Mapping of ESA-CCI land cover data to plant functional types for use in the CLASSIC land model

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1 Abstract

2 Plant functional types (PFTs) are used to represent vegetation distribution in land surface models (LSMs). Previous studies have shown large differences in the geographical distribution of PFTs 3 currently used in various LSMs, which may arise from the differences in the underlying land 4 5 cover products but also the methods used to map or reclassify land cover data to the PFTs that a 6 given LSM represents. There are large uncertainties associated with existing PFT mapping methods since they are largely based on expert judgment and therefore are subjective. In this 7 study, we propose a new approach to inform the mapping or the cross-walking process using 8 9 analyses from sub-pixel fractional error matrices, which allows for a quantitative assessment of 10 the fractional composition of the land cover categories in a dataset. We use the Climate Change 11 Initiative (CCI) land cover product produced by the European Space Agency (ESA). Previous work has shown that compared to fine-resolution maps over Canada, the ESA-CCI product 12 13 provides an improved land cover distribution compared to that from the GLC2000 dataset currently used in the CLASSIC (Canadian Land Surface Scheme Including Biogeochemical 14 Cycles) model. A tree cover fraction dataset and a fine-resolution land cover map over Canada 15 are used to compute the sub-pixel fractional composition of the land cover classes in ESA-CCI, 16 17 which is then used to create a cross-walking table for mapping the ESA-CCI land cover categories to nine PFTs represented in the CLASSIC model. There are large differences between 18 19 the new PFT distributions and those currently used in the model. Offline simulations performed with the CLASSIC model using the ESA-CCI based PFTs show improved winter albedo 20 21 compared to that based on the GLC2000 dataset. This emphasizes the importance of accurate 22 representation of vegetation distribution for realistic simulation of surface albedo in LSMs. Results in this study suggest that the sub-pixel fractional composition analyses are an effective 23

way to reduce uncertainties in the PFT mapping process and therefore, to some extent, objectifythe otherwise subjective process.

26

27 **1. Introduction**

28 Land cover is a critical component of the earth system that affects the exchange of energy, water, 29 and carbon between the land surface and the atmosphere (Pielke et al., 1998; Sterling et al., 2013). Accurate representation of global land cover (LC) is important for land surface models 30 31 (LSMs) which provide the lower boundary conditions to the atmosphere in numerical weather 32 forecasting, climate, and earth system models (ESMs). Plant functional types (PFTs) are groups of plant species that share similar structural, phenological, and physiological traits, and have 33 been commonly used in LSMs to represent vegetation distribution. This simplification has 34 35 allowed the simulation of structural attributes of vegetation dynamically within ESMs (Arora & 36 Boer, 2010; Bonan et al., 2003; Krinner et al., 2005). In order to improve the representation of 37 ecosystem ecology and vegetation demographic processes within ESMs, both species-based and 38 trait-based models have been attempted in LSMs (Fisher et al., 2018; Zakharova et al., 2019). However, these individual-based models are computionally too expensive to model 39 biogeochemical processes, especially photosynthesis and the carbon cycle at the global scale 40 (Bonan et al., 2002; Smith et al., 1997; 2001). As a compromise, "cohort-based" models have 41 been developed where individual plants with similar properties (size, age, functional type) are 42 grouped together and have been implemented in some ESMs (Fisher et al., 2018). Though there 43 are limitations in PFTs-based models (Scheiter et al., 2013; Zakharova et al., 2019), PFTs are 44 commonly used in LSMs that participate routinely in the Global Carbon Project (Friedlingstein 45

et al., 2020) and in ESMs that participate in the Coupled Models Intercomparison Project (CMIP,
Wang et al., 2016).

There are three approaches for modeling PFTs: (1) static, where the fractional coverage of PFTs 48 is prescribed and does not vary through time; (2) forced, where the fractional coverage of PFTs 49 50 is still prescribed but vary through time based on scenarios of land cover/land-use change; and (3) dynamic, where the fractional coverage of PFTs is simulated dynamically with competition 51 for available space and resources between PFTs (Fisher et al., 2018; Melton and Arora, 2016). 52 The number and type of PFTs used in each LSM differ. Global land cover datasets are typically 53 54 used to derive the fractional coverage of PFTs for use in LSMs. However, large differences exist 55 in both the fractional coverage and the geographical distribution of PFTs, which are caused by differences in the LC datasets themselves but also due to the methods used to map LC datasets to 56 the PFTs represented in various models (Fritz et al, 2011; Hartley et al., 2017; Ottle et al., 2013; 57 58 Wang et al., 2016).

59 Since different PFTs are characterized by different physical and biogeochemical processes and parameter values, the spatial distribution and fractional cover of PFTs constitute one of the 60 61 important geophysical fields that are required for realistic simulation of carbon, water, and energy budgets in LSMs (Arora and Boer, 2010; Betts, 2001). For example, the surface 62 roughness for short or tall vegetation is very different, which affects simulated turbulent 63 exchanges. The surface albedos for needleleaf evergreen trees, broadleaf deciduous trees, and 64 grasslands are also very different, especially during winter when deciduous trees are leafless and 65 short vegetation is largely buried by snow (Bartlett and Verseghy, 2015; Moody et al., 2007). 66 67 Wang et al. (2016) found that the bias in winter albedo in selected boreal forest regions among the CMIP5 models was largely related to biases in leaf area index (LAI) and tree cover fraction. 68

Model experiments using the MPI-ESM by Georgievski and Hagemann (2019) suggested that
uncertainties in vegetation distribution may lead to noticeable variations in near-surface climate
variables and large-scale circulation patterns.

The Canadian Land Surface Scheme Including Biogeochemical Cycles (CLASSIC) is an open-72 73 source community land model that is designed to address research questions that explore the role 74 of the land surface in the global climate system (Melton et al., 2020). It is the successor to the coupled modelling framework based on the Canadian Land Surface Scheme (CLASS; Verseghy, 75 1991; 1993) and the Canadian Terrestrial Ecosystem Model (CTEM; Arora and Boer, 2005; 76 Melton and Arora, 2016). The physics and biogeochemistry modules of CLASSIC are based on 77 78 CLASS and CTEM models, respectively. Since the development of CTEM in the early 2000s, 79 the GLC2000 LC product has been used to specify the spatial distribution of PFTs for CLASSIC when employed as the land surface component of the Canadian Earth system model developed 80 81 by Environment and Climate Change Canada (Arora et al., 2009; Wang et al., 2006). The 82 Climate Change Initiative (CCI) LC product recently produced by the European Space Agency (ESA) is available at an annual temporal resolution for the period 1992 to 2018 at 300 m spatial 83 resolution (ESA, 2017). It was produced based on broad user consultation, specifically to address 84 85 the needs of the climate modelling community (Bontemps et al., 2012). Wang et al. (2019) showed that when compared to the finer resolution maps over Canada, the 300 m ESA-CCI 86 product provides much improved LC distribution over Canada compared to that from the 1 km 87 GLC2000 dataset. 88

To map LC classes to PFTs, a cross-walking table (CW-table) is usually created to assign
fractions of each LC class to the different PFTs, such that the sum of the fractions for each class
is always one (including fractions of water and bare ground). Previous methods for creating such

CW-tables are mainly based on LC class descriptions, expert knowledge, and the spatial 92 distribution of global biomes (Ottle et al., 2013; Poulter et al., 2011; 2015; Sun and Liang, 2007; 93 Wang et al., 2006). Because LC maps only provide the types of vegetation, and each class can be 94 associated with a broad range of fractional cover of either one or more vegetation types, there are 95 large uncertainties associated with any cross-walking or reclassification process. Wang et al. 96 (2019) reclassified the 10 PFTs in the default CW-table provided in the ESA-CCI LC product 97 user manual (Table 7-2, ESA, 2017) into PFTs represented in the CLASSIC model, and 98 compared them with those based on the GLC2000 dataset. The results suggest that uncertainties 99 100 in the CW-tables were a major source of large differences in the PFT distributions. In addition, the fractional coverage of tree PFTs based on the default CW-table for the ESA-CCI LC dataset 101 was overestimated along the taiga-tundra transition zone in western Canada, which led to 102 underestimation in winter albedo in CLASSIC offline simulations driven with observation-based 103 reanalysis data (Wang et al., 2018). 104

The objective of this study is to develop a new CW-table for reclassifying the ESA-CCI LC 105 classes into PFTs represented in the CLASSIC model over the model's Canadian domain, and to 106 compare and assess the performance of CLASSIC offline simulations using the new and existing 107 PFT distributions. Given the close link between the bias in winter albedo and the vegetation 108 distribution in the models (Wang et al., 2016), our assessment of model performance focuses on 109 the simulated surface albedo during the maximum snow accumulation period (February-March 110 for the boreal forest). This simplifies our analyses by excluding the fall/spring transition periods 111 112 when biases in snow accumulation and melt timing can have a large influence on surface albedo 113 simulated by LSMs (Wang et al., 2014). In addition, we extend the CW-table for the ESA-CCI LC dataset to the global domain. A comprehensive assessment of the impact of the PFT 114

- distribution based on the new CW-table and the ESA-CCI LC dataset on the performance of the
- 116 CLASSIC model at the global scale is presented in Arora et al. (2022).

117 2. Data and the CLASSIC model

118 2.1 The Hybrid LC map over Canada

The United States Geological Survey archive of Landsat imagery has provided open and free access to georeferenced and spectrally corrected analysis-ready imagery (Wulder et al., 2012), which makes it possible to generate time series of LC maps to study LC change. Recently two of these products based on Landsat imagery were generated over Canada, including the North America Land Change Monitoring System (NALCMS) LC dataset (Latifovic et al., 2017) and the Virtual Land Cover Engine (VLCE) framework-generated LC dataset (Hermosilla et al.,

125 2018).

126 Based on the random forest algorithm and local optimization method, the Canada Centre for 127 Remote Sensing has generated the NALCMS LC maps of Canada for the years 2010 and 2015 at 128 30 m resolution using Landsat imagery (Latifovic et al., 2017). These LC products are the 129 Canadian contribution to the 30 m resolution 2010/2015 LC map of North America to the joint 130 collaborative effort by the Mexican, American, and Canadian government institutions under the NALCMS umbrella. The NALCMS LC map has 19 classes based on the United Nations Land 131 Cover Classification System (LCCS; Di Gregorio, 2005). Assessment based on reference 132 133 samples showed an overall accuracy of 76.6% for the year 2010 data (Latifovic et al., 2017), which is used in this study. 134

