

Atmospheric turbulence observed during a fuel-bed-scale low intensity surface fire

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1 **Abstract.** The ambient atmospheric environment affects the growth and spread of wildland fires,
2 whereas heat and moisture release from the fires and the reduction of the surface drag in the
3 burned areas can significantly alter local atmospheric conditions. Observational studies on fire-
4 atmosphere interactions have used instrumented towers to collect data during prescribed fires,
5 but a few towers in an operational scale burn plot (usually $> 10^3 \text{ m}^2$) have made it extremely
6 challenging to capture the myriad of factors controlling fire-atmosphere interactions, many of
7 which exhibit strong spatial variability. Here, we present analyses of atmospheric turbulence data
8 collected using a 4×4 array of fast-response sonic anemometers during a fire experiment on a 10
9 $\text{m} \times 10 \text{ m}$ burn plot. In addition to confirming some of the previous findings on atmospheric
10 turbulence associated with low-intensity surface fires, our results revealed substantial
11 heterogeneity in turbulent intensity and heat and momentum fluxes just above the combustion
12 zone. Despite the small plot (100 m^2), fire-induced atmospheric turbulence exhibited strong
13 dependence on the downwind distance from the initial line fire and the relative position specific
14 to the fire front as the surface fire spread through the burn plot. This result highlights the
15 necessity for coupled atmosphere-fire behavior models to have 1-2 m grid spacing to resolve
16 heterogeneities in fire-atmosphere interactions that operate on spatiotemporal scales relevant to
17 atmospheric turbulence. The findings here have important implications for modeling smoke
18 dispersion, as atmospheric dispersion characteristics in the vicinity of a wildland fire are directly
19 affected by fire-induced turbulence.

20

21

22 **1 Introduction**

23 Wildland fires are fundamentally linked to atmospheric conditions, with macroscale
24 (thousands of kilometers, weeks to months) factors, such as prolonged periods without
25 substantial precipitation, high temperature, and low humidity, contribute to the drying and pre-
26 heating of fuels, setting the stage for large wildland fire episodes (Potter, 1996; 2012; Finney *et*
27 *al.*, 2015; Littell *et al.*, 2016; Kitzberger *et al.*, 2017). Once ignited, microscale (< 1000 m, < 1 h)
28 conditions, such as local topography and wind speed and direction, take precedence in shaping
29 fire behavior characteristics like burn intensity, ember production, spotting, fire whirls and the
30 rate of spread. Most wildland fires tend to spread in the direction the wind blows, with stronger
31 wind speeds corresponding to faster fire spread (Carrier *et al.*, 1991; Wolff *et al.*, 1991; Clark *et*
32 *al.*, 1996).

33 An essential microscale factor influencing fire behavior is atmospheric turbulence,
34 characterized by irregular microscale air motions in the form of eddies superimposed on mean
35 atmospheric motions (Stull, 1988). Turbulent eddies affect fire behavior as well as the transfer of
36 gaseous and particulate emissions from the fires to the surrounding atmosphere (Clements *et al.*,
37 2008; Seto *et al.*, 2014; Viegas and Neto, 2015; Skowonski and Hom, 2015; Heilman *et al.*,
38 2015; Heilman, 2021). Turbulence in the atmosphere is generated primarily by wind shear as a
39 result of changes in wind speed and/or direction, known as mechanical turbulence, and by
40 convection, referred to as thermal turbulence. Mechanical turbulence is often generated when air
41 flow encounters surface drag, rough terrain or other natural or man-made obstacles and
42 boundaries separating different air masses (e.g., weather fronts), different land cover types (e.g.,
43 grass vs. forested land) or land use types (e.g., agriculture vs. urban). Thermal turbulence is
44 produced when heated surface air rises in the atmosphere, a process known as convection,

45 commonly occurring during daytime when incoming solar radiation exceeds outgoing terrestrial
46 radiation. Fire-induced turbulence, a type of thermal turbulence, results from heat released by
47 combustion, producing buoyant plumes that rise from the combustion zone.

48 Atmospheric turbulence is a pivotal factor influencing fire behavior and the complex
49 exchange of momentum and scalars (e.g., heat, moisture, carbon monoxide, carbon dioxide, and
50 particulate matter) between the combustion zone and the surrounding atmosphere. Existing
51 literature on fire-induced turbulence predominantly draws from data gathered in either
52 management-scale burns, encompassing plots ranging from several to hundreds of hectares, or
53 fine-scale laboratory experiments conducted in burn chambers or wind tunnels under controlled
54 conditions. Notably, a discernible gap exists in observations that seamlessly bridge these two
55 scales (Skwonski, *et al.*, 2021). This study aims to fill this knowledge void by presenting a
56 comprehensive analysis of turbulent data collected from a densely instrumented small-scale (10
57 m x 10 m) burn plot situated in a pitch and loblolly pine plantation. Through this investigation,
58 we seek to augment our understanding of how surface fires modify turbulence and contribute to
59 the dynamic exchange of momentum and scalars between the fire and the surrounding
60 atmosphere.

61 Comprehensive observations of atmosphere turbulence in the presence of wildland fires
62 have only become available in recent decades. For instance, the FireFlux experiment conducted
63 on February 23, 2006 over a 40-hectare plot of native tall-grass prairie in Galveston, Texas
64 represented a significant large-scale field experiment where comprehensive turbulence data were
65 collected above and in the vicinity of a wildland fire front (Clements *et al.*, 2007; Clements *et*
66 *al.*, 2008). The experiment utilized fast-response three-dimensional (3D) sonic anemometers
67 mounted at multiple levels on a tall (43 m) and a short (10 m) tower within the burn plot. This

68 groundbreaking experiment revealed a fivefold increase in turbulence kinetic energy and a
69 threefold increase in surface stress during the fire-front passage, with a rapid return of turbulence
70 to the ambient level behind the fire front. A subsequent field experiment, FireFlux-II, took place
71 at the same site in 2013, aiming to fill gaps in the original FireFlux experiment and provide
72 additional insight on fire-atmosphere interactions and fire-induced turbulence regimes (Clements
73 *et al.*, 2019).

74 While these experiments in Texas provided direct turbulence measurements during
75 intense grass fires, other wildland fire experiments in the New Jersey Pine Barrens provided
76 information on fire-induced turbulence during low-intensity forest understory fires (Heilman *et*
77 *al.* 2015, 2017, 2019 and 2021; Mueller *et al.* 2017, 2019; Clark *et al.* 2020). Conducted between
78 2010 and 2021, these forest fire experiments covered burn plots ranging from approximately 5 to
79 100 hectares, with turbulence data collected using 3D sonic anemometers and thermocouples
80 mounted on 3-, 10-, 20- and 30-m micrometeorological flux towers. The data revealed
81 substantial variations in turbulence intensity, stress, and fluxes across the canopy layer,
82 complicating the understanding of local turbulence regimes and their interaction with the
83 spreading fires. Notably, fire-induced increases in turbulent kinetic energy are considerably
84 larger near the top of the forest canopy layer than within it, suggesting a substantial vertical
85 mixing or transport of fire emissions near the canopy top (Heilman *et al.*, 2015). The
86 observations also highlighted the persistence of an anisotropic turbulence regime throughout the
87 vertical extent of overstory canopy layers, even within highly buoyant plumes during the passage
88 of fire fronts. The results suggested that spreading line fires could significantly affect the
89 skewness of daytime velocity distributions typically found inside forest vegetation layers, and

90 the contributions to turbulence production and evolution from mechanical shear production and
91 diffusion could differ markedly in the pre-fire and post-fire environments (Heilman *et al.*, 2017).

92 The data from both the TX grass fires and NJ forest understory fires have also provided
93 insight into the turbulent momentum and heat transfer processes during fires. Enhanced
94 turbulence updrafts and downdrafts during fires facilitate the transfer of warmer air (or lower
95 momentum air) upward and colder air (or higher momentum air) downward, known as “ejection”
96 and “sweep”, respectively (Heilman *et al.*, 2021). Analyses suggested that wildland fires in grass
97 or forest environments could substantially alter the relative importance of sweep and ejection
98 processes in redistributing momentum, heat and other scalars in the lower atmosphere (Heilman
99 *et al.*, 2021). Sweep events dominate momentum transfer at the fire front, regardless of fire type,
100 despite the stronger updrafts than downdrafts at the front. However, the effect of fires on
101 turbulent heat transfer differ between heading intense grass fires and backing low-intensity
102 forest-understory fires. The former tended to be dominated by ejection events, while in the latter
103 case ejection and sweep events are equally important (Heilman *et al.*, 2021).

104 The TX and NJ wildland fire experiments were conducted over burn plots on relatively flat
105 terrain. However, wildland fire behaviors can be significantly influenced by topography (Werth
106 *et al.*, 2011; Sharples, 2009; Sharples *et al.*, 2012), as topography exerts a strong impact on both
107 weather and fuel conditions (Bennie *et al.*, 2008; Ebel, 2013; Billmire *et al.*, 2014; Calviño-
108 Cancela *et al.*, 2017; Povak *et al.*, 2018). In California, a series of prescribed burn experiments
109 between 2008 and 2012 were conducted in complex terrain with burn plots on a simple slope
110 (Seto and Clements *et al.*, 2011; Seto *et al.*, 2013; Clements and Seto, 2015; Amaya and
111 Clements, 2020) or in a narrow valley (Seto and Clements, 2011), ranging from 2 to 15 hectares
112 in size. Although all burn plots were dominated by grass fuels, data from these experiments

113 provided unique information on the interactions between terrain-induced circulations and fire-
114 induced flows. Results indicated that terrain-induced slope flows and valley winds can interact
115 with fire-induced flows, enhancing horizontal and vertical wind shears that subsequently
116 contribute to turbulence production. Interactions of fire-induced flows with slope winds also
117 produce local convergence or divergence with strong updrafts and downdrafts. Turbulence
118 regimes tend to be anisotropic immediately above fire fronts, transitioning towards isotropic
119 conditions higher up (Seto *et al.*, 2013, Clements and Seto, 2015; Amaya and Clements, 2020).
120 Data from these studies also revealed an increase in turbulent energy in both velocity and
121 temperature spectra at higher frequencies, attributed to small eddies shed by fire fronts, and an
122 increase at lower frequencies related to the strengths of the cross-stream wind component
123 generated by the fire and enhanced by topography (Seto *et al.*, 2013).

124 The aforementioned field experiments were conducted on operational-scale (or
125 management-scale) burn plots, ranging from several to 100 hectares, making it unfeasible to
126 cover such large burn plots with just a few micrometeorological towers. Consequently, the
127 measurement strategy of these experiments was centered around tall towers placed at couple of
128 key spots in the burn plot to provide information on vertical variations of fire-atmosphere
129 interactions. However, the lack of spatial coverage of the complex fuel and atmospheric
130 conditions at these large burn sites makes interpretation of limited observations challenging.
131 Laboratory studies (e.g, Forthofer and Goodrick, 2011; Campbell-Lochrine *et al.*, 2021; Di
132 Cristina *et al.*, 2022) have the advantage of monitoring fires using densely spaced instruments.
133 Nevertheless, laboratory studies are often conducted under controlled conditions that may not be
134 representative of the real fuel and atmospheric environments encountered in outdoor wildland

135 fires. There exists an apparent gap in the observations of fire-atmosphere interactions between
136 operational-scale burns and fine-scale laboratory experiments.

137 In the context, we present analyses of turbulent data collected during a small-scale (10 m
138 $\times 10$ m) experimental burn, which was densely instrumented for the purpose of bridging the gap
139 in our knowledge about fire-atmosphere interactions between operational-scale ($\geq 10^3$ m²) and
140 laboratory-scale ($< 10^1$ m²) fire experiments. The primary question we aim to address is how a
141 low-intensity surface fire may modify turbulence in the atmosphere just above the combustion
142 zone. Specifically, our analyses will explore questions such as: How does the surface fire alter
143 turbulence intensity and turbulent heat and momentum exchanges between the combustion zone
144 and the atmosphere above? Whether and how would the fire change the partitioning of the heat
145 and momentum fluxes into different types of events (both event number and event contribution)?
146 How do the modifications of the fire on turbulence vary spatially across the burn plot? Answers
147 to these questions could prove useful for predicting fire-atmosphere interactions, particularly the
148 momentum and scalar exchanges between the fire and the atmosphere. Moreover, insights into
149 spatial variability could guide the determination of horizontal grid spacing in coupled
150 atmosphere-fire behavior models necessary to capture horizontal variability in near-surface
151 atmospheric turbulence during the presence of surface fires.

152

153 **2 Method**

154

155 **2.1 Experiment and Instrumentation**

156 The experimental burn that this study focuses on took place on May 20, 2019 in a pitch
157 and loblolly pine plantation at the Silas Little Experimental Forest in New Lisbon, New Jersey.

158 This particularly burn was part of broader series of 35, densely instrumented, low-intensity
159 surface fire experiments on 100 m² (10 m x 10 m) plots in this plantation conducted between
160 March 2018 and June 2019 by a research project funded by the Department of Defense's
161 Strategic Environmental Research and Development Program (SERDP). The overall goal of this
162 research project was to collect data using laboratory-scale (10⁰-10¹m²) experiments, intermediate
163 or fuel-bed-scale (10² m²) burns and management-scale (10³⁻⁴ m²) prescribed fires to improve the
164 understanding of combustion processes and fire-atmosphere interactions across scales (Gallagher
165 *et al.*, 2022; Skwonski, *et al.*, 2021).

166 As shown in Figure 1, the 100 m² burn plot was densely monitored by instruments
167 mounted on four parallel east-west-oriented trusses (A, B, C, D). On each truss, four 3D fast-
168 response sonic anemometers (R.M. Young 81000V, Traverse City, MI, USA) were mounted at
169 2.5 m above the ground level (AGL) to collect the east-west (*u*), north-south (*v*) and vertical (*w*)
170 velocity components and temperature at a sampling rate of 10 Hz (Clark *et al.*, 2022a).
171 Additional 10-Hz temperature data were also obtained using fine-wire thermocouples (Omega
172 SSRTC-GG-K-36, Omega Engineering, Inc., Stamford, CT, USA) mounted at a range of heights
173 (0, 5, 10, 20, 30, 50, 100 cm) below the two inner trusses (B and C) (Clark *et al.*, 2022b). A
174 radiometer/visible spectrum camera pair was mounted adjacent to each sonic anemometer to
175 measure radiative heat fluxes and flame arrival times and persistence (Kremens *et al.*, 2022).
176 Spatially explicit fire spread data were derived from infrared data collected by an infrared video-
177 camera (A655SC, FOL6 100.0-650.0 C lens, FLIR Systems Inc., Wilsonville, OR, USA)
178 mounted on top of a 10-m tower in the center of the plot (Skowronski *et al.*, 2022a). A custom
179 field calorimetry hood (labeled TACO next to B2) with an inlet oriented over a portion of the
180 fuel bed was used to sample O₂, CO₂, and CO concentrations in buoyant plumes (Campbell-

181 Lochrie *et al.*, 2022). Gas concentrations were measured at 1 Hz using an Infrared gas analyzer
182 (Crestline NDIR 7911, Crestline, Livermore, CA, USA).

183 The analyses here focused only on the data from the 4×4 sonic anemometer array. All
184 sonic anemometer data underwent a quality assurance and control process to remove spurious
185 values (Clark *et al.*, 2022a). Initially, data that were collected prior to a designated common start
186 time was removed, providing a starting point for the observations for the burn period. Next, the
187 data from sonic anemometers include a self-reporting diagnostic column where any non-zero
188 number is considered an invalid measurement, so any measurement that reported a non-zero
189 diagnostic code was removed. Following these initial steps, data that fell outside the sonic
190 anemometer operating parameters (wind speed: ± 40 m/s; temperature: ± 50 °C) were also
191 removed.

192 The horizontal wind velocities were rotated into a streamwise coordinate system where
193 the u -component (streamwise component) is aligned with the prevailing wind direction, and the
194 v -component (cross-stream component) is perpendicular to the prevailing wind direction pointing
195 to the left. Vertical winds were not corrected for tilt because of the short (<30 min) observational
196 period and because the burn plot was on level ground and each sonic anemometer was carefully
197 mounted and leveled so that the wind sensors were very close to true horizontal and vertical
198 planes. The results (presented below) indeed suggested that the contamination of vertical
199 velocity by horizontal velocities were negligibly small as the average vertical wind component
200 during the pre-burn period was nearly zero.

