Reply to Referee #2 comments on "Daytime-only-mean data can enhance understanding of land-atmosphere coupling"

Zun Yin on behalf of co-authors

This paper demonstrates that the averaging approach employed by models (in this case, ERA5) in generating output diagnostics has an impact on what can be inferred from those diagnostics in the context of land-atmosphere coupling strength. Intuitively, such an impact makes perfect sense. I agree with the authors' call to modeling centers to provide relevant land-atmosphere coupling diagnostics at higher time resolution for improved analysis of land-atmosphere coupling.

AR: Thank you so much for your review and comments. We very much appreciate your agreement with our vision. Please check our point-by-point replies below. Referee's comments are in bold; authors' responses are in regular; and modifications in the manuscript are in blue.

All this being said, I must recommend major revision for this paper. The analysis strategy used is far from intuitive, and after reading the paper several times, I'm left unconvinced that the particular strategy used here is optimal (though I don't pretend to know what the optimal strategy is). It almost goes without saying that daytime-only data can get at the two-legged metric better than full-day or full-month data; still, I can't wrap my head around the idea that the quantile approach is the best way to tell us what we want to know (see comment 2 below).

AR: The quantile analysis is selected because of two features of the two-legged metrics (TLM): (1) The coupling strength estimated by TLM is reflected by the relative spatial difference rather than the numerical values of TLM, because there are not fundamental threshold values that distinguish regions of strong coupling from regions of week coupling. The primary aim of TLM is to discover hot spots where land and atmosphere are strongly coupled over the globe. Therefore, the spatial pattern (i.e., relative difference) of TLM is more important than the sole values of TLM in grid cells. (2) The magnitudes of a TLM, dominanted by the standard deviation term (i.e., ρ in Eq. 1), based on different timing periods are essential different. For instance, the daily amplitude of entire-day-mean sensible heat flux is systematically smaller than that of daytime-only-mean sensible heat flux. Therefore, the TLMs based on the raw values of daytime-mean (D), entire-day-mean (E), and monthly-mean (M) cannot be directly compared to demonstrate how the spatial patterns shift from one diagnosis to another.

The quantile approach can reflect the spatial patterns of TLM and provide the possibility of patterns comparison between TLMs based on different time smoothing data. In fact, other climate-relevant studies have successfully utilized the quantile approach to compare estimates based on different algorithms. For example, because satellite-based and modeled estimations are not suitable to compare to gauge measurements directly, the quantile approach was employed for relevant bias correction

or downscaling in the form of a probability density function (PDF) (Guo et al., 2018; Vrac et al., 2012; Xie et al., 2017). More importantly, the spatial differences highlighted by quantile analysis prompted the discovery of two physical mechanisms obscuring L-A diagnoses in this study (Sect. 3.3 and Fig. 4 in the main text). These findings prove that the quantile analysis is useful for spatial difference investigation in our case.

As the main concern of the L-A coupling is the daytime period, we assume that the D data are more reliable than the E and the M. Again, we found two mechanisms disturbing the L-A coupling diagnoses. However, these disturbances can be dramatically reduced if daytime-only data are utilized, confirming that the D is more reliable for L-A diagnoses than the E.

In the revision, we will explicitly explain the features of TLM and the reason for applying the quantile approach, as:

The TLMs are designed to highlight differences in L-A coupling strength between geographic regions and/or between different times of year in a given region. Those relative differences require subjective decisions to determine the threshold values separating regions of "strong" coupling from regions of weaker coupling. However, direct comparison of the numerical values of TLMs based on different time windows of inputs (i.e., M, E, and D) is not appropriate for three primary reasons. First, the magnitude of the TLMs is strongly affected by the σ term (Eq. 1), and this measure of variability can be quite different for daytime and nighttime processes. For example, H_D has much larger variance than the H_E , which systematically enlarges the \mathcal{A}_D . Additionally, strong L-A coupling signals can be positive or negative, suggesting that the change of TLM's magnitude (its absolute value) is the relevant quantity of interest rather than the magnitude of changes. Finally, L-A coupling processes are not characterized by clear thresholds, but rather by relative spatial and temporal differences.

To overcome these limitations and remove any subjectivity in our assessment of coupling strength, we use quantile to assess coupling strengths and quantify the spatial differences between TLM_M , TLM_E , and TLM_D .

