

# Uncertainties, complexities and possible forecasting of the volcán de Colima energy emissions (México, years 2013-2015) based on the fractal reconstruction theorem

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## 10 **Abstract.**

The effusive-explosive energy emissions process in a volcano is a dynamic and complex physical phenomenon. The importance of quantifying this complexity in terms of the physical and mathematical mechanisms that govern these emissions should be a requirement for deciding to apply a possible forecasting strategy with a sufficient degree of certainty.

15 The complexity of this process is determined in this **research by means of** the reconstruction theorem and statistical procedures **applied to** the effusive-explosive volcanic energy emissions **corresponding to the** activity in the Volcán de Colima (western segment of the Trans-Mexican volcanic belt) **along** the years 2013-2015. The analysis is focused on measuring the degree of persistence, or randomness of the series; the degree of predictability of energy emissions; and the quantification of the degree of complexity and “memory loss” of the physical mechanism throughout an episode of volcanic  
20 emissions. The results indicate that the analysed time series **depict** a high degree of persistence and low memory loss, **becoming the mentioned effusive-explosive volcanic emission structure a** candidate to successfully apply a forecasting strategy.

**Key-words:** **Volcan de Colima explosions, México, fractal reconstruction theorem, Hurst, Lyapunov and Kolmogorov**  
25 **exponents, possible forecasting and nowcasting strategies.**

## **1 Introduction**

A right forecasting of dangerous long drought episodes, high magnitude earthquakes or great volcanic emissions should be one of the main objectives on the scientific fields of climatology, seismology or volcanology to prevent disasters which could affect the environment and the human life. Several examples of forecasting algorithms could be cited, among them the  
30 nowcasting strategy (Rundle et al., 2016, 2017), the multifractal analysis in seismology (Monterrubio-Velasco et al., 2020),

the ARIMA process in climatic research (Lana et al, 2021), and neural algorithms (Lipton et al, 2015 and Lei, 2021) also useful to predict monthly rainfall. These cited algorithms systematically forecast the next episode taking into account a certain number of previous recorded data, being this number strongly associated with the characteristics of the physical mechanism. These forecasting results should be validated by previously analysing the degree of complexity and the “loss of memory” of the physical mechanisms along the evolution of the physical process. In other words, how many previous data would be necessary for a right use of a forecasting algorithm and which could be the range of uncertainties on the predictions.

The database, Section 2, analysed in this research is the time series of explosive volcanic events (Vulcanian explosions) emitted by Volcán de Colima (Western segment of Trans-Mexican volcanic belt) during the years 2013-2015. A Vulcanian explosion is an eruption where fragmented material is expelled to the atmosphere, as result of an overpressure into the conduit or lava dome (Arámbula-Mendoza et al., 2018). The event releases energy in several ways, based on elastic, seismic, acoustic and thermal processes. The next explosion will occur when again the overpressure break again the impermeable cap. The Volcán de Colima has emitted many Vulcanian explosions, some of them with generation of Pyroclastic Density Currents (PDCs) until 5 km of runout (Arámbula-Mendoza et al., 2019). For these reasons, strategies for a right forecasting of the mentioned Vulcanian explosions are important.

The reconstruction theorem (Section 3) is a mathematical strategy that allows quantify the degree of complexity in a time-series and its loss of memory, both required to validate a possible forecasting strategy (Dicks, 1999). Additionally, the nowcasting algorithm (Section 5), a statistical process developed by Rundle et al. (2016, 2017) to detect the risk of imminent high magnitude earthquakes, could be also applied to quantify the probability of an imminent high magnitude volcanic emission.

The main objective of this work is to detect the degree of difficulty for a forecasting of volcanic emissions associated with energies close to or exceeding  $10^8$  J. By means of the reconstruction theory (Diks, 1999) the complexity is measured by means of some parameters such as the persistence degree, the intensity of chaotic behavior of the system, and the “loss of memory” of the physical mechanism. Moreover, the statistical distribution of the effusive-explosive energy of these emissions and their return periods are also analysed. The just mentioned nowcasting strategy is also taken into account to confirm future extreme emissions of energy. The results obtained after applying the reconstruction theorem, analysing the whole time series, 6 consecutive segments and 21 moving window data, are detailed in Section 4.

The most relevant results of the reconstruction theorem and their effects on forecasting algorithms are discussed in Section 6. Finally, the Section 7 (Conclusions), summarizes the most relevant results with respect to the expected success on preventing volcanic energy emissions, based on forecasting and nowcasting processes.

## 2. Database

A time series of volcanic explosions, named as Vulcanian explosions (Bachtell Clarke and Esposti Ongaro, 2015) emitted by Volcán de Colima (Western segment of Trans-Mexican volcanic belt, years 2013-2015) (Arámbula-Mendoza et al., 2018) is analysed. Figure 1a depicts the histogram of the logarithm of the emitted energy. The dataset contains 6182 observations of the emissions equalling to or exceeding approximately  $2 \times 10^6$  J accomplishing the Gutenberg-Richter law (Gutenberg and Richter, 1958) as shows Figure 1b.

Figure 2a describes the six segments to be analysed, where can be observed the highest explosions at the beginning of the first segment [ $\log_{10}(\text{Energy})=8.2$ ], third segment [ $\log_{10}(\text{Energy})=8.4$ ] and at the end of the series, out of the 6th segment [ $\log_{10}(\text{Energy})=8.9$ ]. With the aim of analysing the whole set of volcanic emissions accomplishing the Gutenberg-Richter law, Figure 2b depicts two examples of moving window segments.

The statistical distribution of these emissions is analysed by means of the L-Skewness-Kurtosis formulation (Hosking and Wallis, 1997). The statistical analysis of these emissions shows that the complete series of emissions, including those not accomplishing the Gutenberg-Richter law, are well fitted to the Generalised Logistic, GL, function (Figures 3a and 3b). Additionally, three different empirical distributions of extreme emissions, equalling to or exceeding respectively 90%, 95% and 99% of data (Figure 3b), can be associated with the Generalised Extreme value, GEV, function. Figure 4 shows the evolution of these three expected extreme emissions with the increasing return periods (given in number of events equalling to or exceeding 90, 95 and 99% respectively). For instance, the expected values of emissions for the three percentage levels and return periods up to 200 extreme emissions fit quite well the theoretical evolution, with emissions close to  $1.0 \times 10^8$  J,  $2.0 \times 10^8$  J and  $8.0 \times 10^8$  J for 90, 95 and 99% extreme distributions. This first approach to the possibility of very high explosions and the corresponding expected return period (number of extreme episodes before a very high extreme emission) would be quite similar to a nowcasting analysis, strategy proposed by Rundle et al. (2016, 2017) to detect the risk of imminent high magnitude earthquakes.

## 3. The reconstruction theorem

Previously to the reconstruction theorem (Diks, 1999) based on monofractal theory, the degree of randomness, anti-persistence or persistence of the analysed data is established taking into account the concept of the Hurst exponent (Turcotte 1997) which is defined as the exponent H of the power-law

$$\frac{R(\tau)}{S(\tau)} \propto \tau^H, \quad (1)$$

being  $R(\tau)$  the range of the different chosen segments of length  $\tau$  of a series and  $S(\tau)$  the corresponding standard deviation.  $H$  close to 0.5 implies a strong randomness of the series. Conversely,  $H$  clearly lowering or exceeding 0.5 means anti-persistence or persistence, respectively. Consequently, the Hurst exponent offers a first point of view of the behaviour of the analysed series. It has to be also remembered that the Hurst exponent has to be coincident with a specific value of the generalised Hurst exponent, obtained by means of multifractal analysis (Kantelhardt et al., 2002) applied to the same series. The analysis of the monofractal structure of a series, by means of the reconstruction theorem (Diks, 1999), permits quantifying its complex forecasting by means of the following parameters:

-The necessary minimum number of nonlinear equations governing the physical mechanism, usually referenced as correlation dimension,  $\mu(m)$ , being  $m$  the reconstruction space dimension.

-The embedding dimension,  $d_E$ , the asymptotic value of the correlation dimension, with  $m$  theoretically tending to  $\infty$ .

-The Kolmogorov entropy,  $\kappa$ , which quantifies the loss of memory of the mechanism along the analysed physical process.

The reconstruction theorem process is based on generating a set of  $m$ -dimensional space vectors using the series  $\{x(i)\}$  of data :

$$Z(i) = x_i, x_{i+1}, \dots, x_{i+m-1}, i = 1, \dots, n - m + 1 \quad (2)$$

being  $n$  the length of the series, and the definition of the correlation integral in terms of the Grassberger–Procaccia formulation (Grassberger and Procaccia, 1983a, 1983b)

$$C(m, r) = \lim_{N \rightarrow \infty} \frac{1}{N^2} \sum_{i, j=1}^N H(r - \|z(i) - z(j)\|) \quad (3)$$

$r$  being an Euclidean distance in the  $m$ -dimensional space and  $H\{\cdot\}$  the Heaviside function. The correlation integral can be rewritten as

$$C(m, r) = A_m e^{-mk} r^{\mu(m)} \quad (4)$$

$$\log[C(m, r)] = \log(A_m) - mk + \mu(m) \log(r) \quad (5)$$

being  $K$  the Kolmogorov entropy exponent and  $A_m$  and  $\mu(m)$  the correlation amplitude and the mentioned correlation dimension for every reconstruction dimension  $m$ . A confident quantification of  $\mu(m)$  for every reconstruction dimension has

to be carefully computed, avoiding a very flat evolution of  $C(m,r)$  for small values of  $r$ , caused by the lacunarity (Turcotte, 1997), and the saturation of  $C(m,r)$  for the highest values of  $r$ . With respect to the quantification of the Kolmogorov entropy,  $K$ , by using equation 5, naming  $\mu(m)$  the term  $\log\{C(m, r)\} - \mu(m)\log(r)$ , and after obtaining  $\alpha(m)$ , the equation 5 becomes

$$\alpha(m) = \log(A_m) - mK \quad (6)$$

Equation 6 is characterised by an almost constant value of  $\log(A_m)$  for high reconstruction dimensions  $m$ . Consequently, a very accurate value of the Kolmogorov coefficient  $K$  could be obtained by a linear regression in terms of equation 6, but only for the mentioned set of the highest reconstruction dimensions  $m$ . The same set of  $m$ -dimensional space vectors permits the computation of the Lyapunov exponents  $\lambda_i$ , ( $i=1, \dots, m$ ) (Eckmann et al., 1986; Stoop and Meier, 1988; Wiggins, 2003) which quantify the intensity of chaotic behavior of a system specially the first  $\lambda_1$  exponent) when the results, forthcoming volcanic emissions at the present case, have been estimated by means of some forecasting algorithm. Additionally, the Kaplan-Yorke dimension,  $D_{KY}$ , (Kaplan and Yorke, 1979)

$$D_{KY} = l_0 + \frac{1}{|\lambda_{l_0+1}|} \sum_{j=1}^{l_0} \lambda_j \quad (7)$$

,with  $l_0$  the maximum number of Lyapunov exponents in decreasing order accomplishing

$$\lambda_1 + \lambda_2 + \dots + \lambda_{l_0} \geq 0, \quad (8)$$

quantifies the fractal dimension of the nucleus around of which the consecutive  $m$ -dimensional vectors describe the corresponding orbital trajectories. In short, the highest the value of  $D_{KY}$ , more complex will be to establish the forthcoming value of the analysed physical problem.