201

202 **2.2 Fuel and ambient atmospheric conditions**

203 The primary fuel for this burn was pitch pine needles (*Pinus rigida* Mill.). Based on
204 biometric and terrestrial laser scan measurements collected pre- and post-burn, the fuel mass was
205 estimated to be about 0.5 kg m⁻² and fuel moisture content about 5.5% (Skowronski *et al.*,
206 2022b).

207 The ambient atmospheric conditions on the day of the burn is indicated using the data
208 from a surface weather station located approximately 200 m northeast of the burn plot that has
209 similar type of land cover as the burn plot (Figure 2). Ambient winds were very weak in the
210 morning, varying in direction between south and west. Wind speeds increased in midday to about
211 5 m s⁻¹ along with a direction shift to southwest and west. This wind speed increase was likely
212 due to the mixing of higher winds from above to the surface as the mixing layer grew higher
213 during the day. The growth of the mixing layer was a result of increased turbulent mixing
214 associated with surface heating, as indicated by an increase in surface temperatures from about
215 20 °C in the morning to slightly above 30 °C around 1400 Local Standard Time (LST) and a
216 corresponding decrease in relative humidity from over 80% in the morning to less than 40% in
217 the early afternoon.

218

219 **2.3 Fire spread**

220 The experiment started around 14:25 LST when a single 10-meter cotton cord was
221 soaked in accelerant, ignited and then dropped on the fuel bed to produce a single, near linear
222 ignition across the western border of the plot. Infrared imagery data (Figure 3) captured by the
223 overhead infrared camera is used to evaluate the changes in temperature from just before ignition
224 (Figure 3a), immediately after ignition (Figure 3b), and through the period following the ignition
225 as the line fire spread with winds across the plot (Figure 3c-f). The average fire spread rate

226 throughout the burn was estimated from these data to be approximately 5.4 cm s^{-1} . The ignition
227 produced a line fire parallel to the western boundary of the plot (Figure 3b). The line fire spread
228 in the direction of the west-southwesterly background wind towards the east-northeast over the
229 next few minutes (Figure 3c, d). The initial spread was faster on the northern portion of the
230 domain, as expected from the south-southwesterly wind direction. As the fire burned through the
231 northern portion of the plot, the fire front caught up in the southern portion (Figure 3e). The fire
232 ended at around 14:32:16 LST as the fire front reached the eastern boundary of the plot and ran
233 out of fuel to continue (Figure 3f).

234

235 **2.4 Data Analysis**

236 The quality-controlled 10-Hz wind and temperature data from the 3D sonic anemometers are
237 used to calculate turbulent perturbations defined as the differences between the instantaneous
238 observations and the mean values:

$$239 \quad \varphi' = \varphi - \bar{\varphi} \quad (1)$$

240 where $\bar{\varphi}$ is the mean value that is estimated by block-averages

$$241 \quad \bar{\varphi} = \frac{1}{N} \sum_{n=1}^N \varphi_n \quad (2)$$

242

243 Here, N is the number of samples over the averaging period or the time block and the mean
244 values represent the mean state of the atmospheric flow. In traditional turbulence studies, mean
245 state is usually determined by averaging the data over a period of a few minutes up to 1 hour,
246 depending on atmospheric stability and the scale of interest. However, the block-averaged values
247 during the period of active burning are likely to be contaminated by the fire and therefore poorly
248 represent the mean background flow. To resolve this issue, Seto *et al.* (2013) and Heilman *et al.*
249 (2015) proposed that the block-averaged means for the fire period be replaced by block-averaged
250

251 means calculated during the pre-burn period. In order to adopt this approach, the observational
252 period is divided into three periods representing pre-burn, burn and post-burn, which are
253 described in detail below.

254 The arrival of the fire front at most locations in the sonic anemometer array was clearly
255 marked by a sharp rise in temperature (Figure 4). However, the magnitudes of the temperature
256 increase and the rates of increase vary with the location of the sonic anemometers because the
257 shape of the flame front was irregular (Figure 3). Note that the sonic temperatures are limited to
258 50 °C, which is the operational range for the instruments beyond which data are deemed
259 unreliable. Based on the temperature time series and the time when the fire was ignited along the
260 western boundary (14:25 LST), the 10-min period from 14:15:13 through 14:25:12 LST is
261 defined as the pre-burn period over which the mean values for u , v , S (horizon wind speed), w ,
262 and T are calculated, and these values are used for computing perturbations for the entire
263 experiment. The definition of the burn period, however, is complicated by the fact that the fire
264 front reaches/leaves each sonic anemometer at a different time and consequently the true burn
265 period across the plot varies somewhat depending on the location of each sonic anemometer.

266 To create a robust definition of the burn period that can be applied to all the sonic
267 anemometers in the 4×4 array, and eventually to other burns in the broader burn series, the
268 sharp rise in sonic temperatures associated with fire front is measured using integer (n) multiples
269 of the standard deviation (denoted using σ) of the average temperature over the pre-burn period.
270 A threshold value that is too small (e.g., 1 or 2 times standard deviation) may not distinguish the
271 increase in temperature associated with the fire front from normal temperature fluctuations
272 during the day, but a value that is too large (e.g., 10 time standard deviation) may fail to detect
273 the fire front associated with a small or moderate temperature increase. Figure 5 shows the

274 number of sonic anemometers whose temperatures exceed $n \times \sigma$ as n increases from 1 to 35, and
275 the length of the exceedance period. As n increases from 1 to 8 or the threshold value for fire-
276 induced temperature increase changes from 1σ to 8σ , the number of sonic anemometers drops
277 from 16 to 13 and the period drops sharply from just under 60 min to about 6 min. Continued
278 increases in the threshold values from 8σ to 25σ result in no change in the number of
279 anemometers and very little change in the length of the period (less than 1 min). This analysis
280 suggests that 8σ could be used as the threshold for temperature increases associated with fire
281 front. Thresholds lower than 8σ would imply a burn period of 30- to 60-min long that, according
282 to the time series in Figure 4, would include periods of no fire and therefore de-emphasize the
283 effects of the fire in the resulting analyses. Applying this criterion to all the sonic anemometers
284 and defining the burn period as between the first and last sonic temperature at or above the
285 threshold leads to the selection of the burn period as 14:26:13 to 14:32:29 LST. Finally, the 10
286 min following the burn period (14:32:30 to 14:42:29 LST) is defined as the post-burn period.

287 Following the establishment of the three periods, wind and temperature perturbations are
288 calculated using equations (1) and (2), where the pre-burn averaged values are used as means for
289 the burn and post-burn periods. Strictly speaking, the perturbations calculated for the burn and
290 post-burn periods are not classical turbulent perturbations; to differentiate the features from
291 classical turbulence, they should be interpreted as being primarily fire-induced turbulent
292 perturbations.

293 As noted above, horizontal wind velocity is rotated into a streamwise coordinate where
294 the x -component (streamwise component, u) is aligned with the prevailing wind direction and the
295 y -component (cross-stream component, v) is perpendicular and pointing to the left of the
296 prevailing wind. The prevailing wind direction for the rotation is determined by the 10-min pre-

297 burn period average of wind directions across all 16 sonic anemometers. The average wind
298 directions during the pre-burn period vary slightly across the 16 sonic anemometers, with mean
299 and median wind directions of 225 and 226 degrees, respectively. The subtle variations in wind
300 directions is possibly due to slight error in sensor alignment, rather than actual flow
301 heterogeneity. The 226 degrees is used as the prevailing wind direction for the purpose of
302 coordinate rotation.

303 The quality controlled, coordinate rotated data from the sonic anemometers are analyzed
304 to determine fire-induced changes to turbulence intensity, vertical heat fluxes and vertical fluxes
305 of horizontal momentum also known as shear stress just above the combustion zone by
306 comparing values between the pre-burn and the burn periods. The values are also compared
307 between the pre-burn and post-burn periods to determine how quickly the effects of fire dissipate
308 or how fast the atmosphere returns to the ambient state.

309 Turbulence intensity is measured by the turbulent kinetic energy (*TKE*) defined as the
310 sum of the variance of the three velocity components:

311
312
$$TKE = (\overline{u'^2} + \overline{v'^2} + \overline{w'^2}) / 2 \quad (3)$$

313
314 Turbulent shear stress is commonly measured by shear velocity or friction velocity denoted by
315 u_* and the square of friction velocity is related to the magnitude of the kinematic vertical flux of
316 horizontal momentum:

317
$$u_*^2 = (\overline{u'w'}^2 + \overline{v'w'}^2)^{\frac{1}{2}} \quad (4)$$

318 where $u'w'$ and $v'w'$ are the vertical fluxes of streamwise and cross-stream momentum flux,
319 respectively and the overbar denotes time average. The average period is 1 min for this analysis

320 to be consistent with previous studies on fire-induced turbulence (Seto *et al.*, 2013; Heilman *et*
321 *al.* 2021). Vertical kinematic heat flux is calculated as $\overline{T'w'}$ and the averaging period is also 1
322 min.

323 For the analyses of vertical turbulent fluxes of heat and horizontal momentum, a quadrant
324 analysis technique (Katul *et al.*, 1997, 2006; Heilman *et al.*, 2021) is utilized to delineate the
325 contributions to the turbulent heat or momentum transfer from four types of processes
326 corresponding to the four quadrants of a w' (horizontal) and φ' (vertical) coordinate, where the w'
327 denotes vertical velocity perturbation and φ' denotes perturbations of temperature (T') or
328 horizontal wind speed (S') in heat or momentum flux calculations, respectively. The four
329 quadrants are: Q1: $\varphi'w' > 0, \varphi' > 0, w' > 0$; Q2: $\varphi'w' < 0, \varphi' > 0, w' < 0$; Q3: $\varphi'w' > 0, \varphi' <$
330 $0, w' < 0$; Q4: $\varphi'w' < 0, \varphi' < 0, w' > 0$. Note that the perturbation in horizontal wind speed
331 (S'), rather than the streamwise or cross-wind components (u' or v'), are used for computing
332 momentum flux following Heilman *et al.*, (2021):

$$333 \quad S' = S - \bar{S} \quad (5)$$

$$334 \quad S = \sqrt{u^2 + v^2} \quad (6)$$

335 The quadrant analysis is also known as sweep-ejection analysis (Heilman *et al.*, 2021)
336 which associates each quadrant with a specific type of vertical turbulent transfer events. The
337 names of the events and the associated quadrant designations, which are different for turbulent
338 heat and momentum fluxes, are given in Table 1.

339 Based on the definition in Table1, ejection (Q1) and sweep (Q3) events contribute to
340 positive vertical turbulent heat flux through the upward transfer of warmer air from below
341 (ejection) or the downward transfer of cooler air from above (sweep), while inward interaction

342 (Q2) and outward interaction (Q4) events contribute to negative turbulent heat flux through the
343 downward transfer of warmer air from above (inward interaction) or the upward transfer of
344 cooler air from below (outward interaction). For vertical flux of horizontal momentum, inward
345 interaction and outward interaction events contribute to positive flux through the upward transfer
346 of faster moving air (outward interaction) or the downward transfer of slower moving air (inward
347 interaction), while sweep and ejection events contribute to negative momentum flux through the
348 downward transfer of faster moving air (sweep) or the upward transfer of slower moving air
349 (ejection). Note that the warmer/cooler or faster/slower air is relative to the air in the adjacent
350 layers.

351 The sweep-ejection analysis calculates the proportion of a given type of events by simply
352 counting the number of events or the data points in the 10 Hz time series that fall within the
353 given quadrant. The contributions of the given type of events to the average turbulent fluxes over
354 a given time period (T_p) are calculated, following Heilman *et al.* (2021), by the integral

355

$$356 \quad \overline{\varphi'w'}_Q = \frac{1}{T_p} \int_0^{T_p} \varphi'(\tau)w'(\tau)\varepsilon_Q d\tau \quad (7)$$

357

358 where ε_Q is 1 for the given quadrant and zero otherwise, τ is time and φ' is temperature or
359 horizontal wind speed perturbation for heat or momentum fluxes, respectively.

360

361 **3 Results and Discussion**

362

363 **3.1 Fire-Induced Perturbations to Wind and Temperature**

364 Before we examine fire-induced changes to turbulence in ambient atmosphere, we first
365 take a look at the response of the instantaneous temperature and wind to the surface line fire
366 recorded by the 16 sonic anemometers as the fire spread from west to east across the 10 m ×10 m
367 burn plot (Figure 6). Note that perturbation temperatures (T' , see Eq. 1), instead of actual
368 temperatures, are shown to accommodate the magnitude difference between temperature and
369 wind, facilitating a more coherent visualization of the joint effects of the fire on temperature and
370 wind.

371 The natural or non-fire fluctuation recorded during the pre-burn period are small, with
372 magnitudes generally less than 2.5 m s⁻¹ for u , 1 m s⁻¹ for v and 2.5 °C for T' . The fire impinging
373 upon the sonic anemometers is marked by a sharp increase in T' , but the magnitude of the
374 temperature changes depend heavily on location, from very little change on the western side (A1,
375 B1, C1, D1) of the burn plot where the fire was ignited, to a nearly 20°C increase on the eastern
376 side (A4, B4, C4, D4). This spatial heterogeneity in T' is consistent with the pattern of the fire
377 spread from the western boundary toward the east and northeast by the southwesterly ambient
378 wind (Figure 4). During the burn period, the u fluctuations decreased slightly while the v
379 fluctuations increased. The v -component no longer fluctuated around zero, as in the pre-burn
380 period, but rather it was dominated by negative values, indicating a systematic shift in wind
381 direction. There was a tendency for u and T' to return towards the pre-burn conditions after the
382 burn, but the v component remained negative during the post-burn period.

383 The observed changes in the distribution of wind and temperature values associated with
384 the fire at all 16 sonics are summarized by the box-whisker plots in Figure 7. The pre-burn mean
385 is 1.7 m s⁻¹ for the streamwise wind component u and near zero (-0.04 m s⁻¹) for the cross-stream
386 component v . The pre-burn vertical velocity distribution also has near zero mean, which

387 confirms that the sonic anemometers were well-leveled. During the burn period, the mean of u
388 dropped in magnitude from 1.7 to 1.05 m s⁻¹ while the mean of v increased in magnitude from -
389 0.04 to -0.65 m s⁻¹, indicating an overall shift in wind direction from southwesterly to west-
390 southwesterly. This change in the horizontal wind components suggests that ambient air was
391 drawn towards the fire producing convergence at the fire front. There is also a fire-induced
392 widening of the distributions of the horizontal wind components, particularly the v component,
393 and an increase in the number of outliers with magnitudes that nearly doubled the pre-fire
394 magnitude. The large negative values in v during the burn period reinforce the suggestion of
395 convergence in the vicinity of the fire.

396 Interestingly, there is little evident change in the overall distribution of w during the burn
397 period, except that more and larger outliers are indicated. The maximum updrafts (downdrafts)
398 during the burn period reach speeds of nearly 6 m s⁻¹ (-5 m s⁻¹), which is more than double those
399 of the pre- and post-burn periods, suggesting that intermittent turbulent eddies associated with
400 the fire could have a strong impact on vertical velocity just above fuel bed. The T' distribution
401 also widens substantially during the burn period ($\sigma=4.24$ °C) compared to the pre-burn period
402 ($\sigma=0.48$ °C), with the maximum temperature perturbation reaching nearly 20°C.

403 The influence of the fire on the horizontal wind components continues into the post-burn
404 period, as the post-burn distributions of u and v fall between those of the pre-burn and burn
405 periods. In contrast, the post-burn w distribution returns to a distribution very close to that of the
406 pre-burn period. Similarly, the T' distribution during the post-burn period is very similar to that
407 of the pre-burn period. The similarities between the w' and T' distributions suggest that the two
408 variables are closely related to each other, with large updrafts during the burn period being
409 generated primarily by heating. This result suggests that the fire-induced circulation exhibits

410 behavior more consistent with a buoyant plume than mechanically forced rising motion resulting
411 from converging surface air.

412

413 **3.2 Intensity of Fire-Induced Turbulence**

414 We now explore the modifications of the fire to atmospheric turbulence properties just
415 above the combustion zone. The first question to address is how turbulence intensity quantified
416 by TKE in Eq. (3) is modified by the fire and how the modification may vary with location in the
417 burn plot. Figure 8 shows time series of 1-minute averaged TKE and its three components (the
418 variance of the three velocity components) for each of the sonic anemometers. The time series
419 indicate lower TKE values in the pre-burn period, larger values during the burn period, and
420 values remaining high in the post-burn period. The burn period TKE is primarily driven by an
421 increase in horizontal velocity variance, $\overline{u'^2}$ and $\overline{v'^2}$, particularly the cross-stream component
422 $\overline{v'^2}$. The TKE values remain high into the post-burn period and, at several sonic anemometers
423 (D3 and C4), the post-burn TKE peaks are comparable with or higher than the peaks observed
424 during the burn period.

