Point by point response

1. Lack of adequate scientific justification for the work

We have placed our study in the broader perspective of PSC research. We have broadened the <u>Introduction</u> to include there the fact that the subsidence of large NAT particles is considered one of the main causes of denitrification of the polar winter stratosphere. Their settling time influences this process, which is in turn dependent on NAT particle shape and size, both of which determine their settling speed and lifetime, hence their denitrification efficacy. Of particular interest is the shape, given that non-spherical particles may fall significantly slower than volume equivalent spheres. A positive result of our study, which maybe we did not adequately underline, is that it strongly suggests avoiding ARs too close to 1.0, preferring ARs below 0.55 or above 1.5. This result has been underlined and referred to previous estimates on the asphericity of NAT particles.

We have reported such considerations in the Introduction:

The aims of this effort are both to verify the ability of the T-matrix approach to reproduce the observations from lidar/backscattersonde, once the PSDs are supposed known, and to provide a contribution to the estimation of the shape and size limits of the NAT PSC particles. The question of the shape of NAT particles is in fact far from being clarified, and has important implications for the denitrification mechanisms of the polar stratosphere, an important step in the process that lead to the destruction of stratospheric ozone. In fact, large PSC NAT particles settling down are considered one of the main causes of denitrification of the polar winter stratosphere (Di Liberto et al., 2015). Their settling time influences this process, and it is in turn dependent on NAT particles shape and size, both of which determining their settling speed and lifetime, hence their denitrification efficacy. Woiwode et al. (2014) assumed significantly non-spherical NAT particles to simulate the NAT settling speed leading to a the vertical redistribution of HNO₃ observed between two companion flights during the RECONCILE airborne field campaign in the Arctic (von Hobe et al., 2013). Woiwode et al. (2016, 2019) have also suggested that NAT particles may be highly aspherical based on the infrared spectrometer MIPAS-STR limb observations exhibiting a spectral signature around 820 cm⁻¹ and an overall spectral pattern compatible with large highly aspherical NAT particles. T-Matrix calculations assuming randomly oriented highly aspherical NAT particles (aspect ratios 0.1 or 10 for elongated or disk-like spheroids, respectively) were able to reproduce the MIPAS-STR observations to a large degree. Molleker et al. (2014) hypothesized strongly aspherical NAT particles to reconcile the amount of the condensed HNO3 resulting from PSC cloud spectrometer measurements with the expected stratospheric values, and to provide consistency between particles settling velocities and growth times with back trajectories. Moreover, Grothe et al. (2006) observed highly aspherical NAT in laboratory experiments. This is in contrast with earlier studies that assumed an AR = 0.9 for the NAT spheroids to match microphysical model simulation with airborne (Carslaw et al., 1998) or satellite borne (Hoyle et al., 2013; Engel et al., 2013) lidar observations.

As quoted above concerning NAT asphericities, Molleker et al. (2014) and Woiwode et al. (2014; 2016; 2019) have suggested that NAT particles may be highly aspherical.

Similarly, our study concludes that the best agreement between measurements and optical modeling occurs for strongly aspherical NAT.

This is in contrast with earlier studies that assumed an AR = 0.9 for NAT spheroids (Carslaw et al., 1998, Hoyle et al., 2013, Engel et al. 2013).

Concerning the R_{Th} , the hypothesis of dividing the PSD into a liquid part and a solid part on the basis of size is a hypothesis supported by what we know about PSC particle formation and measurements (Deshler et al., 2003), leads to sensible results, and is in agreement with the depolarization-large particle correlation qualitatively presented in new Figure 1.

2. Simplicistic assumption of a common AR

We have explored the possibility of deriving the AR that best matches computation and experiment, on a case by case basis. This is the main upgrade of our work. In this new version, we have acknowledged that the presented β or δ measured-computed comparison with respect to (R_{th}, AR) can only constrain a range of R_{th}.

In our additional simulations, for each of the experimental (PSD, β , δ) datum we have seeked the AR and R_{th} that best matches the measured δ with the T-matrix calculations. Exploiting the already presented results, AR and R_{th} were allowed to vary respectively in the ranges (0.3-0.6)U(1.5-3) and 0.5-0.8 μ m.

Once these AR and R_{th} have been found, the same were used to calculate β and compare it with the experimental, on a case-by-case basis.

This new procedure leads to an improvement in the agreement between experimental and calculated δs , while it does not change appreciably the agreement between the βs .

We have described this new procedure in the Abstract:

The parameters R_{th} and AR of our model have been varied between 0.1 and 2 μ m and between 0.3 and 3, respectively, and the calculated backscattering coefficient and depolarization were compared with the observed ones.

