

What is the cause(s) of large ozone variabilities in three megacity clusters in eastern China during 2015–2020?

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Abstract. Due to a robust emission control policy, significant reductions in major air pollutants, such as PM_{2.5}, SO₂, NO₂, and CO, were observed in China between 2015 to 2020. On the other hand, during the same period, there was a notable increase in ozone (O₃) concentrations, making it a prominent air pollutant in eastern China. The annual mean concentration of maximum daily 8-hour average (MDA8) O₃ exhibited alarming linear increases of 2.4, 1.1, and 2.0 ppb yr⁻¹ in three megacity clusters: Beijing-Tianjin-Hebei (BTH), Yangtze River Delta (YRD) and Pearl River Delta (PRD), respectively. Meanwhile, there was a significant three-fold increase in the number of O₃-exceeding days, defined as MDA8 O₃ > 75 ppb. Our analysis indicated that the upward increases in the annual mean concentration of MDA8 were primarily driven by the rise in consecutive O₃-exceeding days. There were expansions of high O₃ in urban centers to rural areas accompanied by a saturation effect that MDA8 O₃ concentrations at the high O₃ stations in 2015 remained nearly constant of 100 ppb. Lastly, we found a close association between O₃ episodes with four or more consecutive O₃-exceeding days and the position and strength of tropical cyclones (TCs) in the northwest Pacific and the West Pacific subtropical high (WPSH). The TC and WPSH contributed to meteorological conditions characterized by clear skies, subsiding air motion, high vertical stability in the lower troposphere, increased solar radiation, and positive temperature anomaly at the surface. These favorable meteorological conditions greatly facilitated the formation of O₃. Thus, we propose that the worsening O₃ increases observed in BTH, YRD and PRD from 2015 to 2020 can be mostly attributed to enhanced photochemical O₃ production resulting from an increased occurrence of meteorological conditions with high solar radiation and positive temperature anomalies under the influence of WPSH and TCs.

1 Introduction

Ozone (O₃) is an important greenhouse gas, which can also have adverse effects on human health, vegetation, and materials (Bell et al., 2006; Cohen et al., 2017; Kalabokas et al., 2020; Nuvolone et al., 2018). Surface O₃ is a secondary pollutant produced by photochemical reactions involving O₃ precursors such as volatile organic compounds (VOCs), carbon monoxide

(CO) and nitrogen oxides (NO_x) (Ma et al., 2012; Monks et al., 2015; Wang et al., 2017). In addition to O₃ precursors, meteorological conditions are also crucial factors driving the O₃ formation. Solar radiation, temperature, relative humidity, wind speed, and cloud cover have been found to be closely related to O₃ formation (Dong et al., 2020; Han et al., 2020; Yin et al., 2019). Furthermore, large-scale circulations, such as the East Asian monsoon, West Pacific subtropical high (WPSH) and tropical cyclones (TCs) can influence O₃ concentration as well (Lu et al., 2019; Rowlinso et al., 2019; Yang et al., 2014; Zhao and Wang, 2017).

The concentrations of air pollutants SO₂, NO_x, CO, PM₁₀ and PM_{2.5} in China have been significantly reduced since 2013 (Li M. et al., 2021; Li et al., 2022; Zhai et al., 2019), thanks to the implementation of “Air Pollution Prevention and Control Action Plan”. However, the O₃ concentration has dramatically increased and emerged as a major air pollutant in eastern China (Bian et al., 2019; Fu et al., 2019; Wang et al., 2020; Zheng et al., 2018). O₃ concentrations are particularly high in the three megacity clusters in eastern China, namely Beijing-Tianjin-Hebei (BTH), Yangtze River Delta (YRD) and Pearl River Delta (PRD) (Gao et al., 2020; Guo et al., 2019; Li K. et al., 2021; Liu et al., 2018; Yang et al., 2019).

Annual mean concentrations of maximum daily 8-hour average (MDA8) O₃ in the three megacity clusters are shown in Fig. 1. The linear increases of MDA8 O₃ for BTH, YRD and PRD are 2.4, 1.1 and 2.0 ppb yr⁻¹, respectively during the period 2015–2020. These increases are unusually large compared to the increases in other parts of China as well as the positive trends worldwide (Chen et al., 2020; Lu et al., 2018; Professional Committee of Ozone Pollution Control of Chinese Society for Environmental Sciences, 2022; Zhang et al., 2020). Thus, a crucial scientific question is: What is the cause(s) of these large increases in O₃ concentration? Some recent studies suggested that changing photochemical processes induced by anthropogenic emissions are responsible for these increases (Li et al., 2019; Li et al., 2022; Shao et al., 2021; Wang et al., 2020). However, in our analysis of the O₃ increases at individual stations in eastern China during the period 2015–2020, we noticed that the interannual variations of O₃ concentration were strongly affected by the position and intensity of WPSH and the presence of TCs in the western Pacific and South China Sea, consistent with the results of a number of recent studies (Chang et al., 2019; Mao et al., 2020; Ouyang et al., 2022; Zhao and Wang, 2017). These results suggest that transport/meteorological parameters associated with WPSH and TCs may also play an important role in the large increases of MDA8 O₃.

The significant impact of WPSH on weather patterns and O₃ concentrations over East China is widely recognized (Bachmann, 2015; Chang et al., 2019; Yin et al., 2019; Zhao and Wang, 2017). It is well established that the WPSH plays a critical role in controlling weather conditions, which in turn affects O₃ concentrations. For example, the WPSH is known to contribute to the formation of the East Asian monsoon and influence precipitation patterns in the YRD. It also influences air temperature and precipitation across North and South China (Zhang, 2001; Zhao and Wang, 2017). These changes in meteorological conditions have a profound impact on the photochemical production, dispersion, and accumulation of O₃.

Previous studies have indicated that in the peripheries of TCs, PRD experiences specific atmospheric conditions, e.g., high pressure, low humidity, and intense solar radiation, which are highly conducive to O₃ formation. These conditions often result in consecutive days with elevated levels of O₃, as observed in various case studies (Ouyang et al., 2022; Wei et al., 2016). Furthermore, statistical investigations have established several noteworthy connections between TCs and O₃ concentrations in

the PRD area. For example, the meteorological conditions associated with the TC periphery frequently contributed to the formation of elevated surface O₃ levels and aerosols (Deng et al., 2019). In addition, TCs in the East China Sea had a higher likelihood of causing increased O₃ concentrations in the PRD region (Zhao et al., 2022). Lastly, TCs in the vicinity of Taiwan, China have the greatest influence on air quality in Hong Kong when compared to TCs in other areas, which is primarily because these TCs facilitate the transportation of air pollutants from the PRD region (Lam et al., 2018).

In this study, we focus on exploring possible contributions to the large O₃ increases in the three megacity clusters in eastern China by changes in meteorological parameters associated with WPSH and TCs during the period 2015–2020. The rest of this paper is organized as follows. In Section 2, the data and methodology used in this study are described. Major characteristics of the O₃ interannual variability and increases in the three megacity clusters are discussed in Section 3.1. In Section 3.2, we examine the spatial expansion and saturation of high O₃. The annual change of O₃-exceeding days with different durations are also examined. A hypothesis of the cause of O₃ increases in three megacity clusters in eastern China during 2015–2020 is presented in Section 3.3. Section 4 presents a summary and conclusions.

2 Data and methodology

2.1 Pollutant Data

The observed hourly concentrations of air pollutants, including O₃, NO₂, CO, PM_{2.5}, and SO₂ from 2015 to 2020 are obtained from the Chinese National Environmental Ministry of Environmental Protection (<http://www.cnemc.cn/en/>). Gridded MDA8 O₃ data from Tracking Air Pollution in China dataset (<http://tapdata.org.cn>) with a resolution of 10 km are also used (Xue et al., 2020).

2.2 Meteorological Data

The European Centre for Medium-Range Weather Forecasts (ECMWF) Reanalysis v5 (ERA5) dataset (available at <https://cds.climate.copernicus.eu/>), with a horizontal resolution of 0.25° × 0.25° and a temporal resolution of 1 h, was used to analyze the influence of meteorological parameters on O₃ pollution. The variables used in this study include 2 m temperature (T2m), surface net solar radiation (SSR). In addition, daily mean relative humidity, geopotential height, zonal and meridional wind at 500 hPa, and vertical velocity at 850hPa from the National Center for Environmental Prediction (NCEP) and National Center for Atmospheric Research (NCAR) reanalysis (<https://psl.noaa.gov/data/gridded/data.ncep.reanalysis.html>) at a resolution of 2.5°×2.5° are used.

2.3 Methods

The Chinese National Ambient Air Quality Standard for MDA8 O₃ is 160 μg m⁻³, which corresponds to 75 ppb at 273.15 K and 1 atm. The O₃-exceeding days are defined as MDA8 O₃ concentration >75 ppb, while non-O₃-exceeding days are defined as MDA8 O₃ concentration <75 ppb. According to the duration of O₃ pollution episode, it can be divided into consecutive O₃-

exceeding days with four or more days ($O_3 \text{ days} \geq 4$) and consecutive O_3 -exceeding days less than four days ($O_3 \text{ days} < 4$). In addition, some commonly-used methods are applied in this study, including linear fitting, meteorological synthesis method, and two-tailed Student's t test.

100 The normalized annual mean O_3 concentration of the O_3 -exceeding days is calculated by adding the O_3 concentration of the O_3 -exceeding day each year and dividing it by the total number of days in the year. The normalized annual mean O_3 of the non- O_3 -exceeding days is calculated by the same method except for the non- O_3 -exceeding days.

Table 1 lists the criteria and corresponding numbers of low O_3 and high O_3 stations in the three megacity clusters. This classification is undertaken with the purpose of distinguishing stations with various O_3 levels within the three megacity clusters, and it is based on the number of O_3 exceeding days in 2015. Stations with the number of O_3 -exceeding days fewer than or equal to the low O_3 criterion (second column) are considered as low O_3 stations. When more than or equal to the high O_3 criterion (4th column), they are considered as high O_3 stations. We have tested a few reasonably different criteria and found only some insignificant differences in the results. i.e., the results associated with low O_3 and high O_3 stations are robust against reasonable changes in their selection criteria. For example, the relatively large criterion (37 days) of low O_3 stations in YRD is intended to include a large enough number of stations (about one third of the total of 152 stations) to be fully representative of low O_3 and moderate O_3 stations. These results have been compared to those of a more stringent criterion of nineteen days and found no notable change in the major characteristics (Figs. S1 and S2).

3 Results and discussion

3.1 Major characteristics of O_3 increases

Major characteristics of the large increases in the annual mean O_3 concentration are shown in Figs. 2a, 2b and 2c for BTH, YRD and PRD, respectively, in which the normalized annual mean concentrations of MDA8 O_3 in the three megacity clusters are compared to contributions from two groups: The O_3 -exceeding days and non- O_3 -exceeding days. The increase in O_3 -exceeding days is the primary contributor to the substantial increase in the annual mean O_3 in all three megacity clusters from 2015 to 2020. The contribution of O_3 -exceeding days is affected mostly by the changing number of exceeding days (more than 80%), and secondly but nevertheless significantly by their changes in concentrations (less than 20%) (Tables 2–4). e.g., in BTH the exceeding days were 31, 43, 62, 74, 96 and 78 days in the individual years of 2015–2020, respectively, while their concentrations of those years were 66.42, 64.13, 69.44, 68.21, 70.19 and 69.69, respectively (Table 2 second column). Contributions from non- O_3 -exceeding days are insignificant (p -value > 0.1), except that in BTH (Fig. 2a) which shows a significant declining contribution (p -value = 0.02) due to the reduced number of non- O_3 -exceeding days. Therefore, the following discussions on the O_3 increases will be focused on the O_3 -exceeding days.

125 Annual numbers of single and consecutive O_3 -exceeding days are shown in Figs. 3a, 3b and 3c for BTH, YRD and PRD, respectively. A drastic two to three-fold increase in the annual numbers of consecutive O_3 -exceeding days can be seen in all three regions. In contrast, the numbers of single O_3 -exceeding days show only a slight increase in PRD. These drastic increases

in the annual numbers of consecutive O₃-exceeding days are clearly the primary contributors to the increases in O₃ shown in Figs. 2a, 2b and 2c. This brings up two key scientific questions: What is the cause(s) of the drastic increases in the numbers of consecutive O₃-exceeding days? Is it due to changing emissions of air pollutants or changing meteorological parameters?