1. Because the quantile analysis approach is not intuitive, further exposition in the Methods section would go a long way toward making this study more comprehensible. Perhaps the authors have spent so much time thinking about the analysis approach that it comes as second nature to them, but they should know that this won't be the case for the average reader. Significant additional explanation is needed. For example, I'm guessing that quantiles are based on all land (non land-ice) points across the globe. True? Please clarify. Also, are the quantiles computed separately for each season? If so, why are southern hemisphere JJA points mixed in with northern hemisphere JJA points in determining the quantiles? One would think that seasonal variations in the diagnostics would be hemisphere-specific.

AR: The quantile analysis was applied over land defined by the ERA5 data (e.g., land-sea fraction larger than 0.5), so landice is included. Moreover, the quantile analysis was firstly applied based on seasons (Fig. 1). After demonstrating the largest discrepancies existing in summer, we conducted the following analysis focusing on summer only (JJA and DJF for Northern and Southern Hemisphere, respectively). We will clarify these points in the text as:

We collected ERA5 output over land (land-ice included) every other hour from 1:00 UTC...

To focus on the season and coupling leg with the largest sensitivity to time series averaging window, we select A in summer (JJA and DJF in the Northern and Southern Hemisphere, respectively) as an example to explore the TLM differences in the

following content. Figure 2: (a) Spatial patterns of significant A_M , A_E , and A_D (top 10% quantile of absolute values) in summer (JJA and DJF for northern and southern hemisphere respectively).

2. Assuming that I do know what the authors are doing, I have some misgivings about what the quantile approach can tell us. Would consideration of only northern hemisphere extratropical points (probably a much cleaner approach, given seasonality) give the same results? Would a continental-scale analysis (e.g., North America only) give the same results? There's no way of knowing a priori; one can only speculate. Also, consider two highly hypothetical scenarios: AR: We compared the quantile analysis over specific regions (i.e., the extratropical region in Northern Hemisphere and the North America) and over the globe in Figure R1, R2, R3, and R4. Generally, there is no significant differences in terms of spatial patterns in L-A strongly couple seasons. The patterns fade in DJF in these global analysis because the L-A coupling is weak in Northern Hemisphere in winter. And the quantile analysis at global scale can help us to ignore those L-A weakly coupling regions, which is more advanced than the regional analysis. All in all, the key results based on the quantile analysis will not be very sensitive to changes in the analysis region or the quantile threshold.

a) The TLM values produced with all three averaging approaches are perfectly valid (i.e., are perfectly consistent with each other) except over 20% of the Earth (defined by vegetation type, location on the globe, or whatever). In that 20%, the monthly averaging approach inappropriately assigns a very high coupling strength when the actual coupling strength is very low. In this hypothetical example, the monthly averaging approach would look very bad at the high extreme, as it should, but it would also look bad (20% off) everywhere else, when this example's assumptions say that it actually works just fine. This seems to be a basic limitation of the quantile approach.

AR: The key aim of this study and the quantile analysis is to demonstrate whether diagnoses based on different time smoothing data can provide the same spatial patterns. As mentioned in the hypothesis, spatial difference will be found between M- and D-based (for example) diagnoses by quantile analysis because monthly data fails reflect the high extreme (the top 20% in reality is not fully equal to the top 20% in M-based TLM). Therefore, our aim has been reached. Specifically, through showing spatial differences, we demonstrate that at least one of the diagnosis contains bias. According to our assumption (D-based diagnosis is more accurate than E- and M-based diagnosis), we can conclude that the monthly averaging approach inappropriately assigns coupling strength somewhere.

b) In a separate hypothetical example, suppose that 80% of the globe experiences no land-atmosphere feedback of any relevance at all. In this case, quantile differences found between the averaging approaches within this lower 80% would have no practical meaning, and there'd be no point, e.g., in plotting quantile changes.