## 4. Results

### 4.1 The Hurst exponent.

The results of the Hurst exponent for the whole series and the six data segments is described in Figures 5a and 5b, being obtained a clear sign of persistence for the complete series of Vulcanian explosions, with  $H$  exceeding a value of 0.7, a moderate persistence for the first, second, fourth and fifth segments, a smooth increase of  $H$  from the fifth to the sixth segment and a clear persistence ( $H > 0.70$ ) for the third segment. This third clear persistence is detected for a data segment including the second highest energy emission (Figures 2a and 2b). Conversely, the lowest Hurst exponents for the

first, second, fourth and fifth segments are characterised by more moderate emissions of energy. Finally, the increase of H  
155 for the sixth segment could be caused by the imminence of the highest energy, an emission immediately after this data  
segment. A more detailed evolution of the Hurst exponent is described by the 21 moving windows (Figures 6a and 6b), being  
detected that the increasing of the persistence for the 6 first moving windows is stabilised for the other 15 windows with  
notable signs of persistence, with H varying from 0.72 to 0.76. In short, the factor of persistence from the point of view of  
the Hurst exponent suggests a certain facility of forecasting algorithms, being remarkable that after the first 1000 emissions,  
160 (beginning of the five moving window), the highest persistency with some fluctuations is achieved and the highest emissions  
are included in these windows.

#### 4.2 Embedding dimension.

With respect to the embedding dimension, Figure 7 illustrates five examples of the first segment of 1000 recorded emissions,  
165 where the slope,  $\mu(m)$ , of the  $\log_{10}\{C(r)\}$  with respect to  $\log_{10}\{r\}$  monotonically increases for an interval of r, being then  
described the asymptotic evolution of these slopes towards the definitive embedding dimension  $d_E$ . The embedding  
dimensions for the 21 moving windows and the 6 segments are respectively summarised in Table 1 and Table 2. By  
remembering that this dimension defines the minimum number of non-linear differential equations associated with the  
physical process, the most complex segments from the mathematical point of view would be the first, second, third and sixth,  
170 being not so complex the fourth and fifth ones. Nevertheless, the discrepancies when comparing the different segments are  
not excessive, given that 9 or 10 differential equations would be sufficient to analyse every one of the 6 segments. A quite  
different evolution of  $d_E$  is obtained for the 21 moving windows (Table 1), with dimensions approximately varying from 9.5  
to 6.9. Between the 11<sup>th</sup> and 13<sup>th</sup> moving windows  $d_E$  diminishes (a more simplified mathematical structure should be  
assumed for these volcanic emissions) and for the remaining windows (14<sup>th</sup> -21<sup>th</sup>) their mathematical structures complexity  
175 **return** to moderate values (7.3 – 7.6).

#### 4.3 The Kolmogorov entropy.

The obtained values of the Kolmogorov entropy exponent, based on equation 6 and summarised in Tables 1 and 2, are also  
illustrated with some examples (Figure 8). In these four examples, the “loss of memory” of the physical mechanism is quite  
180 similar for the 6<sup>th</sup> segment and the 10<sup>th</sup> moving window, with values of K which could complicate a bit more the forecasting  
processes, in comparison with the previous 5<sup>th</sup> segment and the 13<sup>th</sup> moving window. In spite of these discrepancies with  
respect to the “loss of memory” for the different segments and moving windows, they are quite similar in many cases, being  
only remarkable two examples of extreme minimum (fifth segment,  $K = 0.258$ ) and extreme maximum (10<sup>th</sup> moving window,  
 $K = 0.410$ ). Consequently, the “loss of memory”, making complex the forecasting process, would not affect in the same way  
185 to all the volcanic explosive emissions.

#### 4.4 The Lyapunov exponents.

190 A right computation of the Lyapunov exponents needs an iterative process, with the aim of minimising the final uncertainty on every exponent. At the present computations, 975 iterations have been good enough to obtain the first fifteen exponents with very small oscillations at the end of the iterative process. An example of this process is shown in Figure 9, where are described the evolution of the exponents for the third segment of emissions up to  $\lambda_{15}$ . A higher number of exponents is not necessary by two questions. First, the possible errors or uncertainties on forecasting processes could be specially associated with the first Lyapunov exponents. Second, observing the evolution at the end of iterations of the exponents of Figure 9, the Kaplan-Yorke dimension can be computed without the necessity of Lyapunov exponents exceeding dimension 15.

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The results are summarised in Table 3, showing the mean and standard deviation for every one of the first ten Lyapunov exponents, obtained for the 21 moving windows and the six data segments after 975 iterations of the corresponding computational algorithm to obtain accurate and confident values. First, in agreement with the results exposed in the mentioned table, every one of the  $\lambda_i$  exponents is quite similar both for segments and moving windows, bearing in mind the very similar average values and small standard deviations. Second, the first small negative Lyapunov values are always detected for  $\lambda_7$  or  $\lambda_8$ . Consequently, the information offered by the Lyapunov exponents, concerning the possible errors on forecasting, should be very similar all along the emissions, being not detected differences between data segments and moving windows. Finally, the Kaplan-Yorke dimension manifests a notable similarity for both, segments and moving window.

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205 Whereas for the six trams,  $D_{KY}$  varies from 12.51 to 12.75, the range is quite similar for the 21 moving windows, varying from 12.70 to 13.04. Consequently, the fractal dimension of the nucleus around which the consecutive m-dimensional reconstructed vectors describe the corresponding trajectories is complex (a fractal dimension exceeding 12.0). Nevertheless, this complexity becomes confined within a short interval, being quite similar for all segments and moving windows.

## 210 5. Some examples of nowcasting

The nowcasting process (Rundle et al., 2016, 2017) is based on the computation of the “natural time”, or in other words, the number of consecutive earthquakes (seismic cycle length) with magnitudes within a determined interval. In this way, the empiric cumulative distribution function, CDF, of these “natural times” would be established by the high magnitude earthquakes interrupting these seismic cycle lengths. Consequently, the nowcasting process does not exactly predict a forthcoming high magnitude, but quantify the probability of an imminent high earthquake magnitude, based on the empiric CDFs curves.

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220 A first illustrative example of the nowcasting algorithm, from the point of view of the seismic activity, is depicted in Figure 10a. It corresponds to the recorded seismic activity in Canterbury (National Earthquake Information Database, <https://www.gns.cri.nz>, years 1990-2020). In spite of the minimum seismic magnitude accomplishing the Gutenberg-Richter law (Wiemer and Wyss, 2000) should be 3.5 (Figure 10b), maximum magnitudes of 4.5, 5.0 and 5.5, as well as a minimum magnitude of 2.5, have been considered necessities to obtain a more detailed evolution of the corresponding cycle lengths.

The probability of forthcoming extreme magnitudes (7.2 and 7.8) interrupting a cycle length exceeds 80%. Consequently, the probability of an earthquake of similar extreme magnitudes should be more or less imminent if the real cycling length is approximately ranging between 100 and 1000 “natural times”, depending on the chosen maximum magnitude  $M_{max}$ .

225 Two examples of nowcasting corresponding to volcanic energy explosions are shown in Figure 10b. The first one corresponds to the volcanic activity of the third segment (Figure 2a) and the second includes the whole series of volcanic emissions. In both cases, the cycle lengths are obtained by considering a minimum and maximum level of volcanic emission of, respectively,  $10^{2.5}$  J and  $10^{4.5}$  J. Conversely to the example of seismic activity, the few extreme levels exceeding  $\log_{10}(\text{Energy}) = 8.0$  are not associated with high CDF values. Consequently, the nowcasting algorithm could be considered  
230 less effective in comparison with the seismic activity results, probably due to the mechanism of the volcanic emissions, which do not include structures such as background activity, swarms, forecastings, mainshocks and aftershocks. Nevertheless, a notable number of high  $\log_{10}(\text{Energy})$  emissions slightly smaller than 8.0, and some of them very close to the highest emissions, are associated with values of CDFs close to 0.6, and even exceeding 0.8. An example of this fact could be  
235 the third segment of emissions, where some of them are not independent but associated with the highest emission close to 8.5. In short, the nowcasting process could be also assumed as an algorithm contributing to the predictability of volcanic explosions, but perhaps not so obvious as for the case of seismic activities.

## 6. Discussion of the results

240 The results obtained by the reconstruction theorem and the possible relationships between the fractal reconstructions exponents ( $H$ ,  $K$ ,  $\lambda$  and  $d_E$ ) and changes on the volcanic emissions are summarized in Figures 11a and 11b. First of all, relevant changes on parameters such as mean, standard deviation, skewness and kurtosis (Table 4) are not detected for the different segments of volcanic emissions. Additionally, the Kolmogorov-Smirnov test (95 and 99% of probability) discards the possible Gaussian distribution of these emissions, in agreement with Figures 3a and 3b, where the Generalised logistic  
245 distribution (GLO) is assumed from the L-Skewness-Kurtosis formulation.

The Hurst exponent (Figure 11a) is characterised by a continuous increase, finally achieving oscillations close to 0.7, with evident structure of persistence since the 7<sup>th</sup> up to the 21<sup>th</sup> moving window, all of them including two high emissions,  $\log_{10}(\text{energy}) = 8.372$  and  $7.937$ . Consequently, the values of Hurst would manifest persistence (convenient for appropriate forecasting) when a high emission becomes included in the moving window. Conversely, the “loss of memory”  
250 (Kolmogorov exponent) of the physical mechanism (not convenient for good forecasting) increases up to the 10<sup>th</sup> window, notably decreasing for the rest of windows. In this case, the influence of a high emission would appear outdated in comparison with the results of the Hurst exponent. With respect to the Lyapunov exponent,  $\lambda_1$ , its changes along the moving windows are small (Figure 11a) with not very remarkable discrepancies with an average value of 0.169 and a standard deviation of 0.013 (Table 3). Consequently, the different forecasting errors on energy emissions could not be different, at  
255 least from the point of view of  $\lambda_1$ . These errors could be also consequence of the degree of complexity of the non-linear



differential equations system, quantified by the embedding dimension,  $d_E$ . Figure 11b depicts the clear reduction of this complexity after the 10<sup>th</sup> moving window, with a fast decrease up to the 13<sup>th</sup> window and values close to 7.5 after this last cited window. Similar to the evolution of the Kolmogorov entropy, close to the 10<sup>th</sup> window, an evident decrease of both fractal parameters is detected. Reinforcing this similarity, it is also noticeable the stabilised values of these two fractal parameters for the 13<sup>th</sup> and 14<sup>th</sup> windows. It is also convenient to observe that, in agreement with the very similar obtained Kaplan-Yorke dimensions (ranging from 12.5 to 13.0) for segments and moving windows, the data-vectors of high dimension  $m$  used for the reconstruction theorem, depict a very similar structure. In other words, the trajectories of these reconstructed vectors around the fractal nucleus are quite similar.

With respect to the results of the nowcasting, the return period curves (90%, 95% and 99% of extreme emissions) could be, as cited before, a strategy relatively similar. Nevertheless, the nowcasting process permits to decide minimum and maximum emission of energy levels to define the best empirical distribution of cycle lengths of “natural waiting time”, detecting in this way the probability, in percentage, of a probable imminent volcanic emission of high energy. In spite of the nowcasting method does not determine a concrete next volcanic emission, given that it is not a forecasting process, it takes into account that future high emissions will be expected with similar “natural waiting times”. Although some emissions, (Figures 10b), exceeding  $\log_{10}(\text{Energy}) = 7.5$  have probabilities close to 50-60%, probabilities close to or exceeding 80% are also obtained. In short, the nowcasting process seems to be more effective with seismic activity than volcanic emissions. Nevertheless, their results could be also compatible and complementary with algorithms forecasting the high emission of volcanic energy.

## 7. Conclusions

The fractal parameters, obtained by means of the reconstruction theorem of 6 segments and 21 moving windows of the analysed volcanic explosions in Volcán de Colima (México), as well as the nowcasting strategy, are the first steps for the application of different forecasting processes.

