

1 A new analytical method for stability analysis of rock blocks with 2 cavities in sub-horizontal strata by the considering eccentricity effect

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9 **Abstract.** The basal cavity of a rock block formed due to differential weathering is an important predisposing factor for rockfall
10 in hard-soft interbedded rocks, which induces eccentricity effect at the base of the rock block. Rock block falling due to the
11 eccentricity effect with the failure modes of toppling or sliding is defined as biased rockfall in this study. Taking into
12 account Considering the non-uniform stress distribution due to the eccentricity effect, a new analytical method is proposed for
13 three-dimensional stability force and stability of biased rockfall analysis of biased rockfall is proposed. The development of
14 non-uniform stress distribution stress calculated by this analytical method was verified by numerical simulation. The biased
15 rockfall progresses from partial damage of the soft underlying layer, caused by non-uniform distributed stress, to toppling and
16 sliding of overhanging hard rock block due to overall unbalanced force. Therefore, a set of factors of safety (Fos) against
17 partial damage (compressive and tensile damage of the soft underlying layer) and overall failure (toppling and sliding of the
18 hard rock block) are used to determine the rockfall susceptibility level. The analytical method is applied and validated with
19 using biased rockfalls on the northeast edge of the Sichuan Basin in Southwest China, where large a significant number amounts
20 of rockfalls consisting of overhanging thick sandstone and underlying mudstone occur have developed, composed of
21 overlying overhanging thick sandstone and underlying mudstone. The evolution process of biased rockfalls is divided into four
22 stages, initial state, cavity formation, partial unstable and failure. The proposed method is validated by calculating Fos of the
23 typical unstable rock blocks in the study area. As the cavity continues to grow, The the continuous retreat of cavity causes
24 stress redistribution between the hard and soft rock layers. This results in damage to the underlying soft rock layer due to the
25 development of the eccentricity effect, ultimately leading to the failure of the hard rock block. The critical cavity retreat ratio
26 is determined to be 0.33, which is used to classify the low and moderate rockfall susceptibility in the eastern Sichuan
27 Basin. Consequently, the development of the eccentricity effect leads to damage to the underlying soft rock layer and further
28 failure of the hard rock block. The critical cavity retreat ratio is determined to be 0.33 to classify the low and moderate rockfall
29 susceptibility in the eastern Sichuan Basin. The proposed analytical method is effective for the early identification of biased
30 rockfall provides insights into the evolution of biased rockfall and a means for early identification and susceptibility assessment
31 of rockfall, which is significant for rockfall prevention and risk mitigation.

32 **List of symbols**

33	a	length of the block along the x direction
34	A	area of contact surfaces
35	b	width of the block along the y direction
36	c	cohesive force of the mudstone
37	d_i	width of the basal cavity in a certain direction
38	e_x	eccentric distance along the x direction
39	e_y	eccentric distance along the y direction
40	E_x	horizontal seismic force along the x direction
41	Fos	factor of safety
42	h	height of the block
43	h_w	height of the water in the fracture
44	H_x	water pressure along the x direction
45	I_x	moment of inertia with respect to the x -axis
46	I_y	moment of inertia with respect to the y -axis
47	k_e	earthquake contribution coefficient
48	k_1	rainfall coefficient, taking 1 in the rainfall scenario and 0 in the non-rainfall scenario
49	k_2	earthquake coefficient, taking 1 in the seismic scenario and 0 in the non-seismic scenario
50	k_3	free surface coefficient, taking 1 for two free surfaces and 0 for three free surfaces
51	M_{bx}	total bending moments with respect to the x -axis on the mudstone foundation
52	M_{by}	total bending moments with respect to the y -axis on the mudstone foundation
53	M_{bEx}	bending moment of E_x with respect to the x -axis on the mudstone foundation
54	M_{bHx}	bending moment of H_x with respect to the x -axis on the mudstone foundation
55	M_{bWx}	bending moment of W with respect to the x -axis on the mudstone foundation
56	M_{Ex}	overturning moment provided by E_x along the x direction
57	M_{Hx}	overturning moment provided by H_x along the x direction
58	M_{px}	stabilizing moment of p_n along the x direction
59	$M_{W_{inx}}$	stabilizing moment provided by W along the x direction
60	$M_{W_{outx}}$	overturning moment provided by W along the x direction
61	N_z	total applied vertical load on the mudstone base
62	O	origin of the (x, y) coordinates
63	$p(x, y)$	pressure magnitude at point (x, y)
64	r_i	the basal cavity retreat ratio equal to the ratio of cavity width to block width in a certain direction

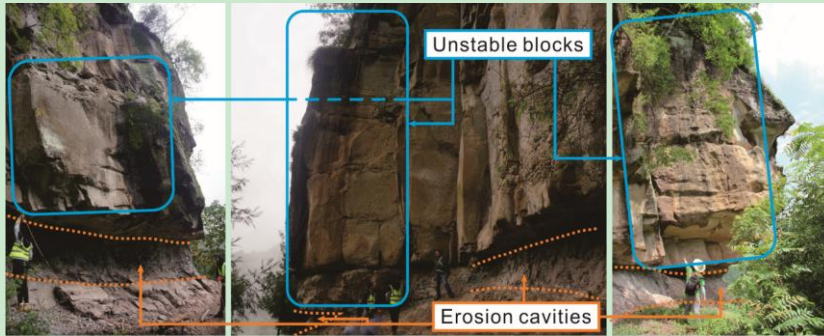
65	W	weight of the block
66	x	distance to O along the x -axis
67	y	distance to O along the y -axis
68	α	true dip of the contact surface
69	γ_s	unit weight of sandstone
70	γ_w	unit weight of water
71	θ_1	apparent dip of α on plane J1
72	θ_2	apparent dip of α on plane J2
73	σ_{cmax}	ultimate compressive strength of the mudstone
74	σ_{tmax}	ultimate tensile strength of the mudstone
75	τ_{max}	ultimate shear strength of the mudstone
76	φ	friction angle of the mudstone
77	ω_1	angle between the trend of the contact surface and the x direction
78	ω_2	angle between the trend of the contact surface and the y direction

79 1 Introduction

80 Rockfall is defined as the detachment of a rock block from a steep slope along a surface, on which little or no shear
81 displacement takes place (Cruden and Varnes, 1996). Rockfalls frequently occur in mountainous ranges, cut slopes, and coastal
82 cliffs, and they may cause significant facility damage and casualties in residential areas and transport corridors (Chau et al.,
83 2003; Volkwein et al., 2011; Corominas et al., 2018). Stability analysis of rock blocks are crucial for risk management and
84 early warning of rockfall (Kromer et al., 2017).

