Modeling the Influence of Chain Length on SOA Formation via Multiphase Reactions of Alkanes

Azad Madhu[1], Myoseon Jang[1], David Deacon[1]

[1]Engineering School of Sustainable Infrastructure and Environment, University of Florida, Gainesville, 32608, United States of America

Correspondence to: Myoseon Jang (mjang@ufl.edu)

Abstract. Secondary Organic Aerosol (SOA) from diesel fuel is known to be significantly sourced from the atmospheric oxidation of aliphatic hydrocarbons. In this study, the formation of linear-alkane SOA was predicted using the Unified Partitioning Aerosol Phase Reaction (UNIPAR) model that simulated multiphase reactions of hydrocarbons. In the model, the formation of oxygenated products from the photooxidation of linear alkanes was simulated using a near-explicit gas kinetic mechanism. Autoxidation paths integrated with alkyl peroxy radicals were added to the Master Chemical Mechanismv3.3.1 to improve the prediction of low volatility products in the gas phase and SOA mass. The resulting gas products were then lumped into volatility-reactivity based groups that are linked to mass-based stoichiometric coefficients. The SOA mass in the UNIPAR model is produced via three major pathways: partitioning of gaseous oxidized products onto both the organic and wet inorganic phases, oligomerization in organic phase, and reactions in the wet inorganic phase (acid-catalyzed oligomerization and organosulfate formation). The model performance was demonstrated for SOA data that were produced through the photooxidation of a homologous series of linear alkanes ranging from C9 to C15 under varying environments (NOx levels, and inorganic seed conditions) in a large outdoor photochemical smog chamber. The product distributions of linear alkanes were mathematically predicted as a function of carbon number using an incremental volatility coefficient (IVC) to cover a wide range of alkane lengths. The prediction of alkane SOA using the incremental volatility-based product distributions, which were obtained with C9-C12 alkanes, was evaluated for prediction of C13 and C15 chamber data and further extrapolated to predict the SOA from longer chain alkanes (≥C15) that can be found in diesel. The model simulation of linear alkanes in diesel fuel suggests that SOA mass is mainly produced by alkanes C15 and higher. Alkane SOA is insignificantly impacted by the reactions of organic species in the wet inorganic phase due to the hydrophobicity of products, but significantly influenced by gas-particle partitioning.

1 Introduction

Secondary organic aerosol (SOA) represents a significant proportion (20%-90%) of fine particulate matter in ambient air (Kanakidou et al., 2005). SOA is formed from the atmospheric photooxidation of hydrocarbon (HC) species and can have significant impacts on climate, and human health (U.S. EPA 2019; Nel, 2005; Shrivastava et al., 2017; World Health, 2016). To properly assess SOA burdens, especially in urban areas, precursor alkane species must be accounted for. Alkanes are released into the atmosphere primarily through varying anthropogenic sources such as fossil fuels and volatile chemical products (i.e. personal care products, paints, pesticides, etc.) (Li et al., 2022; Wu et al., 2019). Alkanes have been shown to be significant components of gasoline exhaust (6-18%) and diesel exhaust (18-31%) and vehicle emissions are the main source in urban areas (Caplain et al., 2006). Whereas the majority of SOA produced from gasoline sources is from its aromatic content, about half of the SOA produced from diesel can be attributed to aliphatic compounds (Gentner et al., 2012). Currently, long chain alkanes are recognized to be important precursors of SOA production with increasing carbon numbers in linear alkanes leading to higher SOA yields (Aumont et al., 2012). Branched and cyclic alkanes will tend to have different yields compared to linear alkanes with the same carbon number. This is partially because branched alkanes tend to produce more volatile products and cyclic alkanes produce less volatile products relative to linear alkanes (Lim and Ziemann, 2009).

Previous studies that model SOA formation have often had a discrepancy between predicted SOA formation and field observations (Shrivastava et al., 2017; Volkamer et al., 2006). This discrepancy can be explained by several reasons. Firstly, the contribution of intermediate volatility organic compounds (IVOCs) has often been ignored during model simulations (Robinson et al., 2007). Multiple studies in recent years have attempted to correct this discrepancy by including IVOCs. Using the CHIMERE regional air quality model, Hodzic et al. (2010) found that the inclusion of IVOCs led to a substantial improvement in predictions of SOA compared to data collected from the MILAGRO field experiment. Alkanes are important IVOCs in urban environments. Lee-Taylor et al. (2011) attempted to simulate the same MILAGRO field data using a near-explicit gas mechanism (GECKO-A) to determine SOA mass based on gas-particle partitioning of possible intermediate products. They reproduced the field data well and they found that more that 75% of SOA produced in their model were from products of larger linear alkane species, which were used as surrogates for larger precursors.

Secondly, currently existing gas mechanisms for precursor hydrocarbons have a large level of uncertainty due to unidentified oxidation paths. For example, Xavier et al. (2019) found that the formation of terpene SOA simulated by using exclusively the MCM mechanism consistently underpredicted experimental values but the inclusion of peroxy radical autoxidation mechanisms (PRAM) improved predictions by increasing the formation of low volatility products. Like terpenes, autooxidation occurring on the products of alkane species can produce low volatility products that readily partition to the particle phase, but it has not been included in the current MCM. The addition of those reaction pathways can improve currently existing mechanisms and consequently, SOA predictions.

Thirdly, relying on an improvement of the gas mechanism alone is not sufficient for improving predictions of SOA production. Lee-Taylor et al. notes that their reasonable agreement for SOA predictions with field data is not necessarily realistic as they do not include alternative pathways for SOA formation such as particle-phase reactions. Whereas the particle-phase reactions of alkane SOA are limited, the organic matter produced by aerosol phase reactions of reactive organic compounds, originating from other precursors, can increase the partitioning of alkane oxidation products. Yang et al. (2019) simulated SOA formation in China using the nested air quality prediction modeling system (NAQPMS) but have a similar issue to Hodzic et al. (2010) as they also exclusively use volatility-based lumping and do not consider particle-phase reactions. Pye and Pouliot (2012) consider particle phase reactions in their simulation of regional SOA formation using the Community Multiscale Air Quality (CMAQ) model. However, the complexity of particle phase reactions in their model is limited as they use a uniform reaction rate constant for the particle phase reactions of semi-volatile SOA. Alkane SOA may not be significantly influenced by particle phase reactions compared to other precursors, but some studies have attempted to account for particle phase reactions to form alkane SOA. For example, Cappa et al. (2013), using the Statistical Oxidation Model (SOM), attempts to implicitly capture particle-phase reactions through fitting model parameters to alkane SOA experimental measurements instead of explicitly modeled particle-phase reactions. Similarly, in the Functional Group Oxidation Model (FGOM) derived by Zhang and Seinfeld (2013), the reactions are controlled by rate constants that are fit to experimental measurements, rather than derived from the properties of explicit or lumped compounds from alkane oxidation. The UNIPAR model was previously developed to improve predictions by considering particle-phase reactions as well as gas-particle partitioning as mechanisms for SOA production (Beardsley and Jang, 2016; Im et al., 2014; Zhou et al., 2019). In this study, autoxidation pathways of alkyl peroxy radicals are added to currently existing alkane MCM oxidation mechanisms to include low volatility oxidation products and applied to the UNIPAR model for SOA prediction. The simulation results are compared to data collected in the UF Atmospheric PHotochemical Outdoor Reactor (UF-APHOR) chamber under the controlled environments (NOx levels, and seed conditions). In order to simulate SOA formation from a wide range of alkane length, a mathematical equation to construct product distributions was equipped in the UNIPAR model by using explicit products of C9-C12 linear alkanes. This equation is extrapolated to predict product distributions from larger alkanes (C13 & C15), for which explicit mechanisms are unavailable in MCM, and applied to SOA prediction. These predictions are also compared to chamber data obtained for C13 and C15. Extrapolation to alkanes larger than C15 is explored and used to project SOA formation from linear alkane species found in diesel fuel. In order to increase the feasibility of the model, UNIPAR was also integrated with the carbon bond mechanism (Emery et al., 2015).

2 Experimental section

Alkane SOA was produced from photooxidation of a homologous series of linear alkanes (C9-C13, C15) using the UF-APHOR dual chamber (52 m^3 each) located at the University of Florida. The detailed description of the operation of the large outdoor smog chamber can be found in a previous study (Im et al., 2014). The detailed experimental conditions of

outdoor chamber experiments are summarized in Table 1. Generally, the precursor alkane HCs were evaporated through

95 heating to introduce them into the chamber. A non-reactive gas, CCl_4 (Sigma-Aldrich; $\geq$99.5%), was introduced into the

chamber as an indicator for chamber dilution. HONO was added to the chamber as a source of hydroxyl radicals. HCs,

HONO, and NO (2% in N_2, Airgas Inc., USA) were introduced into the smog chamber before sunrise. The HCs were studied

under two different NO_x levels (high NO_x: HC/NO_x <5 ppbC/ppb; low NO_x: HC/NO_x >10 ppbC/ppb) and two different seed

conditions (without seed and sulfuric acid, SA).

100 A GC-FID (7820A, Agilent Technologies, Inc., USA) was used to measure the concentration of gas-phase HCs and CCl_4. A

photometric ozone analyzer (400E, Teledyne Technologies, Inc., USA) and a chemiluminescence NO/NO_x analyzer (T201,

Teledyne Technologies, Inc., USA) were used to measure the concentrations of ozone and NO_x within the chamber,

respectively. A Particle into liquid sampler (ADISO 2081, Applikon Inc., USA) integrated with Ion Chromatography

(Compact IC 761, Metrohm Inc., Switzerland) (PILS-IC) was used in experiments performed with inorganic seed to measure

105 the inorganic ion concentration within the chamber. A scanning mobility particle sizer (SMPS 3080, TSI Inc., USA) was

used to measure the size distribution of particles within the chamber. Previous studies that have measured the density of

alkane SOA have found a range from 1 to 1.4 g/cm^3 (Li et al., 2020; Li et al., 2022; Lim and Ziemann, 2009; Loza et al.,

2014). Aerosols from each alkane experiment in this study were assumed to have a density of 1.2 g/cm^3. A hygrometer

(CR1000 measurement and control system, Campbell Scientific Inc., USA) was used to measure meteorological factors

110 (temperature, relative humidity (RH) and an ultraviolet radiometer (TUVR, Eppley Laboratory Inc., USA) was used to

measure sunlight intensity. An organic carbon/elemental carbon analyzer (OC/EC model 4, Sunset Laboratory Inc., USA)

was used to measure the concentration of organic carbon in aerosol every 50 minutes. The concentration of organic matter in

aerosol (OM, $\mu g\ m^{-3}$) was then calculated based on the OC concentration predicted by the UNIPAR model and an OM to OC

ratio. The OM to OC ratio of SOA from alkane species decreased as the chain length increased. The concentrations of OM

115 measured from the chamber were corrected for chamber dilution using a dilution factor and for particle wall loss to the

chamber wall using a particle loss factor. An aerosol chemical speciation monitor (ACSM, Aerodyne Research Inc., USA)

was used to measure the aerosol composition (sulfate, nitrate, ammonium, and OM). The compositions obtained from the

ACSM were compared with measurements from the OC/EC and the PILS-IC. SOA yields (*Y*) were then calculated as the final measured concentration of OM divided by the total consumption of HC precursors.