VLCE is an automated framework to enable change-informed annual LC mapping using time
series of Landsat surface reflectance. Temporally consistent annual LC maps representative of

| 137 | Canada's forested ecosystems from 1984 to 2012 were generated using the VLCE framework, |
|-----|--|
| 138 | characterizing LC dynamics following wildfire and harvesting events by Hermosilla et al. |
| 139 | (2018). The VLCE maps have 12 LC classes in a hierarchical classification structure following |
| 140 | that of the National Forest Inventory. Assessment based on reference samples showed an overall |
| 141 | accuracy of 70.3% for the map of the year 2005 (the year with the greatest number of reference |
| 142 | samples; Hermosilla et al., 2018). Land cover data from the year 2010 are used in this study. |
| 143 | Overall, the 19-class NALCMS product presents a more detailed LC distribution than the 12- |
| 144 | class VLCE map over Canada. For example, areas classified as "Exposed/Barren lands" in the |
| 145 | VLCE map correspond to either "Sub-polar or polar grassland-lichen-moss", "Sub-polar or polar |
| 146 | barren-lichen-moss", or "Barren lands" in the NALCMS map. Areas of cropland are not |
| 147 | separated from grassland in the VLCE map. A recent study showed that the wetland class in |
| 148 | NALCMS suffers from large uncertainty in forest cover mapping because treed-wetland was not |
| 149 | separated from herbaceous wetland in its legend (Wang et al., 2019). To take advantage of both |
| 150 | datasets, we created a hybrid product by combining them through the following steps: (1) |
| 151 | Reproject the VLCE data from its Lambert Conformal Conic projection to the same Lambert |
| 152 | Azimuthal Equal Area projection that is used for the NALCMS data; (2) Replace pixels |
| 153 | classified as "Exposed/Barren lands" and "Bryoids" in the VLCE data with the more specific LC |
| 154 | classes from the NALCMS data; (3) Replace pixels classified as "Herbs" in the VLCE data with |
| 155 | the "Cropland" class in the NALCMS data (remains "Herbs" if not classified as "Cropland" in |
| 156 | NALCMS); (4) and merge the rest of LC classes from NALCMS to the corresponding classes in |
| 157 | the VLCE data. There are a total of 17 classes in this new hybrid product, which we will |
| 158 | henceforth refer to it as the Hybrid LC dataset and is shown in Figure 1. |

2.2 The global LC products

160 The GLC2000 dataset was generated from SPOT/VEG data collected from November 1999 to December 2000 at 1 km resolution (Bartholomé and Belward, 2005). It was produced by 21 161 separate regional expert groups using an unsupervised image classification method. Based on the 162 LCCS, the regional products were merged into one global product with a generalized LCCS 163 legend of 22 classes. Assessment based on a random sampling of reference sites globally 164 estimated an overall accuracy of 68.6% for the GLC2000 product (Mayaux et al., 2006). 165 The annual ESA-CCI LC data at 300 m resolution are available for the period 1992-2018, which 166 were generated from baseline data and annual LC changes (ESA, 2017). The baseline data were 167 generated using a combination of machine learning and unsupervised image classification 168 methods from the entire archive of ENVISAT/Medium Resolution Imaging Spectrometer for the 169 period of 2003-2012. The annual LC changes were detected at 1 km resolution from the 170 Advanced Very High Resolution Radiometer time series between 1992 and 1999, SPOT/VEG 171 172 time series between 1999 and 2013, and the PROBA-V time series between 2013 and 2018. Based on the LCCS legend, the ESA-CCI LC data have 22 level 1 classes, and 15 level 2 sub-173 classes. Assessment based on the GlobCover validation database estimated an overall accuracy 174 of 71% for the ESA-CCI LC product (ESA, 2017). 175

176 **2.3 Other datasets**

Airborne Lidar has been used to monitor forests since the 1980s and is well suited to estimate
vegetation height, volume, and biomass (Hopkinson et al., 2006; Wulder et al., 2008). Vegetation
cover percentage for canopy height above 2 m from airborne Lidar data are used to estimate the
fraction of tall versus low vegetation for LC classes with a mix of woody and herbaceous
vegetation in this study. The Lidar data were collected along 34 survey flights across the boreal

forest of Canada in the summer of 2010 by the Canadian Forest Service (Wulder et al., 2012). A
25 by 25 m tessellation was generated with the approximately 400 m wide Lidar swath, with
each cell treated as an individual Lidar plot.

A tree cover fraction (TCF) dataset for 2010 is also used in this study (Hansen et al., 2013; 185 186 hereafter the Hansen TCF dataset). It was based on Landsat images at 30 m resolution. In contrast to the discrete LC classification datasets (providing a certain number of LC classes) as 187 described above, the Hansen dataset is a vegetation continuous field product (providing tree 188 cover fractions from 0 - 100%, in which the satellite spectral information was used to estimate 189 190 the TCF in each pixel using a regression tree algorithm (Hansen et al., 2002; 2010). This may 191 better represent heterogeneous areas than is possible by discrete LC classification. Tree cover is 192 defined to exist over pixels where canopy closure is observed for vegetation taller than 5 m in height. Forests are generally defined as woody vegetation taller than 3 m in the regional and 193 194 global LC datasets. The different definitions of tree heights should not result in much difference in areas with mature forests, such as most boreal forests in Canada. 195

196 Simulated surface albedo by the CLASSIC model in offline experiments is evaluated against the

197 Moderate Resolution Imaging Spectroradiometer (MODIS) (MCD43C3) broadband (0.3–5.0

198 μ m) white-sky albedo (Schaaf et al., 2002), with quality flags of 0–2 (75% or more full

inversions and 25% or fewer fill values) and solar zenith angles less than 70°. The MODIS

albedo product used in this study is at 0.05 degree resolution, and is regridded to the 0.22 degree

201 resolution used for the CLASSIC simulations (see Section 2.4.2).

202 2.4 The CLASSIC model and simulation setup

203 2.4.1 The CLASSIC model

204 CLASSIC is the successor to the coupled modelling framework based on the Canadian Land Surface Scheme (CLASS; Verseghy, 1991; 1993) and the Canadian Terrestrial Ecosystem Model 205 (CTEM; Arora and Boer, 2005; Melton and Arora, 2016). The physics and biogeochemistry 206 components of CLASSIC are based on CLASS and CTEM, respectively. 207 For the physics component, the default model's vegetation is represented in terms of the 208 209 fractional coverage of the four PFTs (needleleaf trees, broadleaf trees, crops, and grasses). The physics component represents a single snow layer with variable depth and a single vegetation 210 canopy layer. As a first-order treatment of subgrid-scale heterogeneity, each grid cell is divided 211 212 up into four sub-areas, consisting of vegetated and bare soil areas, each with and without snow cover. The visible and near-infrared albedos of each PFT/vegetation category are specified. 213 These albedos are further modified by taking into account the fraction of the ground that is seen 214 from the sky above referred to as the sky view factor (which is modelled as a function of the leaf 215 area index). The albedo of the ground that is seen from the sky above depends on if the ground is 216 snow covered or not but also on the soil moisture of the top soil layer, since wet soil is darker 217 than the dry soil. Canopy snow processes such as interception/unloading, sublimation, and melt 218 are all simulated. The aggregated visible and near-infrared albedos for the bulk canopy are 219 220 incremented using the current values weighted by the fractional coverage of the vegetation categories (Verseghy 1993). More details can be found in Appendix A. The overall surface 221 albedo of a grid cell is computed as a weighted mean using the fractional coverages for the four 222 223 sub-areas. Twenty ground layers represent the soil profile, starting with 10 layers of 0.1 m thickness. The thicknesses of the layers gradually increase to 30 m for a total ground depth of 224 over 61 m. Liquid and frozen soil moisture contents, and soil temperature, are determined 225 226 prognostically for permeable soil layers.

227 The biogeochemistry component of CLASSIC used here represents vegetation in terms of nine PFTs: Needleleaf Evergreen trees (NLE), Needleleaf Deciduous trees (NLD), Broadleaf 228 Evergreen trees (BLE), Broadleaf Cold Deciduous trees (BCD), Broadleaf Dry Deciduous trees 229 (BDD), C₃ and C₄ Crops (C3C/C4C), and C₃ and C₄ Grasses (C3G/C4G). These nine PFTs map 230 directly onto the four PFTs used by CLASSIC's physics component. When the physics and 231 232 biogeochemistry components are coupled together, as in the case of simulations carried out in this study, the structural attributes of vegetation including leaf area index, canopy mass, rooting 233 depth, and vegetation height are simulated dynamically as a function of environmental 234 235 conditions by the biogeochemistry component. When the biogeochemistry component is turned off, specified structural attributes of vegetation for use by the physics component are extracted 236 237 from look-up tables.

238 2.4.2 Simulation set up

239 Gridded meteorological data based on the Climatic Research Unit (CRU,

240 https://crudata.uea.ac.uk/cru/data/hrg/) and Japanese reanalysis (JRA) (CRUJRA) are used to

241 drive CLASSIC simulations. The data are constructed by regridding data from the JRA and

adjusting where possible to align with the CRU TS 4.05 data. The blended product from January,

1901 to December, 2020 has the 6-hourly temporal resolution of the reanalysis product but

monthly means adjusted to match the CRU data (Harris, 2020). The 6-hourly data are

disaggregated on-the-fly within CLASSIC into half-hourly data following the methodology by

246 Melton and Arora (2016) for the following seven meteorological variables that are used to force

the model: 2 m air temperature, total precipitation, specific humidity, downward solar radiation

248 flux, downward longwave radiation flux, surface pressure, and wind speed. Surface temperature,

surface pressure, specific humidity, and wind speed are linearly interpolated. Long-wave

250 radiation is uniformly distributed across a 6-hour period, and shortwave radiation is diurnally distributed over a day based on a grid cell's latitude and day of year with the maximum value 251 occurring at solar noon. Precipitation is treated following Arora (1997), where the total 6-hour 252 precipitation amount is used to determine the number of wet half hours in a 6-hour period. The 6-253 hour precipitation amount is then spread randomly, but conservatively, over the wet half-hourly 254 periods. In CLASSIC, the phase of precipitation is determined by a threshold surface air 255 temperature with three options available (Bartlett et al., 2006). The 0°C air temperature threshold 256 is used to partition precipitation into rain or snow in this study. This choice does not have a 257 258 significant impact on the simulated surface albedo in CLASSIC escpecially during the February-March months when the snow cover is near its maximum (Wang et al. 2014). 259 260 Two simulations over the 1850-2020 historical period are performed using PFTs derived from the ESA-CCI and the GLC2000 datasets respectively, which is the only difference between the 261 two simulations. Static PFTs are used in the simulations where the fractional coverage of PFTs is 262 prescribed and does not vary through time. Besides land cover and meteorological forcings, 263 CLASSIC requires globally averaged atmospheric CO₂ concentration, and geographically 264 varying time-invariant soil texture and soil permeable depth. The atmospheric CO₂ concentration 265 values are provided by the Global Carbon Project protocol 266 (https://www.globalcarbonproject.org/index.htm). The soil texture information consists of the 267 percentage of sand, clay, and organic matter and is derived from the SoilGrids250m dataset 268 (Hengl et al., 2017), and permeable soil depth is based on Shangguan et al. (2017). The 269 270 simulations are performed at a 0.22 degree rotated latitude-longitude grid over a domain 271 including Canada and part of Alaska (Fig. 3). Pre-industrial simulations that correspond to the year 1850 are required prior to doing the historical simulations so that model's carbon pools, 272

273 including leaf biomass which determines leaf area index, are spun up to near equilibrium for each land cover. The pre-industrial simulations use 1901-1920 meteorological data repeatedly 274 with atmospheric CO₂ concentration specified at its 1850 level. Each historical simulation is then 275 276 initialized from its corresponding pre-industrial simulation after it has reached equilibrium (with 277 carbon fluxes to conditions corresponding to the year 1850). For the period 1851-1900, the CRUJRA meteorological data for the first 20 years (1901-1920) are used repeatedly. For the 278 1901-2020 period the meteorological data corresponding to each actual year are used. The period 279 from 2001 to 2015 was selected for analyzing the simulated results. 280