85 Rockfalls are prone to occur in soft-hard rock formations, and the non-uniform stress distribution caused by differential
86 weathering of rock formations is the main reason for the failure of rockfall. In the eastern Sichuan Basin, Southwest China,
87 rockfall is widespread and poses high risk (Chen et al., 2008; Chen and Tang, 2010; Zhang et al., 2016; Zhou et al., 2017;
88 Zhou et al., 2018). The rockfall in this area is attributed to the tectonic setting of Jura-type folds and the stratum sequence,
89 which is characterized by the interbedding of hard and soft layers. An alternation of thick sandstone and thin mudstone layers
90 is formed in the wide and gentle-angle synclines (Zhang et al., 2015; Wu et al., 2018). Weathering is known to be one of the
91 main predisposing factors for rockfall (Jaboyedoff et al., 2021; Zhan et al., 2022). The cliff comprised of hard sandstone is the
92 source of rockfall, and the underlying mudstone is more susceptible to weathering. Along with the retreat of basal cavities in
93 the mudstone layer, the gravity centre of the overlyingoverhanging sandstone block moves outward relative to the mudstone.
94 In this case, the stress distribution in the contact surface of sandstone and mudstone is non-uniform. The mudstone on the outer
95 side bears higher compressive stress than that on the inner side. This phenomenon can be defined as an eccentricity effect,
96 which leads to mudstone damage and failure of the overlyingoverhanging sandstone by toppling or sliding. This type of rockfall

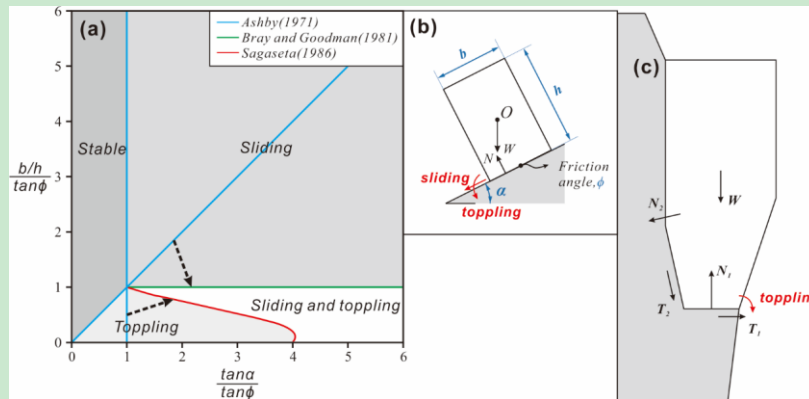
97 is defined as biased rockfall in this study (Fig. 1). Similar rockfall patterns have been widely reported in other regions, such as
98 Joss Bay in England (Hutchinson, 1972), Okinawa Island in Japan (Kogure et al., 2006), and the Colorado Plateau of the
99 southwestern United States (Ward et al., 2011). Retreat of the basal cavity is a main cause for the failure of the
100 overlying overhanging block. Therefore, it is necessary to establish an analytical method, considering the development of the
101 basal cavity, to analyse the stress distribution and stability of rock blocks, which is fundamental to the susceptibility assessment
102 and risk control of biased rockfall.



103
104 **Figure 1** Potential unstable blocks and basal cavities caused by differential weathering.

105 Rockfall stability analysis methods include statistical analysis (Frattoni et al., 2008; Santi et al., 2009), empirical rating systems
106 (Pierson et al., 1990; Ferrari et al., 2016), and mechanical analysis (Jaboyedoff et al., 2004; Derron et al., 2005; Matasci et al.,
107 2018). The statistical analysis and empirical rating systems are suitable for rockfall hazard assessment at a regional scale. The
108 accuracy of statistical analysis depends on the completeness of rockfall inventories (Chau et al., 2003; Guzzetti et al., 2003;
109 D'amato et al., 2016). However, its application to rockfall hazards is limited due to the lack of complete inventory data (Budetta
110 and Nappi, 2013; Malamud et al., 2004). Empirical and semi-empirical rating systems are used where site-specific rockfall
111 inventories are either unavailable or unreliable. Therefore, rockfall susceptibility can be assessed by heuristic ranking of
112 selected predisposing factors (Frattoni et al., 2008; Budetta, 2004). Mechanical analysis based on static equilibrium theory is
113 the main method to analyse the stability of site-specific rockfall using the factor of safety (Fos). Ashby (1971) conducted
114 stability analysis with a parallelepiped block resting on an inclined plane (Fig. 2a), and the solution was subsequently modified
115 by Bray and Goodman (1981) and Sagaseta (1986). Kogure et al. (2006) utilized a cantilever beam model to determine the
116 critical state of limestone cliffs. Frayssines and Hantz (2009) proposed the limit equilibrium method (LEM) to predict block
117 stability against sliding and toppling in steep limestone cliffs (Fig. 2c). Chen and Tang (2010) established a stability analysis
118 method of three types of unstable rocks in the Three Gorges Reservoir area with the LEM. Alejano et al. (2015) studied the
119 influence of rounding of block corners on the block stability. Zhang et al. (2016) defined Fos based on fracture mechanics and
120 studied the progressive failure process by analysing crack propagation. Alejano et al. (2010) and Pérez-Rey et al. (2021)
121 deduced a formula for Fos of blocks with more complex geometry.

