

1 **Modeling of non-structural carbohydrate dynamics by the spatially**
2 **explicitly individual-based dynamic global vegetation model SEIB-**
3 **DGVM (SEIB-DGVM-NSC ver1.0)**

4
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14
15 **Abstract.** Forest dynamics need to be considered when estimating the global carbon budget. The
16 alteration of forest structure and function under a changing climate and expanding human activity
17 could lead to a reduction of forest canopy cover and a spread of lower-biomass ecosystems in
18 warm and dry regions. Non-structural carbohydrate (NSC) acts as a storage buffer between carbon
19 supplied by assimilation and carbon consumed by, inter alia, respiration, reproduction, and pests.
20 Estimation of NSC concentrations in a tree is very important for accurate projection of future
21 forest dynamics. We developed a new NSC module for incorporation into a spatially explicit,
22 individual-based, dynamic global vegetation model (SEIB-DGVM) to validate the simulated NSC
23 dynamics with observations. NSC pools were simulated in three plant organs: leaves, trunk, and
24 roots. The seasonal dynamics of the NSCs varied among plant species, and the sizes of the NSC
25 pools inferred from observations differed between the boreal, temperate, and tropical climates.
26 The NSC models were therefore validated for each of the three climatic regions at both point and
27 global scales to assess the performance of the models. The modeled NSCs showed good
28 agreement in seasonality with the observed NSCs at four sites—Canada (boreal), Austria and
29 Switzerland (temperate), and Panama (tropical)—and in mean values for three climate zones
30 derived from the global NSC dataset. The SEIB-DGVM-NSCv1.0 is expected to enable
31 simulation of biome shifts caused by the changes of NSC dynamics worldwide. These dynamics
32 will contribute to changes of not only the global carbon cycle but also of forest structure and
33 demography at a global scale.

36 **1 Introduction**

37

38 Permanent shifts in forest vegetation dynamics have already been observed and are expected to
39 accelerate under future changes of climate globally (McDowell et al., 2020). Forest dynamics are
40 changing due to anthropogenic drivers, such as rising temperatures and CO₂ partial pressures, and
41 are affected by transient disturbances such as wildfires, droughts, biotic attacks, and land-use
42 changes. The dependence of tree recruitment and growth on anthropogenic drivers and transient
43 disturbances could lead to an increase of tree mortality rates in warm and dry regions (Stevens-
44 Rumann et al., 2018; Xu et al., 2017). These changes will cause forests to become shorter and
45 younger. The result will be a net reduction of forest canopy cover and a shift toward low-biomass
46 ecosystems. Furthermore, higher tree mortality will have a negative impact on global ecosystem:
47 lower biological diversity and altered hydrological and carbon cycles (Adams et al., 2013).
48 Understanding the drivers of vegetation dynamics requires accurately simulating the effect of
49 climate change on global terrestrial biogeochemistry.

50 To increase their chance of survival, trees control their carbon resources and strategically
51 allocate them to growth, respiration, storage, reproduction, and defense (Hoch et al., 2003;
52 Hartmann et al., 2018). When the atmospheric partial pressure of CO₂ increases, trees can allocate
53 surplus carbon to either growth or carbon storage (Hoch et al., 2003; Huang et al., 2020). Changes
54 in tree carbon allocation patterns have been shown to exert large effects on constituents of the
55 terrestrial carbon cycle (Klein and Hoch, 2015). Clarification of the importance of carbon
56 allocation has revealed that non-structural carbohydrates (NSCs) draw much from the other
57 carbon resources because they are the most significant carbon compounds involved in the life
58 processes of trees (He et al., 2020).

59 The NSC is comprised of starch and sugars, which are mobilized mainly for growth and
60 metabolism when sink strength exceeds source activity (Gough et al., 2010; Richardson et al.,
61 2013; Chuste et al., 2020; Herrera-Ramírez et al., 2020). During photosynthesis, freshly
62 assimilated carbon is transported as triose phosphate from the chloroplast to the cytosol, where
63 sucrose is synthesized from it. Some of the sucrose is then changed into starch in the chloroplast,
64 and the starch is consumed to maintain growth and metabolism at times when recently assimilated
65 carbon is not available to the plants (Dietze et al., 2014). Plants that seasonally shed leaves need
66 to rely on stored carbon for maintenance during the leafless season. NSCs play an important role
67 as substrates for the synthesis of compounds in plants and as energy sources for metabolic
68 activities (Hartmann et al., 2018). Moreover, NSCs include key compounds that are used to buffer
69 physiological stress when energy from photosynthesis does not satisfy metabolic demands
70 (Gough et al., 2010; Sala et al., 2012) because carbohydrates such as starch can be easily
71 mobilized and reallocated (Hartmann et al., 2018).

72 In long-lived plants, the ability to store carbon is a key to survival at times when
73 photosynthetic rates are low because of shade, drought, and disturbance (Martínez-Vilalta et al.,
74 2016). As a result, the amount of NSC storage or remobilization depends on the balance between
75 the supply and demand of assimilated carbon and accounts for a large fraction of the annual carbon
76 budget of plants (Richardson et al., 2013). When carbon allocation patterns favor storage over
77 growth, tree growth is limited (Wang et al., 2021). Hence, the dynamics of stored carbon pools
78 can be considered an indicator of the carbon balance of the plant.

79 The decline of stomatal conductance during a drought reduces photosynthetic carbon
80 assimilation and thus decreases the amount of NSC (McDowell et al., 2008; Adams et al., 2017).
81 Although an imbalance of the NSC pool could mechanistically trigger plant mortality, few
82 ecological models predict tree mortality resulting from the role of NSC associated with climate
83 change (Adams et al., 2013; McDowell, 2011). Simulations of the NSC dynamics of plants will
84 elucidate the effects of different drivers on forested ecosystems (Gough et al., 2010).

85 Because the frequency, duration, and severity of droughts are expected to increase globally,
86 the damage to plants through rising temperatures, water vapor pressure deficit, and associated
87 water loss will also increase (IPCC, 2014; Sevanto and Dickman, 2015). Trees can be killed
88 directly by drought, i.e. vial desiccation, or indirectly by associated increases of insect or pathogen
89 attacks. Indirect effects that cause tree mortality include girdling by bark beetles and defoliation
90 events. The frequency and severity of this indirect biotic disturbance from insects and insect–
91 pathogen complexes have been increasing (McDowell et al., 2020; Seidl et al., 2017). According
92 to multiple observational and experimental studies, the resulting imbalance between NSC demand
93 and supply leads to carbon starvation, which is one of the mechanisms that contribute to drought-
94 induced mortality (McDowell, 2011).

95 Dynamic global vegetation models (DGVMs) are often used to represent vegetation dynamics
96 as well as biogeochemical cycles and to simulate the transition of the vegetation structure in
97 response to climatic changes via modeling of competition and disturbance (Hickler et al., 2004;
98 Krinner et al., 2005; Braakhekke et al., 2019). In DGVMs, plant species are classified into plant
99 functional types (PFTs) based on their eco-physiological traits. However, most DGVMs
100 oversimplify individual plant competition by using average values of traits for each PFT (Smith
101 et al., 2001). Most of such models miss the effects of local competition for light, which must be
102 considered when modeling gap population dynamics among individual trees (Sato et al., 2007).

103 In contrast, the Spatially Explicit, Individual-Based, Dynamic Global Vegetation Model
104 (SEIB-DGVM; Sato et al., 2007) can simulate the growth of individual trees on numerous
105 replicate patches and enable observation of how single, large trees can influence nearby trees.
106 Plants in different patches do not interact with each other in terms of physical resources such as
107 light and water. In each patch, the growth and mortality of each tree, and competitive interactions

108 between individual trees are calculated based on environmental conditions. Transient changes in
109 vegetation distribution and dynamics can therefore be examined (Sato et al., 2007). Because the
110 amount of stored NSC depends on the size of individual trees and because the SEIB-DGVM can
111 simulate individual trees, we chose the SEIB-DGVM to estimate the NSC dynamics of plants.

112 SEIB-DGVM has been used to simulate a transient change in the distribution and function of
113 vegetation on the African continent in conjunction with the ranges of dispersal of trees and to
114 address factors that had a strong impact on the transient change (Sato and Ise, 2012). Use of the
115 SEIB-DGVM has enabled reconstruction of the geographical distributions of plant productivity
116 and thermo-hydrology based on observations in eastern Siberia and partial representation of the
117 effect of topography on the abundance of trees in larch forests (Sato et al., 2020). The SEIB-
118 DGVM was coupled with an earth system model (MIROC-ESM; Watanabe et al., 2011).

119 However, the original SEIB-DGVM lacks the ability to compute NSC levels in trees, which
120 means it cannot simulate the plant death caused by an imbalance in the NSC pool, as well as
121 indirect impacts like insect infestations and defoliation. This also hinders the ability to investigate
122 the effects of various drivers on forest ecosystems, such as the intensification of drought, and
123 limits the development of the MIROC-ESM for simulating the global carbon cycle.