281 **3. PFT mapping methods**

The CW-table for the ESA-CCI LC dataset is generated through a multi-step process that 282 283 combines multiple land cover maps at different spatial and categorical resolutions with ancillary data on tree cover and vegetation height (Fig. 2). This includes the following steps: (1) 284 combining two existing land cover maps (NALCMS and VLCE) to produce a harmonized 30 m 285 286 land cover (Hybrid) map with improved categorical precision (as described in Section 2.1); (2) creating a CW-table for the Hybrid land cover map through a direct mapping of classes from the 287 288 Hybrid map onto the CLASSIC PFTs, such that each land cover class corresponds to a particular mix of PFTs as represented in CLASSIC. This step is supported by vegetation height data from 289 an airborne Lidar campaign over parts of Canada; (3) computing the sub-pixel fractional 290 composition for classes in the ESA-CCI land cover map (300 m resolution) based on the 30 m 291 Hybrid land cover dataset and the Hansen tree cover fraction dataset; (4) using the sub-pixel 292 fractional composition analysis to create a CW-table for mapping the ESA-CCI land cover 293 294 classes onto PFTs as represented in CLASSIC; and (5) since the ESA-CCI dataset is global, the CW-table developed over Canada is extended to the whole globe. 295

296

5 **3.1 CW-table for mapping Hybrid LC classes to CLASSIC PFTs**

297 Among the nine CLASSIC PFTs, BLE and BDD forests are not present in Canada. These are primarily tropical PFTs as represented in CLASSIC. NLD accounts for less than 1% of 298 coniferous forests in Canada (Wang et al., 2019). Therefore we do not consider NLD, BLE, and 299 300 BDD from here on in this study. Considering the fine resolution (30 m) of the Hybrid map, especially relative to the model resolution (~ 16 km) used in this study, we assign fractions of 1.0 301 to the two pure forest classes (LC210 and LC220), the cropland (LC15), and the five non-302 vegetative classes (LC16 to LC32) in its CW-table (Table 1). The mixed-wood category (LC230) 303 is split evenly into NLE and BCD in the table based on the definition in the VLCE legend 304 305 (Hermosilla et al., 2018; Wulder et al., 2003). Note that in Table 1, broadleaf deciduous trees (BDD and BCD) are considered together and separated later into their cold and drought 306 deciduous versions. Similarly, crops and grasses (C₃ and C₄) are considered together and 307 308 separated later into their C₃ and C₄ varieties. The reason for this is that the separation of broadleaf trees into their cold and deciduous phenotypes is based on latitude (Wang et al., 2006). 309 The separation of crops and grasses based on their photosynthetic pathway (C_3 or C_4) is done 310 based on the C₄ fraction from Still and Berry (2003), which is available at 1° resolution. 311 CLASSIC explicitly represents shrub PFTs (Meyer et al., 2021), but this work does not use that 312 model version, and therefore the fraction of tall shrubs is assigned to one of the tree PFTs as was 313 done in creating the CW-table for GLC2000 for use with CLASSIC (Wang et al., 2006). Four 314 (LC2 - Sub-polar taiga needleleaf forest, LC50 - Shrubland, LC80 - Wetland, and LC81-315 Wetland-treed) out of the 17 classes in the Hybrid map are characterized by a mosaic of trees, 316 317 shrubs, and herbaceous vegetation. The vegetation coverage for canopy height above 2 m from Lidar plots is used to inform the partitioning of forest (tall vegetation) to non-forest (low 318

vegetation) fractions for these mixed classes. We overlay the Lidar plots on the Hybrid land
cover map in ArcGIS. Samples (20 to 40, note that these classes do not cover large areas in
Canada) for the four mixed classes in the Hybrid map are selected where there are Lidar data.
The vegetation coverage data (for canopy height above 2 m) from Lidar plots for samples of each
class are used to compute an average coverage of tall vegetation (> 2 m) for that class, which is
then used to assign forest fractions for these four classes in Table 1.

The distribution of tree species from Beaudoin et al. (2014) is used to guide the separation of 325 coniferous versus broadleaf forest fractions. For example, for the Wetland-treed category 326 327 (LC81), maps of tree species show that coniferous forests dominate wetland-treed regions, while broadleaf forests are generally non-existent. We, therefore, assign most of the forest fraction to 328 329 NLE (0.55), only 0.05 to BCD, 0.35 to grasses, and the remaining to the bare ground for LC81 (Table 1). The presence of evergreen shrubs is rare in Canada according to National Forest 330 331 Inventory ground plots data (Gillis et al., 2005), thus we only assign an estimated tall shrub fraction (0.20) in the shrub class (LC50) to BCD. The sub-polar or polar classes (LC11 to LC13) 332 are located above the treeline and mainly consist of low shrubs and grass. The fractions of grass 333 (including low shrubs) and bare ground are based on field surveys of fractional vegetation cover 334 and tundra PFT data in Bjorkman et al. (2018) and Macander et al. (2020) (by computing the 335 average fractions at the field sites which overlap with the sub-polar or polar classes in the 336 Hybrid/NALCMS land cover map). High-resolution images from Google Earth engine or Bing 337 Maps are also used to examine the ratio of vegetated versus bare ground for all classes in which 338 bare ground is present. 339

340 **3.2.** CW-table for mapping ESA-CCI LC classes to CLASSIC PFTs over Canada

341 **3.2.1** The error and sub-pixel fractional error matrices

A standard approach for the accuracy assessment of LC products is to use an error matrix. It is a 342 square array or table of numbers arranged in rows and columns, in which the classification from 343 the LC product (usually represented by the rows) is compared to the reference data (usually 344 represented by the columns) for each category (Congalton, 1991). The major diagonal of the 345 matrix presents the number of correct classifications indicating the agreement between the LC 346 and the reference data for each category. In practice, fine-resolution regional maps are often used 347 to assess large-scale LC products derived from coarse-resolution data (Cihlar et al., 2003). In 348 349 such cases, the fine-resolution reference data are aggregated/regridded to match the grid of the coarse-resolution data. Several classes in the reference data may be present in a single coarse-350 resolution pixel depending on the homogeneity of the landscape. In order to compare the 351 reference and the LC data on a one-to-one basis, the dominant LC class (the class with the most 352 abundant fractions based on all fine-resolution pixels in the reference data) is often assigned to 353 the regridded reference pixel. 354

The sub-pixel fractional error matrices have been introduced as a more appropriate way of 355 assessing the accuracy of mixed pixels by Latifovic and Olthof (2004). In contrast with an error 356 357 matrix where only the dominant LC class is used as described above, the sub-pixel fractional error matrix is produced by assigning sub-dominant LC classes from all fine-resolution pixels in 358 the reference data to the corresponding single coarse-resolution pixel. It thus allows a 359 360 quantitative assessment of the fractional composition of the LC classes in the coarse resolution dataset. In this study, both the 30 m Hansen TCF data and the 30 m Hybrid LC map are used to 361 compute the sub-pixel fractional error matrices of the 300 m ESA-CCI dataset (Table 2 and 362 363 Table 3). However, the objective here is not an accuracy assessment as in Latifovic and Olthof

364 (2004) but rather to obtain the fractional composition of the LC classes in the ESA-CCI product and to inform the PFT mapping process. We refer to this process as the sub-pixel fractional 365 composition analyses in the rest of this paper. Sub-pixel fractional composition analyses is first 366 performed for each ecozone and then weighted mean fractions for each ESA-CCI class are 367 computed based on pixel counts in each of the ecozones (see the location of ecozones in Fig. 1). 368 For the Hansen TCF data, results are shown only for the ESA-CCI LC classes containing forests 369 in Canada (Table 2). In the ESA-CCI legend (Table 4), two sub-classes for broadleaf (LC61 and 370 LC62) and needleleaf (LC71 and LC72) forests are included as the closed (>40% forest cover) 371 372 and open (10-40% forest cover) categories apart from the main classes (LC60 and LC70, closed to open (>15%)). As expected, the TCF is larger for the closed classes than for the main and the 373 open classes (Table 2). In Table 2, we also include ratios of TCF between the main class and the 374 closed class, and between the open class and the closed class. We note that the ratios are 375 different for broadleaf (main class vs. closed class: 68.5/86.7=0.8; open class vs. closed class: 376 0.43/86.7=0.43) and needleleaf (main class vs. closed class: 39.3/61.7=0.6; open class vs. closed 377 class: 23.2/61.7=0.38) forests, which need to be taken into account when creating the CW-table 378 for the ESA-CCI dataset. 379

To obtain representative class compositions of the ESA-CCI dataset, only homogenous ESA-CCI pixels are included in the sub-pixel composition analyses based on the Hybrid LC data. The homogenous pixels are defined following the method in Herold et al. (2008). To quantify landscape heterogeneity, 3×3 pixel neighborhoods are assessed for the ESA-CCI data. A neighborhood is considered homogenous if only one LC class is present. The weighted mean fraction for each class is computed from ecozones with more than 10 homogenous ESA-CCI

pixels for that class (Table 3). Only 13 out of the 37 ESA-CCI classes are included in Table 3,

the rest of the classes either have limited presence in Canada or are non-vegetative (Table 4).

In the Hybrid CW-table (Table 1), four LC classes (2, 81, 210, and 230) contribute to the

fractional cover of NLE, and five LC classes (50, 80, 81, 220, and 230) contribute to the

390 fractional cover of BCD. In Table 3, we also include an integrated fractional cover (F) for NLE

and BCD (last two rows) for each of the ESA-CCI classes based on the following formula:

392
$$F = \sum_{i=1}^{N} F \mathbf{1}_{i} * F \mathbf{2}_{i}$$
 (1)

Where $F1_i$ are fractions in Table 3, $F2_i$ are fractions in Table 1, and N is the number of Hybrid LC classes contributing to NLE (N = 4) or BCD (N = 5) as shown in Table 1. As an example, the fraction of NLE for the LC70 (Tree cover needleleaf evergreen closed to open) in ESA-CCI data in Table 3 (see the NLE row and the column for class 70) is obtained as follows: F = $0.02 \times 0.20 +$ $0.17 \times 0.55 + 0.29 \times 1.0 + 0.09 \times 0.5 = 0.44$. This process reduces the subjectivity in assigning the ESA-CCI land cover classes to CLASSIC's two tree PFTs (NLE and BCD) that are present in Canada since the process is based on the high-resolution Hybrid LC data.

