



Mapping of ESA-CCI land cover data to plant functional types for use in the CLASSIC land model

Libo Wang¹, Vivek K. Arora², Paul Bartlett¹, Ed Chan¹, and Salvatore R. Curasi^{3,4}

¹Climate Processes Section, Climate Research Division, Environment and Climate Change Canada, Toronto, ON, Canada

²Canadian Centre for Climate Modelling and Analysis, Climate Research Division, Environment and Climate Change Canada, Victoria, BC, Canada

³Department of Geography and Environmental Studies, Carleton University, Ottawa, ON, Canada.

⁴Climate Processes Section, Climate Research Division, Environment and Climate Change Canada, Victoria, BC, Canada.

Corresponding author: Libo Wang, Libo.Wang@ec.gc.ca



1 **Abstract**

2 Plant functional types (PFTs) are used to represent vegetation distribution in land surface models
3 (LSMs). Large differences are found in the geographical distribution of PFTs currently used in
4 various LSMs. These differences arise from the differences in the underlying land cover products
5 but also the methods used to map or reclassify land cover data to the PFTs that a given LSM
6 represents. There are large uncertainties associated with existing PFT mapping methods since
7 they are largely based on expert judgment and therefore are subjective. In this study, we propose
8 a new approach to inform the mapping or the cross-walking process using analyses from sub-
9 pixel fractional error matrices, which allows for a quantitative assessment of the fractional
10 composition of the land cover categories in a dataset. We use the Climate Change Initiative
11 (CCI) land cover product produced by the European Space Agency (ESA). A previous study has
12 shown that compared to fine-resolution maps over Canada, the ESA-CCI product provides an
13 improved land cover distribution compared to that from the GLC2000 dataset currently used in
14 the CLASSIC (Canadian Land Surface Scheme Including Biogeochemical Cycles) model. A tree
15 cover fraction dataset and a fine-resolution land cover map over Canada are used to compute the
16 sub-pixel fractional composition of the land cover classes in ESA-CCI, which is then used to
17 create a cross-walking table for mapping the ESA-CCI land cover categories to nine PFTs
18 represented in the CLASSIC model. There are large differences between the new PFTs and those
19 currently used in the model. Offline simulations performed with the CLASSIC model using the
20 ESA-CCI based PFTs show improved winter albedo compared to that based on the GLC2000
21 dataset. This emphasizes the importance of accurate representation of vegetation distribution for
22 realistic simulation of surface albedo in LSMs. Results in this study suggest that the sub-pixel



23 fractional composition analyses are an effective way to reduce uncertainties in the PFT mapping
24 process and therefore, to some extent, objectify the otherwise subjective process.

25

26 **1. Introduction**

27 Land cover is a critical component of the earth system that affects the exchange of energy, water,
28 and carbon between the land surface and the atmosphere (Pielke et al., 1998; Sterling et al.,
29 2013). Accurate representation of global land cover (LC) is important for land surface models
30 (LSMs) which provide the lower boundary conditions to the atmosphere in weather forecast,
31 climate, and earth system models. Plant functional types (PFTs) are groups of plant species that
32 share similar structural, phenological, and physiological traits, and have been used in LSMs to
33 represent vegetation distribution. This is because it is far more challenging to model
34 biogeochemical processes, especially photosynthesis and the carbon cycle, at the species level
35 (Bonan et al., 2002; Smith et al., 1997). The number and type of PFTs used in each LSM differ.
36 Global LC datasets are typically used to derive the fractional coverage and geographical
37 distribution of PFTs for use in LSMs. However, large differences exist in fractional coverage and
38 geographical distribution of PFTs, which are caused by differences in the LC datasets themselves
39 but also due to the methods used to map LC datasets to the PFTs represented in various models
40 (Fritz et al, 2011; Hartley et al., 2017; Oettle et al., 2013; Wang et al., 2016).

41 Since different PFTs are characterized by different physical and biogeochemical processes and
42 parameter values, the spatial distribution and fractional cover of PFTs constitute one of the
43 important geophysical fields that are required for realistic simulation of carbon, water, and
44 energy budgets in LSMs (Arora and Boer, 2010; Betts, 2001). For example, the surface albedos



45 for needleleaf evergreen trees, broadleaf deciduous trees, and grasslands are very different,
46 especially during winter when deciduous trees are leafless and short vegetation is largely buried
47 by snow (Bartlett and Verseghy, 2015; Moody et al., 2007). Surface roughness for short or tall
48 vegetation is also very different, which affects simulated turbulent exchanges. Wang et al. (2016)
49 found that the bias in winter albedo in selected boreal forest regions among the models
50 participating in the fifth phase of the Coupled Model Intercomparison Project (CMIP5) was
51 largely related to bias in leaf area index and tree cover fraction. Model experiments using the
52 MPI-ESM by Georgievski and Hagemann (2019) suggested that uncertainties in vegetation
53 distribution may lead to noticeable variations in near-surface climate variables and large-scale
54 circulation patterns.

55 The Canadian Land Surface Scheme Including Biogeochemical Cycles (CLASSIC) is an open-
56 source community land model that is designed to address research questions that explore the role
57 of the land surface in the global climate system (Melton et al., 2020). It is the successor to the
58 coupled modelling framework based on the Canadian Land Surface Scheme (CLASS; Verseghy,
59 1991; 1993) and the Canadian Terrestrial Ecosystem Model (CTEM; Arora and Boer, 2005;
60 Melton and Arora, 2016). The physics and biogeochemistry modules of CLASSIC are based on
61 CLASS and CTEM models, respectively. Since the development of CTEM in the early 2000s,
62 the GLC2000 LC product has been used to specify the spatial distribution of PFTs for CLASSIC
63 when employed as the land surface component of the Canadian Earth system model developed
64 by Environment and Climate Change Canada (Arora et al., 2009; Wang et al., 2006). The
65 Climate Change Initiative (CCI) LC product recently produced by the European Space Agency
66 (ESA) is available at an annual temporal resolution for the period 1992 to 2018 at 300 m spatial
67 resolution (ESA, 2017). It was produced based on broad user consultation, specifically to address



68 the needs of the climate modelling community (Bontemps et al., 2012). Wang et al. (2019)
69 showed that when compared to the finer resolution maps over Canada, the 300 m ESA-CCI
70 product provides much improved LC distribution over Canada compared to that from the 1 km
71 GLC2000 dataset.

72 To map LC classes to PFTs, a cross-walking table (CW-table) is usually created to assign
73 fractions of each LC class to the different PFTs, such that the sum of the fractions for each class
74 is always one (including fractions of water and bare ground). Previous methods for creating such
75 CW-tables are mainly based on LC class descriptions, expert knowledge, and the spatial
76 distribution of global biomes (Ottle et al., 2013; Poulter et al., 2011; 2015; Sun and Liang, 2007;
77 Wang et al., 2006). Because LC maps only provide the types of vegetation, and each class can be
78 associated with a broad range of fractional cover of either one or more vegetation types, there are
79 large uncertainties associated with any cross-walking or reclassification process. Wang et al.
80 (2019) reclassified the 10 PFTs in the default CW-table provided in the ESA-CCI LC product
81 user manual (Table 7-2, ESA, 2017) into four broad PFTs in CLASS, and compared them with
82 those based on the GLC2000 dataset. The results suggest that uncertainties in the CW-tables
83 were a major source of large differences in the PFT distributions. In addition, the fractional
84 coverage of tree PFTs based on the default CW-table for the ESA-CCI LC dataset was
85 overestimated along the taiga-tundra transition zone in western Canada, which led to
86 underestimation in winter albedo in CLASSIC offline simulations driven with observation-based
87 reanalysis data (Wang et al., 2018).

88 The objective of this study is to develop a new CW-table for reclassifying the ESA-CCI LC
89 classes into PFTs used in the CLASSIC model for use over the model's Canada domain and to
90 compare and assess the performance of CLASSIC offline simulations using the new and existing



91 PFTs. Given the close link between the bias in winter albedo and the vegetation distribution in
92 the models (Wang et al., 2016), our assessment of model performance focuses on the simulated
93 surface albedo during the maximum snow accumulation period (February–March for the boreal
94 forest). This simplifies our analyses by excluding the fall/spring transition periods when biases in
95 snow accumulation and melt timing can have a large influence on surface albedo simulated by
96 LSMs (Wang et al., 2014). In addition, we extend the CW-table for the ESA-CCI LC dataset to
97 the global domain. A comprehensive assessment of the impact of the PFT distribution based on
98 the new CW-table and the ESA-CCI LC dataset on the performance of the CLASSIC model at
99 the global scale is presented in Arora et al. (2022).

100 **2. Data and the CLASSIC model**

101 **2.1 The Hybrid LC map over Canada**

102 The open access to the United States Geological Survey archive of Landsat imagery has provided
103 free access to georeferenced and spectrally analysis-ready imagery (Wulder et al., 2012), which
104 makes it possible to generate time series of LC maps to study LC change. Recently two of these
105 products based on Landsat imagery are generated over Canada, including the North America
106 Land Change Monitoring System (NALCMS) LC dataset (Latifovic et al., 2017) and the Virtual
107 Land Cover Engine (VLCE) framework-generated LC dataset (Hermosilla et al., 2018).

108 Based on the random forest algorithm and local optimization method, the Canada Centre for
109 Remote Sensing has generated LC maps of Canada for the years 2010 and 2015 at 30 m
110 resolution using Landsat imagery (Latifovic et al., 2017). These LC products are the Canadian
111 contribution to the 30 m resolution 2010/2015 LC map of North America to the joint
112 collaborative effort by the Mexican, American, and Canadian government institutions under the



113 NALCMS umbrella. The NALCMS LC map has 19 classes based on the United Nations Land
114 Cover Classification System (LCCS; Di Gregorio, 2005). Assessment based on reference
115 samples showed an overall accuracy of 76.6% for the year 2010 data (Latifovic et al., 2017),
116 which is used in this study.

117 VLCE is an automated framework to enable annual LC mapping using time series of Landsat
118 surface reflectance, informed by spatially explicit forest change and a *priori* knowledge of
119 ecological succession. Temporally consistent annual LC maps representative of Canada's
120 forested ecosystems from 1984 to 2012 were generated using the VLCE framework,
121 characterizing LC dynamics following wildfire and harvesting events by Hermosilla et al.
122 (2018). The VLCE maps have 12 LC classes in a hierarchical classification structure following
123 that of the National Forest Inventory. Assessment based on reference samples showed an overall
124 accuracy of 70.3% for the map of 2005 (the year with the greatest number of reference samples;
125 Hermosilla et al., 2018). Land cover data from the year 2010 are used in this study.