The best agreement was found for R_{th} between 0.5 and 0.8 μm , and for AR less than 0.55 and greater than 1.5.

To further constrain the variability of AR within the identified intervals we have sought an agreement with the experimental data by varying AR on a case-by-case basis, and further optimizing the agreement by a proper choice of AR smaller than 0.55 and greater than 1.5, and R_{th} within the interval 0.5 and 0.8 μ m. The ARs identified in this way cluster around the values 0.5 and 2.5.

In paragraph 2.3 Variability with the threshold radius R_{th} and Aspect Ratio AR:

The result of this study allows to identify only the best R_{th} , resulting around 0.5-0.8 μ m, while the ARs compatible with the measurements are all those between 0.3-0.55 and 1.5-3. To further constrain AR we have kept R_{th} at a fixed value, chosen between 0.5 and 0.8 μ m and changed this value with a 0.1 μ m step. For each of these fixed R_{th} , and separately for each PSD, we identified in the intervals (0.3-0.55), (1.5,3) the value of AR which best matched the observed δ with its computed value. Finally, for each PSD we selected the pair R_{th} and AR which provided the best match. Once the ARs and R_{th} have been selected by forcing the agreement between the δ_A , the same ones have been used for the calculation of the β_A .

Figures 4 and 5 have been upgraded with the result of this new approach:



Figure 4. Scatterplot of computed vs measured particle backscattering coefficients β_A . The ARs used for the computations have been selected, case by case, to produce the best agreement between the β computed and measured, and are here represented in color coding. Only ARs in the intervals between 0.3 and 0.55, and between 1.5 and 3, have been considered. R_{th} was also selected within the interval 0.5-0.8 μ m to provide the best match. We report data points with BR greater than 1.2, $\beta_{cross A}$ greater than 5 10⁻⁶km⁻¹sr⁻¹ and temperature at the observation below 200 K.



Figure 5. Scatterplot of computed vs measured particle depolarization δ_A . The ARs used for the computations are those that provided the best match between the δ_A computed and measured, and are here represented in color coding. Only ARs in the intervals between 0.3 and 0.55, and between 1.5 and 3, have been considered. R_{th} was also selected within the interval 0.5-0.8 μ m to provide the best match. We report data points with BR greater than 1.2, β_{cross} A greater than 5 10⁻⁶km⁻¹sr⁻¹ and temperature at the observation below 200 K.

We have also discussed the physical relevance of our results, in the par. 5. Discussion:

The identification of the best R_{th} in the range 0.5 -0.8 µm supports what we already know from the theoretical understanding of NAT particle formation in PSC and from measurements (Deshler et al., 2003b). Concerning particle shape, in our model all solid particles in a single PSD share the same AR, but different PSDs can have different ARs. This approach could suggest that the choice of the AR which, case by case, optimizes the agreement between calculations and measurements, may be the result of chance rather than physics. There are two facts that counter this criticism.

First, it appears that the selected ARs may be related to the shape of the PSD. Figure 6 shows the 2D-histogram by occurrence of ARs and of $N(r > 0.7 \mu m)/N_{tot}$, the ratio between particles with radius greater than 0.7 μ m and total particles, which is a parameter related to the PSD shape. In Figure 6 the AR are not distributed randomly. Conversely, there is a tendency for the AR to grow as the percentage of large particles increases. In fact AR values tend to peak around 0.5 in the lower $N(r > 0.7 \mu m)=N_{tot}$ range, while tend to cluster around 2.5 when $N(r > 0.7 \mu m)=N_{tot}$ is higher. The shape of the PSD mirrors particle formation conditions and history, is linked to the presence of solid particles, as already highlighted in the discussion of Figure 1, and is likely linked to the average particle shape as well. Second, if we consider the sequences of measurements acquired in individual balloon flights, the corresponding sequences of selected ARs do not evolve randomly but, conversely, are autocorrelated. An example of this behavior is shown in Figure 7, where the time series of β and δ are reported respectively with red and blue dots. The ARs that provide the best agreement between experiment and simulation are shown with black dots. It can be seen that temporally contiguous observations often result in the selection of the same AR. Contiguous observations of PSD are likely to have similar characteristics in terms of microphysics, and this seems to be correctly reflected in the constancy of AR. We are therefore confident that our method produces results with a physics-based content.

We report here as well, for completeness, new Figures 6 and 7:



Figure 6. 2D-histogram of occurrence of ARs and of $N(r < 0.7 \mu m)=N_{tot}$, the ratio between particles with radius greater than 0.7 μm and total particles. Only ARs in the intervals between 0.3 and 0.55, and between 1.5 and 3, have been considered.



