



1 **Modeling of non-structural carbohydrate dynamics by the spatially**
2 **explicitly individual-based dynamic global vegetation model SEIB-**
3 **DGVM (SEIB-DGVM-NSC ver1.0)**

4
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15 **Abstract.** Forest dynamics need to be considered when estimating the global carbon budget. The
16 alteration of forest structure and function under a changing climate and expanding human activity
17 could lead to a reduction of forest canopy cover and a spread of lower-biomass ecosystems in
18 warm and dry regions. Non-structural carbohydrate (NSC) acts as a storage buffer between carbon
19 supplied by assimilation and carbon consumed by, inter alia, respiration, reproduction, and pests.
20 Estimation of NSC concentrations in a tree is very important for accurate projection of future
21 forest dynamics. We developed a new NSC module for incorporation into a spatially explicit,
22 individual-based, dynamic global vegetation model (SEIB-DGVM) to validate the simulated NSC
23 dynamics with observations. NSC pools were simulated in three plant organs: leaves, trunk, and
24 roots. The seasonal dynamics of the NSCs varied among plant species, and the sizes of the NSC
25 pools inferred from observations differed between the boreal, temperate, and tropical climates.
26 The NSC models were therefore validated for each of the three climatic regions at both point and
27 global scales to assess the performance of the models. The modeled NSCs showed good
28 agreement in seasonality with the observed NSCs at four sites—Canada (boreal), Austria and
29 Switzerland (temperate), and Panama (tropical)—and in mean values for three climate zones
30 derived from the global NSC dataset. The SEIB-DGVM-NSCv1.0 is expected to enable
31 simulation of biome shifts caused by the changes of NSC dynamics worldwide. These dynamics
32 will contribute to changes of not only the global carbon cycle but also of forest structure and
33 demography at a global scale.

34

35



36 **1 Introduction**

37

38 Permanent shifts in forest vegetation dynamics have already been observed and are expected to
39 accelerate under future changes of climate globally (McDowell et al., 2020). Forest dynamics are
40 changing due to anthropogenic drivers, such as rising temperatures and CO₂ partial pressures, and
41 are affected by transient disturbances such as wildfires, droughts, biotic attacks, and land-use
42 changes. The dependence of tree recruitment and growth on spatial and temporal drivers could
43 lead to an increase of tree mortality rates in warm and dry regions (Stevens-Rumann et al., 2018;
44 Xu et al., 2017). These changes will cause forests to become shorter and younger. The result will
45 be a net reduction of forest canopy cover and a shift toward low-biomass ecosystems. Furthermore,
46 higher tree mortality will have a negative impact on global ecosystem: lower biological diversity,
47 a decrease of wild animal habitat, altered hydrological and carbon cycles, and increased
48 vulnerability to sudden invasions by exotic species (Adams et al., 2013). Understanding the
49 drivers of vegetation dynamics requires accurately simulating the effect of climate change on
50 global terrestrial biogeochemistry.

51 To increase their chance of survival, trees control their carbon resources and strategically
52 allocate them to growth, respiration, storage, reproduction, and defense (Hoch et al., 2003;
53 Hartmann et al., 2018). When the atmospheric partial pressure of CO₂ increases, trees can allocate
54 surplus carbon to either growth or carbon storage (Hoch et al., 2003; Huang et al., 2020). Changes
55 in tree carbon allocation patterns have been shown to exert large effects on constituents of the
56 terrestrial carbon cycle (Klein and Hoch, 2015). Clarification of the importance of carbon
57 allocation has revealed that non-structural carbohydrates (NSCs) draw much from the other
58 carbon resources because they are the most significant carbon compounds involved in the life
59 processes of trees (He et al., 2020).

60 The NSC is comprised of starch and sugars, which are mobilized mainly for growth and
61 metabolism when little recently assimilated carbon is available (Gough et al., 2010; Richardson
62 et al., 2013; Chuste et al., 2020; Herrera-Ramírez et al., 2020). During photosynthesis, freshly
63 assimilated carbon is transported as triose phosphate from the chloroplast to the cytosol, where
64 sucrose is synthesized from it. Some of the sucrose is then changed into starch in the chloroplast,
65 and the starch is consumed to maintain growth and metabolism at times when recently assimilated
66 carbon is not available to the plants (Dietze et al., 2014). Plants that seasonally shed leaves need
67 to rely on stored carbon for maintenance during the leafless season. NSCs play an important role
68 as substrates for the synthesis of compounds in plants and as energy sources for metabolic
69 activities (Hartmann et al., 2018). Moreover, NSCs include key compounds that are used to buffer
70 physiological stress when energy from photosynthesis does not satisfy metabolic demands
71 (Gough et al., 2010; Sala et al., 2012) because carbohydrates such as starch can be easily



72 mobilized and reallocated (Hartmann et al., 2018).

73 In long-lived plants, the ability to store carbon is a key to survival at times when
74 photosynthetic rates are low because of shade, drought, and disturbance (Martínez-Vilalta et al.,
75 2016). As a result, the amount of NSC depends on the balance between the supply and demand of
76 assimilated carbon and accounts for a large fraction of the annual carbon budget of plants
77 (Richardson et al., 2013). When carbon allocation patterns favor storage over growth, tree growth
78 is limited (Wang et al., 2021). Hence, the dynamics of stored carbon pools can be considered an
79 indicator of the carbon balance of the plant.

80 Because the frequency, duration, and severity of droughts are expected to increase globally,
81 the damage to plants through rising temperatures, water vapor pressure deficiency, and associated
82 water loss will also increase (IPCC, 2014; Sevanto and Dickman, 2015). Trees are killed directly
83 by drought or indirectly by associated increases of insect or pathogen attacks. Indirect effects that
84 cause tree mortalities include girdling of the phloem and xylem by bark beetles and defoliation
85 events that delay recovery of trees. The frequency and severity of this indirect biotic disturbance
86 from insects and insect–pathogen complexes have been increasing (McDowell et al., 2020; Seidl
87 et al., 2017). According to multiple observational and experimental studies, the resulting
88 imbalance between NSC demand and supply leads to carbon starvation, which is one of the
89 mechanisms that contribute to drought-induced mortality (McDowell, 2011).

90 The decline of stomatal conductance during a drought reduces photosynthetic carbon
91 assimilation and thus decreases the amount of NSC (McDowell et al., 2008; Adams et al., 2017).
92 Although an imbalance of the NSC pool could mechanistically trigger plant mortality, few
93 ecological models predict tree mortality resulting from the role of NSC associated with climate
94 change (Adams et al., 2013; McDowell, 2011). Simulations of the NSC dynamics of plants will
95 elucidate the effects of different drivers on forested ecosystems (Gough et al., 2010).

96 Dynamic global vegetation models (DGVMs) are often used to represent vegetation dynamics
97 as well as biogeochemical cycles and to simulate the transition of the vegetation structure in
98 response to climatic changes via modeling of competition and disturbance (Hickler et al., 2004;
99 Krinner et al., 2005; Braakhekke et al., 2019). In DGVMs, plant species are classified into plant
100 functional types (PFTs) based on their eco-physiological traits. However, most DGVMs
101 oversimplify individual plant competition by using average values of traits for each PFT (Smith
102 et al., 2001). Most of such models miss the effects of local competition for light, which must be
103 considered when modeling gap population dynamics among individual trees (Sato et al., 2007).

104 In contrast, the Spatially Explicit, Individual-Based, Dynamic Global Vegetation Model
105 (SEIB-DGVM; Sato et al., 2007) can simulate the growth of individual trees on numerous
106 replicate patches and enable observation of how single, large trees can influence nearby trees.
107 Plants in different patches do not interact with each other in terms of physical resources such as



108 light and water. In each patch, the growth, competitive interactions, and mortality of each tree are
109 calculated based on environmental conditions. Transient changes in vegetation distribution and
110 dynamics can therefore be examined (Sato et al., 2007).

111 SEIB-DGVM has been used to simulate a transient change in the distribution and function of
112 vegetation on the African continent in conjunction with the ranges of dispersal of trees and to
113 address factors that had a strong impact on the transient change (Sato and Ise, 2012). Use of the
114 SEIB-DGVM has enabled reconstruction of the geographical distributions of plant productivity
115 and thermo-hydrology based on observations in eastern Siberia and partial representation of the
116 effect of topography on the abundance of trees in larch forests (Sato et al., 2020). The SEIB-
117 DGVM was coupled with a flagship Earth system model (MIROC-ESM; Watanabe et al., 2011).
118 Because the amount of stored NSC depends on the size of individual trees and because the SEIB-
119 DGVM can simulate individual trees, we chose the SEIB-DGVM to estimate the NSC dynamics
120 of plants.