124 The objectives of the research were to 1) incorporate a module to simulate NSC dynamics in
125 the SEIB-DGVM and 2) validate the simulated NSC dynamics with observational data at both
126 point and global scales. We therefore created a new function in the SEIB-DGVM to represent the
127 NSC dynamics of individual trees. How NSC is produced, stored, and distributed among different
128 plant organs under environmental stress is poorly understood (Jones et al., 2019; Rademacher et
129 al., 2021; Wang et al., 2021). Our enhanced model improves the physiological simulation of the
130 leaf life cycle and enhances understanding of how NSC affects the distribution of vegetation,
131 gross primary production (GPP), and net primary production (NPP) as well as tree dynamics (age,
132 height, and trunk diameter) at global scales in the future. By adjusting the NSC accumulation
133 rates of individual trees and the threshold of NSC-induced mortality during drought, the model
134 can simulate the timing, location, and percentage of trees that die in response to moderate drought.
135 Furthermore, the model can increase our understanding of the role of NSCs.

136

137 **2 Model**

138

139 **2.1 Spatially Explicit Individual-Based Dynamic Global Vegetation Model (SEIB-DGVM)**

140

141 The SEIB-DGVM (v3.02; http://seib-dgvm.com/data/seib_code302.zip; Sato et al., 2016) is a
142 carbon budget model that simulates the establishment of individual trees, competition between
143 trees, and the death of individual trees according to input climate data. The default settings follow

144 the structure of a three-dimensional virtual forest on a 30 m × 30 m stand of trees with 1 m × 1 m
 145 simulation grid cells. In each grid cell, a tree belonging to one of 14 woody plant functional types
 146 (PFTs) is assigned depending on conditions, in addition to one of 2 grass PFTs. All physical and
 147 physiological processes are calculated at daily time steps, trunk growth is estimated monthly, and
 148 vegetation dynamics and disturbance such as wildfire and heat stress are assessed annually.
 149 Because of the lack of field observations at the time the model was developed, there is no
 150 mechanism to control the NSC in leaves and roots in the original SEIB-DGVM, and only the NSC
 151 in trunks is simulated after adjusting the available organic matter for reproduction and respiration.
 152 The original SEIB-DGVM therefore cannot represent NSC-induced effects on forest ecosystems
 153 realistically without accounting for the NSC in leaves and roots.

154

155 2.2 NSC components

156

157 2.2.1 NSC pool

158 The new NSC pools are separated into three organs of an individual tree: leaves, trunk, and roots.
 159 In the original SEIB-DGVM, the NSC in trunks is supplemented to 250 g dry matter (DM) from
 160 the litter after seed establishment and is based on the existing leaf biomass after the first 30 days
 161 of the growing season. It is used for foliation after the dormant phase and metabolic processes
 162 such as the synthesis of a storage organ and remobilization of the nutrients within it. In the new
 163 model, the carbon newly assimilated via photosynthesis goes into three NSC pools. The NSC
 164 pools can be later mobilized for growth and respiration as follows (Fig. 1).

165

$$166 \quad \Delta NSC_t = NPP_t = \Delta NSC_{trunk,t} + \Delta NSC_{leaf,t} + \Delta NSC_{root,t}, \quad (1)$$

167

168 when $NSC_{trunk,t-1} < NSC_{trunk,max(t)}$, $NSC_{leaf,t-1} < NSC_{leaf,max(t)}$, and $NSC_{root,t-1} < NSC_{root,max(t)}$,

169

$$170 \quad \begin{cases} \text{parent } NSC_{trunk,t} = \min(NSC_{trunk,max(t)}, \Delta NSC_t), \\ NSC_{leaf,t} = \min(NSC_{leaf,max(t)}, \Delta NSC_t - NSC_{trunk,t}), \\ NSC_{root,t} = \min(NSC_{root,max(t)}, \Delta NSC_t - NSC_{trunk,t} - NSC_{leaf,t}), \end{cases} \quad (2)$$

171

172 when $NSC_{trunk,t-1} < NSC_{trunk,max(t)}$, $NSC_{leaf,t-1} < NSC_{leaf,max(t)}$, and $NSC_{root,t-1} > NSC_{root,max(t)}$

173

$$174 \quad \begin{cases} NSC_{trunk,t} = \min(NSC_{trunk,max(t)}, \Delta NSC_t), \\ NSC_{leaf,t} = \min(NSC_{leaf,max(t)}, \Delta NSC_t - NSC_{trunk,t}), \\ NSC_{root,t} = NSC_{root,max(t)} \end{cases}, \quad (3)$$

175

176 when $NSC_{trunk, t-1} < NSC_{trunk, max(t)}$, $NSC_{leaf, t-1} > NSC_{leaf, max(t)}$, and $NSC_{root, t-1} < NSC_{root, max(t)}$,

177

$$178 \begin{cases} NSC_{trunk, t} = \min(NSC_{trunk, max(t)}, \Delta NSC_t), \\ NSC_{leaf, t} = NSC_{leaf, max(t)}, \\ NSC_{root, t} = \min(NSC_{root, max(t)}, \Delta NSC_t - NSC_{trunk, t} - NSC_{leaf, t}), \end{cases} \quad (4)$$

179

180 when $NSC_{trunk, t-1} < NSC_{trunk, max(t)}$, $NSC_{leaf, t-1} > NSC_{leaf, max(t)}$, and $NSC_{root, t-1} > NSC_{root, max(t)}$,

181

$$182 \begin{cases} \text{parent } NSC_{trunk, t} = \min(NSC_{trunk, max(t)}, \Delta NSC_t), \\ NSC_{leaf, t} = NSC_{leaf, max(t)}, \\ NSC_{root, t} = NSC_{root, max(t)}, \end{cases} \quad (5)$$

183

184 when $NSC_{trunk, t-1} > NSC_{t, max(t)}$, $NSC_{leaf, t-1} < NSC_{leaf, max(t)}$, and $NSC_{root, t-1} < NSC_{root, max(t)}$,

185

$$186 \begin{cases} NSC_{trunk, t} = NSC_{t, max(t)} \\ NSC_{leaf, t} = \min(NSC_{leaf, max(t)}, \Delta NSC_t), \\ NSC_{root, t} = \min(NSC_{root, max(t)}, \Delta NSC_t - NSC_{trunk, t} - NSC_{leaf, t}), \end{cases} \quad (6)$$

187

188 when $NSC_{trunk, t-1} > NSC_{t, max(t)}$, $NSC_{leaf, t-1} < NSC_{leaf, max(t)}$, and $NSC_{root, t-1} > NSC_{root, max(t)}$,

189

$$190 \begin{cases} NSC_{trunk, t} = NSC_{t, max(t)}, \\ NSC_{leaf, t} = \min(NSC_{leaf, max(t)}, \Delta NSC_t), \\ NSC_{root, t} = NSC_{root, max(t)}, \end{cases} \quad (7)$$

191

192 when $NSC_{trunk, t-1} > NSC_{t, max(t)}$, $NSC_{leaf, t-1} > NSC_{leaf, max(t)}$, and $NSC_{root, t-1} < NSC_{root, max(t)}$,

193

$$194 \begin{cases} NSC_{trunk, t} = NSC_{t, max(t)}, \\ NSC_{leaf, t} = NSC_{leaf, max(t)}, \\ NSC_{root, t} = \min(NSC_{root, max(t)}, \Delta NSC_t), \end{cases} \quad (8)$$

195

196 when $NSC_{trunk, t-1} > NSC_{t, max(t)}$, $NSC_{leaf, t-1} > NSC_{leaf, max(t)}$, and $NSC_{root, t-1} > NSC_{root, max(t)}$,

197

$$198 \begin{cases} NSC_{trunk, t} = NSC_{t, max(t)}, \\ NSC_{leaf, t} = NSC_{leaf, max(t)}, \\ NSC_{root, t} = NSC_{root, max(t)}, \end{cases} \quad (9)$$

199

200 where t is the calculation day, $t-1$ is the previous day, NSC_{organ} is the amount of NSC in each organ,
201 and $NSC_{organ, max(t)}$ is the maximum amount of NSC in each organ on day t .

202 The NSC pools of the organs displays unique seasonality for each climatic zone. The NSC
203 seasonality of each organ varies among tree species mainly because the climate and surrounding
204 environment influence the capacity and utilization of NSCs in plants. To take into consideration
205 the fact that field observations of NSC covered whole seasons and that various plant species were
206 scarce, we classify NSC seasonality into three types: tropical, temperate, and boreal. Observations
207 from temperate forests showed that the NSC seasonal cycles were similar among the organs and
208 peaked around late spring–summer, although the NSC_{organ} differed in size (Hoch et al., 2003;
209 Richardson et al., 2013; Woodruff and Meinzer, 2011; Gruber et al., 2012). In contrast, the NSC
210 concentrations in the leaves of boreal trees peak in June (Sveinbjörnsson et al., 2010), and NSCs
211 in the fine roots increase until summer and then decline toward mid-summer and fall because of
212 the initiation of root growth (Landhäusser and Lieffers, 2003). All 14 woody PFTs of the SEIB-
213 DGVM are sorted into one of three NSC types, and carbon assimilated via photosynthesis was
214 allocated to the NSC pool of each organ in temperate and boreal PFTs as follows.