The results suggest that different strategies for future forecasting emissions can be applied, especially for high-energy emissions. These forecasting strategies can be based on different algorithms (Box and Jenkins, 1976; Lipton et al, 2015; Rundel et al., 2017 and Lei, 2021, among others) and multifractal analysis of moving window data (Monterrubio-Velasco et al., 2020). In spite of the uncertainties with respect to the waiting time of an emission and its corresponding energy, estimated by means of forecasting, are expected to be non-negligible, these algorithms should depict reasonably good approaches to real energy emissions bearing in mind the obtained reconstruction theory results. Additionally, the analysis of the multifractal structure is expected to be a warning factor for volcanic activities associated with high emissions of energy, quite similar to the analysis of consecutive seismic magnitudes (Monterrubio-Velasco et al., 2020) and usefully applied to analyse climatic data, thermometric and pluviometric data, (Burgueño et al., 2014; Lana et al., 2023) and also bearing in mind (Shimizu et al., 2002) the concept of multifractal complexity index, which could also contribute to detect imminent extreme volcanic emissions of energy. In short, the reconstruction theory applied in this research, together with nowcasting

290 and forecasting algorithms and multifractal theory could be a very important process to prevent extreme emissions of  
volcanic energy. .

### Competing interests

The contact author has declared that none of the authors has any competing interests

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### Data availability

Data sharing is not applicable to this article as no new data were created or analyzed in this study.

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**Table 1. Embedding dimension and Kolmogorov coefficient for the 21 moving windows.**

<b>MW</b>	<b>1</b>	<b>2</b>	<b>3</b>	<b>4</b>	<b>5</b>	<b>6</b>	<b>7</b>	<b>8</b>	<b>9</b>	<b>10</b>	<b>11</b>
<b>d<sub>E</sub></b>	9.356	9.309	9.195	9.037	9.287	9.237	9.24 6	9.044	9.36 1	9.45 6	8.89 8
<b>K</b>	0.362	0.363	0.355	0.357	0.354	0.362	0.38 3	0.373	0.39 1	0.41 0	0.39 7
<b>MW</b>	<b>12</b>	<b>13</b>	<b>14</b>	<b>15</b>	<b>16</b>	<b>17</b>	<b>18</b>	<b>19</b>	<b>20</b>	<b>21</b>	
<b>d<sub>E</sub></b>	7.188	6.910	7.277	7.270	7.372	7.385	7.49 5	7.550	7.46 7	7.49 5	
<b>K</b>	0.297	0.266	0.279	0.286	0.281	0.284	0.29 5	0.297	0.29 1	0.28 7	

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**Table 2. Embedding dimension and Kolmogorov coefficient for the 6 segments of volcanic emissions.**

<b>SEGME NT</b>	<b>1</b>	<b>2</b>	<b>3</b>	<b>4</b>	<b>5</b>	<b>6</b>
<b>d<sub>E</sub></b>	9.155	8.932	8.727	7.952	8.101	9.340
<b>K</b>	0.359	0.338	0.377	0.375	0.258	0.398

**Table 3. Mean and standard deviation for the first ten Lyapunov exponents, after 975 iterations.**

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<b>SEGME NT</b>	<b>λ1</b>	<b>λ2</b>	<b>λ3</b>	<b>λ4</b>	<b>λ5</b>	<b>λ6</b>	<b>λ7</b>	<b>λ8</b>	<b>λ9</b>	<b>λ10</b>
<b>&lt;λ&gt;</b>	<b>0.156</b>	<b>0.105</b>	<b>0.074</b>	<b>0.053</b>	<b>0.032</b>	<b>0.015</b>	- <b>0.001</b>	- <b>0.018</b>	- <b>0.039</b>	- <b>0.062</b>
<b>σ(λ)</b>	<b>0.011</b>	<b>0.004</b>	<b>0.005</b>	<b>0.004</b>	<b>0.003</b>	<b>0.003</b>	<b>0.004</b>	<b>0.005</b>	<b>0.003</b>	<b>0.004</b>

**Table 4. Basic characteristics of the energy emissions (logarithm of energy), for the whole database, the 6 segments of 1,000 elements and the last segment of 182 elements. The empirical results of the Kolmogorov-Smirnov, K-S, test are compared with the significance levels of 95 and 99%, KS\_0.05 and KS\_0.01, corresponding to the Gaussian distribution.**

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<b>SEGME NT</b>	<b>0001- 6182</b>	<b>0001- 1000</b>	<b>1001- 2000</b>	<b>2001- 3000</b>	<b>3001- 4000</b>	<b>4001- 5000</b>	<b>5001- 6000</b>	<b>6001- 6182</b>
<b>MAXIM UM</b>	8.903	8.187	7.574	8.372	7.937	7.650	7.667	8.903
<b>MEAN</b>	6.610	6.578	6.563	6.641	6.663	6.558	6.633	6.730
<b>ST.DEV</b>	0.283	0.233	0.225	0.313	0.310	0.242	0.292	0.470
<b>SKEWN ESS</b>	1.394	1.252	0.931	1.267	1.035	1.144	0.834	2.085
<b>KURTOS IS</b>	3.462	3.695	0.780	2.033	1.016	1.328	0.110	5.394
<b>K-S TEST</b>	0.102	0.081	0.083	0.106	0.091	0.102	0.095	0.154
<b>K-</b>	0.017	0.043	0.043	0.043	0.043	0.043	0.043	0.100

<b>S 0.05</b>								
<b>K- S 0.01</b>	0.021	0.051	0.051	0.051	0.051	0.051	0.051	0.120

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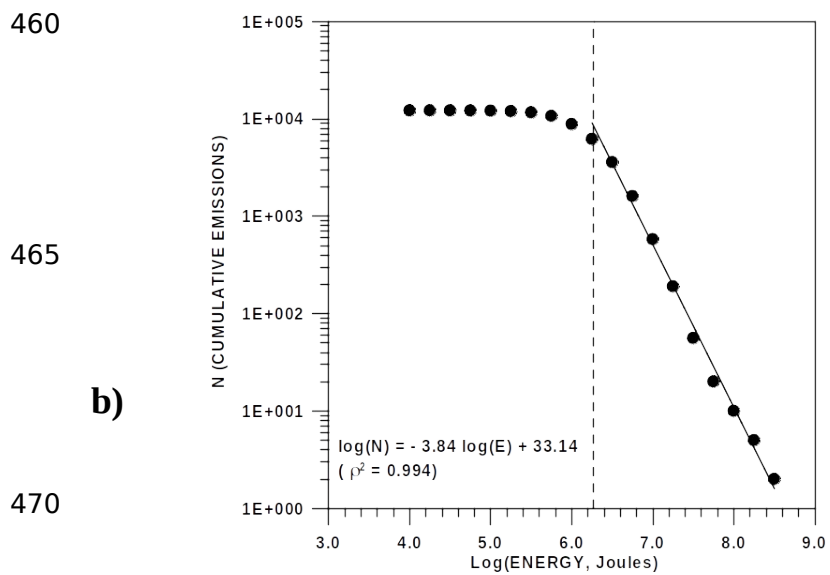
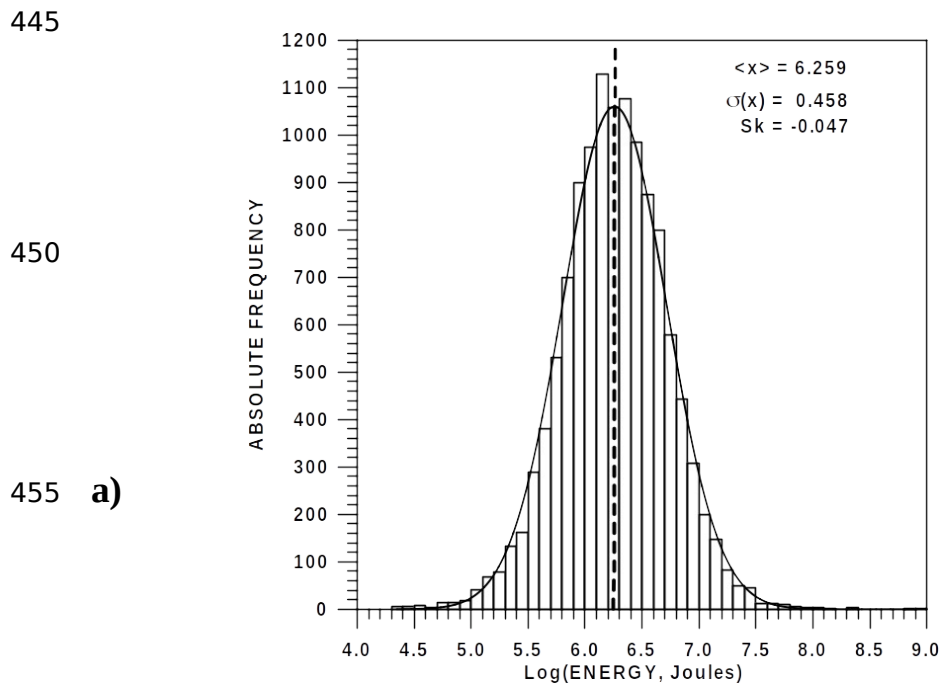
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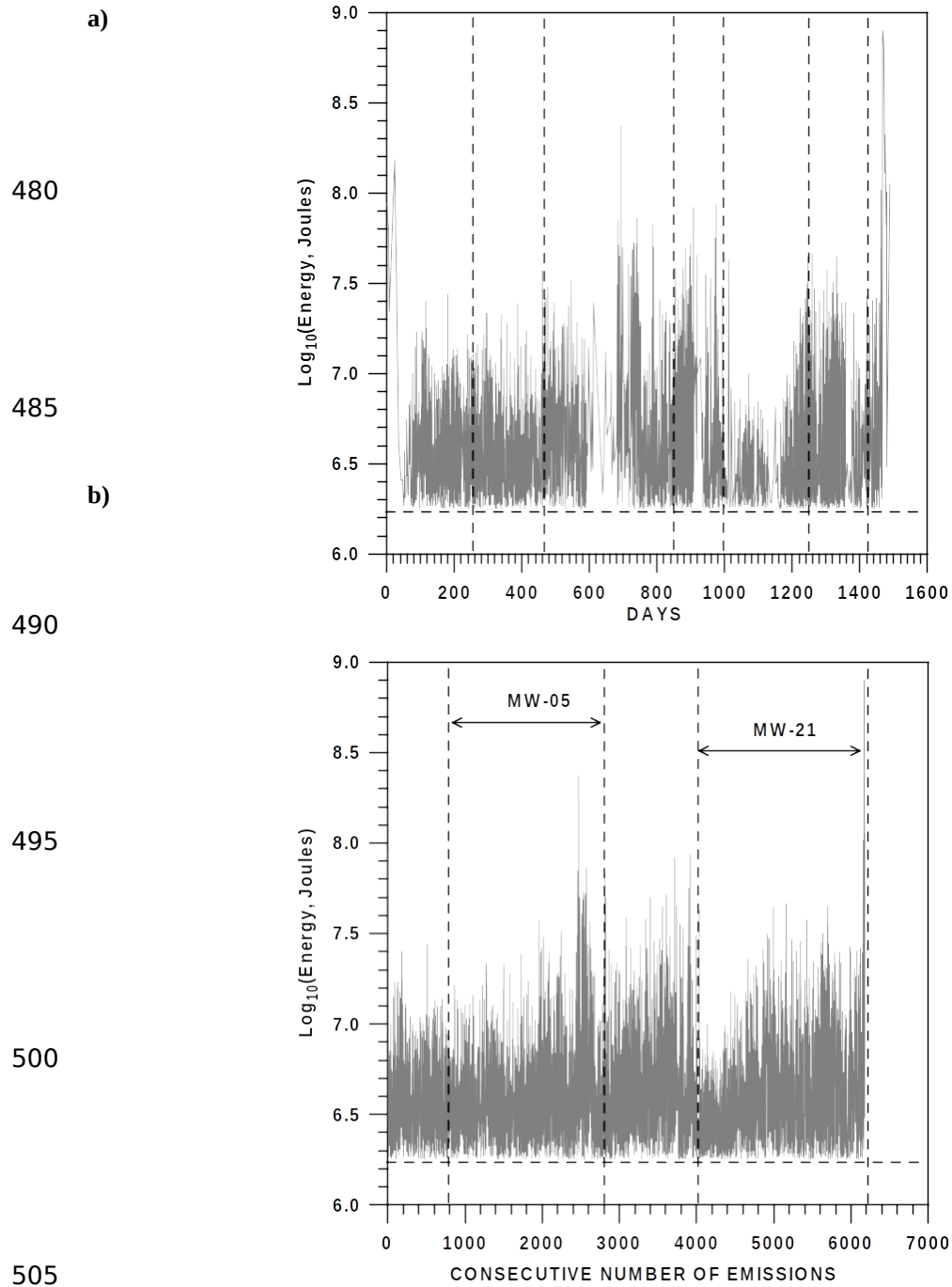
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**Figure 1.** a) Histogram of the volcanic emission energies and b) logarithm of energies accomplishing the Gutenberg-Richter law.