121 The objectives of the research were to 1) incorporate a module to simulate NSC dynamics in
122 the SEIB-DGVM and 2) validate the simulated NSC dynamics with observational data at both
123 point and global scales. We therefore created a new function in the SEIB-DGVM to represent the
124 NSC dynamics of individual trees. How NSC is produced, stored, and distributed among different
125 plant organs under environmental stress is poorly understood (Jones et al., 2019; Rademacher
126 et al., 2021; Wang et al., 2021). Our enhanced model improves the physiological simulation of the
127 leaf life cycle and enhances understanding of how NSC affects the distribution of vegetation,
128 gross primary production (GPP), net primary production (NPP), and carbon stocks as well as tree
129 dynamics (age, height, and trunk diameter) at global scales in the future. By adjusting the NSC
130 accumulation rates of individual trees and the threshold of NSC-induced mortality during drought,
131 the model can simulate the timing, location, and percentage of trees that die in response to
132 moderate drought. Furthermore, the model can increase our understanding of the role of carbon
133 reserves.

134

135 **2 Model**

136

137 **2.1 Spatially Explicit Individual-Based Dynamic Global Vegetation Model (SEIB-DGVM)**

138

139 The SEIB-DGVM (v3.02; http://seib-dgvm.com/data/seib_code302.zip; Sato et al., 2016) is a
140 carbon budget model that simulates the establishment of individual trees, competition between
141 trees, and the death of individual trees according to input climate data. The default settings follow
142 the structure of a three-dimensional virtual forest on a 30 m × 30 m stand of trees with 1 m × 1 m
143 simulation grid cells. In each grid cell, a tree belonging to one of 14 woody plant functional types



144 (PFTs) is assigned if conditions allow in addition to one of 2 grass PFTs. All physical and
 145 physiological processes are calculated at daily time steps, trunk growth is estimated monthly, and
 146 vegetation dynamics and disturbance are assessed annually. Because of the lack of field
 147 observations at the time the model was developed, there is no mechanism to control the carbon
 148 stock in leaves and roots in the original SEIB-DGVM, and only the carbon stock in trunks is
 149 simulated after adjusting the available organic matter for reproduction and respiration. The
 150 original SEIB-DGVM therefore cannot represent carbon starvation effectively without accounting
 151 for the carbon stock in leaves and roots, even though the carbon stock of the trunk depends on the
 152 leaf mass from the previous day.

153

154 2.2 NSC components

155

156 2.2.1 NSC pool

157 The new NSC pools are separated into three organs of an individual tree: leaves, trunk, and roots.
 158 The original SEIB-DGVM simply allocates the surplus carbon that remains after respiration for
 159 the growth of roots and then, if more than 30 days have passed since foliation, for the growth of
 160 leaves and the trunk. In this study, the carbon newly assimilated via photosynthesis goes into three
 161 NSC pools. The NSC pools can be later mobilized for growth and respiration as follows (Fig. 1).

162

$$163 \Delta NSC_t = NPP_t = \Delta NSC_{trunk,t} + \Delta NSC_{leaf,t} + \Delta NSC_{root,t}, \quad (1)$$

164

165 when $NSC_{trunk,t-1} < NSC_{trunk,max(t)}$, $NSC_{leaf,t-1} < NSC_{leaf,max(t)}$, and $NSC_{root,t-1} < NSC_{root,max(t)}$,

166

$$167 \begin{cases} \text{parent } NSC_{trunk,t} = \min(NSC_{trunk,max(t)}, \Delta NSC_t), \\ NSC_{leaf,t} = \min(NSC_{leaf,max(t)}, \Delta NSC_t - NSC_{trunk,t}), \\ NSC_{root,t} = \min(NSC_{root,max(t)}, \Delta NSC_t - NSC_{trunk,t} - NSC_{leaf,t}), \end{cases} \quad (2)$$

168

169 when $NSC_{trunk,t-1} < NSC_{trunk,max(t)}$, $NSC_{leaf,t-1} < NSC_{leaf,max(t)}$, and $NSC_{root,t-1} > NSC_{root,max(t)}$

170

$$171 \begin{cases} NSC_{trunk,t} = \min(NSC_{trunk,max(t)}, \Delta NSC_t), \\ NSC_{leaf,t} = \min(NSC_{leaf,max(t)}, \Delta NSC_t - NSC_{trunk,t}), \\ NSC_{root,t} = NSC_{root,max(t)} \end{cases}, \quad (3)$$

172

173 when $NSC_{trunk,t-1} < NSC_{trunk,max(t)}$, $NSC_{leaf,t-1} > NSC_{leaf,max(t)}$, and $NSC_{root,t-1} < NSC_{root,max(t)}$,

174



$$\begin{cases}
 \text{NSC}_{\text{trunk},t} = \min(\text{NSC}_{\text{trunk},\text{max}(t)}, \Delta \text{NSC}_t), \\
 \text{NSC}_{\text{leaf},t} = \text{NSC}_{\text{leaf},\text{max}(t)}, \\
 \text{NSC}_{\text{root},t} = \min(\text{NSC}_{\text{root},\text{max}(t)}, \Delta \text{NSC}_t - \text{NSC}_{\text{trunk},t} - \text{NSC}_{\text{leaf},t}),
 \end{cases} \quad (4)$$

176

177 when $\text{NSC}_{\text{trunk},t-1} < \text{NSC}_{\text{trunk},\text{max}(t)}$, $\text{NSC}_{\text{leaf},t-1} > \text{NSC}_{\text{leaf},\text{max}(t)}$, and $\text{NSC}_{\text{root},t-1} > \text{NSC}_{\text{root},\text{max}(t)}$,

178

$$\begin{cases}
 \text{parent NSC}_{\text{trunk},t} = \min(\text{NSC}_{\text{trunk},\text{max}(t)}, \Delta \text{NSC}_t), \\
 \text{NSC}_{\text{leaf},t} = \text{NSC}_{\text{leaf},\text{max}(t)}, \\
 \text{NSC}_{\text{root},t} = \text{NSC}_{\text{root},\text{max}(t)},
 \end{cases} \quad (5)$$

180

181 when $\text{NSC}_{\text{trunk},t-1} > \text{NSC}_{t,\text{max}(t)}$, $\text{NSC}_{\text{leaf},t-1} < \text{NSC}_{\text{leaf},\text{max}(t)}$, and $\text{NSC}_{\text{root},t-1} < \text{NSC}_{\text{root},\text{max}(t)}$,

182

$$\begin{cases}
 \text{NSC}_{\text{trunk},t} = \text{NSC}_{t,\text{max}(t)} \\
 \text{NSC}_{\text{leaf},t} = \min(\text{NSC}_{\text{leaf},\text{max}(t)}, \Delta \text{NSC}_t), \\
 \text{NSC}_{\text{root},t} = \min(\text{NSC}_{\text{root},\text{max}(t)}, \Delta \text{NSC}_t - \text{NSC}_{\text{trunk},t} - \text{NSC}_{\text{leaf},t}),
 \end{cases} \quad (6)$$

184

185 when $\text{NSC}_{\text{trunk},t-1} > \text{NSC}_{t,\text{max}(t)}$, $\text{NSC}_{\text{leaf},t-1} < \text{NSC}_{\text{leaf},\text{max}(t)}$, and $\text{NSC}_{\text{root},t-1} > \text{NSC}_{\text{root},\text{max}(t)}$,

186

$$\begin{cases}
 \text{NSC}_{\text{trunk},t} = \text{NSC}_{t,\text{max}(t)}, \\
 \text{NSC}_{\text{leaf},t} = \min(\text{NSC}_{\text{leaf},\text{max}(t)}, \Delta \text{NSC}_t), \\
 \text{NSC}_{\text{root},t} = \text{NSC}_{\text{root},\text{max}(t)},
 \end{cases} \quad (7)$$