215

$$216 \quad NSC_{organ, max} = (a + b \times \text{daily GPP}) \times \text{Biomass}, \quad (10)$$

217

218 where the organ is either a leaf, trunk, or root, a is the minimum value, and b is the seasonality
219 parameter.

220 Tropical species have a different NSC seasonality from temperate and boreal species. The
221 NSC of leaves display a concave upward seasonal pattern that reaches its minimum in late spring–
222 early summer (Würth et al., 2005), which is a dry season when leaf production and flowering
223 deplete NSC pools. Singh and Srivastava (1986) have observed that the NSC of roots is at a
224 minimum level from July to September because the NSC pool is drained to enable survival of the
225 rainy season during that period. The amount of NSCs then increases toward winter, when the fine
226 root biomass declines. Hence, Eq. (10), which is used for temperate and boreal forests, is
227 inadequate for simulation of tropical forests because the NSC depends less on the seasonality of
228 photosynthesis in the tropics. The size of the NSC pool of tropical species therefore accumulates
229 as follows.

230

$$231 \quad NSC_{organ, max} = (a + b) \times \text{Biomass}, \quad (11)$$

232

233 where the organ is either the leaf, trunk, or root, a is the minimum value, and b is the seasonality
234 parameter.

235 First, the surplus carbon that remains after respiration is assigned to the NSC_{trunk} using Eq.
 236 (1). Once NSC_{trunk} has reached its maximum capacity, the rest of assimilated carbon moves
 237 primarily into NSC_{leaf} , secondarily into NSC_{root} . Finally, any remaining carbon is allocated to the
 238 growth of leaves, the trunk, and roots. The sum of the NSCs in the leaves, trunk, and roots (the
 239 total NSC) is maximized in relation to total biomass for each climate region (Table 1). In cases
 240 where the total NSC exceeds this upper limit, the surplus is directly consumed for the growth of
 241 each organ.

242

243

Table 1. Maximum volume of NSC pool

Maximum of NSC pool		
Boreal	10% of total biomass	Martínez-Vilalta et al. (2016)
Temperate	5% of total biomass	Hoch et al. (2003)
Tropical	9% of total biomass	Würth et al. (2005)

244

245 2.2.2 NSC expenditure

246

247 2.2.2.1 Respiration

248

249 Normally, photosynthetically assimilated carbon is used for maintenance respiration without
 250 entering the NSC pool. When the assimilated carbon is insufficient for maintenance respiration,
 251 the NSC compensates for the shortage. The NSC loss is allocated to each organ as follows.

252

$$253 \quad NSC_{leaf, t} = NSC_{leaf, t-1} - R_{a, t-1} \times c_{leaf} \quad (12)$$

$$254 \quad NSC_{trunk, t} = NSC_{trunk, t-1} - R_{a, t-1} \times c_{trunk} \quad (13)$$

$$255 \quad NSC_{root, t} = NSC_{root, t-1} - R_{a, t-1} \times c_{root}, \quad (14)$$

256

257 where autotrophic respiration (R_a) is the difference between assimilated carbon and maintenance
 258 respiration and c is the allocation factor for NSC utilization ($c_{leaf} + c_{trunk} + c_{root} = 1$). If the total
 259 NSC equals the carbon shortfall, the NSC of all organs becomes zero. If an NSC_{organ} is insufficient
 260 to provide the allocated share of R_a , the other organs will supply the difference: the NSC_{leaf} is
 261 supplemented first from the NSC_{trunk} , and if that is not enough, from the NSC_{root} . Similarly, if any
 262 of the other NSC_{organ} pools is unable to cover local shortages, the NSC pools of the remaining
 263 organs will balance the supply and demand. When the total NSC is not enough to pay for the
 264 charges, a 1% of reduction in the biomass of all of the living organs occurs. The removed biomass
 265 of sapwood is transformed into heartwood, while the removed biomass of other organs is placed
 266 into the litter pool. The allocation factors of NSC utilization depend on the climatic region (Table

267 2), and have been adjusted to prevent the allocated share of R_a from hindering an increase in
268 NSC_{organ} during spin-up simulations.

269

270

Table 2. Allocation ratio (c) of NSC to organs

Organ	Boreal	Temperate	Tropical
Leaf	0.20	0.05	0.01
Trunk	0.60	0.90	0.98
Root	0.20	0.05	0.01

271

272 **2.2.2.2 Phenology**

273

274 In SEIB-DGVM, every deciduous PFT has two phenology phases: a growth phase and a dormant
275 phase. The NSCs are consumed for foliation after the dormant phase. The NSC is allocated at the
276 rates shown in Table 2.

277

278 **2.2.2.3 Turnover**

279

280 Part of the NSC pools of leaves and roots is transformed into litter at the same fractional rates as
281 in the turnover of general carbon pools for leaves and roots. This turnover is calculated at daily
282 steps, regardless of the phenology phase.

283

284 **2.2.2.4 Establishment**

285

286 The establishment process is performed on the last day of each simulation year in the SEIB-
287 DGVM. The characteristics of the PFT are determined by five bioclimatic parameters: (1) the
288 maximum temperature in the coldest month; (2) the maximum growing-degree day; (3) the
289 minimum growing-degree day; (4) the minimum photosynthetically active radiation; and (5) the
290 duration of drought. All new trees, independent of their PFT, start with a sapwood diameter of
291 0.01 m and heartwood diameter of 0.00 m. Initially, these new trees have no leaves or fine roots.
292 Their carbon cycle is therefore maintained by initial values of 250 g DM of assimilated carbon
293 and 250 g DM of NSC ($NSC_{leaf} = 10$ g, $NSC_{trunk} = 190$ g, and $NSC_{root} = 50$ g) from the litter pool.

294

295 **2.3 Validation of NSC for point and global simulations**

296

297 Observational NSC data for model validation were derived from Martínez-Vilalta et al. (2016),
298 who reviewed 296 papers and summarized NSC dynamics in forests. Their data include total

299 NSCs in leaves, trunks, and roots of mature terrestrial plants from observations over at least four
300 months. The new model was first validated at the point scale. After confirming that the model can
301 accurately simulate at the point scale, it was then validated at the global scale. These data were
302 used for both point-scale and global-scale model validations. For the point-scale simulation,
303 several field sites were selected, which NSC data were available for all organ. At each site, the
304 seasonality of the NSC was measured for at least four months. The modeled outputs were then
305 compared to the observed data, which were calculated using local climate data on a grid that
306 corresponds to the field site. For the global-scale simulation, global mean NSC values derived by
307 using all observed data from the relevant climate zones. The modeled outputs calculated using
308 $0.5^{\circ} \times 0.5^{\circ}$ gridded climate data were then compared to these values.

309

310 **2.3.1 Validation at a point scale**

311

312 **2.3.1.1 Site descriptions**

313

314 Four countries were used to validate the simulated NSC content in the plant organs: boreal
315 (Canada), temperate (Austria and Switzerland), and tropical (Panama). We used local climate data
316 from meteorological stations gap-filled by corrected gridded climate reanalysis data as the input
317 at these sites. We ran the NSC module including the SEIB-DGVM with the location and climate
318 provided and compared the model output with the observation data.

319 The boreal site is located near Alder Flats, Alberta, Canada ($52^{\circ}58'N$, $114^{\circ}59'W$) in 2000. The
320 site was dominated by boreal winter deciduous plants such as *Populus tremuloides* (Landhäusser
321 and Lieffers, 2003). One of the temperate sites is located in the timberline ecotone at Mt.
322 Patscherkofel to the south of Innsbruck, Austria ($47^{\circ}13'N$, $11^{\circ}27'E$) in 2008 (Gruber et al., 2011).
323 Temperate conifer species such as *Pinus cembra* were the dominating tree species. The other
324 temperate site is at the Mont Noble, Canton Valais, Swiss Central Alps ($46^{\circ}12'N$, $7^{\circ}30'E$) and was
325 dominated by temperate conifers (*P. cembra* L.; Hoch et al., 2003) in 2000. The tropical site is
326 located at the Parque Natural Metropolitano near Panama City, the Republic of Panama in 1996
327 ($85^{\circ}8'N$, $79^{\circ}34'W$; Würth et al., 2005). The site has mixed cover with 17 dominant species,
328 including *Cecropia longipes* and *Anacardium excelsum*.