3.2 Spatial expansion and saturation of high O₃

Another important changing characteristics of O₃ concentrations is illustrated in Fig. 4a, which depicts the annual mean concentrations of MDA8 O₃ in BTH during O₃-exceeding days for all 78 stations (black line), 14 stations in the highest category of O₃ concentration (average 103 ppb) observed in 2015 (red line, denoted as high O₃ stations hereafter, Table 1) and 13 stations in the lowest category of O₃ (average 57 ppb) observed in 2015 (green line, denoted as low O₃ stations hereafter, Table 1). It is remarkable that O₃ concentrations at the low O₃ stations caught up within 12 ppb with other stations in merely two years (an increase of about 30 ppb from 2015 to 2017), and actually equaled the average of other stations in 2019. Meanwhile, the high O₃ stations experienced a slight decrease in O₃ concentration, albeit not statistically significant. This phenomenon suggests strongly that the annual mean concentrations of MDA8 O₃ in BTH experienced a fast (within two years) and widespread spatial expansion of high O₃ from urban centers to surrounding regions where O₃ concentrations were low in 2015. Temporally most of the expansion was accomplished during 2015–2017. This phenomenon of a fast and widespread expansion of high O₃ concentrations from urban centers to surrounding regions were also observed at a slightly less degree in YRD (Fig. 4b) and PRD (Fig. 4c).

The spatial expansion of high O₃ from urban centers to surrounding regions in BTH and YRD during 2015–2017 can be clearly visualized in Figs. 5 and 6, respectively. Figs 5a, 5b and 5c show the spatial distribution of daily mean concentrations of MDA8 O₃ for O₃-exceeding days in BTH in 2015, 2017 and their difference (2017 minus 2015), respectively. Comparing Fig. 5a to 5b, one can see that the area greater than 80-ppb (75 ppb is the O₃ exceeding standard) expanded by about a factor of five from 2015 to 2017. The daily average concentration of MDA8 O₃ within the BTH box increased from 66.42 ppb in 2015 (31 days, Fig. 5a) to 69.44 ppb in 2017 (62 days, Fig. 5b), which was a difference of 3.02 ppb or a merely 4.5% increase between the two years (Fig. 5c). After accounting for the number of O₃-exceeding days, the ratio of normalized MDA8 O₃ in all O₃-exceeding days between 2017 and 2015 became 2.09. This comparison suggests that the increase in O₃ in BTH between 2015 and 2017 was driven primarily by the increase in the number of consecutive O₃-exceeding days. Spatially Fig. 5c shows the expansion is mostly to the south and southwest outside of BTH, with YRD getting a predominant portion of O₃ enhancements. Within the BTH box, the nearly constant concentrations of O₃ inside Beijing City (40°N, 116.5°E) coupled with the southwestward expansion of high O₃ in 2017 (Fig. 5c) suggested that there was a saturation of O₃ inside Beijing City. There were also significant increases in O₃ in YRD and even in southern China as far as the western PRD (Fig. 5c). Nevertheless, the O₃ concentrations in YRD and PRD stayed below 70 ppb during the O₃-exceeding days of BTH in both 2015 (Fig. 5a) and 2017 (Fig. 5b). In other words, the O₃-exceeding days of YRD and PRD are mostly decoupled temporally (i.e., not occurring at the same time) from those of BTH. A logical explanation for this phenomenon is that the atmospheric conditions conducive to high O₃ formation in BTH do not overlap significantly with those conditions of YRD and PRD.

Figs. 6a, 6b and 6c are the same as Figs. 5a, 5b and 5c, respectively, but for YRD. Similar to BTH, one can clearly see the expansion of high O₃ from the vicinity of Shanghai City (31°N, 121.3°E) in the northwestern direction reaching as far as the central BTH box during the period 2015–2017 (Figs. 6b and 6c). Comparing Fig. 6a to 6b, one can see that the area greater than 70-ppb expanded from Shanghai and vicinity northwestward by more than a factor of five from 2015 to 2017. This expansion was in different direction from the southwestern expansion occurred in BTH (Fig. 5c). We note, however, this expansion does not necessarily mean the direct transport of high O₃ or its precursors from the vicinity of Shanghai to central BTH. In fact, the presence of separate rather than contiguous red patches of high O₃ (>70 ppb) in southern BTH and northern YRD in Fig. 6b is a clear indication that the high O₃ are primarily controlled by local photochemical production from local O₃ precursors, rather than the direct upwind-downwind transport of high O₃ and/or its precursors. The daily average concentration of MDA8 O₃ within the YRD box increased from 53.79 ppb in 2015 (31 days, Fig. 6a) to 64.35 ppb in 2017 (40 days, Fig. 6b), which was a difference of 10.56 ppb or a 20% increase between the two years (Fig. 6c). After accounting for the number of O₃-exceeding days, the ratio of normalized MDA8 O₃ in all O₃-exceeding days between 2017 and 2015 became 1.54. This comparison suggests that the increase in O₃ in YRD between 2015 and 2017 was due to both the increases in O₃ concentrations (+20%) and the number of O₃-exceeding days (+34%).

Figs. 7a, 7b and 7c are the same as Figs. 5a, 5b and 5c, respectively, but for PRD. Unlike BTH and YRD, there was only a slight expansion of high O₃ within the PRD box toward the southwest in 2017 compared to 2015 (Fig. 7c). Nevertheless, outside the PRD box there was an extensive expansion of high O₃ in eastern China, substantially greater than the expansion within the PRD box (Fig. 7c). The daily average concentration of MDA8 O₃ within the PRD box increased from 61.16 ppb in 2015 (14 days, Fig. 7a) to 65.18 ppb in 2017 (36 days, Fig. 7b), which was a difference of 4.02 ppb or a merely 6.6% increase between the two years (Fig. 7c). After accounting for the number of O₃-exceeding days, the ratio of normalized MDA8 O₃ in all O₃-exceeding days between 2017 and 2015 became 2.74. This comparison suggests that the increase in O₃ in PRD between 2015 and 2017 was almost entirely (93.4%) due to the increase in the number of O₃-exceeding days.

Figs 7a and 7b reconfirm that O₃-exceeding days in PRD were mostly decoupled from those in BTH (Figs. 5a and 5b) and YRD (Figs. 6a and 6b), as their spatial distributions were characterized by highly distinctive regional features in both 2015 and 2017. These differences suggest that the O₃-exceeding days mostly occur in different days in the three individual regions. On the other hand, a comparison of Figs. 7c, 6c and 5c reveals a striking common feature of high values in southwestern BTH and northwestern YRD, and low values in eastern parts of all three BTH, YRD and PRD boxes. These common features suggest that the difference between 2015 and 2017 in all three individual regions is likely caused by a common mechanism/process that changed from 2015 to 2017. Moreover, as suggested in Fig. 3, this common mechanism/process must be closely related to higher number of consecutive O₃-exceeding days in 2017 over those of 2015.

Comparison of Fig. 8a to Fig. 1 reveals an interesting point: While the yearly average MDA8 O₃ concentrations at all stations in BTH (green line in Fig 1) shows a significant O₃ increase of 2.38 ppb yr⁻¹ with p=0.01, the black line in Fig. 8a (MDA8 O₃ of all stations during O₃ exceeding days) shows an insignificant increase of 1.22 ppb yr⁻¹ with p=0.2. This is because the values in Fig. 8 are those of O₃ exceeding days, of which O₃ concentrations at high O₃ stations (red line in Fig. 8a) have a small

195 decrease (albeit insignificant) due to the saturation effect discussed above. This decrease is the main contributor to the high p value of 0.2 of the black line in Fig. 8a (all stations).

We believe this spatial expansion and saturation of high O₃ is a significant new finding which may provide critical clue to the cause(s) of the large O₃ increases. Its robustness against new data in 2021 and 2022 has been evaluated and reconfirmed in the Supplementary (Figure S14).

200 **3.3 Causes of ozone enhancement at low-concentration sites and saturation at high-concentration sites**

Major findings of subsections 3.1 and 3.2 can be summarized as follows: (1) Increases in O₃ observed in the three megacity clusters in eastern China during 2015–2020 (Fig. 1) were mainly caused by the large increases of approximately two to three-fold increase in the number of consecutive O₃-exceeding days (Figs. 2 and 3). (2) A fast and widespread expansion of high O₃ from urban centers to surrounding regions was observed in the three megacity clusters during 2015–2019 (Fig. 4); and the majority of the expansions were accomplished during the 2015–2017 period (green lines in Fig. 4). And (3), the expansions of high O₃ in the three megacity clusters were accompanied by a saturation effect that O₃ concentrations at the high O₃ stations of approximate 100 ppb in 2015 remained nearly constant or slightly declined throughout the entire period of 2015–2020 (Fig. 4).

3.3.1 Changing emissions as a possible cause of O₃ increases in 2015–2020

210 As mentioned earlier, two emission oriented hypotheses have been proposed as a possible cause of the O₃ increases in 2015–2020. One is changing emissions of O₃ precursors NO_x and VOC (Li et al., 2022). The other is the reduced removal of HO₂ radicals due to diminishing PM_{2.5} (Li K. et al., 2021; Shao et al., 2021). Li et al. (2022) showed that the NO titration effect was the cause of the linear trend in O₃ in PRD (0.5 ppb yr⁻¹) during the relatively long period 2006–2019. But for the period 2015–2020, the NO titration effect could account for only about 10% of the linear increase in O₃ of the low O₃ stations in PRD (5.0 ppb yr⁻¹, green line, Fig.S3a).

The increase of 30 ppb in O₃ at the low O₃ stations in BTH from 2015 to 2017 (green line, Fig. 4a and Fig. 8a) represents about 50% increase in O₃. The titration effect can account for only about 5% (Fig. 8f). If this increase of 30 ppb in O₃ were due to an enhancement in O₃ precursors, the enhancement would have to be substantially greater than 50% because of the well-known less-than-linear relationship between changes in O₃ and its precursors, i.e., substantially more percentage changes in precursors are needed for each percentage change in O₃ (Dodge, 1977; Shafer and Seinfeld, 1985). Figs. 8d and 8f show that CO (a proxy for VOC) and NO_x changed only by a few percent from 2015 to 2017, more than one order of magnitude less than the changes needed. Hence it appears that changes in meteorological conditions conducive to O₃ formation are more likely the major contributing factor to the 50% increase in O₃ at the low O₃ stations in BTH. Similar argument can be extended to YRD and PRD (Figs. S1 and S3).

225 The theory of reduced removal of HO₂ radicals by diminishing PM_{2.5} (25%, green line of Fig. 8c) appeared to be valid qualitatively for the 50% increase in O₃ at the low O₃ stations in BTH from 2015 to 2017 (green line of Fig. 8a). But this theory

was contradicted directly by the phenomenon at the high O₃ stations where a 30% reduction in PM_{2.5} (red line of Fig. 8c) corresponded to a decrease rather than an increase in O₃ (red line of Fig. 8a).

3.3.2 Changes in meteorological parameters as a possible cause of O₃ increases in 2015–2020

230 Fig. 9a shows the mean daily O₃ concentrations of the first group with four or more consecutive O₃-exceeding days (labeled O₃ days_{≥4}) in 2015, Fig. 9b shows the mean daily O₃ concentrations of the second group with less than four consecutive O₃-
exceeding days (labeled O₃ days_{<4}), and Fig. 9c is the difference between the two groups (6.10 ppb, Table 2). Figs. 9d–9f are
the same as Figs. 9a–9c, respectively, but for 2017. The first group in 2017 had 28 days and mean O₃ of 74.43 ppb inside the
BTH box, while the second group had 34 days and 65.32 ppb (Table 2). One of the most remarkable differences between 2017
235 and 2015 in Figs. 9a–9f was the large number of days with four or more consecutive O₃-exceeding days (first group) in 2017
(28 days, Fig. 9d) over that of 2015 (7 days, Fig. 9a), which alone contributed to about 62% of the difference in O₃ between
2017 and 2015 as shown in Fig. 2a (red line). Approximately 30% was contributed by the 10 days' difference (2017 vs. 2015)
in the number of days with less than four consecutive O₃-exceeding days (second group). The contribution by the higher
average concentration of MDA8 O₃ of the first group in 2017 is only about 8% (Table 2). These values of contributions
240 reconfirm what is shown in Fig. 3a, i.e., the greater frequency of episodes with four or more consecutive O₃-exceeding days
contributes the majority (62%) to the higher O₃ in BTH in 2017 vs. 2015, the greater intensity/concentration of O₃ during the
episodes contributes only about 8%, consistent with the expansion and saturation effect discussed earlier.

The phenomena illustrated in Figs. 9a–9f also exist in YRD and PRD as well as in most other years. Figures equivalent to Figs.
9a–9c for all years in the three city clusters are provided in the Supplementary Material (Figs. S4–S6). Essential information
245 derived from those figures is summarized in Tables 2–4.