188

189 when $\text{NSC}_{\text{trunk},t-1} > \text{NSC}_{t,\text{max}(t)}$, $\text{NSC}_{\text{leaf},t-1} > \text{NSC}_{\text{leaf},\text{max}(t)}$, and $\text{NSC}_{\text{root},t-1} < \text{NSC}_{\text{root},\text{max}(t)}$,

190

$$\begin{cases}
 \text{NSC}_{\text{trunk},t} = \text{NSC}_{t,\text{max}(t)}, \\
 \text{NSC}_{\text{leaf},t} = \text{NSC}_{\text{leaf},\text{max}(t)}, \\
 \text{NSC}_{\text{root},t} = \min(\text{NSC}_{\text{root},\text{max}(t)}, \Delta \text{NSC}_t),
 \end{cases} \quad (8)$$

192

193 when $\text{NSC}_{\text{trunk},t-1} > \text{NSC}_{t,\text{max}(t)}$, $\text{NSC}_{\text{leaf},t-1} > \text{NSC}_{\text{leaf},\text{max}(t)}$, and $\text{NSC}_{\text{root},t-1} > \text{NSC}_{\text{root},\text{max}(t)}$,

194

$$\begin{cases}
 \text{NSC}_{\text{trunk},t} = \text{NSC}_{t,\text{max}(t)}, \\
 \text{NSC}_{\text{leaf},t} = \text{NSC}_{\text{leaf},\text{max}(t)}, \\
 \text{NSC}_{\text{root},t} = \text{NSC}_{\text{root},\text{max}(t)},
 \end{cases} \quad (9)$$

196

197 where t is the calculation day, $t-1$ is the previous day, $\text{NSC}_{\text{organ}}$ is the amount of NSC in each organ,

198 and $\text{NSC}_{\text{organ},\text{max}(t)}$ is the maximum amount of NSC in each organ on day t .



199 First, the surplus carbon that remains after respiration is assigned to the NSC trunk pool using
200 Eq. (1). As long as the trunk biomass increases relative to that on the previous day, the overflowing
201 carbon moves primarily into NSC_{leaf} , secondarily into NSC_{root} , and finally into the growth of
202 leaves, the trunk, and roots. The sum of the NSCs in the leaves, trunk, and roots (the total NSC)
203 is maximized for each climate region (Table 1). If the total NSC exceeds this upper limit, the
204 surplus is directly consumed for the growth of each organ.

205
206

Table 1. Maximum volume of NSC pool

Maximum of NSC pool		
Boreal	10% of total biomass	Martínez-Vilalta et al. (2016)
Temperate	5% of total biomass	Hoch et al. (2003)
Tropical	9% of total biomass	Würth et al. (2005)

207

208 The NSC pools of the organs displays unique seasonality for each climatic zone. The NSC
209 seasonality of each organ varies among tree species mainly because the climate and surrounding
210 environment influence the capacity and utilization of NSCs in plants. To take into consideration
211 the fact that field observations of NSC covered whole seasons and that various plant species were
212 scarce, we classify NSC seasonality into three types: tropical, temperate, and boreal. Observations
213 from temperate forests showed that the NSC seasonal cycles were similar among the organs and
214 peaked around late spring–summer, although the NSC_{organ} differed in size (Hoch et al., 2003;
215 Richardson et al., 2013; Woodruff and Meinzer, 2011; Gruber et al., 2012). In contrast, the NSC
216 concentrations in the leaves of boreal trees peak in June (Sveinbjörnsson et al., 2010), and NSCs
217 in the fine roots increase until summer and then decline toward mid-summer and fall because of
218 the initiation of root growth (Landhäusser and Lieffers, 2003). All 14 woody PFTs of the SEIB-
219 DGVM are sorted into one of three NSC types, and carbon assimilated via photosynthesis was
220 allocated to the NSC pool of each organ in temperate and boreal PFTs as follows.

221

$$222 \quad NSC_{organ, max} = (a + b \times \text{daily GPP}) \times \text{Biomass}, \quad (10)$$

223

224 where the organ is either a leaf, trunk, or root, a is the minimum value, and b is the seasonality
225 parameter.

226 Tropical species have a different NSC seasonality from temperate and boreal species. The
227 NSC of leaves display a concave upward seasonal pattern that reaches its minimum in late spring–
228 early summer (Würth et al., 2005), which is a dry season when leaf production and flowering
229 deplete NSC pools. Singh and Srivastava (1986) have observed that the NSC of roots is at a
230 minimum level from July to September because the NSC pool is drained to enable survival of the



231 rainy season during that period. The amount of NSCs then increases toward winter, when the fine
 232 root biomass declines. Hence, Eq. (10), which is used for temperate and boreal forests, is
 233 inadequate for simulation of tropical forests because the NSC depends less on the seasonality of
 234 photosynthesis in the tropics. The size of the NSC pool of tropical species therefore accumulates
 235 as follows.

236

$$237 \text{NSC}_{\text{organ, max}} = (a + b) \times \text{Biomass}, \quad (11)$$

238

239 where the organ is either the leaf, trunk, or root, a is the minimum value, and b is the seasonality
 240 parameter.

241

242 2.2.2 NSC expenditure

243

244 2.2.2.1 Respiration

245

246 Normally, photosynthetically assimilated carbon is used for maintenance respiration without
 247 entering the carbon storage system. When the assimilated carbon is inadequate for maintenance
 248 respiration, the NSC compensates for the shortage. The NSC loss is allocated to each organ as
 249 follows.

250

$$251 \text{NSC}_{\text{leaf, t}} = \text{NSC}_{\text{leaf, t-1}} - R_{a, t-1} \times c_{\text{leaf}} \quad (12)$$

$$252 \text{NSC}_{\text{trunk, t}} = \text{NSC}_{\text{trunk, t-1}} - R_{a, t-1} \times c_{\text{trunk}} \quad (13)$$

$$253 \text{NSC}_{\text{root, t}} = \text{NSC}_{\text{root, t-1}} - R_{a, t-1} \times c_{\text{root}}, \quad (14)$$

254

255 where autotrophic respiration (R_a) is the difference between assimilated carbon and maintenance
 256 respiration and c is the allocation factor for NSC utilization ($c_{\text{leaf}} + c_{\text{trunk}} + c_{\text{root}} = 1$). If the total
 257 NSC equals the carbon shortfall, the NSC of all organs becomes zero. The allocation factors of
 258 NSC utilization depend on the climatic region (Table 2). If an $\text{NSC}_{\text{organ}}$ is inadequate to provide
 259 the allocated share of R_a , the other organs will supply the difference: the NSC_{leaf} is supplemented
 260 first from the $\text{NSC}_{\text{trunk}}$, and if that is not enough, from the NSC_{root} . Similarly, if any of the other
 261 $\text{NSC}_{\text{organ}}$ pools is unable to cover local shortages, the NSC pools of the remaining organs will
 262 balance the supply and demand.

263

264

Table 2. Allocation ratio (c) of NSC to organs

Organ	Boreal	Temperate	Tropical
Leaf	0.20	0.05	0.01



Trunk	0.60	0.90	0.98
Root	0.20	0.05	0.01

265

266 **2.2.2.2 Dormancy**

267

268 In SEIB-DGVM, deciduous PFTs have dormant and growth phases. When the PFTs enter the bud
269 flush phase, the NSC will be consumed. The NSC is allocated at the rates shown in Table 2.

270

271 **2.2.2.3 Turnover**

272

273 Part of the NSC pools of leaves and roots is transformed into litter at the same fractional rates as
274 in the turnover of general carbon pools for leaves and roots. This turnover is calculated at daily
275 steps, regardless of the phenology phase.

276

277 **2.2.2.4 Establishment**

278

279 The establishment process is performed on the last day of each simulation year in the SEIB-
280 DGVM. The characteristics of the PFT are determined by five bioclimatic parameters: (1) the
281 maximum temperature in the coldest month; (2) the maximum growing-degree day; (3) the
282 minimum growing-degree day; (4) the minimum photosynthetically active radiation; and (5) the
283 duration of drought. All new trees, independent of their PFT, start with a sapwood diameter of
284 0.01 m and heartwood diameter of 0.00 m. Initially, these new trees have no leaves or fine roots.
285 Their carbon cycle is therefore maintained by initial values of 250 g dry matter (DM) of
286 assimilated carbon and 250 g DM of NSC ($NSC_{leaf} = 10$ g, $NSC_{trunk} = 190$ g, and $NSC_{root} = 50$ g)
287 from the litter pool.

