

# 1 WRF-Comfort: Simulating micro-scale variability of outdoor heat 2 stress at the city scale with a mesoscale model

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10 **Abstract.** Urban overheating, and its ongoing exacerbation due to global warming and urban development, leads to increased  
11 exposure to urban heat and increased thermal discomfort and heat stress. To quantify thermal stress, specific indices have been  
12 proposed that depend on air temperature, mean radiant temperature (MRT), wind speed, and relative humidity. While  
13 temperature and humidity vary on scales of hundreds of meters, MRT and wind speed are strongly affected by individual  
14 buildings and trees, and vary at the meter scale. Therefore, most numerical thermal comfort studies apply micro-scale models  
15 to limited spatial domains (commonly representing urban neighborhoods with building blocks) with resolutions on the order  
16 of 1 m and a few hours of simulation. This prevents the analysis of the impact of city-scale adaptation/mitigation strategies on  
17 thermal stress and comfort. To solve this problem, we develop a methodology to estimate thermal stress indicators and their  
18 subgrid variability in mesoscale models - here applied to the multilayer urban canopy parametrization BEP-BEM within the  
19 WRF model. The new scheme (consisting of three main steps) can readily assess intra-neighborhood scale heat stress  
20 distributions across whole cities and for time scales of minutes to years. The first key component of the approach is the  
21 estimation of MRT in several locations within streets for different street orientations. Second, mean wind speed, and its subgrid  
22 variability, are parameterized as a function of the local urban morphology based on relations derived from a set of  
23 microscale LES and RANS simulations across a wide range of realistic and idealized urban morphologies. Lastly, we compute  
24 the distributions of two thermal stress indices for each grid square combining all the subgrid values of MRT, wind speed, air  
25 temperature, and absolute humidity. From these distributions, we quantify the high and low tails of the heat stress distribution  
26 in each grid square across the city, representing the thermal diversity experienced in street canyons. In this contribution, we  
27 present the core methodology as well as simulation results for Madrid (Spain), which illustrate strong differences between heat  
28 stress indices and common heat metrics like air or surface temperature, both across the city and over the diurnal cycle.

30 **1 Introduction**

31 The combination of urban development and climate change has increased heat exposure in cities in recent decades (Tuholske  
32 et al., 2021) and a continuation of these trends in the 21st century would be difficult to offset locally from an air temperature  
33 perspective (Broadbent et al., 2020; Krayenhoff et al., 2018; Zhao et al., 2021). Adaptation options that target contributions to  
34 heat exposure other than the air temperature, such as radiation (e.g., via shade) and wind (e.g. via improved street ventilation),  
35 should therefore be considered. Quantification of these contributions relative to air temperature requires the application of  
36 comprehensive thermo-physiological heat stress metrics such as the Universal Thermal Climate Index, UTCI (Jendritzky et  
37 al., 2012), the Physiological Equivalent Temperature, PET (Höppe, 1999), or the Standard Effective Temperature, SET (Gagge  
38 et al., 1986). Moreover, exposure to heat hazards is moderated by infrastructure-based and social/mobility-based adaptations  
39 to heat, and by buildings and associated cooling mechanisms. Here, the focus is the development of a tool to quantify the  
40 outdoor component of heat exposure in cities, accounting for all relevant meteorological variables.

41 Heat exposure in urban areas is affected by several meteorological variables that vary on different spatial and temporal scales  
42 (Nazarian et al., 2022). While temperature and humidity vary on spatial scales on the order of hundreds of meters, shortwave  
43 and longwave radiation and wind speed are strongly affected by individual buildings and vary at the scale of a few meters. For  
44 this reason, most numerical thermal comfort studies in urban areas apply micro-scale models with resolutions on the order of  
45 one m and spatial domains that are limited to an urban block or neighborhood (Nazarian et al., 2017; Zhang et al., 2022; Geletić  
46 et al., 2018). While these studies include substantial detail at the micro-scale, they are very expensive computationally and  
47 therefore can be applied only to a few neighborhoods and they neglect the interactions with larger scale meteorological  
48 phenomena (e.g., land/sea breezes, mountain/valley winds, urban breezes) that often play a relevant role in outdoor thermal  
49 comfort and its variation across cities. On the other hand, contemporary meso-scale numerical models can be applied to the  
50 whole urban area and surrounding regions, and therefore capture these larger-scale phenomena, but have spatial resolutions of  
51 several hundred meters at best. These models use a grid mesh that does not resolve buildings and is therefore too coarse to  
52 capture the fine-scale variation of radiation and wind flow of relevance to outdoor heat exposure and ultimately thermal  
53 comfort.

54 The objective of this work is to fill the aforementioned gap by developing a model that includes the most crucial capabilities  
55 of micro-scale assessments of thermal exposure within meso-scale models. This new model will quantify the spatial variability  
56 (i.e., statistical representation of the microscale distribution) for longwave and shortwave radiation as well as wind speed  
57 within each meso-scale grid square. Subsequently, it will capture the range of thermal exposure, as quantified by the UTCI  
58 and SET thermal stress metrics, within each urban grid square across a city at each time of day. The focus here is on the *range*  
59 of thermal exposure, such that we identify the cool and hot spots within the grid cell without having to resolve the entire spatial  
60 distribution. We argue that this represents the most crucial information for heat management and urban design interventions,  
61 as it identifies whether people can move and search for optimal thermal conditions. For example, if hot spots are experiencing  
62 extreme heat stress but the cool spots are at slight heat stress, pedestrians have the opportunity, and autonomy, to seek shade

63 and thermal respite (i.e., temporal and spatial autonomy as described in Nazarian et al. (2019)). Conversely, if the conditions  
64 in the cool spot are already in extreme heat stress, this can be used to inform urban design interventions or heat advisories to  
65 vulnerable populations to avoid being outside at that place and time. Overall, representing the range of heat exposure at the  
66 neighborhood scale while covering regional-scale phenomena is key to human-centric assessments of urban overheating  
67 (Nazarian et al., 2022).

68 The new model is embedded in the multi-layer urban canopy parameterization BEP-BEM (Martilli et al., 2002; Salamanca et  
69 al., 2010) which simulates the local-scale meteorological effects of the grid-average urban morphology within the Weather  
70 Research and Forecasting (WRF) mesoscale model (Skamarock et al., 2019 version 4.3 has been used in this study). Here,  
71 BEP-BEM is extended to quantify the spatial variation of the mean radiant temperature and wind speed within the grid square  
72 at the pedestrian level. To our knowledge, three schemes in the published literature have attempted to capture thermal exposure  
73 in an urban canopy model. Pigliautile (2020) implemented a scheme to estimate human thermal exposure in the Princeton  
74 Single-Layer Urban Canopy Model. However, the scheme has not been run within a mesoscale model. Jin et al. (2022) calculate  
75 urban mean radiant temperature (MRT) in a mesoscale model, while Lemonsu (2015) and Leroyer et al. (2018) assess UTCI  
76 in mesoscale modeling applications within Paris and Toronto, respectively. Moreover, Giannaros et al (2018, 2023), made an  
77 offline coupling of WRF-BEP\_BEM with RayMan (Matzarakis et al. 2007). However, none of these approaches account for  
78 the within-grid spatial variation of wind speed, and their assessment of sub-grid spatial variation of radiation exposure (i.e.,  
79 mean radiant temperature) is limited. Here, we further extend the BEP-BEM model embedded in the WRF meso-scale model  
80 to overcome these limitations and better assess spatial variation of thermal exposure within each urban grid square.

81 In section 2, the methodology is described in detail, with a focus on model development and implementation in WRF. In  
82 Section 3, we present an example of the type of outputs that can be produced. Conclusions are in section 4.

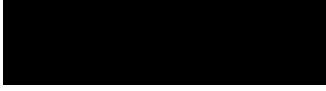
## 83 **2 Methodology**

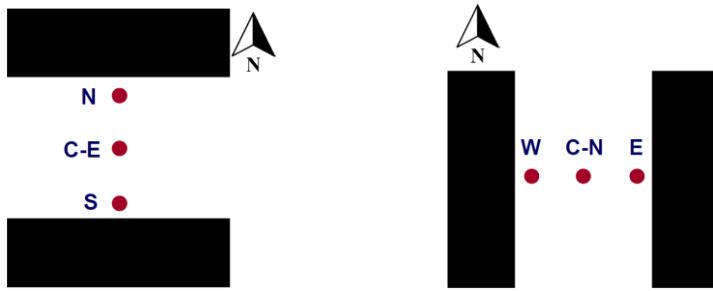
84 The most complete thermal stress indices invariably depend on four meteorological variables: air temperature, mean radiant  
85 temperature (MRT), relative humidity, and wind speed. Among these, MRT and wind speed have the largest spatial variability  
86 in the urban canopy, and this variability is often captured with 3D micro-scale models of urban airflow and radiative heat  
87 transfer. At the meso-scale, however, it is not feasible to incorporate such models, motivating the simplified urban canopy  
88 parameterizations developed here. Below we detail how the BEP-BEM urban canopy model is modified to a) introduce a  
89 simplified model for MRT variation within a meso-scale grid cell (Sec. 2.1) and b) parameterize airflow variability (Sec. 2.2)  
90 in the urban canopy within a grid cell, and make a simple estimate of air temperature variability. These meteorological  
91 parameters are then used to estimate the sub-grid scale variation of thermal stress indices (Sec. 2.3), namely SET and UTCI,  
92 as two of the most commonly used indices for outdoor environments (Potchter et al 2018). Lastly, we discuss how multi-scale

93 temporal and spatial variabilities in thermal exposure can be effectively communicated using the outcomes of the updated  
94 WRF-BEP-BEM model.