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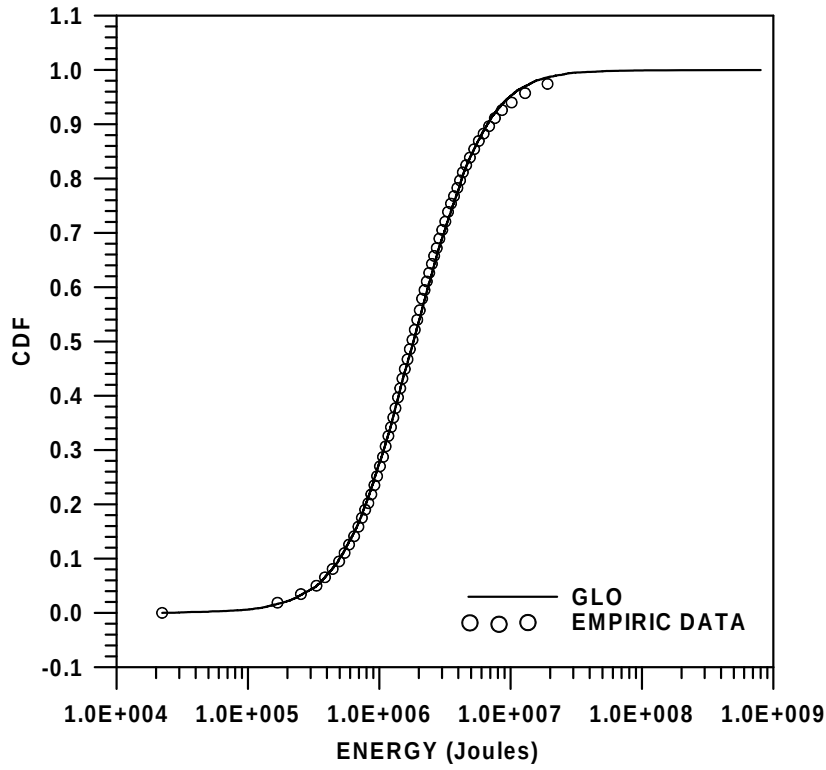


**Figure 2.** a) Evolution of the volcanic energy emission accomplishing the Gutenberg-Richter law. Vertical dashed lines define the segments (intervals of 1000 data) b) Two examples of dataset defined by moving windows of amplitude 2 thousand elements and shift of 200 positions.



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a)



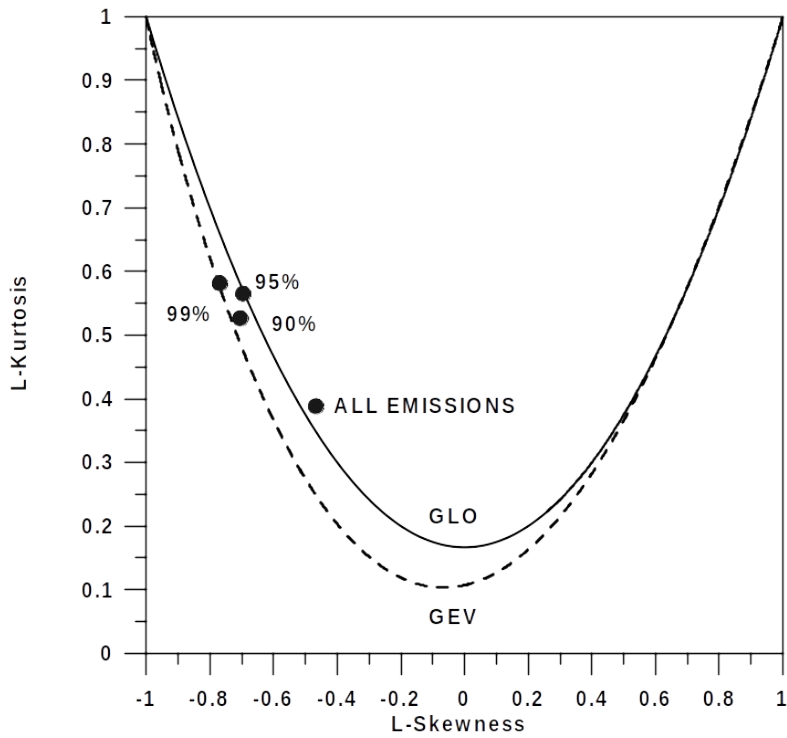
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b)



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**Figure 3.** a) Cumulated distribution function of the effusive volcanic energy emissions, well fitted to the GLO (generalised logistic distribution). b) The emissions equalling to or exceeding 90%, 95% and 99% could be associated with the GEV (generalised extreme values distribution) in agreement with the L-Skewness / L-Kurtosis diagram. The theoretical cumulated distribution GLO of the volcanic emissions is also confirmed by means of the mentioned diagram.