329

330 **2.3.1.2 Input climate data**

331

332 The SEIB-DGVM requires ten climatic variables as environmental drivers: air temperature, soil
333 temperature at a depth of 50 cm (soil layer 1), soil temperature at a depth of 100 cm (soil layer
334 2), soil temperature at a depth of 150 cm (soil layer 3), precipitation, shortwave radiation,

335 longwave radiation, wind velocity, specific humidity, and diurnal range of air temperature. The
336 input climate data were prepared by harmonizing a global reanalysis gridded climate dataset, the
337 WATCH Forcing Data ERA-Interim (WFDEI, 0.5×0.5 degrees, 1979–2016,
338 Weedon et al., 2018), and the climate generated by the SEIB generator (Tei et al., 2017), which
339 is the monthly observation-based climatic datasets produced by Climatic Research Unit (CRU
340 TS4.00, 0.5×0.5 degrees, 1901–2015, Harris et al., 2014) supplemented with the National
341 Centers for Environmental Prediction/National Center for Atmospheric Research
342 (NCEP/NCAR) daily climate datasets (Kalnay et al., 1996) for 1950, with local climatology
343 recorded at meteorological stations near the sites. Local climatology in Panama is measured at
344 the Parque Natural Metropolitano Canopy Crane meteorological station (1995–2019). The
345 climatology in Austria (1979–2008) and Switzerland (1979–2000) was derived from the closest
346 meteorological station to the field site under the European Climate Assessment (Klein et al.,
347 2002, <https://www.ecad.eu>). WFDEI data were used for the climatology in Canada, except for
348 precipitation data, which are measured in the Meteorological Service of Canada (1979–1984,
349 https://climate.weather.gc.ca/historical_data/search_historic_data_e.html).

350 The reanalysis of daily WFDEI and SEIB climate data included daily records, which were
351 corrected by regression models to local climate data. For temperature, humidity, and shortwave
352 radiation values, local climatology were used directly and the daily WFDEI data supplemented
353 by simple linear regression. Precipitation data and wind speeds were first adjusted to monthly and
354 then annual averages and then scaled as a correction. WFDEI precipitation data were scaled after
355 adjusting to the annual climatological precipitation of 995 mm in 2008 for Austria and 630 mm
356 in 2000 for Switzerland. Longwave radiation was calculated using harmonized temperatures and
357 humidities above (Brutsaert, 1975). Missing values were estimated via linear interpolation.
358 Because soil temperature data were unavailable for local sites and for WFDEI, soil layer
359 temperatures were calculated using the SEIB generator by regressing soil layer 1 on atmospheric
360 temperature, soil layer 2 on layer 1, and layer 3 on layer 2. In Austria, humidity data were available
361 from 2005. The WFDEI data were therefore used to estimate missing data via linear interpolation.
362 In Canada, no observational data were available, except for temperature and precipitation.
363 Precipitation in Canada was scaled with WFDEI data after adjusting to the total climatological
364 precipitation for 1979–1984, shortwave radiation was taken from the WFDEI, and humidity data
365 were harmonized in the same way as the humidity data in Austria.

366

367 **2.3.1.3 Simulation scheme**

368

369 To reach equilibrium conditions of the biomes, plant, and soil carbon pools, a 1000-year spin-up
370 simulation was performed by looping the climate data and atmospheric CO₂ concentrations

371 between 1979–2000. Building on the final conditions of the spin-up simulations, continuous
372 simulations corresponding to 1979–2001 in Canada, 1979–2008 in Austria, 1979–2000 in
373 Switzerland, and 1979–1995 in Panama were carried out, and the NSC dynamics were compared
374 with field data.

375

376 **2.3.2 Validation at a global scale**

377

378 In the global-scale simulation, the NSC seasonality in the SEIB-DGVM was validated using
379 CRU/NCEP/MIROC integrated data (0.5×0.5 degrees, 1850–2100, Tei et al., 2017, Watanabe et
380 al., 2011) as climatic input. SEIB-DGVM-NSC ver 1.0 is expected to simulate on future scenarios,
381 thus the different climate data that cover longer period than that of section 2.3.1.2. are used for
382 validation at a global scale. The SEIB-DGVM categorizes plant species into 16 PFTs for global-
383 scale simulations.

384 The outputs of the SEIB-DGVM include two boreal biome types (evergreen and deciduous
385 forests), three temperate biome types (conifer, broad-leaved evergreen, and deciduous forests),
386 and two tropical biome types (evergreen and deciduous forests), whereas the observations
387 included two boreal biome types (conifer and deciduous forests), three temperate biome types
388 (conifer, evergreen, and deciduous forests), and two tropical biome types (evergreen and
389 deciduous forests). The model outputs and observation data were compared for each climate zone.
390 Global climate data were available from 1850 to 2005. The first 30 years (1850–1880) were
391 therefore looped for a 1000-year spin-up simulation. After the spin-up, simulations were run for
392 the period 1850–2005. The NSC dynamics from the period 1975–2005 were used for model
393 validation.

394

395 **2.4 Parameterization of NSC functions**

396

397 Hoch et al. (2003) have reported that the NSC_{leaf} of temperate trees sampled near the village of
398 Hofstetten in Switzerland varies between 7%–20% of the total leaf DM. They determined the
399 seasonal mean of the NSC_{trunk} in sapwood of temperate deciduous trees and temperate evergreen
400 trees to be $4.7\% \pm 0.1\%$ of DM and $1.8\% \pm 0.1\%$ of DM, respectively. There were no significant
401 seasonal differences. The mean NSC_{root} was less than 1.5% of the root DM for forests in Austria
402 throughout the whole season (Gruber et al., 2012), and the total NSC of temperate trees was
403 around 4%–5% of the DM during the growing season (Gruber et al., 2011). For tropical trees
404 collected in Parque Natural Metropolitano in Panama, the NSC_{trunk} and NSC_{root} were 8%–10% of
405 their biomass, whereas the NSC_{leaf} fluctuated within 5%–9% of leaf biomass (Würth et al., 2005).
406 Landhüsser and Lieffers (2003) have reported that the NSC_{root} of boreal trees in Canada, which

407 is used to support leaf flush and root growth, is 3%–4% of their root mass. The stemwood NSC_{trunk}
 408 concentration is $\sim 18 \text{ mg g}^{-1}$ of the DM for sample forests collected by Ameri flux tower (Carbone
 409 et al., 2013). Because of limited observational data, the parameters of the NSC processes were
 410 derived mostly from the values observed at each site used for point-scale validation, and the
 411 maximums of simulated NSCs were corrected so that they were in the range of measured NSCs.

412 First, the parameter a in Eq. (10) controls the base amount of photosynthetically fixed carbon
 413 mobilized for the NSC pools. The parameter b in Eq. (10) controls the seasonal fluctuations of
 414 the NSCs from the parameter a . In temperate zones, the value of b differs before and after July so
 415 that NSC peaks around mid-summer. In contrast, in tropical zones, the amount of NSC in leaves
 416 and trunks decreases throughout the spring–summer.

417 The same parameter a and b were basically used for global-scale validation as for point-scale
 418 validation. However, because the NSCs are influenced by environmental conditions at the field
 419 sites, the observed global mean values used for global-scale validation were different from the
 420 values used for setting parameters for point-scale validation. Therefore, some adjustments were
 421 made to certain parameters to align with the values used in the global-scale validation. Tables 3
 422 and 4 show the parameters used for validation. Parameter values unrelated to the NSC module
 423 remain at the default values of the SEIB-DGVM (Sato et al., 2007).

424

425 **Table 3.** Parameters of NSC pool size function for point-scale simulation

Organ	Canada	Austria	Switzerland	Panama
Leaf	$a: 0.09, b: 0.4 \times 10^{-3}$	$a: 0.04, b: 0.065 \times 10^{-3}$ (Jul–Oct)	$a: 0.13, b: 0.1 \times 10^{-3}$ (Jul–Oct)	$a: 0.06, b: -0.15 \times 10^{-3}$ (Jun–Nov)
		$a: 0.04, b: 0.135 \times 10^{-3}$ (others)	$a: 0.13, b: 0.7 \times 10^{-3}$ (others)	$a: 0.06, b: 0.15 \times 10^{-3}$ (others)
Trunk	$a: 0.06, b: 0.03 \times 10^{-3}$	$a: 0.02, b: 0.005 \times 10^{-3}$	$a: 0.02, b: 0.01 \times 10^{-3}$	$a: 0.1, b: -0.25 \times 10^{-3}$ (Jun–Nov) $a: 0.1, b: 0$ (others)
Root	$a: 0.14, b: 0.06 \times 10^{-3}$	$a: 0.02, b: 0.01 \times 10^{-3}$	$a: 0.06, b: 0.003 \times 10^{-3}$	$a: 0.04, b: 0.5 \times 10^{-3}$