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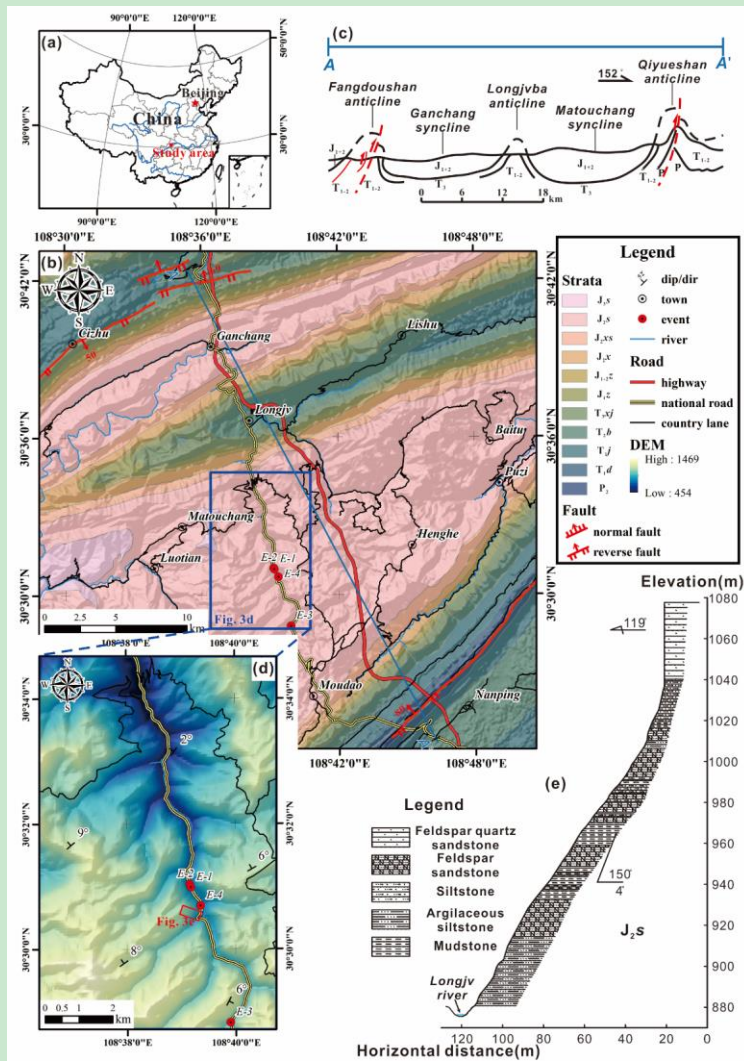
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 123 **Figure 2** Traditional force analysis diagrams of the rock block. (a) and (b) are stability analysis diagrams of rock blocks under dynamic
 124 conditions, resting on an inclined plane with a dip angle of α . The rock block is generalized as a cuboid with dimensions $b \times h$ and weight
 125 W (as modified from Ashby (1971), Bray and Goodman (1981) and Sagaseta (1986)). (c) Force description of the toppling model proposed
 126 by Frayssines and Hantz (2009). In the above assumptions, N , T , and W are regarded as forces applied at a point.
 127 The supporting force on the contact surface is assumed to be applied at a point in the current LEM methods (i.e., N in Fig. 2 b
 128 and c). However, the supporting force is actually a distributed force. The cavity generates an eccentricity effect on the
 129 overlying overhanging rock block and results in a non-uniform distribution of the supporting force on the contact surface, which
 130 is not considered in the traditional LEM. The presence of non-uniform stress distribution plays a critical role in inducing
 131 localized damage within a rock mass. Damage is frequently considered as an indicator or a threshold for the onset of accelerated
 132 failure in rock masses (Zhang et al., 2016). Therefore, it is imperative to consider the non-uniform stress distribution for the
 133 rockfall stability analysis. Furthermore, most studies simplified the three-dimensional geometry of the rock block by one cross-
 134 section, which is used to represent the critical features of the slope structure. Nevertheless, for natural blocks with basal cavities,
 135 the cavities usually present different depths along different directions (Pérez-Rey et al., 2021). Therefore, a three-dimensional
 136 model is necessary to calculate the accurate stability. In addition, when a block has multiple free faces and a complex structure,
 137 its potential failure is dominated by different modes, including rock mass damage and overall block failure. Therefore, the
 138 probable failure modes should be determined prior to the calculation of Fos .
 139 Based on rockfall investigation in the Eastern Sichuan Basin, China, the main objective of this study was to propose a new
 140 three-dimensional method for the determination of failure modes and Fos of biased rockfall, considering the non-uniform
 141 force distribution on the contact surfaces. Compared with the traditional LEM method, this study takes into account the partial
 142 damage of the underlying soft rock and the overall instability of the overlying overhanging hard rock blocks, and can evaluate
 143 the stability of biased rockfall more comprehensively. Fos of the typical unstable rock blocks in the study area are calculated
 144 to validate the proposed method. In addition, the critical cavity retreat ratio in this area is analysed. This study is an extension

145 of the basic LEM for rockfall, which can promote the accuracy of rockfall stability analysis and facilitate rockfall prevention
146 and risk mitigation.

147 **2 Study area**

148 **2.1 Geological setting**

149 The study area is located on the northeastern edge of the Sichuan Basin, China (Fig. 3a). Continuous erosion processes generate
150 moderate-low mountain and valley landforms (Yu et al., 2021). The tectonic structure of this area is characterized by a series
151 of ENE anticlines and synclines (Fig. 3b, c). In the anticline area, the rock layers dip relatively steeply, where translational
152 rockslides are the main mode of slope failure. The syncline area is dominated by gently dipping strata and is prone to rockfall
153 (Zhou et al., 2018). The study area is located in the core of the Matouchang syncline, where the rock layers are sub-horizontal
154 (Fig. 3d, e). In this valley, due to the longstanding fluvial incision, the relative relief is approximately 500 m and the valley
155 flanks are extremely steep (Fig. 3e). In addition, the toes of the hill slopes are reshaped because of the construction of the
156 G318 national road, which is the main traffic line and is always threatened by rockfalls dropping from steep rock slopes (shown
157 in Fig. 3d and Table 1).



158

159 **Figure 3** (a) Location of the study area in China; (b) geological map of the study area; (c) tectonic sketch profile of A-A', whose location is
 160 showed in Fig. 3b; (d) rockfall-prone segment and key investigation areas. The red dots are the positions of historical rockfall events,

161 corresponding to the numbers in Table 1; (e) Geological cross-section of the hillslope in the Jitougou section of G318 national road, which
 162 is marked by a red rectangle in Fig. 3d.

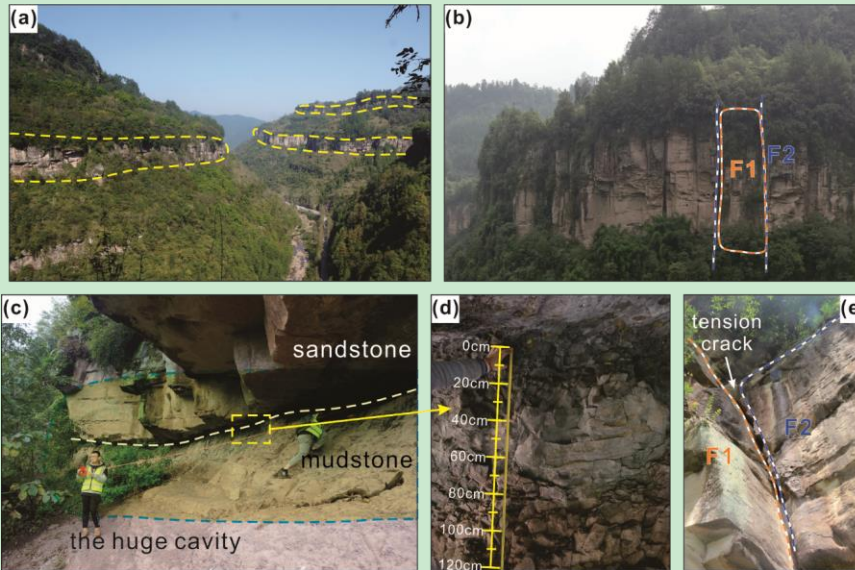
163 **Table 1** Historical rockfall events along G318 national road in the study area

No	Location	Time of occurrence (GMT+8)	Volume [m ³]	Consequence
E-1	K1698+900	2014-05 to 06*	Unknown	The power transmission facilities outside the road were smashed.
E-2	K1699+000	2015-02-14 23:00	About 240	A passing truck was stuck and two people dead.
E-3	K1690+700	2015-06-16	Unknown	The road was interrupted for a day.
E-4	K1698+400	2015-06-18 09:00	About 200	A vehicle was crashed into a gully and four people dead.

164 *Note: The exact time is unknown.

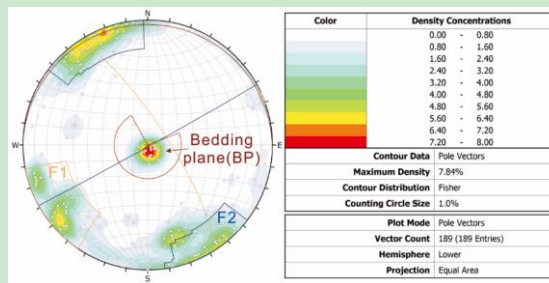
165 **2.2 Rockfall characteristics**

166 The slopes in the study area consist of a sub-horizontally interbedded sandstone and mudstone layer. Therefore, there are
 167 multiple layers of potentially unstable rock blocks in the hill slopes (Fig 4a). The thick sandstone has two sets of sub-vertical
 168 joints (Fig. 5), which cut the rock mass into blocks as the potential rockfall source (Fig. 4b). Cavities have formed in the
 169 underlying mudstone layer (Fig. 4c, d). Joints and bedding planes (BP) constitute the detachment surfaces between the blocks
 170 and steep slope (Fig. 4e). The eccentricity effect produced by the mudstone cavity plays an important role in the evolution
 171 process of rockfall. When the basal mudstone cannot provide adequate supporting force, the blocks detach from the steep slope,
 172 and biased rockfall occurs. Sliding and toppling are two possible failure modes of biased rockfall.