In Figs. 10a and 10b the values of SSR and T2m of the episodes with four or more consecutive O₃-exceeding days are compared
to those of O₃ episodes with less than four consecutive O₃-exceeding days, and to those of clean days (non-O₃-exceeding days).
As expected, the O₃ episodes with four or more consecutive O₃-exceeding days consistently have the highest values of SSR
and T2m, while the clean days have the lowest values. This is the case in nearly all years studied as shown in the Supplementary
250 Material (Fig. S7) and is also generally true in YRD and PRD (Figs. S8 and S9). Coupling the higher values of SSR and T2m
in the O₃ episodes with four or more consecutive O₃-exceeding days depicted in Fig. 10 and greater number of days in the O₃
episodes with four or more consecutive O₃-exceeding days shown in Fig. 3, we therefore propose a hypothesis as follows: the
cause of worsening O₃ increases in BTH, YRD and PRD from 2015 to 2020 could be attributed to enhanced photochemical
O₃ production due to increased occurrence of meteorological conditions of high solar radiation and positive temperature
255 anomaly at the surface.

Quantitatively the coupling of Fig. 10 with Fig. 3 can be performed by multiplying the difference between the red (four or
more consecutive O₃-exceeding days) and green (clean days) values of SSR/T2m in Fig. 10 with the frequency of occurrence
(in percentage of total days) of O₃ episodes with four or more consecutive O₃-exceeding days from Fig. 3. The results are
compared to the yearly O₃-exceeding days in Fig. 11. Correlation between the yearly O₃-exceeding days and weighed SSR is

260 very good with R values 0.88 or greater in all three regions, lending strong support for our hypothesis. Correlation between the yearly O₃-exceeding days and weighed T2m is high correlated in BTH but not correlated in YRD and PRD, which probably suggests that T2m is not as strongly coupled to O₃ formation as SSR. Inclusion of O₃ episodes with less than four consecutive O₃-exceeding days in Fig. 11 did not change the correction coefficients significantly, supporting the robustness of results shown in Fig.11.

265 Hu W. et al. (2023), in collaboration with this study, conducted a statistical analysis to assess processes that contribute to high O₃ formation in PRD when TCs were present in the northwest Pacific. They investigated the impact of the distance between TCs in the northwest Pacific and PRD on the O₃ concentration in the PRD from 2006 to 2020. They found that the large numbers of consecutive O₃-exceeding days in 2017 and 2019 relative to 2015 were primarily attributable to the greater occurrence of downdrafts and stable atmospheric conditions brought about by mid-distance category TCs. This finding clearly
270 establishes that changing frequency of mid-distance category TCs (i.e., changing meteorological conditions) is the cause of the increases in the numbers of consecutive O₃-exceeding days as well as the higher O₃ concentrations in PRD. Ongoing study by our research group further shows that the mid-distance category TCs are predominately those TCs with tracks starting around the southern Philippines and ending near Korea and/or Japan. Since TC tracks in northwestern Pacific are strongly controlled by WPSH, we conclude that both Philippines-to-Korea/Japan track TCs and corresponding distribution and intensity
275 of WPSH contributed to the favorable meteorological conditions leading to higher consecutive O₃-exceeding days in PRD from 2015 to 2020.

Following the analysis by Hu W. et al. (2023), the mean vertical velocity at 850 hPa during all O₃-exceeding days in PRD in 2015 (Fig. 12a) is compared to that of episodes with four or more consecutive O₃-exceeding days in 2017 (Fig. 12b). Major features in Fig. 12 compare very well with those of Fig. 7. E.g., area with positive vertical velocity (downdrafts) in 2017 (red
280 area in Fig. 12b), which was highly conducive to O₃ formation, was by far more widespread and greater in value than that of 2015 (red area in Fig. 12a), agreeing well with the greater high O₃ area of Fig. 7b (2017) than that of Fig. 7a (2015). This agreement confirms that the increase in O₃ in PRD from 2015 to 2017 was caused by increased downdrafts and stable atmospheric conditions (meteorological conditions) brought about by TCs as suggested by Hu W. et al. (2023). The same plots for BTH are shown in Fig. 13. Features of Fig. 13 are highly consistent with those of Fig. 5. The same plot for YRD (Fig. S10)
285 also showed more extensive and greater downdrafts in 2017 than 2015. However, the area of positive vertical velocity in YRD appeared to shift about 500 km to the east compared to the area of high of O₃ in Fig. 6b. Considering the uncertainty in evaluating the vertical velocity and that O₃ formation is also dependent on parameters other than the vertical velocity, the discrepancy is acceptable.

In summary of this section, the increases in O₃ in the three megacity clusters are critically dependent on the number of four or
290 more consecutive O₃-exceeding days. In addition, Hu W. et al. (2023) found that the changing frequency of mid-distance category TCs (i.e., changing meteorological conditions) is the cause of the increases in the numbers of consecutive O₃-exceeding days as well as the O₃ concentrations in PRD. More importantly, our additional analyses of the mean vertical velocity at 850 hPa over the three megacity clusters (Figs. 12, 13 and S10) show that the increases in O₃ in all three megacity clusters

from 2015 to 2017 were caused by enhanced downdrafts and stable atmospheric conditions (meteorological conditions) which
295 were highly conducive to O₃ formation. The enhanced downdrafts and stable atmospheric conditions were brought about by
TCs and associated WPSH. Here we bring up WPSH because it is well known that the tracks of TCs are influenced strongly
by WPSH, and that WPSH affects strongly regional atmospheric dynamics and therefore O₃ formation (Chang et al., 2019;
Mao et al., 2020; Ouyang et al., 2022; Zhao and Wang, 2017).

3.3.3 Contribution of western pacific subtropical high

300 Mao et al. (2020) made a comprehensive study of an 11-day O₃ episode in BTH in 2017 and found it was dominated by the
presence of the WPSH and mid-high latitude wave activities. Depending on the position and intensity, WPSH is well known
to be a crucial factor affecting O₃ concentrations in various parts of eastern China (Chang et al., 2019; Yin et al., 2019; Zhao
and Wang, 2017). During this 11-day O₃ episode, the ridge line of WPSH maintained at approximately 22°N from June 24 to
June 29, which in combination with mid-high latitude wave activities induced meteorological conditions highly conducive to
305 the O₃ production in BTH and northern YRD (Mao et al., 2020).

Following the analysis of Mao et al. (2020), the impact of WPSH on O₃ in BTH in April–September has been analyzed in Fig.
14 which depicts the composite 500 hPa geopotential height contours, humidity and winds in BTH in April–September for O₃-
exceeding days in 2015 (a), clean days in 2015 (b), O₃-exceeding days in 2017 (c), clean days in 2017 (d), O₃-exceeding days
in 2019 (e) and clean days in 2019 (f). The three years 2015, 2017 and 2019 are chosen because their differences in O₃
310 contribute predominately to the overall O₃ increases (Figs. 1–2). The importance of WPSH is clearly visible in all Figs. 14a–
14f when the 5880 and 5900 gpm isolines (green lines) of O₃-exceeding days are compared to those of clean days. In all three
years, the WPSH of the former (O₃-exceeding days) were significantly stronger than the latter (clean days) as evident by the
strong anticyclonic winds and/or the larger areas inside the 5880 gpm isolines. Even in the case of 2017 when the area inside
5880 gpm isolines of the former looked to be similar to that of the latter, the appearance of 5900 line in the former indicated a
315 stronger WPSH. The strong anticyclonic winds in the O₃ exceeding days (Figs. 14a, 14c and 14e) force moist air of South
China Sea northward into southern China and contributed to extensive clouds and precipitation and thus low O₃ formation over
southern China and southern YRD. This difference in the O₃ formation between BTH and southern China provides a good
explanation to why the O₃-exceeding days mostly occur in different time periods in the three megacity clusters as discussed in
Section 3.2. Furthermore, over East China Sea the prevailing westerlies were forced northward, slowed down and lead to
320 meteorological conditions in BTH and northern YRD characterized by cloudless sky, sinking motion and high vertical stability
in the lower troposphere, as well as high SSR and positive T2m anomaly at the surface. These meteorological conditions were
highly conducive to the formation and accumulation of O₃. In contrast, the weaker WPSH of the clean days allowed relatively
strong westerlies to prevail over BTH during clean days in the three years, which tended to disperse O₃ (Figs. 14b, 14d and
14f). e.g., the average wind speed over BTH was about 10 m s⁻¹ in Fig. 14b, while only about 5 m s⁻¹ in Fig. 12a. Quantitatively
325 Fig. 12c had 31 more O₃-exceeding days (93 ppb) than Fig. 14a, the 31 days came at the expense of clean days (52 ppb) (Figs.
14b and 14d). The contribution of these 31 days to the difference in MDA8 O₃ between 2017 and 2015 (6.5 ppb, Fig. 1) can

be calculated as follows: $((93 \times 62) + (56 \times 121)) / (62 + 121) - ((86 \times 31) + (52 \times 152)) / (31 + 152) = 10.8$ ppb. This difference of 10.8 ppb in MDA8 O₃ between 2017 and 2015 was for the period of April to September. It should be divided by 2 and became 5.4 ppb for the yearly difference in MDA8 O₃ between 2017 and 2015. This value of 5.4 ppb accounted for 83% of the observed difference in MDA8 O₃ between 2017 and 2015 (6.5 ppb, Fig. 1). Similar statement can be made for the difference in MDA8 O₃ between 2019 and 2017 (Figs. 14e and 14c, Fig. 1).

We have made the same analysis for other years as well as for YRD and PRD. The results are mostly similar, and thus presented in the Supplementary material (Figs. S11, S12 and S13). Figs. 15a–15d for PRD in 2017 and 2019 are shown because there were interesting anticyclonic circulations over PRD during O₃-exceeding days in both years (Figs. 15a and 15c). The 2017 anticyclone was a direct product of the WPSH as it resided within the western tip of the 5880 gpm isoline. The 2019 anticyclone was also likely associated with the WPSH as the center of anticyclone resided just beneath the 5860 gpm isoline to the west of PRD. The anticyclonic circulations were accompanied by stable downdrafts, low winds, and cloudless sky conditions (short arrows and blue shades in Figs. 15a and 15c), which were highly conducive to the O₃ formation. Cloudless sky conditions also occurred in YRD and BTH in Figs. 15a and 15c, but the high wind speed prevented the accumulation of O₃. This difference in O₃ accumulation between PRD and other two regions provide another good explanation to why the O₃-exceeding days mostly occur in different days in the three megacity clusters as discussed in Section 3.2. Quantitatively Fig. 15c had 27 more O₃-exceeding days (90 ppb) than Fig. 13a, the 27 days came at the expense of clean days (39 ppb) (Figs. 15b and 15d). The contribution of these 27 days to the difference in MDA8 O₃ in PRD between 2019 and 2017 (6.0 ppb, Fig. 1) can be calculated as follows: $((90 \times 62) + (46 \times 182)) / (62 + 182) - ((85 \times 35) + (39 \times 209)) / (35 + 209) = 11.6$ ppb. This difference of 11.6 ppb in MDA8 O₃ between 2019 and 2017 was for the period of April to November. It should be divided by 365/244 and became 7.75 ppb for the yearly difference in MDA8 O₃ between 2019 and 2017. This value of 7.75 ppb was 1.75 ppb more than the observed difference in MDA8 O₃ between 2019 and 2017 (6.3 ppb, Fig. 1), suggesting a reduction of about 3.5 ppb in MDA8 O₃ in the cold months of January–March and December between 2017 and 2019, which was approximately confirmed by the observed reduction of 3.88 ppb.

The presence of anticyclonic circulations over PRD is in good agreement with the results of Ouyang et al. (2022) and Hu W. et al. (2023). The latter authors suggested that the anticyclonic circulations over PRD were primarily caused by TCs in northwestern Pacific. Nevertheless, it is widely acknowledged that the tracks of TCs in the northwestern Pacific are influenced, at least to some extent, by WPSH (Sun et al., 2015; Wang et al., 2017), making it difficult to separate the roles played by the TCs on the anticyclonic circulations and O₃ formation from those of WPSH. Clearly, further investigations is needed to fully understand the complex relationship among WPSH, TCs and O₃. Based on these results, we hypothesize that the increased frequency of these meteorological conditions enabled by the changing intensity and position of WPSH could contribute as a major cause of the O₃ increases in the three megacity clusters in eastern China during 2015–2020.