425 The box-whisker plots in Figure 9 depict the fire-induced changes to the distribution of
426 turbulence intensity as observed by all 16 sonic anemometers. Averaging across all the
427 instruments, the burn period mean TKE is $1.25 \text{ m}^2\text{s}^{-2}$, which is roughly double the pre-burn mean
428 of $0.697 \text{ m}^2\text{s}^{-2}$. The interquartile range of the burn period TKE is nearly three times the pre-burn
429 period range. Despite the increase in the mean and the interquartile range of the TKE from the
430 pre-burn to the burn period, the mean TKE values are still below $3 \text{ m}^2\text{s}^{-2}$, which is a threshold
431 sometimes used as an indicator for substantial boundary-layer turbulence (Stull, 1988; Heilman
432 and Bian, 2013), suggesting that this low-intensity surface line fire fails to produce a

433 substantially turbulent environment at the levels just above the fuel bed. The mean TKE in the
434 post-burn period does not return to that of the pre-burn period and remains elevated ($1.21 \text{ m}^2\text{s}^{-2}$).
435 While the $\overline{w'^2}$ returns to the pre-burn conditions, the horizontal components remain elevated.

436 More specifically, $\overline{u'^2}$ and $\overline{v'^2}$ make up 53.0% and 38.5% of the average pre-burn TKE ,
437 respectively. During the burn period, the contribution to TKE from $\overline{u'^2}$ decreases slightly to
438 49.1% and the contribution from $\overline{v'^2}$ increases substantially to 43.3%. As noted earlier (Figures 6
439 and 7), the burn period also exhibits a larger range of horizontal and vertical wind components,
440 which is consistent with the larger range of TKE values in Figure 9.

441 In the post-burn period, the distribution of vertical velocity variance returns to the pre-
442 burn distribution. However, the range of values in the horizontal components are smaller during
443 the post-burn period than the burn period, but still larger than during the pre-burn period. The
444 medians of the horizontal TKE components are higher in the post-burn period than in either of
445 the other periods. While the $\overline{u'^2}$ outliers (above the 99.3rd percentile) decrease, the $\overline{v'^2}$ outliers
446 increase in magnitude. As was previously discussed, post-burn average wind directions differ
447 slightly from the pre-burn, accompanied by increases in the magnitude of the horizontal winds
448 (Figures 6 and 7). This result is consistent with elevated TKE values persisting into the period
449 after the end of the fire.

450 Additional analysis of the variance of the three velocity components enables an
451 assessment of turbulence anisotropy indicated by the ratio of $\overline{w'^2}$ to $2 \times TKE$. When this ratio
452 approaches $1/3$ for a given time period, the period can be said to experience an isotropic
453 turbulent regime (Heilman *et al.*, 2015). The mean $\overline{w'^2}$ for all the sonic anemometers is 0.0597
454 m^2s^{-2} for the pre-burn period, $0.0931 \text{ m}^2\text{s}^{-2}$ for the burn period, and $0.052 \text{ m}^2\text{s}^{-2}$ for the post-burn

455 period, which yields an anisotropy ratio of 0.042, 0.036, 0.021 for the pre-burn, burn and post-
456 burn periods, respectively. As the anisotropy ratios are well below 1/3 in all three periods, the
457 turbulence regime just above the combustion zone remains anisotropic at all time. It is worth
458 noting that in contrary to the belief that the increase in vertical velocity variance in response to
459 the surface heating during the burn should act to move turbulence towards a more isotropic
460 regime, the ratio here is slightly smaller during the burn period than the pre burn period largely
461 because the fire-induced increase in the cross-stream velocity variance is larger than the increase
462 in the vertical velocity variance. Heilman and Bian (2015) calculated the anisotropy ratios at 3
463 m above ground for two forest understory fires. The ratio decreased from 0.118 to 0.0718 from
464 pre-burn to burn in one experiment, but increased from 0.089 pre-burn to 0.13 in another
465 experiment. Since the sonic anemometers located on the western and southern sides of the burn
466 plot show no clear increase in $\overline{w'^2}$, the anisotropy ratio is also calculated for each sonic to verify
467 that the mean values did not mask anisotropy variations at individual locations in the burn plot.
468 No individual sonic anemometer reaches a ratio of 1/3, and the highest individual ratio (0.133) is
469 found at sonic anemometer A4 during the burn period. This result indicates that overall, the *TKE*
470 just above the combustion zone is highly anisotropic and is dominated by the horizontal
471 components for this burn. This result is not surprising as the sonic anemometers are located only
472 2.5 m above ground where horizontal turbulence would be expected to dominate over vertical
473 turbulence (Heilman *et al.*, 2015).

474

475 **3.3 Fire-Induced Shear Stress**

476 To address the question on how the surface fire alter turbulent momentum transfer
477 between the combustion zone and the atmosphere above, we next explore fire-induced changes
478 to turbulent momentum fluxes or shear stress measured by friction velocity described in Eq. (4).

479 Figure 10 shows time series of 1-minute averaged u_*^2 and the streamwise $\overline{u'w'}$ and cross-
480 stream $\overline{v'w'}$ stress components (the momentum flux), measured by each of the sonic
481 anemometers for the three periods. Kinematic momentum fluxes and u_*^2 are similar across all
482 the sonic anemometers during the pre-burn period, although three of the northernmost
483 instruments (A2, A3, and A4) indicate a negative spike in $\overline{u'w'}$ just before the start of the burn
484 period. These spikes contribute to an increase in u_*^2 at this time as well. It is unclear what
485 caused these features, but candidates include an anomalous burst of wind along the northern edge
486 of the burn plot and possible contamination of the wind data by activities of the burn managers
487 as they prepared to ignite the fire.

488 During the burn period, the values of $\overline{u'w'}$ and $\overline{v'w'}$ increase somewhat, leading to
489 increases in the u_*^2 values. The fire-induced changes generally increase in magnitude from west
490 (left) to east (right) and south to north, consistent with the fire-spread pattern. The largest
491 increase occur at the easternmost (right) locations, particularly A4 and C4 where u_*^2 values
492 nearly doubled. The smallest increases are not found at the westernmost locations, but at C2 and
493 D2. With a few exceptions, $\overline{u'w'}$ and $\overline{v'w'}$ are negative in the beginning of the burn period,
494 turning positive later in the period. The $\overline{u'w'}$ values exhibit the largest burn period variation at
495 A4, followed by B4, and similar patterns are observed for $\overline{v'w'}$. Overall, variations in u_*^2 suggest
496 an increase in shear stress magnitude in the burn period compared to the pre-burn period, with
497 the easternmost sonic anemometers recording 1-minute averaged values that are far greater than
498 the westernmost sonic anemometers.

499 During the post-burn period, some sonic anemometers (A2, B2, C1, C2, D2) recorded
500 higher u_*^2 than during the burn period, while others (A1, B1, B3, C2, C3, D3) recorded values

501 similar to the burn period. In either case, the average values are larger than during the pre-burn
502 period. The maximum post-burn values among all the sonic anemometers occur at A2 for u_*^2
503 and $\overline{v'w'}$ and C1 for $\overline{u'w'}$, both of which are larger than their burn-period peaks.

504 The overall distributions of u_*^2 , $\overline{u'w'}$, and $\overline{v'w'}$ from all 16 sonic anemometers are
505 depicted in Figure 11. During the pre-burn period, $\overline{u'w'}$ is negative, with a mean value of -0.015
506 $\text{m}^2 \text{s}^{-2}$, indicating an overall downward transfer of higher streamwise momentum air, which is
507 expected as wind speed usually increases with height. The mean of the cross-stream momentum
508 flux $\overline{v'w'}$ is near zero ($0.007 \text{ m}^2 \text{s}^{-2}$). However, the spread of the two components is similar, with
509 standard deviations of $0.057 \text{ m}^2 \text{s}^{-2}$ and $0.046 \text{ m}^2 \text{s}^{-2}$ for $\overline{u'w'}$ and $\overline{v'w'}$, respectively. The pre-burn
510 stress u_*^2 of $0.061 \text{ m}^2 \text{s}^{-2}$ ($u_* = 0.25 \text{ m}^2 \text{s}^{-2}$) is typical for daytime surface layers.

511 An increase in the downward (upward) transfer of higher streamwise (cross-stream)
512 momentum is observed during the burn period as the median values become more negative for
513 $\overline{u'w'}$ and more positive for $\overline{v'w'}$. However, the mean values change little from the pre-burn
514 period. The spread is doubled from a standard deviation of 0.046 to $0.098 \text{ m}^2 \text{s}^{-2}$ for $\overline{u'w'}$ and
515 nearly tripled from 0.05 to $0.124 \text{ m}^2 \text{s}^{-2}$ for $\overline{v'w'}$. The stronger upward transfer of cross-stream
516 momentum is consistent with the generation of cross-stream wind and updrafts in the vicinity of
517 the surface fire. Despite this overall fire-induced increase in $\overline{v'w'}$, the distribution of the cross-
518 stream momentum is negatively skewed by large negative outliers, suggesting occasional transfer
519 of higher cross-stream momentum by downdrafts near the vicinity of the fire. Both the mean and
520 standard deviation of u_*^2 values are doubled to $0.13 \text{ m}^2 \text{s}^{-2}$ and $0.086 \text{ m}^2 \text{s}^{-2}$, respectively, over the
521 pre-burn values. The peak 1-min averaged values of u_*^2 exceed $0.4 \text{ m}^2 \text{s}^{-2}$ (or a friction velocity
522 of 0.6 m s^{-1}), which is 2.5 times larger than the pre-burn values. Clements *et al.* (2008) also

523 observed a three-fold increase in friction velocity in their experiment involving a high intensity
524 grass fire, although the absolute values of the friction velocity in their experiment were five
525 times larger (1 and 3 m s⁻¹ before and during the fire) than the current experiment.

526 The mean post-burn u_*^2 value (0.10 m²s⁻²) is lower than that of the burn period but still
527 higher than the pre-burn value, driven primarily by the cross-stream component. The values of
528 the $\overline{v'w'}$ (0.0471 m² s⁻²) in the post-burn period is more than six times the pre-burn average
529 (0.0072 m²s⁻²), with a standard deviation (0.069 m²s⁻²) that is between the pre-burn period
530 (0.046) and burn period (0.096) values. The mean friction velocity therefore does not return to
531 the pre-burn average, although it is lower than the average during the burn period. Other
532 experiments (e.g. Clements *et al*, 2008; Heilman, *et al.* 2019) noted a return of friction velocity
533 to pre-burn values soon after the passage of the fire front, during a period when smoldering was
534 occurring. The results of this analysis suggest that friction velocities do not quickly return to pre-
535 burn values on all fires.

536

537 **3.4 Fire-Induced Turbulent Heat Flux**

538 We proceed to examine the impact of the fire on turbulent heat flux. Time series of 1-
539 minute average kinematic turbulence sensible heat flux $\overline{T'w'}$ for each sonic anemometer are
540 shown in Figure 12 for the three periods, which also shows the overall distribution of heat fluxes
541 for all the sonic anemometers. In the pre-burn period, the sonic anemometers recorded
542 background $\overline{T'w'}$ values that averaged around 5.25×10^{-2} °C m s⁻¹ (or 52.7 W m⁻² after multiplying
543 by the density and heat capacity of air), with a standard deviation of 3.41×10^{-2} °C m s⁻¹ (34 W m<sup>-
544 2</sup>). During the burn period, a fire-induced increase in $\overline{T'w'}$ is evident at all but the westernmost

545 sonic anemometers (A1, B1, C1, and D1), with larger increases appearing at the easternmost
546 locations. The largest $\overline{T'w'}$ values generally occur early in the burn period, with the A4 sonic
547 having the largest $\overline{T'w'}$ value of $2.13 \text{ }^\circ\text{C m s}^{-1}$ (2.138 kW m^{-2}). Based on the IR imaging (Figure
548 4), after the first three minutes of the burn period there is a slight shift in the burn direction
549 towards the southeastern side of the plot. This shift in direction is apparent in the time series for
550 the D4 sonic anemometer, which is located on the southeastern corner of the burn plot, where
551 elevated $\overline{T'w'}$ values are recorded late in the burn period, at a time when the values have
552 dropped at most of the other sonic anemometers. The overall distribution of the burn-period $\overline{T'w'}$
553 is skewed by larger values since the plot mean was 0.268 K m s^{-1} (269 W m^{-2}) but the median
554 was just $0.0974 \text{ }^\circ\text{C m s}^{-1}$ (98 W m^{-2}).

555 Values of $\overline{T'w'}$ during the post-burn period quickly drop back to just slightly above the
556 pre-burn values, with a mean of $6.35 \times 10^{-2} \text{ }^\circ\text{C m s}^{-1}$ (64 W m^{-2}) and a standard deviation of
557 $3.76 \times 10^{-2} \text{ }^\circ\text{C m s}^{-1}$ (38 W m^{-2}). However, the post-burn period contains several outliers (above the
558 99.3% percentile), indicating the influence of smoldering on some of the sonic anemometers
559 even after the fire has exited the burn plot. A specific example of the smoldering effect is the D4
560 sonic anemometer, where the post-burn $\overline{T'w'}$ ($0.126 \text{ }^\circ\text{C m s}^{-1}$ or 126 W m^{-2}) is about twice the
561 pre-burn value. The overall modest increase of $\overline{T'w'}$ in the post-burn period compared to the pre-
562 burn period was also observed in the two wildland fire experiments described in Heilman *et al.*
563 (2019).

564

565 **3.5 Quadrant Analyses**

566 **3.5.1 Turbulent heat fluxes**

567 The analysis above provided a quantitative assessment of fire-induced changes to the
568 turbulent heat and momentum fluxes through comparisons of flux values between the pre-burn
569 and the burn periods. However, such analysis cannot reveal what types of heat or momentum
570 transfer events are mostly affected by the fire. We apply the quadrant analysis method (also
571 known as sweep-ejection analysis) described earlier (Table 1) to the observed turbulent fluxes to
572 provide additional insight into how the fire changes the composition of heat and momentum
573 fluxes. By partitioning the total heat and momentum fluxes into four quadrants representing
574 different types of flux events, the quadrant or sweep-ejection analysis allows for the delineation
575 of the fire influence on specific types of turbulent heat and momentum transfer processes.

576 Figure 13 shows the relative contributions and the proportional number of occurrence of
577 the different heat-flux events (i.e., sweeps, ejections, outward interactions and inward
578 interactions) during each period, observed by each of the 16 sonic anemometers. During the pre-
579 burn period, the partitioning among the four types of events (see Table 1) by contribution and
580 proportion exhibits little variation across the 16 sonic anemometers. At all locations, the ejection
581 and sweep dominate, accounting for over 60% of the total events, with sweep being slightly
582 larger. The rest is split between outward interaction and inward interaction events, with the
583 former slightly outnumbering (20-23%) the latter (14-19%). A similar partitioning is observed
584 for the event contributions for the heat fluxes, but the ejection events, despite being slightly less
585 frequent, contribute more to the heat flux than do the sweep events. This apparent inconsistency
586 between the partitioning of the event number and the event contribution suggests that ejection
587 events likely involve larger eddies and stronger heat transfer compared to sweep events. This
588 pre-burn period partitioning is similar to previous ambient daytime measurements observed in
589 other studies (e.g., Heilman *et al.*, 2021).

590 The burn period is marked by substantial heterogeneity across the 16 sonic anemometers.
591 Despite differences in the magnitudes of contributions to the heat fluxes amongst the sonic
592 anemometers, the increases in the overall positive mean heat flux during the burn period can be
593 largely attributed to increases of ejection events that contribute to positive heat fluxes through
594 upward transfer of warmer air from the combustion zone to the atmosphere above. There is also
595 an increase in the negative contribution from inward interaction events, which represents the
596 downward transfer of warmer air from the atmosphere to the combustion zone. The contributions
597 to the overall mean heat flux by the other two types of events, sweep and outward interaction,
598 show little change from the pre-burn to the burn periods, which suggests that the turbulent heat
599 transfer processes represented by these types of events, namely downward transfer of colder air
600 from above to the surface or upward transfer of colder air from the combustion zone to the
601 atmosphere, are not very sensitive to the presence of a low-intensity fuel-bed-scale surface fire.