3 Model description

The UNIPAR model's ability to accurately simulate SOA formation from various aromatic HCs (Im et al., 2014; Zhou et al., 2019; Han and Jang, 2022a), monoterpenes (Yu et al., 2021), and isoprene (Beardsley and Jang, 2016) has been previously demonstrated. In this study, the UNIPAR model was extended to simulate the formation of SOA from multiphase reactions of linear alkanes including gas (*g*), organic (*org*) and inorganic (*inorg*) phases. The model utilizes a near-explicit gas phase mechanism for the photooxidation of each alkane. The model employed the Master Chemical Mechanism (MCM v3.3.1) and gas autoxidation pathways to yield highly oxidized molecules at given alkane and meteorological conditions under various NOx levels (HC ppbC/NOx ppb= 2-50). The resulting products were then classified into an array of 51 different lumping species based on their volatility and reactivity. Previous studies have a detailed description of the lumping criteria and the mass-based stoichiometric coefficient (α_i) of lumping group *i* as a function of NOx levels and the degree of aging (Zhou et al., 2019). To cover alkanes in a variety of carbon lengths, the product distributions of linear alkanes were represented as a function of carbon number using an incremental volatility coefficient (IVC). The SOA mass in the model forms via three pathways: OM produced via multiphase partitioning of organic products (OM_P),aerosol phase reactions of organic species to form OM_{AR} via oligomerization in the *org* phase, and reactions in the wet *inorg* phase which also form OM_{AR} (acid-catalyzed oligomerization and organosulfate (OS) formation). The inputs to the UNIPAR model include the equations for α_i; physicochemical parameters of lumping species; the consumption of HC (ΔHC); the concentrations of alkyl peroxide radical (RO_2) and hydroperoxyl radical (HO_2); concentrations of ionic species (sulfate and ammonium ions); temperature, and humidity. Fig. 1 illustrates the UNIPAR model frame.

3.1 Gas mechanisms

The atmospheric oxidation of alkanes begins with the reaction with an OH radical, followed by the addition of O_2 to form peroxyl radicals. This oxidative addition can create specific structures in some compounds which have been identified in previous literature to be capable of undergoing autoxidation reactions (Crounse et al., 2013; Bianchi et al., 2019; Roldin et al., 2019). These reactions have been previously modeled (Pye et al., 2019; Xavier et al., 2019). An example of an autoxidation pathway added to the MCM mechanism for a linear alkane is shown in Fig. S1. Most autoxidation products from the alkanes higher than C9 belong to the low volatility group of the UNIPAR model and these products can increase SOA mass via gas-particle partitioning. As the size of the linear alkane increases, there is an increased likelihood that some

of the precursor HC is lost to the chamber wall. The rate at which alkane species partition to the wall of the UF-APHOR chamber was measured using a night-time experiment (Table S1).

3.2 Lumping and aging

The lumping structure of the UNIPAR model, along with a dynamic α_i array which considers aging, has been developed in previous studies (Zhou et al., 2019; Han and Jang, 2020; Yu et al., 2021). According to this structure, the alkane oxidation products, originating from simulations at various NO_x levels (HC ppbC/NO_x ppb = 2~50) using the MCM and alkyl peroxy radical autoxidation mechanisms, were lumped . A standard meteorological profile (14 June 2018) of a day with a clear sky near summer solstice was used and all simulations began before sunrise. The α_i array consists of 6 different reactivity levels (very fast, fast, medium, slow, partitioning only, and multi-alcohol) and 8 different volatility levels based on vapor pressure which represent 48 species, along with 3 explicit species that are lumped separately (glyoxal, methylglyoxal, and epoxydiols). As gas products originating from the precursor HC are atmospherically aged, they become more oxidized, which leads to the addition of functional groups. This atmospheric aging can augment the product distribution, forming more reactive and less volatile products via oxidation, or photolysis products which are more volatile but may be more reactive. To incorporate the impact of aging on SOA formation, the dynamic α_i is reconstructed using two different gas phase oxidation compositions: fresh and highly oxidized. A weighted aging factor is used to dynamically change α_i values based on these two compositions. The aging factor (f_a) at a time t, as detailed in Zhou et al. (2019), is defined as:

$$f_a(t) = \log \frac{[HO_2] + [RO_2]}{[HC]_0} \tag{1}$$

where $[HO_2]$ and $[RO_2]$ and $[HC]_0$ are the concentrations of hydroperoxide radical, organic peroxyl radical, and initial HC, respectively. Both fresh and highly oxidized compositions are calculated for each NO_x level as well as the respective aging factors. $f_a(t)$ is also converted into an aging scale that ranges from 0 (fresh composition) to 1 (highly oxidized composition) as follows (Zhou et al., 2019):

$$f_a'(t) = \frac{f_a(highly\ oxidized) - f_a(t)}{f_a(highly\ oxidized) - f_a(fresh)} \tag{2}$$

$f_a(t)$ is calculated for a given NO_x level and used to dynamically calculate the α_i values for that same NO_x level as follows (Zhou et al., 2019):

$$\alpha_i = \left(1 - f'_a(t)\right)(fresh\ \alpha_i) + \left(f'_a(t)\right)(highly\ oxidized\ \alpha_i) \tag{3}$$

The dynamic updates of physicochemical parameters (molecular weight (MW), O:C ratio, and hydrogen bonding (HB)) of lumping species are also performed in a similar way. Further details about lumping criteria and dynamic updates of physicochemical parameters can be found in the study by Zhou et al. (2019).

3.3 Unification and Extrapolation of lumping species

To predict the SOA formation from linear alkanes larger than C12, for which MCM do not exist, the product distributions of C9-C12 are parameterized using an incremental volatility coefficient (IVC) as a function of the carbon length and then extrapolated. As the carbon number of linear alkanes increases, the vapor pressure of both the precursor HC and the resulting oxidation products will decrease, causing a shift in the product distribution. The IVC is meant to generalize this shift by describing the drop in vapor pressure caused by the addition of a carbon atom. The IVC ($I(n)$) which is used to calculate the product distribution for a linear alkane with carbon number n using the product distribution of a linear alkane with carbon number n-1 is as follows:

$$I(n) = a \left(\frac{1}{b}\right)^{(n-9)} \tag{4}$$

where coefficients a and b are obtained by fitting the predicted α_i values using equation 4 to those constructed from the individual HC by using explicit gas mechanisms in section 3.1. In this study, a and b were calculated to be 0.79 and 1.83, respectively. The α value ($\alpha_{n,j,k}$) with carbon number n, at a given volatility (j) and reactivity (k) is predicted by using α value with carbon number n-1 as follows:

$$\alpha_{n,j,k} = I(n) * \left(\alpha_{n-1,j+1,k}\right) + \left(1 - I(n)\right) * (\alpha_{n-1,j,k}) \tag{5}$$

Term $\alpha_{n-1,j+1,k}$ represents the stoichiometric coefficient of volatility group j+1 adjacent to volatility group j at a given reactivity k with carbon number n-1. Term $\alpha_{n-1,j,k}$ represents the stoichiometric coefficient of volatility group j at a given reactivity k with carbon number n-1. Volatility group j is less volatile than volatility group j+1. Notably, the shifts in composition are only within the vapor pressure groups (j) at a given reactivity group (k).

Using this IVC, the product distribution of C9 is the base (Eq. 4) to predict the product distribution of C10. This process is sequentially repeated for higher carbon numbers. Each repetition will yield the product distribution of a linear alkane with 1

carbon number larger than the last. Fig. S2 illustrates the construction of product distribution based on the incremental

195   volatility approach.

3.4 SOA formation by partitioning

Partitioning coefficients for each lumping species, i, between g and org phase ($K_{or,i}$) or between the g and wet $inorg$ phase

($K_{in,i}$) are calculated using the typical gas-particle partitioning model (Pankow, 1994):

$$K_{or,i} = \frac{7.501RT}{10^9 MW_{org}\, \gamma_{org,i} p_{L,i}^o} \tag{6}$$

200

$$K_{in,i} = \frac{7.501RT}{10^9 MW_{inorg} \gamma_{inorg,i} p_{L,i}^o} \tag{7}$$

where R is the gas constant (8.314 J mol^{-1} K^{-1}), and T is temperature (K). MW_{org} and MW_{inorg} are the average molecular

weights (g mol^{-1}) of the organic and inorganic phases of the aerosol, respectively. $p_{L,i}^o$ is the subcooled liquid vapor pressure

of a species, i. The activity coefficient in the organic phase for each lumping species, $\gamma_{org,i}$, is assumed to be unity (Jang and

Kamens, 1998). The activity coefficient in the inorganic phase for each lumping species, $\gamma_{inorg,i}$, is predicted by a semi-

205   empirical regression equation which was fit to the activity coefficients of various organic compounds as a function of

physicochemical parameters (MW, O:C ratio, and HB) and sulfate fraction (FS). FS is an indicator for aerosol acidity which

is defined as follows:

$$FS = \frac{[SO_4^{2-}]}{[SO_4^{2-}]+ [NH_4^+]} \tag{8}$$

where $[SO_4^{2-}]$ and $[NH_4^+]$ are the concentration of sulfate and ammonium ions, respectively. The semi-empirical equation,

210   derived from activity coefficients estimated using the Aerosol Inorganic-Organic Mixtures Functional Groups Activity

Coefficients (AIOMFAC) model (Zuend et al., 2011) at a given RH is as follows:

$$\gamma_{inorg,i} = e^{0.035 \cdot MW_i - 2.704 \cdot \ln(O:C_i) - 1.121 \cdot HB_i - 0.330 \cdot FS - 0.022 \cdot (\cdot RH)} \tag{9}$$

Further information on the derivation and statistical properties of Eq. (9) can be found in Zhou et al. (2019). The partitioning

coefficients are used to calculate the concentration of each lumping species in the three phases ($C_{g,i}$, $C_{org,i}$, and $C_{inorg,i}$) from

215   the total concentration of each lumping species ($C_{T,i}$). The total SOA mass formed by partitioning in both org and $inorg$

phases (OM_P) is predicted by the following equation which was developed by Schell et al. (2001) and reconstructed to consider mass formed by particle-phase reactions (OM_{AR}) by Cao and Jang (2010):

$$OM_P = \sum_i \left[C_{T,i} - OM_{AR,i} - C_{g,i}^* \frac{\frac{C_{org,i}}{MW_i}}{\sum_i \left(\frac{C_{org,i}}{MW_i} + \frac{OM_{AR,i}}{MW_{oli,i}} \right) + \frac{OM_0}{MW_{oli,i}}} \right] \tag{10}$$

where Cg^* ($1/K_{org,i}$) and OM_0 (mol m^{-3}) represent the effective saturation concentration and pre-existing OM, respectively. $MW_{oli,i}$ and MW_i represent the molecular weights of oligomeric products and lumping species, respectively. Eq. (10) is solved using Newton-Rapson method, which iterates until a convergence is reached (Press et al., 1992).