400 3.2.2 CW-table for the ESA-CCI LC dataset over Canada

401 Table 2 and Table 3 thus form the basis for creating the CW-table for mapping the ESA-CCI LC

402 classes to CLASSIC's PFTs (Fig. 2 and Table 4). For the ESA-CCI class LC61 (Tree cover

403 broadleaved deciduous closed) (not included in Table 3 due to limited presence in Canada),

- ratios of TCF for LC60 vs LC61 in Table 2 and the fractions of LC60 (Tree cover broadleaved
- 405 deciduous closed to open) in Table 3 are used to derive fractions for LC61 in Table 4. The
- 406 remapping of LC62 (Tree cover broadleaved deciduous open) and LC72 (Tree cover needleleaf
- 407 evergreen open) into CLASSIC's PFTs is done in a similar way. Since NLD is not included in

408 either Table 2 or Table 3, the needleleaf deciduous tree cover classes (LC80-82) are assigned to the same fractions as the needleleaf evergreen tree cover classes (LC70-72). For simplicity, the 409 fractions in Table 3 are rounded to values with either "0" or "5" at the hundredth place when 410 used in Table 4. For the rest of the classes not included in either Table 2 or Table 3, values are 411 based on the default CW-table from the ESA-CCI user guide (Table 7-2, ESA, 2017). The spatial 412 413 distribution of LC classes is also taken into consideration when determining the fractions in the CW-table. For example, the sparse vegetation class (LC150) is mainly distributed above the 414 treeline in alpine and Arctic tundra environments, thus we only assign 0.05 to BCD, the rest to 415 416 C3G/C4G and bare ground (Table 4). 417 The six CLASSIC PFTs (those present in Canada) are produced from the Hybrid and the ESA-CCI maps based on Table 1 and Table 4 respectively. The PFTs from the Hybrid map are used as 418 a reference here to map ESA-CCI land cover classes to CLASSIC's PFTs. To make the spatial 419

420 distribution of PFTs from ESA-CCI agree better with those from the Hybrid dataset, fractions for

421 the following classes in Table 4 are adjusted upward by 0.05: LC60 from 0.65 to 0.70 for BCD;

LC71 and LC81 from 0.80 to 0.85 for NLE; and LC120 from 0.10 to 0.15 for BCD. Values for

423 LC10-20 are also slightly adjusted to reduce crop fractions.

424 Based on Table 4, the fractional coverage of nine CLASSIC PFTs are also produced on a global

scale and used in offline CLASSIC simulations in Arora et al. (2022), who carry out a

426 comprehensive assessment of the impact of using two different LC datasets (ESA-CCI versus

427 GLC2000) for representing the nine PFTs in the CLASSIC model. However, some adjustments

- 428 to Table 4 are found to be necessary. This is because fractions of NLE (Needleleaf evergreen
- 429 forests) in Eurasia are found to be too low relative to the Hansen TCF data, with maximum
- 430 values only around 0.45 in most NLE dominated areas, where the maximum TCF from the

431 Hansen dataset is around 0.80. This indicates that the needleleaf evergreen forests classes (LC

432 70-72) in the ESA-CCI map may represent different forest/tree cover fractions in Canada and

433 Eurasia, which is confirmed by sub-pixel fractional composition analyses based on the Hansen

434 TCF dataset. Details are presented in Appendix B.

435 **4. Results**

436 4.1 Comparison of PFTs from Hybrid, ESA-CCI, and GLC2000 data

437 Figure 3 shows the spatial distribution of PFTs derived from the Hybrid, ESA-CCI, and

438 GLC2000 LC datasets respectively. C₄ crops (C4C) and grasses (C4G) are sparse in Canada as

439 would be expected since C4 PFTs grow only in warmer temperatures when the average monthly

temperature exceeds 22 °C (Fox et al., 2018). Based on the fractional distribution of C4

441 vegetation in Still and Berry (2003) and the Hybrid map, the average fraction is 0.5% for C4

442 crops and 0.1% for C4 grasses in Canada. Therefore, only four out of the six PFTs (those present

443 in Canada) are shown in Figure 3. In general, the spatial distributions of the PFTs from the ESA-

444 CCI and the Hybrid datasets agree well except for C₃ grasses (C3G) (Fig. 3j and Fig. 3k). This is

not surprising given that the CW-table for the ESA-CCI dataset is based on the Hybrid map.

446 Areas mapped as C3G in Hybrid (Fig. 3j), were mainly classified as sparse vegetation (LC150)

447 in the ESA-CCI legend (Table 4). However, LC150 from ESA-CCI was also found in some areas

448 of the high Arctic islands, where barren land is the dominant class in the Hybrid map (grey

449 coloured areas in Fig. 1). If too much grass were assigned to LC150, it would yield

- 450 unrealistically large fractional coverage of grass in the high Arctic islands. In Table 4, for
- 451 LC150, 0.05 is assigned to BCD, 0.35 to grasses, and the rest to the bare ground for LC150,
- which yields a total vegetation cover of 40% and is more than the definition (<15% vegetation)

used in the ESA-CCI legend. Yet, this still results in less C3G and less bare ground in the ESACCI map (Fig. 3k and Fig. 3n) than those from the Hybrid map (Fig. 3j and Fig. 3m). This
suggests that it is not ideal to classify areas in the high Arctic islands and in the Arctic tundra
region as being in the same land cover category.

457 There are large differences in the spatial distribution of the PFTs based on the GLC2000 LC

458 product and those based on the Hybrid and ESA-CCI datasets (Fig. 3 and Fig. 4). Relative to

459 PFTs from ESA-CCI, GLC2000 has less NLE and more BCD in northwestern Canada, and more

460 NLE in southern and eastern Canada (Fig. 4a and Fig. 4b). GLC2000 based CLASSIC PFT

461 fractions also exhibit more crops, less grass, and more bare ground (Fig. 4c-4e). These

differences partly stem from the differences in the ESA-CCI and GLC2000 LC datasets, but arealso due to the fact how the fractions in the CW-tables of the two datasets are used to translate

464 LC data to fractional coverage of PFTs as demonstrated in Wang et al. (2019).

465 4.2 Bias in simulated surface albedo and LAI

466 The top row of Figure 5 shows the bias in winter albedo (March) simulated by CLASSIC when using PFT distributions based on the ESA-CCI (Fig. 5a) and GLC2000 products (Fig. 5b). While 467 model biases are the result of both the driving geophysical and meteorological data that are used 468 to force the model, as well as the model itself, the comparison between the two simulations does 469 show the effect of differences in the distribution of PFTs. Relative to observed surface albedo 470 from MODIS, there are relatively large negative biases in the southwest of Hudson Bay and 471 central Quebec, while there are relatively large positive biases in western Canada and Alaska in 472 the simulation when using the GLC2000 product to obtain PFT distributions (Fig. 5b). Both the 473 negative and the positive biases are largely reduced in the simulation using PFT distributions 474

475 based on the ESA-CCI product (Fig. 5a). The lower row of Figure 5 shows the spatial distribution of the difference in surface albedo (Fig. 5c) and leaf area index (Fig. 5d) between the 476 two simulations, which are closely correlated (r = -0.85). Given the same meteorological forcing 477 dataset is used to drive both simulations, the differences in the simulated LAI are due mainly to 478 the different PFT distributions used in the two simulations. Since NLE is the only PFT with 479 LAI > 0 during winter in Canada, the LAI difference in March as shown in Figure 5d is mainly 480 due to the different fractional coverage of NLE based on the ESA-CCI and GLC2000 products 481 (Fig. 4a). 482

In contrast, the large positive albedo biases (up to ~ 0.4) in southern Canada are more or less the 483 same in both simulations (Fig. 5a and Fig. 5b), where the dominant PFT is C3 crops (Fig. 3h and 484 485 Fig. 3i). Those positive albedo biases are likely due to the standing crop stubble and the lack of the representation of blowing snow and its sublimation currently in CLASSIC (Harder et al., 486 487 2018; Pomeroy et al., 1993). Harder et al. (2018) showed that the height of the stubble over 488 wheat and canola field in Saskatchewan, Canada may range from 10 to 40 cm, with a maximum PAI (plant area index) of 1.0. Wang et al. (2016) showed that surface albedo in CLASSIC 489 490 decreased exponentially with increasing PAI for the bare or snow-covered canopy over snow, while most reductions of the albedo were achieved through the increase of PAI from 0 to 1.0. 491 They showed that surface albedo decreased from 0.75 to 0.31 in CLASSIC when PAI increased 492 from 0 to 1.0 for the bare canopy over snow, which appears to account for most of the positive 493 albedo biases in the agricultural areas of southern Canada (Fig. 5a and Fig. 5b). Improvements to 494 495 the crop module of CLASSIC to improve cropland albedo are currently being considered.

496 **5. Summary and conclusions**

A hybrid land cover map at 30 m resolution is created by merging the NALCMS and VLCE land 497 cover products over Canada. Vegetation height data from Lidar plots, tree species, and high 498 resolution images are used to inform the creation of a CW-table for mapping the 17 LC classes 499 of the Hybrid map to six CLASSIC PFTs that are present in Canada. Both the Hybrid map and 500 the Hansen tree cover fraction data are used to compute the sub-pixel fractional composition of 501 502 the LC classes in the ESA-CCI LC dataset, which is then used to create a cross-walking table for mapping the 37 ESA-CCI categories to CLASSIC PFTs over the model's Canadian domain. 503 Based on the new CW-tables, PFT distributions are produced from the Hybrid and the ESA-CCI 504 505 LC products, respectively, and are compared with those based on the GLC2000 dataset currently used in CLASSIC. The results show that the spatial distribution of PFTs from the ESA-CCI 506 dataset is in better agreement with those from the Hybrid map, while there are large differences 507 in the PFTs from the GLC2000 dataset and from the Hybrid/ESA-CCI datasets. The CW-table 508 developed over Canada is adjusted and also used to map PFTs based on the ESA-CCI LC 509 product for use in CLASSIC simulations at the global scale. 510 Our PFT mapping approach for the ESA-CCI dataset is mainly based on sub-pixel fractional 511 composition analyses using the Hybrid map and the Hansen tree cover fraction data, and 512 therefore the accuracy of the latter two datasets affects the PFT mapping process. Some LC 513 categories in the ESA-CCI legend either have limited presence or no presence in Canada, such as 514 the Needleleaf deciduous trees, Broadleaf Evergreen trees, and Broadleaf Dry Deciduous trees 515 etc., and the sub-pixel fractional composition analyses therefore can not be performed for these 516 517 LC categories. The needleleaf deciduous tree cover classes are assigned to the same fractions as the needleleaf evergreen tree cover classes in the CW-table, and values based on the default CW-518 table from the ESA-CCI user guide are used for the other LC categories. Therefore potentially 519

520 large uncertainties may be associated with these classes in the resulting fractional coverage of PFTs especially at the global scale. Similar analyses for other regions (e.g. Eurasia and tropics) 521 for which high quality regional land cover maps are available will be helpful in reducing these 522 uncertainties in the future work. In addition, the exercise of mapping PFTs at the global scale in 523 this study reveals that there are inconsistencies in the representation of fractional coverage for 524 some LC categories in the ESA-CCI map for different regions of the globe. Future improvements 525 in the consistency of the LC categories globally in the ESA-CCI LC product would greatly 526 benefit the land surface and the earth system modelling community. In the meantime, caution 527 528 should be exercised when using this product for mapping PFTs represented in any LSM based on a single cross-walking table at the global scale. 529

CLASSIC simulations driven with meteorological data from the CRU-JRA product show that the 530 simulated winter albedo is improved when using PFT distributions based on the ESA-CCI LC 531 532 product compared to that based on the GLC2000 product, which is consistent with findings from previous studies. While, CLASSIC simulations could also have been performed using its PFT 533 distributions based on the Hybrid LC product, the reason for using the ESA-CCI based PFT 534 fractions for CLASSIC is that ESA-CCI is a global product. CLASSIC simulations are routinely 535 performed at the global scale both in the framework of the Canadian Earth System Model (Swart 536 et al., 2019), where CLASSIC serves as its land component, and offline where global CLASSIC 537 simulations driven with the CRU-JRA meteorological data contribute to the annual global carbon 538 budget assessments of the Global Carbon Project (Friedlingstein et al., 2020; Seiler et al., 2021). 539 540 Untreated crop stubble appears to be contributing to the positive winter albedo biases in southern Canada, which needs to be addressed in a future version of CLASSIC. These results underscore 541

the importance of accurate representation of vegetation distribution in a realistic simulation ofsurface albedo in LSMs.