602 Compared to the partitioning in event contribution, the fire-induced changes to the
603 partitioning in event number are less clear. In general, the sonic anemometers that show an
604 increase in the contribution by inward interaction events also exhibit an increase in the number
605 of inward interaction events from the pre-burn to the burn periods. However, an increased
606 contribution to the overall mean heat flux by ejection events does not correspond to an increase
607 in the number of the ejection events. The increased number of sweep events are in agreement
608 with the increased sweep contributions at several sonics (A2-A4 and B2-B4), although the sweep
609 contributions are overwhelmed by that of the ejection contributions at these sonic anemometers.

610 A key finding from this heat flux sweep-ejection analysis is that turbulent heat fluxes
611 during the burn period are overwhelmingly dominated by ejection events, but there is usually a
612 small or no increase in the number of ejection events. This suggests that the presence of a low-

613 intensity fuel-bed-scale fire does not necessarily produce more upward turbulent heat transfer
614 events, but instead, it produces stronger events that quickly transfer and diffuse the sensible heat
615 generated by combustion into the ambient atmosphere above.

616 During the post-burn period, most sonic anemometers show vertical heat flux values that
617 are smaller than the burn period but still larger than the pre-burn period. The largest contribution
618 to the overall mean heat flux is usually from sweep events, accompanied also by an increase in
619 the number of the events, indicating the occurrence of many events where cold air is transferred
620 downward. The post-burn period also exhibits an increase in the heat-flux contributions from
621 outward interaction events, which represent downward transfer of warm air. Similar to the burn
622 period, inward interaction events, both in contribution and number, vary considerably across the
623 sonic array.

624 Figure 14 shows the partitioning of both the event number and the event contribution to
625 turbulent heat fluxes using data from all 16 sonic anemometers, which highlights more clearly
626 how the fire modifies the overall heat flux regime. Similar to the heat flux quadrant analysis for
627 individual sonic anemometers, the heat flux events averaged across the sonic anemometer array
628 for the pre-burn period is dominated by sweep (32%) and ejection (28%) events. Inward
629 interaction events occur with the least proportion (17%), followed by outward interaction events
630 (23%). The sweep and ejection events, which contribute to positive heat fluxes, are much larger
631 in magnitude than the negative heat flux contributions from the inward and outward interaction
632 events. The dominance of sweep and ejection events for the turbulent heat fluxes during the pre-
633 burn period follows observations made in previous studies (Heilman *et al.*, 2021).

634 The combined proportions of sweep and ejection events (both contributing to positive
635 heat fluxes) and the outward and inward interaction events (both contributing to negative heat

636 fluxes) remain similar between the burn and the pre-burn period. However, between the two
637 types of events in each group, one (sweep, inward interaction) increases and the other (ejection,
638 outward interaction) decreases in proportion. Previous fire experiments also reported an increase
639 in sweep events and a generally proportional decrease in ejection events (Heilman *et al.*, 2021),
640 but the magnitudes of the changes are larger than what is observed here, likely because the
641 previous fires are more intense. Additionally, modest changes in the partitioning of the event
642 number and contributions for this fire could be a byproduct of combining data from sonic
643 anemometers that are not strongly affected by the fire front (i.e. the westernmost sonic
644 anemometers) with those that experience more substantial changes.

645 The large changes in the contributions of the heat flux events during the burn period
646 suggest that this fire has greater impacts on the event contributions to the mean turbulent heat
647 fluxes than on the event number. Specifically, ejection event contributions dominate in the burn
648 period, making up 70.4% of the total contribution, while sweep and outward interaction
649 contributions decrease by a third and a sixth, respectively, compared to their contributions during
650 the pre-burn period. The magnitude of the contribution from inward interaction events increases
651 slightly but is quite similar to the contribution during the pre-burn period.

652 Heat flux events in the post-burn period more closely resemble the pre-burn period than
653 the burn period, but the event contributions and the event number do not return entirely to their
654 pre-burn values. As noted in the analyses of *TKE* and kinematic heat flux (Figures 9 and 11), this
655 result is consistent with smoldering occurring in the burn plot during the post-burn period. The
656 sweep event contribution during the post-burn period is 1.5 times higher than during the pre-burn
657 period and 1.3 times higher than during the burn period. Compared to the pre-burn values, the
658 post-burn period event contributions are slightly higher for outward interaction events and

659 slightly lower for ejection and inward interaction events. Overall, the post-burn period is
660 dominated by contributions from sweep events (37.7%), which is followed by ejection event
661 (25.3%) although lower than pre-burn values. These results differ somewhat from the Heilman *et*
662 *al.* (2021) in that they reported both sweep and ejection events returning to pre-burn values,
663 while only ejection events return to pre-burn values for this fire.

664

665 **3.5.2 Turbulent momentum fluxes**

666 Quadrant analysis is also applied to partition the vertical turbulent kinematic flux of
667 horizontal momentum $\overline{S'w'}$ into four different types and the results for each of the 16 sonic
668 anemometers are shown in Figure 15. During the pre-burn period, the overall mean momentum
669 fluxes are negative at all but two sonic anemometers (C1, C2) where the flux is slightly positive.
670 Between the two types of events that contribute to negative momentum fluxes, the sweep events
671 (downward transfer of higher horizontal momentum air from the atmosphere to the combustion
672 zone) contribute more than the ejection events (upward transfer of lower horizontal momentum
673 air from the combustion zone to the atmosphere above), which is consistent with the slightly
674 higher number of sweep events than ejection events. Between the two types of events that
675 contribute to positive momentum fluxes, the outward interaction events (upward transfer of
676 higher horizontal momentum air from the combustion zone to the atmosphere above) contribute
677 more than the inward interaction events (downward transfer of lower horizontal momentum air
678 from the atmosphere to the combustion zone), although the number of the inward and outward
679 interaction events is similar.

680 The changes from the pre-burn period to the burn period vary substantially by location,
681 but the sign of the overall mean momentum fluxes remains unchanged at most locations. The
682 most pronounced and consistent change across the anemometer array is a substantial increase in
683 the proportional number of inward interaction events and, to a lesser degree, the contribution
684 from these events. The ejection events also exhibit an increase in the number and the
685 contribution at most of the sonic anemometer locations. There is a general decrease in the
686 number of sweep and outward interaction events, but the contributions are not consistent, with
687 some sonic anemometers showing an increase while others experience a decrease in contribution.

688 An exception to the above general observations between the pre-burn and burn periods is
689 B4, where the overall momentum flux shifts from negative to positive due to an increase in
690 outward interaction contribution by as much as 5 times the pre-burn magnitude. The amount of
691 increase in the contribution from the outward interaction events, however, does not match the
692 small increase (approximately 10%) in the event number, which suggests that the increase in the
693 overall momentum flux magnitude at this location is likely due a small number of extremely
694 strong events of upward transfer of higher horizontal momentum air associated with large,
695 energetic eddies generated by the surface fire.

696 The large heterogeneity in the event contribution values for the momentum fluxes across
697 the sonic anemometer array during the burn period dissipated substantially into the post-burn
698 period. The event contribution and event number distributions once again become less dependent
699 on the locations of the sonic anemometers. Despite this tendency to return to the pre-burn
700 distribution, the post-burn period experiences larger contributions from, and higher number of
701 ejection and inward interaction events than sweep and outward interaction events, which is
702 opposite to the pre-burn period and similar to the burn period.

703 Figure 16 shows a quadrant analysis that combines data from all the sonic anemometers,
704 which allows for an assessment of how the fire modified the momentum flux turbulence regime
705 for the entire burn plot. Overall, sweep (31.9%) and outward interaction (26.6%) events
706 dominate the momentum flux contributions in the pre-burn period. The increases in the
707 proportion of inward interaction and ejection events from the pre-burn to the burn periods make
708 the contributions more balanced across the four quadrants, suggesting that the different event
709 contributions are more similar to each other during the burn than the pre-burn period. In the post-
710 fire period, inward interaction events contribute more to the mean momentum flux (25.7%) than
711 during the pre-fire period (18.1%). The event number distributions in the combined analysis
712 echoes the results from the individual sonic anemometers, with the pre-burn period showing
713 similar values for all four quadrants, a sharp increase in inward interaction events and decrease in
714 outward interaction events during the burn period, and fewer inward interaction events during the
715 post-burn period than during the burn period but more numerous than during the pre-burn period.

716 The results of the quadrant analysis of momentum fluxes presented above are somewhat
717 different from those of previous studies involving operational-scale prescribed burns. Heilman *et*
718 *al.* (2021) showed that during an intense grass fire and two low-intensity forest understory fires,
719 there can be substantial increase in the number and contribution of sweep and outward
720 interaction events and that the increase in the positive momentum flux from outward interaction
721 events largely offset the increase in the negative flux associated with sweep events. Whereas in
722 the small fuel-bed scale burn here, inward interactions occur most frequently, followed by
723 ejection events. However, the ejection event contributions to the mean momentum flux are larger
724 (32.3%), with the inward interaction event contributions (24.2%) more similar to the outward
725 interaction (23.4%) contributions. The feature of increased frequency of inward interaction

726 events and their increased contribution to the mean momentum flux compared to previous burns
727 is further observed in the post-burn period.

728 The event number and event contributions during the post-burn period also differ with
729 increased ejection and inward interactions events, 32.8% and 20.6%, while the large-scale burns
730 in Heilman *et al.* (2021) showed a closer return to pre-fire periods, with sweep and ejection
731 events making up the majority of event number and contributions. The contributions from sweep,
732 inward interaction, and ejection events remain elevated during the post-burn period, while the
733 contributions from outward interaction decrease during post-burn to values lower than the values
734 of the pre-burn period.

735

736 **4. Summary**

737 This study presents the atmospheric turbulence dynamics observed through a 4×4 array
738 of fast-response 3D sonic anemometers during a low-intensity fire experiment on a 10 m x 10 m
739 burn plot in the Silas Little Experimental Forest in New Jersey, USA. The density of turbulence
740 measurements is unprecedented for fire experiments, allowing for a deeper analysis of
741 heterogeneities as the surface line-fire spread through the burn plot than was previously possible.
742 The analysis focuses on assessments of the fire impacts on turbulence intensity, as measured by
743 *TKE*, turbulent momentum flux or shear stress as measured by friction velocity, and turbulent
744 heat flux.

745 The influence of the low-intensity surface line-fire on the atmosphere above the
746 combustion zone is evidenced by an increase in temperature up to 20 °C, the generation of strong
747 updrafts up to 6 m s^{-1} and downdrafts up to -5 m s^{-1} and a decrease in the streamwise velocity

748 coupled with an increase in the cross-stream velocity indicating horizontal convergence in the
749 vicinity of the fire front. The observed fire exhibited behavior more consistent with a buoyant
750 plume than mechanically forced rising motion resulting from converging surface air. The
751 influence of the fire on horizontal velocity components persisted longer after fire front passage
752 while the influence on vertical velocity subsided rapidly behind the fire front.

753 The fire modified turbulence characteristics at the fuel bed-atmosphere interface. There
754 was an increase in the turbulence intensity, with *TKE* values 2-3 times higher than the ambient
755 environment, due primarily to the increase in cross-stream velocity variance and, to a lesser
756 degree, the increase in the vertical velocity and streamwise velocity variance. Heilman *et al.*
757 (2017) also reported two to threefold increases in *TKE* values during two operational-scale low-
758 intensity forest understory prescribed fires. It is interesting to note that this increase in *TKE* is
759 only slightly smaller than what was observed during the intense grass fire during FireFlux
760 (Clements *et al.*, 2007), although the magnitude of *TKE* of the intense grass fire is substantially
761 larger than that of the low-intensity fires. Despite this increase in *TKE*, the value of *TKE* was still
762 smaller than what is expected in an environment of substantial turbulence. Additionally, despite
763 the increase in the vertical velocity variance during the fire, the *TKE* was still dominated by the
764 horizontal velocity variance, indicating that the turbulence regime remained anisotropic
765 (anisotropic ratio $\ll 1/3$) above the combustion zone of this low-intensity fuel-bed-scale surface
766 fire.

767 The fire enhanced upward sensible heat fluxes substantially by as much as 40 times the
768 flux in the ambient atmosphere (from 50 W m^{-2} to 2 kW m^{-2}). This change in the sensible heat
769 flux is largely attributable to an increased contribution of upward transfer by turbulent eddies of
770 warmer air from the combustion zone to the atmosphere above, which is also known as ejection

771 events for vertical turbulent heat transfer. This increase in the contribution of the ejection events
772 to turbulent heat fluxes was not caused by a corresponding increase in the number of ejection
773 events that changed little from the pre-burn to burn periods. This mismatch between the ejection
774 event contribution and event number suggests that the presence of a low-intensity fuel-bed-scale
775 fire may not necessarily produce more upward turbulent heat transfer events, but rather, it can
776 produce strong ejection events associated with large, energetic eddies. The warmer air
777 transported upward by the ejection events can also be transported downward by inward
778 interaction events, which also increased somewhat during the fire.

779 Compared to the turbulent heat flux, the impact of the fire on turbulent momentum flux
780 or shear stress was less pronounced. In general, an increase in momentum fluxes was observed
781 during the burn, with friction velocity, a measure of total shear stress on horizontal wind, 2-3
782 times the ambient value (from $\sim 0.25 \text{ ms}^{-1}$ to 0.6 ms^{-1}). Previous studies of operational-scale
783 grass fire or forest understory fires also found up to a 3-fold increase in friction velocity despite
784 that the scale of this fire is much smaller than the previous fires and that the absolute values of
785 friction velocity during the intense grass fire were 5 times higher than the low-intensity fire here
786 (Clements *et al.*, 2007; Heilman *et al.*, 2017; 2021). The fire was accompanied by an increase in
787 the downward transfer of lower horizontal momentum air, also known as inward interaction
788 events, along with a smaller increase in the upward transfer of lower horizontal momentum air
789 referred to as ejection events. This finding differs from previous observations during an
790 operational-scale forest understory fire where an increase in sweep (downward transfer of higher
791 horizontal momentum air) and outward interaction (upward transfer of higher horizontal
792 momentum air) contributions to the mean momentum fluxes were detected (Heilman *et al.*,
793 2021).

794 These findings directly address the initial research inquiries: How does the surface fire
795 impact turbulence intensity and the exchanges of turbulent heat and momentum between the
796 combustion zone and the atmosphere above? Additionally, the investigation delves into whether
797 and how the fire alters the distribution of heat and momentum fluxes into different event types,
798 considering both event number and contribution.

799 Perhaps the most significant finding from this study is the large variations in the observed
800 fire-induced perturbations across the sonic anemometer array in the burn plot. This directly
801 corresponds to the third question raised in the introduction: How do the fire-induced
802 modifications on turbulence vary spatially across the burn plot? The anemometers on the western
803 side of the burn plot where a surface line-fire was ignited picked up very weak or no signals of
804 the fire despite the proximity to the initial fire line. In contrast, the sonic anemometers in the
805 center or eastern side of the burn plot picked up clear fire signals. Although the features of fire-
806 induced turbulence regime (e.g., anisotropy, sweep-ejection dynamics) revealed by the sonic
807 anemometers are similar, the magnitudes vary with downwind distance and the relative position
808 of the sonic anemometers to the impinging fire front. Considering the size of the burn plot (10 m
809 x 10 m) and the homogeneity of consumed fuels, this finding suggests that considerable care
810 should be taken when comparing, contrasting, and combining data from multiple fires or from
811 multiple instruments on the same fire to ensure that significant fire signals are not being over- or
812 under-represented in the analyses that inform the conclusions of the studies. This also calls into
813 question of using numerical simulations from coupled atmosphere-fire behavior models with
814 horizontal grid spacing ≥ 10 m. The results presented here suggest that 1-2 m grid spacing is
815 necessary for model simulations to capture atmospheric turbulent circulations that have
816 spatiotemporal scales similar to the scales associated with flame dynamics in the combustion

817 zone. It is however, impractical for operational applications to use such fine resolution.
818 Operational models, with resolutions ranging from tens to hundreds of meters, often fall within
819 the so called 'gray zone' where turbulence is partially resolved and existing turbulence closure
820 schemes designed to parameterize all turbulent motions are inadequate. Advancements in
821 computing technology have brought this zone to the forefront of operational model simulations.
822 Developing turbulence closure schemes for this scale is an active area of research. Large-eddy
823 simulation (LES) models, validated using laboratory data, are instrumental in this endeavor. The
824 experiments described in this study, capturing fire-induced turbulence on a 10 m x 10 m plot, can
825 play a crucial role in developing turbulence parameterizations for the gray zone when combined
826 with LES models.