173

174 **Figure 4** Characteristics of biased rockfalls in the study area. (a) Multiple-layers of rockfall sources, which consist of thick sandstone. (b)
 175 Two sets of sub-vertical joints (F1 and F2) recognized by the UAV photos. (c) Large basal cavity developed in the underlying mudstone. (d)
 176 Dense fractures on the mudstone surface generated by weathering and compression. (e) Vertical tension crack in the rear of the block, through
 177 which precipitation can infiltrate.
 178 According to the historical rockfall events in this area, precipitation is considered a triggering effect of rock instability. The
 179 precipitation mainly infiltrates along the sub-vertical joints or cracks of the sandstone (Fig. 4e). However, the drainage of
 180 fissure water is hysteretic due to the obstruction of basal mudstone. Therefore, transient steady flow exists in vertical cracks
 181 during heavy rainfall, and the hydrostatic pressure triggers the detachment of rock blocks. Thus, typical scenarios (such as
 182 rainfall intensity and earthquake) need to be considered in the stability analysis model.



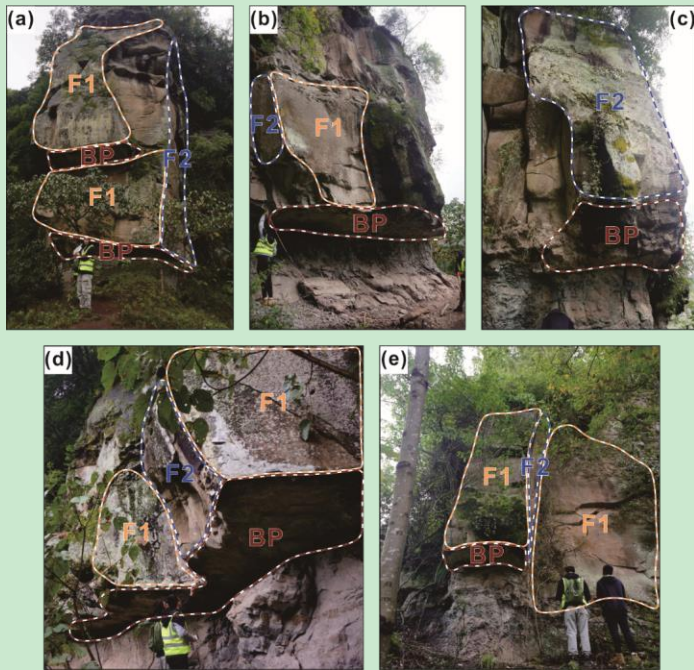
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 184 **Figure 5** Stereo net produced using compass-clinometer survey data, which shows the densities and orientations of five clusters. The data
 185 were collected in the rockfall-prone area shown in Fig. 3d.

186 3 Calculation method

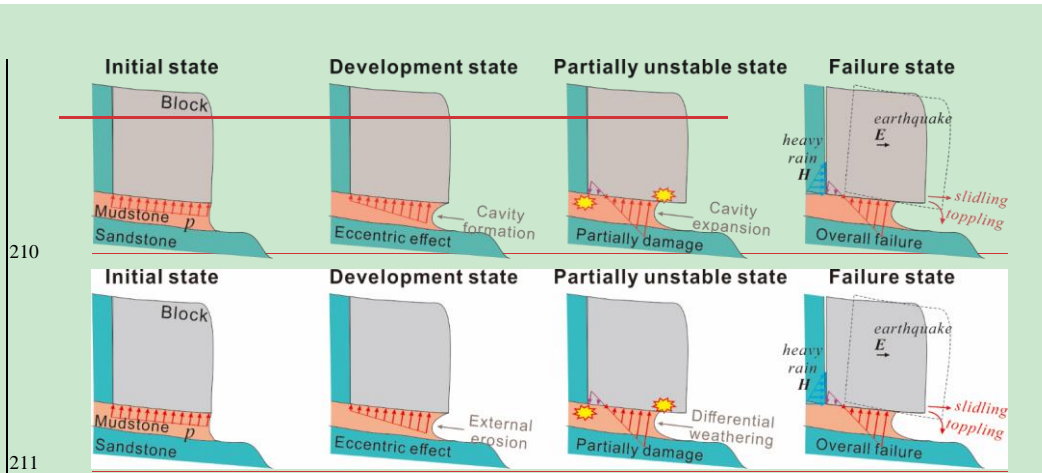
187 3.1 Geological models and assumptions

188 A detailed geological investigation of unstable rock blocks was carried out in the study area (Fig. 6). The geological model of
 189 the rock block is mainly composed of the **overlyingoverhanging** sandstone and the underlying mudstone. The sandstone block
 190 is assumed to be a rigid body, which is divided by two sets of orthogonal vertical smooth joints without friction resistance.
 191 According to the relatively persistent sub-vertical fractures observed in the field, the vertical joints are assumed to be fully
 192 persistent in the geological model. The sandstone block is assumed to be a complete body without persistent discontinuity, and
 193 it will not disintegrate before it falls. Due to the cavity in mudstone, the contact surface between sandstone and mudstone
 194 exhibits an eccentricity effect where non-uniform stresses are distributed at different positions. Mudstone is mainly loaded by
 195 compressive stress and tensile stress. When the compressive stress of mudstone exceeds its strength on the outer side, some
 196 initial damage appears. The effective contact surface between mudstone and sandstone is reduced, which aggravates the non-
 197 uniform distribution of stress. In this way, the ability of mudstone to resist the sliding and toppling of **overlyingoverhanging**
 198 sandstone is reduced. In the field, compression deformation of mudstone can be observed, which usually manifests as micro-

199 fractures and cleavages (Fig. 4d). The deformation is very slight and slow in the short term. In addition, the LEM is essentially
 200 a force/stress approach that does not take into account the deformation. Therefore, in this study, it is assumed that the mudstone
 201 is not subjected to deformation. The rock block remains in the state of static equilibrium prior to the final overall failure. Fig.
 202 7 displays the four evolution stages of biased rockfall. In the initial stage, the base cavity has not yet formed, and the normal
 203 force acting on the contact surface is uniform in different positions. The eccentricity effect leads to a non-uniform supporting
 204 force as the cavity grows, and partial damage gradually develops when the non-uniform stress exceeds the compressive or
 205 tensile strength of the mudstone. Under the triggering effects of rainfall or earthquakes, the rock blocks are separated by sliding
 206 or toppling.