426

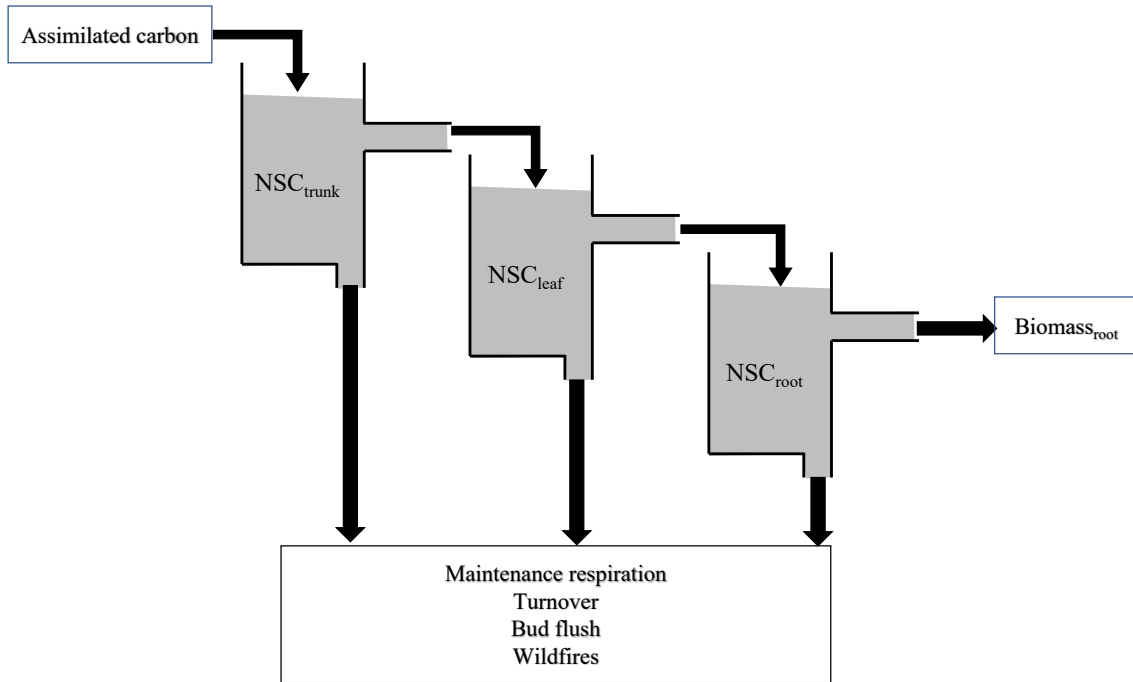
427

428 **Table 4.** Parameters of NSC pool size functions for global-scale simulation

Organ	Boreal	Temperate	Tropical
Leaf	$a: 0.09, b: 0.4 \times 10^{-3}$	$a: 0.13, b: 0.1 \times 10^{-3}$ (Jul–Oct) $a: 0.13, b: 0.9 \times 10^{-3}$ (others)	$a: 0.06, b: -0.15 \times 10^{-3}$ (May–Jul) $a: 0.06, b: 0.15 \times 10^{-3}$ (others)

Trunk	$a: 0.06, b: 0.03 \times 10^{-3}$	$a: 0.04, b: 0.01 \times 10^{-3}$	$a: 0.1, b: -0.13 \times 10^{-3}$ (May–Jul) $a: 0.1, b: 0$ (others)
Root	$a: 0.14, b: 0.06 \times 10^{-3}$	$a: 0.06, b: 0.003 \times 10^{-3}$	$a: 0.04, b: 0.5 \times 10^{-3}$

429



430

431 **Figure 1.** Schematic model structure of NSC pool. The assimilated carbon satisfies the NSC_{trunk},
 432 and then the excess assimilated carbon satisfies the next NSC_{leaf} and NSC_{root}. The accumulated
 433 carbon in NSC pool is used to compensate for the shortage of maintenance respiration and help
 434 with bud flush, and is sometimes lost due to turnover. When the biomass of plants is lost due to
 435 wildfires, the NSC also decreases.

436

437

438

439

440 **3 Results**

441

442 **3.1 Seasonality of NSC at the point scale**

443

444 **3.1.1 Boreal**

445 In Canada, the fact that the dominant PFT in the simulations was boreal deciduous trees was
446 consistent with observations at the site in Canada. The model simulated an increase in the NSC
447 of leaves from 80 mg g⁻¹ to 203 mg g⁻¹, whereas the observed NSC values were 89 mg g⁻¹ in May
448 and 185 mg g⁻¹ in August (Fig. 2a). The modeled NSCs in leaves therefore captured the increasing
449 trend during the growing season, but the simulations overestimated the maximum NSC a little.
450 The observed NSCs in trunks fluctuated from 90 mg g⁻¹ to 192 mg g⁻¹ during a year with no
451 specific seasonal trend (Fig. 2e). The model outputs in trunks were in the range 56–76 mg g⁻¹.
452 Although the observed fluctuations exceeded the modeled outputs, the modeled outputs were
453 within one standard deviation of the observations. The range of the observed NSCs in roots was
454 97–138 mg g⁻¹, whereas the range of the modeled NSCs was 117–132 mg g⁻¹ (Fig. 2i). However,
455 the observations peaked in August 2001 and in October 2002. The modeled NSCs of roots differed
456 from the observed NSCs because the former peaked during August in both years. Overall, the
457 simulated NSCs agreed well with the observed data (Fig. 3; RMSE = 69.92 mg g⁻¹, $r = 0.21$).

458

459 **3.1.2 Temperate**

460

461 In Austria, the fact that the dominant PFT in the simulations was temperate conifer forests was
462 consistent with observations at the site in Austria. The modeled NSCs in leaves accumulated until
463 July with a maximum of 142 mg g⁻¹. This pattern was similar to the observed seasonality of the
464 NSCs, which peaked at 150 mg g⁻¹ (Fig. 2b). The modeled NSCs in trunks were stable in the
465 range 19–26 mg g⁻¹, and the observations were within the range 18–38 mg g⁻¹, with no specific
466 seasonality (Fig. 2f). The modeled values were interspersed between the observations. The
467 modeled NSCs in roots varied in a curvilinear manner from 18 to 26 mg g⁻¹, a range that was
468 similar to the range of the observed NSCs, 13–32 mg g⁻¹ (Fig. 2j). The seasonality and magnitudes
469 of the modeled NSCs were consistent with observations (Fig. 3; RMSE = 9.52 mg g⁻¹, $r = 0.95$).

470 In Switzerland, the dominant PFT in the simulations corresponded to the temperate conifers
471 observed at the field site. The NSCs in the tree leaves accumulated during early spring and reached
472 up to 222 mg g⁻¹ (Fig. 2c). The decrease of the NSCs after July to a minimum of 135 mg g⁻¹ was
473 similar to the decline of the observed NSCs to a minimum of 124 mg g⁻¹. The modeled NSCs in
474 trunks fell in the range 13–16 mg g⁻¹, which was overlapped with the range of the observed NSCs
475 in trunks, 15–33 mg g⁻¹ (Fig. 2g), and the modeled NSCs all fell within one standard deviation of

476 the observations. The modeled NSCs in roots increased gradually from 45 to 62 mg g⁻¹, which is
477 similar to the observed range of observations, 48–64 mg g⁻¹ (Fig. 2k). The simulations captured
478 the amounts and seasonal patterns of the NSCs in the different organs and produced results that
479 compared well with observations (Fig. 3; RMSE = 25.83 mg g⁻¹, $r = 0.91$).

480

481 **3.1.3 Tropical**

482

483 In Panama, while a wide range of woody species was found at the Panama site, in the simulation
484 the tropical evergreen PFT became dominant. The simulations showed that the NSCs in leaves
485 were stored during winter and were then gradually consumed from July to October, when they
486 reached a minimum of 52 mg g⁻¹ (Fig. 2d). The observed NSCs in leaves likewise decreased from
487 69 to 48 mg g⁻¹ between August and October. The model therefore followed the observed
488 seasonality of the leaf NSCs. The modeled NSCs in trunks fell in the range 35–73 mg g⁻¹ (Fig.
489 2h). The slight decrease of the modeled NSCs in trunks during the summer was not apparent in
490 the observations. However, the simulated values fell within the range of the observed NSCs, 27–
491 97 mg g⁻¹. The simulated NSCs in roots fell in the range 23–55 mg g⁻¹; the observed NSCs ranged
492 from 43 to 70 mg g⁻¹ (Fig. 2l). Despite the weak correlation between simulated and observed
493 NSCs, the model results were within the acceptable margin of error (Fig. 3; RMSE = 20.75 mg
494 g⁻¹, $r = 0.08$).

495

496 **3.2 Comparison of annual mean NSC concentrations at a global scale**

497

498 For validation at a global scale, the mean annual NSCs from the new model were compared with
499 the observed mean annual NSCs in boreal, temperate, and tropical regions (Table 5). The model
500 simulated the amounts of NSCs in forest tree trunks in all climate regions with high accuracy. The
501 modeled NSCs in the trunks of trees in boreal forests averaged 47.48 ± 18.35 mg g⁻¹, which
502 compared favorably with the observed average of 76.67 ± 23.68 mg g⁻¹. In temperate forests, the
503 modeled NSCs of trunks averaged 44.78 ± 6.82 mg g⁻¹, which was close to the observed average of
504 51.59 ± 22.63 mg g⁻¹. The modeled NSCs of trunks in tropical forests averaged 66.68 ± 18.79 mg g⁻¹,
505 which was close to the average of the observations, 106.23 ± 32.52 mg g⁻¹. Although the modeled
506 NSCs in leaves of temperate and tropical forests were close to observed values, the modeled NSCs
507 in leaves of boreal forests underestimated the observed values. Moreover, the modeled NSCs in
508 roots of tropical forests were smaller than the observed NSCs. Overall, the simulated NSCs of all
509 organs of forest trees in all climate regions agreed reasonably well with observations (Fig. 4;
510 RMSE = 66.75 mg g⁻¹, $r = 0.17$). The model could simulate the NSCs with high accuracy, with
511 the exception of the NSCs of tree leaves in boreal forests and of tree roots in tropical forests (Fig.