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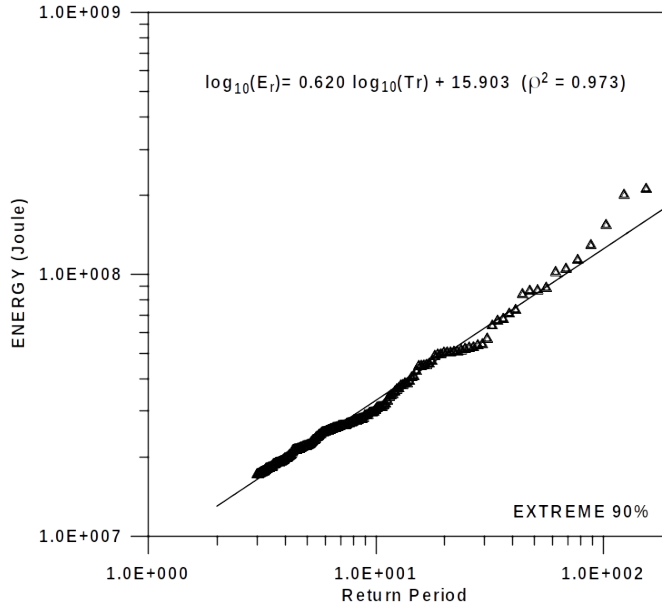
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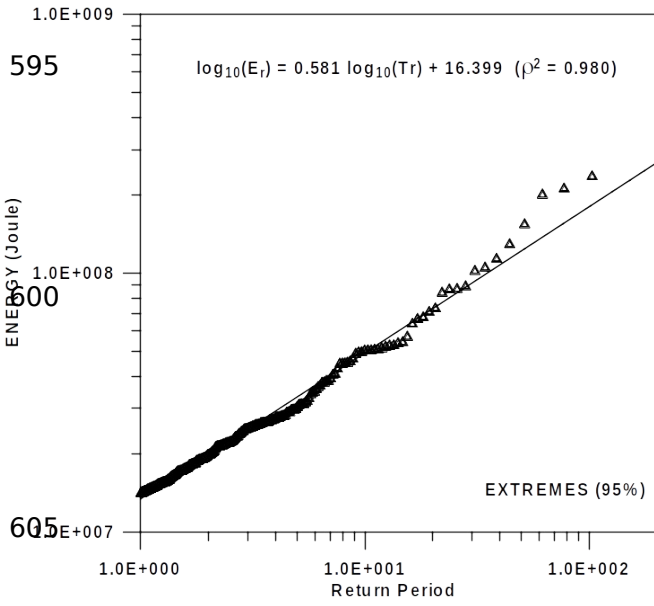
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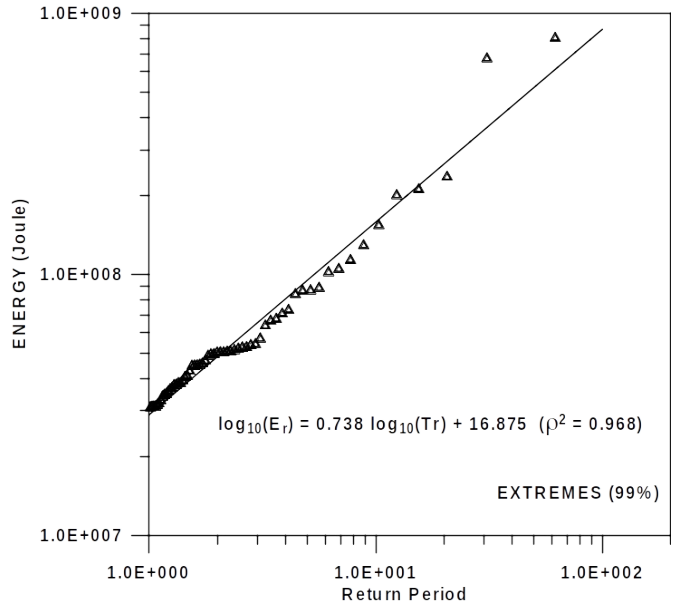
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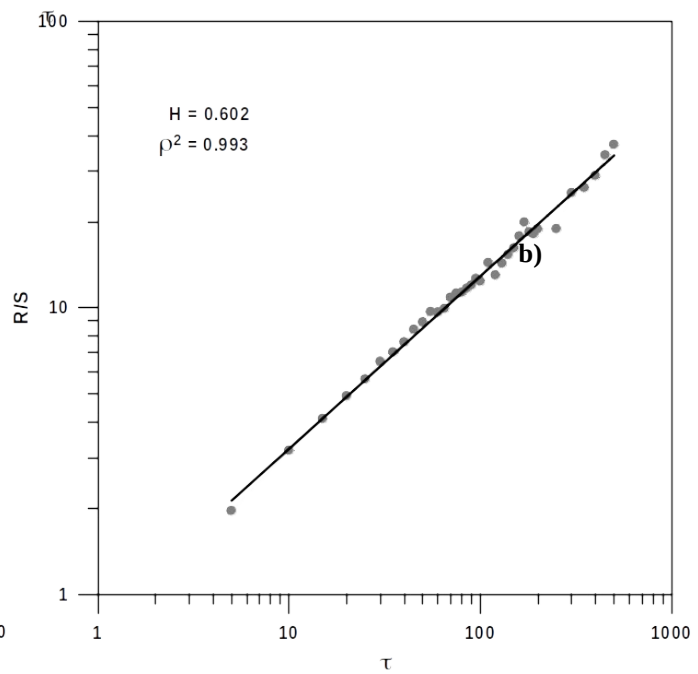
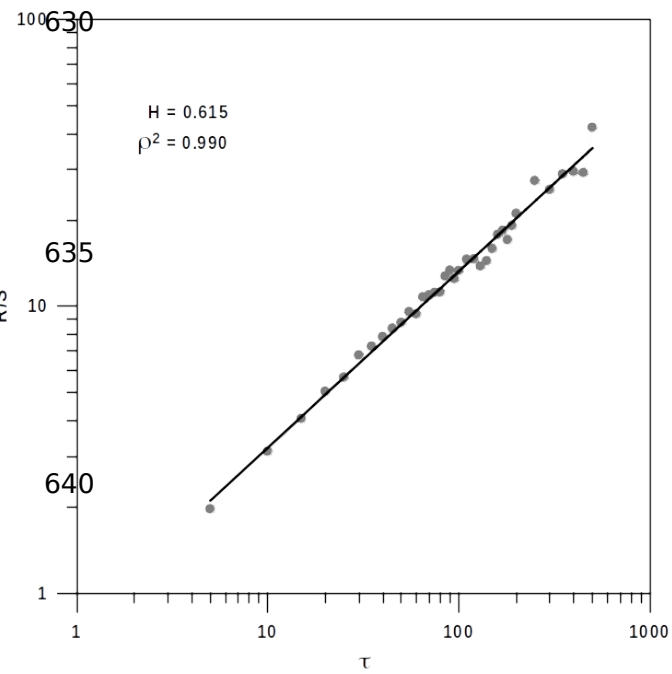
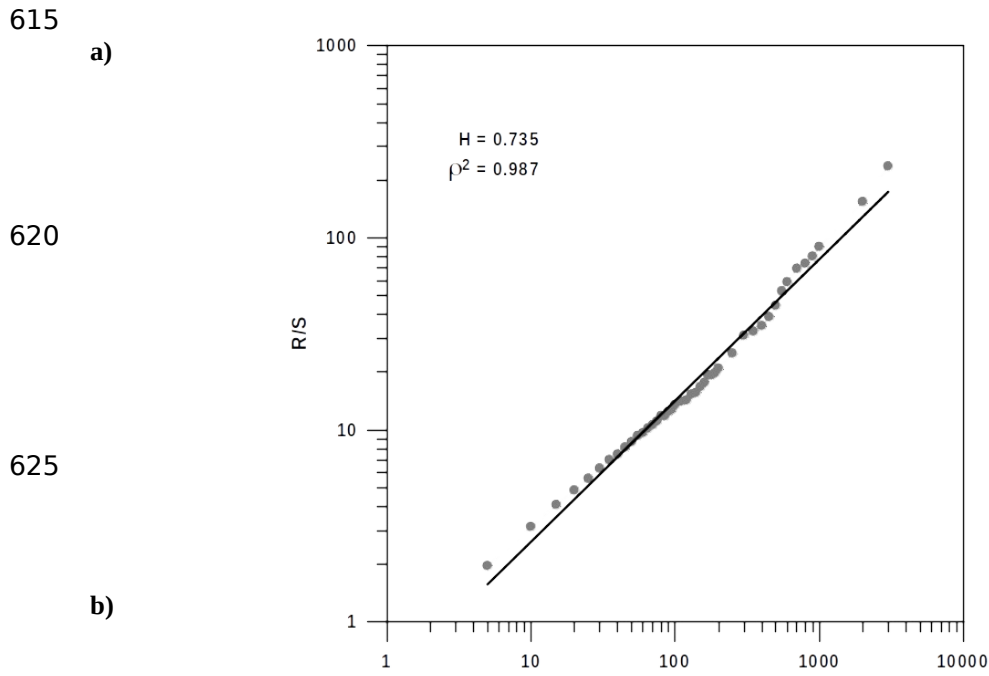
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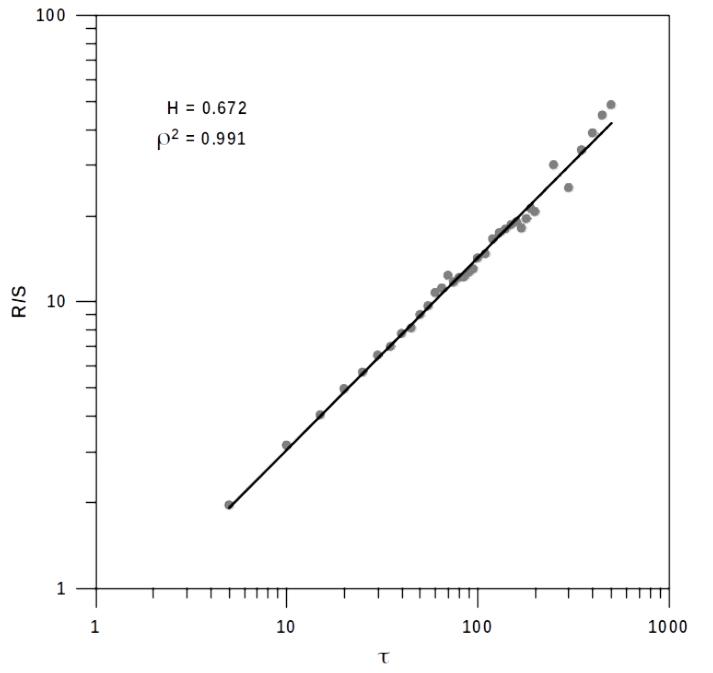
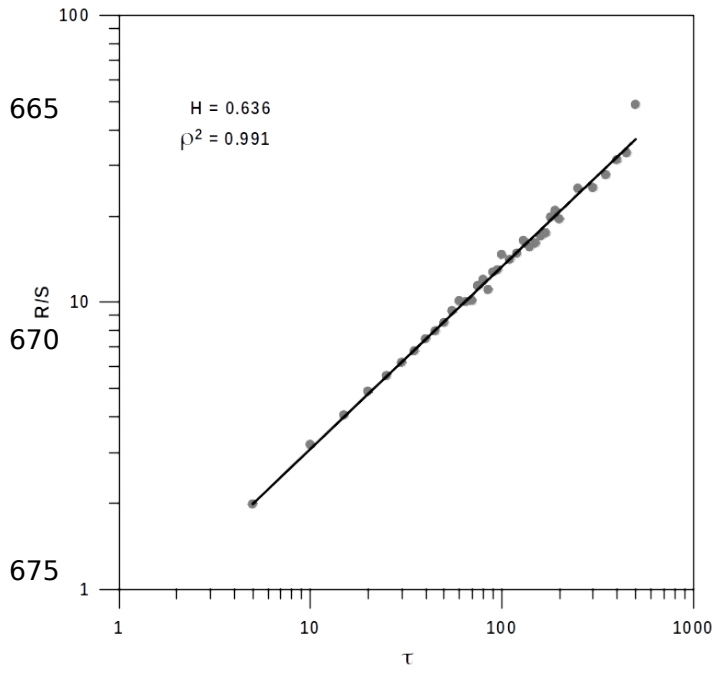
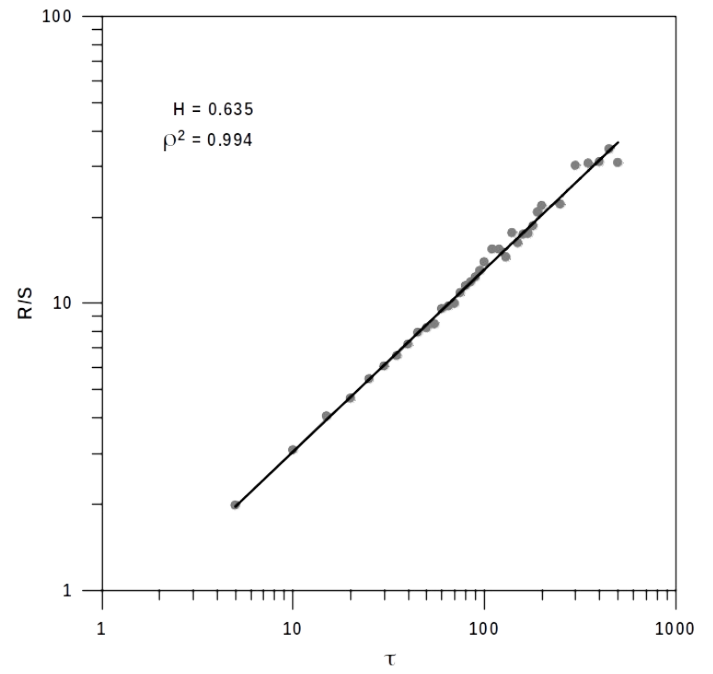
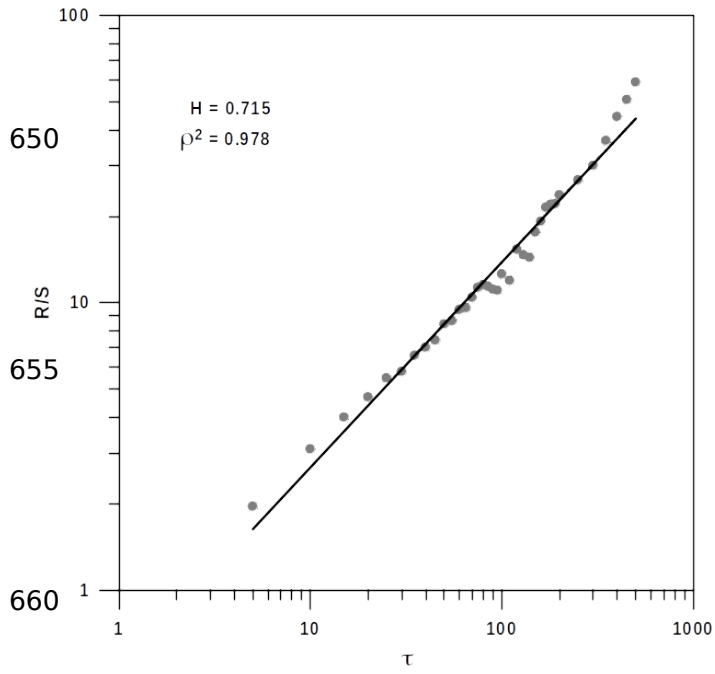
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**Figure 4.** Return period curves (90, 95 and 99%) of extreme emissions.



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(Figure 5 continue)



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**Figure 5.** Some examples of the Hurst exponent for a) the whole series of effusive emissions of energy accomplishing the Gutenberg-Richter law and b) the same series fragmented on six trams of equal number of records. In agreement with the definition of equation (1),  $R/S$  has not units and  $\tau$  represents the different lengths (number of data) of the segments considered.