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 208 **Figure 6** The unstable blocks were labelled W02, W08, W18, W04, and W21, which are detached by the dominating discontinuities in Fig.
 209 5. Basal cavities can be identified under the bedding planes of sandstone.



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Figure 7 The evolution process of rock blocks from stable state to failure.

Fig. 8 represents the mechanical model of the force equilibrium analysis of a rock block with two or three free faces. The rock block (the overlyingoverhanging sandstone) is generalized as a parallelepiped block. The underlying mudstone is impermeable, so rainfall can fill the joints and transmit horizontal hydrostatic pressure. The shear strength of the underlying mudstone is assumed to obey the Mohr–Coulomb criterion. Rainfall and earthquakes decrease Fos by generating hydrostatic pressure H in the vertical crack and horizontal seismic force E on the block.

A Cartesian coordinate system is established in three-dimensional space for the force analysis. The origin O is located at the centre of the contact surface between sandstone and mudstone. For the case with two free surfaces, the orientation of the free surfaces is set to be the positive direction of the x -axis and y -axis. For the case with three free surfaces, the negative direction of the x -axis is also a free surface. Joint J2 is perpendicular to the x -axis, and joint J1 is perpendicular to the y -axis.

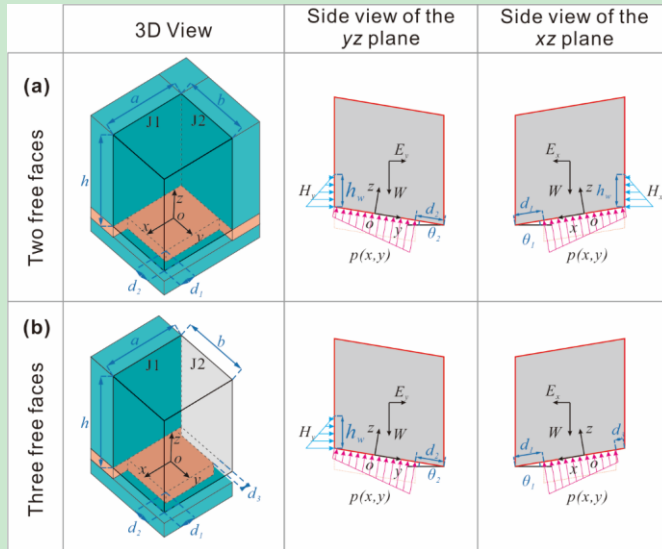


Figure 8 Diagram of the force equilibrium analysis of the rock block model. (a) and (b) represent the case of unstable rock blocks with two or three free vertical surfaces, respectively.

3.2 Calculation processes

3.2.1 Stress distribution at the block base

The following formulas are used to calculate the apparent dip of α (θ_1 and θ_2):

$$\theta_1 = \arctan(\tan \alpha \cdot \cos \omega_1) \quad (1)$$

$$\theta_2 = \arctan(\tan \alpha \cdot \cos \omega_2) \quad (2)$$

where ω_1 and ω_2 are the angles between the trend of the contact surface and the x direction or y direction, respectively.

As shown in Fig. 8b, with respect to the x -axis, gravity, seismic forces, and hydrostatic pressure create a non-symmetrical stress distribution on the foundation. The bending moment of gravity with respect to the x -axis (M_{bWx}) is

$$M_{bWx} = W \cdot \frac{d_1 - d_3}{2} \cos \theta_1 \quad (3)$$

Assuming that the height of the water in the fracture is h_w , the hydrostatic pressure along the x direction (H_x) and its bending moment (M_{bHx}) are respectively expressed as

$$H_x = \frac{\gamma_w h_w^2}{2} (b - d_2) \quad (4)$$

237
$$M_{bHx} = \int_{-\frac{b-d_2}{2}}^{\frac{b-d_2}{2}} \int_0^{h_w \cos \theta_1} \gamma_w \left(h_w - \frac{z}{\cos \theta_1} \right) \left(\frac{z}{\cos \theta_1} + \frac{a-d_1-d_3}{2} \cdot \sin \theta_1 \right) dz dy \quad (5)$$

238 The horizontal seismic force along x direction (E_x) and its bending moment (M_{bEx}) are respectively expressed as

239
$$E_x = k_e W \quad (6)$$

240
$$M_{bEx} = E_x \left(\frac{h}{2} - \frac{d_1-d_3}{2} \sin \theta_1 \right) \quad (7)$$

241 The total applied vertical load (N_z) and the total bending moments along the x direction (M_{bx}) can be derived as

242
$$N_z = W \cos \alpha - (H_x \cdot k_1 \cdot k_3 + E_x \cdot k_2) \sin \theta_1 - (H_y \cdot k_1 + E_y \cdot k_2) \sin \theta_1 \quad (8)$$

243
$$M_{bx} = M_{bWx} + M_{bHx} \cdot k_1 \cdot k_3 + M_{bEx} \cdot k_2 \quad (9)$$

244 where k_1 , k_2 and k_3 are the coefficients set to make Eq. (8) and Eq. (9) compatible with different calculation scenarios.

245 Therefore, Eqs. (8) and (9) and the following formulas can be expressed in a unified form. In the natural scenario, k_1 and k_2
246 are both equal to 0. In the rainfall scenario, $k_1 = 1$. In the earthquake scenario, $k_2 = 1$. For the case of two free faces, $k_3 = 1$.

247 For the case of three free surfaces, $k_3 = 0$.

248 Based on bending theory (Adrian, 2010), the eccentricity distance along the x direction (e_x) can be expressed as

249
$$e_x = \frac{M_{bx}}{N_z} = \frac{M_{bWx} + M_{bHx} \cdot k_1 \cdot k_3 + M_{bEx} \cdot k_2}{W \cos \alpha - (H_x \cdot k_1 \cdot k_3 + E_x \cdot k_2) \sin \theta_1 - (H_y \cdot k_1 + E_y \cdot k_2) \sin \theta_1} \quad (10)$$

250 The same method can be used to obtain e_y :

251
$$e_y = \frac{M_{by}}{N_z} = \frac{M_{bWy} + M_{bHy} \cdot k_1 + M_{bEy} \cdot k_2}{W \cos \alpha - (H_x \cdot k_1 \cdot k_3 + E_x \cdot k_2) \sin \theta_1 - (H_y \cdot k_1 + E_y \cdot k_2) \sin \theta_1} \quad (11)$$

252 According to the stress distribution of a rectangular shaped foundation (Adrian, 2010), the stress in the (x, y) coordinates,

253 $p(x, y)$, is

254
$$p(x, y) = \frac{N}{A} + \frac{N e_x}{I_y} x + \frac{N e_y}{I_x} y \quad (12)$$

255 with the formulas

256
$$I_x = \frac{(a-d_1)(b-d_2)^3}{12} \quad (13)$$

257
$$I_y = \frac{(b-d_2)(a-d_1)^3}{12} \quad (14)$$

258
$$A = (a-d_1-d_3)(b-d_2) \quad (15)$$

259 By substituting Eq. (13-15) into Eq. (12), $p(x, y)$ can be derived as

260
$$p(x, y) = \frac{N}{A} \left[1 + \frac{12 e_x}{(a-d_1-d_3)^2} x + \frac{12 e_y}{(b-d_2)^2} y \right] \quad x \in \left[-\frac{a-d_1-d_3}{2}, \frac{a-d_1-d_3}{2} \right], y \in \left[-\frac{b-d_2}{2}, \frac{b-d_2}{2} \right] \quad (16)$$