512 4; RMSE = 34.15 mg g⁻¹, $r = 0.71$). The original SEIB-DGVM only calculated NSCs in the trunks
 513 of trees with, an average value of 63.70 ± 44.64 mg g⁻¹ in boreal forests, 20.87 ± 15.91 mg g⁻¹ in
 514 temperate forests, and 16.61 ± 10.22 mg g⁻¹ tropical forests. Although the NSC in trunks of boreal
 515 forests from the original SEIB-DGVM was close to observation, the old model underestimated
 516 the NSC in trunks of temperate and tropical forests. The simulated NSCs from the original SEIB-
 517 DGVM in all climate regions were found to be poorly less correlated with observations (Fig. 4;
 518 RMSE = 55.37 mg g⁻¹, $r = 0.01$).

519

520 **Table 5.** Comparison of modelled and observed annual mean NSC concentrations (mg g⁻¹) on a
 521 global scale. The observed results are represented as the mean ± 1 standard deviation.

	Boreal		Temperate		Tropical	
	Observation	Model	Observation	Model	Observation	Model
Leaf	202.80 ± 19.97	94.91 ± 42.91	127.10 ± 25.6	170.90 ± 46.54	86.42 ± 20.21	46.92 ± 16.20
Trunk	76.67 ± 23.68	47.48 ± 18.35	51.59 ± 22.63	44.78 ± 6.82	106.23 ± 32.52	66.68 ± 18.79
Root	118.49 ± 13.24	105.80 ± 40.82	67.65 ± 18.79	23.58 ± 10.57	170.40 ± 36.49	44.55 ± 15.15

522

523 3.3 Woody biomass and total NSCs on a global scale

524

525 The average of the total GPP simulated from the new model during 1976–2005 was 123 PgC
 526 year⁻¹. The model estimated the mean total woody biomass to be 282 PgC year⁻¹ in boreal zones,
 527 100 PgC year⁻¹ in temperate zones, and 337 PgC year⁻¹ in tropical zones globally during 1976–
 528 2005. In boreal zones, the new model estimated the mean concentration of total NSCs to be 4.98%
 529 ± 1.87% of total woody biomass, while the original SEIB-DGVM estimated it to be 6.37% ±
 530 4.46% of total woody biomass (Fig. 5). The new model's estimation of the percentage of NSCs to
 531 total woody biomass in North America and North Russia was lower than the original SEIB-
 532 DGVM. In temperate zones, the mean concentration of total NSCs was 4.67% ± 0.54% of total
 533 woody biomass from the new model, while 2.09% ± 1.59% from the original SEIB-DGVM. The
 534 NSCs in the temperate forests of Asia and South America accounted for a larger fraction of total
 535 biomass in the new model compared to the original SEIB-DGVM. Total NSCs of tropical forests
 536 in South America and Africa from the new model were 6.19% ± 1.66% of their total woody
 537 biomass, the original SEIB-DGVM estimated it to be 1.66% ± 1.02% of the total biomass. The
 538 new model estimated a larger percentage of NSC to total biomass across tropical regions compared to
 539 the original SEIB-DGVM.

540 Mean values of the simulated total NSCs relative to total woody biomass from the new model
541 were close to previous estimates for temperate and tropical forests (Table 6). The total NSCs of
542 temperate, broad-leaved, evergreen forests from the new model were $4.63\% \pm 0.50\%$, which
543 corresponded to the woody biomass reported by Smith et al. (2018). Furthermore, in the new model,
544 the total NSCs of temperate conifer forests were $4.72\% \pm 0.58\%$ of total woody biomass, which was
545 close to the figure of 4% reported by Körner (2003). While, the original SEIB-DGVM calculated the
546 total NSCs of temperate broad-leaved evergreen forests to be $2.64\% \pm 1.24\%$, and the total NSCs of
547 temperate conifer forests were $5.30\% \pm 2.68\%$ of the total woody biomass, which closely matched
548 the observations. However, the original SEIB-DGVM only considered NSCs in the trunks, whereas
549 the new model allocates the total NSC into three organs, resulting in a close match to the observed
550 total NSC. According to Würth et al. (2005), the percentages of woody biomass contributed by
551 NSCs are 4%–8% in tropical forests. The new model calculated to be $4.66\% \pm 1.28\%$ in tropical
552 deciduous forests, and $7.11\% \pm 1.08\%$ in tropical evergreen forests. In contrast, the total NSCs of
553 tropical deciduous forests from the original SEIB-DGVM were $1.66\% \pm 1.35\%$ and the total
554 NSCs of tropical evergreen forests were $1.66\% \pm 0.71\%$, which were different from the observed
555 values. These observed percentages are close to our simulated values from the new model.

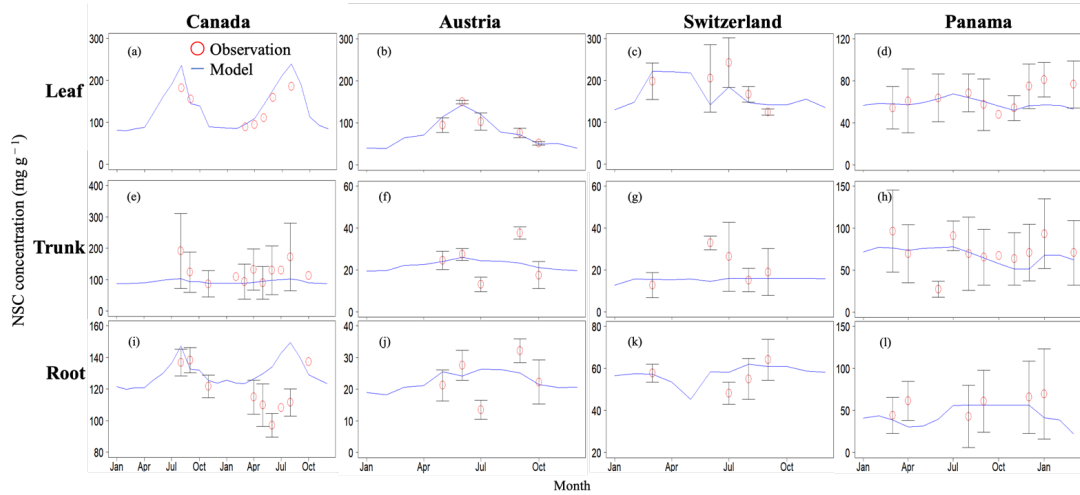
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557 **Table 6.** Comparison of modeled mean NSC concentrations to observed total NSC concentrations (%)
558 for different types of biomes. The simulated results are expressed as the mean \pm 1 standard deviation.

	Total NSC	Leaf	Trunk	Root	Total NSC (Original SEIB- DGVM)	Observation
Boreal deciduous	3.41 ± 1.58	0.05 ± 0.09	3.06 ± 1.23	0.30 ± 0.62	2.47 ± 3.32	
Boreal evergreen	6.06 ± 1.16	0.75 ± 0.38	4.73 ± 1.29	0.58 ± 0.37	7.24 ± 4.22	
Temperate deciduous	2.30 ± 0.33	0.02 ± 0.01	2.25 ± 0.31	0.03 ± 0.01	1.45 ± 0.93	1.0–12.5 (Gough et al., 2009)
Temperate broad- leaved evergreen	4.63 ± 0.50	0.49 ± 0.20	4.10 ± 0.56	0.04 ± 0.03	2.64 ± 1.24	2.6–4.4 (Smith et al., 2018)
Temperate conifer	4.72 ± 0.58	0.89 ± 0.38	3.77 ± 0.73	0.08 ± 0.04	5.30 ± 2.68	4.0 (Körner, 2003)

Tropical deciduous	4.66 ± 1.28	0.04 ± 0.03	4.60 ± 1.27	0.03 ± 0.02	1.66 ± 1.35	4.0–8.0 (Würth et al., 2005)
Tropical evergreen	7.11 ± 1.08	0.08 ± 0.03	7.00 ± 1.08	0.02 ± 0.01	1.66 ± 0.71	4.0–8.0 (Würth et al., 2005)