95 **2.1 A simplified model for MRT variability in the urban canopy**

96 The mean radiant temperature is a measure of the total radiation flux absorbed by the human body, including both shortwave  
97 (from the sun, either directly or after reflection on the walls or road) and longwave (emitted from solid bodies like walls or  
98 road, or from the sky) radiation. Whether pedestrians are shaded or in the sunshine, as well as their distance from warm surfaces  
99 emitting radiation, is therefore very important. BEP-BEM applies a simple urban morphology: two street canyons of different  
100 orientations, each with the same street width and building height distribution on each side of the canyon (Martilli et al. 2002).  
101 To capture the within-grid spatial extremes of mean radiant temperature, we assess pedestrian locations at the center of the  
102 street for two canyon orientations considered in BEP-BEM and at positions located at a distance of 1.5 m from the building  
103 wall on each side of the street, representing the sidewalks. Thus, there are 6 positions (three for each street direction) in each  
104 urban grid square where we compute the mean radiant temperature (shown for the example of North-South and East-West  
105 streets in Fig. 1). For shortwave and longwave radiation exchange, the standard BEP view factor and shading routines (Martilli  
106 et al. 2002) are used to estimate the amount of shortwave (direct and diffuse) and longwave radiation reaching a vertical  
107 segment 1.80 m tall and located in each of the six positions previously mentioned (Fig. 1, Appendix A). Reflection of shortwave  
108 radiation and emission and reflection of longwave radiation from both building walls and the street surface are accounted for  
109 via these view factors. The pedestrian is “transparent” from the perspective of the urban facets, meaning that its presence does  
110 not alter the shortwave and longwave radiation reaching the building walls and road. The mean radiant temperature is computed  
111 by weighting the radiation reaching each side of the vertical segment by 0.44, and the radiation reaching the downward- and  
112 upward-facing (at 1.80 m height) surfaces of the pedestrian by 0.06 each. This approach follows the six-directional weighting  
113 method (Thorsson et al. 2007) and aggregates the four lateral weightings of 0.22 into two lateral weightings of 0.44 since BEP-  
114 BEM is a two-dimensional model (e. g. the streets are considered infinitely long). Namely,

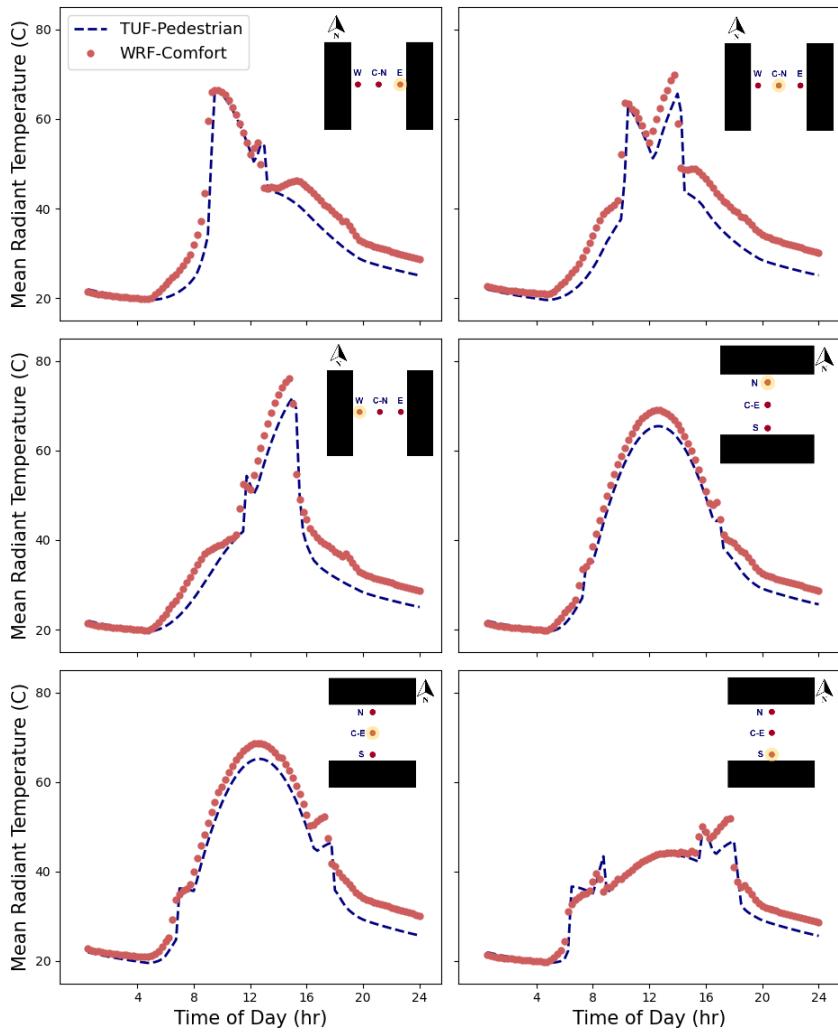
115  (1)  
116 where, for a N-S oriented street,  $i=1,2$  are for the vertical sides of the pedestrian looking East, and West respectively, and  $i=3,4$   
117 are for the top and bottom. Therefore,  $W_{1,2}=0.44$ , while  $W_{3,4}=0.06$ , while the absorptivity of the pedestrian in the shortwave  
118 and longwave,  $a_K$  (the absorption coefficient for short-wave radiation of the human body) and  $a_L$  (the absorption coefficient  
119 for long-wave radiation, or emissivity, of the human body), respectively, are  $a_K=0.7$ , and  $a_L=0.97$ ,  $K_{1,2}$  and  $L_{1,2}$  are the short  
120 and longwave radiation reaching the vertical segment, and  $K_{3,4}$  and  $L_{3,4}$  are short and longwave radiation reaching the top and  
121 bottom respectively, and  $\sigma$  is the Stefan-Boltzmann constant (see Appendix A for details about how the radiation components  
122 are computed).



**Figure 1:** Two street directions (left: E-W canyon, right: N-S canyon) and pedestrian locations considered for Mean Radiant Temperature calculations.

123 The diurnal progression of the mean radiant temperature computed by this new model in BEP-BEM is subsequently compared  
 124 with that obtained from TUF-Pedestrian, a more detailed three-dimensional model that has been evaluated against  
 125 measurements (Lachapelle et al. 2022; Jiang et al. 2023). TUF-Pedestrian is configured with identical input parameters and  
 126 meteorological forcing, and with long canyons that approximate the two-dimensional BEP-BEM canyon geometry. The new  
 127 model clearly captures the relevant details of the diurnal progression of MRT at all six locations (Fig. 2), with a mean absolute  
 128 difference of 3.4 K, and a root mean square difference of 4.3 K across all locations. A comparison of the shortwave radiation  
 129 loading on the pedestrian between the two models reveals very good agreement (Appendix B Fig. B1, B2), considering the  
 130 highly simplified urban morphology used by BEP-BEM, with biggest errors limited to short periods of time; thus, most of the  
 131 model disagreement arises from differences between longwave loading on the pedestrian as a result of different methods for  
 132 computation of surface temperature between the models. Overall, the new model of mean radiation temperature in BEP-BEM  
 133 provides satisfactory results.

134



135  
 136 **Figure 2:** Comparison of diurnal variation of Mean Radiant Temperature (MRT) between the new model in BEP-BEM and TUF-  
 137 Pedestrian for each of the six locations in Fig. 1. [TUF pedestrian acts here as a reference.](#)

## 138 2.2 Parameterize airflow variability in the urban canopy

139 Mesoscale models solve conservation equations for the three components of momentum. From these, it is possible to derive  
 140 the spatially averaged wind velocity in each grid cell, at the grid resolution of the mesoscale model, commonly of the order of  
 141 300m-1km. The spatially averaged wind velocity in the urban canopy  $\langle V \rangle$ , close to the pedestrian height (~2.5m), is the square  
 142 root of the sum of the spatial average of the two horizontal components  $u$ , and  $v$ , (neglecting the vertical component, which is  
 143 usually at least one or two orders of magnitude smaller than the horizontal),

$$144 \langle V \rangle = \frac{1}{V_{air}} \sqrt{\left( \int_{V_{air}} u dV \right)^2 + \left( \int_{V_{air}} v dV \right)^2} \quad (2)$$

145 where here  $V_{air}$  is the volume of the grid cell occupied by air (e. g. without the buildings)

146 However, the wind velocity calculated in mesoscale models is different from the average wind speed that would be experienced  
 147 by a person in the grid cell. This is better represented by the spatial average of the wind speed  $\langle U \rangle$  (e. g. the modulus of the  
 148 vector), written as

$$149 \langle U \rangle = \frac{1}{V_{air}} \int_{V_{air}} \sqrt{u^2 + v^2} dV \quad (3)$$

150 To assess the impact of airflow on human thermal comfort, the wind speed should be estimated from the wind velocity  
 151 computed in the mesoscale models. Additionally, it is critical to parameterize and estimate the spatial variability of mean wind  
 152 speed in the urban canopy. Accounting for these factors, the range of wind speed variability at the pedestrian level is estimated,  
 153 which is critical for the quantification of spatial variability of outdoor thermal stress and comfort.