261 p_{max} and p_{min} can be derived from Eq. (16) as

$$p_{max} = p\left(\frac{a-d_1-d_3}{2}, \frac{b-d_2}{2}\right) \quad (17)$$

$$p_{min} = p\left(-\frac{a-d_1-d_3}{2}, -\frac{b-d_2}{2}\right) \quad (18)$$

The mudstone foundation has both compressive strength and tensile strength, so the value of $p(x, y)$ is modified to obtain the two piecewise functions

$$p_p(x, y) = \begin{cases} \sigma_{cmax}, & p(x, y) \geq \sigma_{cmax} \\ p(x, y), & 0 < p(x, y) \leq \sigma_{cmax} \\ 0, & p(x, y) < 0 \end{cases} \quad (19)$$

$$p_n(x, y) = \begin{cases} 0, & p(x, y) < -\sigma_{tmax} \\ p(x, y), & -\sigma_{tmax} \leq p(x, y) < 0 \\ 0, & p(x, y) \geq 0 \end{cases} \quad (20)$$

Here, $p_p(x, y)$ provides support normal force for the overlyingoverhanging sandstone, and $p_n(x, y)$ provides tension force.

3.2.2 Calculation of factors of safety

According to the Mohr-Coulomb criterion-principle-of-friction, the ultimate shear strength τ_{max} is

$$\tau_{max} = \int_{\frac{a-d_1-d_3}{2}}^{\frac{a-d_1-d_3}{2}} \int_{\frac{b-d_2}{2}}^{\frac{b-d_2}{2}} [p_p(x, y) \tan \varphi + c] dy dx \quad (21)$$

Therefore, Fos against sliding, Fos_{sl} , can be defined as

$$Fos_{sl} = \frac{S_{stabilizing}}{S_{sliding}} = \frac{\tau_{max}}{W|\sin \alpha_s| + H_x \cdot \cos \omega_s \cdot \cos \alpha_s \cdot k_1 \cdot k_3 + H_y \cdot |\sin \omega_s| \cdot \cos \alpha_s \cdot k_1 + E \cdot \cos \alpha_s \cdot k_2} \quad (22)$$

When the block can slide freely, $\alpha_s = \alpha$, $\omega_s = 0$; when the block is constrained to slide along a joint plane (e.g., J1), $\alpha_s = \theta_1$ or θ_2 , $\omega_s = \omega_1$ or ω_2 . For the case of an anaclinal slope, the sliding direction is opposite to the free surface. Therefore, the rock block does not slide, and Fos_{sl} is not considered in the model.

With regard to stability against toppling, along the x direction, the part of the block above the mudstone base provides the stabilizing moment $M_{W_{inx}}$, and the part of the block above the cavity provides the overturning moment $M_{W_{outx}}$. When tension exists, there is an additional stabilizing moment. M_{px} , $M_{W_{inx}}$, $M_{W_{outx}}$ and M_{px} can be derived as

$$M_{W_{inx}} = W \frac{a-d_1}{a} \cos \theta_1 \cdot \left(\frac{a-d_1}{2}\right) \quad (23)$$

$$M_{W_{outx}} = W \frac{d_1}{a} \cos \theta_1 \cdot \frac{d_1}{2} \quad (24)$$

$$M_{px} = - \int_{\frac{b-d_2}{2}}^{\frac{b-d_2}{2}} \int_{\frac{a-d_1-d_3}{2}}^{\frac{a-d_1-d_3}{2}} p_n(x, y) \cdot \left(\frac{a}{2} - d_1 - x\right) dx dy \quad (25)$$

and M_{Hx} and M_{Ex} can be derived as

$$M_{Hx} = \int_{\frac{b-d_2}{2}}^{\frac{b-d_1}{2}} \int_0^{h_w \cos \theta_1} \gamma_w \left(h_w - \frac{z}{\cos \theta_1} \right) \left(\frac{z}{\cos \theta_1} + (a - d_1) \sin \theta_1 \right) dz dy \quad (26)$$

$$M_{Ex} = E_x \left(\frac{h}{2} + \left(\frac{a}{2} - d_1 \right) \sin \theta_1 \right) \quad (27)$$

Therefore, the *Fos* against toppling along the *x* direction, Fos_{tox} , results in

$$Fos_{tox} = \frac{M_{stabilizing}}{M_{overturning}} = \frac{M_{W_{inx}} + M_{px}}{M_{W_{outx}} + M_{Hx} \cdot k_1 \cdot k_3 + M_{Ex} \cdot k_2} \quad (28)$$

Similarly, Fos_{toy} can be obtained as

$$Fos_{toy} = \frac{M_{stabilizing}}{M_{overturning}} = \frac{M_{W_{iny}} + M_{py}}{M_{W_{outy}} + M_{Hy} \cdot k_1 + M_{Ey} \cdot k_2} \quad (29)$$

The smaller value is selected as the *Fos* of the toppling failure mode Fos_{to} :

$$Fos_{to} = \min(Fos_{tox}, Fos_{toy}) \quad (30)$$

When the stress on mudstone exceeds its strength, it causes partial damage and decreases the stability of the rock block.

Therefore, *Fos* with the consideration of compressive strength (Fos_{co}) and tensional strength (Fos_{te}) can be derived as

$$Fos_{co} = \frac{\sigma_{cmax}}{p_{max}} \quad (31)$$

$$Fos_{te} = \frac{\sigma_{tmax}}{-p_{min}} \quad (32)$$

Fos_{co} and Fos_{te} represent the current damage degree of mudstone due to compressive stress and tensile stress, respectively.

When the stress exceeds the ultimate strength, the strength of the mudstone is reduced to the residual value, and the initial

deformation appears. The ability of mudstone to provide resistance to the sliding and toppling of sandstone blocks is thus

reduced, and Fos_{sl} and Fos_{to} subsequently decline. The smaller the value of Fos_{co} and Fos_{te} , the greater the damage to the

underlying mudstone. The effective contact area between sandstone and mudstone becomes smaller as the development of