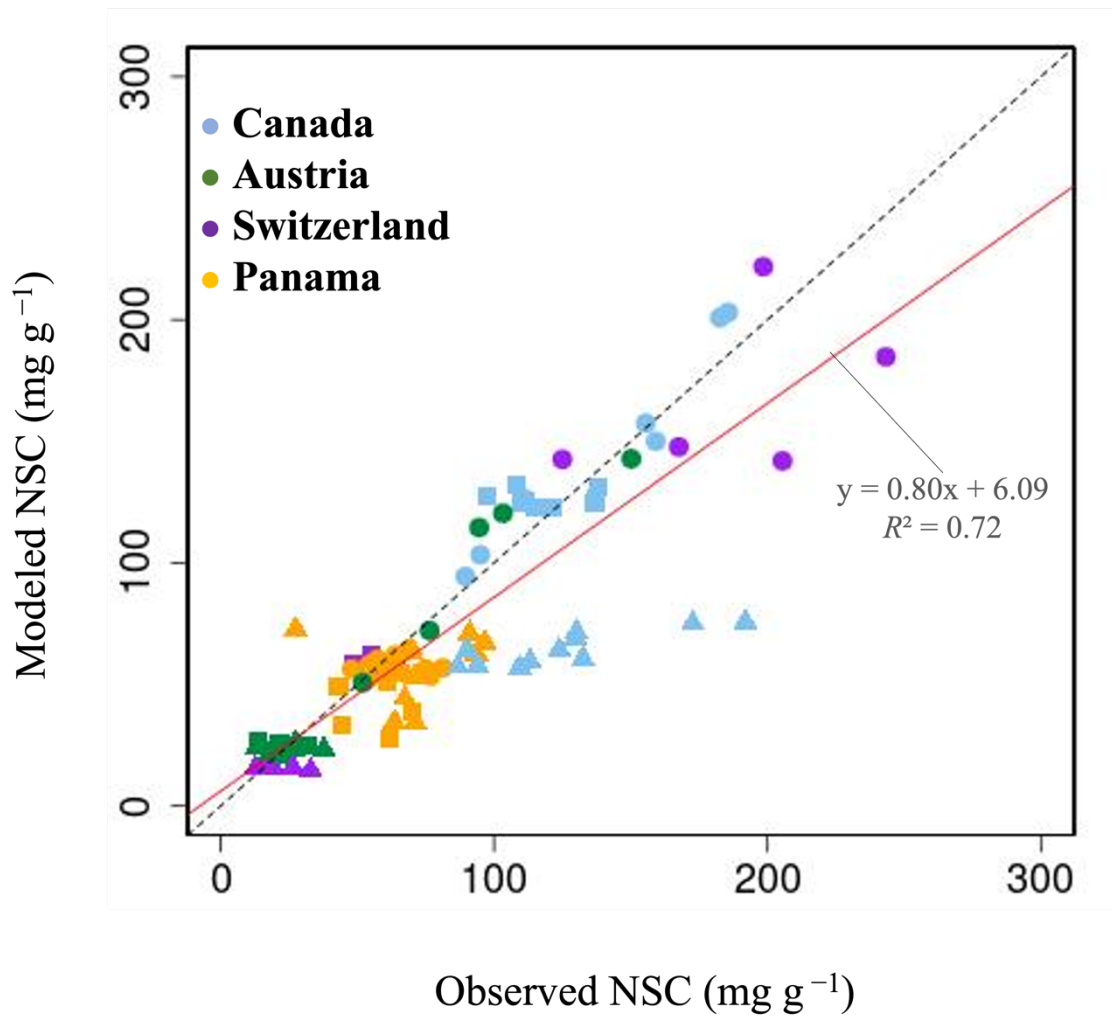
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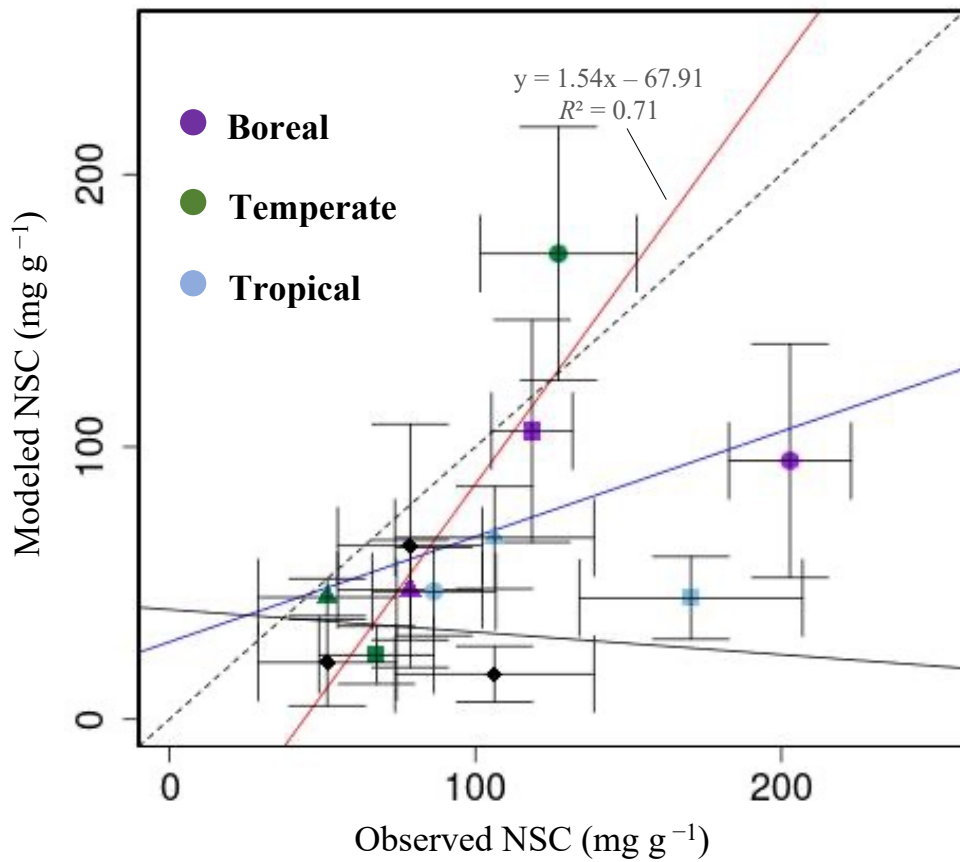
561 **Figure 2.** Validation of the modeled NSC with observed NSC data (mg g^{-1}) at sites in Canada,
 562 Austria, Switzerland, and Panama. Red circles indicate the observed data, and blue lines indicate
 563 the modeled NSC. The observed results are represented as mean \pm 1 standard deviation. Observed
 564 data are derived from Martínez-Vilalta et al. (2016).

565



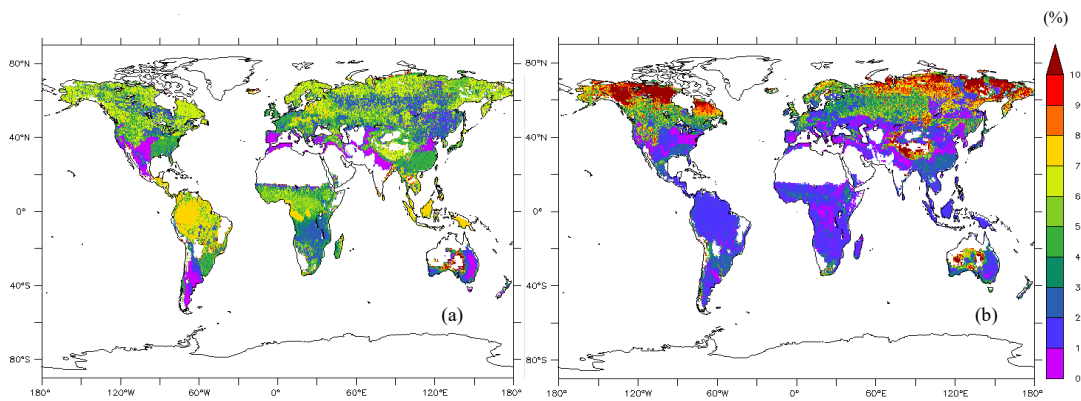
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Figure 3. Plot of modeled NSC (mg g⁻¹) with observed NSC (mg g⁻¹) at a point scale. ●, leaves; ▲, trunks; ■, roots. For all data, r is 0.72, and RMSE is 29.65 mg g⁻¹.



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Figure 4. Plot of modeled NSC (mg g^{-1}) with observed NSC (mg g^{-1}) at a global scale. ●, leaves; ▲, trunks; ■, roots; ◆, trunks in the original SEIB-DGVM. Red line represents the regression line of the plot that compares the modeled NSC from the new model with the observed NSC, except for the NSCs of tree leaves in boreal forests and of tree roots in tropical forests. Blue line represents the regression line of all plot from the new model with the observed NSC. Black solid line represents the regression line of modeled NSC from the original SEIB-DGVM with the observed NSC.



579

580 **Figure 5.** The global map of percentage of total NSC concentration relative to total dry woody
581 biomass averaged during 1976–2005 (%) (a) from the new model (b) from the original SEIB-
582 DGVM.
583
584

585 **4 Discussion**

586

587 At the point scale, the modeled NSCs for boreal forests in Canada were close to the observed
588 NSCs. The seasonality of the modeled NSCs in leaves was consistent with observations.
589 However, the seasonality of NSC in roots differed from the observations because there were
590 insufficient observations in boreal regions that enabled assessment of the seasonality of NSCs in
591 all organs. The seasonality of NSCs in roots is therefore still unclear. In temperate zones, the
592 model simulated the observed NSCs very accurately. The simulated NSCs of temperate forests
593 were close to observed values in Austria and Switzerland. The simulations showed that the
594 NSCs in leaves were consumed in winter for bud flush, and the leaves accumulated NSCs
595 during the growing season. This pattern corresponded to the seasonality reported in Asaadi et al.
596 (2018) and Furze et al. (2019). In the tropical zones, the model also captured a seasonality of
597 NSCs that was similar to observations. The NSC concentration in the canopy of tropical forests
598 decreased from June to August to satisfy increased maintenance demands (Signori-Müller et al.,
599 2022; Würth et al., 2005). The simulated NSCs in leaves followed a similar pattern from June to
600 August, and the simulated NSCs in leaves, trunks, and roots were close to observed values.