154 Here, we describe the parameterization of a) wind speed-to-velocity ratio and b) wind speed distribution, based on urban  
 155 density parameters. Data are considered from over 173 microscales CFD simulations of urban airflow over realistic and  
 156 idealized urban configurations, spanning a wide range of building plan area ( $\lambda_p$ ), frontal area ( $\lambda_f$ ), and wall area ( $\lambda_w$ ) densities  
 157 representative of realistic urban neighbourhoods in different types of cities. CFD simulations are conducted using 162 large-  
 158 eddy simulations (LES) and 11 Reynolds-averaged Navier-Stokes (RANS) schemes detailed in Appendix C.

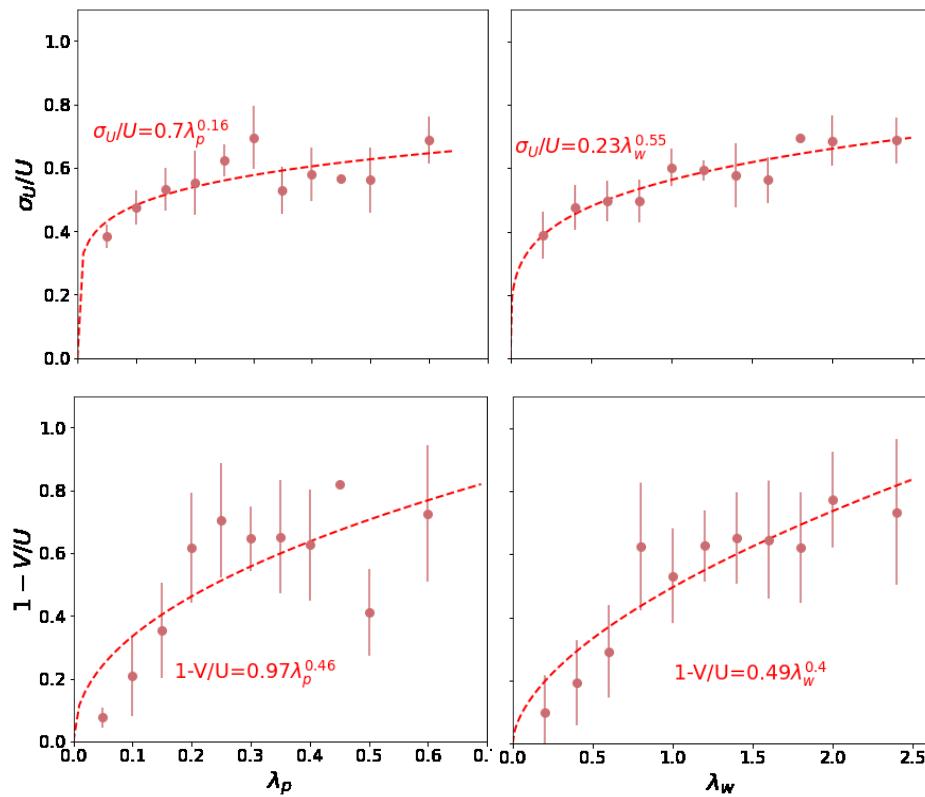
159 Mean wind velocity  $\langle V \rangle$ , speed  $\langle U \rangle$  and its spatial standard deviation ( $\sigma_U$ ) are computed at a horizontal cross-section at  
 160 pedestrian height for each CFD simulation and used for deriving parameterizations (Fig. 3). An additional data point is added  
 161 at  $\lambda_p=\lambda_w=0$ , ensuring that wind speed is equal to wind velocity, and its standard deviation is set to zero, for the non-urban case.

162 It is important to remark here, that we are dealing with the standard deviation of the spatial distribution of the mean wind  
 163 speed. With the term *mean* we indicate the result of an ensemble (over many realizations) or time average (over time scales  
bigger/larger than the turbulence time scale, but smaller than the time scale of the mesoscale motions), but not a spatialee  
average. The urban canopy, in fact is spatially heterogeneous, and, for this reason, the time and ensemble averages are different  
than the spatial average. Only when  $\lambda_p=\lambda_w=0$  (e. g. there are no buildings), and the horizontal homogeneity is recovered,  
and must therefore the variability must be zero. This  $\sigma_U$ , therefore should not be confounded with the turbulent  $\sigma$ , that indicates

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168 the variability in instantaneous wind speed induced by turbulent motions, that which indeed is not zero even when there are no  
 169 buildings.

170



171 **Figure 3:** Relationship between  $1-V/U$  (bottom row), and  $\sigma_U/U$  (top row), and two morphological parameters,  $\lambda_p$  (left column), and  
 172  $\lambda_w$  (right column) based on the CFD simulations. Dots represent the average of the value among all the simulations that share the same  
 173 morphological parameter, and the vertical bar indicates the standard deviation. The dashed line and the formula indicate the best fit.

174

175 Parameterizations are derived (shown in Fig. 3) for two density parameters ( $\lambda_p=Ap/Atot$ , and  $\lambda_w=Aw/Atot$ , where  $Ap$  is the  
 176 area of the horizontal surface occupied by buildings, or the roof area,  $Aw$  is the area of vertical (wall) surfaces, and  $Atot$  is the  
 177 total horizontal area). We find that  $\lambda_w$  better predicts mean wind speed and its spatial variability at the pedestrian height,  
 178 because it represents both horizontal and vertical heterogeneities in the urban canopy. Note that  $\lambda_F$  has not been included in  
 179 the study, given the difficulty to estimate it for real urban areas, and to translate it to the simplified 2D urban morphology used  
 180 by BEP-BEM. In any case,  $\lambda_F$  is closely related to  $\lambda_w$ . Therefore, the following parameterizations are implemented at the  
 181 pedestrian height as a function of the wall area density  $\lambda_w$

182

183 
$$\langle U \rangle = \frac{\langle V \rangle}{1 - 0.49\lambda_w^{0.4}} \quad (4)$$

184 
$$\sigma_U = \langle U \rangle (0.25\lambda_w^{0.55}) \quad (5)$$

185 We, therefore, assign three values of wind speed in each grid cell,

186 
$$\langle speed \rangle_1 = \max(0.01, \langle U \rangle (1 - 0.25\lambda_w^{0.55}))$$

187 
$$\langle speed \rangle_2 = \langle U \rangle \quad (6)$$

188 
$$\langle speed \rangle_3 = \langle U \rangle (1 + 0.25\lambda_w^{0.55})$$

189 Note that here we consider the three values equally likely, in order to realistically span the range of possible values that the  
 190 wind speed can take in each grid cell. Since UTCI has been designed for 10m wind speeds, a simple log law is used to  
 191 rescale wind speed at 10m, before passing it to the UTCI routine.

192 **2.3 Calculation of the thermal comfort index**

193 To represent the subgrid spatial variability of air temperature, detailed CFD simulations are not available, so we simply use a  
 194 variability of 1 degree Celsius, which we consider to be a conservative estimate of the spatial variability of air temperature  
 195 over a spatial scale of the order of one  $\text{km}^2$  – This value is consistent with the range obtained in the few non-urban  
 196 simulations available, like Santiago et al. (2014), and Nazarian et al. (2018) over idealized arrays, as well as that one  
 197 obtained by Rivas-Ramos over a realistic neighbourhood of Madrid (2024, personal communication). A better determination  
 198 of the variability is left to future studies. Therefore, for each grid cell, we have three values for air temperature:

199 
$$Temp_1 = Temp_{WRF} - 1$$

200 
$$Temp_2 = Temp_{WRF} \quad (7)$$

201 
$$Temp_3 = Temp_{WRF} + 1$$

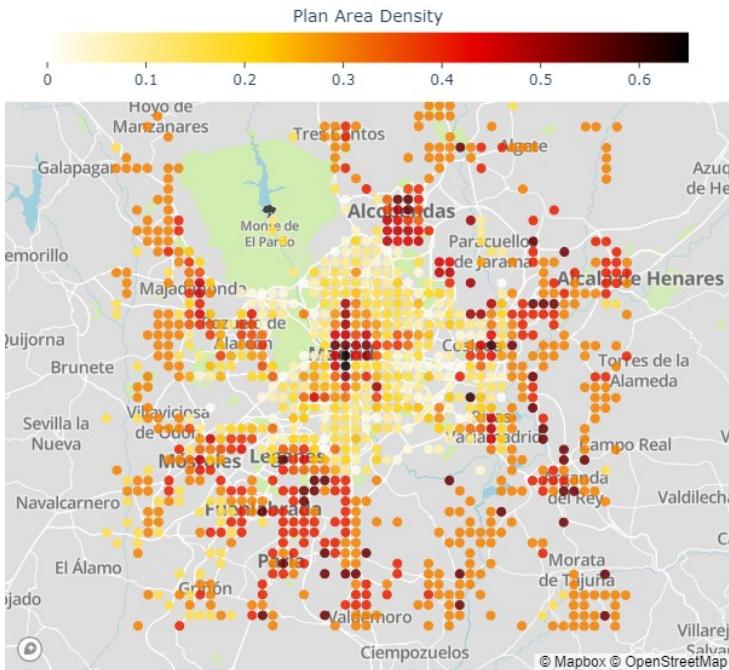
202 where  $Temp_{WRF}$  is the air temperature provided by WRF.