3.5 SOA formation by particle-phase reactions

OM_{AR} is attributed to both the *org* and *inorg* phases. In the *org* phase, SOA formation is attributed to oligomerization as organic compounds undergo self-dimerization reactions (Im et al., 2014; Zhou et al., 2019). In the *inorg* phase, oligomerization of organic compounds can be acid-catalyzed (Jang et al., 2002). Oligomerization reactions are expressed as second-order reactions (Odian, 2004) with rate constants $k_{AR,org,i}$ and $k_{AR,inorg,i}$ (L mol^{-1} s^{-1}) in the *org* and *inorg* phases, respectively. A semi-empirical model developed by Jang et al. (2005) is used to estimate $k_{AR,in,i}$ as follows:

$$k_{AR,inorg,i} = 10^{0.25pK_{BH_i^+} + 1.0X + 0.95R_i + \log(a_w[H^+]) - 2.58} \tag{11}$$

where R_i represents species reactivity, $pK_{BH^+}{}_i$ represents the protonation equilibrium constant, a_w represents the activity of water, X represents excess acidity (Cox and Yates, 1979), and $[H^+]$ represents the concentration of protons which are estimated using the extended aerosol inorganic model (E-AIM (Clegg et al., 1998)) $k_{AR,org,i}$ is determined as follows:

$$k_{AR,org,i} = 10^{\left[0.25pK_{BH_i^+} + 0.95R_i + 1.2\left(1 - \frac{1}{1+e^{0.005(300-MW_{org})}}\right) + \frac{2.2}{1+e^{6(0.75-O:C)}} - 10.07 \right]} \tag{12}$$

For the oligomerization in the *org* phase, the terms related to acidity (X, and $a_w[H^+]$) are excluded. Studies have previously demonstrated that aerosol viscosity can influence the mobility of chemical species and thus, apparent reaction rates, which can be limited by slow bulk diffusion in the particle-phase (De Schrijver and Smets, 1966; Reid et al., 2018). The molecular weight of species in the organic phase (MW_{org}) and the O:C ratio, which are important predictors for viscosity, are considered to calculate $k_{AR,org,i}$ (Han and Jang, 2022).

In the wet *inorg* phase of the aerosol, sulfuric acid can react with reactive organic compounds to form dialkyl sulfate (diOS). This formation of diOS can contribute to SOA mass production but can also lead to a reduction in [H$^+$] which decrease the

240 rate of SOA mass produced by acid-catalyzed oligomerization in the *inorg* phase. The formation of diOS is predicted in the UNIPAR model and applied to reduce [H$^+$] in *inorg* phase as previously reported (Im et al., 2014; Beardsley and Jang, 2016; Zhou et al., 2019).

3.6 Correction of intermediate organic vapor deposition to walls

In addition to wall loss of precursor HC, intermediate oxidized products derived from the precursor HC can deposit to the

245 wall. The organic vapor deposition to wall is kinetically described with the deposition ($k_{on,i}$) and desorption ($k_{off,i}$) rate constants of each lumping species, i, to the UF-APHOR wall by using the method detailed by Han and Jang (2020); Han and Jang (2022b). $k_{on,i}$ is expressed as a fractional loss rate (Mcmurry and Grosjean, 1985):

$$k_{on,i} = \left(\frac{A}{V}\right) \frac{\alpha_{w,i}\overline{v}_i/4}{1+\frac{\pi\alpha_{w,i}\overline{v}_i}{8(K_eD)^{1/2}}} \tag{13}$$

where D (1.0×10^{-6} m^2 s^{-1}) and K_e (0.12 s^{-1}) are the diffusion coefficient and coefficient of eddy diffusion applied as a fixed

250 number, respectively. $\left(\frac{A}{V}\right)$ represents the surface area to volume ratio of the chamber. $\overline{v}_i$ and $\alpha_{w,i}$ represent the mean thermal speed of the gas molecules, and accommodation coefficient of i to the wall, respectively. Equations for the calculation of $\overline{v}_i$ and $\alpha_{w,i}$ can be found in section S4. $K_{w,i}$ ($K_{w,i} = k_{on,i}/k_{off,i}$) is calculated as follows:

$$\ln(K_{w,i}) = -\ln(\gamma_{w,i}) - \ln(p^o_{L,i}) + \ln\left(\frac{7.501RTOM_{wall}}{10^9MW_{OM}}\right) \tag{14}$$

where $p^o_{L,i}$ (mmHg) is the liquid vapor pressure of each lumping group, i. R (8.314 J mol^{-1} K^{-1}) is the ideal gas constant and

255 T (K) is temperature. OM_{wall} (mg m^{-3}) and MW_{OM} are the concentration of organic matter on the wall, and the molecular weight of organic matter on the wall, respectively. The activity coefficient, $\gamma_{w,i}$, is calculated using the quantitative structure-activity relationship (QSAR) approach with the physicochemical properties $H_{d,i}$, $H_{a,i}$, and P_i which represent hydrogen bond acidity, hydrogen bond basicity, and polarizability of each lumping group i, respectively (Abraham et al., 1991; Abraham and Mcgowan, 1987; Leahy et al., 1992; Platts et al., 1999; Puzyn et al., 2010). Eq. (14) can be rewritten as:

260 $$\ln(K_{w,i}) = -(a_pH_{d,i} + b_pH_{a,i} + r_pP_i + c_p) - \ln(p^o_{L,i}) + \ln\left(\frac{7.501RTOM_{wall}}{10^9MW_{OM}}\right) \tag{15}$$

The values of $H_{d,i}$, $H_{a,i}$, and P_i (Table S2) were calculated using the PaDEL-Descriptor, (Yap, 2011). The value of $K_{w,i}$ is used along with the $k_{on,i}$ to predict intermediate product wall loss using an analytical equation from the study by Han and Jang (2020) as follows:

$$C_{g,i} = \frac{K_{w,i}C_{T,i}}{K_{w,i}+1}e^{-k_{on,i}\left(1+\frac{1}{K_{w,i}}\right)t} + \frac{C_{T,i}}{K_{w,i}+1}$$
(16)

where $C_{g,i}$ (µg m^{-3}) is the gas-phase concentration of a lumping species, i, after time step t (360 s). $C_{T,i}$ (µg m^{-3}) is the sum of $C_{g,i}$ and the concentration of lumping species i on the chamber wall ($C_{w,i}$ (µg m^{-3})).

3.7 UNIPAR procedure for SOA mass production each time step

At each step, $C_{T,i}$ is estimated by using the newly introduced ΔHC and α_i and is combined with the previous step's concentration of lumping species except those used for the formation of OM_{AR} and organic vapor deposition to walls for the simulation of chamber data. Then, the updated $C_{T,i}$ is split into $C_{g,i}$, $C_{org,i}$, and $C_{inorg,i}$ based on multiphase partitioning coefficients as seen in Eq. (6) and Eq. (7). C_{inorg} and C_{org} are then used to form OM_{AR} via oligomerization in both the *inorg* and *org* phases with the rate constants calculated in Eq. (11) and Eq. (12), respectively. In the model, the quantity of the sulfate associated with the esterification of sulfuric acid to form organosulfates in the *inorg* phase is also estimated. Following the process to form OM_{AR}, the remaining concentration of lumping species is used to estimate the organic vapor deposition to the wall using Eq. (16). OM_P is calculated using a Newtonian approach (Eq. 10) in the presence of OM_{AR} and the preexisting OM_0 at the end of each time step. For the total SOA mass, OM_{AR}, OM_P and OM_0 are combined.

4 Results and Discussion

4.1 Chamber data vs. model prediction

UNIPAR model predictions were compared to data collected under various environmental conditions in the UF-APHOR chamber. As seen in Fig. 2, the C10 SOA mass prediction based on the product distributions created using the pre-existing MCM was almost negligible and significantly underestimated from observation. The addition of the alkyl peroxy radical autoxidation mechanism (Section 3.1) was able to considerably improve the prediction of alkane SOA formation. Fig. 2

indicates that alkane SOA mass is mainly attributed to low volatility products produced from alkyl peroxy-radical autoxidation.

285   Fig. 3 together with Fig. 2 shows that UNIPAR, equipped with product distributions derived from the MCM mechanisms with alkyl peroxy radical autoxidation reactions, can reasonably predict final SOA mass produced under various environmental conditions for alkanes C9, C10, C11 and C12. The HC/NO_x (ppbC/ppb) levels in chamber studies range from 1.8 to 22.4 (Table 1). Most experiments in this study were performed with initial HONO concentrations that were about 1/3 of the total NO_x concentration. Thus, higher NO_x experiments had higher concentrations of HONO and this experimental

290   condition apparently skewed the impact of NO_x levels on SOA formation. However, the model is noticeably able to predict SOA mass for a variety of NO_x levels of this study while also considering various HONO concentrations. Under the controlled environment, the impact of NO_x levels on SOA yields is discussed in sect. 4.4 below.

In addition to NO_x levels, the influence of seed conditions on alkane SOA was demonstrated as seen in C9C, C9D, C10B, C10D, C11B, and C11C. No significant impact of acidic seed on SOA formation appeared under similar experimental

295   conditions. For example, non-seeded SOA growth is similar to SA-seeded SOA for C10 alkane (i.e., C10A vs. C10B; C10C vs. C10D, respectively) suggesting that most alkane oxidation products are inert and less polar compared to other oxygenated products from biogenic (Yu et al., 2021; Beardsley and Jang, 2016; Han and Jang, 2022a) and aromatic HCs (Im et al., 2014; Zhou et al. 2019; Han and Jang 2022).

4.2 SOA prediction via product distributions with IVC

300   To predict the SOA formation from a wide range of alkane lengths, the product distributions of linear alkanes were mathematically predicted as a function of carbon number using an IVC as discussed in section 3.3. The resulting product distributions were applied to the UNIPAR model (Fig. 4). Compared to the prediction from the product distribution with MCM and alkyl peroxy radical autoxidation (Fig. 3), the unified simulation of C10 degraded model performance, causing a slightly underprediction for the lower NO_x conditions and an overprediction for the higher NO_x conditions. The unification

305   did not cause significant changes to predictions for C11A and C11C but somewhat improved model performance in case of C11B. C12C was somewhat improved in the simulation with the unified stoichiometric coefficient array but C12A was

overpredicted. It is possible that the cold temperature for experiment C12A (Table 1) caused significant organic vapor wall loss that was unaccounted for in this simulation.

As discussed in sect. 3.3, Eq. 4 for the IVC was fitted to the product distributions of C10, C11, and C12 using unified coefficients a and b to account for the vapor pressure drop in products that results from the increase of the carbon number of the precursor. The vapor pressure drop is not linear as the addition of a carbon to a longer carbon chain causes a smaller decrease in volatility (i.e. the reduction of volatility from adding a carbon to C11 will be smaller compared to the reduction in volatility from adding a carbon to C10). Eq. (4) is non-linear but may not be the ideal equation to model the vapor pressure drop within the product array. Thus, as a consequence of the fitting process, the model performance of the C10 array may have been sacrificed to improve model performance in the case C11, and C12. Furthermore, the current explicit gas mechanism cannot include every possible product, which can be associated with the oxidation on all carbons in each alkane chain. The products of this study are obtained from oxidation at a given carbon position, which is representative of the whole set of gas oxidation. This can cause some deviation of prediction from the true atmospheric oxidation process. Thus, the product distribution originating from these explicit mechanisms can also cause variation in prediction of SOA mass.