827 Future work will compare results from this case with those of other burns in the SERDP
828 10 m x 10 m fuel-bed-scale burn series to delineate the effect of fuel and ambient atmospheric
829 conditions on fire-atmosphere interactions and with results from other prescribed-fire
830 experiments to help scale up or scale down the results between small-scale and operational scale
831 fires. Future work will also include the reanalysis of 10 Hz sonic anemometer data from other
832 fire experiments using some or all of the methodologies employed here, which could contribute
833 to the identification and documentation of a series of steps, protocols, standards, and
834 methodologies by which 10-Hz sonic anemometer data collected during fire experiments can be
835 compared and contextualized. Additionally, forthcoming analyses will integrate data collected
836 from other instruments deployed during these fuel-bed-scale fire experiments. For instance,
837 examining the high-frequency thermocouple vertical profile (0, 5, 10, 20, 30, 50, 100 cm) in
838 conjunction with infrared data can offer significant insights into the vertical variation of
839 temperature between the combustion zone and the atmosphere immediately above. Finally yet

840 importantly, employing spectral and co-spectral analyses will be essential in revealing the
841 temporal and spatial scale of turbulence regimes at the fuel-bed and atmosphere interface. These
842 analyses will simultaneously enable a holistic exploration of the oscillatory behavior tied to line
843 fires.

844 Another facet to delve into in future research involves the generation of vorticity, a
845 consequential byproduct of fires that significantly influences fire behavior. Estimating fire-
846 induced vorticity from field observations presents a formidable challenge, necessitating a
847 carefully designed instrument array capable of capturing both horizontal and vertical variations
848 in wind velocity. Despite these challenges, the utilization of the 4x4 sonic anemometer array in
849 the 10m x 10m burn plot provides a distinctive opportunity. This array captures horizontal
850 variations in wind velocity as the line fire spreads through the plot, offering a unique opportunity
851 for estimating vertical vorticity associated with line fires. However, it is important to note that
852 estimating horizontal vorticity is not feasible due to the sonic anemometer array's velocity
853 measurement on a single vertical level (2.5 m), which does not capture the necessary vertical
854 variations of velocity for horizontal vorticity calculation. Future experiments will require
855 deploying a densely spaced sonic anemometer similar to the current one but at multiple vertical
856 levels to comprehensively evaluate vorticity associated with fires.

857 Because the burn period was chosen to be between the time when the first and the last
858 sonic anemometers have temperatures satisfying the threshold value (eight standard deviations in
859 these analyses), the burn period included time after the fire has passed the sonic anemometer
860 location, which likely yielded an underestimation of the fire effect. Similarly, the inclusion of all
861 16 sonic anemometers in the analysis, including those that registered little fire signal, likely
862 contributed to an underestimation. Consequently, fire-induced turbulent circulations and the

863 associated turbulent heat and momentum fluxes are likely to be stronger than what has been
864 reported here.

865

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872 instrument deployment of the 10 m x 10 m burn series. Our gratitude also extends to the two
873 anonymous reviewers for their insightful and constructive comments, which have undeniably
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875

876 **Code and Data Availability**

877 Python language was used for all analyses and data management, with the Pandas package
878 (<https://zenodo.org/record/7037953#.Yw-at3bMIp4>) used for data processing, NumPy package
879 (<https://numpy.org/>) used for most statistical calculations and Matplotlib visualization package
880 (<https://matplotlib.org/>) used for plotting, , all of which are open source packaged in the Python
881 environment. The computer codes and the data are hosted on software sharing and version
882 control website and service GitHub. [https://github.com/JosephSeitz/SERDP-10x10meter-Burn-](https://github.com/JosephSeitz/SERDP-10x10meter-Burn-Cleaner)
883 [Cleaner](#).

884 All data used in this study are publicly archived and available via the USFS Data Archive (in
885 press, links to be included in revised version).

886

887 **Author Contributions**

888 All authors contributed to the research design. K.C., N.S., M.G., M.P., R.H. and E.M. conducted
889 the fire experiment and collected the data. J.C. and M.P., with assistance from K.C., did the
890 initial process and formatting of the data. J.S., with assistance and guidance from J.J.C. and
891 discussions and feedback from S.Z., W.H., X.B. M.K., performed the data analysis and produced
892 all the plots. S.Z. wrote the manuscript and was responsible for the revision. M.G., W.H., K.C.
893 and N.S. edited the initial manuscript.

894

References

- Amaya, M.A. and Clements C.B. (2020) Evolution of plume core structures and turbulence during a wildland fire experiment. *Atmosphere*, **11**, 842.
- Bennie, J., Huntley, B., Wiltshire, A., Hill, M.O. and Baxter, R. (2008) Slope, aspect and climate: Spatially explicit and implicit models of topographic microclimate in chalk grassland. *Ecological Modeling*, **216**, 47-59.
- Billmire, M., Frenc, N.H.F., Loboda, T., Owen, R.C. and Tyner, M. (2014) Santa Ana winds and predictors of wildfire progression in southern California, *International Journal of Wildland Fire*, **23**, 1119-1129.
- Calviño-Cancela, M, Chas-Amil, M.L., García-Martínez, E.D. and Touza, J. (2017) Interacting effects of topography, vegetation, human activities and wildland-urban interfaces on wildfire ignition risk. *Forest Ecology and Management*, **397**, 10-17.
- Campbell-Lochrie, Z., Walker-Ravena, C., Gallagher, M., Skowonski, N., Mueller, E., Hadden, R.M. (2021) Investigation of the role of bulk properties and in-bed structure in the flow regime of buoyancy-dominated flame spread in porous fuel bed. *Fire Safety Journal*, **120**, <https://doi.org/10.1016/j.firesaf.2020.103035>
- Campbell-Lochrie, Zakary J.; Hadden, Rory M.; Mueller, Eric V.; Walker-Ravena, Carlos; Gallagher, Michael R.; Clark, Kenneth L.; Hom, John L.; Kremens, Robert L.; Cole, Jason A.; Patterson, Matthew M.; Everland, Alexis I.; Skowronski, Nicholas S. 2022. Multi-scale analyses of wildland fire combustion processes: Small-scale field experiments -

Transportable Analyzer for Calorimetry Outside (TACO). Fort Collins, CO: Forest Service Research Data Archive.

Carrier, G.F., Fendell, F.E. and Wolff, M.F. (1991) Wind-aided fire spread across arrays of discrete fuel elements. I. Theory. *Combustion Science and Technology*, **75**, pp.31-51.

Clark, Kenneth L.; Gallagher, Michael R.; Mueller, Eric V.; Hadden, Rory M.; Walker-Ravena, Carlos; Campbell-Lochrie, Zakary J.; Cole, Jason A.; Patterson, Matthew M.; Everland, Alexis I.; Skowronski, Nicholas S. 2022a. Multi-scale analyses of wildland fire combustion processes: Small-scale field experiments - three-dimensional wind and temperature. Fort Collins, CO: Forest Service Research Data Archive.

Clark, Kenneth L.; Gallagher, Michael R.; Mueller, Eric V.; Hadden, Rory M.; Walker-Ravena, Carlos; Campbell-Lochrie, Zakary J.; Cole, Jason A.; Patterson, Matthew M.; Everland, Alexis I.; Skowronski, Nicholas S. 2022b. Multi-scale analyses of wildland fire combustion processes: Small-scale field experiments - temperature profile. Fort Collins, CO: Forest Service Research Data Archive.

Clark, K.L., Heilman, W.E., Skowronski, N.S., Gallagher, M.R., Mueller, E., Hadden, R.M., and Simeoni, A. (2020) Fire behavior, fuel consumption, and turbulence and energy exchange during prescribed fires in pitch pine forests. *Atmosphere*, **11**, 242.

Clark, T.L., Jenkins, M.A., Coen, J.L. and Packham, D.R., (1996) A coupled atmosphere-fire model: Role of the convective Froude number and dynamic fingering at the fireline. *International Journal of Wildland Fire*, **6**, pp.177-190.

Clements, C.B., and Seto, D. (2015) Observations of fire-atmosphere interactions and near-surface heat transport on a slope. *Boundary-Layer Meteorology*, **154**, 409-426.

Clements, C.B., Kochanski, A.K., Seto, D., Davis, B., Camacho, C., Lareau, N.P., Contezac, J., Restaino, J., Heilman, W.E., Krueger, S.K. and Butler, B. (2019) The FireFlux II experiment: a model-guided field experiment to improve understanding of fire–atmosphere interactions and fire spread. *International Journal of Wildland Fire*, **28**, 308-326.

Clements, C.B., Kochanski, A.K., Seto, D., Davis, B., Camacho, C., Lareau, N.P., Contezac, J., Restaino, J., Heilman, W.E., Krueger, S.K., Butler, B., Ottmar, R.D., Vihnanek, R., Flynn, J., Filippi, J.B., Barboni, T., Hall, D.E., Mandel, J., Jenkins, M.A., O'Brien, J., Hornsby, B., and Teske, C. (2019) The FireFlux II experiment: a model-guided field experiment to improve understanding of fire–atmosphere interactions and fire spread. *International Journal of Wildland Fire*, **28**, 308-326.

Clements, C.B., Zhong, S., Bian, X., Heilman, W.E., and Byun, D.W. (2008), First observations of turbulence generated by grass fires. *Journal of Geophysical Research*, **113**, D22102.

Clements, C.B., Zhong, S., Goodrick, S., Li, J., Potter, B.E., Bian, X., Heilman, W.E., Charney, J.J., Perna, R., Jang, M. and Lee, D. (2007) Observing the dynamics of wildland grass fires: FireFlux—A field validation experiment. *Bulletin of the American Meteorological Society*, **88**, 1369-1382.

Di. Christina, G., Gallagher, M., Skowonski, N., Simeoni, In, S.-K. (2022) Design and implementation of a portable large-scale wind tunnel for wildfire research. *Fire Safety Journal*, **131**, 103607.

Ebel, B.A. (2013) Simulated unsaturated flow processes after wildfire and interactions with slope aspect. *Water Resources Research*, **49**, 8090-8107

- Finney, M.A., Cohen, J.D., Forthofer, J.M., McAllister, S.S., Golner, M.J., Gorham, D.J., Saito, K., Akafuah, N.K., Adam, B.A., and English, J.D. (2015) Role of buoyant flame dynamics in wildfire spread. *Proceedings of the National Academy of Sciences*, **112**, 9833-9838.
- Forthofer, J.M., and Goodrick, S.L. (2011) Review of vortices in wildland fire. *Journal of Combustion*, **2011**, Article ID 984363.
- Gallagher, Michael R.; Skowronski, Nicholas S.; Hadden, Rory M.; Mueller, Eric V.; Clark, Kenneth L.; Campbell-Lochrie, Zakary J.; Walker-Ravena, Carlos; Kremens, Robert L.; Everland, Alexis I.; Patterson, Matthew M.; Cole, Jason A.; Heilman, Warren E.; Charney, Joseph J.; Bian, Xindi; Mell, William E.; Hom, John L.; Im, Seong-kyun; Kiefer, Michael T.; Zhong, Shiyuan; Simeoni, Albert J.; Rangwala, Ali; Di Cristina, Giovanni. (2022) Multi-scale analyses of wildland fire combustion processes: Small-scale field experiments – plot layout and documentation. Fort Collins, CO: Forest Service Research Data Archive
- Heilman, W.E. (2021) Atmospheric turbulence in wildland fire environments: implications for fire behavior and smoke dispersion. *Fire Management Today*. **79**, pp.24-29.
- Heilman, W.E. and Bian, X. (2013) Climate variability of near surface turbulent kinetic energy over the United States: Implications for fire weather prediction. *Journal of Applied Meteorology and Climatology*, **52**, 753-772.
- Heilman, W.E., Barnerjee, T., Clements, C.B., Clark, K.L., Zhong, S., and Bian X. (2021) Observations of sweep-ejection dynamics for heat and momentum fluxes during wildland fires in forested and grassland environments. *Journal of Applied Meteorology and Climatology*, **60**, 185-199.

- Heilman, W.E., Bian, X., Clark, K.L. and Zhong, S. (2019) Observations of turbulent heat and momentum fluxes during wildland fires in forested environments. *Journal of Applied Meteorology and Climatology*, **58**, pp.813-829.
- Heilman, W.E., Bian, X., Clark, K.L., Skowronski, N.S., Hom, J.L. and Gallagher, M.R. (2017) Atmospheric turbulence observations in the vicinity of surface fires in forested environments. *Journal of Applied Meteorology and Climatology*, **56**, 3133-3150.
- Heilman, W.E., Clements, C.B., Seto, D., Clark, K.L., Skowronski, N.S., and Hom, L.J. (2015) Observations of fire-induced turbulence regimes during low-intensity wildland fires in forested environments: Implications for smoke dispersion. *Atmospheric Science Letters*, **16**, 453-460.
- Heilman W.E. and Bian, X. (2013)
- Katul, G., Poggi, D., Cava, D., and Finnigan, J. (2006) The relative importance of ejections and sweeps to momentum transfer in the atmospheric boundary layer. *Bound.-Layer Meteor.*, **120**, 367–375.
- Katul, G., Kuhn, G., Schieldge, J., and Hsieh, C.-I. (1997) The ejection sweep character of scalar fluxes in the unstable surface layer. *Bound.-Layer Meteor.*, **83**, 1–26.
- Kitzberger, T., Falk, D.A., Westerling, A.L., and Swetnam T.W. (2017) Direct and indirect climate controls predict heterogeneous early-mid 21st century wildfire burned area across western and boreal North America. *PLOS ONE*, **12**, e0188486.
- Kremens, Robert L.; Gallagher, Michael R.; Clark, Kenneth L.; Mueller, Eric V.; Hadden, Rory M.; Heilman, Warren E.; Charney, Joseph J.; Hom, John L.; Campbell-Lochrie, Zakary J.;

- Walker-Ravena, Carlos; Everland, Alexis I.; Cole, Jason A.; Patterson, Matthew M.; Skowronski, Nicholas S. 2022. Multi-scale analyses of wildland fire combustion processes: Small-scale field experiments - fire radiative power. Fort Collins, CO: Forest Service Research Data Archive.
- Littell, J.S., Peterson, D.L., Riley, K.L., Liu, Y. and Luce, C.H. (2016). A review of the relationships between drought and forest fire in the United States. *Global Change Biology*, **22**, 2353-2369.
- Moody, M.J., Gibbs, J.A., Kruger, S., Mallia, D., Pardyjak, E.R., Kochanski, A.K., Bailey, B.N., Stoll, R. (2022) QES-Fire, a dynamically coupled fast-response wildfire model. *International Journal of Wildland Fire*, **31**, 306-325.
- Mueller, E.V., Skowronski, N., Clark, K., Gallagher, M., Kremens, R., Thomas, J.C., El Houssami, M., Filkov, A., Hadden, R.M., Mell, W.; et al. (2017) Utilization of remote sensing techniques for the quantification of fire behavior in two pine stands. *Fire Safety Journal*, **91**, 845–854, doi:10.1016/j.firesaf.2017.03.076.
- Potter, B.E. (1996) Atmospheric properties associated with large wildfires. *International Journal of Wildland Fire* **6**, 71–76.
- Potter, B.E. (2012): Atmospheric interactions with wildland fire behavior – I: Basic surface interactions, vertical profiles and synoptic structures. *International Journal of Wildland Fire*, **21**, 779-801.
- Povak, N.A., Hessburg, P.F. and Salter, R.B. (2018) Evidence for scale-dependent topographic controls on wildfire spread. *Ecosphere*, **9**(10): e02443.