126 Overall, the 19-class NALCMS product presents a more detailed LC distribution than the 12-
127 class VLCE map over Canada. For example, areas classified as “Exposed/Barren lands” in the
128 VLCE map correspond to either “Sub-polar or polar grassland-lichen-moss”, “Sub-polar or polar
129 barren-lichen-moss”, or “Barren lands” in the NALCMS map. Areas of cropland are not
130 separated from grassland in the VLCE map. A recent study showed that the wetland class in
131 NALCMS suffers from large uncertainty in forest cover mapping because treed-wetland was not
132 separated from herbaceous wetland in its legend (Wang et al., 2019). To take advantage of both
133 datasets, we created a hybrid product by combining them through the following steps: (1)
134 Reproject the VLCE data from its Lambert Conformal Conic projection to the same Lambert
135 Azimuthal Equal Area projection that is used for the NALCMS data; (2) Replace pixels



136 classified as “Exposed/Barren lands” and “Bryoids” in the VLCE data with the more specific LC
137 classes from the NALCMS data; (3) Replace pixels classified as “Herbs” in the VLCE data with
138 the “Cropland” class in the NALCMS data; (4) and merge the rest of LC classes from NALCMS
139 to the corresponding classes in the VLCE data. There are a total of 17 classes in this new hybrid
140 product, which we will henceforth refer to it as the Hybrid LC dataset and is shown in Figure 1.

141 **2.2 The global LC products**

142 The GLC2000 dataset was generated from SPOT/VEG data collected from November 1999 to
143 December 2000 at 1 km resolution (Bartholomé and Belward, 2005). It was produced by 21
144 separate regional expert groups using an unsupervised image classification method. Based on the
145 LCCS, the regional products were merged into one global product with a generalized LCCS
146 legend of 22 classes. Assessment based on a random sampling of reference sites globally
147 estimated an overall accuracy of 68.6% for the GLC2000 product (Mayaux et al., 2006).

148 The annual ESA-CCI LC data at 300 m resolution are available for the period 1992-2018, which
149 were generated from baseline data and annual LC changes (ESA, 2017). The baseline data were
150 generated using a combination of machine learning and unsupervised image classification
151 methods from the entire archive of ENVISAT/ Medium Resolution Imaging Spectrometer for the
152 period of 2003-2012. The annual LC changes were detected at 1 km resolution from the
153 Advanced Very High Resolution Radiometer time series between 1992 and 1999, SPOT/VEG
154 time series between 1999 and 2013, and the PROBA-V time series between 2013 and 2018.
155 Based on the LCCS legend, the ESA-CCI LC data have 22 level 1 classes, and 15 level 2 sub-
156 classes. Assessment based on the GlobCover validation database estimated an overall accuracy
157 of 71% for the ESA-CCI LC product (ESA, 2017).



158 2.3 Other datasets

159 Airborne Lidar has been used to monitor forests since the 1980s and is well suited to estimate
160 vegetation height, volume, and biomass (Hopkinson et al., 2006; Wulder et al., 2008). Vegetation
161 cover percentage for canopy height above 2 m from airborne Lidar data are used to estimate the
162 fraction of tall versus low vegetation for LC classes with a mix of woody and herbaceous
163 vegetation in this study. The Lidar data were collected along 34 survey flights across the boreal
164 forest of Canada in the summer of 2010 by the Canadian Forest Service (Wulder et al., 2012). A
165 25 by 25 m tessellation was generated with the approximately 400 m wide Lidar swath, with
166 each cell treated as an individual Lidar plot.

167 A tree cover fraction (TCF) dataset for 2010 is also used in this study (Hansen et al., 2013;
168 hereafter the Hansen TCF dataset). It was based on Landsat images at 30 m resolution. In
169 contrast to the discrete LC classification datasets (providing a certain number of LC classes) as
170 described above, the Hansen dataset is a vegetation continuous field product (providing tree
171 cover fractions from 0 – 100%), in which the satellite spectral information was used to estimate
172 the TCF in each pixel using a regression tree algorithm (Hansen et al., 2002; 2010). This may
173 better represent heterogeneous areas than is possible by discrete LC classification. Tree cover is
174 defined to exist over pixels where canopy closure is observed for vegetation taller than 5 m in
175 height. Forests are generally defined as woody vegetation taller than 3 m in the regional and
176 global LC datasets. The different definitions of tree heights should not result in much difference
177 in areas with mature forests, such as most boreal forests in Canada.

178 Simulated surface albedo by the CLASSIC model in offline experiments is evaluated against the
179 Moderate Resolution Imaging Spectroradiometer (MODIS) (MCD43C3) broadband (0.3–5.0



180 μm) white-sky albedo (Schaaf et al., 2002), with quality flags of 0–2 (75% or more full
181 inversions and 25% or fewer fill values) and solar zenith angles less than 70° . The MODIS
182 albedo product used in this study is at 0.05 degree resolution, and is regridded to the 0.22 degree
183 model grid (see Section 2.4.2).

184 **2.4 The CLASSIC model and simulation setup**

185 **2.4.1 The CLASSIC model**

186 CLASSIC is the successor to the coupled modelling framework based on the Canadian Land
187 Surface Scheme (CLASS; Verseghy, 1991; 1993) and the Canadian Terrestrial Ecosystem Model
188 (CTEM; Arora and Boer, 2005; Melton and Arora, 2016). The physics and biogeochemistry
189 components of CLASSIC are based on CLASS and CTEM, respectively.

190 For the physics component, vegetation is represented in terms of the fractional coverage of the
191 four PFTs (needleleaf trees, broadleaf trees, crops, and grasses). The physics component
192 represents a single snow layer with variable depth and a single vegetation canopy layer. As a
193 first-order treatment of subgrid-scale heterogeneity, each grid cell is divided up into four sub-
194 areas, consisting of vegetated and bare soil areas, each with and without snow cover. The grid
195 cell albedo is computed as a weighted mean using the fractional coverages for each surface type.
196 Canopy snow processes such as interception/unloading, sublimation, and melt are simulated by
197 the physics module of CLASSIC. Twenty ground layers represent the soil profile, starting with
198 10 layers of 0.1 m thickness. The thicknesses of the layers gradually increase to 30 m for a total
199 ground depth of over 61 m. Liquid and frozen soil moisture contents, and soil temperature, are
200 determined prognostically for permeable soil layers.

201 The biogeochemistry component of CLASSIC represents vegetation in terms of nine PFTs:
202 Needleleaf Evergreen trees (NLE), Needleleaf Deciduous trees (NLD), Broadleaf Evergreen



203 trees (BLE), Broadleaf Cold Deciduous trees (BCD), Broadleaf Dry Deciduous trees (BDD), C₃
204 and C₄ Crops (C3C/C4C), and C₃ and C₄ Grasses (C3G/C4G). These nine PFTs map directly
205 onto the four PFTs used by CLASSIC's physics component. When the physics and
206 biogeochemistry components are coupled together, the structural attributes of vegetation
207 including leaf area index, canopy mass, rooting depth, and vegetation height are simulated
208 dynamically as a function of environmental conditions by the biogeochemistry component.
209 When the biogeochemistry component is turned off, specified structural attributes of vegetation
210 for use by the physics component are extracted from look-up tables.

211 **2.4.2 Simulation set up**

212 Gridded meteorological data based on the Climatic Research Unit (CRU,
213 <https://crudata.uea.ac.uk/cru/data/hrg/>) and Japanese reanalysis (JRA) (CRUJRA) are used to
214 drive CLASSIC simulations. The data are constructed by regridding data from the JRA and
215 adjusting where possible to align with the CRU TS 4.05 data. The blended product from year
216 January 1901 to December 2020 has the 6-hourly temporal resolution of a reanalysis product but
217 monthly means adjusted to match the CRU data (Harris, 2020). The 6-hourly data are
218 disaggregated into half-hourly data for the following seven meteorological variables that are used
219 to force the model: 2 m air temperature, total precipitation, specific humidity, downward solar
220 radiation flux, downward longwave radiation flux, surface pressure, and wind speed. In
221 CLASSIC, the phase of precipitation is determined by a threshold surface air temperature with
222 three options available (Bartlett et al., 2006). The 0°C air temperature threshold is used to
223 partition precipitation into rain or snow in this study. This choice should not have a significant
224 impact on the simulated surface albedo in CLASSIC (Wang et al. 2014).



225 Two simulations over the 1850-2020 historical period are performed using PFTs derived from
226 the ESA-CCI and the GLC2000 datasets respectively, which is the only difference between the
227 two simulations. Besides land cover and meteorological forcings, CLASSIC requires globally
228 averaged atmospheric CO₂ concentration, and geographically varying time-invariant soil texture
229 and soil permeable depth. The atmospheric CO₂ concentration values are provided by the Global
230 Carbon Project protocol (<https://www.globalcarbonproject.org/index.htm>). The soil texture
231 information consists of the percentage of sand, clay, and organic matter and is derived from the
232 SoilGrids250m dataset (Hengl et al., 2017), and permeable soil depth is based on Shangguan et
233 al. (2017). The simulations are performed at a 0.22 degree rotated latitude-longitude grid over a
234 domain including Canada and part of Alaska (Fig. 3). Pre-industrial simulations that correspond
235 to the year 1850 are required prior to doing the historical simulations so that model's carbon
236 pools, including leaf biomass which determines leaf area index, are spun up to near equilibrium
237 for each land cover. The pre-industrial simulations use 1901-1920 meteorological data repeatedly
238 with atmospheric CO₂ concentration specified at its 1850 level. Each historical simulation is then
239 initialized from its corresponding pre-industrial simulation after it has reached equilibrium. For
240 the period 1851-1900, the CRUJRA meteorological data for the first 20 years (1901-1920) are
241 used repeatedly. For the 1901-2020 period the meteorological data corresponding to each actual
242 year are used. The period from 2001 to 2015 was selected for analyzing the simulated results.

243 **3. PFT mapping methods and results**

244 A CW-table for the Hybrid LC map based on the Lidar plots, tree species, field surveys, and
245 high-resolution images is first created. We then compute the sub-pixel fractional composition of
246 the LC classes in the 300 m ESA-CCI data using the 30 m Hansen TCF and the 30 m Hybrid LC
247 data respectively, which are in turn used to create the CW-table for the ESA-CCI dataset (Fig. 2).



248 Finally the CW-table developed for the ESA-CCI dataset over Canada is extended to the whole
249 globe as explained in Section 3.3.