4.3 Extrapolation of UNIPAR to longer chain alkanes

The simulation of alkane SOA using the IVC-based product distributions, which were demonstrated with C10-C12 alkanes in section 4.2, was evaluated for prediction of C13 and C15 SOA. Because the explicit mechanisms for the alkanes higher than C12 are not currently available in MCM, the CB6 ozone model (Emery et al., 2015) was implemented to predict the consumption of alkane and the concentrations of HO_2 and RO_2 and used with the UNIPAR model integrated with the IVC-based product distributions. In Figure 5, the simulated C13 and C15 SOA masses were compared to chamber data. Overall, SOA mass was reasonably predicted by using the UNIPAR model integrated with the IVC-based product distribution. Furthermore, SOA simulation was improved when considering the organic-product vapor deposition to the chamber wall. The impact of the chamber wall artifact is generally greater with the longer chain alkanes. For example, the gap between the simulations of SOA mass with and without the wall artifact is larger for C15 compared to C13 as seen in Fig. 5. Notably, the UNIPAR model is unable to predict the early peak in SOA formation that is seen in the chamber data for C12 (Fig. 4), C13 and C15 alkanes (Fig. 5). This initial spike in observed SOA mass is likely because of the fast SOA formation due to the

reaction with a high concentration of hydroxyl radicals, which is produced via photolysis of the high concentration HONO introduced into chamber. This initial spike is generally larger with the longer alkane. The low and mid-volatility products that are formed in the morning will be rapidly condensed via nucleation and non-equilibrium partitioning. The viscosity of

335   the oxidized products tends to be very high, especially for larger alkanes, due to the high molecular weight and the low O:C ratios. The redistribution of organic products between the gas and particle phase to reach equilibrium is achieved via their slow evaporation. This will typically not occur in the ambient atmosphere as the hydroxyl radical is gradually produced via the small amount of HONO in the morning or the ozone mechanism in daytime. A similar trend of rapid SOA formation early in the morning has been observed previously in chamber studies for highly concentrated d-limonene (Yu et al., 2021),

340   which can be rapidly oxidized and form low volatility products. When both the concentrations of HONO and the long-chain alkane are reduced, this spike can be supressed as seen in Fig. S3.

4.4 Sensitivity of SOA formation to NO_x levels, temperature, OM_0, and seed condition

Fig. 6 illustrates the sensitivity of 6 different alkane SOA yields to 6 HC/NO_x levels at a given temperature (298K) under the same sunlight profile (Fig. S4) obtained on 01/20/20 and 30% RH. All sensitivity simulations for a series of alkanes (C10,

345   C12, C14, C16, C18, and C20) were performed with the CB6-UNIPAR model (Table S3) integrated with the IVC-based unified product distribution. The impact of NO_x conditions on linear alkane SOA formation has not been studied extensively in the past. In general, aromatic SOA yields rapidly increase with decreasing the NOx level up to HC ppbC/NOx ppb ≈ 5, reaching to a plateau (Im et al., 2014), while SOA yields for each alkane generally increased as the HC /NO_x level increased over a broad range of the HC ppbC/NOx ppb ratio. Zhang and Seinfeld (2013) found that SOA yields for Dodecane (C12)

350   were higher in lower NO_x conditions (<6 ppb). Cappa et al. (2013) found significant differences in SOA yields between high and low NO_x conditions after 10 hours of oxidation, with higher yields for the low NO_x condition (<1 ppb).

As discussed in Section 4.1, a significant part of alkane SOA is comprised of the low volatility products originating from autoxidation mechanisms. The fraction of autoxidation products to the total SOA mass generally increases with decreasing carbon number or decreasing NO_x level due to gas-particle partitioning of non-autoxidation low volatility products as seen in

355   Fig. S45. Additionally, the impact of the autoxidation reactions on RO_2 chemistry is examined using the Integrated Reaction Rates in Fig S6. Generally, the autoxidation reactions of selected RO_2 species displayed (HO3C106O2, and HO3C126O2

from the MCM mechanisms of C10 and C12, respectively) were insignificant compared to reactions with NO. Additionally, the autoxidation reaction rate was slightly increased under low NO_x conditions compared to high NO_x conditions, which is consistent with Fig. S5.

360   Fig. 7 shows the simulated SOA yields, using UNIPAR with IVC-based product distributions, for a series of alkanes at 3 different temperatures (278K, 288K, 298K) and 2 different NO_x levels (HC ppbC/NO_x ppb= 15, 3) at $OM_0 = 5$. For both high and low NO_x levels, each alkane produces a higher SOA yield at a lower temperature. The volatility of non-autoxidation products from the larger alkane such as eicosane (C20) is low enough to exist primarily in the particle phase, causing the lower sensitivity to temperature compared to smaller alkanes. This result agrees with other studies that have shown a

365   significant impact of temperature on the SOA formation of n-dodecane and n-undecane under low NO_x conditions (Li et al., 2020; Takekawa et al., 2003). For high NO_x conditions, the result also agrees with past work that has found a decent impact of temperature on the SOA formation from n-dodecane at the low OM_0, which is relevant to ambient aerosol (Lamkaddam et al., 2017). Fig. 7 also includes the contribution of OM_P and OM_{AR} to SOA in the absence of inorganic seed at atmospherically relevant conditions ($OM_0 = 5$ $\mu g/m^3$). Overall, the relative contribution of OM_{AR} is higher with lower

370   temperatures and higher carbon numbers. Notably, the levels of SOA loading found in the chamber experiments which were used to validate model performance are much greater than those observed in the ambient environment (up to 32 times higher than the annual PM air quality standard set by the EPA of 12 $\mu g/m^3$ (Epa, 2012)).

Figs. S7 and S8 illustrate the sensitivity of alkane SOA to various pre-existing aerosol mass (OM_0) ranging from 2.5 to 10 $\mu g/m^3$ and aerosol acidity (no seed, wet-AS, and wet-AHS at 60% RH), respectively. SOA yields are significantly increased

375   for each linear alkane as the initial organic matter in the simulation increases. Overall, smaller alkanes have a greater sensitivity to *OM_0*. Alkane oxidation products are generally hydrophobic with a low O:C ratio and very weakly soluble in the wet *inorg* phase. For example, the simulated O:C ratios of alkane SOA in this study ranges from 0.48 and 0.55. Thus, alkane SOA is insignificantly impacted by reactions in the wet *inorg* phase, as seen in Fig. S8, but significantly influenced by gas-aerosol partitioning.

4.5 Uncertainty of model rate constants

In the autoxidation mechanism for C12, three products originating from autoxidation of an alkyl peroxy radical (SR1) dominate the formation of low volatility products to form SOA mass as seen in Table S4: SR3, SR34, and SR40. Previous studies have reported that calculated rate constants for autoxidation reactions have an uncertainty of a factor of 5 (Møller et al., 2016; Praske et al., 2018). To evaluate the uncertainty in SOA formation via autoxidation, the rate constant for the 1st step of autoxidation of an alkyl peroxy radical (SR1) was increased/decreased by a factor of 5. Fig. S9 illustrates the resulting mass changes due to change in this rate constant (SR1): 37.7% increase and 12.4% decrease in SOA mass formed by increasing/decreasing by a factor of 5, respectively for 100 µg/m^3 C12 consumption at HC ppbC/NO$_x$ ppb = 3, and RH = 30%. The rate constant ($k_{AR,in,i}$) of acid-catalyzed oligomerization of organic species i in inorganic phase was increased/decreased by a factor of 2 but the resulting SOA mass shows no significant change for 100 µg/m^3 of C10, C12, C14, C16, C18, and C20 consumptions under the high NO$_x$ conditions (HC ppbC/NO$_x$ ppb = 3). This is expected as alkane oxidation products are lowly soluble in salted *inorg* phase with low reactivity.

4.6 Application of IVC-base product distributions to SOA simulation from diesel-linear alkanes

Diesel fuel is comprised of various lengths of linear alkanes ranging from C9 to C24. In order to simulate diesel SOA formation, the composition of mass weighted linear alkanes (Gentner et al., 2012) in diesel fuel was applied to UNIPAR model in Fig. 8. About 80% SOA mass from diesel linear alkanes is produced from long-chain alkanes (≥C15), with the highest percentage (≈11.8%) from C18, while long chain gaseous alkanes comprise only about 60% of the linear alkane mass. Thus, we conclude that long-chain alkanes are an important source of SOA formation in air parcels originating from diesel fuel combustions. Gentner et al. (2012) also predicted SOA formation from various alkanes in diesel and found that long-chain alkanes make up the vast majority of the SOA mass. Notably, Gentner et al. (2012) found that the peak of SOA mass production from linear alkanes in diesel appears between C19 and C22 which is a slightly higher carbon number compared to our prediction.

4.7 Summary and Conclusions

Several studies in the regional scale simulation have considered alkanes as an important source of SOA (Hodzic et al., 2010; Lee-Taylor et al., 2011; Pye and Pouliot, 2012; Yang et al., 2019). In this study, the UNIPAR model was employed to predict SOA formation from linear alkanes originating from both gas-particle partitioning and particle-phase reactions. Unlike the SOA formation from other precursors (aromatics and biogenics) in which both OM_P and OM_{AR} are significant, alkane SOA is predominantly produced via gas-particle partitioning of organic products. The inclusion of alkyl peroxy-radical autoxidation reactions in the gas mechanism was found to significantly improve SOA mass predictions as seen in Fig. 3. Furthermore, the product distributions created using the MCM mechanisms with alkyl peroxy radical autoxidation reactions were extended to larger alkanes using the IVC approach (Fig. S2). Then, these product distributions were coupled with organic vapor deposition to walls to predict chamber data for C13 and C15 (Fig. 6).

NO_x levels and temperature were found to be the two most influential factors on alkane SOA production (Fig. 6 and 7). At lower NO_x levels, the SOA yield from each of C10, C12, C14, C16, C18, and C20 were found to be higher compared to higher NO_x levels. The impact of temperature was found to be more significant for smaller alkane species as they were more likely to have intermediate volatility products which have partitioning processes that are sensitive to changes in temperature. No significant seed impact appeared regardless of any acidity (Fig. S8). Thus, the reduction of SO_2 emissions would not have a significant impact on linear alkane SOA formation. Furthermore, the reduction of NO_x emissions would serve to increase the alkane SOA yields in ambient air. Although the impact of aerosol-phase reactions on the SOA mass created by linear alkane precursors is insignificant, the SOA mass produced by aerosol-phase reactions from products of other precursors (aromatics and biogenics) in the ambient environment can serve to increase partitioning of the products derived from alkanes. Accordingly, each of the series of linear alkanes, ranging from C10-C20, were sensitive to concentrations of pre-existing OM_0 (Fig. S7).

IVC-based product distributions were extended to C24 in order to simulate the relative SOA formation from the various linear alkanes in diesel fuels. The UNIPAR model predicted that about 80% SOA mass from diesel linear alkanes is produced from long-chain alkanes ($\geq$C15). In addition to linear alkanes, branched and cyclic alkanes are important components of diesel fuel (Gentner et al., 2012) and must be considered when predicting SOA formation in urban

environments. Improvements can also be made on oxidized product distributions originating from linear alkanes. The MCM mechanisms used for linear alkanes in this study are simplified in such a way that, if a specific reaction (i.e. reaction with hydroxyl radical) can occur on multiple locations on a specific compound, only one instance of this reaction pathway is
430 included. The inclusion of each of these different reaction pathways can possibly modulate the product distribution and ultimately increase/decrease the predicted SOA production from the UNIPAR model.

Code availability
435   Code to run the SOA model in this study is available upon request.

Data availability
    The chamber data and simulated results used in this study are available upon request

440 ## Author contributions
MJ designed the experiments, and AM, DD, and MJ carried them out. AM prepared the manuscript with contributions from MJ.

Competing interests
445 The authors declare that they have no conflict of interest.

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References

U.S. EPA. Integrated Science Assessment (ISA) for Particulate Matter 2019.

Abraham, M. H. and McGowan, J. C.: The Use of Characteristic Volumes to Measure Cavity Terms in Reversed Phase
455 Liquid Chromatography, Chromatographia, 23, 243-246, 1987.