288

289 **2.3 Validation of NSC for local and global simulations**

290

291 Observational NSC data for model validation were derived from Martínez-Vilalta et al. (2016),
292 who reviewed 296 papers and summarized NSC dynamics in forests. Their data include total
293 NSCs in leaves, trunks, and roots of mature terrestrial plants from observations lasting at least
294 four months. These data were used for both the local-scale and the global-scale model validations
295 mixed with other sources described below.

296

297 **2.3.1 Validation at field points**

298



299 **2.3.1.1 Site descriptions**

300

301 For the point-scale validation, we selected locations where NSC data were available for all organs.
302 Four countries satisfied these criteria and were used to validate the simulated NSC content in the
303 plant organs: Canada (boreal), Austria and Switzerland (temperate), and Panama (tropical). At
304 each site, the seasonality of the NSC was measured for at least four months. We used local climate
305 data from meteorological stations gap-filled by corrected gridded climate reanalysis data as the
306 input at these sites. We ran the NSC module including the SEIB-DGVM with the location and
307 climate provided and compared the model output with the observation data.

308 The boreal site was located near Alder Flats, Alberta, Canada (52°58'N, 114°59'W) in 2000.
309 The site was dominated by boreal winter deciduous plants such as *Populus tremuloides*
310 (Landhäuser and Lieffers, 2003). One of the temperate sites was located in the timberline ecotone
311 at Mt. Patscherkofel to the south of Innsbruck, Austria (47°13'N, 11°27'E) in 2008 (Gruber et al.,
312 2011). Temperate conifer species such as *Pinus cembra* were the dominating tree species. The
313 mean annual air temperature is 2.5°C, with a maximum of 26.0°C in summer and a minimum of
314 -28.0°C in winter. The mean annual precipitation is 995 mm. The other temperate site was at the
315 Mont Noble, Canton Valais, Swiss Central Alps (46°12'N, 7°30'E) and was dominated by
316 temperate conifers (*P. cembra* L.; Hoch et al., 2003) in 2000. The mean annual precipitation is
317 630 mm. The growing season lasts from early May to early October, and the mean temperature
318 during that time is 6.5°C. The tropical site was located at the Parque Natural Metropolitano near
319 Panama City, the Republic of Panama in 1996 (85°8'N, 79°34'W; Würth et al., 2005). The site
320 has mixed cover with 17 dominant species, including *Cecropia longipes* and *Anacardium*
321 *excelsum*. The mean annual temperature is 27°C, and the mean annual precipitation was 2,120
322 mm during 1993–1995 measured at 1.6 km from the site with a canopy crane (data from the
323 Panama Canal Commission, meteorological and hydrological branch).

324

325 **2.3.1.2 Input climate data**

326

327 The SEIB-DGVM requires ten climatic variables as environmental drivers: air temperature, soil
328 temperature at a depth of 50 cm (soil layer 1), soil temperature at a depth of 100 cm (soil layer
329 2), soil temperature at a depth of 150 cm (soil layer 3), precipitation, shortwave radiation,
330 longwave radiation, wind velocity, specific humidity, and diurnal range of air temperature. The
331 input climate data were prepared by harmonizing a global reanalysis gridded climate dataset, the
332 WATCH Forcing Data ERA-Interim (WFDEI, 0.5 × 0.5 degrees, 1979–2016,
333 Weedon et al., 2018), and the climate generated by the SEIB generator (Tei et al., 2017), which
334 is the monthly observation-based climatic datasets produced by Climatic Research Unit (CRU



335 TS4.00, 0.5×0.5 degrees, 1901-2015, Harris et al., 2014) supplemented with the National
336 Centers for Environmental Prediction/National Center for Atmospheric Research
337 (NCEP/NCAR) daily climate datasets (Kalnay et al., 1996) for 1950, with local climatology
338 recorded at meteorological stations near the sites. Local climatology in Panama is measured at
339 the Parque Natural Metropolitano Canopy Crane meteorological station (1995–2019). The
340 climatology in Austria (1979–2008) and Switzerland (1979–2000) was derived from the closest
341 meteorological station to the field site under the European Climate Assessment (Klein et al.,
342 2002, <https://www.ecad.eu>). WFDEI data were used for the climatology in Canada, except for
343 precipitation data, which are measured in the Meteorological Service of Canada (1979–1984,
344 https://climate.weather.gc.ca/historical_data/search_historic_data_e.html).

345 The reanalysis of daily WFDEI and SEIB climate data included daily records, which were
346 corrected by regression models to local climate data. For temperature, humidity, and shortwave
347 radiation values, local climatology were used directly and the daily WFDEI data supplemented
348 by simple linear regression. Precipitation data and wind speeds were first adjusted to monthly and
349 then annual averages and then scaled as a correction. WFDEI precipitation data were scaled after
350 adjusting to the annual climatological precipitation of 995 mm in 2008 for Austria and 630 mm
351 in 2000 for Switzerland. Longwave radiation was calculated using harmonized temperatures and
352 humidities above (Brutsaert, 1975). Missing values were estimated via linear interpolation.
353 Because soil temperature data were unavailable for local sites and for WFDEI, soil layer
354 temperatures were calculated using the SEIB generator by regressing soil layer 1 on atmospheric
355 temperature, soil layer 2 on layer 1, and layer 3 on layer 2. In Austria, humidity data were available
356 from 2005. The WFDEI data were therefore used to estimate missing data via linear interpolation.
357 In Canada, no observational data were available, except for temperature and precipitation.
358 Precipitation in Canada was scaled with WFDEI data after adjusting to the total climatological
359 precipitation for 1979–1984, shortwave radiation was taken from the WFDEI, and humidity data
360 were harmonized in the same way as the humidity data in Austria.

361

362 **2.3.1.3 Simulation scheme**

363

364 To reach equilibrium conditions of the biomes, plant, and soil carbon pools, a 1000-year spin-up
365 simulation was performed by looping the climate data and atmospheric CO₂ concentrations
366 between 1979–2000. Building on the final conditions of the spin-up simulations, continuous
367 simulations corresponding to 1979–2001 in Canada, 1979–2008 in Austria, 1979–2000 in
368 Switzerland, and 1979–1995 in Panama were carried out, and the NSC dynamics were compared
369 with field data.

370



371 **2.3.2 Validation at a global scale**

372

373 In the global-scale simulation, the NSC seasonality in the SEIB-DGVM was validated using
374 CRU/NCEP/MIROC integrated data (0.5×0.5 degrees, 1850–2100, Tei et al., 2017, Watanabe et
375 al., 2011) as climatic input. SEIB-DGVM-NSC ver 1.0 is expected to simulate on future scenarios,
376 thus the different climate data that cover longer period than that of section 2.3.1.2. are used for
377 validation at a global scale. The SEIB-DGVM categorizes plant species into 16 PFTs for global-
378 scale simulations.

379 For the validation, we used all available NSC data from the given climate zone. The outputs
380 of the SEIB-DGVM include two boreal biome types (evergreen and deciduous forests), three
381 temperate biome types (conifer, broad-leaved evergreen, and deciduous forests), and two tropical
382 biome types (evergreen and deciduous forests), whereas the observations included two boreal
383 biome types (conifer and deciduous forests), three temperate biome types (conifer, evergreen, and
384 deciduous forests), and two tropical biome types (evergreen and deciduous forests). The model
385 outputs and observation data were compared for each climate zone. Global climate data were
386 available from 1850 to 2005. The first 30 years (1850–1880) were therefore looped for a 1000-
387 year spin-up simulation. After the spin-up, simulations were run for the period 1850–2005. The
388 NSC dynamics from the period 1975–2005 were used for model validation.