203 We therefore have, for each urban grid cell, *three* values of wind speed, *three* values of temperature, and *six* values of mean  
204 radiant temperature. No variability of the absolute humidity is considered, but the relative humidity is computed using the  
205 three values of air temperature.

206 Based on the variation of these climate variables, assumed uncorrelated, 54 possible combinations of the air temperature, mean  
207 radiant temperature, and wind speed values can be formed. For each one of these combinations, we calculate the corresponding  
208 SET or UTCI value. Based on the resulting distribution, we estimate the value of the 10th, 50th, and 90th percentile SET or  
209 UTCI for each grid square (at each output time). Increasing the number of points where the mean radiant temperature is  
210 computed, or adding more values for the wind speed, does not change significantly the values of the percentiles (not shown).

211 **3. Characterization of thermal comfort in regional-scale models: Madrid case**

212 To illustrate the capabilities of the new scheme, a typical heat wave day in the city of Madrid (Spain) is simulated with WRF.  
213 Madrid is located on a plateau at 500-700m above sea level, in the middle of the Iberian Peninsula. It experiences hot summers,  
214 with frequent heat waves that increasing cause severe heat stress in the population, and it is therefore considered a relevant  
215 case study. Four nested domains have been used, with resolutions of 27, 9, 3, and 1km respectively. The city morphology (Fig.  
216 4) is derived from high-resolution LIDAR data that covers most of the metropolitan area of Madrid (Martilli et al., 2022),  
217 while the morphology of the surrounding towns is determined based on Local Climate Zone maps (Brousse et al., 2016). It is  
218 also important to mention that the city is located on a hilly terrain, with higher elevations in the N-W part of the urban area  
219 (around 700m a.s.l.) dropping to 500m a.s.l. or less in the S-E. Moreover, there are two topographical depressions on the two  
220 sides of the city centre, caused by the rivers Jarama and Manzanares (for a detailed description of the topography see also  
221 Martilli et al. 2022, where the same set-up was used). Other model configurations are the NOAH vegetation model for the  
222 non-urban grid points and the Bougeault and Lacarrere (1989) PBL scheme for turbulence parameterization. WRF coupled  
223 with BEP-BEM has previously been successfully used to simulate a heat wave period in Madrid (Salamanca et al., 2012). The  
224 period used in this paper is three days (14-16 July 2015). In particular, the analysis will focus on the 15th, when the maximum  
225 simulated temperature was above 40 Celsius. More information about the validation and a sensitivity study to select the optimal  
226 set-up can be found in Rodriguez-Sanchez (2020).



227

228 **Figure 4.** Map of the plan area building density over the Madrid region. The underlying map was created with Mapbox OpenStreetMap.  
 229 [The map is oriented so that left is west, and up is north and the size is 50x50km.](#)

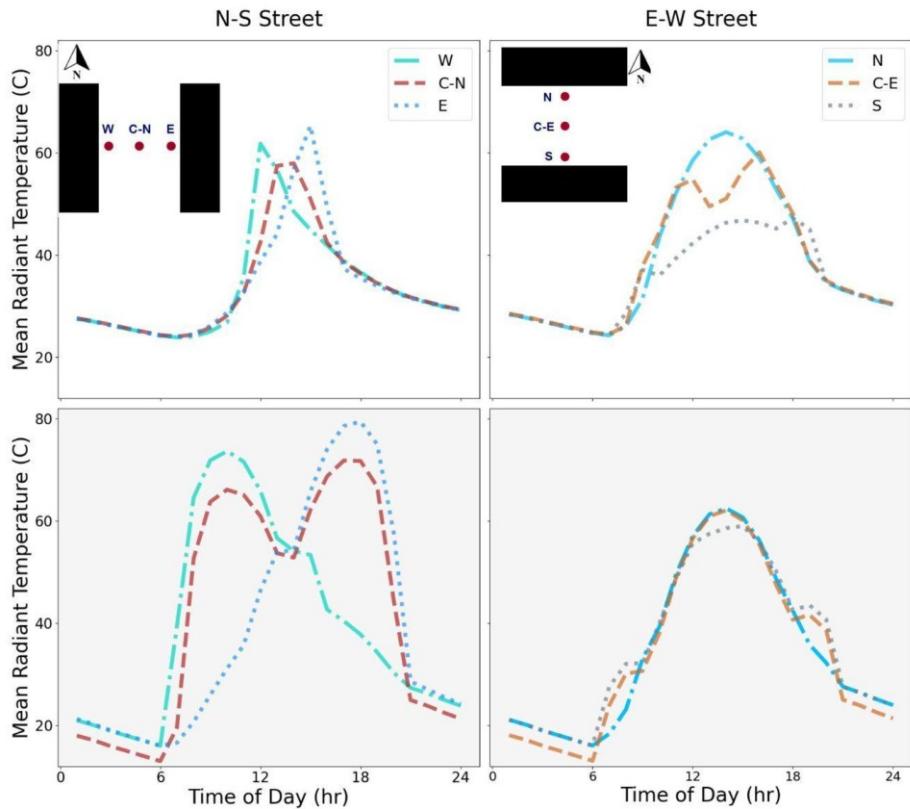
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### 3.1 Sub-grid scale variability of MRT and thermal comfort.

231 In order to understand how urban morphology affects the simulated heat stress, we focus on two grid points with very different  
 232 urban morphology. One is located in the dense core of the city, with a building plan area density of  $\lambda_p = 0.69$ , and a height-to-  
 233 width ratio (H/W) value of 1.6. The second is located in the southern part of the urban area, in a residential neighbourhood  
 234 with a much lower building density ( $\lambda_p = 0.2$ ) and a H/W=0.1.

235 In Figure 5, the diurnal evolution of the mean radiant temperature in the six points (three per street direction) is presented for  
 236 the high urban density point and the low urban density point. During the daytime, the impact of the shadowing is clear, with  
 237 reduced mean radiant temperature in the high-density point compared to the more exposed low-density. On the other hand,  
 238 during night-time, the reduced sky-view factor in the high-density point slows down the cooling compared to the more open  
 239 low-density location.