3.3.4 Saturation at high-concentration sites

Why the favorable meteorological conditions only cause O₃ increase at low O₃ stations, but not at high O₃ stations? And why the saturation O₃ level is around 100 ppb as shown in Fig. 4? These questions can be best addressed by examining Fig. 16 which depicts the time series of individual processes (where DDEP denotes dry deposition, CHEM the net photochemical production of O₃, HTRA the horizontal transport and VTRA the vertical transport) contributing to O₃ budget in PRD (averaged over 56 stations in PRD) calculated by the WRF-CMAQ model for the O₃ episode of September 24–October 1, 2019 (Ouyang et al., 2022). This episode was one of the most sever O₃ episodes since the official O₃ observation started in PRD in 2006. MDA8 O₃ exceeded the 75 ppb standard on all eight days of the episode. Hourly O₃ reached as high as 110 ppb, yet all MDA8 O₃ stayed approximately between 75 and 100 ppb. This suggests a ceiling/saturation level of approximately 100 ppb for MDA8 O₃, consistent with what was observed in Fig. 4 for PRD as well as BTH and YRD. Since this episode was one of the most sever episodes, we can assume that the 100 ppb saturation level would also be applicable to other O₃ episodes in Guangdong. More importantly, the saturation effect was also a common feature in the results of other three-dimensional models for other megacity clusters, in which MDA8 O₃ usually saturated at 100 ppb, e.g., in YRD (Li et al., 2012) and in Beijing (Zhang et al., 2023). This explains why the saturation O₃ level is around 100 ppb as shown in Fig. 4.

In regard to the first question: Why the favorable meteorological conditions only cause O₃ increase at low O₃ stations, but not at high O₃ stations? It can be understood as follows: At a low O₃ station of 65 ppb MDA8 O₃ in PRD in 2015 (Fig. 4c), Fig. 16 shows that MDA8 O₃ can readily increase to 75–100 ppb in a few hours from an early morning low ozone of about 50 ppb under favorable meteorological conditions. However, at a high O₃ station of 100 ppb MDA8 O₃ in 2015 (Fig. 4c) under the same favorable meteorological conditions, MDA8 O₃ would also reach 75–100 ppb in a few hours from an early morning low ozone of about 50 ppb (Note here we assume all stations start the day with an early morning low ozone of 50 ppb, consistent with the value in Fig. 16). In other words, the saturation levels at all stations are the same at 75–100 ppb, independent of the ozone concentration in 2015.

In terms of contributing processes, the saturation level of 75–100 ppb is controlled primarily by photochemical loss, dry deposition and dispersion to the free troposphere. This can be clearly seen in Fig. 16, on all eight days in the mid-morning when O₃ is approaching toward its peak value, CHEM declines sharply due to photochemical loss, and HTRA, VTRA and DDEP all become greater. Near noontime O₃ starts to drop sharply.

3.4 Uncertainty and cautionary statements

It is worth noting that the analyses conducted in Sections 3.1–3.3 have predominantly relied on correlation or regression analysis techniques, which do not imply a cause-and-effect relationship. To establish a cause-and-effect link between the proposed changes in meteorological parameters and O₃ increases, it is necessary to employ a mechanistic model that is based on the proposed causes and can accurately reproduce the observed O₃ increase. Until such model reproduction is achieved, all correlation or regression findings should be considered as a potential maximum cause-and-effect relationship (Wu et al., 2022). However, current mechanistic models suffer from significant uncertainties, making it difficult to credibly simulate critical

atmospheric processes that regulate O₃ formation. These processes include atmospheric transport parameterizations, the sources and sinks of OH, HO₂ and RO₂ radicals, and the photochemistry of VOCs and OVOCs.

4 Summary and Conclusions

395 Thanks to a strong emission control policy, major air pollutants in China, including PM_{2.5}, SO₂, NO₂ and CO had shown remarkable reductions during 2015–2020. However, O₃ concentration had increased significantly and emerged as a major air pollutant in eastern China during the same time period. The annual mean concentration of MDA8 in three megacity clusters in eastern China, namely BTH, YRD and PRD, showed alarming large upward linear increases of 25%, 10% and 19%, respectively during 2015–2019. Identifying the causes of these worsening O₃ increases is urgently required for air pollution prevention and management.

400 Some recent studies suggested that enhanced photochemical processes induced by changing anthropogenic emissions were responsible for these increases (Li et al., 2019; Li et al., 2022; Shao et al., 2021; Wang et al., 2020). However, we noticed that there was independent evidence, including the spatial distribution of the expansion of high O₃ (Figs. 5 and 6) and inter-annual variations in O₃, Ox, NO₂, CO and PM_{2.5} (Fig. 8), suggesting that transport/meteorological conditions rather than emissions of O₃ precursors were more likely to be the major contributor to the O₃ increases. Moreover, we found that the increases in O₃ observed in the three megacity clusters during 2015–2020 (Fig. 1) were mainly caused by the large increase of approximately two to three-fold increase in the number of consecutive O₃-exceeding days (Fig. 3), during that time a fast and widespread expansion of high O₃ from urban centers to surrounding regions was observed (Fig. 4), and the majority of the expansions was accomplished during the two-year 2015–2017 period (green lines in Fig. 4). Furthermore, the expansions of high O₃ in the three megacity clusters were accompanied by a saturation effect that O₃ concentrations at the high O₃ stations (high O₃ in 2015) of approximate 100 ppb remained nearly constant throughout the entire period of 2015–2020 (even to 2022, Fig. S14), while the low O₃ stations (low O₃ in 2015) with O₃ less than 75 ppb in all three megacity clusters experienced a significant enhancement in O₃ (>5 ppb yr⁻¹) during 2015–2017 (Figs. 4a, 4b and 4c). Finally, greater frequency of episodes with four or more consecutive O₃-exceeding days contributed the majority to the higher O₃ in all three megacity clusters in 2017 vs. 2015, the greater intensity/concentration of O₃ during the episodes contributes only about 10% (Fig. 9), consistent with the expansion and saturation effect discussed earlier.

415 Coupling the higher values of SSR and T2m in the O₃ episodes with four or more consecutive O₃-exceeding days depicted in Fig. 10 and greater occurrence (number of days) in the O₃ episodes with four or more consecutive O₃-exceeding days shown in Fig. 3, we hypothesize that the cause of the worsening O₃ increases in BTH, YRD and PRD from 2015 to 2020 could be attributed to enhanced photochemical O₃ production due to the increased occurrence of meteorological conditions of high solar radiation and positive temperature anomaly under the influence of WPSH and TCs. The hypothesis is substantiated in Fig. 11, which shows excellent correlation between the yearly O₃-exceeding days and SSR with R values 0.88 or greater in all three

regions. Correlation between the yearly O₃-exceeding days and T2m is good in BTH but poor in YRD and PRD, which probably suggests that T2m is not as strongly coupled to O₃ formation as SSR.

The increases in O₃ in the three megacity clusters are found to be critically dependent on the number of four or more consecutive O₃-exceeding days. In collaboration with this study, Hu W. et al. (2023) found that the changing frequency of mid-distance category TCs (i.e., changing meteorological conditions) is the cause of the increases in the numbers of consecutive O₃-exceeding days as well as the O₃ concentrations in PRD. Our additional analyses of the mean vertical velocity at 850 hPa in the three megacity clusters (Figs. 12, 13 and S10) show that the increases in O₃ in all three megacity clusters from 2015 to 2017 were associated with enhanced downdrafts and stable atmospheric conditions (meteorological conditions) which were highly conducive to O₃ formation. Finally, the enhanced downdrafts and stable atmospheric conditions were most likely brought about by TCs and associated WPSH.

Therefore, we propose that the O₃ concentrations at the high O₃ stations stayed close to a saturation level of about 100 ppb throughout 2015 to 2020, even under more frequent conditions conducive to O₃ formation, was the result of a relatively high rates of atmospheric dispersion, dry deposition and photochemical loss due to the high O₃ concentration (Fig. 16). While the low O₃ stations, where O₃ production were relatively small in 2015, experienced significant enhancements in the O₃ production in 2017 and 2019 because of more frequent downdrafts and stable atmospheric conditions associated with TCs and WPSH in the northwestern Pacific, which were highly conducive to O₃ photochemical production and accumulation (Hu W. et al., 2023). Following the analysis of Mao et al. (2020), the impact of WPSH on O₃ in BTH in April–September has been analyzed in Fig. 14. We found that the increased frequency of these meteorological conditions enabled by the changing intensity and position of WPSH could contribute as a major cause of the O₃ increases in the three megacity clusters in eastern China during 2015–2020.

Nevertheless, it is crucial to recognize that the examinations carried out in Sections 3.1–3.3 primarily utilized correlation or regression analysis techniques, which do not inherently establish causal relationships. To attribute cause and effect between the suggested alterations in meteorological parameters and O₃ increases, it is necessary to employ a mechanistic model that accurately replicates the observed O₃ increase based on the proposed cause(s). Until the model successfully reproduces the phenomenon, all correlation or regression findings should be treated as merely indicating the highest potential cause-and-effect relationship (Wu et al., 2022).

In conclusion, we hypothesize that the cause of the worsening O₃ increases in BTH, YRD and PRD from 2015 to 2020 is attributable to enhanced photochemical O₃ production due to the increased occurrence of meteorological conditions of high solar radiation and positive temperature anomaly under the influence of WPSH and TCs. Therefore, we suggest that future O₃ pollution prevention and control policies should pay more attention to changes in the meteorological/climate conditions, particularly changes in the large-scale circulations, including WPSH and TCs.

Data availability. Hourly surface O₃, NO₂, CO, PM_{2.5}, and SO₂ data were obtained from China National Environmental Centre (<http://www.cnemc.cn/en/>). Hourly meteorological data are obtained from European Centre for Medium-Range Weather

Forecasts ERA5 reanalysis (<https://cds.climate.copernicus.eu/>). Daily meteorological data are obtained from National Center for Environmental Prediction (NCEP) and National Center for Atmospheric Research (NCAR) (<https://psl.noaa.gov/data/gridded/data.ncep.reanalysis.html>). The data of this study are available upon request to Shaw Chen Liu (shawliu@jnu.edu.cn).

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Author Contributions. SL and RL proposed the essential research idea. TH, and YL performed the analysis. TH, YL, RL, and SL drafted the manuscript. YX, SO, BW, and YZ helped analysis and offered valuable comments. All authors have read and agreed to the published version of the manuscript.

465 *Competing interests.* The authors declare that they have no conflict of interest.

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Table 1. Criteria and corresponding numbers of low O₃ and high O₃ stations in the three megacity clusters in 2015. The criterion listed for each megacity cluster was based on the number of MDA8 O₃ exceeding days in 2015. For instance, the criterion for a low O₃ site in BTH was the number of MDA8 O₃ exceeding days in 2015 being less than or equal to 19 days, while for high O₃ site was the number of MDA8 O₃ exceeding days in 2015 being greater than or equal to 71 days.

	Criterion of low O ₃ stations	Number of low O ₃ stations	Criterion of high O ₃ stations	Number of high O ₃ stations	Total number of stations
BTH	≤ 19 days	13	≥ 71 days	14	78
YRD	≤ 37 days	54	≥ 67 days	13	152
PRD	≤ 12days	10	≥ 46 days	10	48

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Table 2. Mean O₃ concentrations (ppb) and number of days of all O₃-exceeding days (2nd column), consecutive O₃-exceeding days with less than four days (3rd column), consecutive O₃-exceeding days with four or more days (4th column) and the difference between (≥ 4 days) and (< 4 days) (5th column) within the BTH box in 2015–2020.