Previous methods for mapping PFTs from LC datasets have mainly been based on class 544 descriptions, expert knowledge, and the spatial distribution of global biomes, which is a largely 545 546 subjective process. As a consequence, a PFT method developed for mapping one LC dataset to PFTs represented in one model can not be easily transferred to other LC datasets even for 547 deriving PFTs in the same model. The development of satellite and computing technology has 548 enabled the creation of more detailed global LC products at finer spatial resolutions in recent 549 550 years, however, the lack of an objective PFT mapping method impedes the implementation of the new improved LC products in LSMs. Here, we have proposed a method to inform the cross-551 walking process using sub-pixel fractional composition analyses based on a tree cover fraction 552 dataset and a fine-resolution LC map. Our results suggest that the sub-pixel fractional 553 554 composition analyses provide an effective way to reduce uncertainties in the cross-walking process and therefore, to some extent, objectifies the otherwise subjective process. The PFT 555 mapping approach developed in this study can also be applied to other LC datasets for mapping 556 PFTs used in other LSMs. 557

558

559 Appendix A

560 In CLASSIC, the surface albedo for a canopy over snow (α) is:

561
$$\alpha = \alpha_c (1 - \chi) (1 - f_{snow}) + \alpha_{c,snow} (1 - \chi) (f_{snow}) + \alpha_{snow} \chi \tau_c$$
(1)

562
$$\chi = \exp\left(-K^*PAI\right) \tag{2}$$

calculated using separate parameters (α_c , $\alpha_{c,snow}$, τ_c and K) for both the visible (VIS) and near 564 infrared (NIR) bands, where α_c is the snow-free canopy albedo, $\alpha_{c,snow}$ the snow-covered canopy 565 albedo, f_{snow} the fraction of the canopy with snow on it, α_{snow} the snowpack albedo. τ_c is canopy 566 transmissivity and is modeled using a Beer's law approach, ignoring multiple reflections 567 (Verseghy et al. 1993). K is an extinction coefficient that varies with vegetation type. The 568 appearance of τ_c in the last term of Eq.1 accounts for the shading of the snowpack by the canopy, 569 570 converting the simulated snowpack albedo to an effective value of the canopy gaps. PAI is plant area index which is the sum of leaf area index and stem area index. 571

572

573 Appendix B

Based on Table 4, the fractional coverage of nine CLASSIC PFTs are also produced on a global 574 scale. However, some adjustments to Table 4 were found necessary. This is because fractions of 575 576 NLE (Needleleaf evergreen forests) in Eurasia are found to be too low relative to the Hansen TCF data, with maximum values of only around 0.45 in most NLE dominated areas, where the 577 maximum TCF from the Hansen dataset is around 0.80. Needleleaf evergreen forests are 578 579 represented by LC classes 70 (closed to open), 71 (closed), and 72 (open). Examining the ESA-CCI LC map shows that in Eurasia nearly all needleleaf evergreen forests are classified as LC70 580 581 (closed to open), with only less than 400 pixels as LC71 (closed), and none as LC72 (open). In 582 contrast, in Canada 36% of needleleaf evergreen forest are classified as LC70 (closed to open), 64% as LC71 (closed), and less than 1% as LC72 (open). This is understandable given that sub-583 584 classes were only assigned where surface samples were available (ESA, 2017). Sub-pixel 585 fractional composition analyses of the ESA-CCI classes based on the Hansen TCF dataset show

| 586 | that in Eurasia TCF for LC70 (closed to open) is 66% and for LC71 (closed) is 35% (note the |
|-----|--|
| 587 | few pixels within this class). This is in contrast with those in Canada where the TCF for LC70 |
| 588 | (closed to open) is 39% and for LC71 (closed) is 62%, explaining the too low NLE fractions in |
| 589 | Eurasia when mapping PFTs based on Table 4, and also the too high TCF in northwestern |
| 590 | Canada when mapping PFTs based on the default CW-table (Wang et al., 2018). In order to |
| 591 | apply Table 4 globally, the original LC70 (closed to open) was split into two classes: LC73 (a |
| 592 | new class) which is the same as LC70 over Canada (and zero everywhere else), and LC70 |
| 593 | (revised) which is the same as before except zero over Canada. The fractions for the new LC70 |
| 594 | class are made the same as for LC71 in Table 4, which applies to NLE outside of Canada. |
| 595 | Essentially, the closed-to-open needleleaf forest LC70 class over Eurasia is treated as the closed |
| 596 | needleleaf forest. |

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References

- Arora, V.: Land surface modelling in general circulation models: a hydrological perspective,
- PhD thesis, Department of Civil and Environmental Engineering, University of Melbourne,1997.
- Arora, V. K., Boer, G. J.: A parameterization of leaf phenology for the terrestrial ecosystem
 component of climate models. Glob. Chang. Biol., 11, 39–59, 2005.
- 610 Arora, V. K. and Boer, G. J.: Uncertainties in the 20th century carbon budget associated with
- 611 land use change, Glob. Change Biol., 16, 3327–3348, https://doi.org/10.1111/j.1365-
- 612 2486.2010.02202.x, 2010.
- Arora, V. K., Boer, G. J., Christian, J. R., Curry, C. L., Denman, K. L., Zahariev, K., Flato, G.
- 614 M., Scinocca, J. F., Merryfield, W. J., and Lee, W. G.: The Effect of Terrestrial Photosynthesis
- Down Regulation on the Twentieth-Century Carbon Budget Simulated with the CCCma Earth
- 616 System Model, J. Clim., 22, 6066–6088, https://doi.org/10.1175/2009JCLI3037.1, 2009.
- Arora, V. K., Seiler, C., Wang, L., and Kou-Giesbrecht, S.: Towards an ensemble-based
- evaluation of land surface models in light of uncertain forcings and observations, EGUsphere
- 619 [preprint], https://doi.org/10.5194/egusphere-2022-641, 2022.
- Beaudoin, A., Bernier, P.Y., Guindon, L., Villemaire, P., Guo, X.J., Stinson, G., Bergeron, T.,
- 621 Magnussen, S., and Hall, R.J.: Mapping attributes of Canada's forests at moderate resolution
- through kNN and MODIS imagery, Can. J. For. Res. 44: 521–532, dx.doi.org/10.1139/cjfr-2013-
- 623 0401, 2014.
- 624 Bartholomé, E., and Belward, A. S.: GLC2000: A new approach to global land cover mapping
- from Earth Observation data, Int. J. Remote Sens., 26, 1959–1977, 2005.

- Bartlett, P. A., MacKay, M. D., and Verseghy, D. L.: Modified snow algorithms in the Canadian
- land surface scheme: Model runs and sensitivity analysis at three boreal forest stands, Atmos.–
 Ocean, 44, 207–222, 2006.
- Bartlett, P. A., and Verseghy, D. L.: Modified treatment of intercepted snow improves the
- 630 simulated forest albedo in the Canadian land surface scheme, Hydrol. Process., 29, 3208–3226,
- 631 https://doi:10.1002/hyp.10431, 2015.
- Betts, R. A.: Biogeophysical impacts of land use on present-day climate: near-surface
- temperature change and radiative forcing, Atmos. Sci. Lett., 2, 39–51,
- 634 https://doi.org/10.1006/asle.2001.0037, 2001.
- Bjorkman, A. D., Myers-Smith, I. H., Elmendorf, S.C. et al.: Plant functional trait change across
- a warming tundra biome, Nature, 562, 57–62, https://doi.org/10.1038/s41586-018-0563-7, 2018.
- Bonan, G. B., Levis, S., Kergoat, L., Oleson, K. W.: Landscapes as patches of plant functional
- types: An integrating concept for climate and ecosystem models, Global Biogeochem. Cycles,
- 639 16(2), 1021, https://doi:10.1029/2000GB001360, 2002.
- Bonan, G. B., Levis, S., Sitch, S., Vertenstein, M., and Oleson, K. W.: A dynamic global
- 641 vegetation model for use with climate models: Concepts and description of simulated vegetation
- 642 dynamics. Global Change Biology, 9, 1543–1566, 2003.
- Bontemps, S., Herold, M., Kooistra, L., van Groenestijn, A., Hartley, A., Arino, O., Moreau, I.,
- 644 Defourny, P.: Revisiting land cover observation to address the needs of the climate modeling
- 645 community, Biogeosciences, 9, 2145–2157, https://doi:10.5194/bg-9-2145-2012, 2012.

- 646 Di Gregorio, A.: Land Cover Classification System Classification concepts and user manual
- for Software version 2, FAO Environment and Natural Resources Service Series, No. 8, Rome,
 208pp, 2005.
- ESA: Land Cover CCI Product User Guide Version 2. Tech. Rep., Available at:
- 650 maps.elie.ucl.ac.be/CCI/viewer/download/ESACCI-LC-Ph2-PUGv2_2.0.pdf, 2017.
- Fisher, R. A., Koven, C. D., Anderegg, W. R. L. et al.: Vegetation demographics in Earth System
- Models: A review of progress and priorities. Glob Change Biol., 24, 35–54,
- 653 https://doi.org/10.1111/gcb.13910, 2018.
- Fox, D.L., Pau, S., Taylor, L., Strömberg, C. A. E., Osborne, C.P., Bradshaw, C., Conn, S.,
- Beerling, D.J., and Still, C.J.: limatic Controls on C4 Grassland Distributions During the
- Neogene: A Model-Data Comparison, Frontiers in Ecology and Evolution, 6,
- 657 <u>https://doi:10.3389/fevo.2018.00147</u>, 2018.
- Friedlingstein, P., O'Sullivan, M., Jones, M. W., Andrew, R. M., Hauck, J., Olsen, A., Peters, G.
- P., Peters, W., Pongratz, J., Sitch, S., Le Quéré, C., Canadell, J. G., Ciais, P., Jackson, R. B.,
- Alin, S., Aragão, L. E. O. C., Arneth, A., Arora, V., Bates, N. R., Becker, M., Benoit-Cattin, A.,
- Bittig, H. C., Bopp, L., Bultan, S., Chandra, N., Chevallier, F., Chini, L. P., Evans, W., Florentie,
- L., Forster, P. M., Gasser, T., Gehlen, M., Gilfillan, D., Gkritzalis, T., Gregor, L., Gruber, N.,
- Harris, I., Hartung, K., Haverd, V., Houghton, R. A., Ilyina, T., Jain, A. K., Joetzjer, E., Kadono,
- 664 K., Kato, E., Kitidis, V., Korsbakken, J. I., Landschützer, P., Lefèvre, N., Lenton, A., Lienert, S.,
- 665 Liu, Z., Lombardozzi, D., Marland, G., Metzl, N., Munro, D. R., Nabel, J. E. M. S., Nakaoka, S.-
- 666 I., Niwa, Y., O'Brien, K., Ono, T., Palmer, P. I., Pierrot, D., Poulter, B., Resplandy, L.,
- 667 Robertson, E., Rödenbeck, C., Schwinger, J., Séférian, R., Skjelvan, I., Smith, A. J. P., Sutton,

- A. J., Tanhua, T., Tans, P. P., Tian, H., Tilbrook, B., van der Werf, G., Vuichard, N., Walker, A.
- 669 P., Wanninkhof, R., Watson, A. J., Willis, D., Wiltshire, A. J., Yuan, W., Yue, X., and Zaehle,
- 670 S.: Global Carbon Budget 2020, Earth Syst. Sci. Data, 12, 3269–3340,
- 671 https://doi.org/10.5194/essd-12-3269-2020, 2020.
- 672 Fritz, S., See, L., McCallum, I., Schill, C., Obersteiner, M., van der Velde, M., Boettcher, H.,
- 673 Havlík, P., Achard, F.: Highlighting continued uncertainty in global land cover maps for the user
- 674 community, Environ. Res. Lett., 6, 44005. https://dx.doi.org/10.1088/1748-9326/6/4/044005,
- **675** 2011.
- 676 Georgievski, G., Hagemann, S.: Characterizing uncertainties in the ESA-CCI land cover map of
- the epoch 2010 and their impacts on MPI-ESM climate simulations, Theor. Appl. Climatol., 137,

678 1587–1603, https://doi.org/10.1007/s00704-018-2675-2, 2019.