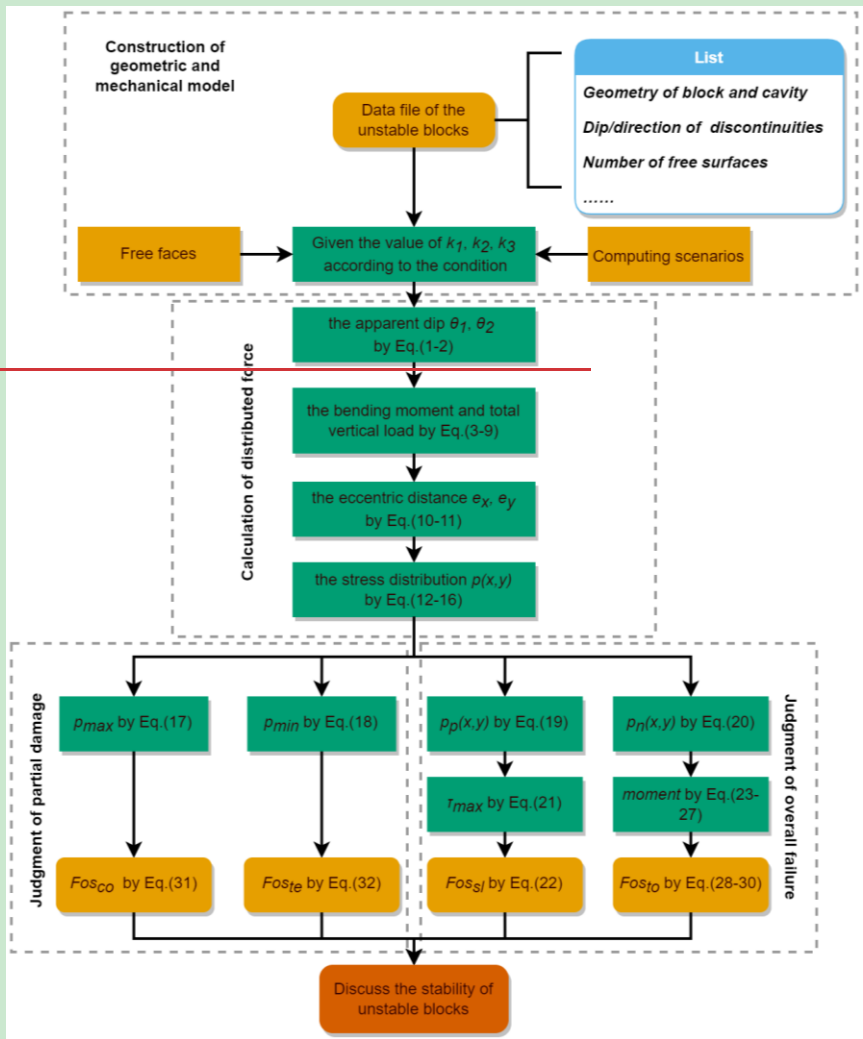
compressive and tension damage, which significantly affects the stability of the overlyingoverhanging sandstone block.

Finally, four *Fos* of unstable rock block are obtained. Fos_{sl} and Fos_{to} are routine indicators directly representing the stability

of sandstone blocks. Fos_{co} and Fos_{te} are two indicators proposed in this study for the stability analysis of biased rockfall,

which describe the damage state of the underlying mudstone base. It is necessary to simultaneously consider four *Fos* to

evaluate the stability of unstable biased rockfall. The entire calculation process is shown in Fig. 9.



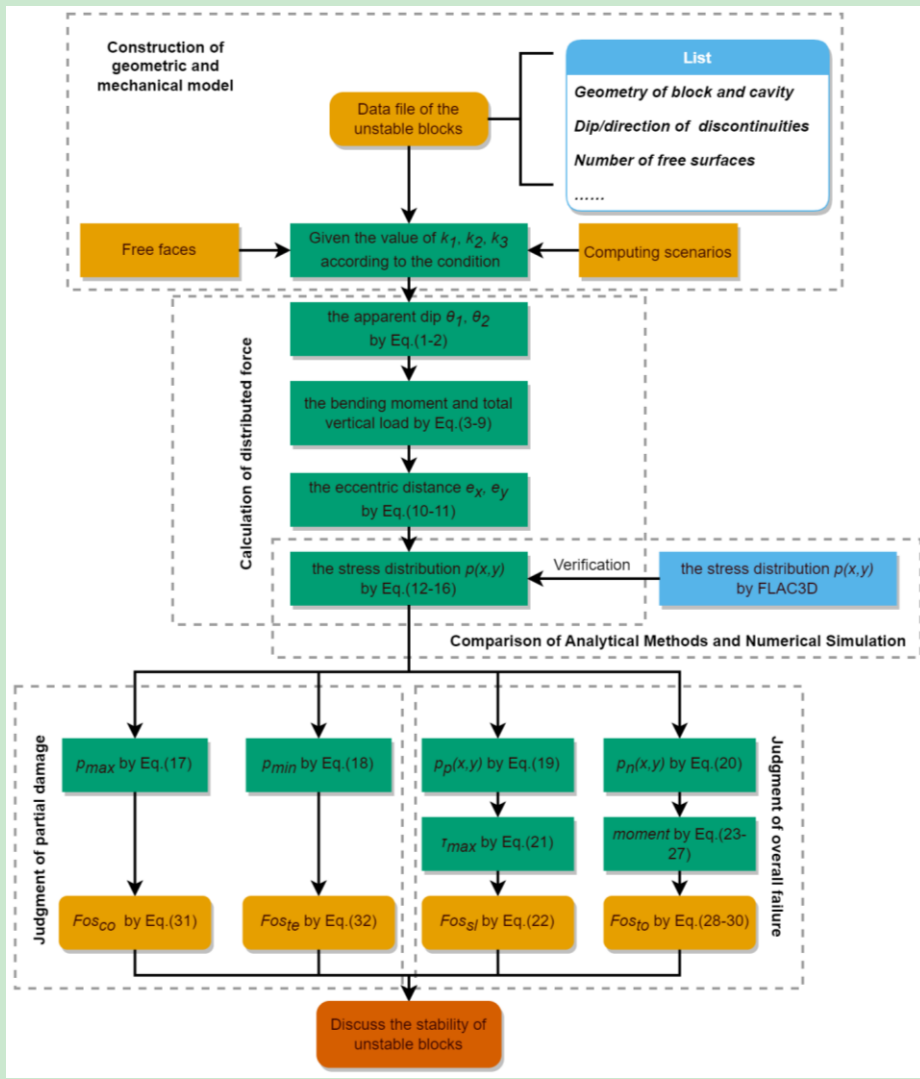


Figure 9 Calculation process of FOS of the unstable rock blocks.

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4 Validation of analytical methods by numerical simulation

The damage mechanisms at the base of the rock block play an important role in the rockfall evolution process. However, the stress distribution on the contact surface calculated by the proposed analytical methods is difficult to be validated by the field data. Therefore, numerical simulation of a biased rockfall was conducted in this study to determine the stress distribution on the contact surface between overhanging sandstone and underlying mudstone. By comparing the results of the proposed analytical methods with those obtained from the numerical simulation, the reliability of the analytical methods can be validated. FLAC3D, a professional software that utilizes the finite difference method (FDM) for three-dimensional analysis of rocks, soils, and other materials, was employed for the 3D numerical simulation. Based on the geological models, a 3D numerical simulation model was conducted with FLAC3D 6.00 to analyse the stress distribution on the contact surface (Fig. 10).