389

390 **2.4 Parameterization of NSC functions**

391

392 Hoch et al. (2003) have reported that the NSC_{leaf} of temperate trees varies between 7%–20% of
393 the total leaf DM. They determined the seasonal mean of the NSC_{trunk} in sapwood of temperate
394 deciduous trees and temperate evergreen trees to be $4.7\% \pm 0.1\%$ of DM and $1.8\% \pm 0.1\%$ of DM,
395 respectively. There were no significant seasonal differences. The mean NSC_{root} was less than 1.5%
396 of the root DM throughout the whole season (Gruber et al., 2012), and the total NSC of temperate
397 trees was around 4%–5% of the DM during the growing season (Gruber et al., 2011). For tropical
398 trees, the NSC_{trunk} and NSC_{root} were 8%–10% of their biomass, whereas the NSC_{leaf} fluctuated
399 within 5%–9% of leaf biomass (Würth et al., 2005). Landhäusser and Lieffers (2003) have
400 reported that the NSC_{root} of boreal trees, which is used to support leaf flush and root growth, is
401 3%–4% of their root mass. The stemwood NSC_{trunk} concentration is $\sim 18 \text{ mg g}^{-1}$ of the DM
402 (Carbone et al., 2013). Because of limited observational data, the parameters of the NSC processes
403 were derived mostly from the values observed at each site used for point-scale validation, and the
404 maximums of simulated NSCs were corrected so that they were in the range of measured NSCs.

405 First, the parameter b in Eq. (10) was determined from the percentage of NSC and the biomass
406 in January. The parameter a was then adjusted so that the fluctuations of the NSCs did not exceed



407 the measured mean seasonal NSCs. The parameter a in Eq. (10) controls the minimum amount of
 408 photosynthetically fixed carbon mobilized for the NSC pools. In temperate zones, the value of a
 409 differs before and after July, and the NSC peaks around mid-summer. In contrast, in tropical zones,
 410 the amount of NSC in leaves and trunks decreases throughout the spring–summer.

411 Because the NSCs depend on the environmental conditions at the field sites, which could not
 412 be incorporated into the global simulation, different parameters were used when validating point-
 413 scale and global-scale results. Tables 3 and 4 show the parameters used for validation. Parameter
 414 values unrelated to the NSC module remain at the default values of the SEIB-DGVM (Sato et al.,
 415 2007).

416

417 **Table 3.** Parameters of NSC pool size function for point-scale simulation

Organ	Canada	Austria	Switzerland	Panama
Leaf	$a: 0.4 \times 10^{-3}, b: 0.09$	$a: 0.065 \times 10^{-3}, b:$ 0.04 (Jul–Oct)	$a: 0.1 \times 10^{-3}, b: 0.13$ (Jul–Oct)	$a: 0.06, b: -0.15 \times$ 10^{-3} (Jun–Nov)
		$a: 0.135 \times 10^{-3}, b:$ 0.04 (others)	$a: 0.8 \times 10^{-3}, b: 0.13$ (others)	$a: 0.06, b: 0.15 \times 10^{-3}$ (others)
Trunk	$a: 0.03 \times 10^{-3}, b: 0.06$	$a: 0.05 \times 10^{-3}, b: 0.02$	$a: 0.01 \times 10^{-3}, b:$ 0.02	$a: 0.1, b: -0.35 \times 10^{-3}$ (Jun–Nov)
Root	$a: 0.06 \times 10^{-3}, b: 0.14$	$a: 0.01 \times 10^{-3}, b: 0.02$	$a: 0.003 \times 10^{-3}, b:$ 0.06	$a: 0.1, b: 0$ (others)
				$a: 0.04, b: 0.005$

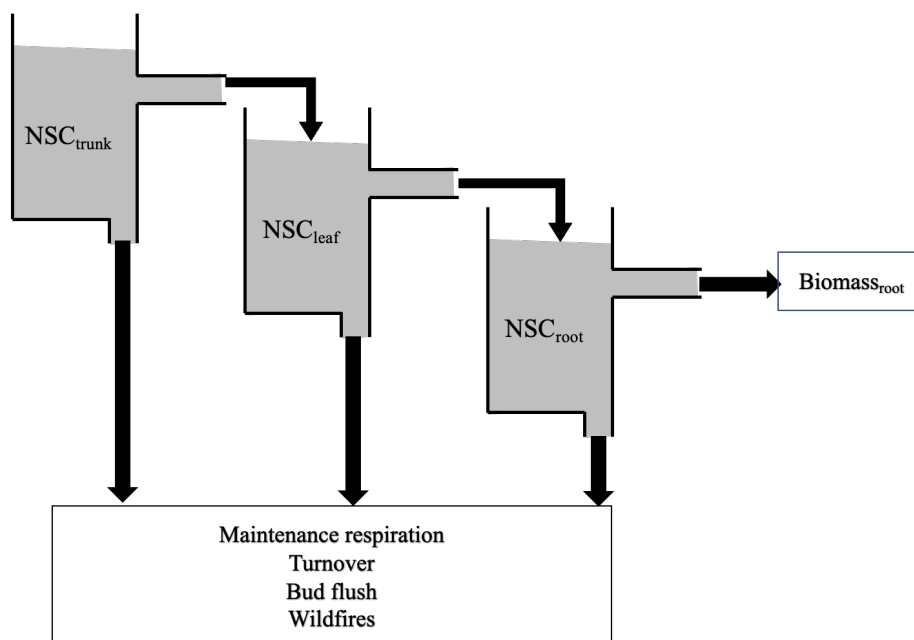
418

419

420 **Table 4.** Parameters of NSC pool size functions for global-scale simulation

Organ	Boreal	Temperate	Tropical
Leaf	$a: 0.4 \times 10^{-3}, b: 0.09$	$a: 0.1 \times 10^{-3}, b: 0.13$ (Jul–Oct)	$a: 0.06, b: -0.15 \times 10^{-3}$ (May– Jul)
		$a: 0.7 \times 10^{-3}, b: 0.13$ (others)	$a: 0.06, b: 0.15 \times 10^{-3}$ (others)
Trunk	$a: 0.03 \times 10^{-3}, b: 0.06$	$a: 0.01 \times 10^{-3}, b: 0.04$	$a: 0.1, b: -0.35 \times 10^{-3}$ (May– Jul)
Root	$a: 0.06 \times 10^{-3}, b: 0.14$	$a: 0.003 \times 10^{-3}, b: 0.06$	$a: 0.1, b: 0$ (others)
			$a: 0.06 \times 10^{-3}, b: 0.0145$

421



422

423 **Figure 1.** Schematic model structure of NSC pool.

424

425



426 **3 Results**

427

428 **3.1 Seasonality of NSC at the point scale**

429

430 **3.1.1 Canada**

431

432 The fact that the dominant PFT in the simulations was boreal deciduous trees was consistent with
433 observations at the site in Canada. The model simulated an increase in the NSC of leaves from 80
434 mg g^{-1} to 203 mg g^{-1} , whereas the observed NSC values were 89 mg g^{-1} in May and 185 mg g^{-1}
435 in August (Fig. 2a). The modeled NSCs in leaves therefore captured the increasing trend during
436 the growing season, but the simulations overestimated the maximum NSC a little. The observed
437 NSCs in trunks fluctuated from 90 mg g^{-1} to 192 mg g^{-1} during a year with no specific seasonal
438 trend (Fig. 2e). The model outputs in trunks were in the range 56–76 mg g^{-1} . Although the
439 observed fluctuations exceeded the modeled outputs, the modeled outputs were within one
440 standard deviation of the observations. The range of the observed NSCs in roots was 97–138 mg g^{-1}
441 g^{-1} , whereas the range of the modeled NSCs was 117–132 mg g^{-1} (Fig. 2i). However, the
442 observations peaked in August 2001 and in October 2002. The modeled NSCs of roots differed
443 from the observed NSCs because the former peaked during August in both years. Overall, the
444 simulated NSCs agreed well with the observed data (Fig. 3; RMSE = 69.92 mg g^{-1} , $r = 0.21$).

445

446 **3.1.2 Austria**

447

448 The fact that the dominant PFT in the simulations was temperate conifer forests was consistent
449 with observations at the site in Austria. The modeled NSCs in leaves accumulated until July with
450 a maximum of 142 mg g^{-1} . This pattern was similar to the observed seasonality of the NSCs,
451 which peaked at 150 mg g^{-1} (Fig. 2b). The modeled NSCs in trunks were stable in the range 19–
452 26 mg g^{-1} , and the observations were within the range 18–38 mg g^{-1} , with no specific seasonality
453 (Fig. 2f). The modeled values were interspersed between the observations. The modeled NSCs in
454 roots varied in a curvilinear manner from 18 to 26 mg g^{-1} , a range that was similar to the range of
455 the observed NSCs, 13–32 mg g^{-1} (Fig. 2j). The seasonality and magnitudes of the modeled NSCs
456 were consistent with observations (Fig. 3; RMSE = 9.52 mg g^{-1} , $r = 0.95$).