AR: First, we demonstrated that L-A coupling existed in a broad regions where the ρ term in TLMs is significant (colored regions in Fig. S4), unless the referee believes that a significant correlation does not mean coupling at all. Even if the assumption holds, we provided spatial difference analysis (the fitting rate) based on different quantile thresholds (from 50% to 95% Fig. 1). As there is no specific threshold to determine whether L-A feedback exists or not, we let the readers interpret the result by selecting the threshold by themselves. Nevertheless, we think the top 10% TLMs indicates that the coupling reaches a certain

I'm not saying that these scenarios are realistic; I'm just saying that it's easy to come up with scenarios that call into question the understanding that can be gained from a quantile-based analysis. The authors should provide significant discussion about the limitations of dealing with quantiles like this.

AR:In fact, the quantile analysis can reach our aim in both scenarios proposed above. True. Many scenarios can be proposed to debate that the quantile analysis may over/underestimate coupling strength at a certain level. However, the key aim of the quantile analysis is to demonstrate whether spatial differences exist between different diagnoses. Thus we think the quantile analysis is suitable for our study unless a case showing that two diagnoses have the same spatial patterns but result in spatial pattern differences in quantile analysis.

3. I disagree with the conclusion on lines 234-236, in reference to the Koster et al. study. That study did not use the two-legged approach to quantify coupling; it simply quantified the impact of soil moisture variations on precipitation variability at the multi-day time-scale. For the particular coupling characterization it was after, the calculation was exact and was not limited in any way by daytime-only vs. all-day vs. multi-day considerations. The results of the present study are best considered in relation to studies that use the two-legged metric.

AR: We agree with the referee that TLM and Koster's approach provide L-A interaction diagnoses from different angles and different temporal scales. We will modify the manuscript accordingly to focus on a broad topic regarding monthly smoothing, as:

..., which may result in the overestimation of L-A coupling strength in some climatic transition zones where climatic intermonthly variations are larger than intra-monthly variations.

4. Section 2.4 came off as opaque to me. What does the "top 25% quantile" refer to - if it refers to the ACF values, why are the lower values being ignored? Why is the ratio of the sigmas relevant? What is meant by "numerator of the rho term"? Why is the relevance of the ratio of the N terms? Also, though I can kind of guess what are the authors getting at when they talk about signal attenuation in the first place, I can't be sure. A major rewrite is needed here.

AR: We apologize for failing to include relevant details in the main text. The relevant information was available in the supplemental materials, but we fully agree with the author that the main text should be interpretable without needing to refer to the supplement. We will move this material to the main text in the revision. Moreover, the σ_M/σ_E does not indicate attenuation rate but the rate of information maintenance. In the revised manuscript, "attenuation resistance" is used to replace "attenuation rate".

In Sect. 2.4, we will add,

The TLMs contain a correlation term ρ and a variance term σ (Eq. ??). First, we investigate the difference of the σ term between D- and M-based TLMs. To keep the symbols simple, we denote a_i and b_i (*i* is day index) as detrended and seasonal removed daily time series. A_j and B_j (*j* is the month index) are corresponding monthly time series. As the long-term average of b_i (i.e., \bar{b}) is zero, the σ_b can be expressed as

$$\sigma_{b} = \left(\frac{1}{DMY}\sum_{i=1}^{DMY} b_{i}^{2} - \bar{b}^{2}\right)^{\frac{1}{2}}$$
$$= \left(\frac{1}{MY}\sum_{i=1}^{MY} \left[\frac{b_{i}^{2} + b_{i+1}^{2} + b_{i+2}^{2} + \dots + b_{i+D}^{2}}{D}\right]_{j}\right)^{\frac{1}{2}}$$
(R1)

D, M, and Y are the number of days, months, and years, respectively. The σ_B can be written as

$$\sigma_{B_{j}} = \left(\frac{1}{MY}\sum_{j=1}^{MY}B_{j}^{2} - \bar{B}^{2}\right)^{\frac{1}{2}}$$

$$= \left(\frac{1}{MY}\sum_{j=1}^{MY}\left(\frac{\sum_{i\in j}^{D}b_{i}}{D}\right)^{2}\right)^{\frac{1}{2}}$$

$$= \left(\frac{1}{MY}\sum_{j=1}^{MY}\left[\frac{(b_{i}+b_{i+1}+b_{i+2}+\ldots+b_{i+D})^{2}}{D^{2}}\right]_{j}\right)^{\frac{1}{2}}$$
(R2)