Abraham, M. H., Whiting, G. S., Doherty, R. M., and Shuely, W. J.: Hydrogen bonding: XVI. A new solute salvation parameter, π2H, from gas chromatographic data, Journal of Chromatography A, 587, 213-228, https://doi.org/10.1016/0021-9673(91)85158-C, 1991.

460 Aumont, B., Valorso, R., Mouchel-Vallon, C., Camredon, M., Lee-Taylor, J., and Madronich, S.: Modeling SOA formation from the oxidation of intermediate volatility *n*-alkanes, Atmos. Chem. Phys., 12, 7577-7589, 10.5194/acp-12-7577-2012, 2012.

Beardsley, R. L. and Jang, M.: Simulating the SOA formation of isoprene from partitioning and aerosol phase reactions in the presence of inorganics, Atmos. Chem. Phys., 16, 5993-6009, 10.5194/acp-16-5993-2016, 2016.

465 Bianchi, F., Kurtén, T., Riva, M., Mohr, C., Rissanen, M. P., Roldin, P., Berndt, T., Crounse, J. D., Wennberg, P. O., Mentel, T. F., Wildt, J., Junninen, H., Jokinen, T., Kulmala, M., Worsnop, D. R., Thornton, J. A., Donahue, N., Kjaergaard, H. G., and Ehn, M.: Highly Oxygenated Organic Molecules (HOM) from Gas-Phase Autoxidation Involving Peroxy Radicals: A Key Contributor to Atmospheric Aerosol, Chemical Reviews, 119, 3472-3509, 10.1021/acs.chemrev.8b00395, 2019.

Cao, G. and Jang, M.: An SOA Model for Toluene Oxidation in the Presence of Inorganic Aerosols, Environmental Science & Technology, 44, 727-733, 10.1021/es901682r, 2010.

470 Caplain, I., Cazier, F., Nouali, H., Mercier, A., Déchaux, J.-C., Nollet, V., Joumard, R., André, J.-M., and Vidon, R.: Emissions of unregulated pollutants from European gasoline and diesel passenger cars, Atmospheric Environment, 40, 5954-5966, https://doi.org/10.1016/j.atmosenv.2005.12.049, 2006.

Cappa, C. D., Zhang, X., Loza, C. L., Craven, J. S., Yee, L. D., and Seinfeld, J. H.: Application of the Statistical Oxidation Model (SOM) to Secondary Organic Aerosol formation from photooxidation of C_{12} alkanes, Atmos. Chem. Phys., 13, 1591-1606, 10.5194/acp-13-1591-2013, 2013.

475 Clegg, S. L., Brimblecombe, P., and Wexler, A. S.: Thermodynamic Model of the System H+−NH4+−SO42-−NO3-−H2O at Tropospheric Temperatures, The Journal of Physical Chemistry A, 102, 2137-2154, 10.1021/jp973042r, 1998.

Cox, R. A. and Yates, K.: Kinetic equations for reactions in concentrated aqueous acids based on the concept of "excess acidity", Canadian Journal of Chemistry, 57, 2944-2951, 10.1139/v79-479, 1979.

480 Crounse, J. D., Nielsen, L. B., Jørgensen, S., Kjaergaard, H. G., and Wennberg, P. O.: Autoxidation of Organic Compounds in the Atmosphere, The Journal of Physical Chemistry Letters, 4, 3513-3520, 10.1021/jz4019207, 2013.

Han, S. and Jang, M.: Simulating the impact of gas-wall partitioning on SOA formation using the explicit gas mechanism integrated with aqueous reactions containing electrolytes, Science of The Total Environment, 748, 141360, https://doi.org/10.1016/j.scitotenv.2020.141360, 2020.

485 De Schrijver, F. and Smets, G.: Polymerization kinetics in highly viscous media, Journal of Polymer Science Part A-1: Polymer Chemistry, 4, 2201-2210, https://doi.org/10.1002/pol.1966.150040914, 1966.

Emery, C., Jung, J., Koo, B., and Yarwood, G.: Improvements to CAMx Snow Cover Treatments and Carbon Bond Chemical Mechanism for Winter Ozone, 2015.

EPA, U. S.: 2012 National Ambient Air Quality Standards (NAAQS) for Particulate Matter (PM), 2012.

490 Gentner Drew, R., Isaacman, G., Worton David, R., Chan Arthur, W. H., Dallmann Timothy, R., Davis, L., Liu, S., Day Douglas, A., Russell Lynn, M., Wilson Kevin, R., Weber, R., Guha, A., Harley Robert, A., and Goldstein Allen, H.: Elucidating secondary organic aerosol from diesel and gasoline vehicles through detailed characterization of organic carbon emissions, Proceedings of the National Academy of Sciences, 109, 18318-18323, 10.1073/pnas.1212272109, 2012.

Han, S. and Jang, M.: Modeling Diurnal Variation of SOA Formation via Multiphase Reactions of Biogenic Hydrocarbons, 495 Atmos. Chem. Phys. Discuss., 2022, 1-25, 10.5194/acp-2022-327, 2022a.

Han, S. and Jang, M.: Prediction of secondary organic aerosol from the multiphase reaction of gasoline vapor by using volatility–reactivity base lumping, Atmospheric Chemistry and Physics, 22, 625-639, 2022b.

Han, S. and Jang, M.: Simulating the impact of gas-wall partitioning on SOA formation using the explicit gas mechanism integrated with aqueous reactions containing electrolytes, Science of The Total Environment, 748, 141360, https://doi.org/10.1016/j.scitotenv.2020.141360, 2020.

500

Hodzic, A., Jimenez, J. L., Madronich, S., Canagaratna, M. R., DeCarlo, P. F., Kleinman, L., and Fast, J.: Modeling organic aerosols in a megacity: potential contribution of semi-volatile and intermediate volatility primary organic compounds to secondary organic aerosol formation, Atmos. Chem. Phys., 10, 5491-5514, 10.5194/acp-10-5491-2010, 2010.

Im, Y., Jang, M., and Beardsley, R. L.: Simulation of aromatic SOA formation using the lumping model integrated with explicit gas-phase kinetic mechanisms and aerosol-phase reactions, Atmos. Chem. Phys., 14, 4013-4027, 10.5194/acp-14-4013-2014, 2014.

505

Jang, M. and Kamens, R. M.: A Thermodynamic Approach for Modeling Partitioning of Semivolatile Organic Compounds on Atmospheric Particulate Matter: Humidity Effects, Environmental Science & Technology, 32, 1237-1243, 10.1021/es970773w, 1998.

Jang, M., Czoschke, N. M., and Northcross, A. L.: Semiempirical Model for Organic Aerosol Growth by Acid-Catalyzed Heterogeneous Reactions of Organic Carbonyls, Environmental Science & Technology, 39, 164-174, 10.1021/es048977h, 2005.

510

Jang, M., Czoschke, N. M., Lee, S., and Kamens, R. M.: Heterogeneous Atmospheric Aerosol Production by Acid-Catalyzed Particle-Phase Reactions, Science, 298, 814-817, doi:10.1126/science.1075798, 2002.

Kanakidou, M., Seinfeld, J. H., Pandis, S. N., Barnes, I., Dentener, F. J., Facchini, M. C., Van Dingenen, R., Ervens, B., Nenes, A., Nielsen, C. J., Swietlicki, E., Putaud, J. P., Balkanski, Y., Fuzzi, S., Horth, J., Moortgat, G. K., Winterhalter, R., Myhre, C. E. L., Tsigaridis, K., Vignati, E., Stephanou, E. G., and Wilson, J.: Organic aerosol and global climate modelling: a review, Atmos. Chem. Phys., 5, 1053-1123, 10.5194/acp-5-1053-2005, 2005.

515

Lamkaddam, H., Gratien, A., Pangui, E., Cazaunau, M., Picquet-Varrault, B., and Doussin, J.-F.: High-NOx Photooxidation of n-Dodecane: Temperature Dependence of SOA Formation, Environmental Science & Technology, 51, 192-201, 10.1021/acs.est.6b03821, 2017.

520

Leahy, D. E., Morris, J. J., Taylor, P. J., and Wait, A. R.: Model solvent systems for QSAR. Part 3. An LSER analysis of the 'critical quartet.' New light on hydrogen bond strength and directionality, Journal of the Chemical Society, Perkin Transactions 2, 705-722, 10.1039/p29920000705, 1992.

Lee-Taylor, J., Madronich, S., Aumont, B., Baker, A., Camredon, M., Hodzic, A., Tyndall, G. S., Apel, E., and Zaveri, R. A.: Explicit modeling of organic chemistry and secondary organic aerosol partitioning for Mexico City and its outflow plume, Atmos. Chem. Phys., 11, 13219-13241, 10.5194/acp-11-13219-2011, 2011.

525

Li, J., Li, K., Li, H., Wang, X., Wang, W., Wang, K., and Ge, M.: Long-chain alkanes in the atmosphere: A review, Journal of Environmental Sciences, 114, 37-52, https://doi.org/10.1016/j.jes.2021.07.021, 2022.

Li, J., Wang, W., Li, K., Zhang, W., Peng, C., Zhou, L., Shi, B., Chen, Y., Liu, M., Li, H., and Ge, M.: Temperature effects on optical properties and chemical composition of secondary organic aerosol derived from n-dodecane, Atmos. Chem. Phys., 20, 8123-8137, 10.5194/acp-20-8123-2020, 2020.

530

Lim, Y. B. and Ziemann, P. J.: Effects of Molecular Structure on Aerosol Yields from OH Radical-Initiated Reactions of Linear, Branched, and Cyclic Alkanes in the Presence of NOx, Environmental Science & Technology, 43, 2328-2334, 10.1021/es803389s, 2009.

535

Loza, C. L., Craven, J. S., Yee, L. D., Coggon, M. M., Schwantes, R. H., Shiraiwa, M., Zhang, X., Schilling, K. A., Ng, N. L., Canagaratna, M. R., Ziemann, P. J., Flagan, R. C., and Seinfeld, J. H.: Secondary organic aerosol yields of 12-carbon

alkanes, Atmos. Chem. Phys., 14, 1423-1439, 10.5194/acp-14-1423-2014, 2014.McMurry, P. H. and Grosjean, D.: Gas and aerosol wall losses in Teflon film smog chambers, Environmental science & technology, 19, 1176-1182, 1985.

540   Møller, K. H., Otkjær, R. V., Hyttinen, N., Kurtén, T., and Kjaergaard, H. G.: Cost-Effective Implementation of Multiconformer Transition State Theory for Peroxy Radical Hydrogen Shift Reactions, The Journal of Physical Chemistry A, 120, 10072-10087, 10.1021/acs.jpca.6b09370, 2016.

Nel, A.: Air Pollution-Related Illness: Effects of Particles, Science, 308, 804-806, 10.1126/science.1108752, 2005.

Odian, G.: Principles of polymerization, John Wiley & Sons2004.

545   Pankow, J. F.: An absorption model of gas/particle partitioning of organic compounds in the atmosphere, Atmospheric Environment, 28, 185-188, https://doi.org/10.1016/1352-2310(94)90093-0, 1994.

Platts, J. A., Butina, D., Abraham, M. H., and Hersey, A.: Estimation of Molecular Linear Free Energy Relation Descriptors Using a Group Contribution Approach, Journal of Chemical Information and Computer Sciences, 39, 835-845, 10.1021/ci980339t, 1999.