250 **3.1 CW-table for mapping Hybrid LC classes to CLASSIC PFTs**

251 Among the nine CLASSIC PFTs, BLE and BDD forests are not present in Canada. These are
252 primarily tropical PFTs as represented in CLASSIC. NLD only accounts for less than 1% of
253 coniferous forests in Canada (Wang et al., 2019). Therefore we do not consider NLD, BLE, and
254 BDD from here on in this study. Considering the fine resolution (30 m) of the Hybrid map,
255 especially relative to the model resolution (~16 km) used in this study, we assign fractions of 1.0
256 to the two pure forest classes (LC210 and LC220), the cropland (LC15), and the five non-
257 vegetative classes (LC16 to LC32) in its CW-table (Table 1). The mixed-wood category (LC230)
258 is split evenly into NLE and BCD in the table based on the definition in the VLCE legend
259 (Hermosilla et al., 2018; Wulder et al., 2003). Note that in Table 1, BDD and BCD, and C₃ and
260 C₄ crops and grasses, are considered together. The reason for this is that the separation of
261 broadleaf trees into their cold and deciduous phenotypes is based on latitude (Wang et al., 2006).
262 The separation of crops and grasses based on their photosynthetic pathway (C₃ or C₄) is done
263 based on the C₄ fraction from Still et al. (2003).

264 CLASSIC does not yet explicitly represent shrub PFTs (research on including shrubs as a
265 separate PFT is ongoing), and therefore the fraction of tall shrubs is assigned to one of the tree
266 PFTs as was done in creating the CW-table for GLC2000 for use with CLASSIC (Wang et al.,
267 2006). Four (LC2, LC50, and LC80-81) out of the 17 classes in the Hybrid map are characterized
268 by a mosaic of trees, shrubs, and herbaceous vegetation. The vegetation coverage for canopy
269 height above 2 m from Lidar plots is used to inform the partitioning of forest (tall vegetation) to



270 non-forest (low vegetation) fractions for these classes. The distribution of tree species from
271 Beaudoin et al. (2014) is used to guide the splitting of coniferous versus broadleaf forest
272 fractions. For example, for the Wetland-treed category (LC81), maps of tree species show that
273 coniferous forests dominate wetland-treed regions, while broadleaf forests are generally non-
274 existent. We, therefore, assign most of the forest fraction to NLE (0.55), only 0.05 to BCD, 0.35
275 to grasses, and the remaining to the bare ground for LC81 (Table 1). The presence of evergreen
276 shrubs is rare in Canada according to National Forest Inventory ground plots data (Gillis et al.,
277 2005), thus we only assign an estimated tall shrub fraction (0.20) in the shrub class (LC50) to
278 BCD. The sub-polar or polar classes (LC11 to LC13) are located above the treeline and mainly
279 consist of low shrubs and grass. The fractions of grass (including low shrubs) and bare ground
280 are based on field surveys of fractional vegetation cover and tundra PFTs data in Bjorkman et al.
281 (2018) and Macander et al. (2020). High-resolution images from Google Earth engine or Bing
282 Maps are also used to examine the ratio of vegetated versus bare ground for all classes in which
283 bare ground is present.

284 **3.2. CW-table for mapping ESA-CCI LC classes to CLASSIC PFTs over Canada**

285 **3.2.1 The error and sub-pixel fractional error matrices**

286 A standard approach for the accuracy assessment of LC products is to use an error matrix. It is a
287 square array or table of numbers arranged in rows and columns, in which the classification from
288 the LC product (usually represented by the rows) is compared to the reference data (usually
289 represented by the columns) for each category (Congalton, 1991). The major diagonal of the
290 matrix presents the number of correct classifications indicating the agreement between the LC
291 and the reference data for each category. In practice, fine-resolution regional maps are often used
292 to assess large-scale LC products derived from coarse-resolution data (Cihlar et al., 2003). In



293 such cases, the fine-resolution reference data are aggregated/regridded to match the grid of the
294 coarse-resolution data. Several classes in the reference data may be present in a single coarse-
295 resolution pixel depending on the homogeneity of the landscape. In order to compare the
296 reference and the LC data on a one-to-one basis, the dominant LC class (the class with the most
297 abundant fractions based on all fine-resolution pixels in the reference data) is often assigned to
298 the regridded reference pixel.

299 The sub-pixel fractional error matrices have been introduced as a more appropriate way of
300 assessing the accuracy of mixed pixels by Latifovic and Olthof (2004). In contrast with an error
301 matrix where only the dominant LC class is used as described above, the sub-pixel fractional
302 error matrix is produced by assigning sub-dominant LC classes from all fine-resolution pixels in
303 the reference data to the corresponding single coarse-resolution pixel. It thus allows a
304 quantitative assessment of the fractional composition of the LC classes in the coarse resolution
305 dataset. In this study, both the 30 m Hansen TCF data and the 30 m Hybrid LC map are used to
306 compute the sub-pixel fractional error matrices of the 300 m ESA-CCI dataset (Table 2 and
307 Table 3). However, the objective here is not an accuracy assessment as in Latifovic and Olthof
308 (2004) but rather to obtain the fractional composition of the LC classes in the ESA-CCI product
309 and to inform the PFT mapping process. We refer this process as the sub-pixel fractional
310 composition analyses in the rest of this paper. Sub-pixel fractional composition analyses is first
311 performed for each ecozone and then weighted mean fractions for each ESA-CCI class are
312 computed based on pixel counts in each of the ecozones (see the location of ecozones in Fig. 1).
313 For the Hansen TCF data, results are shown only for the ESA-CCI LC classes containing forests
314 in Canada (Table 2). In the ESA-CCI legend (Table 4), two sub-classes for broadleaf (LC61 and
315 LC62) and needleleaf (LC71 and LC72) forests are included as the closed (>40% forest cover)



316 and open (10-40% forest cover) categories apart from the main classes (LC60 and LC70). As
317 expected, the TCF is larger for the closed classes than for the main and the open classes (Table
318 2). In Table 2, we also include ratios of TCF between the main class and the closed class, and
319 between the open class and the closed class. We note that the ratios are different for broadleaf
320 (main class vs. closed class: $68.5/86.7=0.8$; open class vs. closed class: $0.43/86.7=0.43$) and
321 needleleaf (main class vs. closed class: $39.3/61.7=0.6$; open class vs. closed class:
322 $23.2/61.7=0.38$) forests, which need to be taken into account when creating the CW-table for the
323 ESA-CCI dataset.

324 To obtain representative class compositions of the ESA-CCI dataset, only homogenous ESA-CCI
325 pixels are included in the sub-pixel composition analyses based on the Hybrid LC data. The
326 homogenous pixels are defined following the method in Herold et al. (2008). To quantify
327 landscape heterogeneity, 3×3 pixel neighborhoods are assessed for the ESA-CCI data. A
328 neighborhood is considered homogenous if only one LC class is present. The weighted mean
329 fraction for each class is computed from ecozones with more than 10 homogenous ESA-CCI
330 pixels for that class (Table 3). Only 13 out of the 37 ESA-CCI classes are included in Table 3,
331 the rest of the classes either have limited presence in Canada or are non-vegetative (Table 4).

332 In the Hybrid CW-table (Table 1), four LC classes (2, 81, 210, and 230) contribute to the
333 fractional cover of NLE, and five LC classes (50, 80, 81, 220, and 230) contribute to the
334 fractional cover of BCD. In Table 3, we also include an integrated fractional cover (F) for NLE
335 and BCD (last two rows) for each of the ESA-CCI classes based on the following formula:

336
$$F = \sum_{i=1}^N F1_i * F2_i \quad (1)$$



337 Where $F1_i$ are fractions in Table 3, $F2_i$ are fractions in Table 1, and N is the number of Hybrid
338 LC classes contributing to NLE ($N = 4$) or BCD ($N = 5$) as shown in Table 1. As an example, the
339 fraction of NLE for the LC70 (Tree cover needleleaf evergreen closed to open) in ESA-CCI data
340 in Table 3 (see the NLE row and the column for class 70) is obtained as follows: $F = 0.02 \times 0.20 +$
341 $0.17 \times 0.55 + 0.29 \times 1.0 + 0.09 \times 0.5 = 0.44$. This process reduces the subjectivity in assigning the
342 ESA-CCI land cover classes to CLASSIC's two tree PFTs (NLE and BCD) that are present in
343 Canada since the process is based on the high-resolution Hybrid LC data.

344 **3.2.2 CW-table for the ESA-CCI LC dataset over Canada**

345 Table 2 and Table 3 thus form the basis for creating the CW-table for mapping the ESA-CCI LC
346 classes to CLASSIC's PFTs (Fig. 2 and Table 4). For the ESA-CCI class LC61 (Tree cover
347 broadleaved deciduous closed) (not included in Table 3 due to limited presence in Canada),
348 ratios of TCF for LC60 vs LC61 in Table 2 and the fractions of LC60 (Tree cover broadleaved
349 deciduous closed to open) in Table 3 are used to derive fractions for LC61 in Table 4. The
350 remapping of LC62 (Tree cover broadleaved deciduous open) and LC72 (Tree cover needleleaf
351 evergreen open) into CLASSIC's PFTs is done in a similar way. Since NLD is not included in
352 either Table 2 or Table 3, the needleleaf deciduous tree cover classes (LC80-82) are assigned to
353 the same fractions as the needleleaf evergreen tree cover classes (LC70-72). For simplicity, the
354 fractions in Table 3 are rounded to values with either "0" or "5" at the hundredth place when
355 used in Table 4. For the rest of the classes not included in either Table 2 or Table 3, values are
356 based on the default CW-table from the ESA-CCI user guide (Table 7-2, ESA, 2017). The spatial
357 distribution of LC classes is also taken into consideration when determining the fractions in the
358 CW-table. For example, the sparse vegetation class (LC150) is mainly distributed above the



359 treeline in alpine and Arctic tundra environments, thus we only assign 0.05 to BCD, the rest to
360 C3G/C4G and bare ground (Table 4).

361 The six CLASSIC PFTs (those present in Canada) are produced from the Hybrid and the ESA-
362 CCI maps based on Table 1 and Table 4 respectively. The PFTs from the Hybrid map are used as
363 a reference here to map ESA-CCI land cover classes to CLASSIC's PFTs. To make the spatial
364 distribution of PFTs from ESA-CCI agree better with those from the Hybrid dataset, fractions for
365 the following classes in Table 4 are adjusted upward by 0.05: LC60 from 0.65 to 0.70 for BCD;
366 LC71 and LC81 from 0.80 to 0.85 for NLE; and LC120 from 0.10 to 0.15 for BCD. Values for
367 LC10-20 are also slightly adjusted to reduce crop fractions.