 σ_b contains all squared b_i , but σ_B contains averaged products of all combinations of b_i within a month. It is not difficult to proof that $D^2 \sum_{i=1}^N b_i^2 \ge (b_i + b_{i+1} + ... + b_N)^2$. The equal relation stands when $b_i = b_{i+1} = ... = b_N$, indicating all daily variables are the same within a month. Considering all months, the σ_B is larger if b_i follows the Matthew principle better, that is large values assemble together in specific months and small values assemble together in other months. As b_i is a time series of variables in a natural process. b_i is somehow correlated with itself at a certain time scale, that is the memory of b_i . It implies that if b_i is large, its neighbours (e.g., b_{i-1} and b_{i+1}) are large as well. Thus, the memory (characterized by auto-correlation) may determine the information maintained from σ_b to σ_B , if the σ_b is considered as the accurate information we want.

The ρ term based on daily time series can be written as:

$$\rho(a,b) = \frac{\sum_{i=1}^{DMY} (a_i - \bar{a}) (b_i - \bar{b})}{\sigma_a \sigma_b}$$
$$= \frac{\sum_{i=1}^{DMY} a_i b_i}{\sigma_a \sigma_b}.$$
(R3)

 \bar{a} and \bar{b} are mean of a_i and b_i , respectively. Similarly, we can get $\rho(A, B)$ as

$$\rho(A,B) = \frac{\sum_{j=1}^{MY} (A_j - \bar{A}) (B_j - \bar{B})}{\sigma_A \cdot \sigma_B}$$
$$= \frac{1}{\sigma_A \sigma_B} \sum_{j=1}^{MY} \left(\frac{\left(\sum_{i \in j} a_i\right) \left(\sum_{i \in j} b_i\right)}{D^2} \right).$$
(R4)

The ρ term contains σ terms, which has been discussed. If we focus on the numerator, we can find that the difference of numerator between E and M has a similar structure as the ρ difference between E and M. Thus, we deduct that the cross-covariance between a_i and b_i is the key contributor to the difference of the ρ 's numerator between E and M.

According to our deduction, we infer that the memory of the L-A state (i.e., the auto-correlation for a single variable and the cross-covariance for paired variables) can characterize the coupling signal attenuation due to the monthly smoothing of daily time series. Thus, for a single variable (i.e., the σ term), we calculate its auto-correlation function (ACF) with a maximum lag 30 days (within a month). Then we average the ACF values belonging to the top 25% quantile as an indicator of the attenuation resistance. And the attenuation resistance is characterized by the ratio of $\sigma_{\rm M}$ to $\sigma_{\rm E}$. For paired variables (i.e., the numerator of the ρ term $N(\rho)$, e.g., $N(\rho) = \sum_{i=1}^{DMY} a_i b_i$ in Eq. R3), we calculate the cross-covariance function (CCF) instead, but with a maximum lag ±30 days. For negatively correlated variables, we select the mean of the lowest 25% CCF as the indicator. For positively correlated variables, we select top 25% as the quantile threshold as the ACF case. Instead of $N(\rho_{\rm M})/N(\rho_{\rm E})$, we use $N(\rho_{\rm M})/(|N(\rho_{\rm E})| + |N(\rho_{\rm M})|)$ to characterize associated signal attenuation resistance, in order to avoid uncertainties due to phase shift from $N(\rho_{\rm E})$ to $N(\rho_{\rm M})$.

5. The correlations in Figure 3 are undoubtedly statistically significant, but they are far from "high" (line 185) or even "moderately large" (line 191). Those in panels (a) and (b) indicate only a 10% explanation of variance, and those in the remaining panels indicate well less than half the variance explained. The text, though, presents these fields as clear indications that the authors have identified the main controls on various quantities ("Significant correlation coefficients suggest that our indicator adequately explains the attenuation..."). To be honest, I got very little out of Figure 3 and the associated discussion.