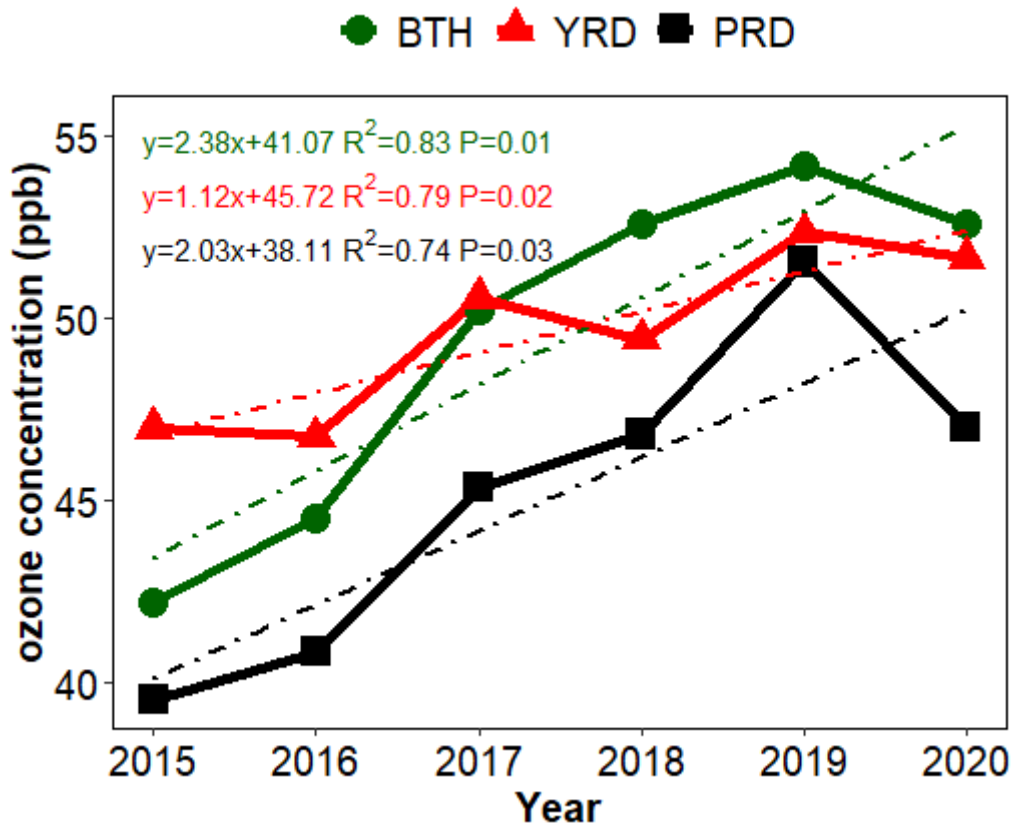
	All days Concentration (days) ppb	<4 days Concentration (days) ppb	≥ 4 days Concentration (days) ppb	Difference (≥ 4 days) – (<4 days) ppb
2015	66.42(31)	65.04(24)	71.14(07)	6.10
2016	64.13(43)	62.65(26)	66.39(17)	3.74
2017	69.44(62)	65.32(34)	74.43(28)	9.11
2018	68.21(74)	65.43(27)	69.80(47)	4.37
2019	70.19(96)	65.28(30)	72.42(66)	7.14
2020	69.69(78)	65.52(40)	74.08(38)	8.56

625 **Table 3.** Mean O₃ concentrations (ppb) and number of days of all O₃-exceeding days (2nd column), consecutive O₃-exceeding days with less than four days (3rd column), consecutive O₃-exceeding days with four or more days (4th column) and the difference between (≥4days) and (<4days) (5th column) within the YRD box in 2015–2020.

	All days Concentration (days) ppb	<4 days Concentration (days) ppb	≥4 days Concentration (days) ppb	Difference (≥4 days) – (<4 days) ppb
2015	53.79(31)	53.59(19)	54.11(12)	0.52
2016	58.87(27)	58.03(23)	63.73(04)	5.70
2017	64.35(40)	62.62(25)	67.22(15)	4.60
2018	63.33(43)	62.49(32)	65.75(11)	3.26
2019	67.18(49)	66.09(27)	68.51(22)	2.42
2020	65.84(38)	64.12(27)	70.06(11)	5.94

630 **Table 4.** Mean O₃ concentrations (ppb) and number of days of all O₃-exceeding days (2nd column), consecutive O₃-exceeding days with less than four days (3rd column), consecutive O₃-exceeding days with four or more days (4th column) and the difference between (≥4days) and (<4days) (5th column) within the PRD box in 2015–2020.

	All days Concentration (days) ppb	<4 days Concentration (days) ppb	≥4 days Concentration (days) ppb	Difference (≥4 days) – (<4 days) ppb
2015	61.16(14)	61.16(14)	---(0)	---
2016	58.44(19)	58.44(19)	---(0)	---
2017	65.18(36)	64.60(23)	66.20(13)	1.60
2018	65.82(31)	63.27(16)	68.55(15)	5.28
2019	69.80(62)	65.96(29)	73.16(33)	7.20
2020	65.08(37)	63.87(22)	66.84(15)	2.97



635 Figure 1: Annual mean concentrations of maximum daily 8-hour average O₃ in BTH (green), YRD (red) and PRD (black).

● Non-ozone days ▲ Ozone days ■ Total days

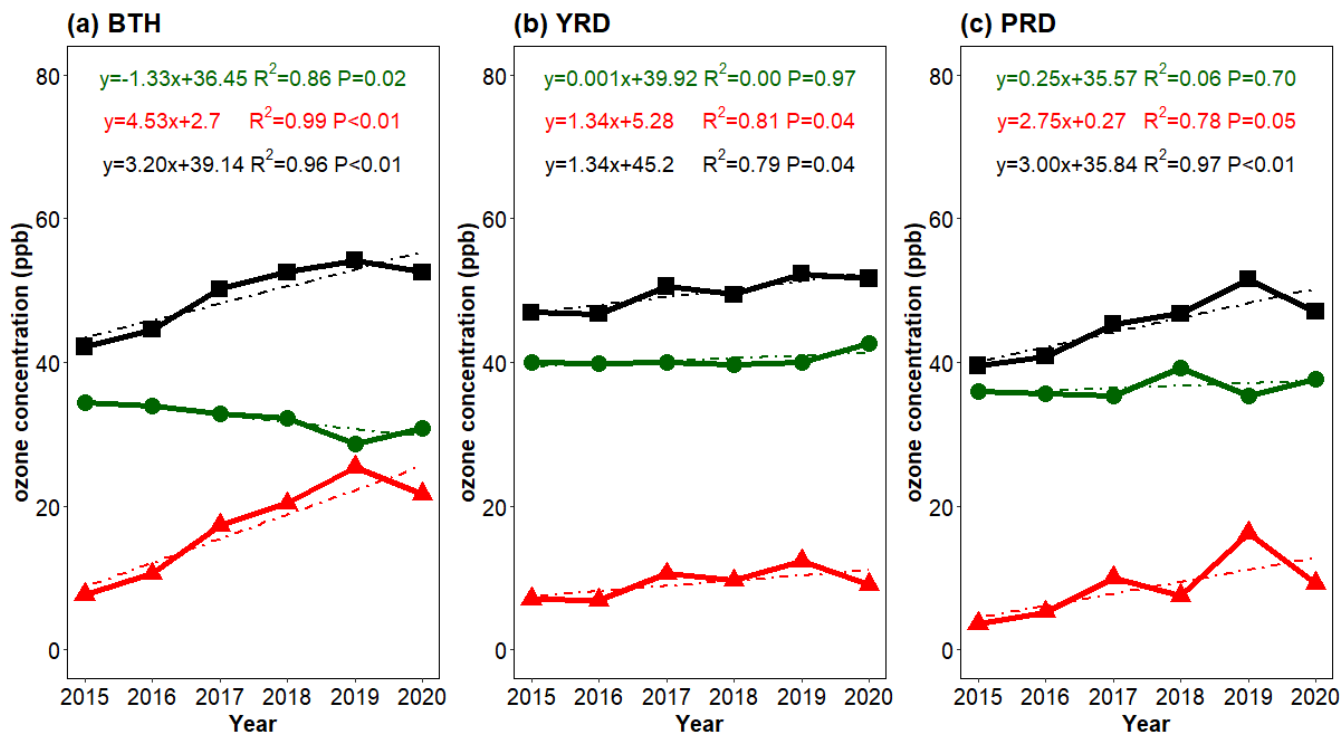
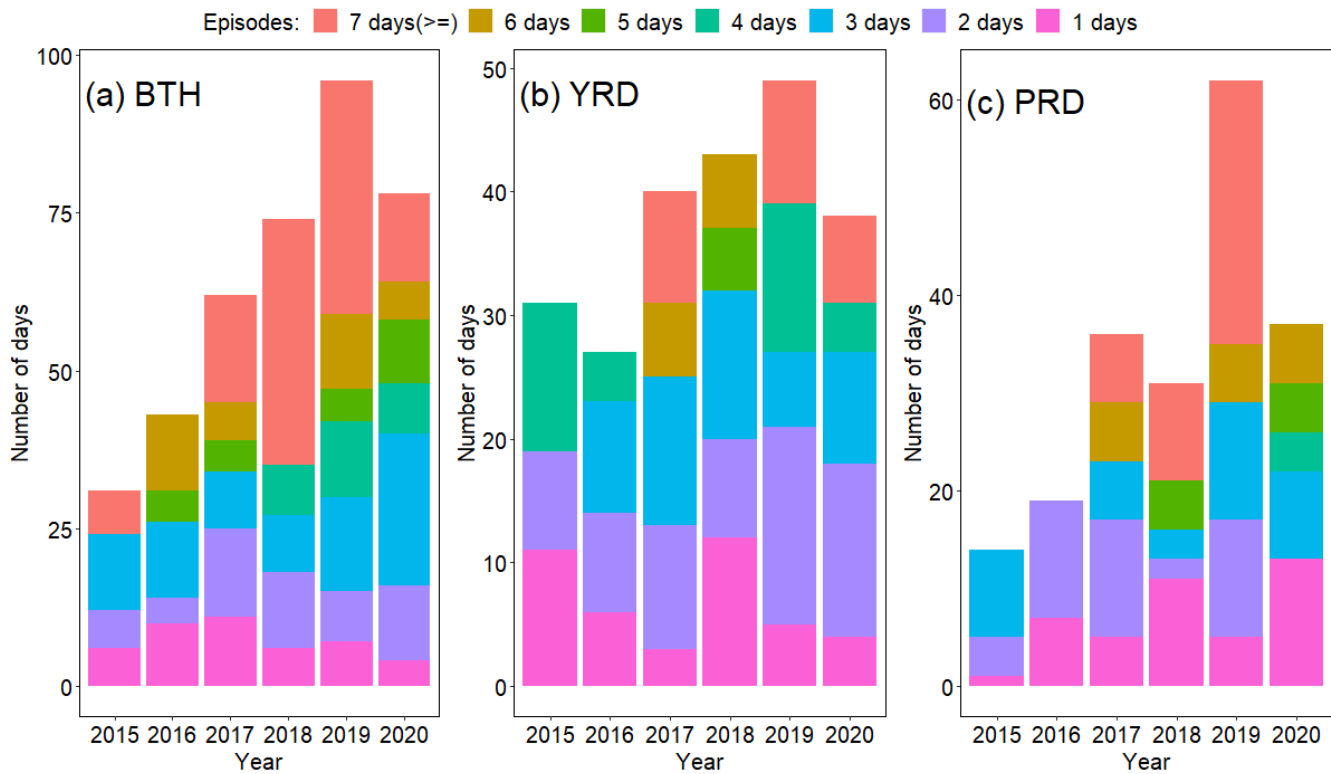


Figure 2: Contributions from the O₃-exceeding days (red) and non-O₃-exceeding days (green) to the annual mean concentration of maximum daily 8-hour average O₃ (black) in BTH (a), YRD (b) and PRD (c).



640 **Figure 3: Annual numbers of various consecutive O₃-exceeding days in BTH (a), YRD (b) and PRD (c). Individual colors denote different numbers of consecutive O₃-exceeding days.**