368 **3.3. CW-table for the ESA-CCI LC dataset over the globe**

369 Based on Table 4, the fractional coverage of nine CLASSIC PFTs are also produced on a global
370 scale. However, fractions of NLE in Eurasia are found to be too low relative to the Hansen TCF
371 data, with maximum values only around 0.45 in most NLE dominated areas, where the
372 maximum TCF from the Hansen dataset is around 0.80. Needleleaf evergreen forests are
373 represented by LC classes 70 (closed to open), 71 (closed), and 72 (open). Examining the ESA-
374 CCI LC map shows that in Eurasia nearly all needleleaf evergreen forests are classified as LC70
375 (closed-to-open), with only less than 400 pixels as LC71 (closed), and none as LC72 (open). In
376 contrast, in Canada 36% of needleleaf evergreen forest is classified as LC70 (closed-to-open),
377 64% as LC71 (closed), and less than 1% as LC72 (open). Sub-pixel fractional composition
378 analyses of the ESA-CCI classes based on the Hansen TCF dataset show that in Eurasia TCF for
379 LC70 (closed-to-open) is 66% and for LC71 (closed) is 35% (note few pixels with this class).
380 This is in contrast with those in Canada where TCF for LC70 (closed-to-open) is 39% and for



381 LC71 (closed) is 62%, explaining the too low NLE fractions in Eurasia when mapping PFTs
382 based on Table 4, and also the too high TCF in northwestern Canada when mapping PFTs based
383 on the default CW-table (Wang et al., 2018). In order to apply Table 4 globally, the original
384 LC70 (closed-to-open) was split into two classes: LC73 (a new class) which is the same as LC70
385 over Canada (and zero everywhere else), and LC70 (revised) which is the same as before except
386 zero over Canada. The fractions for the new LC70 class are made the same as for LC71 in Table
387 4, which applies to NLE outside of Canada. Essentially, the closed-to-open needleleaf forest
388 LC70 class over Eurasia is treated as the closed needleleaf forest.

389 The global PFTs based on Table 4 are used in offline CLASSIC simulations in Arora et al.
390 (2022), who carry out a comprehensive assessment of the impact of using two different LC
391 datasets (ESA-CCI versus GLC2000) for representing the nine CLASSIC PFTs. They also assess
392 the effect of using land cover reconstructions based on the ESA-CCI and GLC2000 datasets on
393 the simulated surface energy, water, and CO₂ fluxes in the CLASSIC model.

394 **3.4 Comparison of PFTs from Hybrid, ESA-CCI, and GLC2000 data**

395 Figure 3 shows the spatial distribution of PFTs derived from the Hybrid, ESA-CCI, and
396 GLC2000 LC datasets respectively. C₄ crops (C₄C) and grasses (C₄G) are sparse in Canada as
397 would be expected since C₄ PFTs grow only in warmer temperatures when the average monthly
398 temperature exceeds 22 °C (Fox et al., 2018). Therefore, only four out of the six PFTs (those
399 present in Canada) are shown in Figure 3. In general, the spatial distributions of the PFTs from
400 the ESA-CCI and the Hybrid datasets agree well except for C₃ grasses (C₃G) (Fig. 3j and Fig.
401 3k). This is not surprising given that the CW-table for the ESA-CCI dataset is based on the
402 Hybrid map. Areas mapped as C₃G in Hybrid (Fig. 3j), were mainly classified as sparse



403 vegetation (LC150) in the ESA-CCI legend (Table 4). However, LC150 from ESA-CCI was also
404 found in some areas of the high Arctic islands, where barren land is the dominant class in the
405 Hybrid map (grey coloured areas in Fig. 1). If too much grass were assigned to LC150, it would
406 yield unrealistically large fractional coverage of grass in the high Arctic islands. In Table 4, for
407 LC150, 0.05 is assigned to BCD, 0.35 to grasses, and the rest to the bare ground for LC150,
408 which yields a total vegetation cover of 40% and is more than the definition (<15% vegetation)
409 used in the ESA-CCI legend. Yet, this still results in less C3G and less bare ground in the ESA-
410 CCI map (Fig. 3k and Fig. 3n) than those from the Hybrid map (Fig. 3j and Fig. 3m). This
411 suggests that it is not ideal to classify areas in the high Arctic islands and in the Arctic tundra
412 region as being in the same land cover category.

413 There are large differences in the spatial distribution of the PFTs based on the GLC2000 LC
414 product and those based on the Hybrid and ESA-CCI datasets (Fig. 3 and Fig. 4). Relative to
415 PFTs from ESA-CCI, GLC2000 has less NLE and more BCD in northwestern Canada, and more
416 NLE in southern and eastern Canada (Fig. 4a and Fig. 4b). GLC2000 based CLASSIC PFT
417 fractions also exhibit more crops, less grass, and more bare ground (Fig. 4c-4e). These
418 differences partly stem from the differences in the ESA-CCI and GLC2000 LC datasets, but are
419 also due to the fact how the fractions in the CW-tables of the two datasets are used to translate
420 LC data to fractional coverage of PFTs as demonstrated in Wang et al. (2019).

421 **3.5 Bias in simulated surface albedo and LAI**

422 The top row of Figure 5 shows the bias in winter albedo (March) simulated by CLASSIC when
423 using PFT distributions based on the ESA-CCI (Fig. 5a) and GLC2000 products (Fig. 5b). While
424 model biases are the result of both the driving geophysical and meteorological data that are used



425 to force the model, as well as the model itself, the comparison between the two simulations does
426 show the effect of differences in the distribution of PFTs. Relative to observed surface albedo
427 from MODIS, there are relatively large negative biases in the southwest of Hudson Bay and
428 central Quebec, while there are relatively large positive biases in western Canada and Alaska in
429 the simulation when using the GLC2000 product to obtain PFT distributions (Fig. 5b). Both the
430 negative and the positive biases are largely reduced in the simulation using PFT distributions
431 based on the ESA-CCI product (Fig. 5a). The lower row of Figure 5 shows the spatial
432 distribution of the difference in surface albedo (Fig. 5c) and leaf area index (Fig. 5d) between the
433 two simulations, which are closely correlated ($r = -0.85$). Given the same meteorological forcing
434 dataset is used to drive both simulations, the differences in the simulated LAI are due mainly to
435 the different PFT distributions used in the two simulations. Since NLE is the only PFT with
436 LAI > 0 during winter in Canada, the LAI difference in March as shown in Figure 5d is mainly
437 due to the different fractional coverage of NLE based on the ESA-CCI and GLC2000 products
438 (Fig. 4a).

439 In contrast, the large positive albedo biases (up to ~ 0.4) in southern Canada are more or less the
440 same in both simulations (Fig. 5a and Fig. 5b), where the dominant PFT is C3 crops (Fig. 3h and
441 Fig. 3i). Those positive albedo biases are likely due to the standing crop stubble not treated in
442 CLASSIC currently. Harder et al. (2018) showed that the height of the stubble over wheat and
443 canola field in Saskatchewan, Canada may range from 10 to 40 cm, with a maximum PAI (plant
444 area index) of 1.0. Wang et al. (2016) showed that surface albedo in CLASSIC decreased
445 exponentially with increasing PAI for the bare or snow-covered canopy over snow, while most
446 reductions of the albedo were achieved through the increase of PAI from 0 to 1.0. They showed
447 that surface albedo decreased from 0.75 to 0.31 in CLASSIC when PAI increased from 0 to 1.0



448 for the bare canopy over snow, which appears to fully account for the positive albedo biases in
449 the agricultural areas of southern Canada (Fig. 5a and Fig. 5b). Improvements to the crop module
450 of CLASSIC to improve cropland albedo are currently being considered.

451 **4. Summary and conclusions**

452 A hybrid land cover map at 30 m resolution is created by merging the NALCMS and VLCE land
453 cover products over Canada. Vegetation height data from Lidar plots, tree species, and high
454 resolution images are used to inform the creation of a CW-table for mapping the 17 LC classes
455 of the Hybrid map to six CLASSIC PFTs that are present in Canada. Both the Hybrid map and
456 the Hansen tree cover fraction data are used to compute the sub-pixel fractional composition of
457 the LC classes in the ESA-CCI LC dataset, which is then used to create a cross-walking table for
458 mapping the 37 ESA-CCI categories to CLASSIC PFTs. Based on the new CW-tables, PFT
459 distributions are produced from the Hybrid and the ESA-CCI LC products, respectively, and are
460 compared with those based on the GLC2000 dataset currently used in CLASSIC. The results
461 show that the spatial distribution of PFTs from the ESA-CCI dataset is in better agreement with
462 those from the Hybrid map, while there are large differences in the PFTs from the GLC2000
463 dataset and from the Hybrid/ESA-CCI datasets.

464 CLASSIC simulations driven with meteorological data from the CRU-JRA product show that the
465 simulated winter albedo is improved when using PFT distributions based on the ESA-CCI LC
466 product compared to that based on the GLC2000 product, which is consistent with findings from
467 previous studies. While, CLASSIC simulations could also have been performed using its PFT
468 distributions based on the Hybrid LC product, the reason for using the ESA-CCI based PFT
469 fractions for CLASSIC is that ESA-CCI is a global product. CLASSIC simulations are routinely



470 performed at the global scale both in the framework of the Canadian Earth System Model (Swart
471 et al., 2019), where CLASSIC serves as its land component, and offline where global CLASSIC
472 simulations driven with the CRU-JRA meteorological data contribute to the annual global carbon
473 budget assessments of the Global Carbon Project (Friedlingstein et al., 2020). Untreated crop
474 stubble appears to be the cause for positive winter albedo biases in southern Canada, which
475 needs to be addressed in a future version of CLASSIC. These results underscore the importance
476 of accurate representation of vegetation distribution in a realistic simulation of surface albedo in
477 LSMs.

478 Previous methods for mapping PFTs from LC datasets are mainly based on class descriptions,
479 expert knowledge, and the spatial distribution of global biomes, which is a largely subjective
480 process. As a consequence, a PFT method developed for mapping one LC dataset to PFTs
481 represented in one model can not be easily transferred to other LC datasets even for deriving
482 PFTs in the same model. The development of satellite and computing technology has enabled the
483 creation of more detailed global LC products at finer spatial resolutions in recent years, however,
484 the lack of an objective PFT mapping method impedes the implementation of the new improved
485 LC products in LSMs. Here, we have proposed a method to inform the cross-walking process
486 using sub-pixel fractional composition analyses based on a tree cover fraction dataset and a fine-
487 resolution LC map. Our results suggest that the sub-pixel fractional composition analyses
488 provide an effective way to reduce uncertainties in the cross-walking process and therefore, to
489 some extent, objectifies the otherwise subjective process. The PFT mapping approach developed
490 in this study can also be applied to other LC datasets for mapping PFTs used in other LSMs.