AR: We agree that some adjectives such as "high", "moderately large" are subjective. In the revised manuscript, we will use "significant" instead. Through Figure 3 we, for the first time as far as we know, demonstrated why information is missing through monthly smoothing and how to characterize the degree of missing by an understandable concept. As far as we knew, what the difference between daily correlation and monthly correlation represent has not been clearly answered yet. Through analyzing their formulas (see our reply to the previous comment), we demonstrate how the correlation information is weakened by monthly smoothing mathematically. Based on this analysis, we propose the indicator $\overline{ACF_{\%}}$ representing L-A memory to characterize the information loss. Although the indicator is not perfect, it is the best way to compress memory information (an array of correlation coefficients) into one value currently. More importantly, Figure 3 shows that the indicator is able to capture the information loss worldwide regardless of geophysical and atmospheric complexities. In addition, this indicator firstly links the time series memory to the correlation attenuation due to coarser temporal smoothing, which has potential implications in broad fields.

In the revision, the novel indicator and its advantages will be added in the discussion, as:

Although monthly-based and daily-based correlation coefficients capture the synchronized fluctuations of two variables from different perspectives, their linkage is yet unclear. In this study, for the first time as far as we know, we demonstrate how the correlation is weakened by monthly smoothing mathematically. Moreover, we propose indicators based on the auto-correlation

function and cross-correlation function representing L-A memory to characterize the information loss. And these indicators are able to capture the information loss worldwide regardless of geophysical and atmospheric complexities (Fig. 3). In addition, these indicators first link the memory of time series to the correlation attenuation due to coarser temporal smoothing, which has potential implications in broad fields.

6. Would it be appropriate to at least mention that the daytime-only diagnostics may produce different results from midday-only diagnostics (e.g., 10AM-2PM)? Presumably not much coupling occurs at dusk and dawn. I'm not suggesting that midday diagnostics be examined in this paper; it's just that the overall problem of optimal averaging time goes beyond simply comparing all-day diagnostics to daytime-only diagnostics.

AR: Thanks for the proposal. We agree that the diagnoses based on different sub-time periods of the daytime may have significant differences. However, in comparison to the vast differences between the daytime and the nighttime, the diagnostic difference induced by different sub-time periods of the daytime is not the primary question, but is worth exploring in the following studies. Moreover, it is difficult to assume that midday-only data can further improve the diagnostics, because the key period of interest may vary with the specific process. For instance, the soil moisture in early morning may be coupled with the convective precipitation in the afternoon or early evening. Therefore, the optimal averaging time should be carefully investigated according to the process of interest.

In the revision, we will modify the text, as:

In addition to daytime mean values, separate averages throughout the local morning, midday, afternoon, and nighttime would be interesting as well depending on the specific perspectives of interest (Taylor et al., 2012; Guillod et al., 2015).

References

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Figure R1. Spatial patterns of significant \mathcal{L}_M , \mathcal{L}_E , and \mathcal{L}_D (top 10% quantile of absolute values) of different seasons in the extratropic region of the Northern Hemisphere. Euler diagrams show the colors for specific relationships (intersections, unions, or disjoints) among \mathcal{L}_M , \mathcal{L}_E , and \mathcal{L}_D . (a), (b), (e), and (f) are screenshots from the global quantile analysis. (c), (d), (g), and (h) are based on quantile analysis of the illustrated region.



Figure R2. Spatial patterns of significant A_M , A_E , and A_D (top 10% quantile of absolute values) of different seasons in the extratropic region of the Northern Hemisphere. Euler diagrams show the colors for specific relationships (intersections, unions, or disjoints) among A_M , A_E , and A_D . (a), (b), (e), and (f) are screenshots from the global quantile analysis. (c), (d), (g), and (h) are based on quantile analysis of the illustrated region.



Figure R3. Spatial patterns of significant \mathcal{L}_M , \mathcal{L}_E , and \mathcal{L}_D (top 10% quantile of absolute values) of different seasons in the North America. Euler diagrams show the colors for specific relationships (intersections, unions, or disjoints) among \mathcal{L}_M , \mathcal{L}_E , and \mathcal{L}_D . (a), (b), (e), and (f) are screenshots from the global quantile analysis. (c), (d), (g), and (h) are based on quantile analysis of the illustrated region.



Figure R4. Spatial patterns of significant A_M , A_E , and A_D (top 10% quantile of absolute values) of different seasons in the North America. Euler diagrams show the colors for specific relationships (intersections, unions, or disjoints) among A_M , A_E , and A_D . (a), (b), (e), and (f) are screenshots from the global quantile analysis. (c), (d), (g), and (h) are based on quantile analysis of the illustrated region.