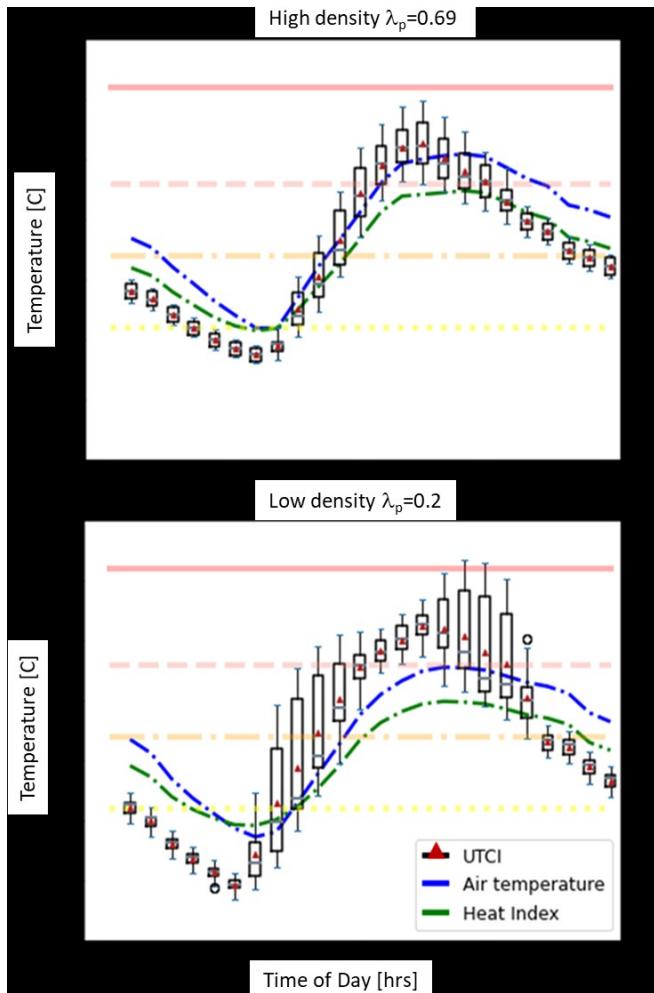


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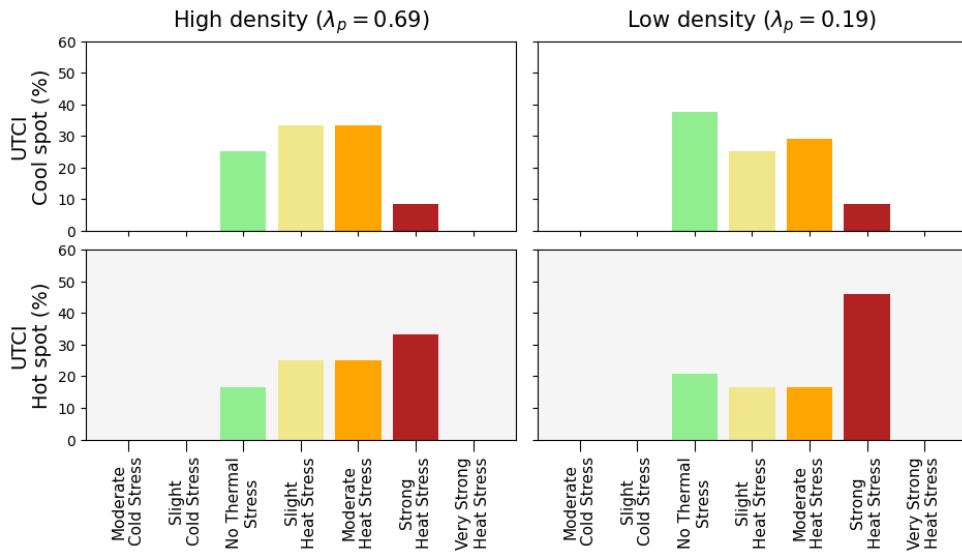
241 **Figure 5.** Diurnal evolution of MRT for 6 points in the urban canopy. The top row (white background) corresponds to a grid point with the  
242 highest building density in the centre of Madrid ( $\lambda_p = 0.69$ ) while the bottom row (with grey background) shows MRT in a low-density  
243 neighbourhood ( $\lambda_p = 0.19$ ). The left column is for an N-S street, while the right column shows an E-W street.

244 This behaviour helps to explain the heat stress index (Figure 6), which is introduced here as an example of an index that can  
245 be computed with standard outputs from meteorological models, i.e., without information related to the radiation environment

246 (e.g., MRT) and urban morphology. The air temperature indicates hotter values both during the day and the night in the high  
247 urban density point compared to the low-density location. The Heat Index, which considers air temperature and humidity only,  
248 and does not include mean radiant temperature or wind, shows the same tendency. On the other hand, the UTCI behaviour  
249 communicates a different and more complete result. In the low-density neighbourhood, more exposed to the sun, the UTCI  
250 shows a stronger sub-grid spatial variability, in particular during the morning and afternoon, with the potential for stronger  
251 heat stress than in the high-density neighbourhood. During night-time, the spatial variability is reduced, due to reduced MRT  
252 variation as the shadowing effect disappears, and higher UTCI values are found at the high urban density location. This  
253 difference in behaviour between the two locations can be seen also in Fig. 7, where the fractions of the 10th percentile of UTCI  
254 values (i.e. representative of one of the coolest spots in the grid cell) and the 90th percentile (i.e., one of the hottest) in the  
255 different heat stress regimes are shown for the two points. Here we can see that in the low-density urban point, the cool location  
256 is in a comfortable UTCI range most of the time, while the hot (90th percentile UTCI) sub grid location is under stress most  
257 of the time. On the other hand, less variability is present in the high-density neighbourhood, with fewer extreme values, and  
258 most of the time in the strong or moderate heat stress regime for both the cool and hot locations within the grid square. This  
259 kind of detail is not available from the Heat Index distribution which does not account for the mean radiant temperature, wind,  
260 or their variabilities (Fig. 8).



261  
262 **Figure 6.** Diurnal evolution of UTCI compared with 2-m air temperature and Heat Index calculated from air temperature and relative  
263 humidity at each grid point). The UTCI boxplot at each hour represents the subgrid-scale distribution calculated based on 6 MRT, 3 wind  
264 speeds, and 3 air temperature values (54 combinations in total). The horizontal lines represent the thermal comfort zones for UTCI (i.e.  
265 above +46C: extreme heat stress; +38 to +46: very strong heat stress; +32 to +38: strong heat stress; +26 to +32: moderate heat stress; and  
266 +9 to +26: no thermal stress).

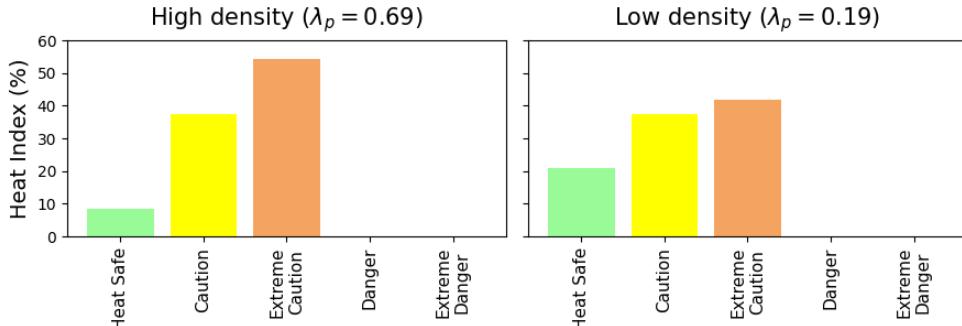


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**Figure 7.** From top to bottom, the frequency of UTCI class over a 24-hour period, for a subgrid location that is cooler (i.e. 10th percentile of UTCI in the urban canopy, top), and for a subgrid location that is hotter (i.e. 90th percentile of UTCI in the urban canopy, bottom), for the high-density (left) and low-density (right) points.



**Figure 8.** same as Figure 7, but for the Heat Index

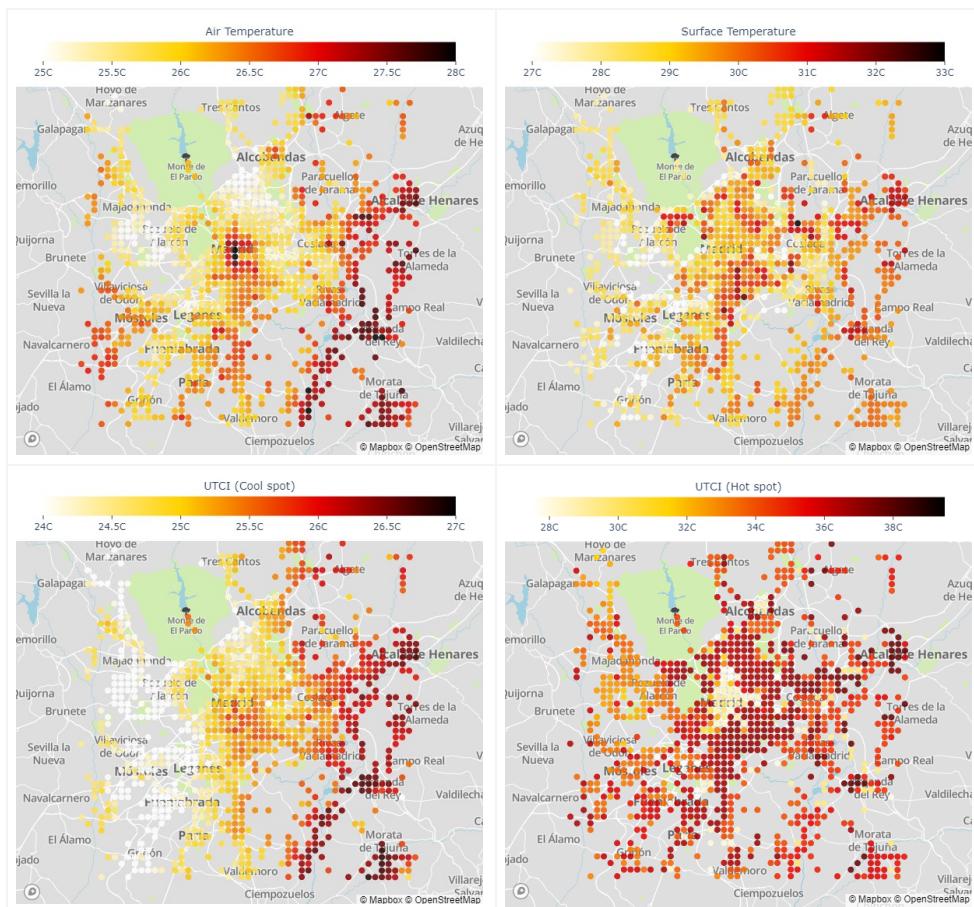
275 **3.2 City-scale maps of outdoor thermal comfort and heat stress indicators.**