- 679 Gillis, M. D., Omule, A. Y., and Brierley, T.: Monitoring Canada's forests: The National Forest
- 680 Inventory, The Forestry Chronicle, 81(2): 214-221, <u>https://doi.org/10.5558/tfc81214-2</u>, 2005.
- Harris, I. C.: CRU JRA v2.1: A forcings dataset of gridded land surface blend of Climatic
- Research Unit (CRU) and Japanese reanalysis (JRA) data; Jan. 1901 Dec. 2019, Centre for
- 683 Environmental Data Analysis, University of East Anglia Climatic Research Unit,
- 684 https://catalogue.ceda.ac.uk/uuid/10d2c73e5a7d46f4ada08b0a26302ef7, 2020
- Hartley, A. J., MacBean, N., Georgievski, G., Bontemps, S.: Uncertainty in plant functional type
- distributions and its impact on land surface models, Remote Sens. Environ., 203, 71–89,
- 687 https://doi.org/10.1016/j.rse.2017.07.037, 2017.

- Harder, P., Warren D. H., and Pomeroy, J. W.: Modeling the Snowpack Energy Balance during
- 689 Melt under Exposed Crop Stubble, J. Hydrometeorol., 19, 1191-1214, https://DOI:
- 690 10.1175/JHM-D-18-0039.1, 2018.
- Hansen, M. C., DeFries, R. S., Townshend, J. R. G., Sohlberg, R., Dimiceli, C., Carroll, M. L.:
- 692 Towards an operational MODIS continuous field of percent tree cover algorithm: Examples us-
- 693 ing AVHRR and MODIS data, Remote Sens. Environ., 83, 303–319, 2002.
- Hansen, M. C., Stehman, S. V., Potapov, P. V.: Quantification of global gross forest cover loss,
- 695 P. Natl. Acad. Sci., 107, 8650–8655, https://doi.org/10.1073/pnas.0912668107, 2010.
- Hansen, M. C., Potapov, P. V., Moore, R., Hancher, M., Turubanova, S. A., Tyukavina, A.,
- 697 Thau, D., Stehman, S. V., Goetz, S. J., Loveland, T.R., et al.: Highresolution global maps of
- 21st-century forest cover change, Science, 342, 850–853, https://DOI: 10.1126/science.1244693,
 2013.
- Hengl, T., Mendes de Jesus, J., Heuvelink, G. B. M., Ruiperez Gonzalez, M., Kilibarda, M.,
- 701 Blagotić, A., Shangguan, W., Wright, M. N., Geng, X., Bauer-Marschallinger, B., Guevara, M.
- A., Vargas, R., MacMillan, R. A., Batjes, N. H., Leenaars, J. G. B., Ribeiro, E., Wheeler, I.,
- 703 Mantel, S., and Kempen, B.: SoilGrids250m: Global gridded soil information based on machine
- 704 learning, PLOS ONE, 12, 1–40, https://doi.org/10.1371/journal.pone.0169748, 2017
- 705 Hermosilla, T., Wulder, M. A., White, J. C., Coops, N. C., Hobart, G.W.: Disturbance informed
- annual land cover classification maps of Canada's forested ecosystems for a 29-year Landsat
- time series. Can. J. Remote. Sens., 44, 67–87. https://doi.org/10.1080/07038992.2018.1437719,
- 708 2018.

- 709 Herold, M., Mayaux, P., Woodcock, C. E., BaESACCIni, A., Schmullius, C.: Some Challenges
- in Global Land Cover Mapping: An Assessment of Agreement and Accuracy in Existing 1 km
- 711 Datasets, Remote Sens. Environ., 112, 2538–2556, 2008.
- Hopkinson, C., Chasmer, L., Lim, K., Treitz, P., and Creed, I.: Towards a universal lidar canopy
- height indicator, Canadian Journal of Remote Sensing, 32 (2), 139-152, https://doi:
- 714 10.5589/m06-006, 2006.
- 715 Krinner, G., Viovy, N., de Noblet-Ducoudre, N., Ogee, J., Polcher, J., Friedlingstein, P. et al.: A
- dynamic global vegetation model for studies of the coupled atmosphere-biosphere system,
- 717 Global Biogeochemical Cycles, 19(1), GB1015, 2005.
- Latifovic, R., Pouliot, D., and Olthof, I.: Circa 2010 Land Cover of Canada: Local Optimization
- 719 Methodology and Product Development, Remote Sens., 9, 1098,
- 720 https://doi.org/10.3390/rs9111098, 2017.
- Latifovic, R., and Olthof, I.: Accuracy assessment using sub-pixel fractional error matrices of
- global land cover products derived from satellite data, Remote Sens. Environ., 90, 153–165,
 2004.
- Macander, M. J., Frost, G. V., Nelson, P. R., and Swingley, C. S.: ABoVE: Tundra Plant
- Functional Type Continuous-Cover, North Slope, Alaska, 2010-2015, ORNL DAAC, Oak
- Ridge, Tennessee, USA, https://doi.org/10.3334/ORNLDAAC/1830, 2020.
- 727 Mayaux, P., Eva, H., Gallego, J., Strahler, A. H., Herold, M., Agrawal, S., Naumov, S., De
- 728 Miranda, E. E., Di Bella, C. M., Ordoyne, C., et al.: Validation of the Global Land Cover 2000
- 729 Map. IEEE Trans. Geosci. Remote Sens., 44, 1728–1739, 2006.

- 730 Melton, J. R. and Arora, V. K.: Competition between plant functional types in the Canadian
- 731 Terrestrial Ecosystem Model (CTEM) v. 2.0, Geosci Model Dev, 9, 323–361,
- 732 https://doi.org/10.5194/gmd-9-323-2016, 2016.
- 733 Melton, J. R., Arora, V. K., Wisernig-Cojoc, E., Seiler, C., Fortier, M., Chan, E., and
- 734 Teckentrup, L.: CLASSIC v1.0: the open-source community successor to the Canadian Land
- Surface Scheme (CLASS) and the Canadian Terrestrial Ecosystem Model (CTEM) Part 1:
- Model framework and site-level performance, Geosci. Model Dev. Discuss., 1–40,
- 737 https://doi.org/10.5194/gmd-2019-329, 2019.
- 738 Meyer, G., Humphreys, E. R., Melton, J. R., Cannon, A. J., and Lafleur, P. M.: Simulating
- shrubs and their energy and carbon dioxide fluxes in Canada's Low Arctic with the Canadian
- Land Surface Scheme Including biogeochemical Cycles (CLASSIC), Biogeosciences, 18, 3263–
 3283, 2021.
- 742 Moody, E. G., King, M. D., Schaaf, C. B., Hall, D. K., Platnick, S.: Northern Hemisphere five-
- year average (2000–2004) spectral albedos of surfaces in the presence of snow: statistics
- computed from Terra MODIS land products, Remote Sens. Environ., 111, 337–345, https://DOI:
 10.1016/j.rse.2007.03.026, 2007.
- 746 Ottlé, C., Lescure, J., Maignan, F., Poulter, B., Wang, T., and Delbart, N.: Use of various remote
- sensing land cover products for plant functional type mapping over Siberia, Earth Syst. Sci. Data,
- 748 5, 331–348, https://doi:10.5194/essd-5-331-2013, 2013.
- 749 Pielke, R. A., Avissar, R., Raupach, M., Dolman, A. J., Zeng, X., Denning, S.: Interactions
- between the atmosphere and terrestrial ecosystem: influence on weather and climate, Global
- 751 Change Biol., 4, 461–475, 1998.

- 752 Poulter, B., Ciais, P., Hodson, E., Lischke, H., Maignan, F., Plummer, S., and Zimmermann, N.
- E.: Plant functional type mapping for earth system models, Geosci. Model Dev., 4, 993–1010,
- 754 https://doi.org/10.5194/gmd-4-993-2011, 2011.
- Poulter, B., MacBean, N., Hartley, A., Khlystova, I., Arino, O., Betts, R., Bontemps, S.,
- 756 Boettcher, M., Brockmann, C., Defourny, P., Hagemann, S., Herold, M., Kirches, G., Lamarche,
- 757 C., Lederer, D., Ottlé, C., Peters, M., and Peylin, P.: Plant functional type classification for earth
- system models: results from the European Space Agency's Land Cover Climate Change
- 759 Initiative, Geosci. Model Dev., 8, 2315–2328, https://doi.org/10.5194/gmd-8-2315-2015, 2015.
- 760 Pomeroy, J. W., Gray, D. m., and Landine, P. G.: The Prairie Blowing Snow Model:
- Characteristics, validation, operation. J. Hydrol., 144, 165–192, <u>https://doi.org/10.1016/0022-</u>
 <u>1694(93)90171-5</u>, 1993.
- 763 Schaaf, C. B., et al.: First operational BRDF, albedo nadir reflectance products from MODIS,
- 764 Remote Sens. Environ., 83, 135–148, https://doi:10.1016/S0034-4257(02)00091-3, 2002.
- 765 Scheiter, S., Langan, L., Higgins, S.I.: Next-generation dynamic global vegetation models:
- rearning from community ecology, New Phytol., 198, 957–969,
- 767 <u>https://doi.org/10.1111/nph.12210</u>, 2013.
- 768 Seiler, C., Melton, J. R., Arora, V. K., and Wang, L.: CLASSIC v1.0: the open-source
- 769 community successor to the Canadian Land Surface Scheme (CLASS) and the Canadian
- 770 Terrestrial Ecosystem Model (CTEM) Part 2: Global Benchmarking, Geoscientific Model
- 771 Development, https://doi.org/10.5194/gmd-14-2371-2021, 2021.

- Shangguan, W., Hengl, T., Mendes de Jesus, J., Yuan, H., and Dai, Y.: Mapping the global depth
- to bedrock for land surface modeling, J. Adv. Model. Earth Syst., 9, 65–88,
- 774 https://doi.org/10.1002/2016MS000686, 2017.
- 775 Smith, T. M., Shugart, H. H., and Woodward, F. I.: Plant functional types: their relevance to
- ecosystem properties and global change, Cambridge University Press, New York, 369 pp., 1997.
- 577 Smith, B., Prentice, I. C., & Sykes, M. T. : Representation of vegetation dynamics in the
- modelling of terrestrial ecosystems: Comparing two contrasting approaches within European
- climate space. Global Ecology & Biogeography, 10, 621–637, 2001.
- 780 Sterling, S. M., Ducharne, A., Polcher, J.: The impact of global land-cover change on the
- terrestrial water cycle. Nat. Clim. Chang., 3, 385–390. http://dx.doi.org/10.1038/nclimate1690,
 2013.
- 783 Still, C. J., and Berry, J. A.: Global distribution of C3 and C4 vegetation: Carbon cycle
- implications, Global Biogeochem. Cycles, 17(1), 1006, doi:10.1029/2001GB001807, 2003.
- Sun, W., Liang, S., Xu, G., Fang, H., and Dickinson, R. E.: Mapping plant functional types from
- MODIS data using multisource evidential reasoning, Remote Sens. Environ., 112, 1010–1024,
- 787 https://doi.org/10.1016/j.rse.2007.07.022, 2008.
- Swart, N. C., Cole, J. N. S., Kharin, V. V., Lazare, M., Scinocca, J. F., Gillett, N. P., Anstey, J.,
- Arora, V., Christian, J. R., Hanna, S., Jiao, Y., Lee, W. G., Majaess, F., Saenko, O. A., Seiler, C.,
- Seinen, C., Shao, A., Sigmond, M., Solheim, L., von Salzen, K., Yang, D., and Winter, B.: The
- 791 Canadian Earth System Model version 5 (CanESM5.0.3), Geosci. Model Dev., 12, 4823–4873,
- 792 https://doi.org/10.5194/gmd-12-4823-2019, 2019.