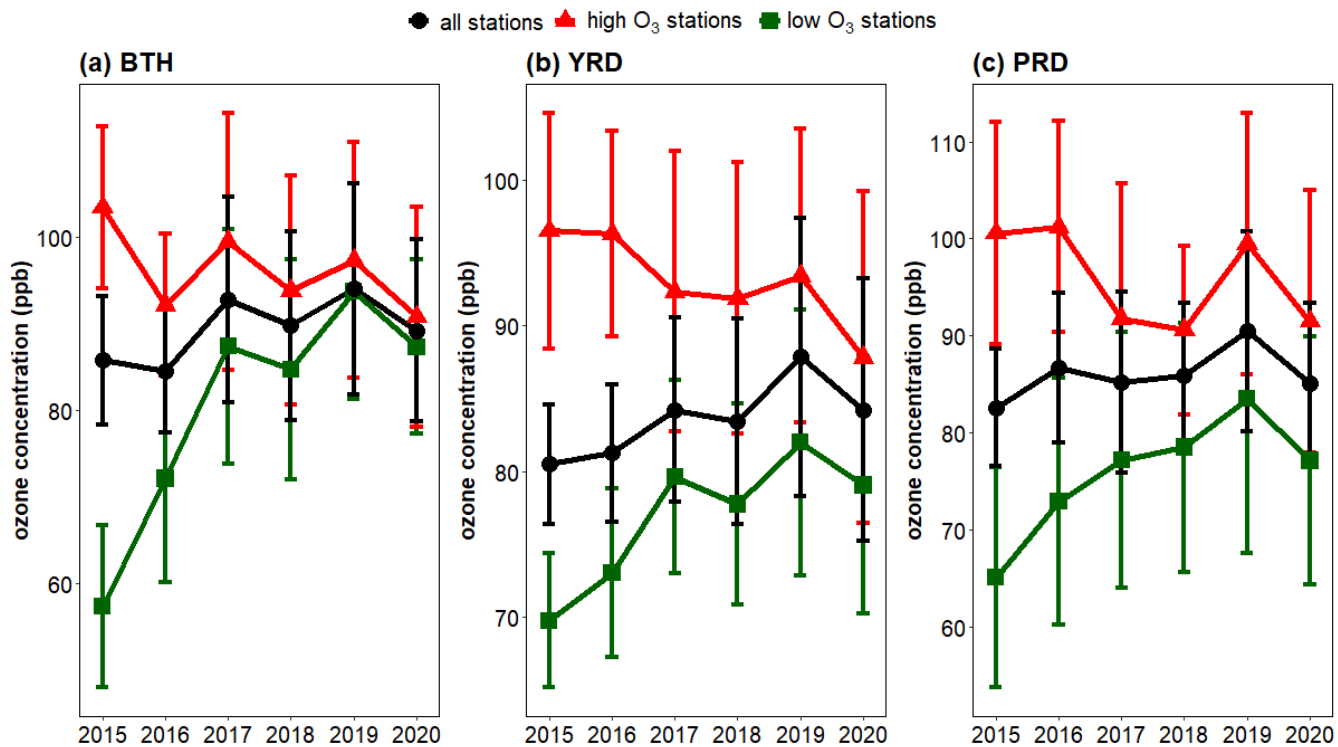
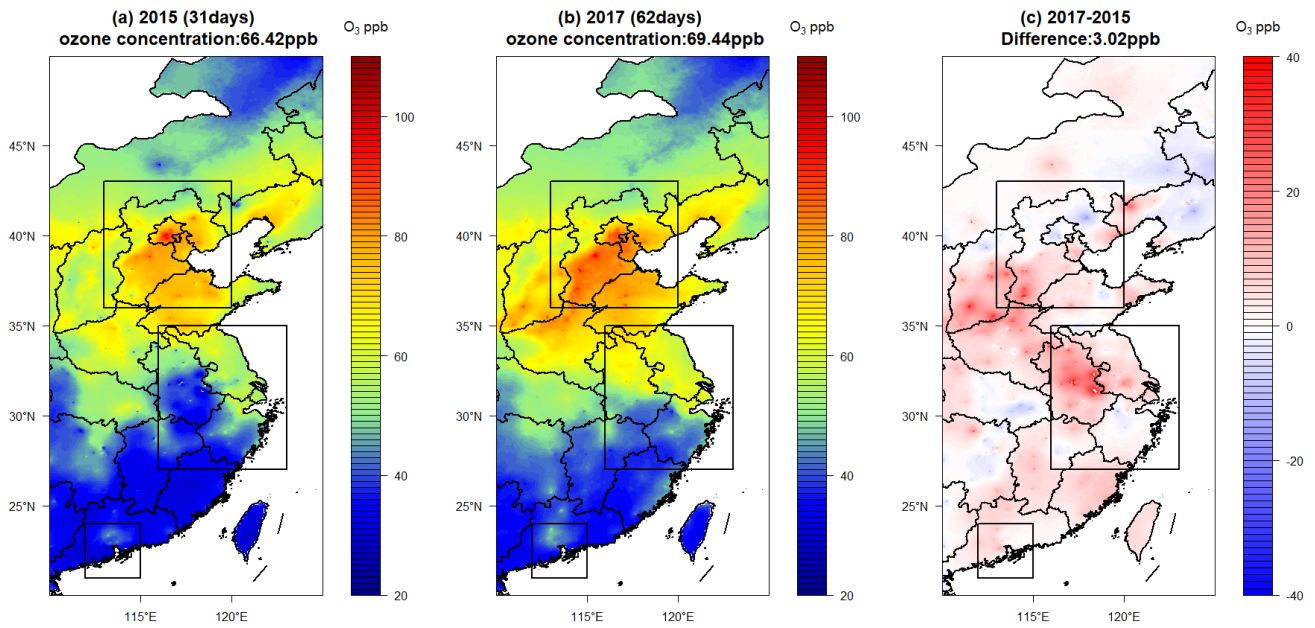
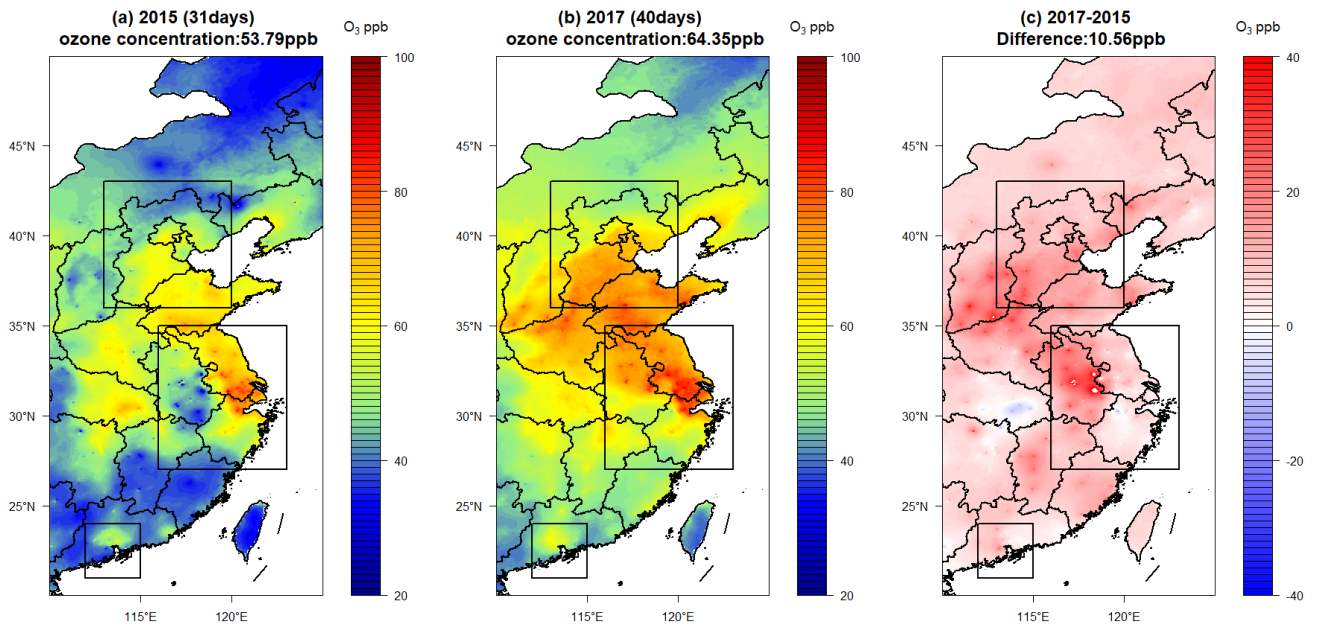


Figure 4: Annual mean concentrations (2015-2020) of maximum daily 8-hour average O₃ during O₃-exceeding days for all stations (black), high O₃ stations (red) and low O₃ stations (green) in BTH (a), YRD (b) and PRD (c).

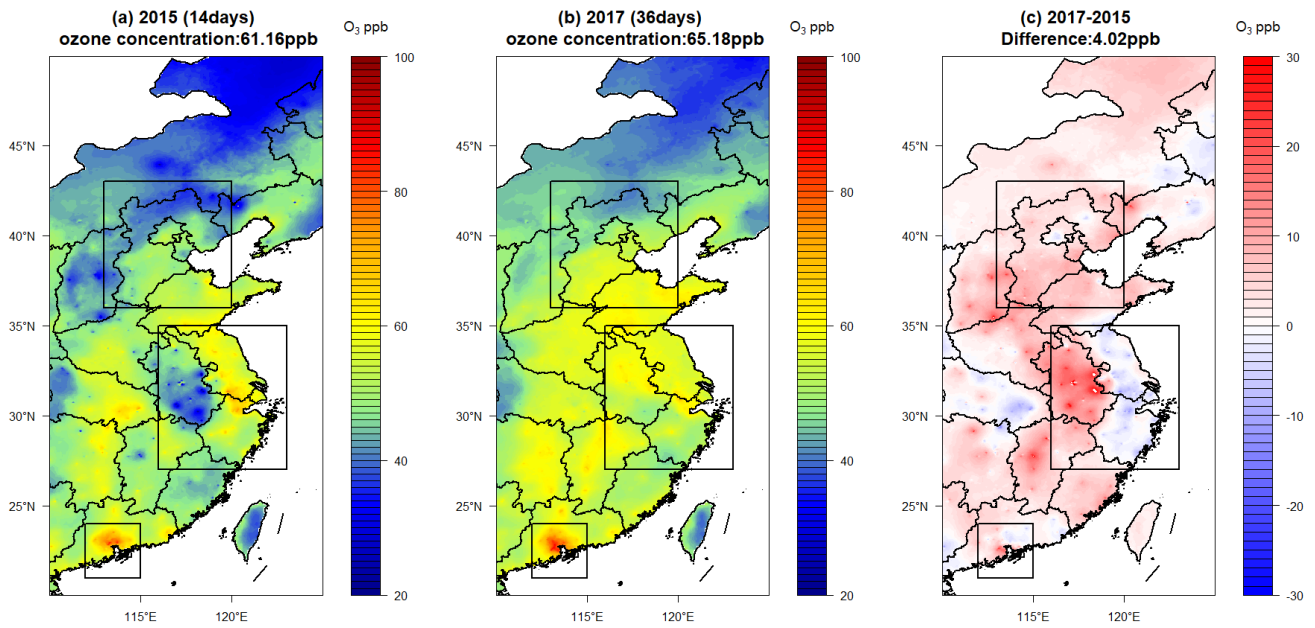


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Figure 5: Spatial distribution of annual mean concentrations of maximum daily 8-hour average O_3 for O_3 -exceeding days in BTH in 2015 (a), 2017 (b) and their difference (2017 - 2015) (c). The top, middle and bottom rectangle boxes denote BTH, YRD and PRD districts, respectively. The number inside the parenthesis behind 2015 or 2017 denotes the number of O_3 -exceeding days.



650 **Figure 6: Spatial distribution of annual mean concentrations of maximum daily 8-hour average O_3 for O_3 -exceeding days in YRD in 2015 (a), 2017 (b) and their difference (2017 - 2015) (c). The top, middle and bottom rectangle boxes denote BTH, YRD and PRD districts, respectively. The number inside the parenthesis behind 2015 or 2017 denotes the number of O_3 -exceeding days.**



655 **Figure 7: Spatial distribution of annual mean concentrations of maximum daily 8-hour average O_3 for O_3 -exceeding days in PRD in 2015 (a), 2017 (b) and their difference (2017 - 2015) (c). The top, middle and bottom rectangle boxes denote BTH, YRD and PRD districts, respectively. The number inside the parenthesis behind 2015 or 2017 denotes the number of O_3 -exceeding days.**