550   Praske, E., Otkjær, R. V., Crounse, J. D., Hethcox, J. C., Stoltz, B. M., Kjaergaard, H. G., and Wennberg, P. O.: Atmospheric autoxidation is increasingly important in urban and suburban North America, Proceedings of the National Academy of Sciences, 115, 64-69, 10.1073/pnas.1715540115, 2018.

Press, W. H., Flannery, B. P., Teukolsky, S. A., and Vetterling, W. T.: Newton-Raphson method using derivative and Newton-Raphson method for nonlinear systems of equations. Numerical Recipes in FORTRAN: The Art of Scientific 555   Computing, 1992.

Puzyn, T., Leszczynski, J., and Cronin, M. T.: Recent advances in QSAR studies: methods and applications, 2010.

Pye Havala, O. T., D'Ambro Emma, L., Lee Ben, H., Schobesberger, S., Takeuchi, M., Zhao, Y., Lopez-Hilfiker, F., Liu, J., Shilling John, E., Xing, J., Mathur, R., Middlebrook Ann, M., Liao, J., Welti, A., Graus, M., Warneke, C., de Gouw Joost, A., Holloway John, S., Ryerson Thomas, B., Pollack Ilana, B., and Thornton Joel, A.: Anthropogenic enhancements to 560   production of highly oxygenated molecules from autoxidation, Proceedings of the National Academy of Sciences, 116, 6641-6646, 10.1073/pnas.1810774116, 2019.

Pye, H. O. T. and Pouliot, G. A.: Modeling the Role of Alkanes, Polycyclic Aromatic Hydrocarbons, and Their Oligomers in Secondary Organic Aerosol Formation, Environmental Science & Technology, 46, 6041-6047, 10.1021/es300409w, 2012.

Reid, J. P., Bertram, A. K., Topping, D. O., Laskin, A., Martin, S. T., Petters, M. D., Pope, F. D., and Rovelli, G.: The 565   viscosity of atmospherically relevant organic particles, Nature Communications, 9, 956, 10.1038/s41467-018-03027-z, 2018.

Robinson Allen, L., Donahue Neil, M., Shrivastava Manish, K., Weitkamp Emily, A., Sage Amy, M., Grieshop Andrew, P., Lane Timothy, E., Pierce Jeffrey, R., and Pandis Spyros, N.: Rethinking Organic Aerosols: Semivolatile Emissions and Photochemical Aging, Science, 315, 1259-1262, 10.1126/science.1133061, 2007.

Roldin, P., Ehn, M., Kurtén, T., Olenius, T., Rissanen, M. P., Sarnela, N., Elm, J., Rantala, P., Hao, L., Hyttinen, N., 570   Heikkinen, L., Worsnop, D. R., Pichelstorfer, L., Xavier, C., Clusius, P., Öström, E., Petäjä, T., Kulmala, M., Vehkamäki, H., Virtanen, A., Riipinen, I., and Boy, M.: The role of highly oxygenated organic molecules in the Boreal aerosol-cloud-climate system, Nature Communications, 10, 4370, 10.1038/s41467-019-12338-8, 2019.

Schell, B., Ackermann, I. J., Hass, H., Binkowski, F. S., and Ebel, A.: Modeling the formation of secondary organic aerosol within a comprehensive air quality model system, Journal of Geophysical Research: Atmospheres, 106, 28275-28293, 575   https://doi.org/10.1029/2001JD000384, 2001.

Shrivastava, M., Cappa, C. D., Fan, J., Goldstein, A. H., Guenther, A. B., Jimenez, J. L., Kuang, C., Laskin, A., Martin, S. T., Ng, N. L., Petaja, T., Pierce, J. R., Rasch, P. J., Roldin, P., Seinfeld, J. H., Shilling, J., Smith, J. N., Thornton, J. A., Volkamer, R., Wang, J., Worsnop, D. R., Zaveri, R. A., Zelenyuk, A., and Zhang, Q.: Recent advances in understanding

secondary organic aerosol: Implications for global climate forcing, Reviews of Geophysics, 55, 509-559, https://doi.org/10.1002/2016RG000540, 2017.

Takekawa, H., Minoura, H., and Yamazaki, S.: Temperature dependence of secondary organic aerosol formation by photo-oxidation of hydrocarbons, Atmospheric Environment, 37, 3413-3424, https://doi.org/10.1016/S1352-2310(03)00359-5, 2003.

Volkamer, R., Jimenez, J. L., San Martini, F., Dzepina, K., Zhang, Q., Salcedo, D., Molina, L. T., Worsnop, D. R., and Molina, M. J.: Secondary organic aerosol formation from anthropogenic air pollution: Rapid and higher than expected, Geophysical Research Letters, 33, https://doi.org/10.1029/2006GL026899, 2006.

World Health, O.: Ambient air pollution: a global assessment of exposure and burden of disease, World Health Organization, Geneva2016.

Wu, L., Wang, X., Lu, S., Shao, M., and Ling, Z.: Emission inventory of semi-volatile and intermediate-volatility organic compounds and their effects on secondary organic aerosol over the Pearl River Delta region, Atmos. Chem. Phys., 19, 8141-8161, 10.5194/acp-19-8141-2019, 2019.

Xavier, C., Rusanen, A., Zhou, P., Dean, C., Pichelstorfer, L., Roldin, P., and Boy, M.: Aerosol mass yields of selected biogenic volatile organic compounds – a theoretical study with nearly explicit gas-phase chemistry, Atmos. Chem. Phys., 19, 13741-13758, 10.5194/acp-19-13741-2019, 2019.

Yang, W., Li, J., Wang, W., Li, J., Ge, M., Sun, Y., Chen, X., Ge, B., Tong, S., Wang, Q., and Wang, Z.: Investigating secondary organic aerosol formation pathways in China during 2014, Atmospheric Environment, 213, 133-147, https://doi.org/10.1016/j.atmosenv.2019.05.057, 2019.

Yap, C. W.: PaDEL-descriptor: An open source software to calculate molecular descriptors and fingerprints, Journal of Computational Chemistry, 32, 1466-1474, https://doi.org/10.1002/jcc.21707, 2011.

Yu, Z., Jang, M., Zhang, T., Madhu, A., and Han, S.: Simulation of Monoterpene SOA Formation by Multiphase Reactions Using Explicit Mechanisms, ACS Earth and Space Chemistry, 2021.

Zhang, X. and Seinfeld, J. H.: A functional group oxidation model (FGOM) for SOA formation and aging, Atmos. Chem. Phys., 13, 5907-5926, 10.5194/acp-13-5907-2013, 2013.

Zhou, C., Jang, M., and Yu, Z.: Simulation of SOA formation from the photooxidation of monoalkylbenzenes in the presence of aqueous aerosols containing electrolytes under various NOx levels, Atmos. Chem. Phys., 19, 5719-5735, 10.5194/acp-19-5719-2019, 2019.

Zuend, A., Marcolli, C., Booth, A. M., Lienhard, D. M., Soonsin, V., Krieger, U. K., Topping, D. O., McFiggans, G., Peter, T., and Seinfeld, J. H.: New and extended parameterization of the thermodynamic model AIOMFAC: calculation of activity coefficients for organic-inorganic mixtures containing carboxyl, hydroxyl, carbonyl, ether, ester, alkenyl, alkyl, and aromatic functional groups, Atmos. Chem. Phys., 11, 9155-9206, 10.5194/acp-11-9155-2011, 2011.

Table 1. Summary of experimental conditions and observed data for experiments performed in the UF-APHOR outdoor chamber.

ID	Date	Alkane	Seed type[a]	Seed mass (μg/m³)	Initial HC (ppb)	HC/NOx (ppbC/ppb)	HONO (ppb)	Temp (K)	RH (%)	OM (μg/m³)	Yield[b] (%)	Comments
C9A	05/07/21	C9	None	0	130	7.2	45	286-312	34-94	2.0	0.8	Figs 3, 4
C9B	05/07/21	C9	None	0	147	1.9	200	287-310	43-99	4.1	2.4	Figs 3, 4
C9C	05/18/21	C9	SA	81	121	18.8	50	291-316	22-77	13.0	4.1	Figs 3, 4
C9D	05/18/21	C9	SA	100	149	7.2	150	291-315	29-84	16.6	3.2	Figs 3, 4
C10A	05/29/21	C10	None	0	313	14.8	70	292-317	24-81	25.6	2.3	Figs 2, 4
C10B	05/29/21	C10	SA	34	323	15.8	63	292-317	31-90	31.6	1.9	Figs 3, 4
C10C	06/01/21	C10	None	0	312	4.5	250	290-316	23-84	37.7	2.5	Figs 2, 4
C10D	06/01/21	C10	SA	35	313	4.1	250	291-314	29-92	28.1	1.7	Figs 3, 4
C11A	09/05/20	C11	None	0	140	9.3	80	297-321	21-52	47.4	7.0	Figs 3, 4
C11B	09/18/20	C11	SA	52	202	22.4	40	297-325	12-42	44.3	7.0	Figs 3, 4
C11C	09/18/20	C11	SA	60	209	4.0	106	298-321	21-52	30.1	3.1	Figs 3, 4
C12A	01/08/20	C12	None	0	374	5.4	175	277-302	22-72	44.1	2.8	Figs 3, 4
C12B	06/07/22	C12	None	0	159	2.3	97	293-321	23-91	19.0	2.4	Figs 3, 4
C12C	06/07/22	C12	None	0	135	4.9	200	293-320	32-99	123.2	20.6	Figs 3, 4
C13A	10/18/20	C13	None	0	207	17.1	45	290-315	26-89	286.9	11.3	Fig 5
C13B	10/18/20	C13	SA	35	202	16.6	45	291-315	31-95	195.8	8.7	Fig 5
C15A	02/17/22	C15	None	0	202	4.3	265	288-307	30-78	388.8	46.8	Fig 5
C15B	02/17/22	C15	None	0	202	8.7	117	289-316	27-95	206.7	28.5	Fig 5