276 The previous analysis helps to understand the spatial distribution of the different variables presented in Fig. 9 at 10 and 16  
277 UTC (note that Madrid is at Longitude 3W, so UTC is essentially equal to solar time). In the dense city centre, the distribution  
278 of 2m air temperature at 0900 UTC shows a hot region, with cooler areas in the less dense regions around it. This effect is due  
279 to the fact that in the dense region, the reduced sky-view factor of the streets (high H/W), as well as the larger thermal storage  
280 capacity in the buildings, reduce the nocturnal cooling, and increase the vertical mixing in that part of the city compared to the  
281 surroundings. Such a difference is still visible in the morning. The higher temperatures in the S-E part of the urban area, and  
282 cool temperatures in the N-W are the result of the topographical differences. The spatial distribution of air temperature is  
283 qualitatively similar to the spatial distribution of the 10-percentile of UTCI (e. g. the cool spot in the grid cell), even if the  
284 differences between the centre and the surrounding urban areas are not as intense as for 2m air temperature. On the other hand,  
285 the 90-percentile map (hot spot), shows a completely different pattern; on the city centre, at that time of the day, the whole  
286 street is still in the shadow, while in the surrounding, less dense urban areas, there are points completely exposed to the sun.  
287 As a comparison, the map of surface temperature (a variable often used to represent the spatial distribution of heat in cities) as  
288 seen from a satellite, i.e. based only on a weighted average of roof, street, and vegetation temperatures (see full equations in  
289 Martilli et al. 2021), does not show a clear pattern, and it is uncorrelated with the other maps. This is a clear indication that  
290 this variable should not be used for the assessment of the heat hazard or heat stress in urban areas.

291 At 1600 UTC the air temperature shows again higher values in the city centre, lower in the urban surroundings, and a gradient  
292 from hotter S-E at lower elevations to cooler N-W at higher elevations (Fig. 10). Such a tendency is present also for the 10th  
293 percentile (cool spot), but with less variability. The 90th percentile map (hot spot) indicates that the area with elevated heat  
294 stress extends well beyond the city centre, including lower-density regions that, even if they have lower air temperatures, are  
295 fully exposed to the sun. Finally, as it was the case for 0900 UTC, the surface temperatures have a map uncorrelated with  
296 neither the air temperatures nor the UTCI maps.

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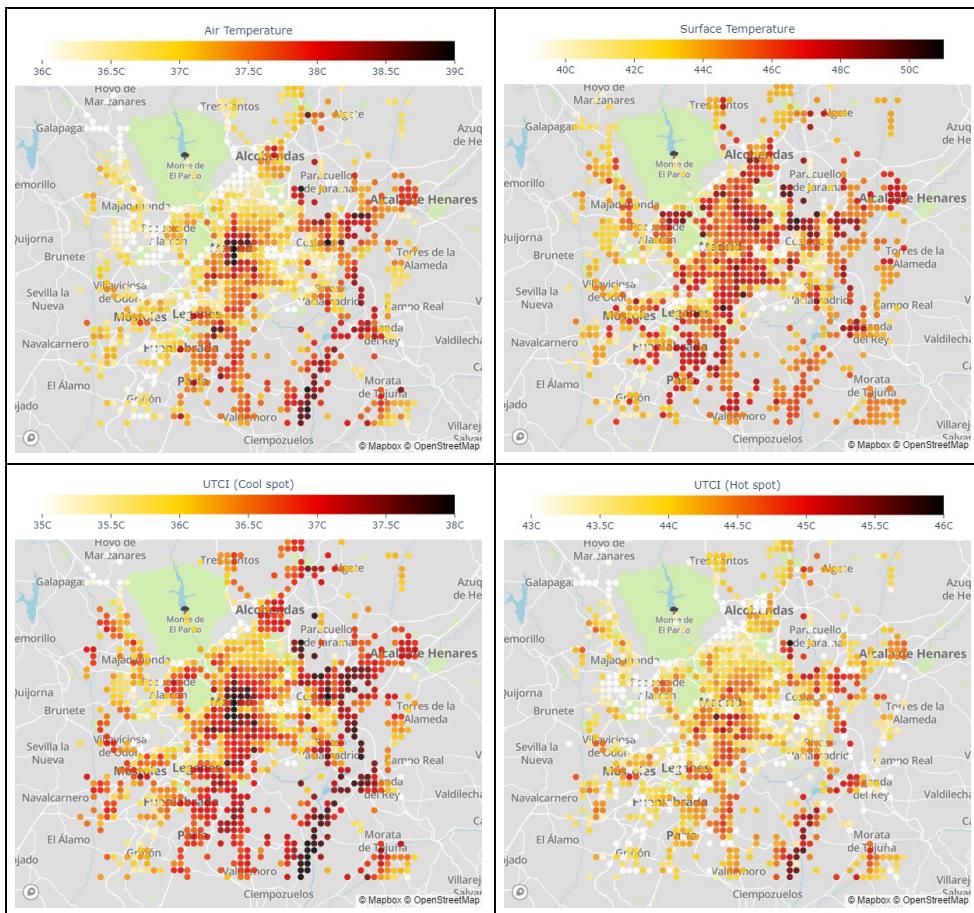
299

300 **Figure 9.** Spatial maps at 0900 UTC for 2-m air temperature (top left), surface temperature (top right), UTCI cool spot e. g. the 10  
 301 percentile of UTCI captured in the urban canopy model (bottom left), and UTCI hot spot e. g. 90 percentile of UTCI in the urban canopy  
 302 (bottom right). Surface temperature is equivalent to that seen by a nadir-view satellite sensor (i.e., an area-weighted average of canopy  
 303 ground temperature, roof temperature, and vegetation temperature in non-urban fractions is considered). The underlying maps were created  
 304 with Mapbox OpenStreetMap

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**Figure 10.** Same as Figure 9, but at 1600 UTC.

310 **4. Limitations**

311 The main limitation of the approach we propose here to account for the sub-grid variability of mean radiant temperature is the  
312 idealization of the urban morphology adopted by the urban canopy parameterization BEP-BEM. This consists of representing  
313 the urban morphology as a series of infinite urban canyons, all with the same width, separated by buildings of constant width,  
314 and variable building height. Two street orientations are considered for each grid cell: North-South, and East-West. The  
315 dimensions of the buildings and street canyons are determined such that the building plan area density, the density of urban  
316 vertical surfaces per horizontal area, and the mean building height are equal to those of the real morphology of the grid cell.  
317 As a result, the total surface areas of walls, roads and roofs in the idealized morphology used by BEP-BEM closely approximate  
318 the corresponding surface areas in the real neighbourhood, and – to a certain extent – the street and buildings of the idealized  
319 morphology can be considered representatives of an average street and set of buildings present in the grid cell. The advantage  
320 of this approach, common among the most widely-used urban canopy parameterizations (Masson, 2000, Kusaka et al. 2001),  
321 is that it allows accurate estimation of shadowing and radiation trapping effects in the urban canopy with low computational  
322 cost, without considering the real urban morphology. Keeping the computational cost low was an essential requirement  
323 considering the computational resources available when these urban canopy parametrizations were developed (about 20 years  
324 ago). With today's computational resources, there may be potential to account for more complexity in the urban morphology.  
325 However, this would require deep changes in the structure of the urban canopy parameterization BEP-BEM that are beyond the  
326 scope of the present article. For this reason we decided to keep the idealized morphology of BEP-BEM and estimate the mean  
327 radiant temperature in six locations representative of the middle of the street and the sidewalks. So, the mean radiant  
328 temperatures computed are representatives of those six points of an “average” street in the grid cell. Indeed, in a grid cell of a  
329 mesoscale model (that typically has a size of the order of one  $\text{km}^2$ ) there is a variety of street and building dimensions and  
330 orientations, so the present approach cannot capture the full spatial variability of mean radiant temperature, a variability that  
331 increases with the heterogeneity of the real urban morphology. Nevertheless, it represents a step forward, since it accounts for  
332 the range (and to some extent, the variability) of mean radiant temperature within the “average” idealized street canyon, that  
333 can be reasonably considered the most likely street typology within the grid cell, something that previous approaches does not.  
334 Overall, the current approach is likely to accurately quantify the mean radiant temperature of at least one “average” shaded  
335 pedestrian and one “average” sunlit pedestrian (during periods with direct shortwave irradiance), and thus capture the largest  
336 source of spatial variation of both MRT and UTCI (Middel and Krayenhoff, 2019). Another limitation of the approach  
337 presented here is the lack of street trees. Currently work is in progress to introduce trees in the version of BEP-BEM  
338 implemented in WRF via implementation of the BEP-Tree model (Krayenhoff et al. 2020), and in this way account for their  
339 impacts on mean radiant temperature as well as on air temperature, humidity, and wind.  
340 The approach used to estimate the mean wind speed and its sub-grid variability is grounded on a large number of CFD  
341 simulations over a variety of urban morphologies. Indeed, as shown in Fig. 3, the sub-grid variability of wind speed can be  
342 quite large, and certainly strongly influenced by the relative arrangements of buildings and streets. So, the approach presented  
343 here will likely underestimate the sub-grid variability of wind speed – and this is why we decided to give the same likelihood

344 to the three values of wind speed estimated in (6), instead of assuming a Gaussian [or Weibull](#) distribution of the probabilities  
345 of wind speed in the grid cell. To fully capture this variability a complete coupling between the mesoscale and a detailed CFD  
346 model would be needed - something that we may be able to do in the near future, but is still unavailable with current  
347 computational resources. Another limitation of the present approach is that the CFD simulations used to build the database  
348 from which the parametrization has been derived are all for a neutral atmosphere, so thermal effects on wind speed and its sub-  
349 grid variability are neglected.