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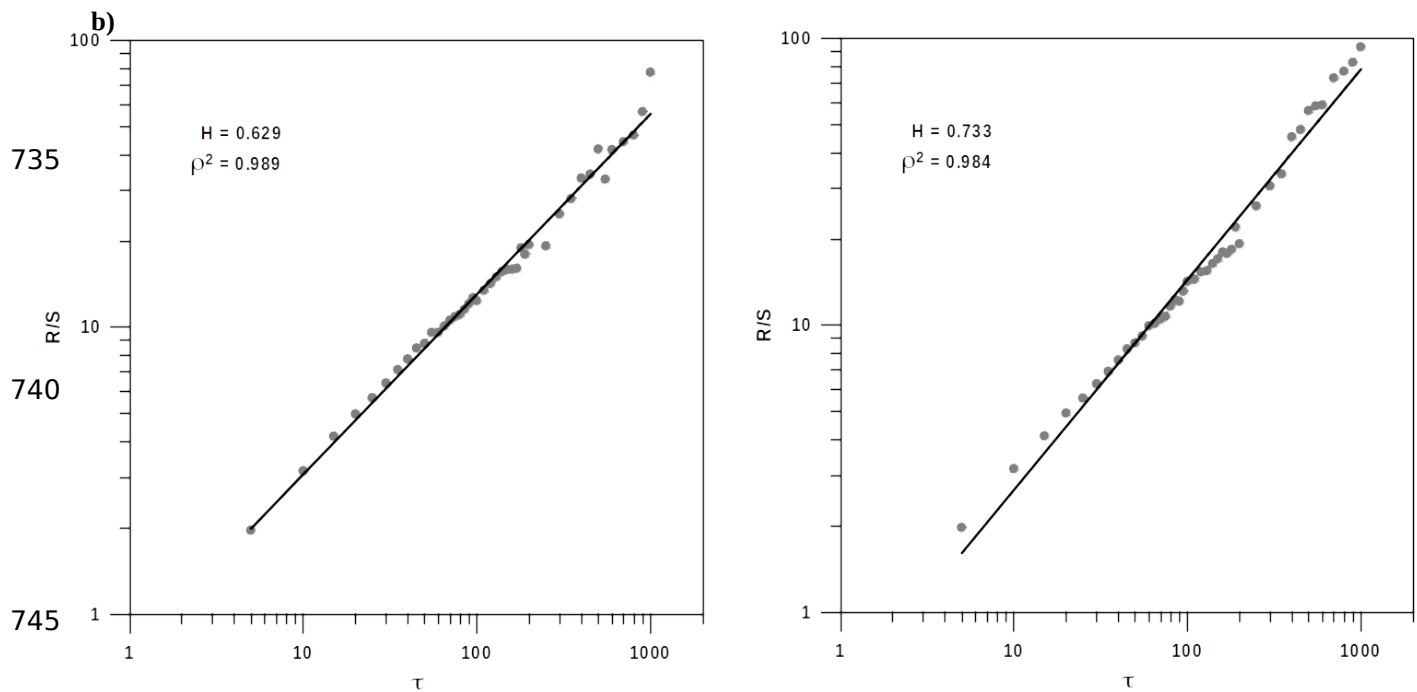
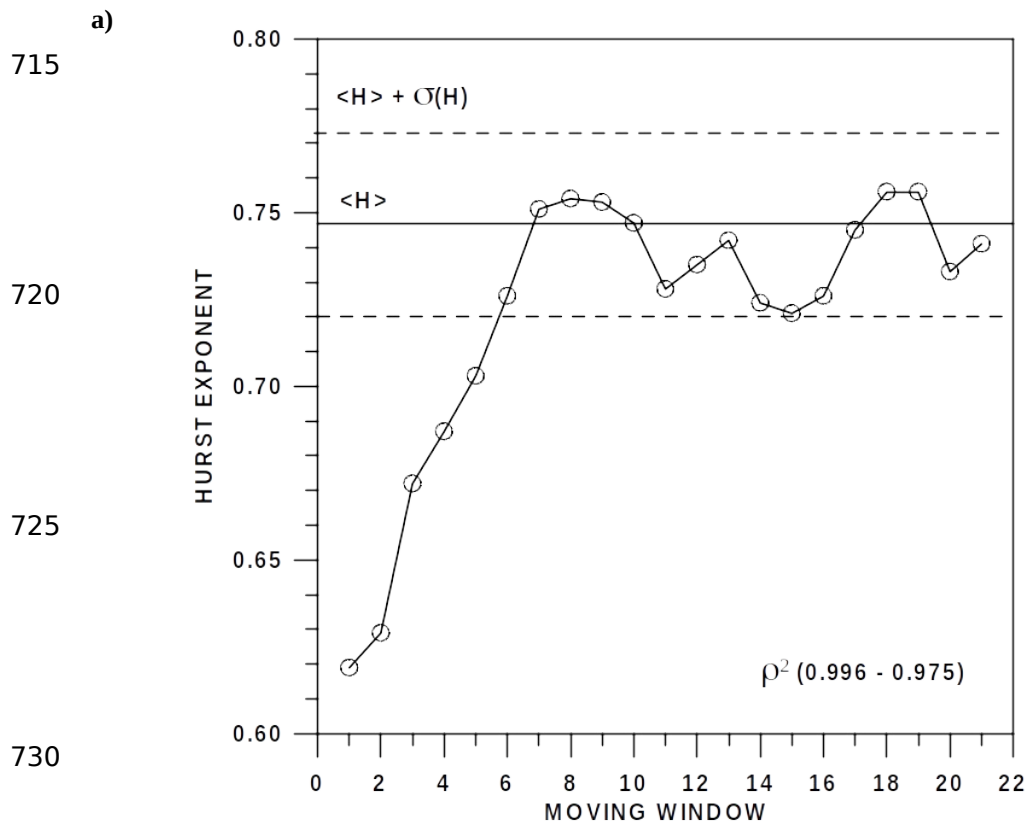
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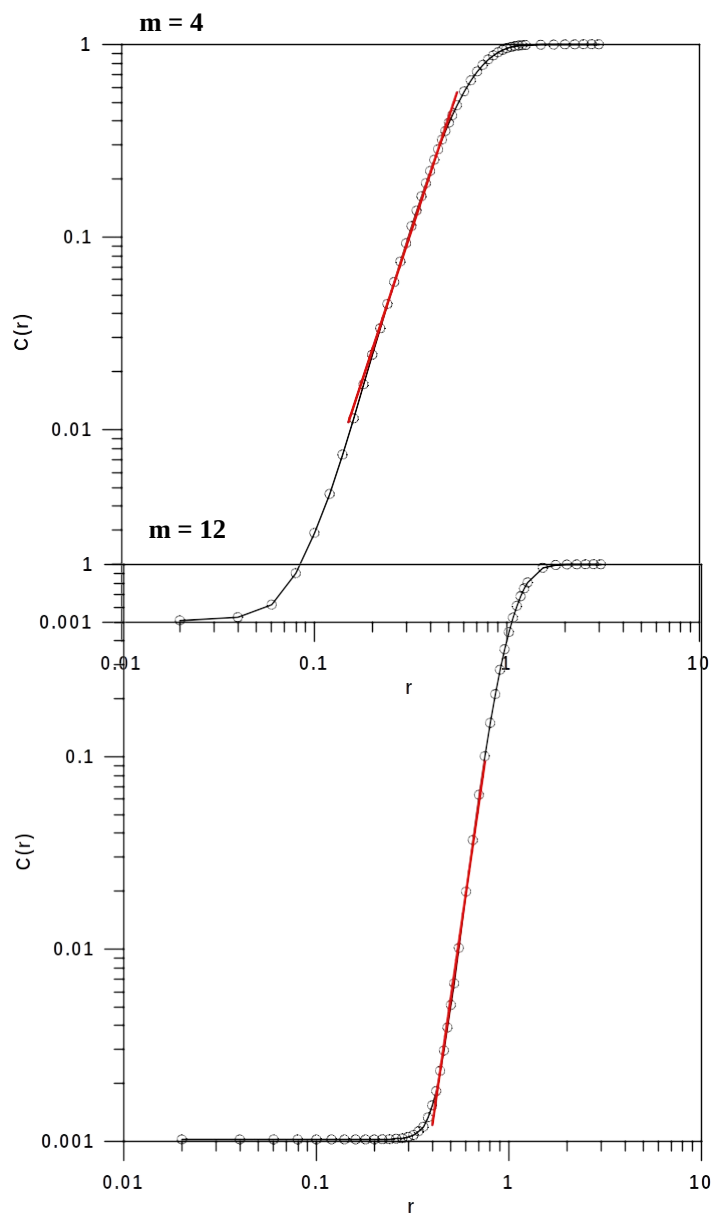
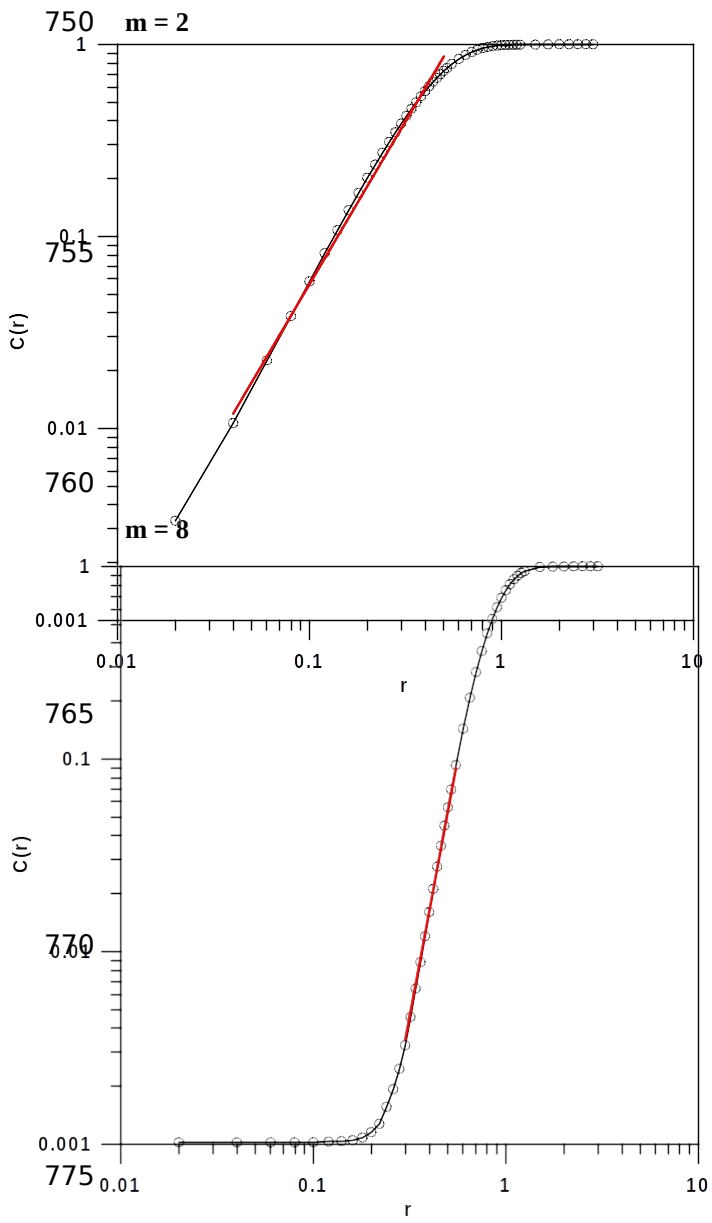
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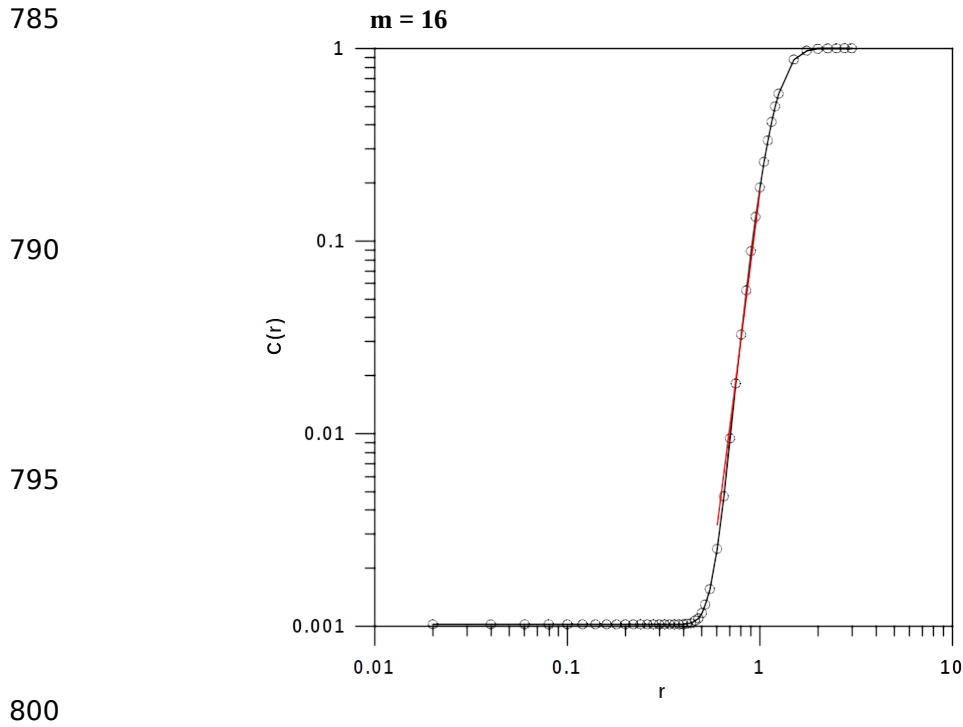
**Figure 6.** a) Evolution of the Hurst exponent for the 21 moving windows and b) two examples for windows 2 and 6.