- Seto, D., Strand, T.M., Clements, C.B., Thistle, H., and Mickler, R. (2014) Wind and plume thermodynamic structures during low-intensity subcanopy fires. *Agricultural and Forest Meteorology*, **198-199**, 53-61.
- Seto D., Clements, C.B., and Heilman, W.E. (2013) Turbulence spectra measured during fire front passage. *Agricultural and Forest Meteorology*, **169**, 195-210.
- Seto, D., and Clements, C.B. (2011) Fire whirl evolution observed during a valley wind-sea breeze reversal. *Journal of Combustion*, **2011**, 12pp <https://doi.org/10.1155/2011/569475>
- Sharples, J.J. (2009) An overview of mountain meteorological effects relevant to fire behaviour and bushfire risk. *International Journal of Wildland Fire* **18**, 737-754.
- Sharples, J.J., McRae, R.H.D., Wilkes, S.R. (2012) Wind–terrain effects on the propagation of wildfires in rugged terrain: Fire channelling. *International Journal of Wildland Fire*, **21**, 282-296.
- Skowronski, N.S. (2021) Multi-scale analysis of wildland fire combustion processes in open canopy forests using coupled iteratively informed laboratory-, field- and model-based approach. Final Technical Report, SERDP Project RC-2641. Available at <https://www.serdp-estcp.org/Program-Areas/Resource-Conservation-and-Resiliency/Air-Quality/Fire-Emissions/RC-2641>
- Skowronski, Nicholas S.; Charney, Joseph J; Clark, Kenneth L.; Gallagher, Michael R.; Hadden, Rory M.; Heilman, Warren E.; Hom, John L.; Kremens, Robert L.; Cole, Jason A.; Campbell-Lochrie, Zakary J.; Walker-Ravena, Carlos; Mueller, Eric V.; Everland, Alexis I.; Patterson, Matthew M. 2022a. Multi-scale analyses of wildland fire combustion

processes: Small-scale field experiments - infrared data. Fort Collins, CO: Forest Service Research Data Archive.

Skowronski, Nicholas S.; Charney, Joseph J; Clark, Kenneth L.; Gallagher, Michael R.; Hadden, Rory M.; Heilman, Warren E.; Hom, John L.; Kremens, Robert L.; Cole, Jason A.; Campbell-Lochrie, Zakary J.; Walker-Ravena, Carlos; Mueller, Eric V.; Everland, Alexis I.; Patterson, Matthew M. 2022b. Multi-scale analyses of wildland fire combustion processes: Small-scale field experiments – terrestrial laser scans. Fort Collins, CO: Forest Service Research Data Archive.

Skowronski, N., and Hom, J.L. (2015): Observations of fire-induced turbulence regimes during low-intensity wildland fires in forested environments: Implications for smoke dispersion. *Atmospheric Sciences Letters*, **16**, 453–460.

Stull, R.B., (1988) An introduction to boundary layer meteorology (Vol. 13). Springer Science & Business Media.

Viegas, D.X., and Neto, L.P. (1991) Wall shear stress as a parameter to correlate the rate of spread of a wind-induced forest fire. *International Journal of Wildland Fire*, **1**, 177–188.

Werth, P.A., Potter, B.E., Clements, C.B., Finney, M.A., Goodrick, S.L., Alexander, M.E., Cruz, M.G., Forthofer, J.A., and McAllister, S.S. (2011) Synthesis of knowledge of extreme fire behavior: For fire managers. General Technical Report PNW-GTR-854, US Department of Agriculture, Forest Service, Pacific Northwest Research Station, Vol. I. Portland, OR, 144.

Table 1. Vertical turbulent transfer events and the associated quadrat designations.

Q	$\phi'w'$	ϕ'	w'	Heat flux	Momentum flux
1	>0	>0	>0	Ejection: upward flux of warmer air	Outward Interaction: upward flux of lower horizontal momentum air
2	<0	<0	>0	Inward Interaction: downward flux of warmer air	Sweep: downward flux of higher horizontal momentum air
3	>0	<0	<0	Sweep: downward flux of cooler air	Inward Interaction: downward flux of lower horizontal momentum air
4	<0	<0	>0	Outward Interaction: upward flux of cooler air	Ejection: upward flux of higher horizontal momentum air

LIST OF FIGURES

Figure 1. Sketch of the burn plot and the instruments deployed to the plot. The four capital letters (A, B, C and D) denote the four trusses and the four numbers (1, 2, 3, 4) refer to the 3D sonic anemometers on the trusses. Posts hanging on trusses B and C show the heights and location of thermocouples. The center post indicates the position of the infrared camera. The boxes next to the sonic anemometers indicate the radiometer/spectral camera pairs. The rectangular box on the ground indicates fuel cells for fuel loading estimation. The symbol near B2 indicates the TACO for emission data collection

Figure 2. Surface meteorological condition on May 20, 2019, the day of the experimental burn, observed by the weather station approximately 200 m northeast of the burn plot.

Figure 3. Infrared images taken at 10 m above the center of the burn plot showing fuel bed temperature before a), near b) and after c-f) ignition.

Figure 4. Time series of 10-Hz observations of temperature (T), horizontal wind speed (S) and vertical wind component (w) observed by the 16 sonic anemometers.

Figure 5. The number of sonic anemometers that recorded temperatures at or above a given threshold value (left) and the length of period over which the threshold was reached or exceeded (right). The symbol σ denotes pre-burn period temperature standard deviation.

Figure 6. Time series of 10 Hz streamwise (u , blue) and cross-stream (v , green) wind velocity components and temperature perturbations (T' , red) recorded by each sonic anemometer at 2.5 m above the ground. The vertical dashed black lines indicate the burn period determined by the first and last occurrence of $T' \geq 8\sigma$. Time is the minutes since the start of the pre-burn period.

Figure 7. Distributions of 10 Hz streamwise (u), cross-stream (v), and vertical (w) wind velocity components, and temperature perturbations (T') from all 16 sonic anemometers during pre-burn, burn and post-burn periods. The box represents the 25th and 75th percentile of the data, with data inside the whiskers representing 99.3% of the data. The orange line in the boxes is the median value, the green triangle is the mean, and the blue shading is the density of values of the data.

Figure 8. Time series of 1-minute averaged turbulent kinetic energy (TKE) (red) for each sonic anemometer and the three components of velocity variance, $u'^2/2$ (yellow), $v'^2/2$ (blue) and $w'^2/2$ (green), that make up the TKE . The vertical dashed black lines indicate the burn period determined by the first and last occurrence of $T' \geq 8\sigma$. Time is the minutes since the start of the pre-burn period.

Figure 9. Distributions of turbulent kinetic energy (TKE) and the three components of velocity variance ($u'^2/2$, $v'^2/2$ and $w'^2/2$) that make up the TKE from all 16 sonic anemometers during the pre-burn, burn and post-burn periods. The box represents the 25th and 75th percentile of the data, with data inside the whiskers representing 99.3% of the data. The orange line in the boxes is the median value, the green triangle is the mean, and the blue shading is the density of values of the data.

Figure 10. Time series of 1-minute averaged friction velocity squared (u_*^2 , pink pluses) and its two components, the streamwise kinematic momentum flux, $\overline{u'w'}$ (yellow circle) and the cross-stream kinematic momentum flux, $\overline{v'w'}$ (blue diamonds), for each of the 16 sonic anemometers. The vertical dashed black lines indicate the burn period determined by the first and last occurrence of $T' \geq 8\sigma$. Time is the minutes since the start of the pre-burn period.

Figure 11. Distributions of friction velocity squared (u_*^2) and its two components ($\overline{u'w'}$ and $\overline{v'w'}$) from all 16 sonic anemometers during the pre-burn, burn, and post-burn periods. The box represents the 25th and 75th percentile of the data, with data inside the whiskers representing 99.3% of the data. The orange line in the boxes is the median value, the green triangle is the mean, and the blue shading is the density of values of the data.

Figure 12. Time series of 1-minute averaged heat flux for each of the 16 sonic anemometers (left) and the distribution of heat fluxes from all 16 sonic anemometers during the pre-burn, burn, and post-burn periods (right). The box represents the 25th and 75th percentile of the data, with data inside the whiskers representing 99.3% of the data. The orange line in the boxes is the median value, the green triangle is the mean, and the blue shading is the density of values of the data.

Figure 13. Quadrant analysis of the instantaneous vertical kinematic turbulent heat fluxes showing the contributions to the total flux from (top row), and the percent of (bottom row) the four types of events: outward interaction (green), ejection (red), inward interaction (blue), and sweep (orange) for each of the 16 sonic anemometers during the pre-burn, burn, and post-burn periods. The black diamonds in the top row indicate the total heat flux values. The sonic anemometers are arranged from west to east roughly following the fire spread across the burn plot.

Figure 14. Quadrant analysis of the instantaneous vertical kinematic turbulent heat fluxes showing the contributions to the total flux from (top row), and the percent of (bottom row) the four types of events: outward interaction (green), ejection (red), inward interaction (blue), and sweep (orange) for all 16 sonic anemometers during the pre-burn, burn, and post-burn periods. The black diamonds in the top row indicate the total heat flux values. The sonic anemometers are arranged from west to east roughly following the fire spread across the burn plot.

Figure 15. Quadrant analysis of the instantaneous vertical kinematic turbulent fluxes of horizontal momentum showing the contributions to the total flux from (top row), and the percent of (bottom row) the four types of events: outward interaction (red), sweep (green), inward interaction (orange), and ejection (blue) for each of the 16 sonic anemometers during the pre-burn, burn, and post-burn periods. The black diamonds in the top row indicate the total flux values. The sonic anemometers are arranged from west to east roughly following the fire spread across the burn plot.

Figure 16. Quadrant analysis of the instantaneous vertical kinematic turbulent fluxes of horizontal momentum showing the contributions to the total flux from (top row), and the percent of (bottom row) the four types of events: outward interaction (red), sweep (green), inward

interaction (orange), and ejection (blue) for all 16 sonic anemometers during the pre-burn, burn, and post-burn periods. The black diamonds in the top row indicate the total flux values. The sonic anemometers are arranged from west to east roughly following the fire spread across the burn plot.

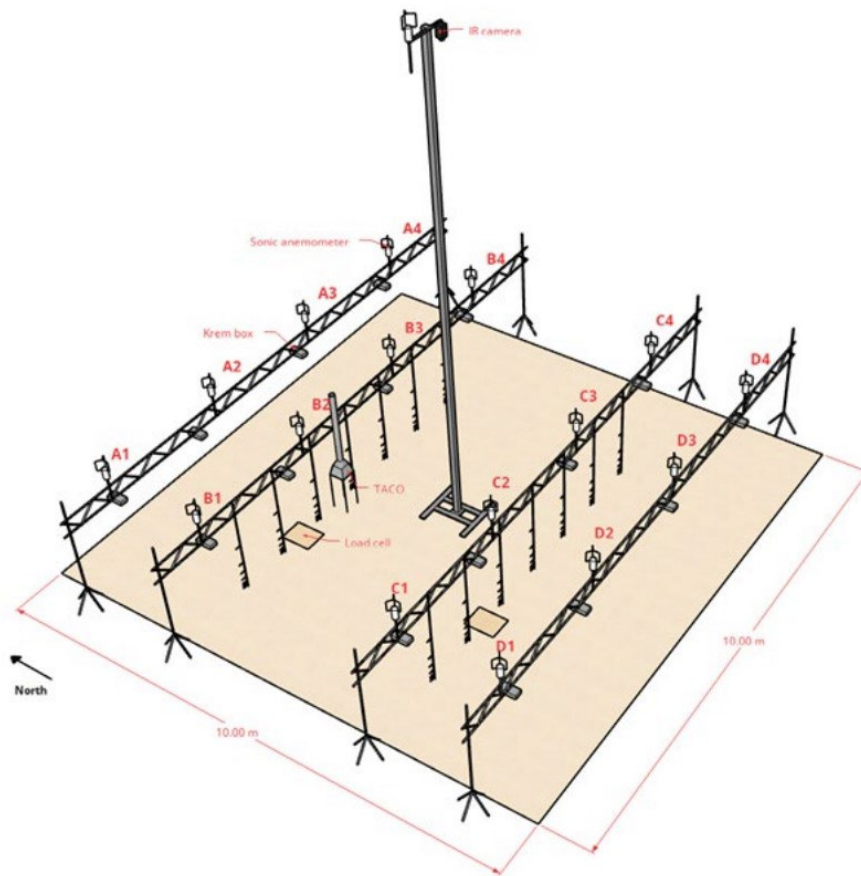


Figure 1. Sketch of the burn plot and the instruments deployed to the plot. The four capital letters (A, B, C and D) denote the four trusses and the four numbers (1, 2, 3, 4) refer to the 3D sonic anemometers on the trusses. Posts hanging on trusses B and C show the heights and location of thermocouples. The center post indicates the position of the infrared camera. The boxes next to the sonic anemometers indicate the radiometer/spectral camera pairs. The rectangular box on the ground indicates fuel cells for fuel loading estimation. The symbol near B2 indicates the TACO for emission data collection

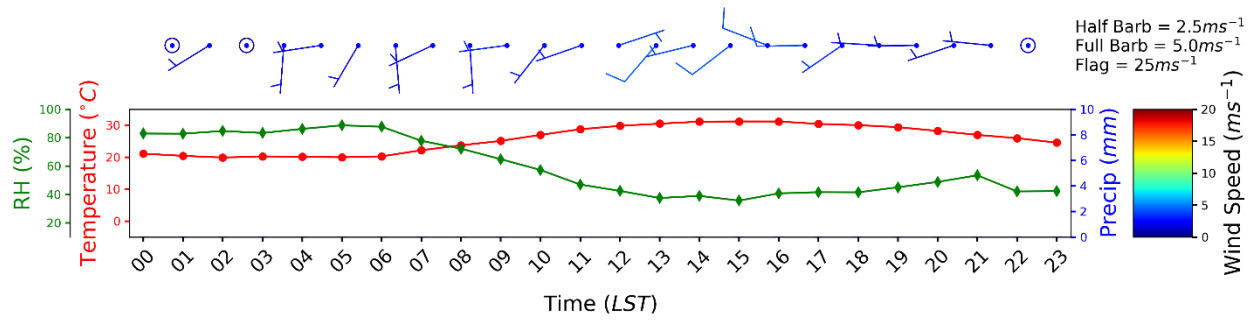


Figure 2. Surface meteorological condition on May 20, 2019, the day of the experimental burn, observed by the weather station approximately 200 m northeast of the burn plot.

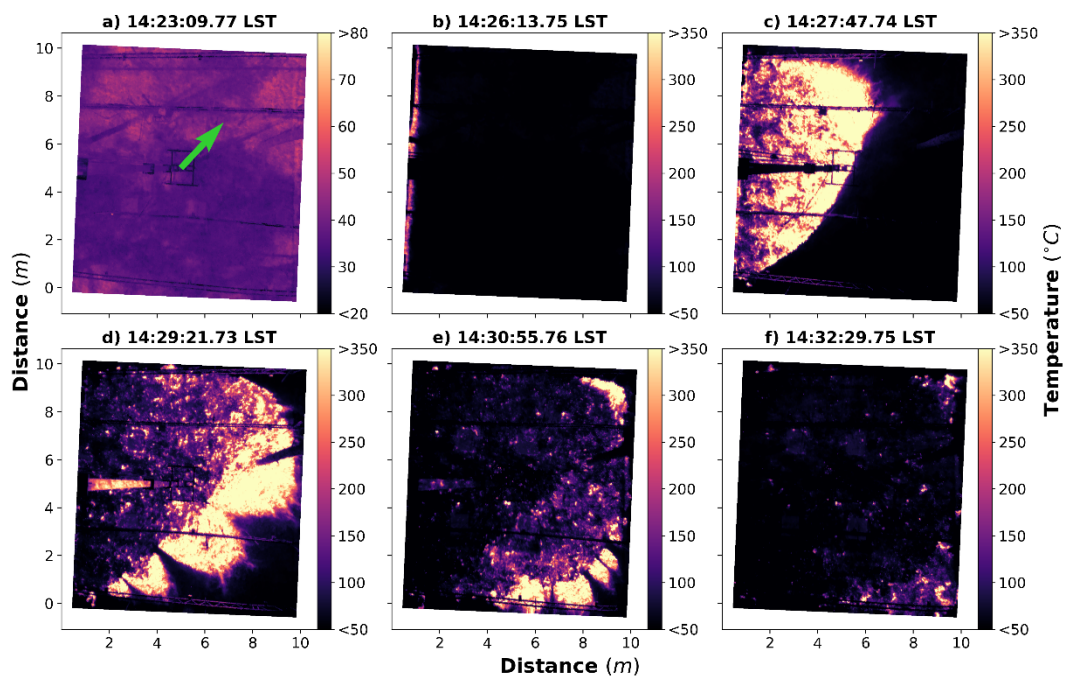


Figure 3. Infrared images taken at 10 m above the center of the burn plot showing fuel bed temperature before a), near b) and after c-f) ignition. The green arrow indicates the direction of background wind.