350 **5. Conclusions**

351 A new parameterization to quantify intra-neighbourhood heat stress variability in urban areas using a mesoscale model is  
352 presented. This approach is based on two primary developments: 1) calculation of mean radiant temperature at several locations  
353 within the idealized urban morphology used by the urban canopy model BEP-BEM; and 2) parameterization of mean wind  
354 speed and its sub-grid spatial variability as a function of the local urban morphology and the mean wind velocity computed by  
355 the WRF mesoscale model, using relations developed from a large suite of CFD simulations over a range of realistic and  
356 idealized urban neighbourhoods. The components of the new parameterization have been validated against microscale model  
357 results. From this approach the sub-grid variability of a heat stress index (i.e. UTCI or SET) can be computed for every grid  
358 point, permitting quantification of the heat exposure at both cool and hot locations within each grid square at each time.

359 The new parameterization has been implemented in the multilayer scheme BEP-BEM in WRF and used to simulate a heatwave  
360 day over Madrid (Spain) as proof of concept. The results of this initial application demonstrate the following:

- 361 I. The new parameterization gives information that is more suitable for the evaluation of heat stress than the air  
362 temperature, being based on an index (UTCI or SET) that also combines air humidity, wind speed, and mean radiant  
363 temperature.
- 364 II. The new parameterization provides substantively more information than air temperature alone (or any other index  
365 that does not account for the mean radiant temperature). It provides information about the sub-grid variability (such  
366 that heat stress in both cool and hot locations in each grid square is quantified). To our knowledge, this has not been  
367 done before with a mesoscale model.
- 368 III. The results for the investigated case indicate a strong intraurban variability, both in air temperature and UTCI values,  
369 that can be linked to the differences in urban morphology and elevation above sea level. The ability to assess the  
370 differential impacts of urban morphology on heat stress is key to the provision of guidance for urban planning  
371 strategies that mitigate urban overheating.
- 372 IV. Nadir-view surface temperature (i.e., as seen from a satellite-mounted remote sensor) is poorly correlated with both  
373 air temperature and UTCI maps, indicating that, despite its ubiquitous use at present, it is unlikely to be an adequate  
374 metric for heat impact assessment studies.

375 Finally, we consider that this new development introduces a new methodology for deploying mesoscale models to assess urban  
376 overheating mitigation strategies.

377

378

379 ***Code Availability***

380 The code of WRF-comfort can be obtained here:

381 <https://doi.org/10.5281/zenodo.7951433>

382 The results of the simulation over Madrid shown in the manuscript are stored here:

383 <https://zenodo.org/record/8199017>

384

385 ***Competing interests***

386 The authors declare that they have no conflict of interest

387

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527 Appendix A. *Computation of Radiation for Mean Radiant Temperature*

528 As explained in the text, the mean radiant temperature at pedestrian level is represented using formula (1). The full expression  
 529 of the longwave radiation components for the vertical faces of the pedestrian ( $L_1, L_2$ ), for the case of an urban morphology  
 530 with buildings of constant height and walls with no windows, is as follows:

$$531 \quad L_1 = \sum_{i=1,n} \psi_{1i,p} \varepsilon_W (Rl_{1W_i} + \sigma T_{1i}^4) + \psi_{1G,p} \varepsilon_G (Rl_G + \sigma T_G^4) + \psi_{1S,p} Rl_S$$

$$532 \quad L_2 = \sum_{i=1,n} \psi_{2i,p} \varepsilon_W (Rl_{2W_i} + \sigma T_{2i}^4) + \psi_{2G,p} \varepsilon_G (Rl_G + \sigma T_G^4) + \psi_{2S,p} Rl_S$$

533 Where (see Fig A1).:

534  $\psi_{1i,p}$  = is the view factor from wall section  $i$  of building 1 to the side 1 of the pedestrian

535  $\varepsilon_W$  = is the emissivity of the wall

536  $Rl_{1W_i}$  = is the long wave radiation reaching the section  $i$  of the wall of building 1

537  $T_{1i}$  = is the surface temperature of the section  $i$  of the wall of building 1

538  $\psi_{1G,p}$  = is the view factor from the ground (or street) to the side 1 of the pedestrian

539  $\varepsilon_G$  = is the emissivity of the ground

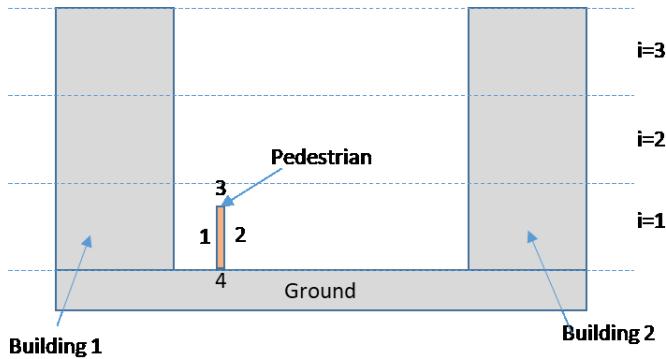
540  $Rl_G$  = is the longwave radiation reaching the ground (street)

541  $T_G$  = is the surface temperature of the ground (street)

542  $\psi_{1S,p}$  = is the view factor from the sky to side 1 of the pedestrian

543  $Rl_S$  = longwave radiation from the sky

544  $\sigma$  = is the Stefan-Boltzmann constant.



545

546 *Figure A1. Schematic of the Street canyon.*

547 Similar meaning applies for side and building 2.

548 The values of the surface temperatures and the longwave radiations are computed with BEP\_BEM. The view factors are  
549 estimated based on formulas A13-A19 of Martilli et al. 2002, using a height for the pedestrian of 1.8 m.

550 For the longwave radiation reaching the top of the pedestrian, we made the simple assumption that it is equal to the radiation  
551 coming from the sky,  $L_3 = Rl_S$ , while for the longwave radiation reaching the bottom of the pedestrian, the assumption is that  
552 it is equal to the radiation emitted and reflected by the ground, or  $L_4 = \varepsilon_G Rl_G + \varepsilon_G \sigma T_G^4$ . We consider that these assumptions  
553 are reasonable, giving that the contribution of the radiation reaching the top and bottom of the pedestrian is only 6% each to  
554 the final value of the mean radiant temperature.

555 A similar approach is followed for the short wave radiation, leading to:

556

$$557 K_1 = \sum_{i=1,n} \psi_{1i,p} \alpha_i R s_{1W_i} + \psi_{1G,p} \alpha_G R s_G + R s_{1S}$$

$$558 K_2 = \sum_{i=1,n} \psi_{2i,p} \alpha_i R s_{2W_i} + \psi_{2G,p} \alpha_G R s_G + R s_{2S}$$

559 Where

560  $R s_{1W_i}$ =short wave radiation reaching the section  $i$  of the wall of building 1

561  $\alpha_i$ =albedo of the section  $i$  of the wall of the building

562  $R s_G$ = the short wave radiation reaching the ground

563  $\alpha_G$  = is the albedo of the ground

564  $Rs_{15}$  = is the short wave radiation from the sun reaching directly side 1 of the pedestrian, computed using formula A10 of  
565 Martilli et al. 2002, using a height of the pedestrian of 1.8m.

566 Similar meaning for side and wall 2.

567 Regarding the radiation reaching the top of the pedestrian,  $K_3$ , for simplicity only the radiation coming directly from the sun  
568 is considered, without accounting for the reflection from the walls. So the value is zero if the pedestrian is in full shadow, and  
569 to estimate it, the formula used is from A11 of Martilli et al. 2002. The value of the radiation reaching the bottom of the  
570 pedestrian is the value reflected by the ground, or  $K_3 = \alpha_G Rs_G$ ,