601 At the global scale, the new model simulated NSC values in each organ of all climate
602 regions that agreed with the observed data, except for the leaves of boreal forests and roots of
603 tropical forests. As for the total NSCs of biome types, the modeled total NSCs of all temperate
604 and tropical biomes matched the observed ranges well. In contrast, the original SEIB-DGVM
605 only calculated NSC in trunks, and the modeled NSC in trunks of temperate and tropical forests
606 were underestimated compared to observations. Total NSCs of tropical biomes were lower than
607 observations, while those of temperate biomes were close to observations, but the original
608 SEIB-DGVM did not assign the total NSC into leaves and roots. Therefore, these findings
609 reveal that the new model can simulate NSC more accurately than the original SEIB-DGVM. In
610 the original SEIB-DGVM, the NSC in trunks depended on the existing leaf biomass, which
611 could not be applied to all biome types and climate zones, especially tropical forests. In the new
612 model, the NSC_{organ} is determined by the biomass of the organ and photosynthesis in some
613 climate zones. The new function, which was validated at the point scale, could therefore
614 perform well on a global scale. The NSCs in trunks, which contain the greatest amounts of
615 carbon in trees, were simulated accurately in all climate regions. The new function could
616 therefore calculate the total NSCs in trees with great accuracy.

617 The model with the new function calculated the global GPP to be $123 \text{ PgC year}^{-1}$, which is
618 close to the previous estimates of $106.2 \pm 2.9 \text{ PgC year}^{-1}$ by Zheng et al. (2020) and 130 ± 1.6
619 PgC year^{-1} by Madani et al. (2020). Moreover, the simulated mean total woody biomass for boreal
620 forests was $282 \text{ PgC year}^{-1}$, which is within the range of $249\text{--}295 \text{ PgC year}^{-1}$ reported by Pan et

621 al. (2011). The simulated woody biomass of 100 PgC year⁻¹ for temperate forests was within the
622 observed range of 59–139 PgC year⁻¹ (Hui et al., 2020) and a little lower than the range of 113–
623 125 PgC year⁻¹ for other temperate forests (Pan et al., 2011). The calculated total woody biomass
624 of 337 PgC year⁻¹ for tropical forests was within the range of 212–340 PgC year⁻¹ reported by
625 Hui et al. (2020) and was not very different from the estimates of 378–564 PgC year⁻¹ by Pan et
626 al. (2011), and 200–300 PgC year⁻¹ by Mitchard (2018). Furthermore, the total NSCs relative to
627 total biomass output from the new function for temperate and tropical biome types agreed with
628 previous research. The total NSC of boreal biome types could not be compared with observations
629 due to lack of data.

630 The new model allows for simulation of various biotic effects on terrestrial ecosystems by
631 calculating the NSC dynamics within each plant organ. The NSCs stored in the trunk and roots
632 help to compensate for the deficit of CO₂ uptake in trees under stress, and the NSC stored in roots
633 is potentially indispensable for tree recovery after disturbances (Herrera-Ramírez et al., 2020).
634 Therefore, the NSC changes in the trunk and roots provide better indicators of carbon source–
635 sink relationships under elevated CO₂ conditions and are more closely related to the carbon
636 balance of plant bodies (Körner, 2003). While the original SEIB-DGVM was unable to simulate
637 biotic effects due to the lack of consideration for NSCs in all organs, simulation of the dynamics
638 of NSC in the three compartments in this research contributes to a better understanding plant
639 growth and the response of carbon dynamics in each organ to increasing atmospheric CO₂.

640 Carbon starvation may also be one of the causes of plant death during drought when
641 photosynthesis decreases and water stress increases (McDowell et al., 2008). If reduced
642 photosynthetic rates cannot supply enough carbon for NSC accumulation during drought, there
643 will be greater canopy dieback in the next season (McDowell, 2011; Chen et al., 2017). The new
644 model can simulate the dieback of long-lived temperate and tropical forests during drought,
645 because it can represent the total NSC in plant bodies at a global scale. Additionally, insect pests
646 have a significant impact on forest ecosystems, especially in temperate biomes, and their
647 outbreaks have increased with climate change (Canelles et al., 2021). To recover from defoliation
648 caused by insect pests and avoid decreased growth rates and lower survival rates, plants allocate
649 carbon for NSC defense mechanisms. The new model accurately simulates the amount of total
650 NSCs in temperate biomes, and therefore, it can be used to estimate the impact of insect pests on
651 a global and future scale.

652 The new model introduced NSC compartments in leaves, trunk, and roots that were validated
653 at the point and global scales. Use of the model developed here enabled simulation of the
654 environmental effects on forests resulting from the changing amount of NSC in each organ. The
655 simulations depicted the amount of NSC in the trunk at a global scale especially well, which
656 constitutes a significant portion of the total NSC. The model could thus be used as an indicator of

657 the carbon cycle in terrestrial ecosystems to understand the effect of climate change. Simulation
658 of photosynthetic carbon allocated into NSC storage in leaves, trunks, and roots enables a more
659 dynamic simulation of the carbon cycle between terrestrial ecosystems and the atmosphere.

660 However, there were still some limitations to this research. We considered two potential
661 limitations that could lead to some discrepancies between the modeled and measured NSC values.
662 First, the relatively coarse spatial resolution of $0.5^{\circ} \times 0.5^{\circ}$ gridded climate data at the global scale
663 could not depict the details of local climates derived from observations. These differences were
664 especially important in the case of temperature and short radiation, which play a key role in NSC
665 dynamics. Temperature surrounding plants is a key factor for the rate of plant growth (Hatfield
666 and Prueger, 2015). And different plant species has a specific temperature range. The short
667 radiation is used for a calculation of photosynthesis rate. These differences of two parameters
668 between $0.5^{\circ} \times 0.5^{\circ}$ gridded climate data and local climate data affect the ability of the trees to
669 accumulate NSCs.

670 Second, the scarcity of ground-measured NSC seasonality prevented us from having more
671 average information on NSC concentrations, especially in the tropical and boreal regions, where
672 there were fewer available data. The NSC seasonality differs between biome types, but because
673 it is difficult to measure NSC dynamics, there is a lack of long-term data for each biome type.
674 Hence, we adjusted the new NSC process and its related parameters based on climate zones rather
675 than biome types in our study. The fact that the NSC allocation was further influenced by
676 environmental conditions caused the allocation patterns to change within the same biome type.
677 The NSC allocation to roots was favored over aboveground allocations when soil resources were
678 lacking, and tree size was considered an important determinant of carbon allocation as well as
679 aridity (Hartmann et al., 2020). As we used data from different measurement sites for global-scale
680 validation, we could not account for the potential influence of varying surrounding conditions on
681 the data collected. In addition to the above factors, the number of samples and duration of
682 observations differed between the various studies. These differences led to no explicit NSC
683 seasonality. These potential sources of error in the field measurements jeopardized the model
684 performances.

685

686 **5 Conclusions**

687

688 In this study, a new NSC model was incorporated into the SEIB-DGVM to understand the effect
689 of NSC allocation on global forest dynamics through competition and establishment among
690 individual trees. The new module calculated the NSC dynamics of three organs—leaves, trunk,
691 and roots—and the general NSC seasonality based on ground measurements was determined for
692 biome types in three climate zones: boreal, temperate, and tropical. The NSC seasonality was

693 validated at four sites: Canada (boreal), Austria and Switzerland (temperate), and Panama
694 (tropical). The mean values of simulated NSC concentration agreed reasonably well with
695 observed data on a global scale.

696 The model enabled us to simulate the biotic effects resulting from insufficient NSC caused
697 by factors such as carbon starvation and insect pests that are otherwise difficult to measure in
698 terrestrial ecosystems globally. The difference of the NSC dynamics in the organs under
699 elevated CO₂ conditions highlighted the importance of modeling the organs separately when
700 studying environmental stresses. As more observation data about NSC dynamics become
701 available, the model can be further improved and can contribute to the simulations of the
702 passive biome shifts that may occur globally.

703

704 *Code and data availability*

705 The model code used in this study is archived at <https://doi.org/10.5281/zenodo.8080601>.

706

707 *Author contributions.* T.K. conceived and supervised this study and acquired the funding. H.N.
708 developed the model code and carried out the analysis and produced the figures. H.N. prepared
709 the original draft, and T.K., and L.V. reviewed it. L.W. prepared the modeling environment. All
710 authors have read and agreed to the published version of the manuscript.

711

712 *Competing interests.* The authors declare that they have no conflicts of interest.

713

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720

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