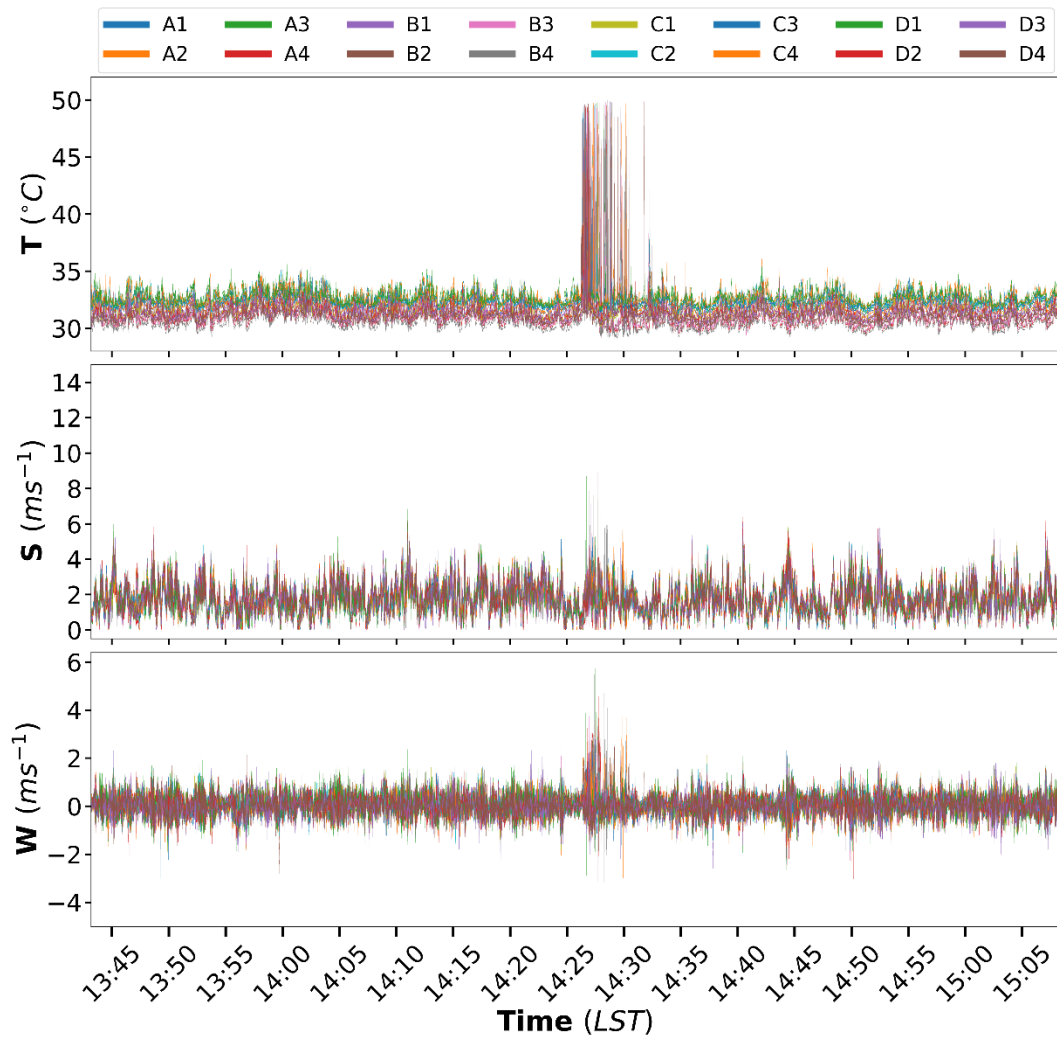


Figure 4. Time series of 10-Hz observations of temperature (T), horizontal wind speed (S) and vertical wind component (w) observed by the 16 sonic anemometers.

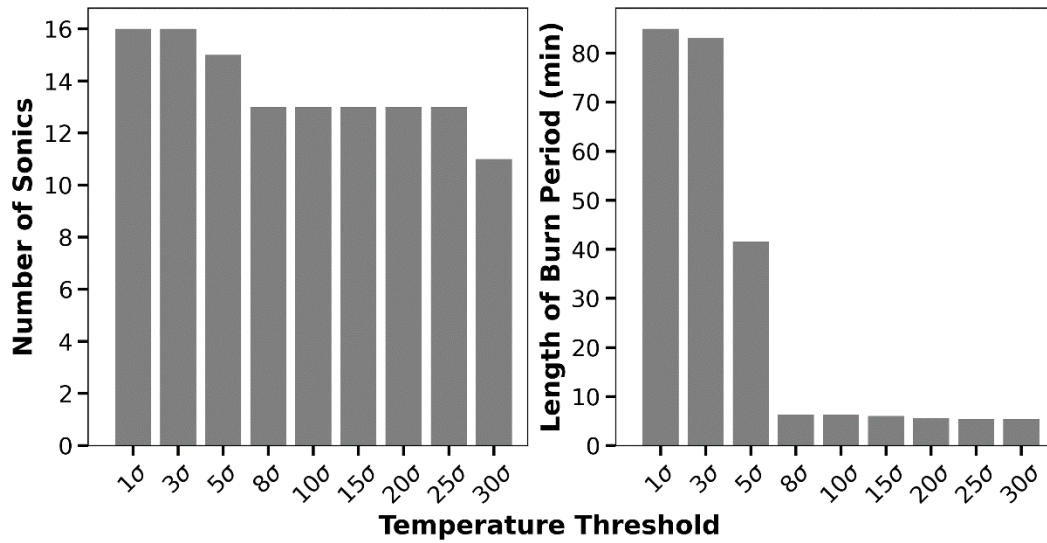


Figure 5. The number of sonic anemometers that recorded temperatures at or above a given threshold value (left) and the length of period over which the threshold was reached or exceeded (right). The symbol σ denotes pre-burn period temperature standard deviation.

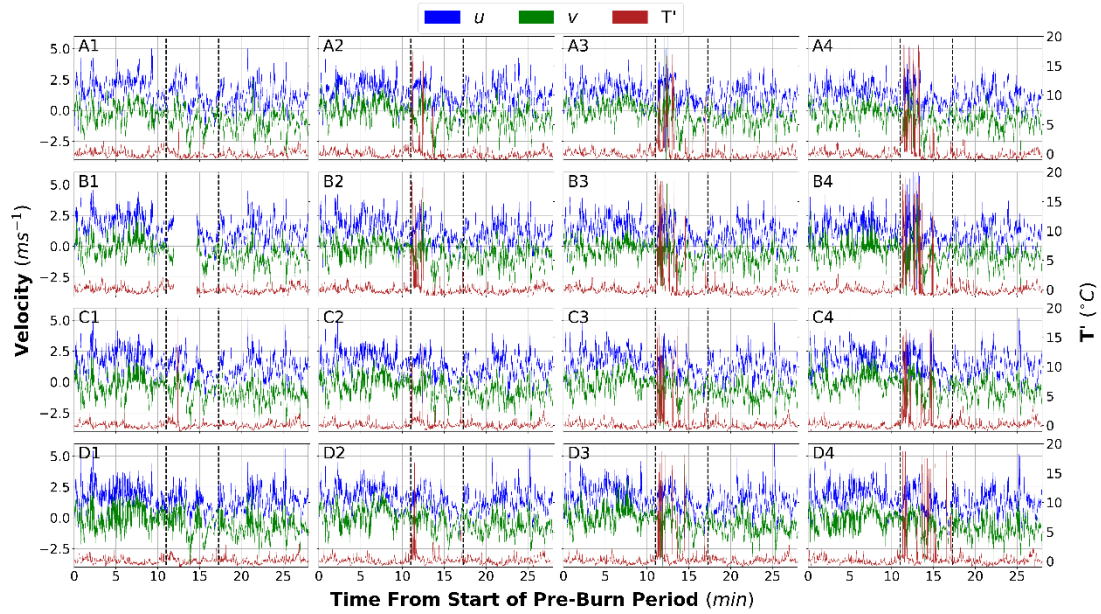


Figure 6. Time series of 10 Hz streamwise (u , blue) and cross-stream (v , green) wind velocity components and temperature perturbations (T' , red) recorded by each sonic anemometer at 2.5 m above the ground. The vertical dashed black lines indicate the burn period determined by the first and last occurrence of $T' \geq 8\sigma$. Time is the minutes since the start of the pre-burn period.

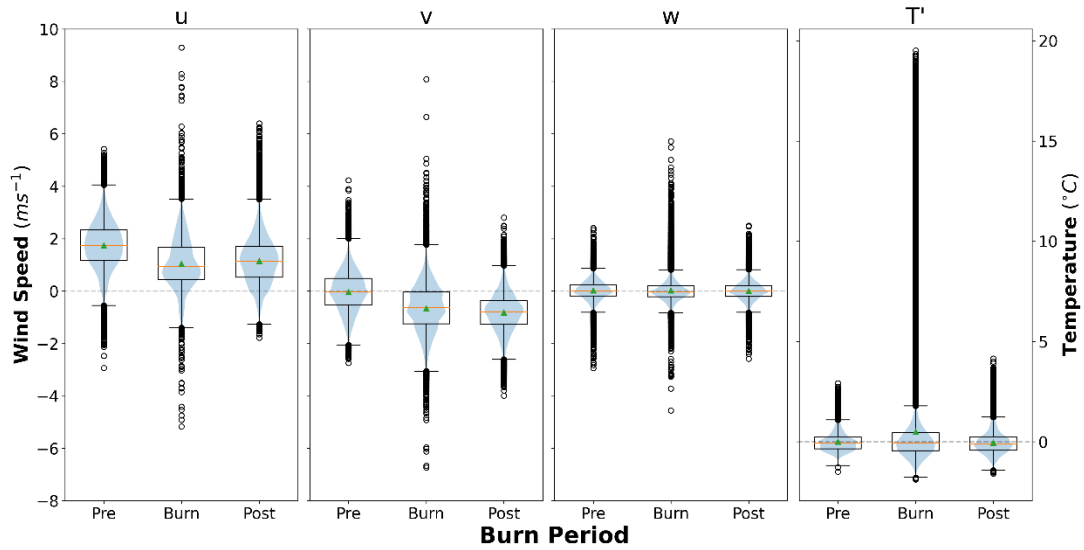


Figure 7. Distributions of 10 Hz streamwise (u), cross-stream (v), and vertical (w) wind velocity components, and temperature perturbations (T') from all 16 sonic anemometers during pre-burn, burn and post-burn periods. The box represents the 25th and 75th percentile of the data, with data inside the whiskers representing 99.3% of the data. The orange line in the boxes is the median value, the green triangle is the mean, and the blue shading is the density of values of the data.

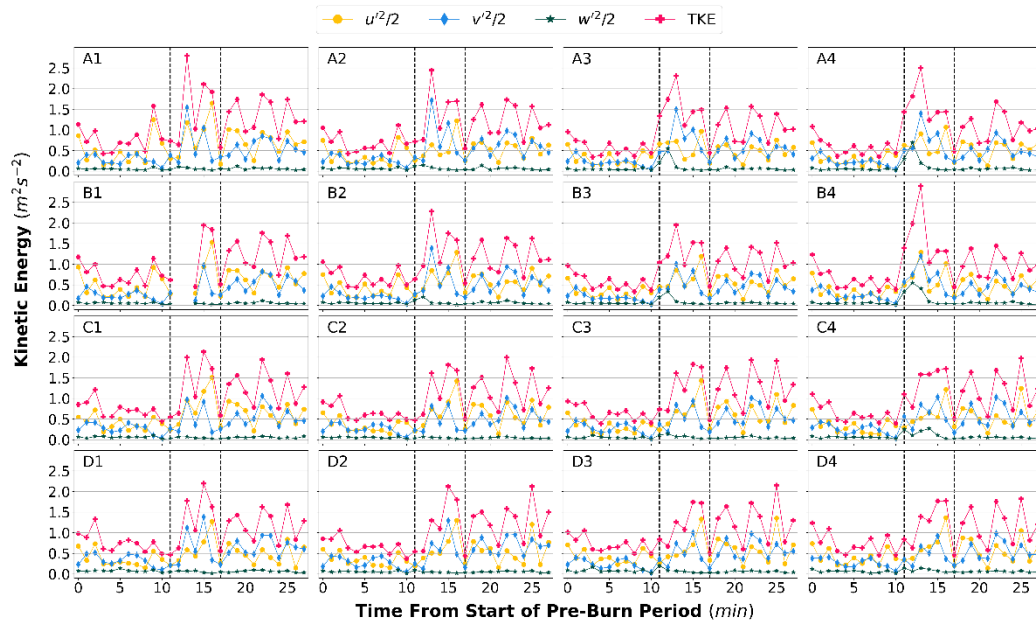


Figure 8. Time series of 1-minute averaged turbulent kinetic energy (TKE) (red) for each sonic anemometer and the three components of velocity variance, $u'^2/2$ (yellow), $v'^2/2$ (blue) and $w'^2/2$ (green), that make up the TKE . The vertical dashed black lines indicate the burn period determined by the first and last occurrence of $T' \geq \delta\sigma$. Time is the minutes since the start of the pre-burn period

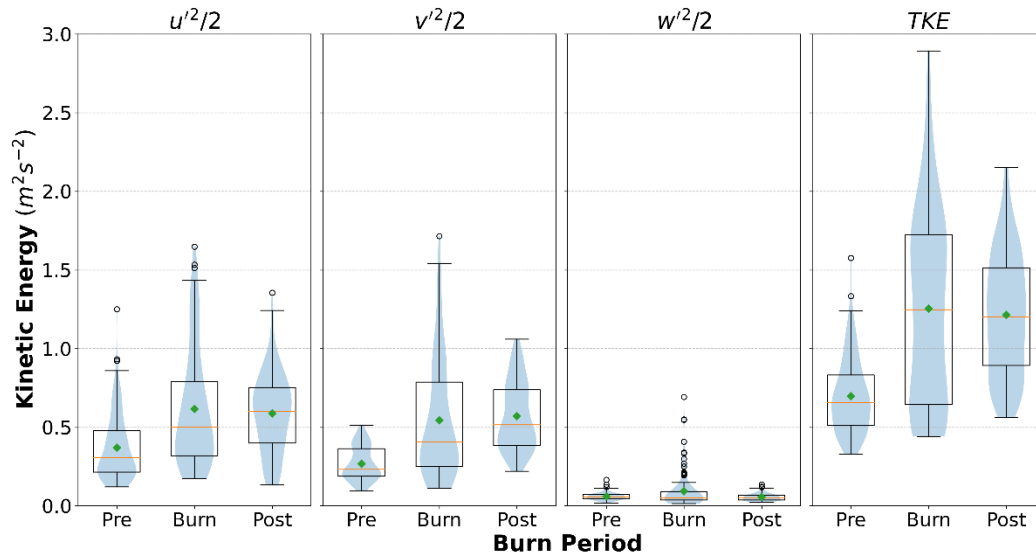


Figure 9. Distributions of turbulent kinetic energy (TKE) and the three components of velocity variance ($u'^2/2$, $v'^2/2$ and $w'^2/2$) that make up the TKE from all 16 sonic anemometers during the pre-burn, burn and post-burn periods. The box represents the 25th and 75th percentile of the data, with data inside the whiskers representing 99.3% of the data. The orange line in the boxes is the median value, the green triangle is the mean, and the blue shading is the density of values of the data.