- 793 Tsendbazar, N. E., de Bruin, S., Mora, B., Schouten, L., Herold, M.: Comparative assessment of
- thematic accuracy of GLC maps for specific applications using existing reference data, Int. J.

Appl. Earth Obs. Geoinf., 44, 124–135, https://doi.org/10.1016/j.jag.2015.08.009, 2016.

- 796 Wulder, M.A., Nelson, T.: EOSD land cover classification legend report: Version 2, Natural
- 797 Resources Canada, Canadian Forest Service, Pacific Forestry Centre, Victoria, British Columbia,
- 798 Cana-da, 13 January 2003, 83pp., Available from
- 799 http://www.pfc.forestry.ca/eosd/cover/EOSD_Legend_Report-v2.pdf.
- 800 Wulder, M.A., Bater, C.W., Coops, N.C., Hilker, T., and White, J.C.: The role of LiDAR in
- sustainable forest management, The Forestry Chronicle, 84 (6), 807-826, 2008.
- 802 Wulder, M.A., White, J.C., Bater, C.W., Coops, N.C., Hopkinson, C., Chen, G.: Lidar plots a
- new large-area data collection option: context, concepts, and case study. Can. J. Remote. Sens.
- 804 38, 600–618, https://doi.org/10.5589/m12-049, 2012.
- 805 Verseghy, D. L.: CLASS—A Canadian Land Surface Scheme for GCMs, I. Soil model, Int. J.
- 806 Climatol., 11, 111–133, https://doi.org/10.1002/joc.3370110202, 1991.
- 807 Verseghy, D., McFarlane, N., and Lazare, M.: Class A Canadian land surface scheme for
- 608 GCMs. II: Vegetation model and coupled runs, Int. J. Climatol., 13, 347–370,
- 809 https://doi:10.1002/joc.3370130402, 1993.
- 810 Wang, A., Price, D. T., Arora, V. K.: Estimating changes in global vegetation cover (1850-2100)
- for use in climate models, Global Biogeochem. Cy., 20, https://doi:10.1029/2005GB002514,
- 812 2006.

| 813 | Wang, L., MacKay, M., Brown, R., Bartlett, P., Harvey, R., and Langlois, A.: Application of |
|-----|---|
| 814 | satellite data for evaluating the cold climate performance of the Canadian Regional Climate |
| 815 | model over Québec, Canada, J. Hydrometeorol., 15, 614–630, https://doi.org/10.1175/JHM-D- |
| 816 | 13-086.1, 2014. |

- 817 Wang, L., Cole, J. N.S., Bartlett, P., Verseghy, D., Derksen, C., Brown, R., von Salzen, K.:
- 818 Investigating the spread in surface albedo for snow-covered forests in CMIP5 models, J.
- 619 Geophys. Res. Atmos., 121, 1104–1119, https://doi:10.1002/2015JD023824, 2016.
- 820 Wang, L., Bartlett, P., Chan, E., and Xiao, M.: Mapping of Plant Functional Type from Satellite-
- 821 Derived Land Cover Datasets for Climate Models, In Proceedings of the 2018 IEEE International
- 822 Geoscience and Remote Sensing Symposium, Valencia, Spain, 22–27, https://DOI:
- 823 10.1109/IGARSS.2018.8518046, 2018.
- Wang, L., Bartlett, P., Pouliot, D., Chan, E., Lamarche, C., Wulder, M.A., Defourny, P., Brady,
- 825 M.: Comparison and Assessment of Regional and Global Land Cover Datasets for Use in
- 826 CLASS over Canada, Remote Sens., 11(19), 2286; https://doi.org/10.3390/rs11192286, 2019.
- Zakharova, L., Meyer, K. M., Seifan, M.: Trait-based modelling in ecology: A review of two
- decades of research, Ecological Modelling, 407,
- 829 https://doi.org/10.1016/j.ecolmodel.2019.05.008, 2019.
- 830
- 831
- 832
- 833
- 834

Table 1. Cross-walking table for mapping the 30 m Hybrid land cover map to CLASSIC PFTs in Canada.

836 Nine PFTs in CLASSIC: NLE - Needleleaf Evergreen trees, NLD - Needleleaf Deciduous trees, BLE -

- 837 Broadleaf Evergreen trees, BCD Broadleaf Cold Deciduous trees, BDD Broadleaf Dry Deciduous
- 838 trees, C3C C3 Crops, C4C C4 Crops, C3G C3 Grasses, and C4C C4 Grasses.

| ID | Map description | 1 NLE | 2 NLD | 3 BLE | 4+5 BCD BDD | 6+7 C3C C4C | 8+9 C3G C4G | Urban | Lake | Bare |
|-----|---|----------|----------|----------|-------------------|-------------------|-------------------|-------|------|------|
| 2 | Sub-polar taiga needleleaf forest | 0.20 | | | | | 0.60 | | | 0.20 |
| 11 | Sub-polar or polar shrubland-lichen- moss | | | | | | 0.65 | | | 0.35 |
| 12 | Sub-polar or polar grassland-lichen- moss | | | | | | 0.45 | | | 0.55 |
| 13 | Sub-polar or polar barren-lichen- moss | | | | | | 0.10 | | | 0.90 |
| 15 | Cropland | | | | | 1.0 | | | | |
| 16 | Barren lands | | | | | | | | | 1.0 |
| 17 | Urban | | | | | | | 1.0 | | |
| 20 | Water | | | | | | | | 1.0 | |
| 31 | Snow_ice | | | | | | | | | 1.0 |
| 32 | Rock_rubble | | | | | | | | | 1.0 |
| 50 | Shrubland | | | | 0.20 | | 0.60 | | | 0.20 |
| 80 | Wetland | | | | 0.05 | | 0.85 | | | 0.10 |
| 81 | Wetland-treed | 0.55 | | | 0.05 | | 0.35 | | | 0.05 |
| 100 | Herbs | | | | | | 0.80 | | | 0.20 |
| 210 | Coniferous | 1.0 | | | | | | | | |
| 220 | Broadleaf | | | | 1.0 | | | | | |
| 230 | Mixedwood | 0.50 | | | 0.50 | | | | | |

- **Table 2**. The sub-pixel fractional tree cover fraction for ESA-CCI (European Space Agency Climate
- 850 Change Initiative) land cover classes (with forest cover) based on the Hansen TCF (Tree Cover Fraction)
- dataset in Canada. Ratios of TCF between the main class and the closed class, and between the open class
- and the closed class are also included.

| ESA- CCI class | ESA-CCI class description | Tree cover Fraction (%) | Ratio of TCF relative to closed class |
|-------------------|---|----------------------------|---|
| 30 | Mosaic cropland (>50%) / natural vegetation (<50%) | 13.7 | |
| 40 | Mosaic natural vegetation (>50%) / cropland (<50%) | 45 | |
| 60 | Tree cover broadleaved deciduous closed to open (>15%) | 68.5 | 0.8 |
| 61 | Tree cover broadleaved deciduous closed (>40%) | 86.7 | 1 |
| 62 | Tree cover broadleaved deciduous open (15-40%) | 37.4 | 0.43 |
| 70 | Tree cover needleleaf evergreen closed to open (>15%) | 39.3 | 0.6 |
| 71 | Tree cover needleleaf evergreen, closed (>40%) | 61.7 | 1 |
| 72 | Tree cover needleleaf evergreen open (15-40%) | 23.2 | 0.38 |
| 90 | Tree cover Mixed | 80.9 | |
| 100 | Mosaic tree and shrub (>50%) / herbaceous cover (<50%) | 37.3 | |
| 110 | Mosaic herbaceous cover (>50%) / tree and shrub (<50%) | 19.6 | |
| 120 | Shrubland | 28.1 | |
| 150 | Sparse vegetation (tree shrub herbaceous cover) (< 15%) | 4 | |
| 160 | Tree cover, flooded fresh/brackish | 43 | |
| 180 | Shrub or herbaceous cover, flooded | 26.9 | |

- 861 Table 3. The sub-pixel fractional composition for ESA-CCI classes (columns, homogenous ESA-CCI
- 862 pixels) based on the Hybrid land cover map (rows) for dominant land cover classes in Canada. The

| 863 | fractions for NLE and BCD | are computed based | on equation (1). |
|-----|---------------------------|--------------------|------------------|
|-----|---------------------------|--------------------|------------------|

| Hybrid/ ESACCI Class | Hybrid description | 30 | 40 | 60 | 70 | 71 | 90 | 100 | 120 | 130 | 140 | 150 | 160 | 180 |
|----------------------------|--|------|------|------|------|------|------|------|------|------|------|------|------|------|
| 2 | Sub-polar taiga needleleaf forest | | | | 0.02 | | | 0.01 | | 0.01 | | | | |
| 11 | Sub-polar/polar shrubland-lichen- moss | | | | | | | | | | 0.01 | 0.05 | | |
| 12 | Sub-polar/polar grassland-lichen- moss | | | | 0.04 | | | | 0.03 | 0.01 | 0.24 | 0.27 | 0.03 | 0.04 |
| 13 | Sub-polar/polar barren-lichen-moss | | | | 0.02 | | | 0.01 | 0.02 | 0.01 | 0.34 | 0.09 | | 0.02 |
| 15 | Cropland | 0.92 | 0.37 | 0.02 | | | | | | 0.1 | | | | |
| 16 | Barren lands | | | | | | | | | 0.01 | 0.15 | 0.17 | | |
| 50 | Shrubland | 0.01 | 0.07 | 0.06 | 0.13 | 0.05 | 0.04 | 0.32 | 0.46 | 0.09 | 0.14 | 0.25 | 0.06 | |
| 80 | Wetland | | 0.03 | 0.08 | 0.2 | 0.05 | 0.03 | 0.27 | 0.2 | 0.02 | 0.06 | 0.09 | 0.37 | 0.75 |
| 81 | Wetland treed | | 0.01 | 0.01 | 0.17 | 0.07 | 0.03 | 0.11 | 0.12 | | | | 0.43 | 0.15 |
| 100 | Herbs | 0.06 | 0.27 | 0.08 | 0.02 | | 0.02 | 0.06 | 0.09 | 0.72 | 0.01 | 0.03 | 0.01 | 0.01 |
| 210 | Coniferous | | 0.01 | 0.02 | 0.29 | 0.72 | 0.07 | 0.04 | 0.03 | | 0.01 | 0.02 | 0.06 | |
| 220 | Broadleaf | 0.01 | 0.13 | 0.57 | 0.02 | 0.01 | 0.28 | 0.07 | 0.01 | 0.01 | | | 0.01 | |
| 230 | Mixedwood | | 0.1 | 0.14 | 0.09 | 0.07 | 0.52 | 0.12 | 0.03 | | | | 0.02 | |
| NLE | Needleleaf evergreen | | 0.07 | 0.09 | 0.44 | 0.8 | 0.32 | 0.19 | 0.16 | 0.01 | 0.02 | 0.05 | 0.31 | 0.08 |
| BCD | Broadleaf cold deciduous | 0.01 | 0.19 | 0.66 | 0.09 | 0.06 | 0.57 | 0.18 | 0.09 | 0.02 | 0.02 | 0.03 | 0.05 | 0.03 |