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(Figure 7 continue)



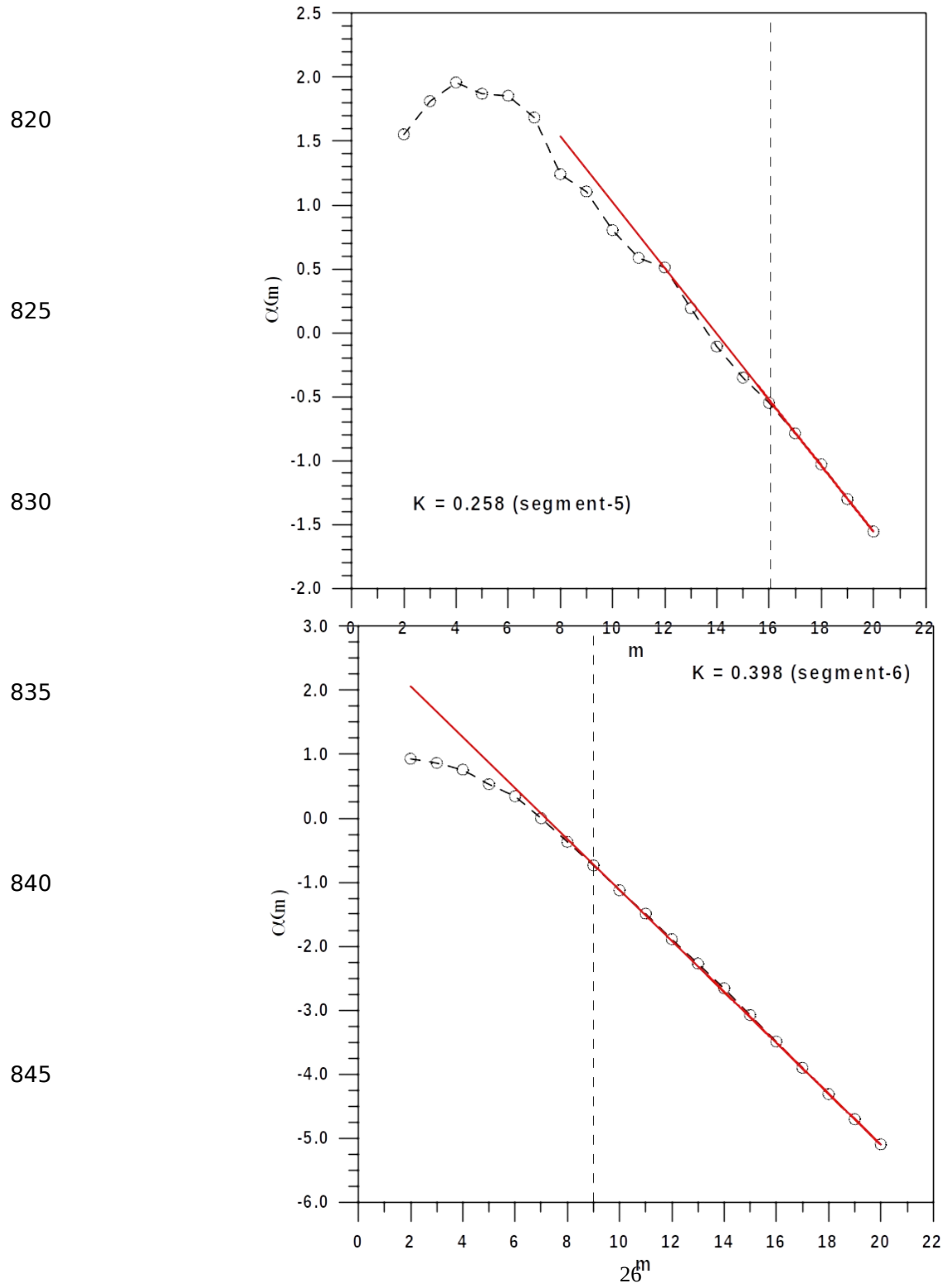
**Figure 7.** An example of the evolution of embedding dimension (first segment of 1000 elements) for reconstruction dimensions  $m = 2, 4, 8, 12, 16$ . The straight red line represents the interval of  $r$  values for which can be determined every one of the embedding dimensions.

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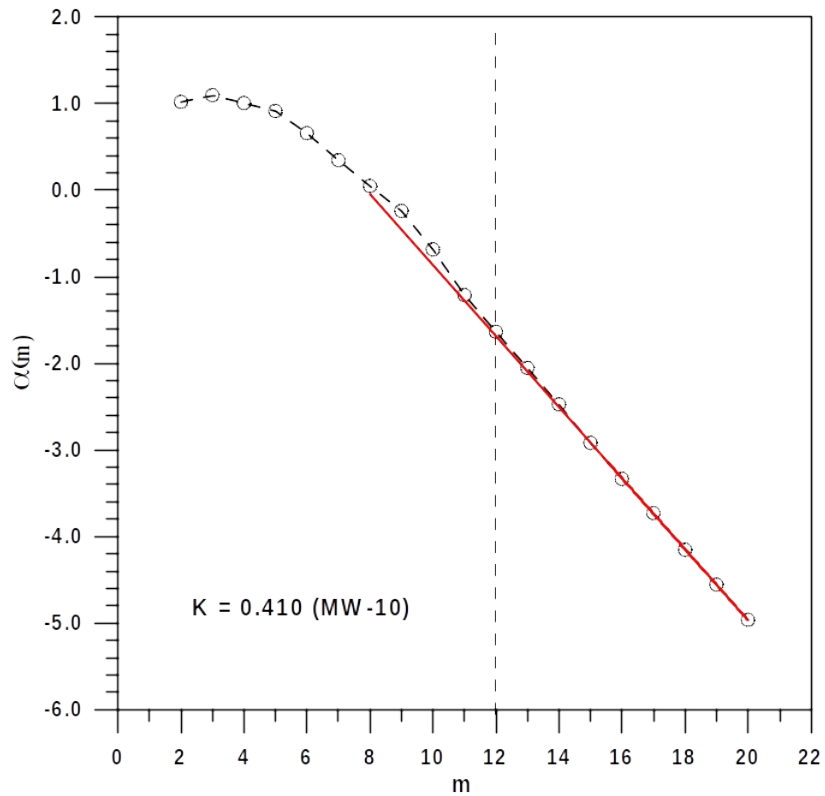
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a)



850 (Figure 8 continue)



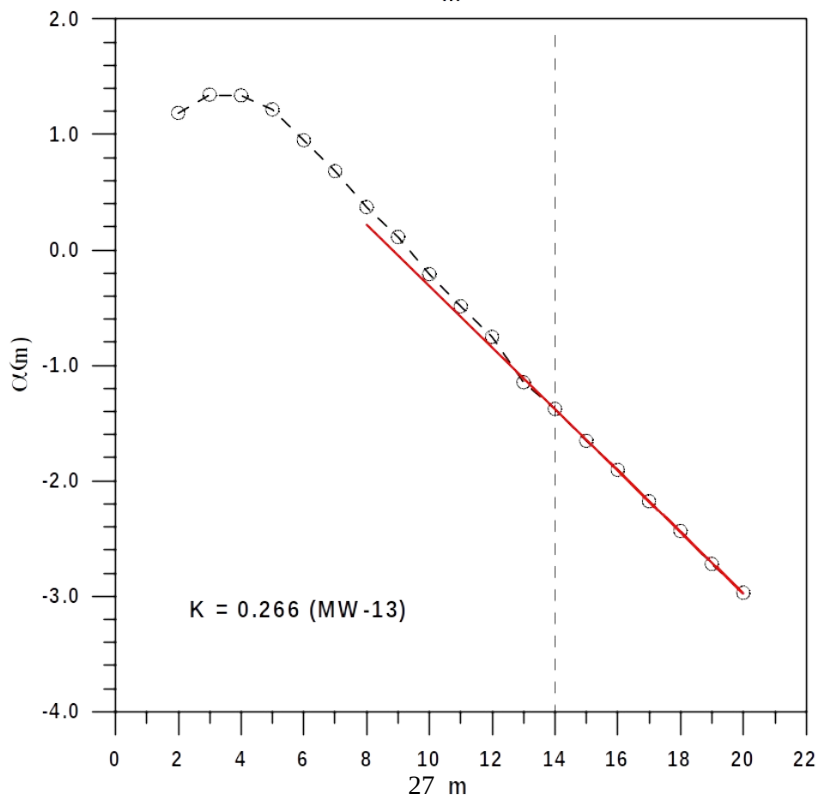
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(b)

885 **Figure 8.** a) Two examples of Kolmogorov entropy exponents,  $K$ , for two segments of 1000 elements and b) for two  
examples for moving windows. The vertical dashed lines define the parameter  $m$  interval used to determine the Kolmogorv  
entropy for every data segment or moving window.

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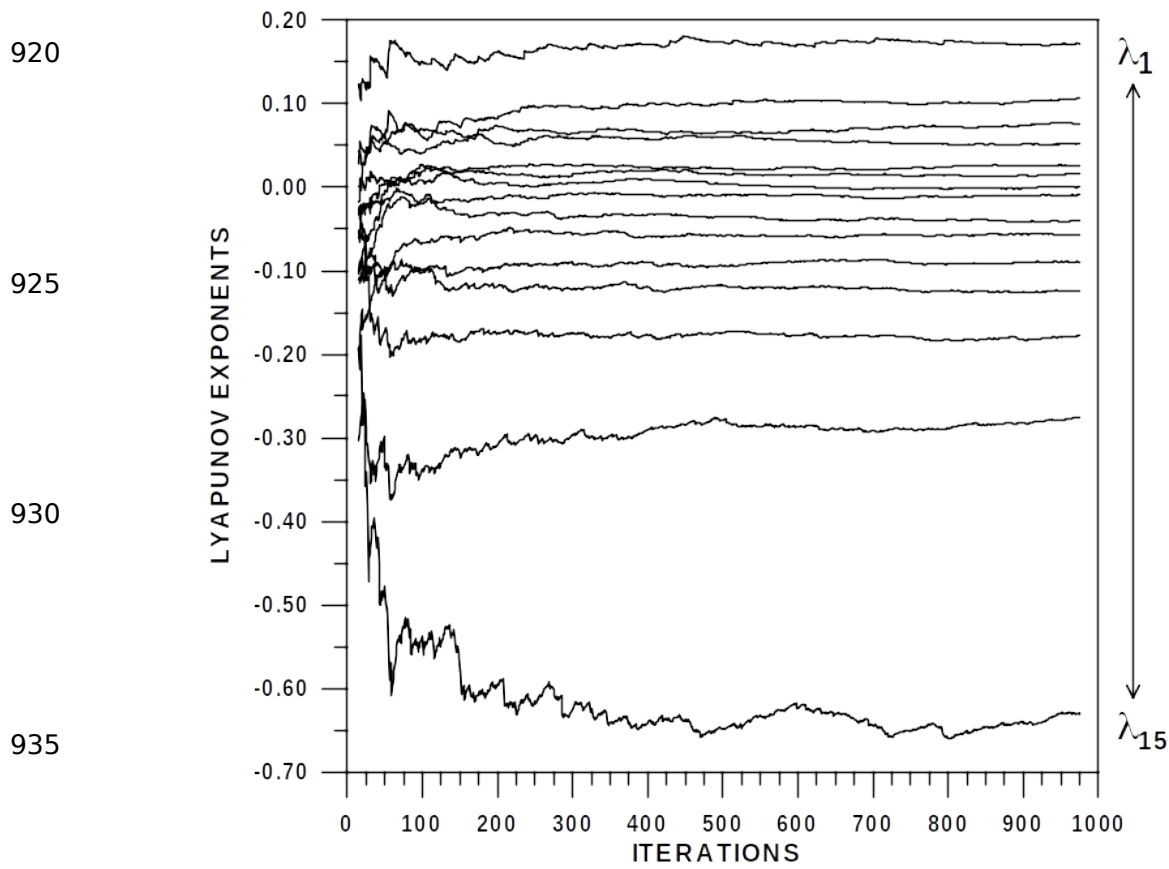
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**Figure 9.** Fifteen Lyapunov exponents for the third segment of the effusive-explosive volcanic emissions.

