameters changed from v2 to v1; b) E3SMv2 with MG2 related parameters changed from v2 to v1; c) E3SMv2 with ZM related parameters changed from v2 to v1; and d) E3SMv2 with the dCAPE_ULL trigger turned off. All results are from 6-year AMIP-style climatology runs.

645   **Figure A2.** Difference in optically intermediate low cloud fraction from MISR simulator (3.6 < Tau < 23) between sensitivity tests and the default E3SMv2 run. a) E3SMv2 with CLUBB related parameters changed from v2 to v1; b) E3SMv2 with MG2 related parameters changed from v2 to v1; c) E3SMv2 with ZM related parameters changed from v2 to v1; and d) E3SMv2 with the dCAPE_ULL trigger turned off. All results are from 6-year AMIP-style climatology runs.

650

Figure A3. Difference in optically think low cloud fraction from MISR simulator (Tau > 23) between sensitivity tests and the default E3SMv2 run. a) E3SMv2 with CLUBB related parameters changed from v2 to v1; b) E3SMv2 with MG2 related parameters changed from v2 to v1; c) E3SMv2 with ZM related parameters changed from v2 to v1; and d) E3SMv2 with the dCAPE_ULL trigger turned off. All results are from 6-year AMIP-style climatology runs.

655

Figure A4. Difference in optically thin high cloud fraction from MODIS simulator (0.3 < Tau < 3.6) between sensitivity tests and the default E3SMv2 run. a) E3SMv2 with CLUBB related parameters changed from v2 to v1; b) E3SMv2 with MG2 related parameters changed from v2 to v1; c) E3SMv2 with ZM related parameters changed from v2 to v1; and d) E3SMv2 with the dCAPE_ULL trigger turned off. All results are from 6-year AMIP-style climatology runs.

 Figure A5. Difference in optically intermediate high cloud fraction from MODIS simulator (3.6<Tau <23) between sensitivity tests and the default E3SMv2 run. a) E3SMv2 with CLUBB related parameters changed from v2 to v1; b) E3SMv2 with MG2 related parameters changed from v2 to v1; c) E3SMv2 with ZM related parameters changed from v2 to v1; and d) E3SMv2 with the dCAPE_ULL trigger turned off. All results are from 6-year AMIP-style climatology runs.

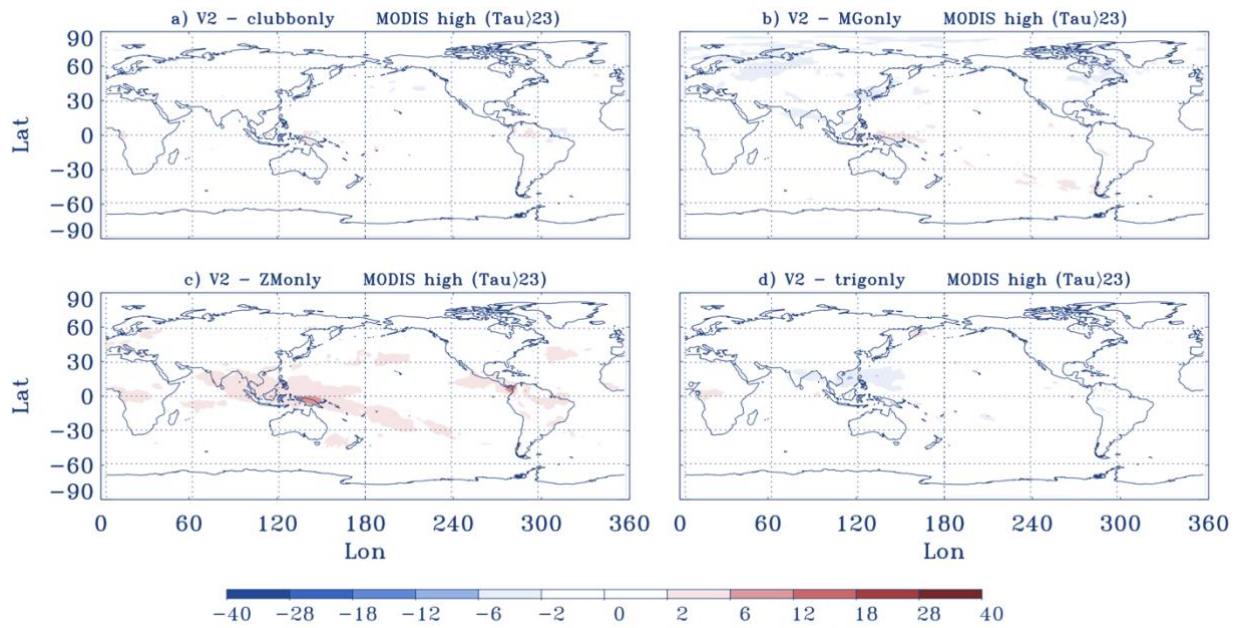
670

675

Figure A6. Difference in optically thick high cloud fraction from MODIS simulator (Tau >23) between sensitivity tests and the default E3SMv2 run. a) E3SMv2 with CLUBB related parameters changed from v2 to v1; b) E3SMv2 with MG2 related parameters changed from v2 to v1; c) E3SMv2 with ZM related parameters changed from v2 to v1; and d) E3SMv2 with the dCAPE_ULL trigger turned off. All results are from 6-year AMIP-style climatology runs.

680