491



492 **Acknowledgements.**

493 We'd like to thank Mike Brady (ECCC) for his help with data processing and especially for
494 creating the Hybrid map, and Mike Wulder from the Canadian Forest Service of Natural
495 Resources Canada for providing the Lidar plots data and for helping interpret the Lidar data and
496 the VLCE land cover map.

497

498 **References**

499 Arora, V. K., Boer, G. J.: A parameterization of leaf phenology for the terrestrial ecosystem
500 component of climate models. *Glob. Chang. Biol.*, 11, 39–59, 2005.

501 Arora, V. K. and Boer, G. J.: Uncertainties in the 20th century carbon budget associated with
502 land use change, *Glob. Change Biol.*, 16, 3327–3348, [https://doi.org/10.1111/j.1365-](https://doi.org/10.1111/j.1365-2486.2010.02202.x)
503 [2486.2010.02202.x](https://doi.org/10.1111/j.1365-2486.2010.02202.x), 2010.

504 Arora, V. K., Boer, G. J., Christian, J. R., Curry, C. L., Denman, K. L., Zahariev, K., Flato, G.
505 M., Scinocca, J. F., Merryfield, W. J., and Lee, W. G.: The Effect of Terrestrial Photosynthesis
506 Down Regulation on the Twentieth-Century Carbon Budget Simulated with the CCCma Earth
507 System Model, *J. Clim.*, 22, 6066–6088, <https://doi.org/10.1175/2009JCLI3037.1>, 2009.

508 Beaudoin, A., Bernier, P.Y., Guindon, L., Villemaire, P., Guo, X.J., Stinson, G., Bergeron, T.,
509 Magnussen, S., and Hall, R.J.: Mapping attributes of Canada's forests at moderate resolution
510 through kNN and MODIS imagery, *Can. J. For. Res.* 44: 521–532, [dx.doi.org/10.1139/cjfr-2013-](https://doi.org/10.1139/cjfr-2013-0401)
511 [0401](https://doi.org/10.1139/cjfr-2013-0401), 2014.



- 512 Bartholomé, E., and Belward, A. S.: GLC2000: A new approach to global land cover mapping
513 from Earth Observation data, *Int. J. Remote Sens.*, 26, 1959–1977, 2005.
- 514 Bartlett, P. A., MacKay, M. D., and Verseghy, D. L.: Modified snow algorithms in the Canadian
515 land surface scheme: Model runs and sensitivity analysis at three boreal forest stands, *Atmos.–*
516 *Ocean*, 44, 207–222, 2006.
- 517 Bartlett, P. A., and Verseghy, D. L.: Modified treatment of intercepted snow improves the
518 simulated forest albedo in the Canadian land surface scheme, *Hydrol. Process.*, 29, 3208–3226,
519 <https://doi.org/10.1002/hyp.10431>, 2015.
- 520 Betts, R. A.: Biogeophysical impacts of land use on present-day climate: near-surface
521 temperature change and radiative forcing, *Atmos. Sci. Lett.*, 2, 39–51,
522 <https://doi.org/10.1006/asle.2001.0037>, 2001.
- 523 Bjorkman, A. D., Myers-Smith, I. H., Elmendorf, S.C. et al.: Plant functional trait change across
524 a warming tundra biome, *Nature*, 562, 57–62, <https://doi.org/10.1038/s41586-018-0563-7>, 2018.
- 525 Bonan, G. B., Levis, S., Kergoat, L., Oleson, K. W.: Landscapes as patches of plant functional
526 types: An integrating concept for climate and ecosystem models, *Global Biogeochem. Cycles*,
527 16(2), 1021, <https://doi.org/10.1029/2000GB001360>, 2002.
- 528 Bontemps, S., Herold, M., Kooistra, L., van Groenestijn, A., Hartley, A., Arino, O., Moreau, I.,
529 Defourny, P.: Revisiting land cover observation to address the needs of the climate modeling
530 community, *Biogeosciences*, 9, 2145–2157, <https://doi.org/10.5194/bg-9-2145-2012>, 2012.



531 Di Gregorio, A.: Land Cover Classification System – Classification concepts and user manual
532 for Software version 2, FAO Environment and Natural Resources Service Series, No. 8, Rome,
533 208pp, 2005.

534 ESA: Land Cover CCI Product User Guide Version 2. Tech. Rep., Available at:
535 maps.elie.ucl.ac.be/CCI/viewer/download/ESACCI-LC-Ph2-PUGv2_2.0.pdf, 2017.

536 Fox, D.L., Pau, S., Taylor, L., Strömberg, C. A. E., Osborne, C.P., Bradshaw, C., Conn, S.,
537 Beerling, D.J., and Still, C.J.: climatic Controls on C4 Grassland Distributions During the
538 Neogene: A Model-Data Comparison, *Frontiers in Ecology and Evolution*, 6,
539 <https://doi.org/10.3389/fevo.2018.00147>, 2018.

540 Friedlingstein, P., O'Sullivan, M., Jones, M. W., Andrew, R. M., Hauck, J., Olsen, A., Peters, G.
541 P., Peters, W., Pongratz, J., Sitch, S., Le Quéré, C., Canadell, J. G., Ciais, P., Jackson, R. B.,
542 Alin, S., Aragão, L. E. O. C., Arneeth, A., Arora, V., Bates, N. R., Becker, M., Benoit-Cattin, A.,
543 Bittig, H. C., Bopp, L., Bultan, S., Chandra, N., Chevallier, F., Chini, L. P., Evans, W., Florentie,
544 L., Forster, P. M., Gasser, T., Gehlen, M., Gilfillan, D., Gkritzalis, T., Gregor, L., Gruber, N.,
545 Harris, I., Hartung, K., Haverd, V., Houghton, R. A., Ilyina, T., Jain, A. K., Joetzjer, E., Kadono,
546 K., Kato, E., Kitidis, V., Korsbakken, J. I., Landschützer, P., Lefèvre, N., Lenton, A., Lienert, S.,
547 Liu, Z., Lombardozi, D., Marland, G., Metzl, N., Munro, D. R., Nabel, J. E. M. S., Nakaoka, S.-
548 I., Niwa, Y., O'Brien, K., Ono, T., Palmer, P. I., Pierrot, D., Poulter, B., Resplandy, L.,
549 Robertson, E., Rödenbeck, C., Schwinger, J., Séférian, R., Skjelvan, I., Smith, A. J. P., Sutton,
550 A. J., Tanhua, T., Tans, P. P., Tian, H., Tilbrook, B., van der Werf, G., Vuichard, N., Walker, A.
551 P., Wanninkhof, R., Watson, A. J., Willis, D., Wiltshire, A. J., Yuan, W., Yue, X., and Zaehle,
552 S.: Global Carbon Budget 2020, *Earth Syst. Sci. Data*, 12, 3269–3340,
553 <https://doi.org/10.5194/essd-12-3269-2020>, 2020.



554 Fritz, S., See, L., McCallum, I., Schill, C., Obersteiner, M., van der Velde, M., Boettcher, H.,
555 Havlík, P., Achard, F.: Highlighting continued uncertainty in global land cover maps for the user
556 community, *Environ. Res. Lett.*, 6, 44005. <https://dx.doi.org/10.1088/1748-9326/6/4/044005>,
557 2011.

558 Georgievski, G., Hagemann, S.: Characterizing uncertainties in the ESA-CCI land cover map of
559 the epoch 2010 and their impacts on MPI-ESM climate simulations, *Theor. Appl. Climatol.*, 137,
560 1587–1603, <https://doi.org/10.1007/s00704-018-2675-2>, 2019.

561 Gillis, M. D., Omule, A. Y., and Brierley, T.: Monitoring Canada's forests: The National Forest
562 Inventory, *The Forestry Chronicle*, 81(2): 214-221, <https://doi.org/10.5558/tfc81214-2>, 2005.

563 Harris, I. C.: CRU JRA v2.1: A forcings dataset of gridded land surface blend of Climatic
564 Research Unit (CRU) and Japanese reanalysis (JRA) data; Jan. 1901 - Dec. 2019, Centre for
565 Environmental Data Analysis, University of East Anglia Climatic Research Unit,
566 <https://catalogue.ceda.ac.uk/uuid/10d2c73e5a7d46f4ada08b0a26302ef7>, 2020

567 Hartley, A. J., MacBean, N., Georgievski, G., Bontemps, S.: Uncertainty in plant functional type
568 distributions and its impact on land surface models, *Remote Sens. Environ.*, 203, 71–89,
569 <https://doi.org/10.1016/j.rse.2017.07.037>, 2017.

570 Harder, P., Warren D. H., and Pomeroy, J. W.: Modeling the Snowpack Energy Balance during
571 Melt under Exposed Crop Stubble, *J. Hydrometeorol.*, 19, 1191-1214, [https://DOI:](https://DOI:10.1175/JHM-D-18-0039.1)
572 [10.1175/JHM-D-18-0039.1](https://DOI:10.1175/JHM-D-18-0039.1), 2018.



- 573 Hansen, M. C., DeFries, R. S., Townshend, J. R. G., Sohlberg, R., Dimiceli, C., Carroll, M. L.:
574 Towards an operational MODIS continuous field of percent tree cover algorithm: Examples us-
575 ing AVHRR and MODIS data, *Remote Sens. Environ.*, 83, 303–319, 2002.
- 576 Hansen, M. C., Stehman, S. V., Potapov, P. V.: Quantification of global gross forest cover loss,
577 *P. Natl. Acad. Sci.*, 107, 8650–8655, <https://doi.org/10.1073/pnas.0912668107>, 2010.
- 578 Hansen, M. C., Potapov, P. V., Moore, R., Hancher, M., Turubanova, S. A., Tyukavina, A.,
579 Thau, D., Stehman, S. V., Goetz, S. J., Loveland, T.R., et al.: Highresolution global maps of
580 21st-century forest cover change, *Science*, 342, 850–853, <https://DOI: 10.1126/science.1244693>,
581 2013.
- 582 Hengl, T., Mendes de Jesus, J., Heuvelink, G. B. M., Ruiperez Gonzalez, M., Kilibarda, M.,
583 Blagotić, A., Shangguan, W., Wright, M. N., Geng, X., Bauer-Marschallinger, B., Guevara, M.
584 A., Vargas, R., MacMillan, R. A., Batjes, N. H., Leenaars, J. G. B., Ribeiro, E., Wheeler, I.,
585 Mantel, S., and Kempen, B.: SoilGrids250m: Global gridded soil information based on machine
586 learning, *PLOS ONE*, 12, 1–40, <https://doi.org/10.1371/journal.pone.0169748>, 2017
- 587 Hermosilla, T., Wulder, M. A., White, J. C., Coops, N. C., Hobart, G.W.: Disturbance informed
588 annual land cover classification maps of Canada’s forested ecosystems for a 29-year Landsat
589 time series. *Can. J. Remote. Sens.*, 44, 67–87. <https://doi.org/10.1080/07038992.2018.1437719>,
590 2018.
- 591 Herold, M., Mayaux, P., Woodcock, C. E., BaESACCI, A., Schullius, C.: Some Challenges
592 in Global Land Cover Mapping: An Assessment of Agreement and Accuracy in Existing 1 km
593 Datasets, *Remote Sens. Environ.*, 112, 2538–2556, 2008.