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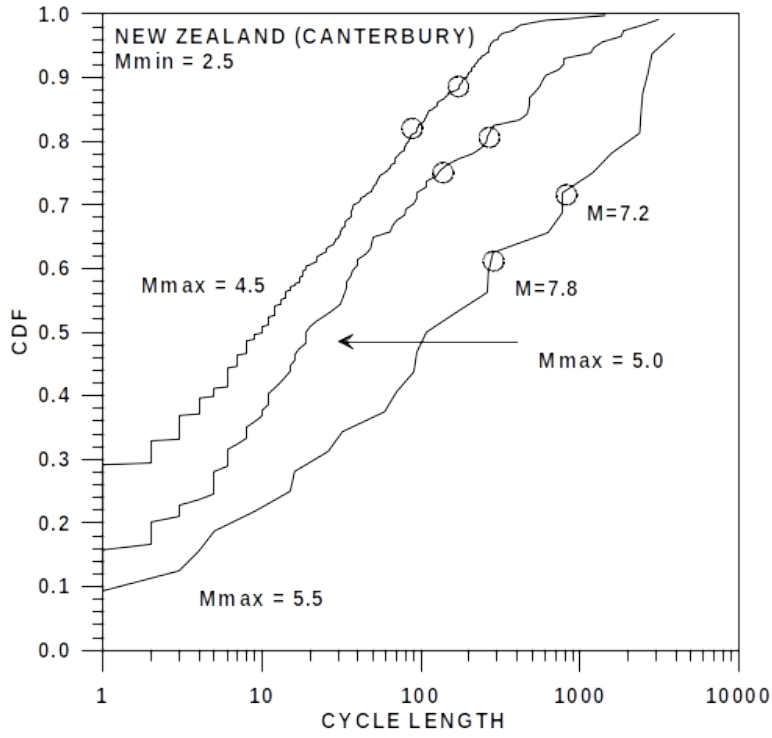
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a)

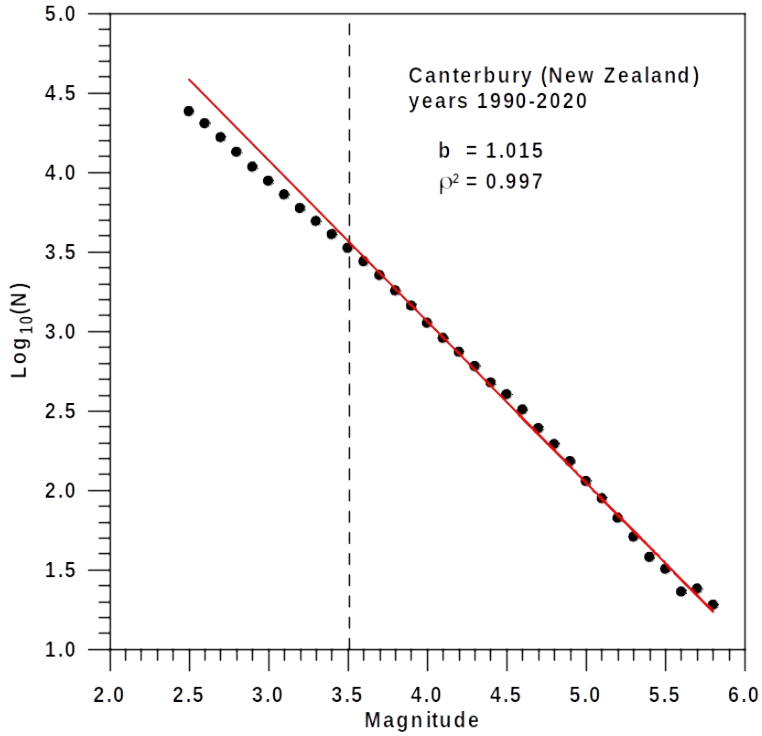


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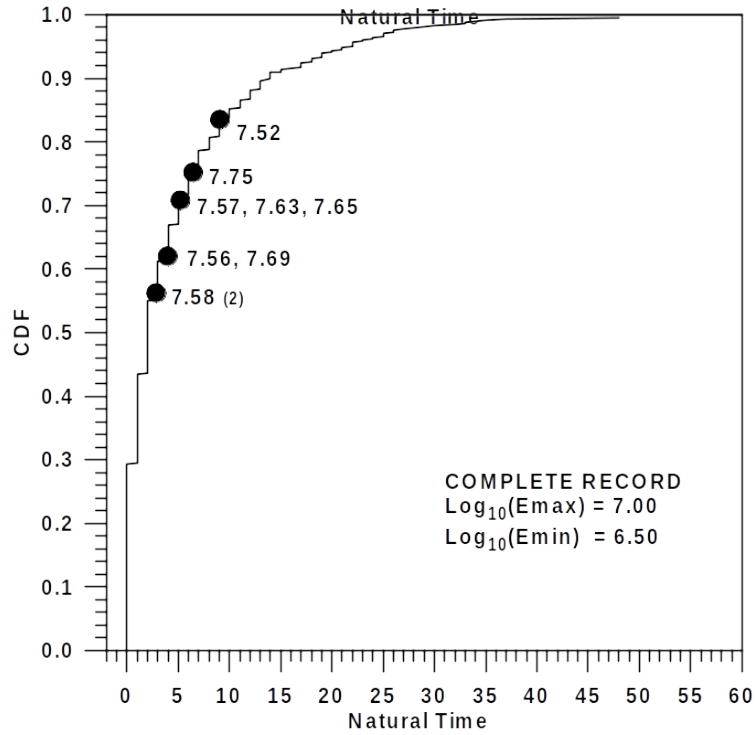
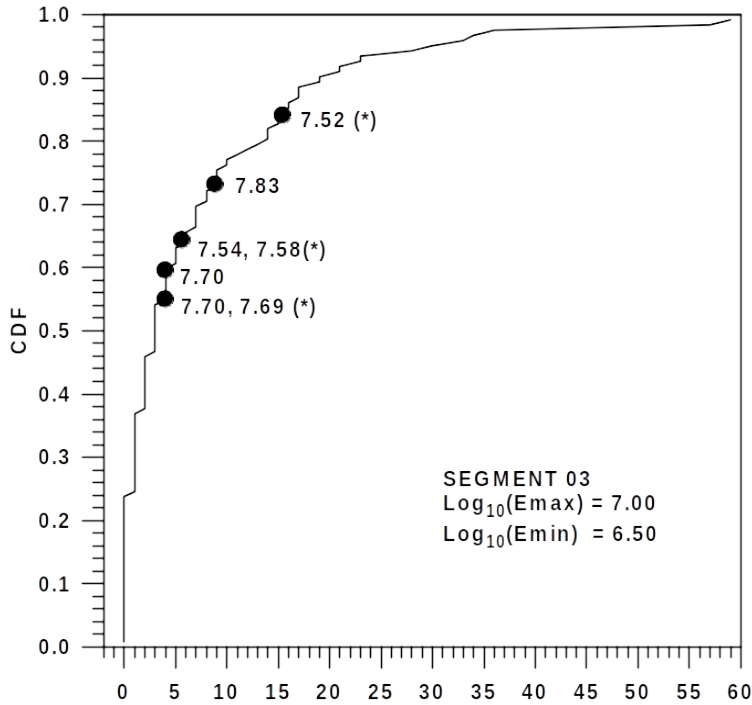
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b)



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Figure 10 (continue)



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995

1000

c)

1005 **Figure 10.** A comparison of nowcasting results for a) the seismic activity of Canterbury (New Zealand), years 1990-2021,  
and b) the corresponding Gutenberg-Richter plot and c) the energy of the volcanic emissions of Colima (México), years  
2013-2015, bearing in mind the SEGMENT 03 of volcanic emissions and all the volcanic record. The extreme three volcanic  
records designed by an asterisk (Figure 10b) are the same detected in the Figure 10c.

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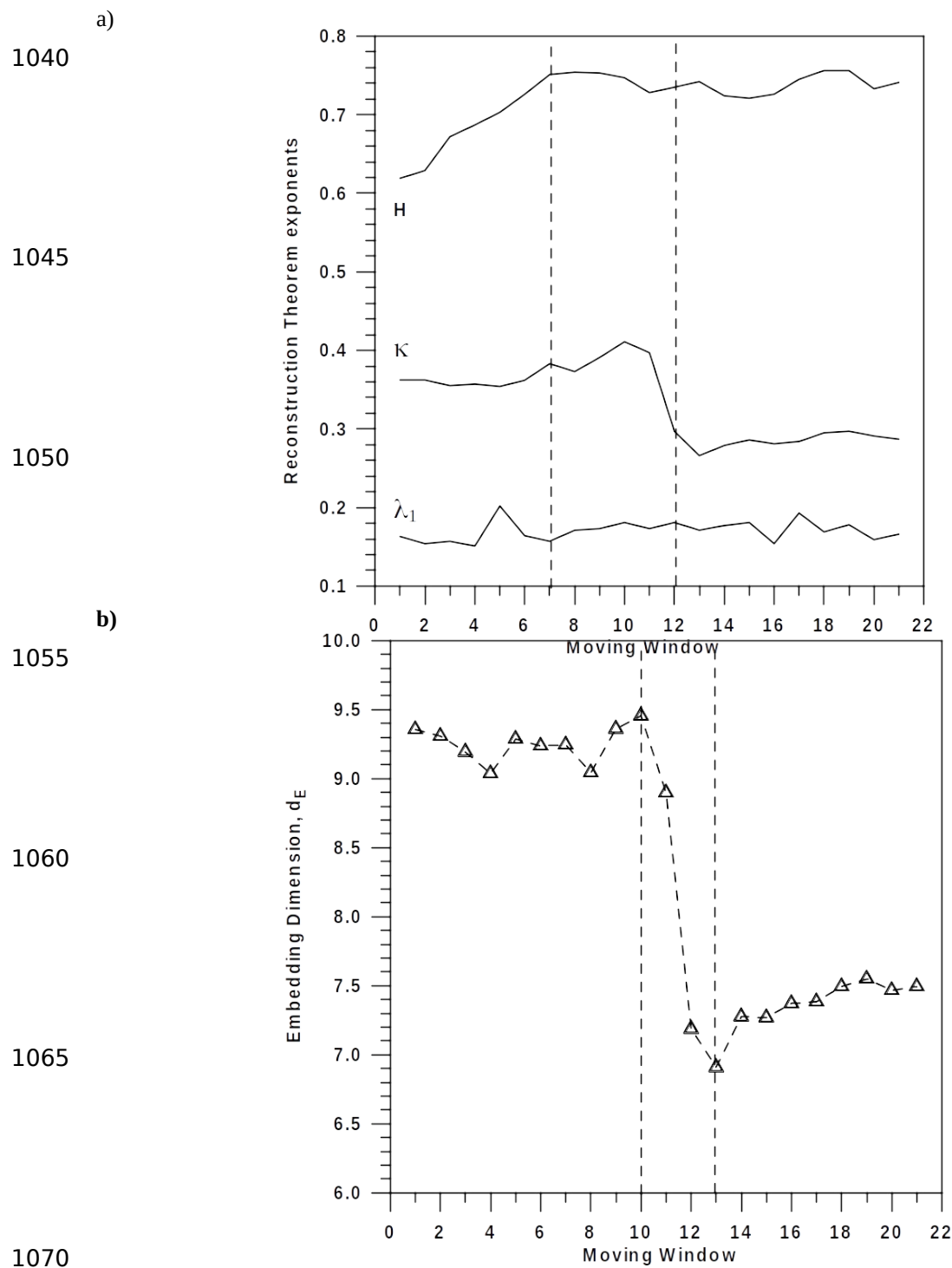
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**Figure 11.** a) Evolution of the Hurst exponent,  $H$ , the Kolmogorov entropy exponent,  $K$ , and the first Lyapunov exponent,  $\hat{\lambda}_1$ . b) Embedding dimension  $d_E$ , for the 21 moving windows.

1075 **Competing interests**

The contact author has declared that none of the authors has any competing interests

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101093038, ChEESE-CoE

**Data availability**

Data sharing is not applicable to this article as no new data were created or analyzed in this study.

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