615 [a] Experiments were performed with no seed (none) and sulfuric acid seed (SA). [b] Yield was calculated as the ratio between the concentration of the final measured SOA mass (μg/m³) and the concentration of precursor alkane consumed (μg/m³).
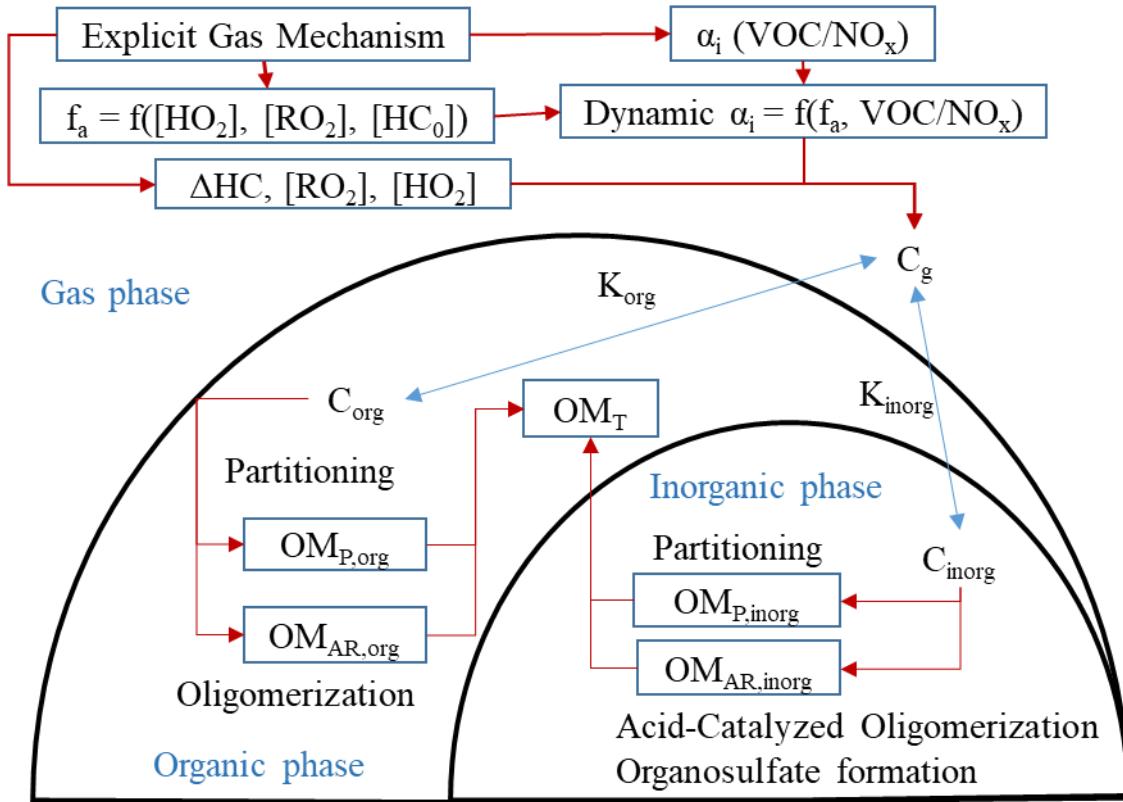
Figure 1. A simplified scheme of the UNIPAR model. [HC]$_0$ represents the initial hydrocarbon (HC) concentration. A gas kinetic mechanism (MCM v3.3.1 + autoxidation of alkyl peroxy radicals) is used to simulate the consumption of HC (ΔHC) and the concentrations of hydroperoxide radical ([HO$_2$]) and organic peroxyl radical ([RO$_2$]). The explicit gas mechanism of each alkane is used to derive the dynamic mass-based stoichiometric coefficient (dynamic α_i) of lumping species i as a function of HC/NO$_x$ and the aging factor f_a. f_a is represented as a function of [HO$_2$], [RO$_2$], and [HC]$_0$ (Zhou et al., 2019). Subscripts g, org, and $inorg$ denote the gas, organic, and inorganic phases, respectively. K_{org} and K_{inorg} represent the partitioning coefficients of lumping species to the org phase and $inorg$ phase, respectively. C_{org} and C_{inorg} represent the concentrations of lumping species in the org and $inorg$ phases, respectively. $OM_{P,org}$ and $OM_{P,inorg}$ represent the mass of organic matter (OM) present in org and $inorg$ phases, respectively due to partitioning. $OM_{Ar,org}$ and $OM_{AR,inorg}$ represent the OM formed in the org phase due to in-particle chemistry such as oligomerization, and $inorg$ phase due to acid-catalyzed oligomerization and organosulfate formation (Beardsley and Jang, 2016; Im et al., 2014; Zhou et al., 2019). OM_T represents the total SOA mass formed due to partitioning and aerosol-phase reactions in both the $inorg$ and org phases.
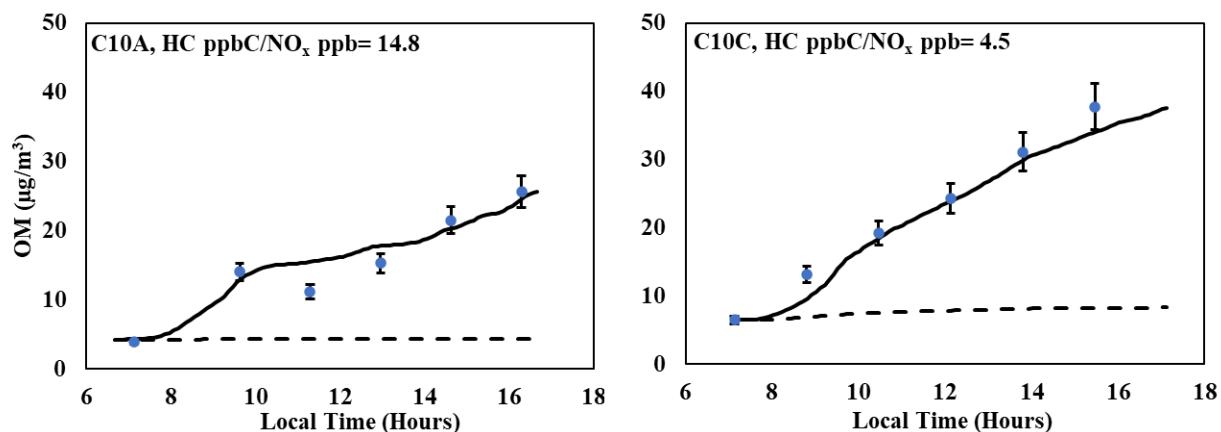
Figure 2. Comparison of SOA mass produced between simulations using a product distribution for C10 derived from the MCM (dashed line) and one from the MCM with alkyl peroxy radical autoxidation reactions added (solid line). The blue dots represent observed SOA data collected that are corrected for particle wall loss to the chamber. Error bars represent 9% uncertainty associated with the OC/EC and particle wall loss correction.

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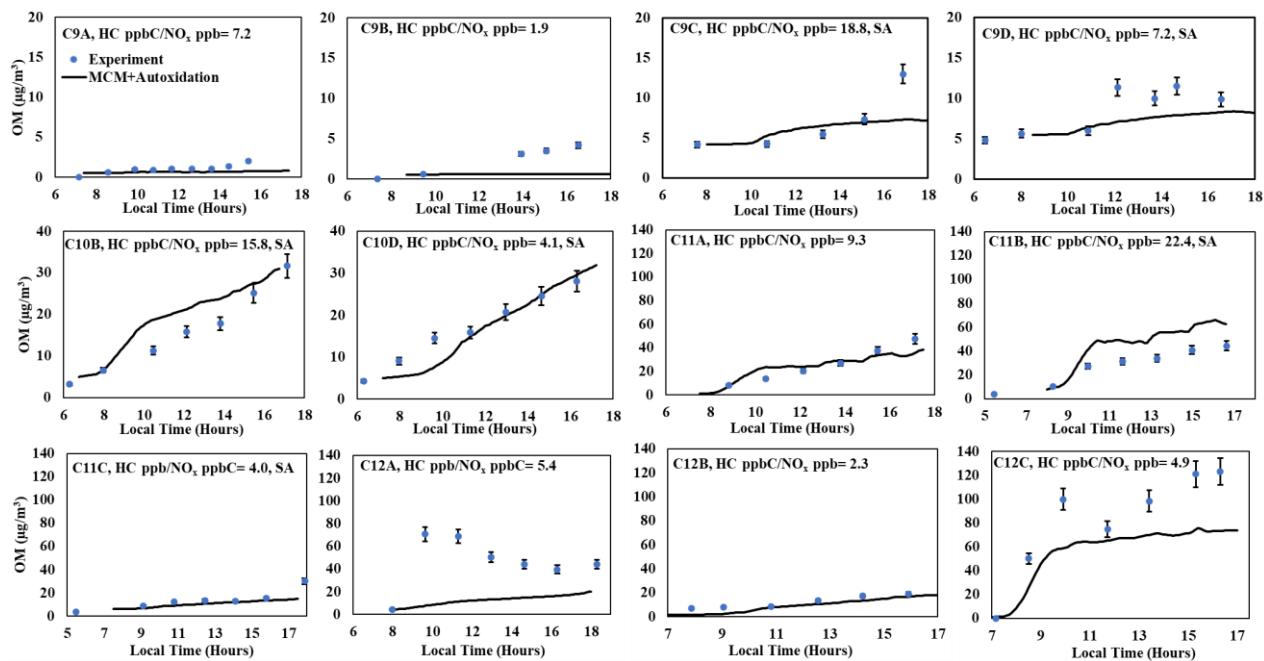
Figure 3. Simulated SOA mass for alkanes C9, C10, and C12 using a product distribution derived from the MCM mechanisms with alkyl peroxy radical autoxidation reactions added. The blue dots represent observed SOA data collected that are corrected for particle wall loss to the chamber. Error bars represent 9% uncertainty associated with the OC/EC measurement and particle wall loss correction.
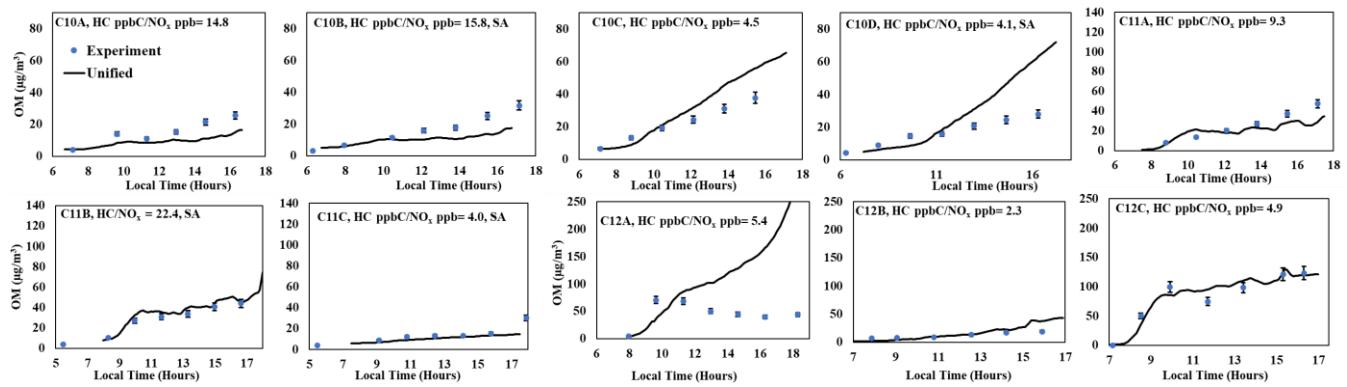
Figure 4. Comparison of SOA mass produced between simulations, for C10, C11, and C12, using the unified product distributions based on the IVC (black line) to chamber data. The blue dots represent observed SOA data collected that are corrected for particle wall loss to the chamber. Error bars represent 9% uncertainty associated with the OC/EC and particle wall loss correction.