| ID | ESA-CCI class description | 1 NLE | 2 NLD | 3 BLE | 4+5 BCD BDD | 6+7 C3C C4C | 8+9 C3G C4G | Urban | Lake | Ocean | Bare |
|-----|--|----------|----------|----------|-------------------|-------------------|-------------------|-------|------|-------|------|
| 10 | Cropland, rainfed (CR) | | | | | 0.80 | 0.20 | | | | |
| 11 | CR Herbaceous cover | | | | | 0.90 | 0.10 | | | | |
| 12 | CR Tree or shrub cover | | | | 0.60 | | 0.30 | | | | 0.10 |
| 20 | Cropland, irrigated or post-flood | | | | 0.05 | 0.85 | 0.10 | | | | |
| 30 | Mosaic cropland (>50%) / natural vegetation (tree, shrub, herb) | 0.05 | | | 0.15 | 0.60 | 0.20 | | | | |
| 40 | Mosaic natural vegetation (tree,shrub, herb) >50% / crop | 0.10 | | | 0.20 | 0.40 | 0.30 | | | | |
| 50 | Tree cover broadleaved evergreen closed to open | | | 0.95 | 0.05 | | 0.0 | | | | |
| 60 | Tree cover broadleaved deciduous closed to open | | | | 0.70 | | 0.25 | | | | 0.05 |
| 61 | Tree cover broadleaved deciduous closed | | | | 0.90 | | 0.10 | | | | |
| 62 | Tree cover broadleaved deciduous open | | | | 0.40 | | 0.40 | | | | 0.20 |
| 70 | Tree cover needleleaf evergreen closed to open | 0.85 | | | 0.05 | | 0.10 | | | | |
| 71 | Tree cover needleleaf evergreen, closed | 0.85 | | | 0.05 | | 0.10 | | | | |
| 72 | Tree cover needleleaf evergreen open | 0.35 | | | 0.10 | | 0.40 | | | | 0.15 |
| 73 | Replace LC70 in Canada | 0.45 | | | 0.10 | | 0.30 | | | | 0.15 |
| 80 | Tree cover needleleaf deciduous closed to open | 0.05 | 0.40 | | 0.10 | | 0.35 | | | | 0.10 |
| 81 | Tree cover needleleaf deciduous closed | 0.05 | 0.80 | | 0.05 | | 0.15 | | | | |
| 82 | Tree cover needleleaf deciduous open | 0.05 | 0.30 | | 0.10 | | 0.45 | | | | 0.15 |
| 90 | Tree cover Mixed | 0.25 | 0.05 | | 0.60 | | 0.10 | | | | |
| 100 | Mosaic tree and shrub (>50%) / herbaceous cover (<50%) | 0.15 | 0.05 | | 0.20 | | 0.45 | | | | 0.15 |
| 110 | Mosaic herbaceous cover (>50%) / tree and shrub (<50%) | 0.05 | 0.05 | | 0.10 | | 0.65 | | | | 0.15 |
| 120 | Shrubland | | | | 0.30 | | 0.45 | | | | 0.25 |
| 121 | Shrubland evergreen | 0.15 | | 0.15 | | | 0.45 | | | | 0.25 |
| 122 | Shrubland deciduous | | | | 0.30 | | 0.45 | | | | 0.25 |

| 871 | Table 4. Cross-walking table for mapping ESA-CCI land cover dataset to CLASSIC PFTs. | |
|-----|--|--|
| | | |

| 120 | G 1 1 | <u>г</u> | 1 | 0.70 | | | | 0.00 |
|-----|--|----------|-------|------|------|------|------|------|
| 130 | Grassland | | | 0.70 | | | | 0.30 |
| 140 | Lichens and mosses | | | 0.20 | | | | 0.80 |
| 150 | Sparse vegetation (tree shrub herbaceous cover) (< 15%) | | 0.05 | 0.35 | | | | 0.60 |
| 151 | Sparse tree (<15%) | | 0.05 | 0.35 | | | | 0.60 |
| 152 | Sparse shrub (<15%) | | | 0.30 | | | | 0.70 |
| 153 | Sparse herbaceous cover (<15%) | | | 0.30 | | | | 0.70 |
| 160 | Tree cover, flooded fresh/brackish | 0.30 | 0.10 | 0.45 | | 0.1 | | 0.05 |
| 170 | Tree cover, flooded saline water | 0.30 | 0.10 | 0.40 | | | 0.1 | 0.10 |
| 180 | Shrub or herbaceous cover, flooded | 0.10 | 0.05 | 0.45 | | 0.15 | 0.15 | 0.10 |
| 190 | Urban areas | 0.02 5 | 0.025 | 0.15 | 0.75 | 0.05 | | |
| 200 | Bare areas | | | | | | | 1.0 |
| 201 | Consolidated bare areas | | | | | | | 1.0 |
| 202 | Unconsolidated bare areas | | | | | | | 1.0 |
| 210 | Water bodies | | | | | 1.0 | | |
| 220 | Permanent snow and ice | | | | | | | 1.0 |

- -

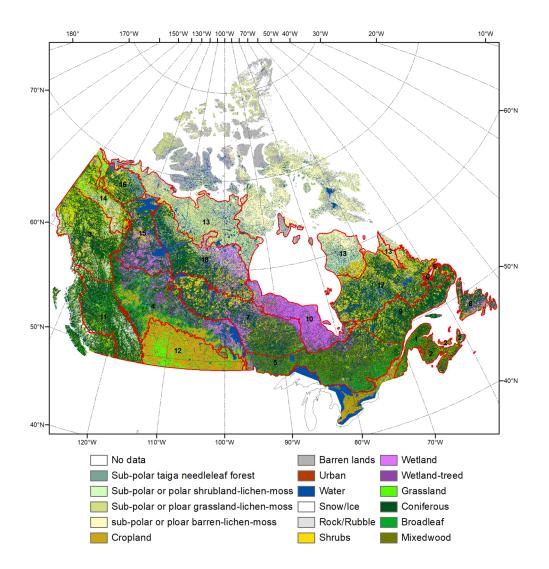


Figure 1. The Hybrid land cover map of Canada based on VLCE and NALCMS land cover maps for2010. The red polygons represent 18 ecozones used in this study.

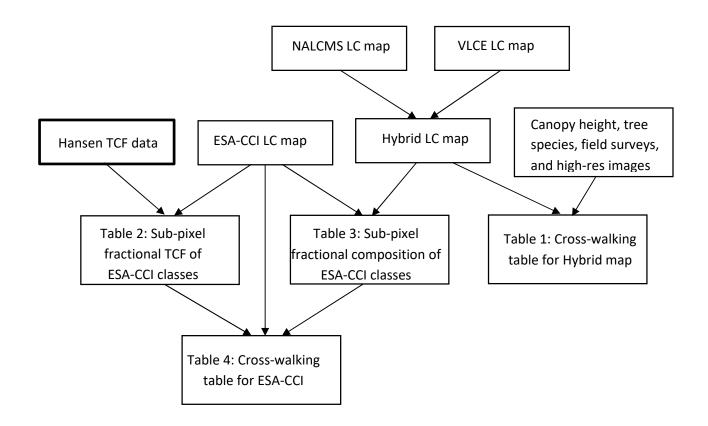
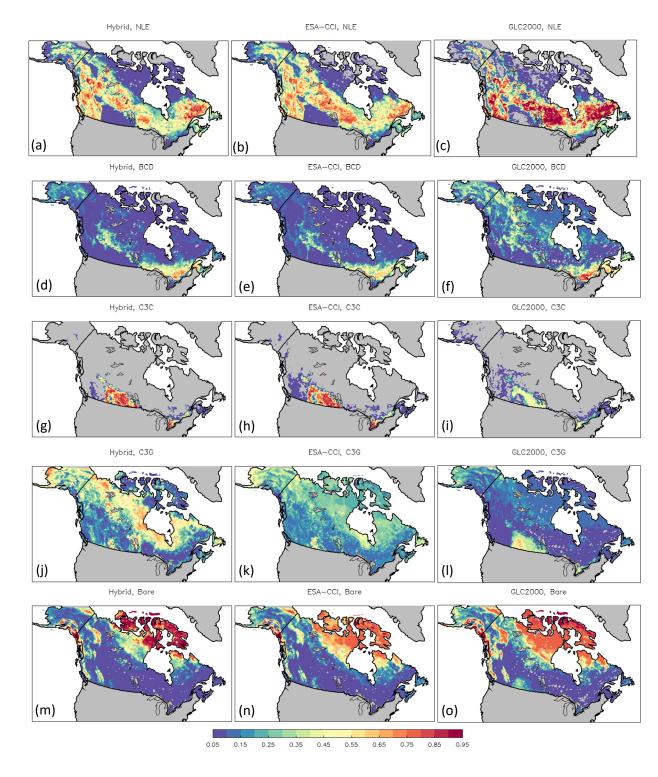


Figure 2. Schematic flow chart of the process for creating the cross-walking table for ESA-CCI land
cover (LC) dataset. NALCMS: the North America Land Change Monitoring System; VLCE: the Virtual
Land Cover Engine; TCF: Tree Cover Fraction.



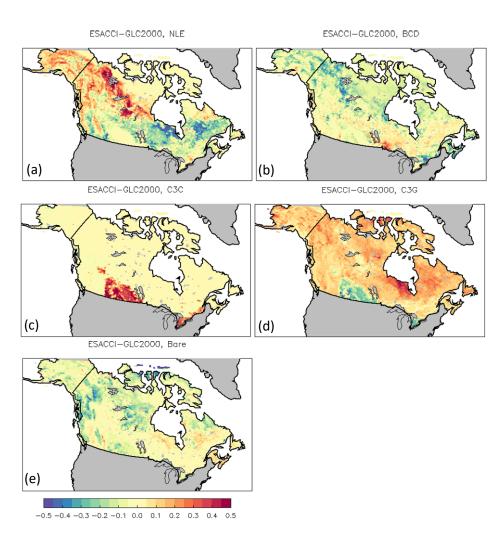
913

Figure 3. The spatial distribution of CLASSIC PFTs based on the Hybrid (left), ESA-CCI (middle), and

915 GLC2000 (right) land cover datasets respectively. The maps for C4C and C4G are not shown for their

916 fractions are small (0.5% for C4 crops and 0.1% for C4 grasses) in Canada. The last panel shows fractions

917 for bare ground from the three datasets.





- 936 (b) BCD, (c) C3C, (d) C3G, and (e) Bare.

- 5.0

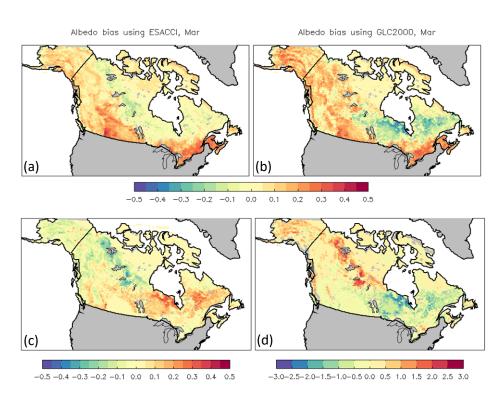




Figure 5. Surface albedo bias (relative to MODIS) in CLASSIC simulations using PFT distributions

- based on (a) ESA-CCI, and (b) GLC2000 land cover products. Panels (c) and (d) show the difference in
- 961 simulated surface albedo (c) and leaf area index (d) between the two simulations.