457

458 **3.1.3 Switzerland**

459

460 The dominant PFT in the simulations corresponded to the temperate conifers observed at the field
461 site. The NSCs in the tree leaves accumulated during early spring and reached up to 222 mg g^{-1}



462 (Fig. 2c). The decrease of the NSCs after July to a minimum of 135 mg g^{-1} was similar to the
463 decline of the observed NSCs to a minimum of 124 mg g^{-1} . The modeled NSCs in trunks fell in
464 the range $13\text{--}16 \text{ mg g}^{-1}$, which was overlapped with the range of the observed NSCs in trunks,
465 $15\text{--}33 \text{ mg g}^{-1}$ (Fig. 2g), and the modeled NSCs all fell within one standard deviation of the
466 observations. The modeled NSCs in roots increased gradually from 45 to 62 mg g^{-1} , which is
467 similar to the observed range of observations, $48\text{--}64 \text{ mg g}^{-1}$ (Fig. 2k). The simulations captured
468 the amounts and seasonal patterns of the NSCs in the different organs and produced results that
469 compared well with observations (Fig. 3; RMSE = 25.83 mg g^{-1} , $r = 0.91$).

470

471 **3.1.4 Panama**

472

473 While a wide range of woody species was found at the Panama site, in the simulation the tropical
474 evergreen PFT became dominant. The simulations showed that the NSCs in leaves were stored
475 during winter and were then gradually consumed from July to October, when they reached a
476 minimum of 52 mg g^{-1} (Fig. 2d). The observed NSCs in leaves likewise decreased from 69 to 48
477 mg g^{-1} between August and October. The model therefore followed the observed seasonality of
478 the leaf NSCs. The modeled NSCs in trunks fell in the range $35\text{--}73 \text{ mg g}^{-1}$ (Fig. 2h). The slight
479 decrease of the modeled NSCs in trunks during the summer was not apparent in the observations.
480 However, the simulated values fell within the range of the observed NSCs, $27\text{--}97 \text{ mg g}^{-1}$. The
481 simulated NSCs in roots fell in the range $23\text{--}55 \text{ mg g}^{-1}$; the observed NSCs ranged from 43 to 70
482 mg g^{-1} (Fig. 2l). Despite the weak correlation between simulated and observed NSCs, the model
483 results were within the acceptable margin of error (Fig. 3; RMSE = 20.75 mg g^{-1} , $r = 0.08$).

484

485 **3.2 Seasonality at a global scale**

486

487 For validation at a global scale, the mean annual NSCs from the model were compared with the
488 observed mean annual NSCs in boreal, temperate, and tropical regions (Table 5). The model
489 simulated the amounts of NSCs in forest tree trunks in all climate regions with high accuracy. The
490 modeled NSCs in the trunks of trees in boreal forests averaged $47.48 \pm 18.35 \text{ mg g}^{-1}$, which
491 compared favorably with the observed average of $76.67 \pm 23.68 \text{ mg g}^{-1}$. In temperate forests, the
492 modeled NSCs of trunks averaged $44.78 \pm 6.82 \text{ mg g}^{-1}$, which was close to the observed average of
493 $51.59 \pm 22.63 \text{ mg g}^{-1}$. The modeled NSCs of trunks in tropical forests averaged $66.68 \pm 18.79 \text{ mg g}^{-1}$,
494 which was close to the average of the observations, $106.23 \pm 32.52 \text{ mg g}^{-1}$. Although the modeled
495 NSCs in leaves of temperate and tropical forests were close to observed values, the modeled NSCs
496 in leaves of boreal forests underestimated the observed values. Moreover, the modeled NSCs in
497 roots of tropical forests were smaller than the observed NSCs. Overall, the simulated NSCs of all



498 organs of forest trees in all climate regions agreed reasonably well with observations (Fig. 4;
 499 RMSE = 66.75 mg g⁻¹, $r = 0.17$). The model could simulate the NSCs with high accuracy, with
 500 the exception of the NSCs of tree leaves in boreal forests and of tree roots in tropical forests (Fig.
 501 4; RMSE = 34.15 mg g⁻¹, $r = 0.71$).

502
 503 **Table 5.** Comparison of modelled and observed annual mean NSC concentrations (mg g⁻¹) on a
 504 global scale. The observed results are represented as the mean ± 1 standard deviation

	Boreal		Temperate		Tropical	
	Observation	Model	Observation	Model	Observation	Model
Leaf	202.80 ± 19.97	94.91 ± 42.91	127.10 ± 25.6	170.90 ± 46.54	86.42 ± 20.21	46.92 ± 16.20
Trunk	76.67 ± 23.68	47.48 ± 18.35	51.59 ± 22.63	44.78 ± 6.82	106.23 ± 32.52	66.68 ± 18.79
Root	118.49 ± 13.24	105.80 ± 40.82	67.65 ± 18.79	23.58 ± 10.57	170.40 ± 36.49	44.55 ± 15.15

505

506 3.3 Woody biomass and total NSCs on a global scale

507

508 The average of the total GPP simulated from the new model during 1976–2005 was 123 PgC
 509 year⁻¹. The model estimated the mean total woody biomass to be 282 PgC year⁻¹ in boreal zones,
 510 100 PgC year⁻¹ in temperate zones, and 337 PgC year⁻¹ in tropical zones globally during 1976–
 511 2005. In boreal zones such as North America and Russia, the mean concentration of total NSCs
 512 was 4.98% ± 1.87% of total woody biomass (Fig. 5). In temperate zones such as Asia, the mean
 513 concentration of total NSCs was 4.67% ± 0.54% of total woody biomass. Total NSCs of tropical
 514 forests in South America and Africa were 6.19% ± 1.66% of their total woody biomass. Mean
 515 values of the simulated total NSCs relative to total woody biomass were close to previous
 516 estimates for temperate and tropical forests (Table 6). The total NSCs of temperate, broad-leaved,
 517 evergreen forests were 4.63% ± 0.50% of the corresponding woody biomass reported by Smith et al.
 518 (2018). The total NSCs of temperate conifer forests were 4.72% ± 0.58% of total woody biomass,
 519 which was close to the figure of 4% reported by Körner (2003). According to Würth et al. (2005),
 520 the percentages of woody biomass contributed by NSCs are 4%–8% in tropical forests, 4.66% ±
 521 1.28% in tropical deciduous forests, and 7.11% ± 1.08% in tropical evergreen forests. These
 522 observed percentages are close to our simulated values.

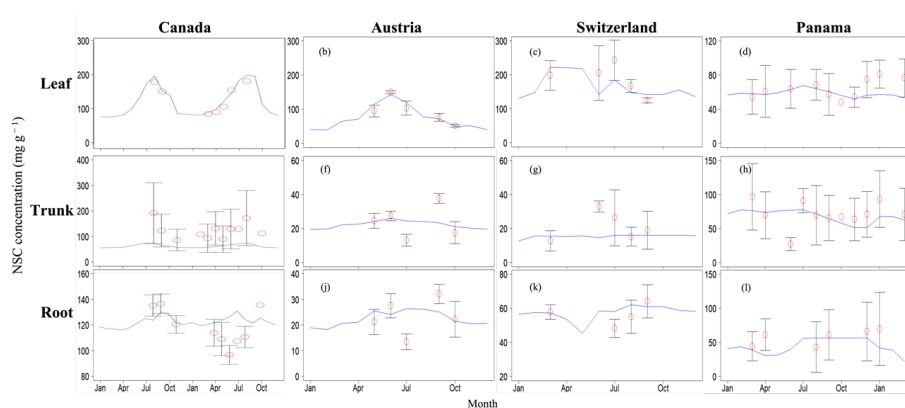
523

524 **Table 6.** Comparison of modeled mean NSC concentrations to observed total NSC concentrations (%)
 525 for different types of biomes. The simulated results are expressed as the mean ± 1 standard deviation