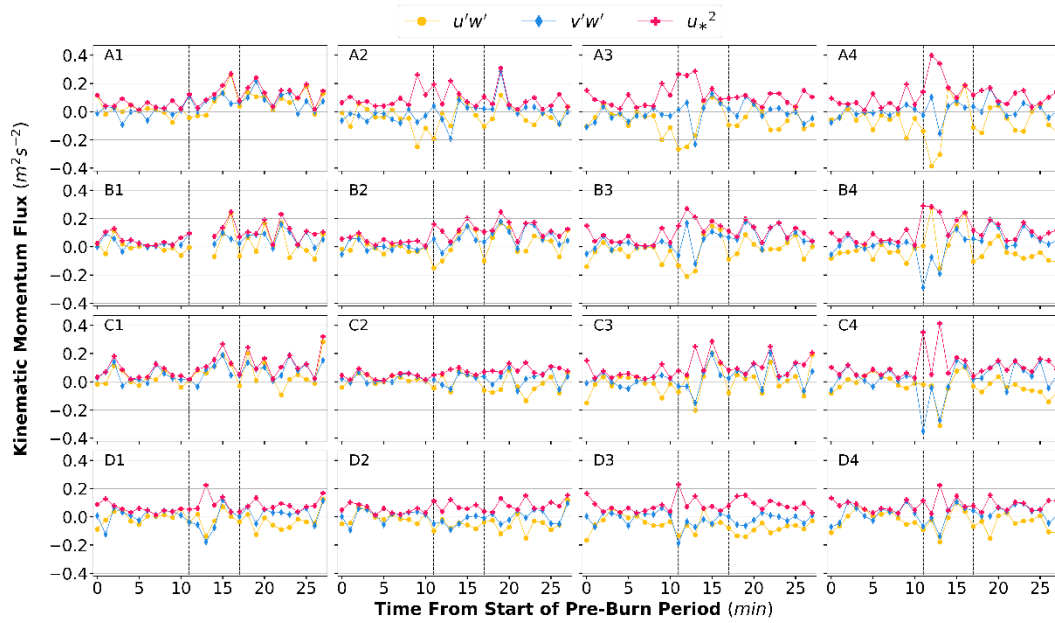


Figure 10. Time series of 1-minute averaged friction velocity squared (u_*^2 , pink pluses) and its two components, the streamwise kinematic momentum flux, $\overline{u'w'}$ (yellow circle) and the cross-stream kinematic momentum flux, $\overline{v'w'}$ (blue diamonds), for each of the 16 sonic anemometers. The vertical dashed black lines indicate the burn period determined by the first and last occurrence of $T' \geq 8\sigma$. Time is the minutes since the start of the pre-burn period.

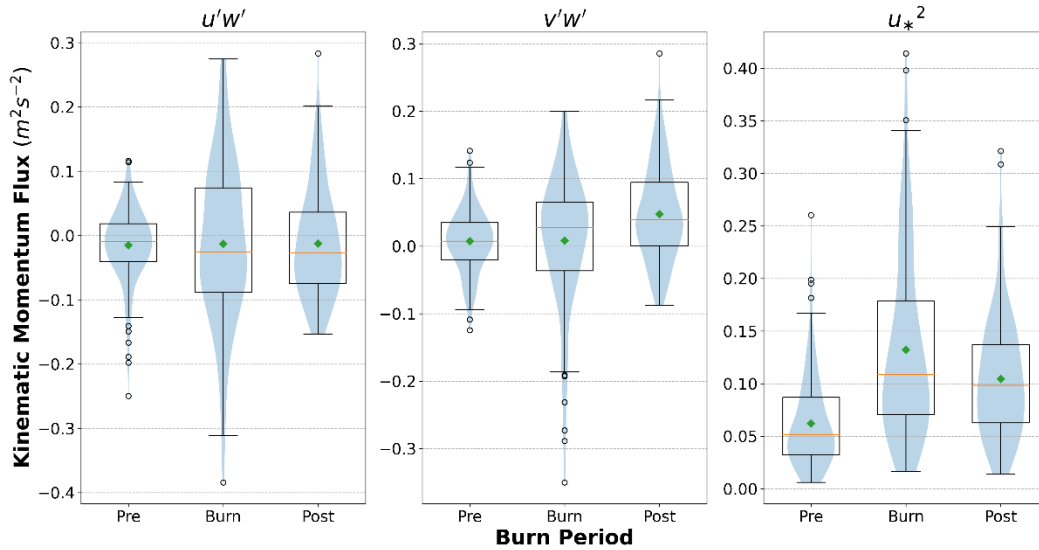


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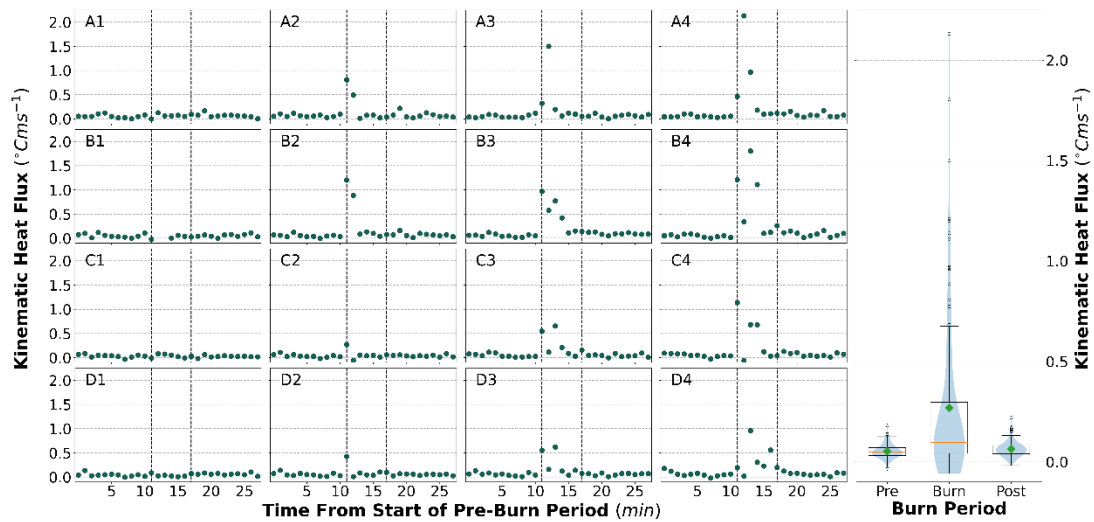


Figure 12. Time series of 1-minute averaged heat flux for each of the 16 sonic anemometers (left) and the distribution of heat fluxes from all 16 sonic anemometers during the pre-burn, burn, and post-burn periods (right). The box represents the 25th and 75th percentile of the data, with data inside the whiskers representing 99.3% of the data. The orange line in the boxes is the median value, the green triangle is the mean, and the blue shading is the density of values of the data.

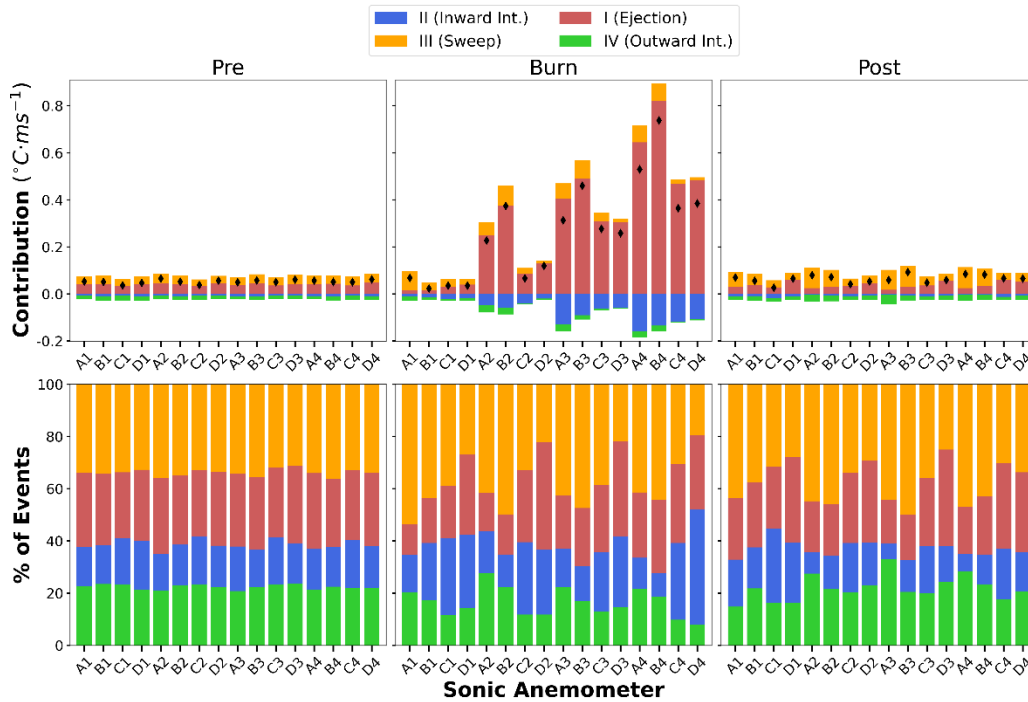


Figure 13. Quadrant analysis of the instantaneous vertical kinematic turbulent heat fluxes showing the contributions to the total flux from (top row), and the percent of (bottom row) the four types of events: outward interaction (green), ejection (red), inward interaction (blue), and sweep (orange) for each of the 16 sonic anemometers during the pre-burn, burn, and post-burn periods. The black diamonds in the top row indicate the total heat flux values. The sonic anemometers are arranged from west to east roughly following the fire spread across the burn plot.

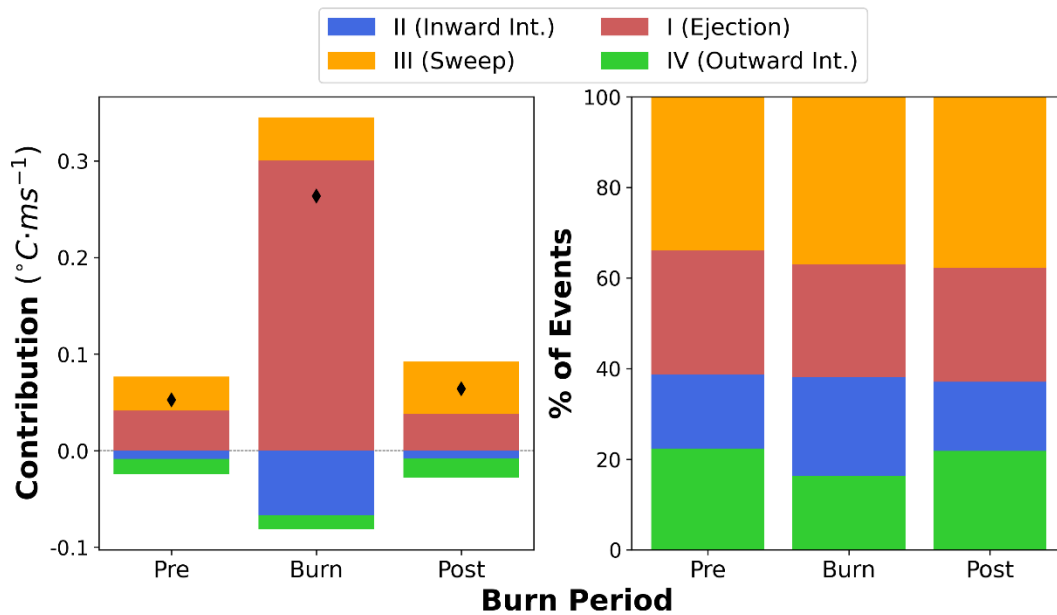


Figure 14. Quadrant analysis of the instantaneous vertical kinematic turbulent heat fluxes showing the contributions to the total flux from (top row), and the percent of (bottom row) the four types of events: outward interaction (green), ejection (red), inward interaction (blue), and sweep (orange) for all 16 sonic anemometers during the pre-burn, burn, and post-burn periods. The black diamonds in the top row indicate the total heat flux values. The sonic anemometers are arranged from west to east roughly following the fire spread across the burn plot.

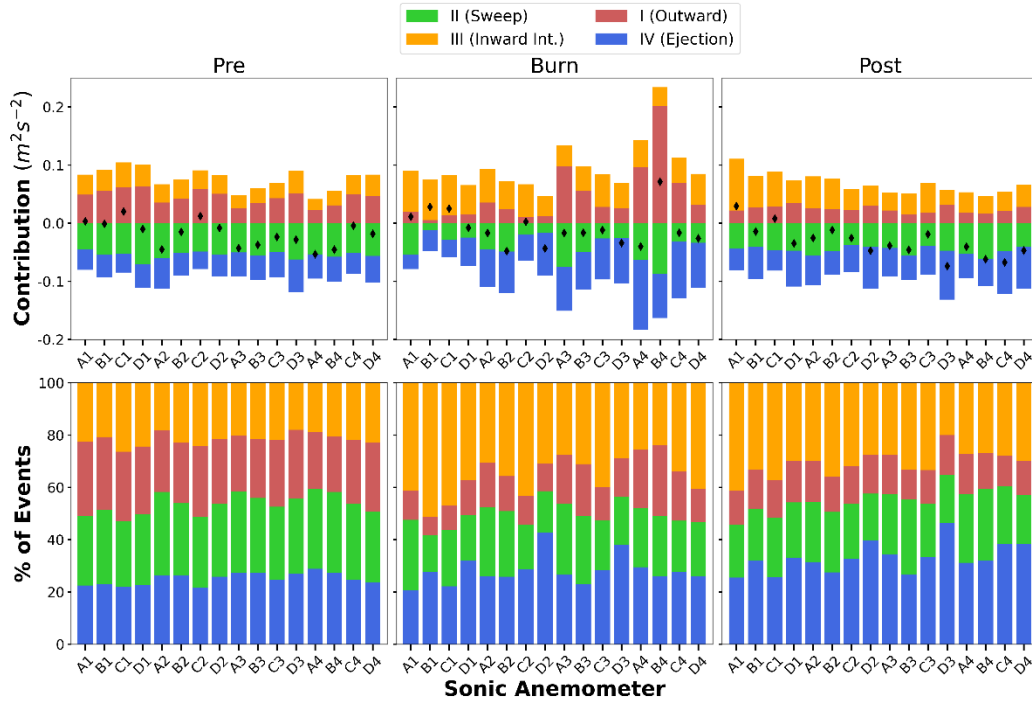


Figure 15. Quadrant analysis of the instantaneous vertical kinematic turbulent fluxes of horizontal momentum showing the contributions to the total flux from (top row), and the percent of (bottom row) the four types of events: outward interaction (red), sweep (green), inward interaction (orange), and ejection (blue) for each of the 16 sonic anemometers during the pre-burn, burn, and post-burn periods. The black diamonds in the top row indicate the total flux values. The sonic anemometers are arranged from west to east roughly following the fire spread across the burn plot.

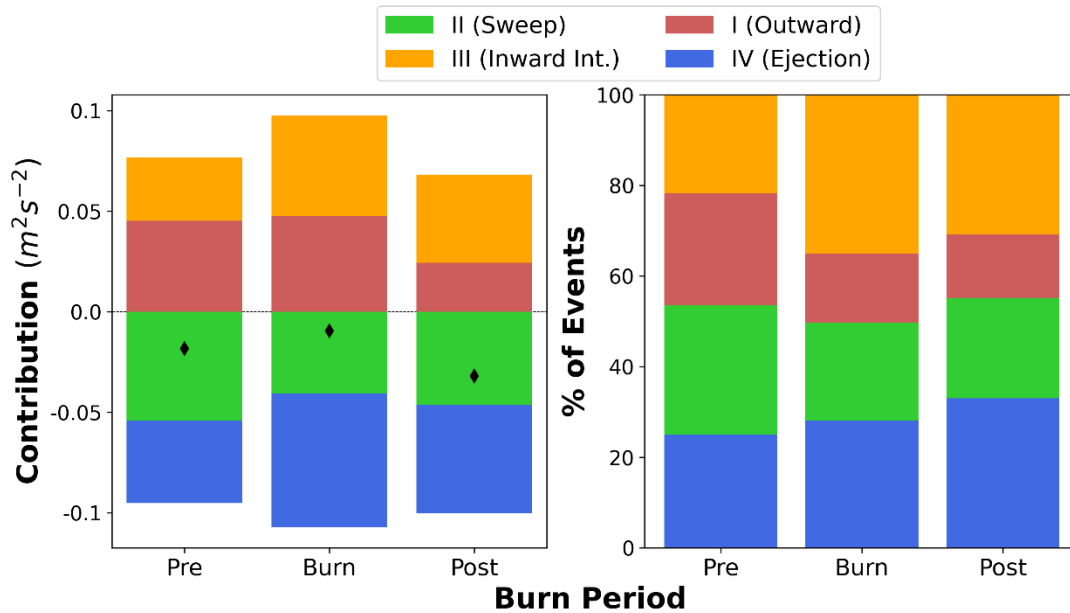


Figure 16. Quadrant analysis of the instantaneous vertical kinematic turbulent fluxes of horizontal momentum showing the contributions to the total flux from (top row), and the percent of (bottom row) the four types of events: outward interaction (red), sweep (green), inward interaction (orange), and ejection (blue) for all 16 sonic anemometers during the pre-burn, burn, and post-burn periods. The black diamonds in the top row indicate the total flux values. The sonic anemometers are arranged from west to east roughly following the fire spread across the burn plot.