- 594 Hopkinson, C., Chasmer, L., Lim, K., Treitz, P., and Creed, I.: Towards a universal lidar canopy
595 height indicator, *Canadian Journal of Remote Sensing*, 32 (2), 139-152, [https://doi:](https://doi:10.5589/m06-006)
596 10.5589/m06-006, 2006.
- 597 Latifovic, R., Pouliot, D., and Olthof, I.: Circa 2010 Land Cover of Canada: Local Optimization
598 Methodology and Product Development, *Remote Sens.*, 9, 1098,
599 <https://doi.org/10.3390/rs9111098>, 2017.
- 600 Latifovic, R., and Olthof, I.: Accuracy assessment using sub-pixel fractional error matrices of
601 global land cover products derived from satellite data, *Remote Sens. Environ.*, 90, 153–165,
602 2004.
- 603 Macander, M. J., Frost, G. V., Nelson, P. R., and Swingley, C. S.: ABoVE: Tundra Plant
604 Functional Type Continuous-Cover, North Slope, Alaska, 2010-2015, ORNL DAAC, Oak
605 Ridge, Tennessee, USA, <https://doi.org/10.3334/ORNLDAAC/1830>, 2020.
- 606 Mayaux, P., Eva, H., Gallego, J., Strahler, A. H., Herold, M., Agrawal, S., Naumov, S., De
607 Miranda, E. E., Di Bella, C. M., Ordoyne, C., et al.: Validation of the Global Land Cover 2000
608 Map. *IEEE Trans. Geosci. Remote Sens.*, 44, 1728–1739, 2006.
- 609 Melton, J. R. and Arora, V. K.: Competition between plant functional types in the Canadian
610 Terrestrial Ecosystem Model (CTEM) v. 2.0, *Geosci Model Dev*, 9, 323–361,
611 <https://doi.org/10.5194/gmd-9-323-2016>, 2016.
- 612 Melton, J. R., Arora, V. K., Wisernig-Cojoc, E., Seiler, C., Fortier, M., Chan, E., and
613 Teckentrup, L.: CLASSIC v1.0: the open-source community successor to the Canadian Land
614 Surface Scheme (CLASS) and the Canadian Terrestrial Ecosystem Model (CTEM) – Part 1:



- 615 Model framework and site-level performance, *Geosci. Model Dev. Discuss.*, 1–40,
616 <https://doi.org/10.5194/gmd-2019-329>, 2019.
- 617 Moody, E. G., King, M. D., Schaaf, C. B., Hall, D. K., Platnick, S.: Northern Hemisphere five-
618 year average (2000–2004) spectral albedos of surfaces in the presence of snow: statistics
619 computed from Terra MODIS land products, *Remote Sens. Environ.*, 111, 337–345, [https://DOI:](https://doi.org/10.1016/j.rse.2007.03.026)
620 [10.1016/j.rse.2007.03.026](https://doi.org/10.1016/j.rse.2007.03.026), 2007.
- 621 Ottlé, C., Lescure, J., Maignan, F., Poulter, B., Wang, T., and Delbart, N.: Use of various remote
622 sensing land cover products for plant functional type mapping over Siberia, *Earth Syst. Sci. Data*,
623 5, 331–348, [https://doi:10.5194/essd-5-331-2013](https://doi.org/10.5194/essd-5-331-2013), 2013.
- 624 Pielke, R. A., Avissar, R., Raupach, M., Dolman, A. J., Zeng, X., Denning, S.: Interactions
625 between the atmosphere and terrestrial ecosystem: influence on weather and climate, *Global*
626 *Change Biol.*, 4, 461–475, 1998.
- 627 Poulter, B., Ciais, P., Hodson, E., Lischke, H., Maignan, F., Plummer, S., and Zimmermann, N.
628 E.: Plant functional type mapping for earth system models, *Geosci. Model Dev.*, 4, 993–1010,
629 <https://doi.org/10.5194/gmd-4-993-2011>, 2011.
- 630 Poulter, B., MacBean, N., Hartley, A., Khlystova, I., Arino, O., Betts, R., Bontemps, S.,
631 Boettcher, M., Brockmann, C., Defourny, P., Hagemann, S., Herold, M., Kirches, G., Lamarche,
632 C., Lederer, D., Ottlé, C., Peters, M., and Peylin, P.: Plant functional type classification for earth
633 system models: results from the European Space Agency's Land Cover Climate Change
634 Initiative, *Geosci. Model Dev.*, 8, 2315–2328, <https://doi.org/10.5194/gmd-8-2315-2015>, 2015.



- 635 Schaaf, C. B., et al.: First operational BRDF, albedo nadir reflectance products from MODIS,
636 Remote Sens. Environ., 83, 135–148, [https://doi.org/10.1016/S0034-4257\(02\)00091-3](https://doi.org/10.1016/S0034-4257(02)00091-3), 2002.
- 637 Shangguan, W., Hengl, T., Mendes de Jesus, J., Yuan, H., and Dai, Y.: Mapping the global depth
638 to bedrock for land surface modeling, J. Adv. Model. Earth Syst., 9, 65–88,
639 <https://doi.org/10.1002/2016MS000686>, 2017.
- 640 Smith, T. M., Shugart, H. H., and Woodward, F. I.: Plant functional types: their relevance to
641 ecosystem properties and global change, Cambridge University Press, New York, 369 pp., 1997.
- 642 Sterling, S. M., Ducharne, A., Polcher, J.: The impact of global land-cover change on the
643 terrestrial water cycle. Nat. Clim. Chang., 3, 385–390. <http://dx.doi.org/10.1038/nclimate1690>,
644 2013.
- 645 Still, C. J., and Berry, J. A.: Global distribution of C3 and C4 vegetation: Carbon cycle
646 implications, Global Biogeochem. Cycles, 17(1), 1006, doi:10.1029/2001GB001807, 2003.
- 647 Sun, W., Liang, S., Xu, G., Fang, H., and Dickinson, R. E.: Mapping plant functional types from
648 MODIS data using multisource evidential reasoning, Remote Sens. Environ., 112, 1010–1024,
649 <https://doi.org/10.1016/j.rse.2007.07.022>, 2008.
- 650 Swart, N. C., Cole, J. N. S., Kharin, V. V., Lazare, M., Scinocca, J. F., Gillett, N. P., Anstey, J.,
651 Arora, V., Christian, J. R., Hanna, S., Jiao, Y., Lee, W. G., Majaess, F., Saenko, O. A., Seiler, C.,
652 Seinen, C., Shao, A., Sigmond, M., Solheim, L., von Salzen, K., Yang, D., and Winter, B.: The
653 Canadian Earth System Model version 5 (CanESM5.0.3), Geosci. Model Dev., 12, 4823–4873,
654 <https://doi.org/10.5194/gmd-12-4823-2019>, 2019.



- 655 Tsendbazar, N. E., de Bruin, S., Mora, B., Schouten, L., Herold, M.: Comparative assessment of
656 thematic accuracy of GLC maps for specific applications using existing reference data, *Int. J.*
657 *Appl. Earth Obs. Geoinf.*, 44, 124–135, <https://doi.org/10.1016/j.jag.2015.08.009>, 2016.
- 658 Wulder, M.A., Nelson, T.: EOSD land cover classification legend report: Version 2, Natural
659 Resources Canada, Canadian Forest Service, Pacific Forestry Centre, Victoria, British Columbia,
660 Cana-da, 13 January 2003, 83pp., Available from
661 http://www.pfc.forestry.ca/eosd/cover/EOSD_Legend_Report-v2.pdf.
- 662 Wulder, M.A., Bater, C.W., Coops, N.C., Hilker, T., and White, J.C.: The role of LiDAR in
663 sustainable forest management, *The Forestry Chronicle*, 84 (6), 807- 826, 2008.
- 664 Wulder, M.A., White, J.C., Bater, C.W., Coops, N.C., Hopkinson, C., Chen, G.: Lidar plots - a
665 new large-area data collection option: context, concepts, and case study. *Can. J. Remote. Sens.*
666 38, 600–618, <https://doi.org/10.5589/m12-049>, 2012.
- 667 Verseghe, D. L.: CLASS—A Canadian Land Surface Scheme for GCMs, I. Soil model, *Int. J.*
668 *Climatol.*, 11, 111–133, <https://doi.org/10.1002/joc.3370110202>, 1991.
- 669 Verseghe, D., McFarlane, N., and Lazare, M.: Class - A Canadian land surface scheme for
670 GCMs. II: Vegetation model and coupled runs, *Int. J. Climatol.*, 13, 347–370,
671 <https://doi:10.1002/joc.3370130402>, 1993.
- 672 Wang, A., Price, D. T., Arora, V. K.: Estimating changes in global vegetation cover (1850-2100)
673 for use in climate models, *Global Biogeochem. Cy.*, 20, <https://doi:10.1029/2005GB002514>,
674 2006.



675 Wang, L., MacKay, M., Brown, R., Bartlett, P., Harvey, R., and Langlois, A.: Application of
676 satellite data for evaluating the cold climate performance of the Canadian Regional Climate
677 model over Québec, Canada, *J. Hydrometeorol.*, 15, 614–630, [https://doi.org/10.1175/JHM-D-](https://doi.org/10.1175/JHM-D-13-086.1)
678 13-086.1, 2014.

679 Wang, L., Cole, J. N.S., Bartlett, P., Verseghy, D., Derksen, C., Brown, R., von Salzen, K.:
680 Investigating the spread in surface albedo for snow-covered forests in CMIP5 models, *J.*
681 *Geophys. Res. Atmos.*, 121, 1104–1119, <https://doi:10.1002/2015JD023824>, 2016.

682 Wang, L., Bartlett, P., Chan, E., and Xiao, M.: Mapping of Plant Functional Type from Satellite-
683 Derived Land Cover Datasets for Climate Models, In Proceedings of the 2018 IEEE International
684 Geoscience and Remote Sensing Symposium, Valencia, Spain, 22–27, [https://DOI:](https://DOI:10.1109/IGARSS.2018.8518046)
685 10.1109/IGARSS.2018.8518046, 2018.