	Total NSC	Leaf	Trunk	Root	Observation
Boreal deciduous	3.41 ± 1.58	0.05 ± 0.09	3.06 ± 1.23	0.30 ± 0.62	
Boreal evergreen	6.06 ± 1.16	0.75 ± 0.38	4.73 ± 1.29	0.58 ± 0.37	
Temperate deciduous	2.30 ± 0.33	0.02 ± 0.01	2.25 ± 0.31	0.03 ± 0.01	1.0–12.5 (Gough et al., 2009)
Temperate broad-leaved evergreen	4.63 ± 0.50	0.49 ± 0.20	4.10 ± 0.56	0.04 ± 0.03	2.6–4.4 (Smith et al., 2018)
Temperate conifer	4.72 ± 0.58	0.89 ± 0.38	3.77 ± 0.73	0.08 ± 0.04	4.0 (Körner, 2003)
Tropical deciduous	4.66 ± 1.28	0.04 ± 0.03	4.60 ± 1.27	0.03 ± 0.02	4.0–8.0 (Würth et al., 2005)
Tropical evergreen	7.11 ± 1.08	0.08 ± 0.03	7.00 ± 1.08	0.02 ± 0.01	4.0–8.0 (Würth et al., 2005)

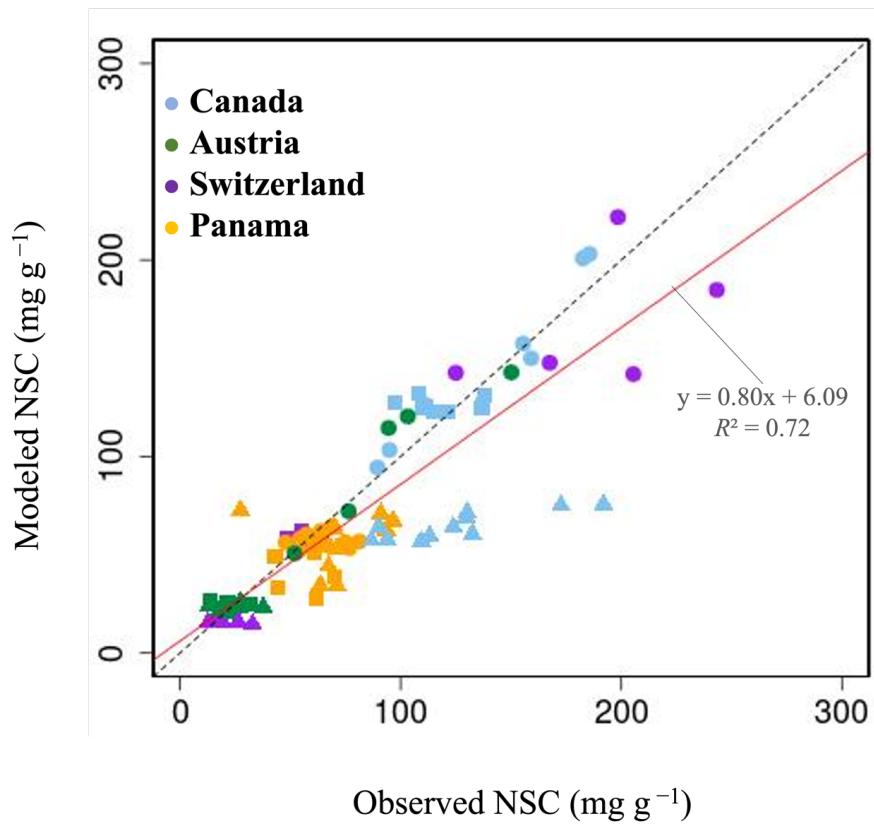
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527

528 **Figure 2.** Validation of the modeled NSC with observed NSC data (mg g^{-1}) at sites in Canada,
 529 Austria, Switzerland, and Panama. Red circles indicate the observed data, and blue lines indicate
 530 the modeled NSC. The observed results are represented as mean \pm 1 standard deviation.

531

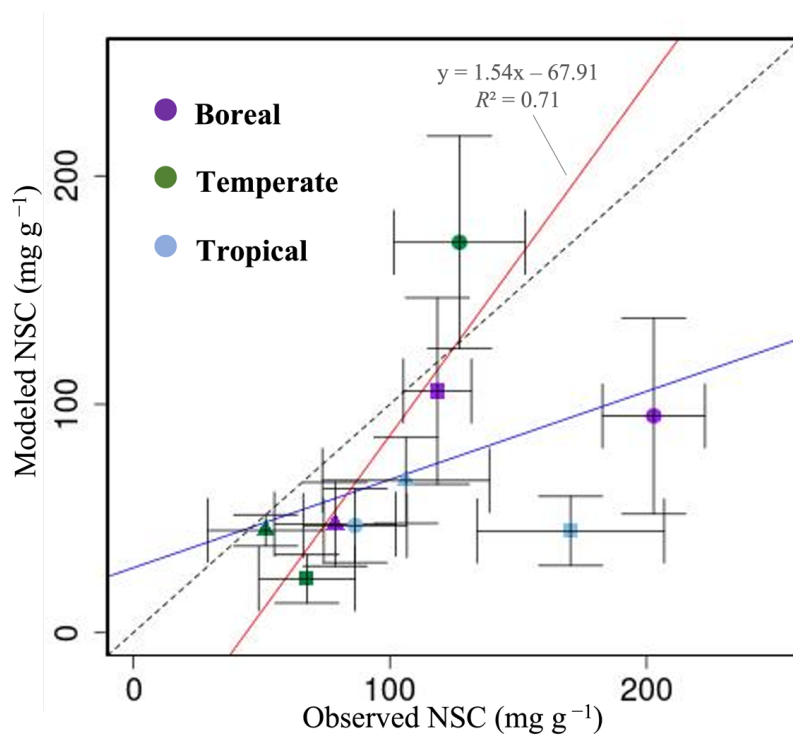


532

533 **Figure 3.** Plot of modeled NSC (mg g^{-1}) with observed NSC (mg g^{-1}) at a point scale. ●, leaves;

534 ▲, trunks; ■, roots. For all data, r is 0.72, and RMSE is 29.65 mg g^{-1} .

535

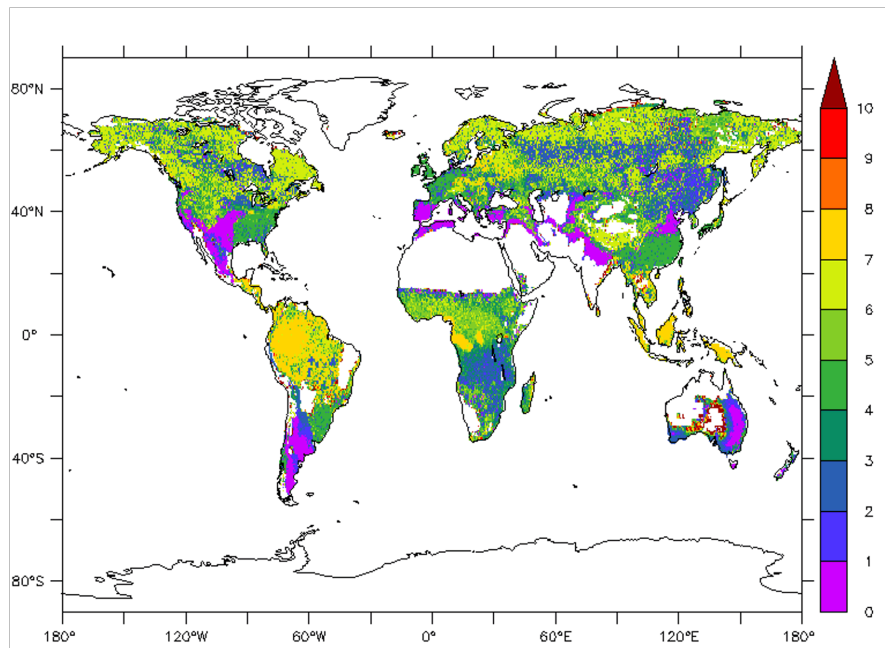


536

537 **Figure 4.** Plot of modeled NSC (mg g⁻¹) with observed NSC (mg g⁻¹) at a global scale. ●, leaves;

538 ▲, trunks; ■, roots.

539



540

541 **Figure 5.** Mean total NSC concentration relative to total woody biomass during 1976–2005 (%).