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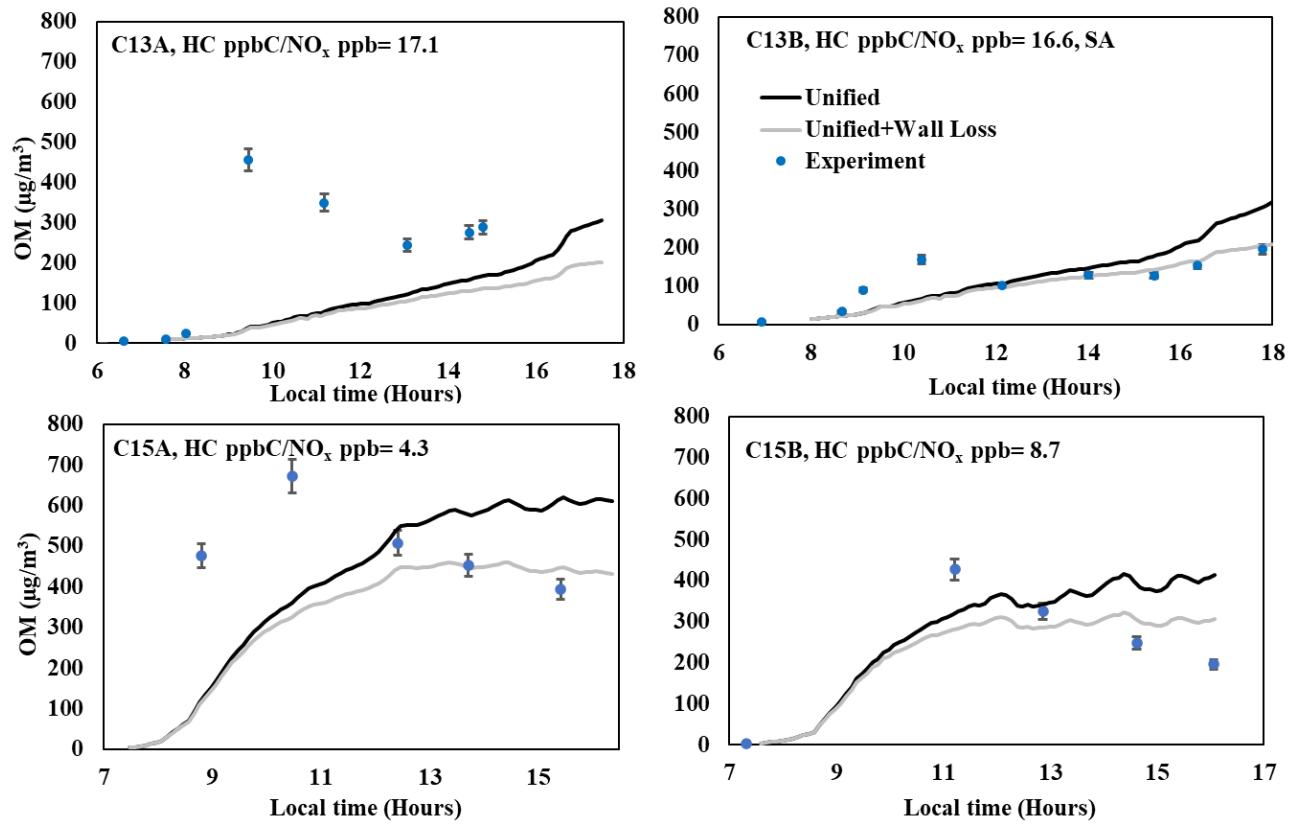
Figure 5. Comparison of the observed and the predicted SOA mass produced from the photooxidation of C13 and C15 alkanes under different experimental conditions. The SOA simulation was performed by using the unified product distributions that is constructed with the IVC approach in the presence of precursor vapor deposition to the chamber wall. In addition, SOA simulation was performed in the presence (gray line) and the absence (black line) of

665 **organic product loss to the wall. The observed SOA mass (dots) was corrected for gas dilution and the particle loss to the chamber wall. ASCM was used to measure SOA mass concentrations for C13 and C15 instead of OC/EC data as the artifacts via off-gassing of semi-volatiles to the denuder upstream OC for C13 and C15 oxidation products. The collection efficiency of organic alkane aerosol for the ACSM was assumed to be 66% for C13 and 63% for C15. The collection efficiency was determined by comparison SMPS data with ACSM data with the aerosol originating from**

670 **non-volatile organics. The collection efficiency of C13B aerosol, which was produced in the presence of acidic seed, was assumed to be 100%. The error associated with OM (6%) is calculated with ACSM data.**

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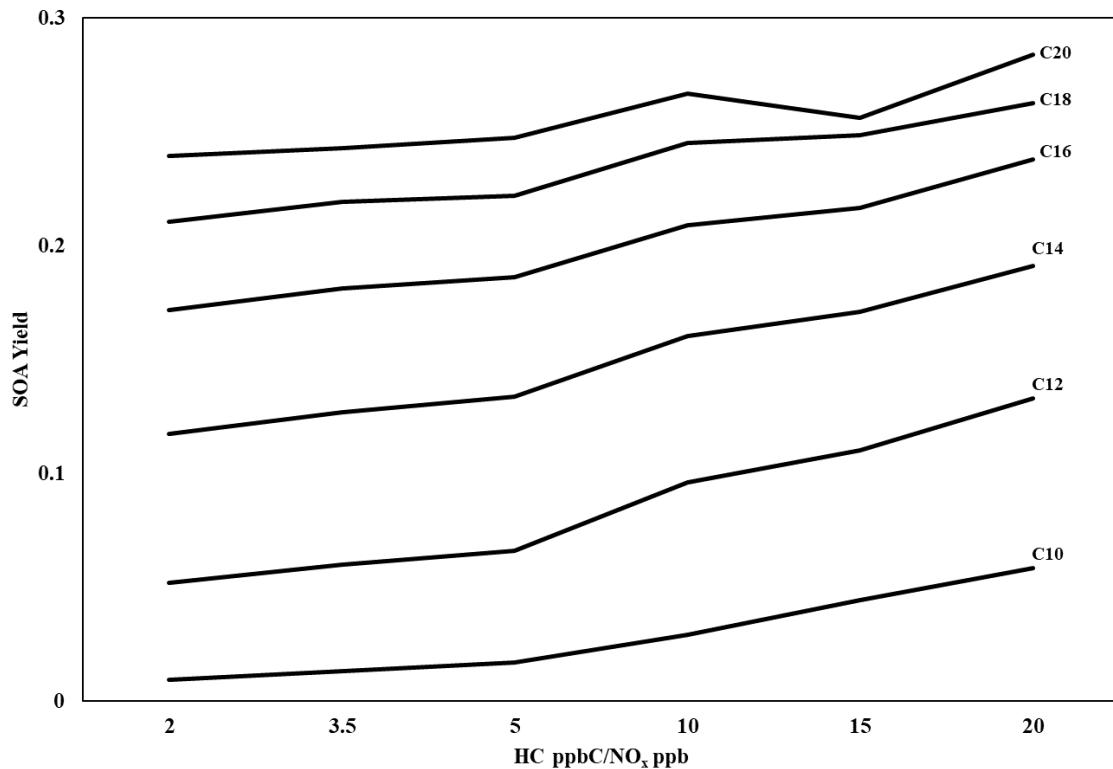
Figure 6. Simulated SOA Yields for a series of linear alkanes at 6 different HC/NO$_x$ conditions using the IVC-based product distributions. OM$_0$ = 5 μg/m^3, 298 K, RH = 30%.

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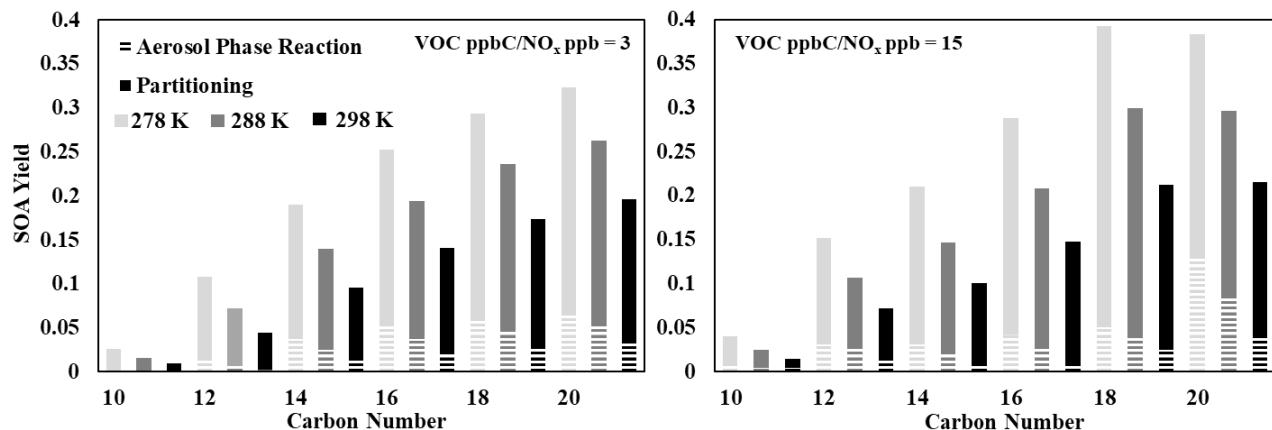
Figure 7. Simulated SOA Yields, separated by OM$_P$ and OM$_{AR}$, for a series of linear alkanes at 3 different temperatures and 2 different HC ppbC/NO$_x$ ppb conditions using the IVC-based product distributions. OM$_0$ = 5 µg/m^3, RH = 30%.

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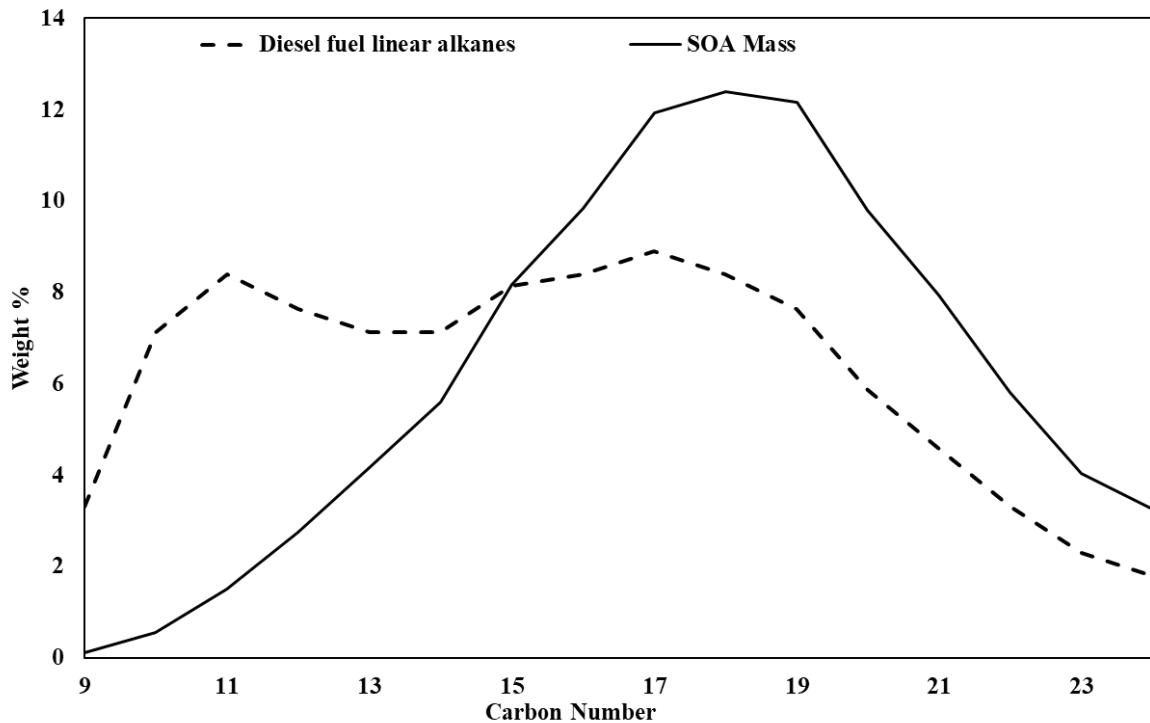
Figure 8. The predicted relative SOA mass formation from the composition of linear alkanes in diesel fuel. The simulation was performed using the CB6-UNIPAR model equipped with IVC-based product distributions. The relative composition of linear alkanes in diesel fuel used is from Gentner et al. (2012). HC ppbC /NO_x ppb = 3, RH = 60%, OM_0 = 5 μg/m^3.