Figure 7: Sequences of b (red dots) and d (blue dots) measured on a balloon flight on December 9th 2001, from Kiruna, Sweden. Each data point represents an average over 60s. Black dots represents the ARs providing the best match between the d and those computed from concomitant measurements of PSD.

3. Lack of proper discussion of model results

The novel approach we have pursued in the revision of our manuscript has led to new results. This has led us to a major revision of paragraph 4. Discussion. We report here the full text, which is now discussing the new results, addressing more closely the reviewers remarks.

The identification of the best R_{th} in the range 0.5 -0.8 μ m supports what we already know from the theoretical understanding of NAT particle formation in PSC and from measurements (Deshler et al., 2003b). Concerning particle shape, in our model all solid particles in a single PSD share the same AR, but different PSDs can have different ARs. This approach could suggest that the choice of the AR which, case by case, optimizes the agreement between calculations and measurements, may be the result of chance rather than physics. There are two facts that counter this criticism.

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In general, our model leads to good correlations between measured and modeled β s. For the δ s the measurements are well reproduced by the calculations in many instances, as is the case for many of the selected ARs in the range 0.3-0.55. However, there are other cases in which the agreement is worse (when the best ARs have been selected in the range 1.5-3), or does not occur at all, as in the cases of observed depolarizations greater than 30%. In these latter cases, the impossibility of reproducing the observed values even under the hypothesis of a completely solid particles implies that, for those PSDs, our model is not able to produce the observed depolarizations. In these particular cases in which the model performs particularly badly, there may be problems of inhomogeneities of the cloud. These cases come from Antarctic observations, for which the microphysical observations from the balloon and the optical ones from ground-based lidar are separated geometrically, so that the two instruments sample air masses separated by several tens of kilometres, and it may be the case that some clouds were not homogeneous on such spatial scales.

4. Lack of a size-dependent AR

For what concerns the possibility of implementing a size-dependent AR approach, we aknowledge to have used a simplified assumption, given our limited knowledge of NAT crystallization. There has been speculation of anisotropic growth, favoring large asphericities when particles grow to large sizes, due to less depleted vapor mixing ratios close to the extremity of the crystals (Grothe et al., 2006). Also, various particle shapes and habits might coexist due to different nucleation and growth histories. Thus, it is certainly possible that a choice of particle size-dependent AR exists, which may improve the agreement with backscattering/depolarization measurements. However, such an agreement, if found, would be easily open to further criticism since there is no basis to make complicated assumptions about the size dependent AR. By suitably adjusting the AR, we would regard such a result as a selection of the most desirable outcome. For these reason, in looking for the AR intervals that best match the backscattering measurements, we prefer to consider only an average AR, as the most conservative assumption. Nevertheless, we have commented on such option in the paragraph <u>4 Discussion</u>: write:

Different shapes produce different polarization, according to T-Matrix. This has also been proven experimentally since the early work of Sassen and Hsueh (1998) and Freudenthaler et al. (1996) that showed how lidar depolarization ratios in persisting contrails ranged from 10% to 70%, depending on the stage of their growth and on temperature. In the T-matrix theory, for fixed AR, the depolarization depends on the particle size and maximizes for particular sizes. There is certainly a way to assume a particle size-dependent AR in our PSDs so as to reconcile the computations with the observed values. However, such an approach

would have little physical basis and could only be justified to maximize the agreement of calculations. Therefore, we have not explored this possibility further, although it is possible that our simplified hypothesis of a common AR for every particle may be the cause of the bad agreement between data and calculations in some case.

5. Suggestion to further explore the model-data mismatch in view of PSD characteristics

Concerning the shape of the PSDs which is suggested to display, our study is based on 473 data points (i.e. 473 triplets of PSD, backscattering coefficient, and depolarization). A three panel plot showing: i. the time series of PSD (color plot), ii. the corresponding backscatter coefficient (line plot); iii. the corresponding depolarization (line plot) could be produced, but given the range of the observations such representation may be difficult to interpret. However, in the revision of the manuscript, when discussing the AR distribution in new Figure 6, we have correlated it with a parameter characterizing the shape of the PSD (see above).

6. Lack of proper error analysis

The treatment of uncertainty has been elaborated, and error bars have been used in the new figures 4 and 5. Errors have been discussed in <u>2.1 Dataset:</u>

Experimental errors in the particle Backscatter Ratio (R-1) are estimated to be 5%, but not less than 0.05 in absolute value, while the error in volume depolarization is about 10%–15%. Additional uncertainty comes from the determination of pressure and temperature from radiosoundings, needed to compute β_A and δ_A (Adriani et al., 2004). [...]