686 Wang, L., Bartlett, P., Pouliot, D., Chan, E., Lamarche, C., Wulder, M.A., Defourny, P., Brady,
687 M.: Comparison and Assessment of Regional and Global Land Cover Datasets for Use in
688 CLASS over Canada, *Remote Sens.*, 11(19), 2286; <https://doi.org/10.3390/rs11192286>, 2019.

689
690
691
692
693
694
695
696
697



698 Table 1. Cross-walking table for mapping the 30 m Hybrid land cover map to CLASSIC PFTs in Canada.

ID	Map description	1 NLE	2 NLD	3 BLE	4+5 BCD BDD	6+7 C3C C4C	8+9 C3G C4G	Urban	Lake	Bare
2	Sub-polar taiga needleleaf forest	0.20					0.60			0.20
11	Sub-polar or polar shrubland-lichen- moss						0.65			0.35
12	Sub-polar or polar grassland-lichen- moss						0.45			0.55
13	Sub-polar or polar barren-lichen- moss						0.10			0.90
15	Cropland					1.0				
16	Barren lands									1.0
17	Urban							1.0		
20	Water								1.0	
31	Snow_ice									1.0
32	Rock_rubble									1.0
50	Shrubland				0.20		0.60			0.20
80	Wetland				0.05		0.85			0.10
81	Wetland-treed	0.55			0.05		0.35			0.05
100	Herbs						0.80			0.20
210	Coniferous	1.0								
220	Broadleaf				1.0					
230	Mixedwood	0.50			0.50					

699

700

701

702

703

704

705

706

707

708

709



710 Table 2. The sub-pixel fractional tree cover fraction for ESA-CCI LC classes (with forest cover) based on
 711 the Hansen TCF dataset in Canada. Ratios of TCF between the main class and the closed class, and
 712 between the open class and the closed class are also included.

ESA-CCI class	ESA-CCI class description	Tree cover Fraction (%)	Ratio of TCF relative to closed class
30	Mosaic cropland (>50%) / natural vegetation (<50%)	13.7	
40	Mosaic natural vegetation (>50%) / cropland (<50%)	45	
60	Tree cover broadleaved deciduous closed to open (>15%)	68.5	0.8
61	Tree cover broadleaved deciduous closed (>40%)	86.7	1
62	Tree cover broadleaved deciduous open (15-40%)	37.4	0.43
70	Tree cover needleleaf evergreen closed to open (>15%)	39.3	0.6
71	Tree cover needleleaf evergreen, closed (>40%)	61.7	1
72	Tree cover needleleaf evergreen open (15-40%)	23.2	0.38
90	Tree cover Mixed	80.9	
100	Mosaic tree and shrub (>50%) / herbaceous cover (<50%)	37.3	
110	Mosaic herbaceous cover (>50%) / tree and shrub (<50%)	19.6	
120	Shrubland	28.1	
150	Sparse vegetation (tree shrub herbaceous cover) (< 15%)	4	
160	Tree cover, flooded fresh/brackish	43	
180	Shrub or herbaceous cover, flooded	26.9	

713

714

715

716

717

718

719

720



721 Table 3. The sub-pixel fractional composition for ESA-CCI classes (columns, homogenous ESA-CCI
 722 pixels) based on the Hybrid LC map (rows) for dominant LC classes in Canada. The fractions for NLE
 723 and BCD are computed based on equation (1).

Hybrid/ ESACCI Class	Hybrid description	30	40	60	70	71	90	100	120	130	140	150	160	180
2	Sub-polar taiga needleleaf forest				0.02			0.01		0.01				
11	Sub-polar/polar shrubland-lichen- moss										0.01	0.05		
12	Sub-polar/polar grassland-lichen- moss				0.04				0.03	0.01	0.24	0.27	0.03	0.04
13	Sub-polar/polar barren-lichen-moss				0.02			0.01	0.02	0.01	0.34	0.09		0.02
15	Cropland	0.92	0.37	0.02						0.1				
16	Barren lands									0.01	0.15	0.17		
50	Shrubland	0.01	0.07	0.06	0.13	0.05	0.04	0.32	0.46	0.09	0.14	0.25	0.06	
80	Wetland		0.03	0.08	0.2	0.05	0.03	0.27	0.2	0.02	0.06	0.09	0.37	0.75
81	Wetland treed		0.01	0.01	0.17	0.07	0.03	0.11	0.12				0.43	0.15
100	Herbs	0.06	0.27	0.08	0.02		0.02	0.06	0.09	0.72	0.01	0.03	0.01	0.01
210	Coniferous		0.01	0.02	0.29	0.72	0.07	0.04	0.03		0.01	0.02	0.06	
220	Broadleaf	0.01	0.13	0.57	0.02	0.01	0.28	0.07	0.01	0.01			0.01	
230	Mixedwood		0.1	0.14	0.09	0.07	0.52	0.12	0.03				0.02	
NLE	Needleleaf evergreen		0.07	0.09	0.44	0.8	0.32	0.19	0.16	0.01	0.02	0.05	0.31	0.08
BCD	Broadleaf cold deciduous	0.01	0.19	0.66	0.09	0.06	0.57	0.18	0.09	0.02	0.02	0.03	0.05	0.03

724

725

726

727

728

729

730



731 Table 4. Cross-walking table for mapping ESA-CCI land cover dataset to CLASSIC PFTs.

ID	ESA-CCI class description	1 NLE	2 NLD	3 BLE	4+5 BCD BDD	6+7 C3C C4C	8+9 C3G C4G	Urban	Lake	Ocean	Bare
10	Cropland, rainfed (CR)					0.80	0.20				
11	CR Herbaceous cover					0.90	0.10				
12	CR Tree or shrub cover				0.60		0.30				0.10
20	Cropland, irrigated or post-flood				0.05	0.85	0.10				
30	Mosaic cropland (>50%) / natural vegetation (tree, shrub, herb)	0.05			0.15	0.60	0.20				
40	Mosaic natural vegetation (tree, shrub, herb) >50% / crop	0.10			0.20	0.40	0.30				
50	Tree cover broadleaved evergreen closed to open			0.95	0.05		0.0				
60	Tree cover broadleaved deciduous closed to open				0.70		0.25				0.05
61	Tree cover broadleaved deciduous closed				0.90		0.10				
62	Tree cover broadleaved deciduous open				0.40		0.40				0.20
70	Tree cover needleleaf evergreen closed to open	0.85			0.05		0.10				
71	Tree cover needleleaf evergreen, closed	0.85			0.05		0.10				
72	Tree cover needleleaf evergreen open	0.35			0.10		0.40				0.15
73	Replace LC70 in Canada	0.45			0.10		0.30				0.15
80	Tree cover needleleaf deciduous closed to open	0.05	0.40		0.10		0.35				0.10
81	Tree cover needleleaf deciduous closed	0.05	0.80		0.05		0.15				
82	Tree cover needleleaf deciduous open	0.05	0.30		0.10		0.45				0.15
90	Tree cover Mixed	0.25	0.05		0.60		0.10				
100	Mosaic tree and shrub (>50%) / herbaceous cover (<50%)	0.15	0.05		0.20		0.45				0.15
110	Mosaic herbaceous cover (>50%) / tree and shrub (<50%)	0.05	0.05		0.10		0.65				0.15
120	Shrubland				0.30		0.45				0.25
121	Shrubland evergreen	0.15		0.15			0.45				0.25
122	Shrubland deciduous				0.30		0.45				0.25

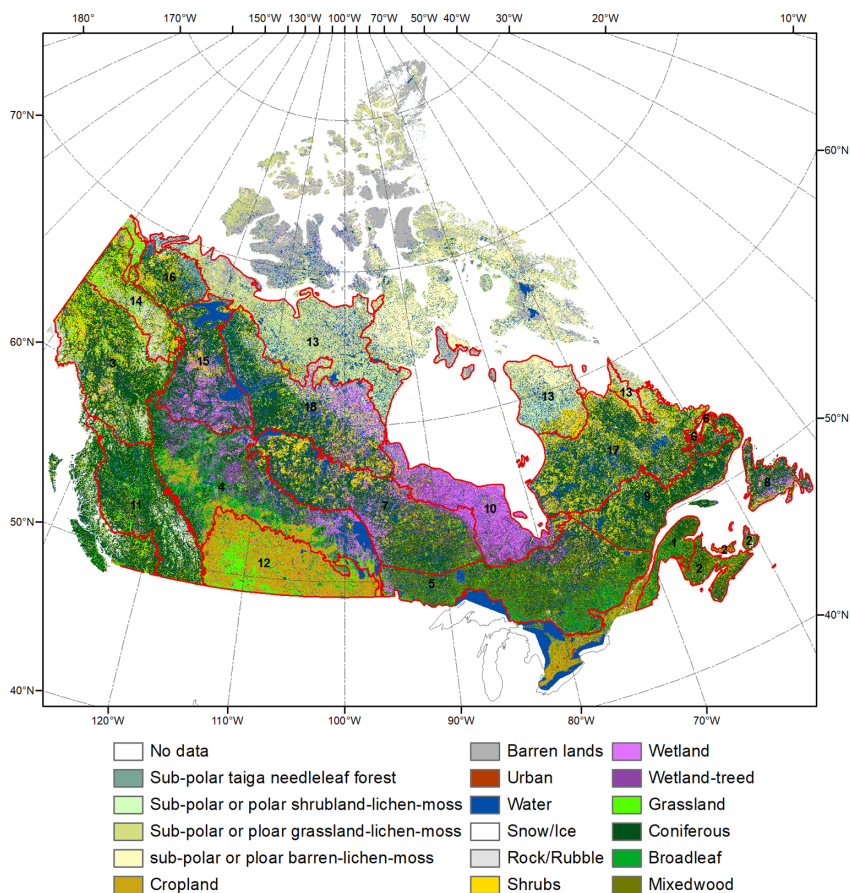


130	Grassland					0.70				0.30
140	Lichens and mosses					0.20				0.80
150	Sparse vegetation (tree shrub herbaceous cover) (<15%)				0.05	0.35				0.60
151	Sparse tree (<15%)				0.05	0.35				0.60
152	Sparse shrub (<15%)					0.30				0.70
153	Sparse herbaceous cover (<15%)					0.30				0.70
160	Tree cover, flooded fresh/brackish	0.30			0.10	0.45		0.1		0.05
170	Tree cover, flooded saline water	0.30			0.10	0.40			0.1	0.10
180	Shrub or herbaceous cover, flooded	0.10			0.05	0.45		0.15	0.15	0.10
190	Urban areas	0.025			0.025	0.15	0.75	0.05		
200	Bare areas									1.0
201	Consolidated bare areas									1.0
202	Unconsolidated bare areas									1.0
210	Water bodies							1.0		
220	Permanent snow and ice									1.0

732
 733
 734
 735
 736
 737
 738
 739
 740
 741
 742
 743
 744
 745
 746



747



748

749 Figure 1. The Hybrid map of Canada based on VLCE and NALCMS land cover maps for 2010. The red
 750 polygons represent 18 ecozones used in this study.