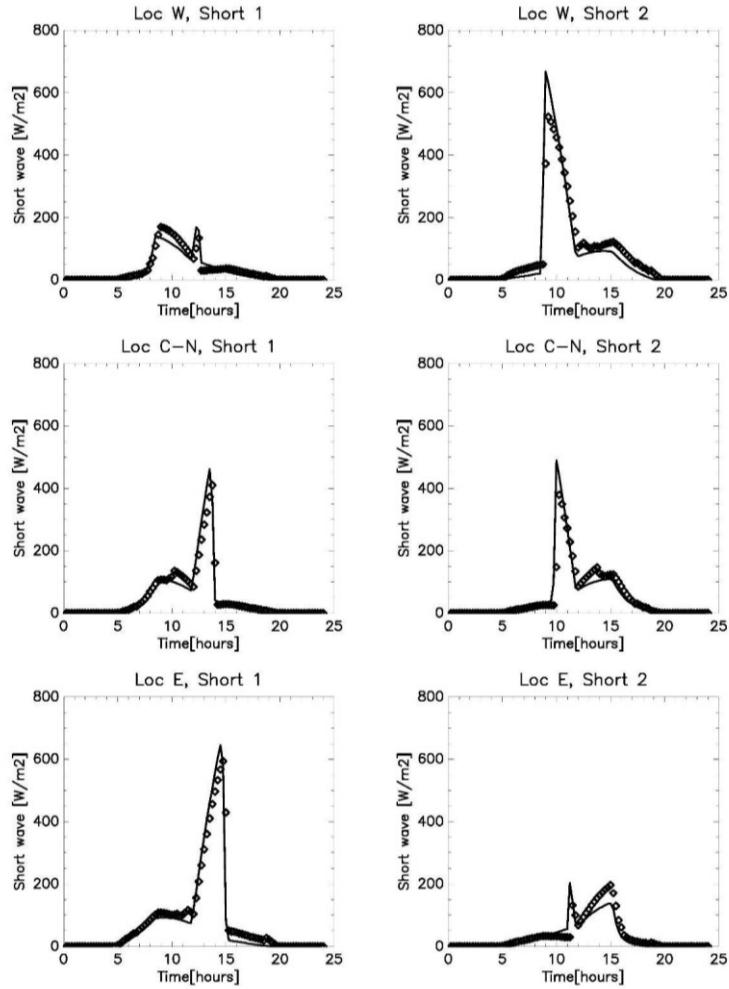
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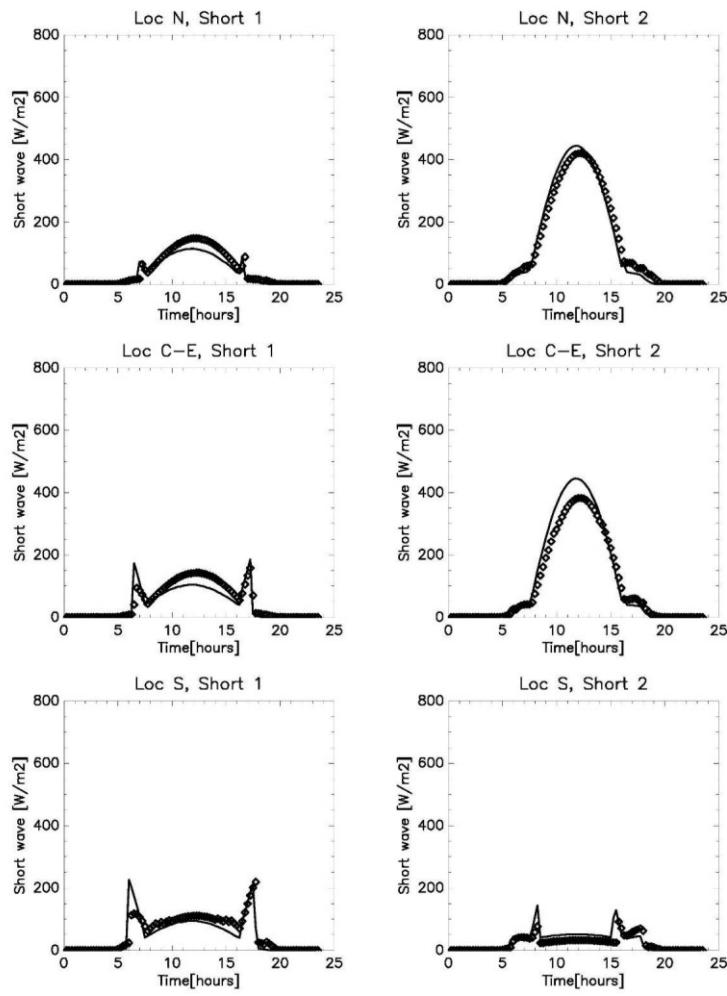
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574 Appendix B. *Comparison of Short wave calculation in BEP-BEM and TUF-pedestrian.*

575 Short wave radiation is an essential component of the MRT. Below we compare the short wave radiation reaching the vertical  
576 sides of the segment representing the human body computed by BEP-BEM vs those estimated with the more detailed model  
577 TUF-pedestrian.



578  
 579 Figure B1. Comparison of short wave radiation at the two sides of the vertical segment representing the pedestrian for the N-  
 580 S oriented street. Solid line is the WRF, while diamonds are TUF. Short 1 means the side 1 of the pedestrian, while Short 2  
 581 the side 2.



582  
 583  
 584 Figure B2. Same as B1, but for an E-W oriented street  
 585

586

587      *Appendix C. CFD simulations for wind speed variability*

588      Data from over 173 microscales CFD simulations of urban airflow are considered over realistic and idealized urban  
 589      configurations, spanning a wide range of building plan area ( $\lambda_p$ ), frontal area ( $\lambda_F$ ), and wall area ( $\lambda_w$ ) densities representative  
 590      of realistic urban neighborhoods in different types of cities. CFD simulations are conducted using 162 large-eddy simulations  
 591      (LES) and 11 Reynolds-averaged Navier–Stokes (RANS) schemes detailed in Table B.1.

Table B.1 Details of CFD microscale simulation cases considered in this study. Simulations are classified based on the configuration (urban form) used. These classifications include **UA** (Uniform height with Aligned configuration), **US** (Uniform height with Staggered configuration), **VA** (Variable height with Aligned configuration), **VS** (Variable height with Staggered configuration), **UR** (Uniform height with Realistic configuration), and **VR-WD** (Variable height with Realistic configuration and multiple Wind Directions considered).

Model	Classification	$H_m$ [m]	$H_{max}$ [m]	$\lambda_p$ range	Count	Source	Example
LES	UA	16	16	[0.0625 - 0.64]	7	Nazarian et al. 2020 Lu et al. 2022	
LES	US	16	16	[0.0625 - 0.64]	7	Nazarian et al. 2020 Lu et al. 2022	
LES	VA	16	20, 24	[0.0625 - 0.64]	42	Lu et al. 2022 Lu et al. 2023	
LES	VS	16	20, 24	[0.0625 - 0.64]	42	Lu et al. 2022 Lu et al. 2023	
LES	UR	16	16	[0.057 - 0.536]	64	Lu et al. 2022	
RANS	VR-WD	14.5-34	variable	[0.190 - 0.680]	11	Sanchez et al. (2017) Santiago et al. (2017) Kracht et al. (2017) Borge et al. (2018) Kracht et al. (2019) Santiago et al. (2020) Sanchez et al. (2021)	

592

593      In the LES simulations, airflow over idealized and realistic urban arrays to determine the model parameters (Nazarian et al.,  
 594      2020; Lu et al., 2022, 2023). Realistic urban layouts are prepared by rasterizing building footprints from an open-source  
 595      dataset OpenStreetMap using OSM2LES (Lu et al., 2022). 64 realistic urban neighbourhoods were obtained assuming uniform  
 596      building height (Table B.1) from several major cities such as Sydney and Melbourne (Australia), Barcelona (Spain), Detroit,

597 Los Angeles, and Chicago (United States). Idealized urban arrays are considered in aligned and staggered arrangement that  
598 follows (Coceal et al., 2007) with varying urban density ( $\lambda_p$  in [0.0625,0.64]) and height variability ( $H_{std}$ =[0m,2.8m,5.6m]).  
599 Simulations are conducted in the Parallelized Large-eddy Simulation Model (PALM, version r4554) (Maronga et al., 2020)  
600 following the same setup in (Nazarian et al., 2020), which has validated results against Direct Numerical Simulation (Coceal  
601 et al., 2007) and wind tunnel experiments (Brown et al., 2001). The computational domain is discretized using the second-  
602 order central differences (Piacsek and Williams, 1970) where the horizontal grid spacing is uniform and the vertical spacing  
603 follows the staggered Arakawa C-grid. The minimal storage scheme is employed in the time integration to solve the filtered  
604 prognostic incompressible Boussinesq equations where the pressure perturbation was calculated in Poisson's equation and was  
605 solved by the FFTW scheme (Frigo and Johnson, 1998).

606 The RANS dataset is derived from steady-state CFD-RANS simulations performed with the Realizable  $k$ - $\epsilon$  turbulence model  
607 (STAR-CCM+, Siemens) over realistic urban areas. The size of the computational domains is determined following the best  
608 practice guideline of COST Action 732 (Franke et al., 2010). The horizontal area covers around 1-1.5 km<sup>2</sup> and the domain top  
609 is at around 8H, being H the mean height of buildings. The resolution of the irregular polyhedral mesh used in all CFD-RANS  
610 simulations goes from 0.5 m close to buildings to 6 m out of the built-up area, which results in between 3 and 8 million grid  
611 points depending on the complexity of the geometry. Inlet vertical profiles for wind speed, turbulent kinetic energy ( $k$ ), and  
612 its dissipation ( $\epsilon$ ), are established in neutral atmospheric conditions. The evaluation of the CFD-RANS simulations was  
613 addressed in previous studies summarized in Table B2 and more information is provided in previous publications.

614