Particle size histograms are fitted to unimodal or bimodal lognormal size distributions, which are the representation of size distribution used in this work. The uncertainties on the determination of the parameters of the mono/bimodal lognormals were determined by Deshler et al. (2003b) with Monte Carlo simulations. These were 20% for distribution width, 30% for median radii and 10% for modal particle concentrations.

And in <u>3. Results</u>:

The uncertainties associated with the measured β_A and δ_A derive from the error analysis for the single lidar data, which can be found in Adriani et al. (2004) or from the standard deviation for the averaged data, depending on which is greater. The uncertainties on the calculated β_A and δ_A , are 40% as determined by Deshler et al. (2003a) for any moment of a PSD derived from the OPC measurements. Deshler et al. determined this through a Monte Carlo simulation which used the uncertainties of the OPC size and concentration measurements to quantify the uncertainties in the PSD parameters and their subsequent moments.

5. Lack of quantitative estimation of goodness of fit

We have now provided the Pearson correlation coefficient (resulting to be 0.56) for the goodness of the 1:1 fit for the β comparison and added 1:1 lines to the data in Figs. 4 and 5. We did not perform goodness-of-fit tests for the comparison of δ s. In this case it is clear that there is a set of well-aligned points along the 1:1 line, and sets of points that deviate from it in a non-

random way. We have discussed the different characteristics of these sets in the <u>3 Result</u> paragraph, totally rewritten, which we report here in its entirety:

3 Results

Figure 4 reports the scatterplot of measured vs computed β_{A_r} colour coded in terms of AR. The figure represents the analogue of Figure 4 in Snels et al. (2021), where in the present case we have used a larger dataset, including now four Arctic balloon flights, and used T-Matrix instead of a factor 0.5 reduction in the Mie backscattering. Figure 5 reports the scatterplot of measured vs computed δ_A similarly color coded in terms of AR. The uncertainties associated with the measured β_A and δ_A derive from the error analysis for the single lidar data, which can be found in Adriani et al. (2004) or from the standard deviation for the averaged data, depending on which is greater. The uncertainties associated with the measured β_A and δ_A derive from the error analysis for the single lidar data, which can be found in Adriani et al. (2004) or from the standard deviation for the averaged data, depending on which is greater. The uncertainties associated with the measured β_A and δ_A derive from the error analysis for the single lidar data, which can be found in Adriani et al. (2004) or from the standard deviation for the averaged data, depending on which is greater. The uncertainties on the calculated β_A and δ_A , are 40% as determined by Deshler et al. (2003a) for any moment of a PSD derived from the OPC measurements. Deshler et al. determined this through a Monte Carlo simulation which used the uncertainties of the OPC size and concentration measurements to quantify the uncertainties in the PSD parameters and their subsequent moments.

Despite the dispersion in Figure 4 the points cluster around the straight line $\beta_{calc}=\beta_{meas}$, indicating the agreement between computation and measurements can be considered fine for β_A with the exception of β values below 4 10^{-5} km⁻¹sr⁻¹ where the β_{calc} underestimate the measurements. Such underestimation seems to be of the order of 10^{-5} km⁻¹sr⁻¹, a magnitude compatible with possible inaccuracies in the calibration of the lidar data. The Pearson correlation coefficient for the entire dataset is 0.56, and increases if the lower values of β are neglected.

The δ_A scatterplot shows the presence of a good number of points that align along the $\delta_{calc} = \delta_{meas}$ correlation line, with AR selected mainly around the value 0.5. However, for depolarization values greater than 30% there is no AR that will reproduce the measurements. These points correspond to those presented in Figure 1, with low values of BR and high values of the concentration ratio of large to total particles. They mainly come from three single observational periods of about one minute each, characterized by air temperatures between 184-188 K. Given the magnitude of the depolarization, it is possible that those observations are not referable to clouds in mixed phase, but rather to clouds of predominantly solid particles. For that particular set of points, we also explored the possibility that all particles were solid, but even under this assumption the comparison with the experimental data did not improve appreciably.

In Figure 5 for depolarizations lower than 15%, the points which deviate, by excess or defect, from the 1:1 straight line have predominantly AR greater than 1.5. So it seems that selected ARs greater than 1.5 generally produce a worse correlation. From Figure 4 we observe that AR values in the range (0.3-0.55) tend to be associated with medium-low β values, while AR values in the range (1.5-3) are mainly associated with medium-high β .