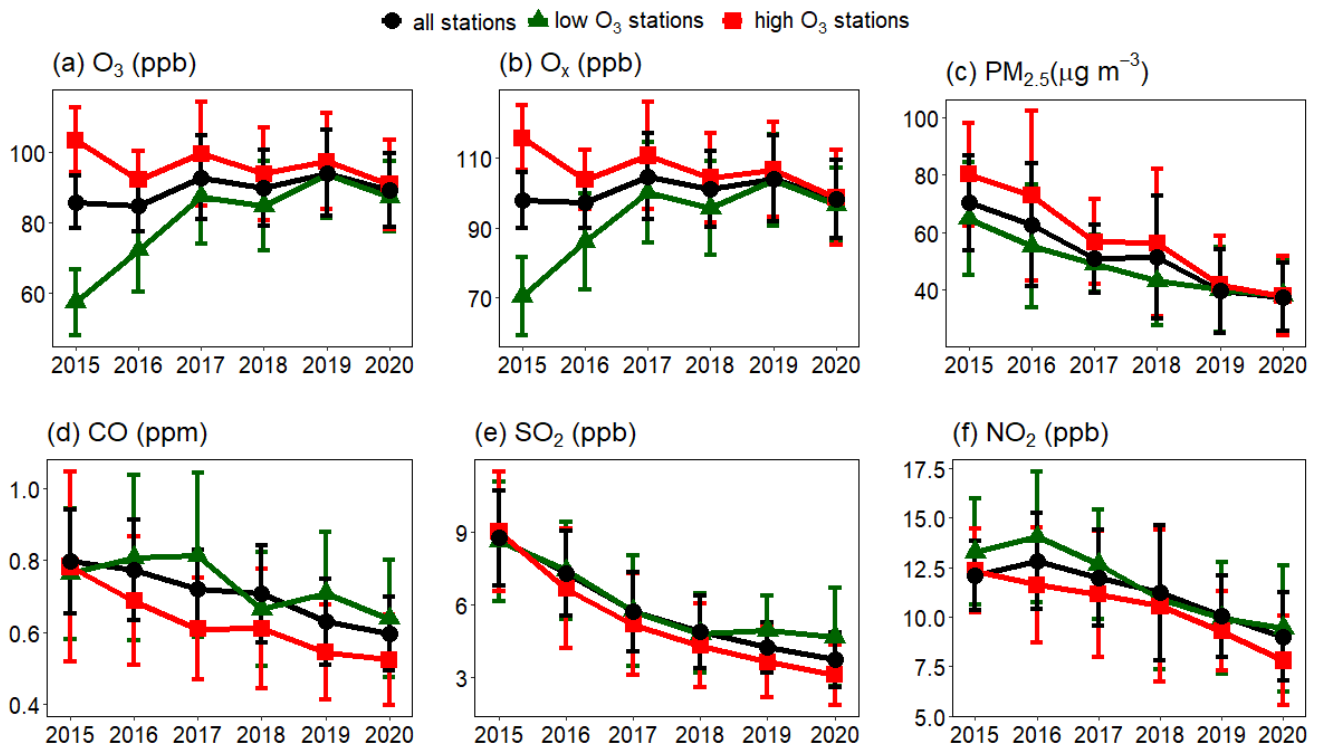
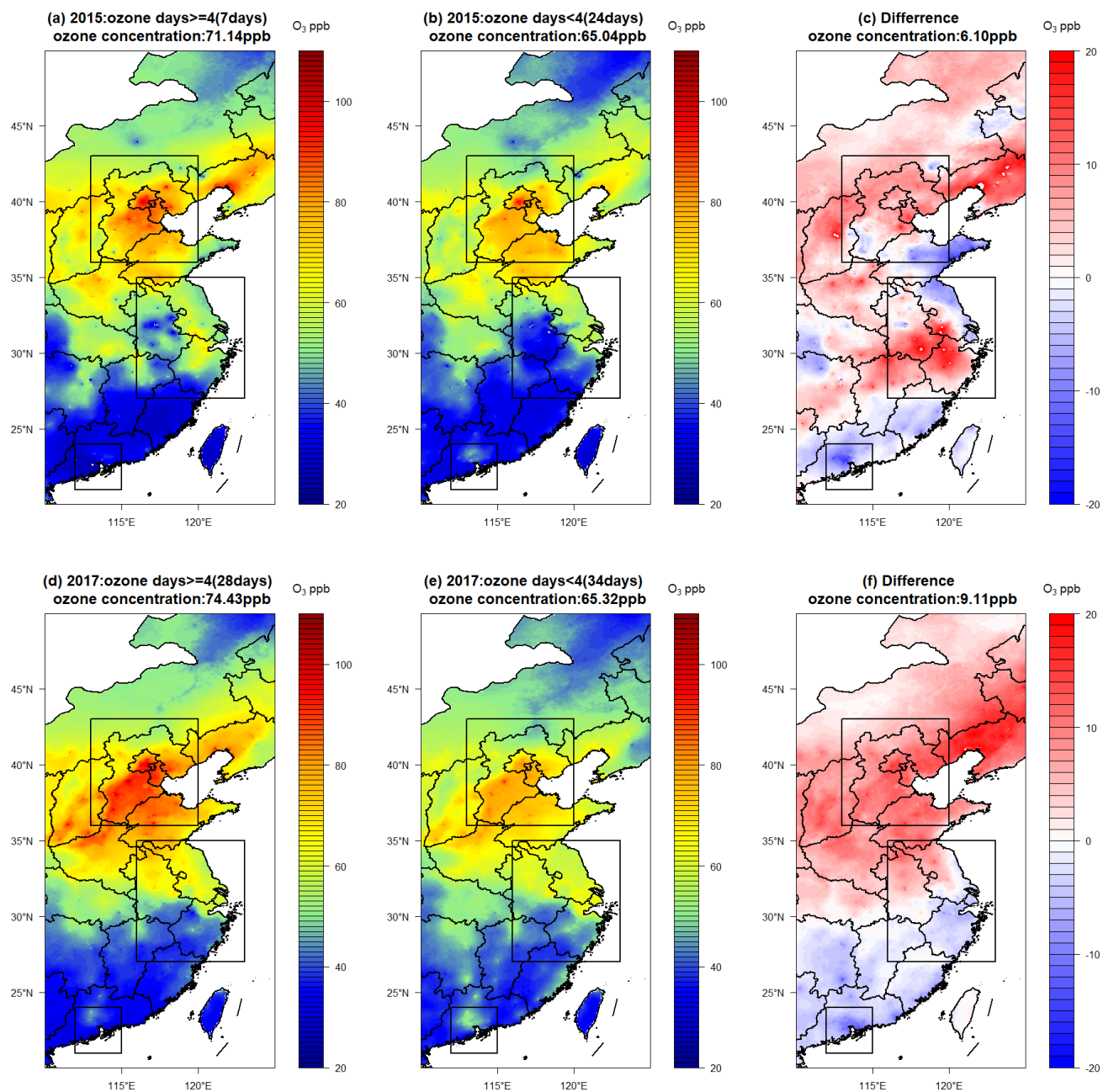
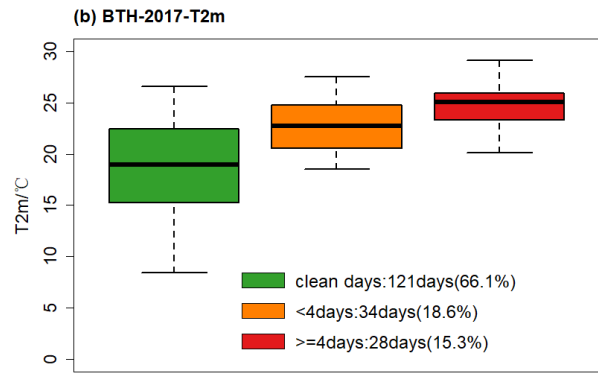
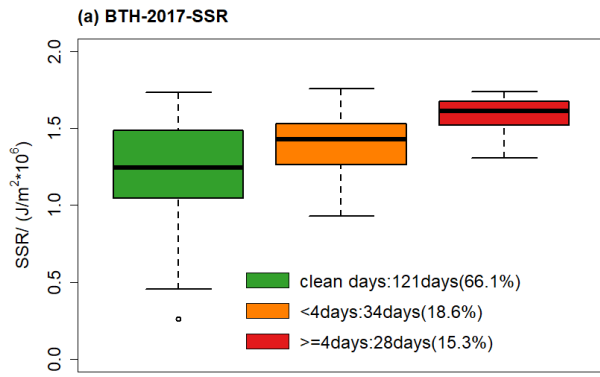


Figure 8: Annual mean concentrations of maximum daily 8-hour average O₃ in BTH during O₃-exceeding days for all stations (black), high O₃ stations (red) and low O₃ stations (green) (a), same as (a) but for O_x (b), PM_{2.5} (c), CO (d), SO₂ (e), NO₂ (f).

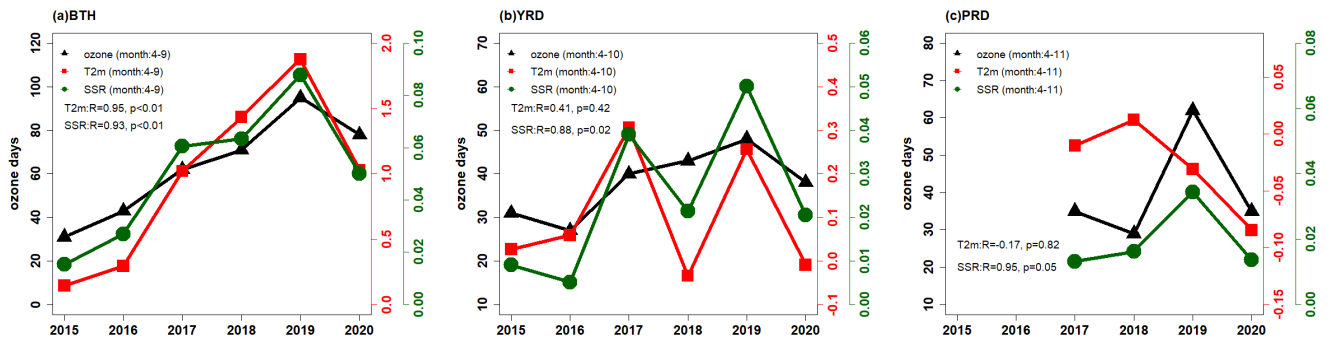


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Figure 9: Spatial distribution of daily mean MDA8 O₃ of O₃-exceeding days in BTH for O₃ episodes with four or more consecutive O₃-exceeding days in 2015 (a), O₃ episodes with less than four consecutive O₃-exceeding days in 2015 (b), and (a minus b) (c); (d, e and f) are the same as (a, b and c), respectively, but for 2017.



665 **Figure 10: Surface solar radiation (SSR) (a) and temperature (T2m) (b) in BTH in April–September 2017 for four episodes with four or more consecutive O₃-exceeding days (red), clean days (non-O₃-exceeding days) (green) and O₃ episodes with less than four consecutive O₃-exceeding days (orange).**



670 **Figure 11: Correlations among annual O₃-exceeding days, surface solar radiation (SSR) and temperature (T2m) in BTH (a), YRD (b) and PRD (c).**

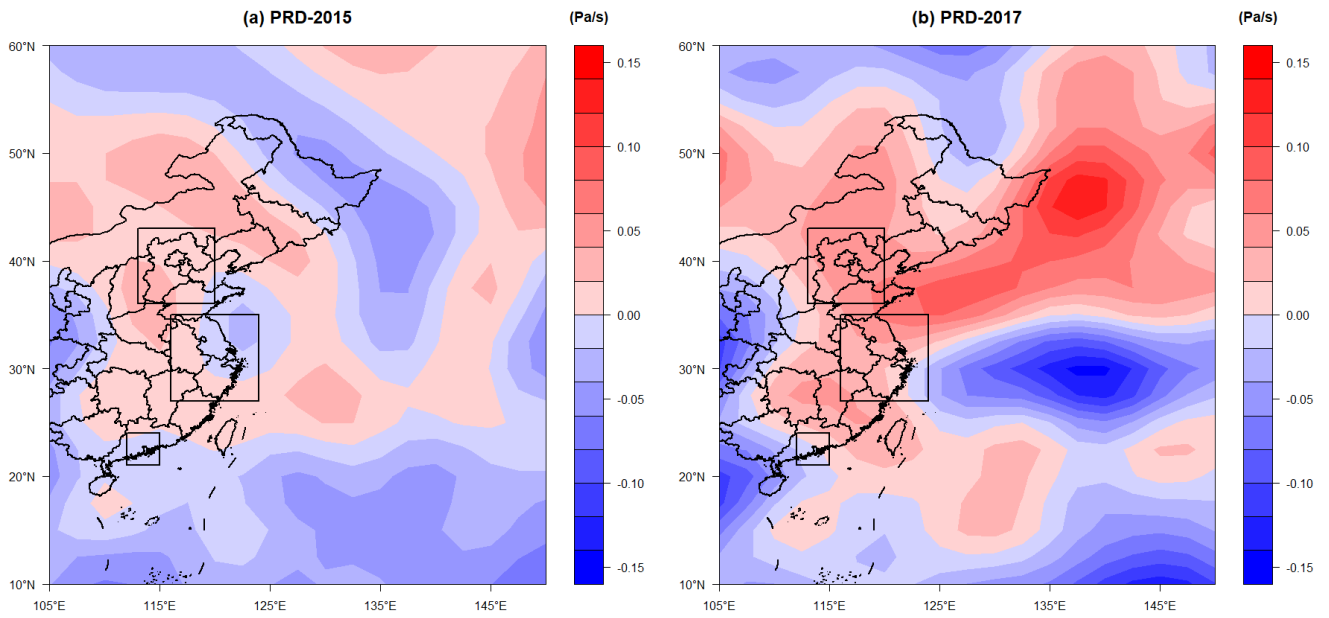
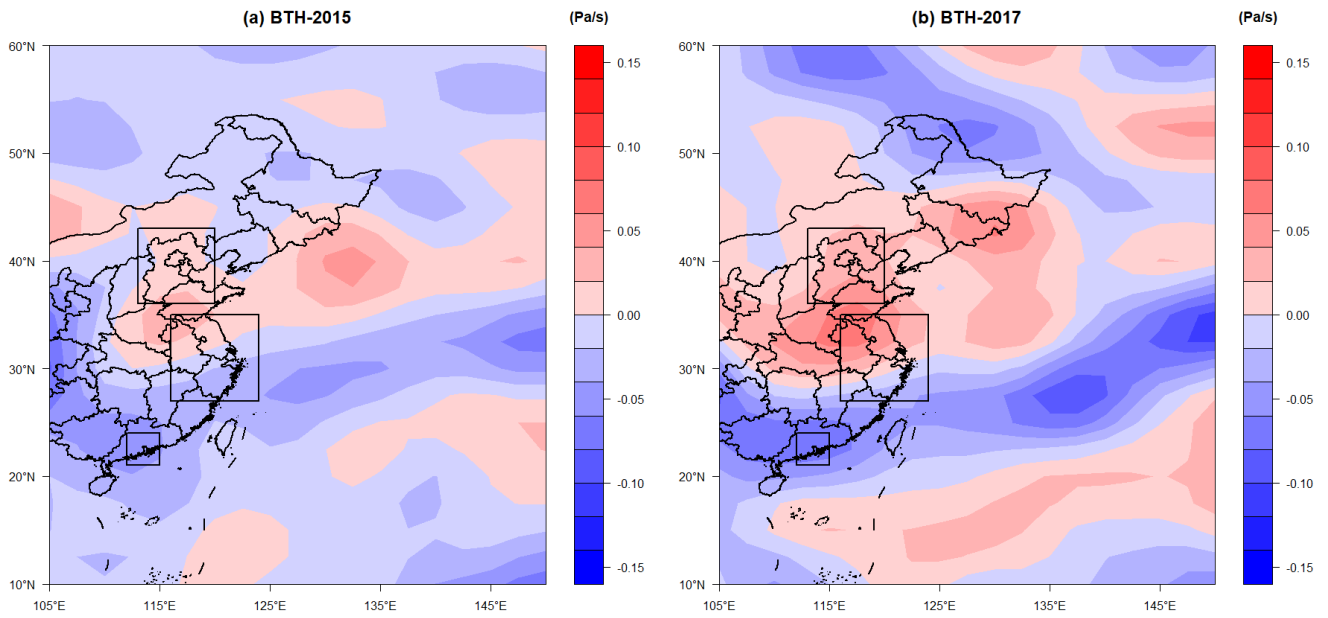