751

752

753

754

755

756

757



758

759

760

761

762

763

764

765

766

767

768

769

770 Figure 2. Schematic flow chart of the process for creating the cross-walking table for ESA-CCI LC
771 dataset.

772

773

774

775

776

777

778

779

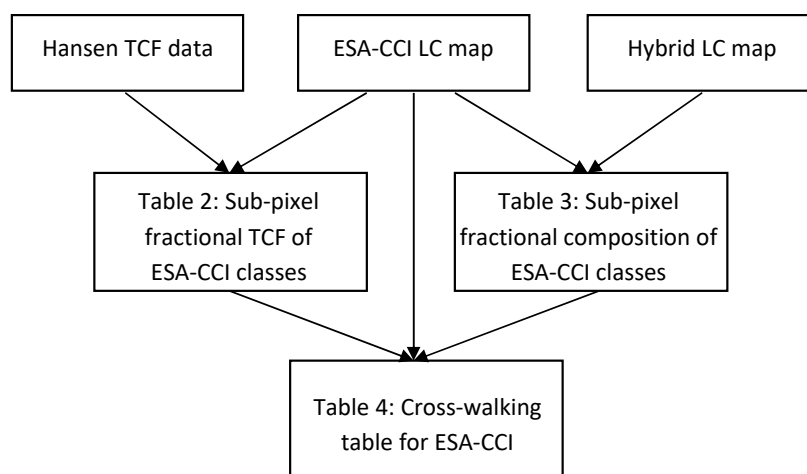
780

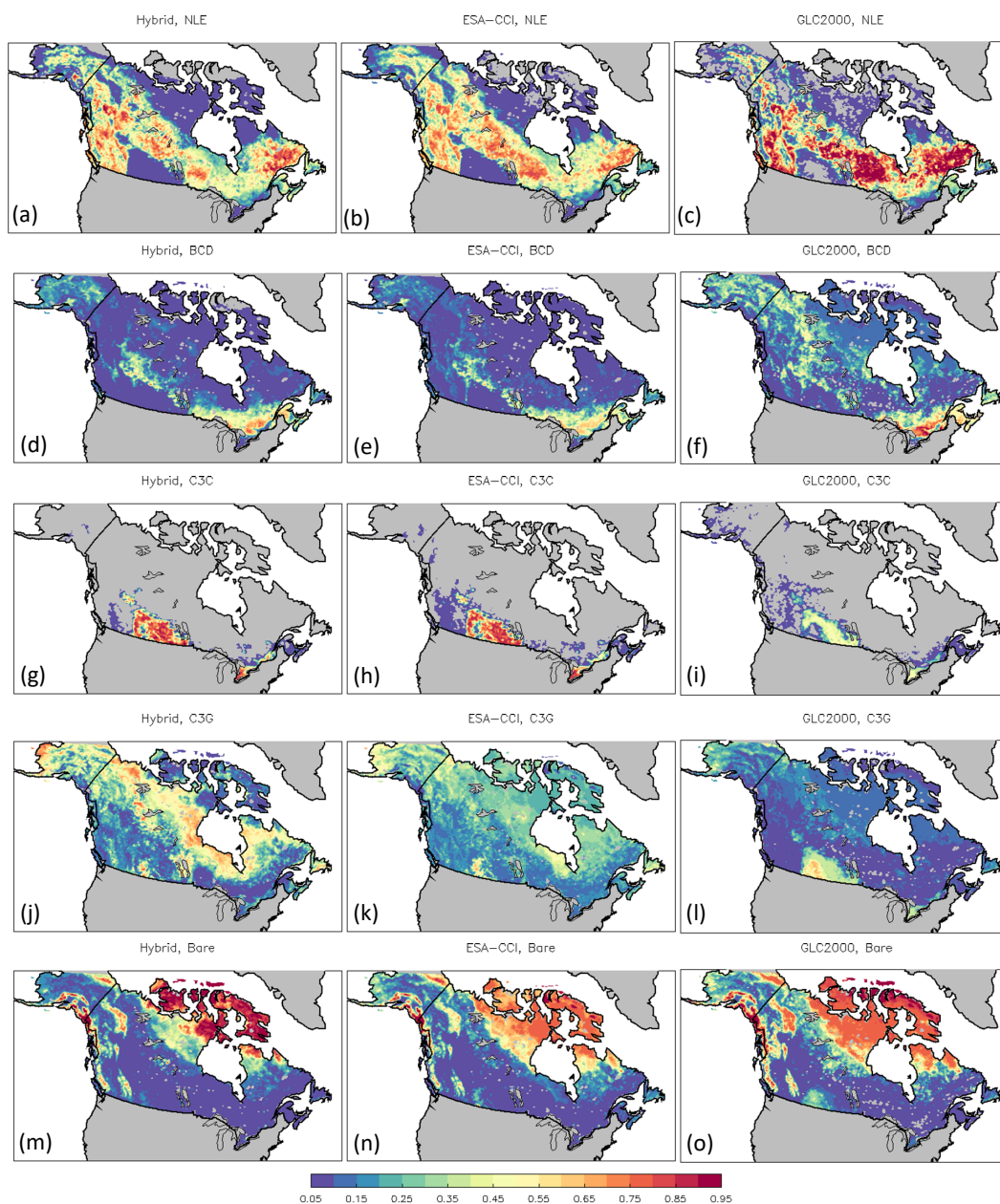
781

782

783

784

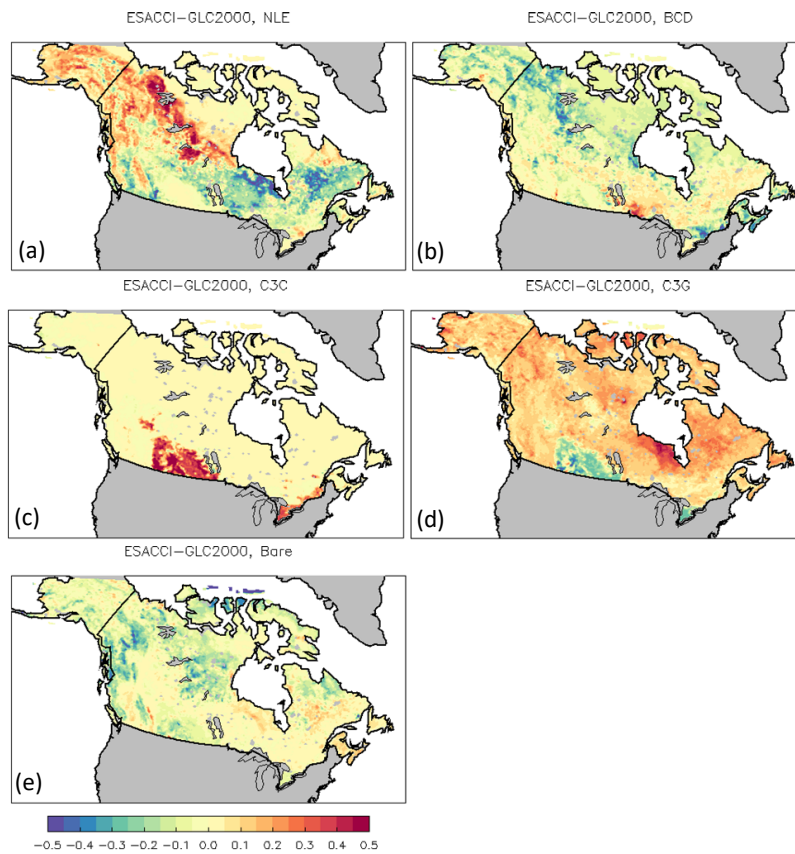




785

786 Figure 3. The spatial distribution of CLASSIC PFTs based on the Hybrid (left), ESA-CCI (middle), and
787 GLC2000 (right) land cover datasets respectively. The maps for C4C and C4G are not shown for their
788 fractions are negligible in Canada. The last panel shows fractions for bare ground from the three datasets.

789



805

806 Figure 4. The difference in PFTs based on ESA-CCI and GLC2000 datasets for selected PFTs (a) NLE,
807 (b) BCD, (c) C3C, (d) C3G, and (e) Bare.

808

809

810

811

812

813

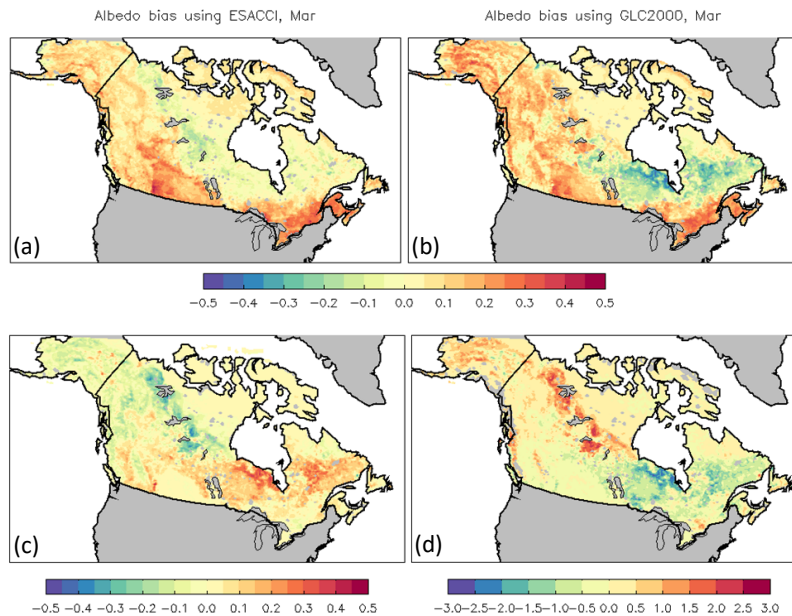
814

815

816



817



829

830 Figure 5. Surface albedo bias in CLASSIC simulations using PFT distributions based on (a) ESA-CI,
831 and (b) GLC2000 land cover products. Panels (c) and (d) show the difference in simulated surface albedo
832 (c) and leaf area index (d) between the two simulations.

833