To conclude, the choice of R_{th} in a range between 0.5 and 0.8 μ m leads to a reasonably good agreement between the β 's, but there seems to be a discrepancy between the calculated value and the measurements in their lower range of variability.

From Figure 4 such mismatch, which makes the measurements larger than the calculations, seems to be of the order of 10^{-5} km⁻¹sr⁻¹. The selection of the AR that produces the best agreement with the observed δ 's leads to three results: i. The ARs in the range 0.3-0.55 tend to be selected in correspondence with medium-low β 's, the ARs in the range 1.5-3 in

correspondence with medium-high β 's. ii. ARs in the 0.3-0-5 range reproduce the measurements well, except for some observations where the depolarizations are greater than 30%; iii. the ARs in the 1.5-3 range reproduce the measurements less well; iv. There is no AR that will allow the calculations to reproduce the measurements for depolarizations greater than 30%.

6. Lack of sensitivity analysis

The sensitivity of the method can be estimated from the results already presented. Specifically, the sensitivity of our method can be obtained from the range of variability of the RMSEs according to the variability of AR and R_{Th}, as shown in figures 2 and 3. We note that those results are relative to the analysis of real particle size distributions (PSDs), from measurements. We don't see how a similar study, done on simulated PSDs, would provide additional information.

However, to pursue the reviewer's requests and understand how critically the uncertainties on the dataset impact the methodology, we have estimated the uncertainties to be attributed to both computed and measured β s and δ s (see new figures 4 and 5); We have added in the revised par. <u>3</u><u>Results</u>:

The uncertainties associated with the measured β_A and δ_A derive from the error analysis for the single lidar data, which can be found in Adriani et al. (2004) or from the standard deviation for the averaged data, depending on which is greater. The uncertainties on the calculated β_A and δ_A , are 40% as determined by Deshler et al. (2003a) for any moment of a PSD derived from the OPC measurements. Deshler et al. determined this through a Monte Carlo simulation which used the uncertainties of the OPC size and concentration measurements to quantify the uncertainties in the PSD parameters and their subsequent moments.

7. Suggestion to use an iterative methodology to vary the mixtures of shapes and sizes until most closely matched by the measurements.

In our study both the AR and the R_{Th} were independently varied in order to simulate different mixtures of shapes and sizes, thus the parameter ranges that best matches the measures were identified. Figures 2 and 3 in our manuscript provide what the reviewer is asking here. This analysis was effective in delimiting the variability of R_{th} , but not of AR, therefore in the revision of the manuscript we proceeded to look for the best match on a case-by-case basis, identifying for each PSD the best AR within the wide ranges of variability previously identified. The result is, for R_{th} within the limits 0.5-0.8 μ m, a distribution of AR within these intervals, which clusters around the values 0.5 and 2.5, approximately, as reported in the new figure 6 where the AR distribution is reported in terms of a parameter characterizing the shape of the PSD.



Figure 6. 2D-histogram of occurrence of ARs and of $N(r < 0.7 \mu m)=N_{tot}$, the ratio between particles with radius greater than 0.7 μm and total particles. Only ARs in the intervals between 0.3 and 0.55, and between 1.5 and 3, have been considered.

8. Suggestion to use the actual scattering measured by the OPC, rather than converting scattering to equivalent optical diameters and then computing scattering, as was done by Baumgardner and Clark (1998).

In Baumgardner and Clarke (1998) the authors infer the total single particle scattering coefficient from the forward scattering coefficient measured between 4 and 12 degrees. This inference is made by calculating, with the aid of Mie's theory, the relationship between the scattering, calculated in the above angle interval, and the total scattering. The inference is then that the total particle scattering can be inferred from the FSSP measured scattering, seemingly on a particle by particle basis. The OPC we employ is not a single particle scatterer, but rather discriminator levels are used to collect all photo multiplier pulses larger than a preset level. Thus the number concentration in any discriminator bracket is the result of all particles which provide a light signal above the lower level and less than the next discriminator level. These OPCs measure a maximum of 12 sizes between 0.19 and 10.0 μ m, much less than the FSSP, and thus require fitting of size distributions to obtain estimates of backscatter. Estimating backscatter from 12 sizes would be insufficient to compare with the lidar measurements which are inherently ensemble measurements. In addition, the OPC uses white light to measure scattering at 40 degrees, an optical signal not directly comparable to the lidar. Furthermore, instead of using Mie theory, we should use T-matrix calculations and replicate them for many different AR and R_{Th}.