Figure 12: Mean vertical velocity at 850hPa during O₃-exceeding days in PRD in 2015 (a) and during episodes with four or more consecutive O₃-exceeding days in 2017 (b).



675 **Figure 13: Mean vertical velocity at 850 hPa during O₃-exceeding days in BTH in 2015 (a) and during episodes with four or more consecutive O₃-exceeding days in 2017 (b).**

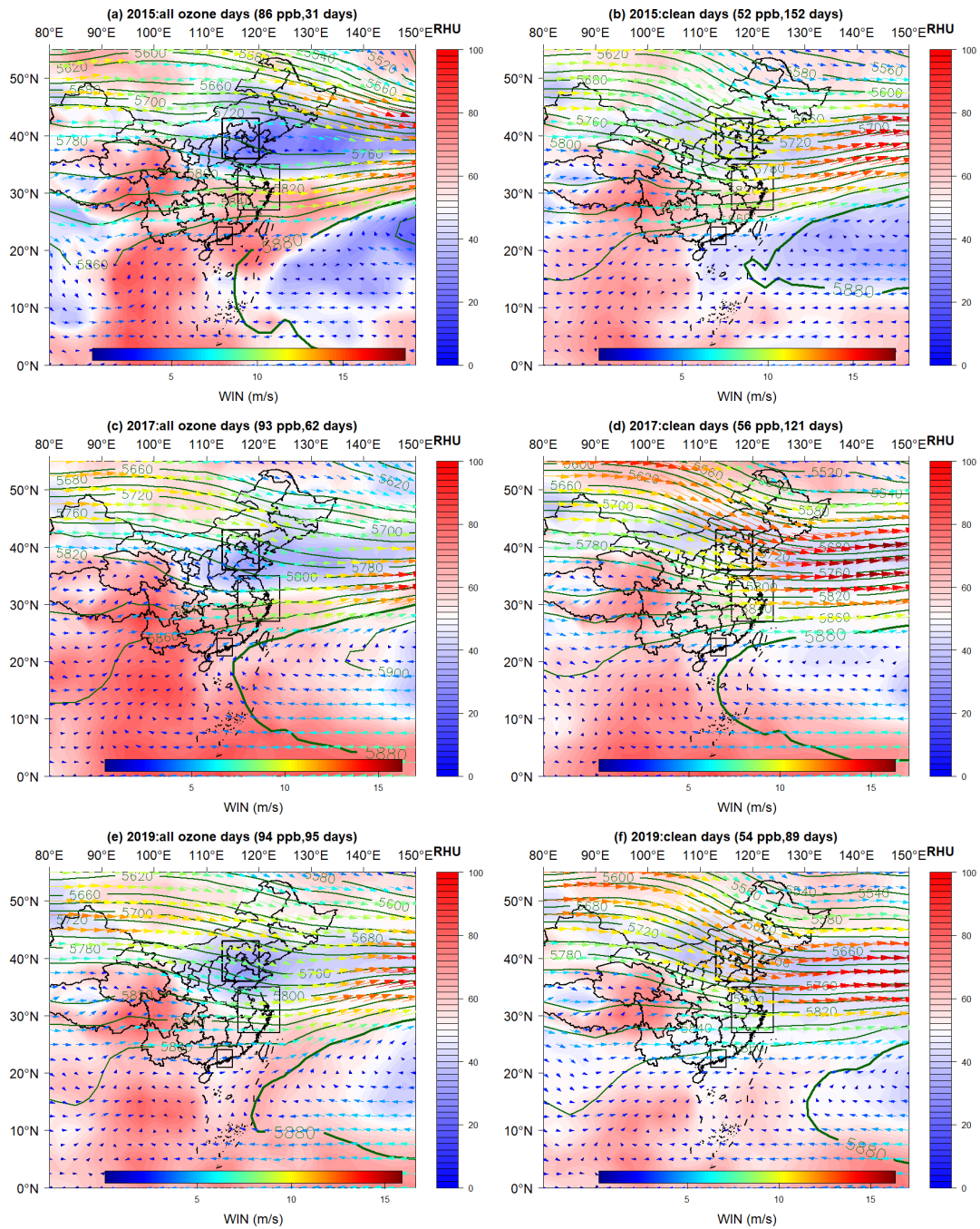


Figure 14: Composite 500 hPa geopotential height contours, humidity and winds in BTH in April-September for O₃-exceeding days in 2015 (a), clean days in 2015 (b), O₃-exceeding days in 2017 (c), clean days in 2017 (d), O₃-exceeding days in 2019 (e) and clean days in 2019 (f).

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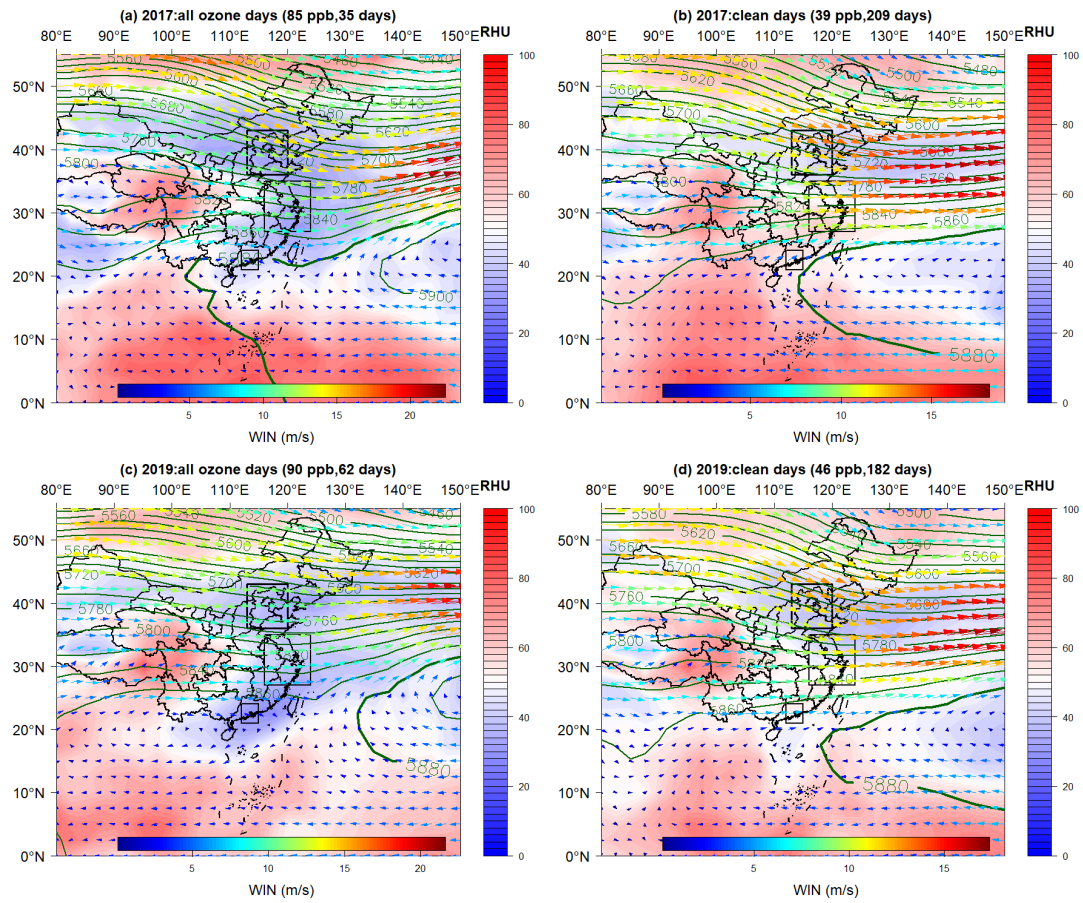
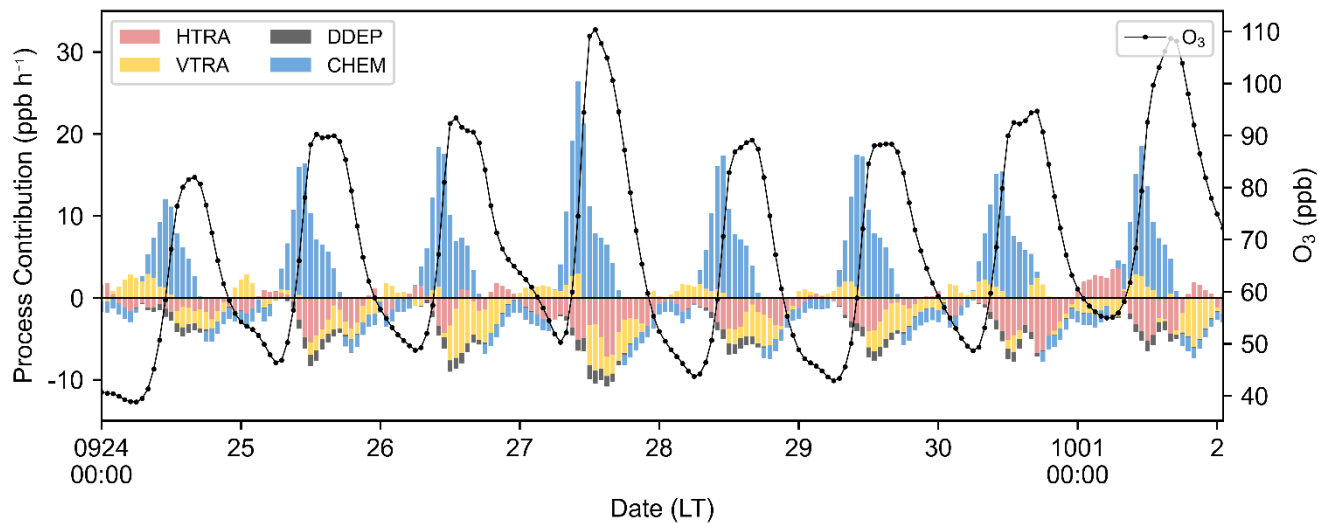


Figure 15: Composite 500 hPa geopotential height contours, humidity and winds in PRD in April-November for O₃-exceeding days in 2017 (a), clean days in 2017 (b), O₃-exceeding days in 2019 (c), clean days in 2019 (d).



685 **Figure 16: Time series of individual processes contributing to O₃ budget in PRD calculated by the WRF-CMAQ model for the O₃ episode of September 24–October 1, 2019. The black line (O₃) represents the averaged O₃ concentrations in the layers below 1260m. Where DDEP denotes dry deposition, CHEM denotes chemical processes, HTRA denotes the horizontal transport and VTRA denotes the vertical transport.**