542

543



544 **4 Discussion**

545

546 On a point scale, the modeled NSCs for boreal forests in Canada were close to the observed
547 NSCs. The seasonality of the modeled NSCs in leaves was consistent with observations.
548 However, the seasonality of NSC in roots differed from the observations because there were
549 insufficient observations in boreal regions that enabled assessment of the seasonality of NSCs in
550 all organs. The seasonality of NSCs in roots is therefore still unclear. In temperate zones, the
551 model simulated the observed NSCs very accurately. The simulated NSCs of temperate forests
552 were close to observed values in Austria and Switzerland. The simulations showed that the
553 NSCs in leaves were consumed in winter for bud flush, and the leaves accumulated NSCs
554 during the growing season. This pattern corresponded to the seasonality reported in Asaadi et al.
555 (2018) and Furze et al. (2019). In the tropical zones, the model also captured a seasonality of
556 NSCs that was similar to observations. The NSC concentration in the canopy of tropical forests
557 decreased from June to August to satisfy increased maintenance demands (Signori-Müller et al.,
558 2022; Würth et al., 2005). The simulated NSCs in leaves followed a similar pattern from June to
559 August, and the simulated NSCs in leaves, trunks, and roots were close to observed values.

560 On a global scale, the simulated NSC values of all climate regions agreed with the observed
561 data, except for the leaves of boreal forests and roots of tropical forests. The new function,
562 which was validated at the point scale, could therefore perform well on a global scale. The
563 NSCs in trunks, which contain the greatest amounts of carbon in trees, were simulated
564 accurately in all climate regions. The new function could therefore calculate the total NSC
565 storage in trees with great accuracy.

566 The model with the new function calculated the global GPP to be $123 \text{ PgC year}^{-1}$, which is
567 close to the previous estimates of $106.2 \pm 2.9 \text{ PgC year}^{-1}$ by Zheng et al. (2020) and 130 ± 1.6
568 PgC year^{-1} by Madani et al. (2020). Moreover, the simulated mean total woody biomass for
569 boreal forests was $282 \text{ PgC year}^{-1}$, which is within the range of $249\text{--}295 \text{ PgC year}^{-1}$ reported by
570 Pan et al. (2011). The simulated woody biomass of $100 \text{ PgC year}^{-1}$ for temperate forests was
571 within the observed range of $59\text{--}139 \text{ PgC year}^{-1}$ (Hui et al., 2020) and a little lower than the
572 range of $113\text{--}125 \text{ PgC year}^{-1}$ for other temperate forests (Pan et al., 2011). The calculated total
573 woody biomass of $337 \text{ PgC year}^{-1}$ for tropical forests was within the range of $212\text{--}340 \text{ PgC}$
574 year^{-1} reported by Hui et al. (2020) and was not very different from the estimates of $378\text{--}564$
575 PgC year^{-1} by Pan et al. (2011), and $200\text{--}300 \text{ PgC year}^{-1}$ by Mitchard (2018). Furthermore, the
576 total NSCs relative to total biomass output from the new function for temperate and tropical
577 biome types agreed with previous research. The total NSC of boreal biome types could not be
578 compared with observations due to lack of data.

579 The new model has a high potential to simulate various biotic effects on terrestrial ecosystems



580 more accurately by calculating the NSC dynamics within each plant organ. The NSCs stored in
581 the trunk and roots compensate for the deficit of CO₂ uptake in trees under stress, and the NSC
582 stored in roots is potentially indispensable for tree recovery after disturbances (Herrera-Ramírez
583 et al., 2020). Furthermore, the NSC changes in the trunk and roots are better indicators of carbon
584 source–sink relationships under elevated CO₂ conditions than the NSC changes in leaves because
585 the NSC concentrations in leaves increase under elevated CO₂ conditions irrespective of growth
586 conditions. The NSC changes in the trunk and roots are therefore more tied to the carbon balance
587 of plant bodies (Körner, 2003). Simulation of the dynamics of NSC in the three compartments in
588 this research contributes to understanding plant growth and the response of carbon dynamics in
589 each organ to increasing atmospheric CO₂.

590 Carbon starvation may also be the main cause of plant death during drought when
591 photosynthesis decreases and water stress increases (McDowell et al., 2008). If reduced
592 photosynthetic rates cannot supply enough carbon for NSC accumulation during drought, there
593 will be greater canopy dieback in the next season. Plants therefore prioritize NSC storage, even
594 when no excess carbon is available (Hartmann et al., 2020). At times, plants are unable to allocate
595 carbon for this NSC defense mechanism, and their reduced ability to recover from biotic attacks
596 such as defoliation caused by insect pests leads to decreased growth rates, less restoration of NSC,
597 and lower survival rates. These processes may culminate in broad-scale tree mortalities
598 (McDowell, 2011; Chen et al., 2017).

599 The new model introduced NSC compartments in leaves, trunk, and roots that were validated
600 at the point and global scales. Use of the model developed here enabled simulation of the
601 environmental effects on forests resulting from the changing amount of NSC in each organ. The
602 simulations depicted the NSC changes in the trunk especially well. The model could thus be used
603 as an indicator of the carbon cycle in terrestrial ecosystems to understand the effect of climate
604 change. Simulation of photosynthetic carbon allocated into NSC storage in leaves, trunks, and
605 roots enables a more dynamic simulation of the carbon cycle between terrestrial ecosystems and
606 the atmosphere.

607 However, there were still some limitations to this research. We considered two potential
608 limitations that could lead to some discrepancies between the modeled and measured NSC values.
609 First, the relatively coarse spatial resolution of 0.5°×0.5° gridded climate data at the global scale
610 could not depict the details of local climates derived from observations. These differences were
611 especially important in the case of precipitation and soil properties, which play a key role in NSC
612 dynamics. Lowering of the water level in soils causes damage to the hydraulic conductivity of the
613 phloem tube, which leads to a decline of phloem conductance at the stem level (Dannoura et al.,
614 2018; Sevanto, 2014). This cessation of phloem transport could change the allocation of
615 photosynthetic products to plant growth and affect the ability of the trees to accumulate NSCs



616 (Dannoura et al., 2018).
617 Second, the scarcity of ground-measured NSC seasonality prevented us from having more average
618 information on NSC concentrations, especially in the tropical and boreal regions, where there
619 were fewer available data. The NSC seasonality differs between biome types, but because it is
620 difficult to measure NSC dynamics, there is a lack of long-term data for each biome type. Hence,
621 parameters were tuned to simulate the same seasonality for all the biome types in a given climate
622 zone in our study. The fact that the NSC allocation was further influenced by environmental
623 conditions caused the allocation patterns to change within the same biome type. The NSC
624 allocation to roots was favored over aboveground allocations when soil resources were lacking,
625 and tree size was considered an important determinant of carbon allocation as well as aridity
626 (Hartmann et al., 2020). Because we pooled data for each organ from different measurement sites
627 for global-scale validation, environmental effects on the data could not be measured. In addition
628 to the above factors, the number of samples and duration of observations differed between the
629 various studies. These differences led to no explicit NSC seasonality. These potential sources of
630 error in the field measurements jeopardized the model performances.

631

632 **5 Conclusions**

633

634 In this study, a new NSC model was incorporated into the SEIB-DGVM to understand the effect
635 of NSC allocation on global forest dynamics through competition and establishment among
636 individual trees. The new module calculated the NSC dynamics of three organs—leaves, trunk,
637 and roots—and the general NSC seasonality based on ground measurements was determined for
638 biome types in three climate zones: boreal, temperate, and tropical. The NSC seasonality was
639 validated at four sites: Canada (boreal), Austria and Switzerland (temperate), and Panama
640 (tropical). The mean values of simulated NSC concentration agreed reasonably well with
641 observed data on a global scale.

642 The model enabled us to simulate the biotic effects resulting from insufficient NSC caused
643 by factors such as carbon starvation and insect pests that are otherwise difficult to measure in
644 terrestrial ecosystems globally. The difference of the NSC dynamics in the organs under
645 elevated CO₂ conditions highlighted the importance of modeling the organs separately when
646 studying environmental stresses. As more observation data about NSC dynamics become
647 available, the model can be further improved and can contribute to the simulations of the
648 passive biome shifts that may occur globally.

649

650 *Code and data availability*

651 The model code used in this study is archived at <https://doi.org/10.5281/zenodo.7021459>.



652

653 *Author contributions.* T.K. conceived and supervised this study and acquired the funding. H.N.
654 developed the model code and carried out the analysis and produced the figures. H.N. prepared
655 the original draft, and T.K., and L.V. reviewed it. L.W. prepared the modeling environment. All
656 authors have read and agreed to the published version of the manuscript.

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658 *Competing interests.* The authors declare that they have no conflicts of interest.

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