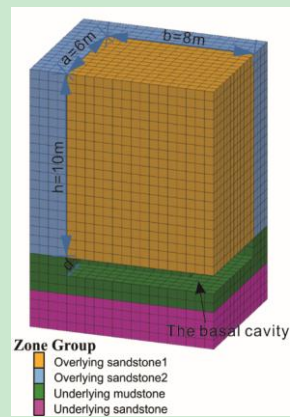
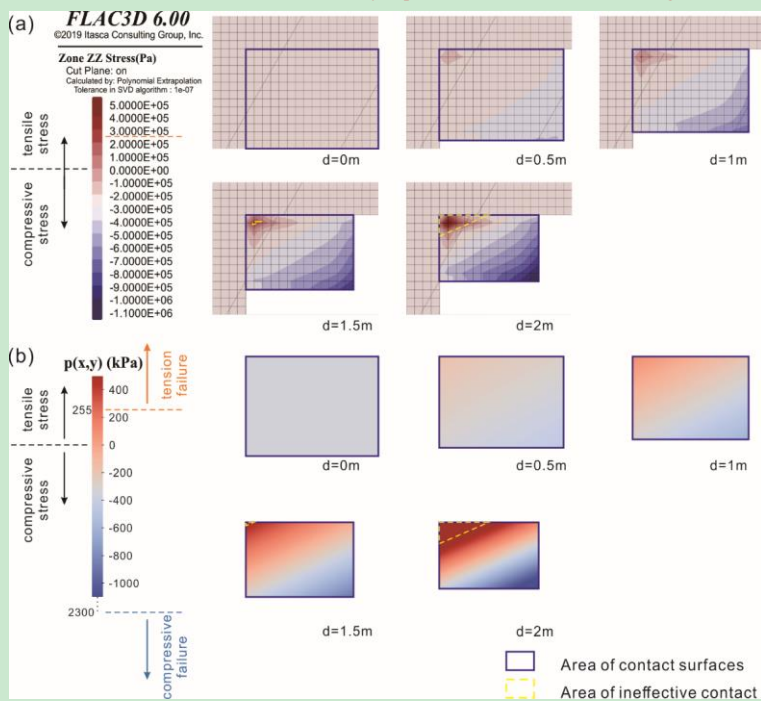


Figure 10 Numerical model built in FLAC3D

The model is mainly composed of sandstone and mudstone, which the Overhanging sandstone1 represents a unstable rock block (dimensions $a \times b \times h$ are 6m, 8m, 10m respectively), and the weathering process of the cavity is represented by excavating in stages in the underlying mudstone. Sandstone was considered as elastic model, and mudstone was assigned Mohr-Coulomb model. Material properties were determined by referring to published literature and investigation reports in the study area. The unit weight of the sandstone block (γ_s) is 25 kN/m³ (Tang et al., 2010), and the mudstone is 22.54 kN/m³. The friction angle of the contact surface (ϕ) is set to 25° and the cohesion (c) is set to 70 kPa (Zhang et al., 2016). Because of the strength degradation of mudstone foundations due to intense weathering, the maximum compressive stress of mudstone (σ_{cmax}) is replaced by the bearing capacity of mudstone foundations (2300 kPa), which is obtained through plate load tests in adjacent areas (Zheng et al., 2021). In addition, the maximum tensile stress of mudstone (σ_{tmax}) is valued as one-ninth of σ_{cmax} . The west, north and bottom boundaries of the model are constrained by roller boundary conditions. The cohesion and internal friction angle of the interface between Overhanging sandstone1 and Overhanging sandstone2 are set to 0. After reaching the

331 initial force-equilibrium state, the mudstone was excavated to simulate the weathering process, and the vertical stress
 332 distribution on the sand-mudstone interface at different cavity depths was obtained, as shown in Figure 11.



333
 334 Fig.11 Diagram of stress distribution in the vertical direction on the contact interface through different methods, (a) the
 335 results of numerical simulation by FLAC3D, (b) the results of of proposed analytical method.
 336 When there is no cavity present, represented by d=0m, the stress distribution is uniform compressive stress (According to the
 337 FLAC3D software, compressive stresses are negative). At d=0.5m, the stress remains entirely compressive, but non-uniform
 338 stress distribution occurs on the contact surfaces. At d=1m, the vertical stress value in the upper left corner of the contact
 339 interface surpasses 0 (Fig.11), indicating the presence of tensile stress. As d increases to 1.5m or 2m, the tensile stress in the
 340 upper left corner gradually intensifies, exacerbating the non-uniform stress distribution. The results obtained from the
 341 numerical simulation align with those from the analytical method, confirming the existence of tensile stress at the contact
 342 interface in the biased rockfall due to external erosion development (Fig.11). Tensile stress commonly emerges within the
 343 contact surface, making it challenging to observe directly in the field.

In the context of the limit equilibrium method, the contact area plays a vital role in stability analysis, as shown in Eq. (21)–(30) in Section 3. The numerical simulation process provides an intuitive understanding of the influence of non-uniform stress distribution on the contact surfaces on the stability of rock blocks. Whether subjected to tension or compression, the rock layer has an ultimate strength. In Fig.11, when $d=1.5\text{m}$ or 2m , the tensile stress exceeds the ultimate tensile strength, leading to tensile failure in the upper left corner of the stress distribution diagram. The region enclosed by a yellow dotted line represents ineffective contact, where no anti-slip force or overturning moment can be generated due to tension failure at the contact surface. Therefore, this area needs to be subtracted from the total contact area when calculating $[[Fos]]_{sl}$ and $[[Fos]]_{to}$. Similar situations occur when the compressive stress exceeds the ultimate compressive strength. The current maximum compressive stress has not reached the ultimate compressive strength in Figure 11. However, As d continues to increase, the area of compression failure will appear in the lower right corner of diagram in Figure 11. This occurrence diminishes the area capable of providing anti-slip force or overturning moment, thereby reducing the stability of the rock blocks. The traditional LEM method does not account for distributed forces and fails to consider changes in the contact surface. The method proposed in this study addresses this issue and is applied to the calculation of the $[[Fos]]_{sl}$ and $[[Fos]]_{to}$ as presented in Eq. (21), (25) and (26)).

4 Parameters and 5 results Results

A detailed field investigation was carried out in the source area of rockfall (Fig. 3d). The size of the blocks was determined by on-site measurement with tape and a laser rangefinder. The basal cavities in mudstone were measured with a steel ruler, and the morphological characteristics of mudstone foundation were mainly described with the average erosion depth of the cavity. The attitude of discontinuities was measured by compass. The mechanical parameters for the Fos calculation of rock blocks were determined by referring to published literature and investigation reports in this area. The unit weight of the sandstone block (γ_s) is 25 kN/m^3 (Tang et al., 2010), the friction angle of the contact surface (φ) is set to 25° and the cohesion (c) is set to 70 kPa (Zhang et al., 2016). Because of the strength degradation of mudstone foundations due to intense weathering, the maximum compressive stress of mudstone (σ_{emax}) is replaced by the bearing capacity of mudstone foundations (2300 kPa), which is obtained through plate load tests in adjacent areas (Zheng et al., 2021). In addition, the maximum tensile stress of mudstone (σ_{tmax}) is value as one-ninth of σ_{emax} . The mechanical parameters have been given in Section.4. The height of the water level (h_w) is set to be one-third of h , and an earthquake contribution coefficient k_e of 0.05 is considered in stability calculations. The data obtained from the field survey were organized according to the coordinate system of the geological model in Section 3.1, and Fos was calculated according to the calculation steps in Section 3.2. The calculated geometric parameters and Fos results are shown in Table 2.

Table 2 Geometric parameters of rock blocks in the study area and Fos results.

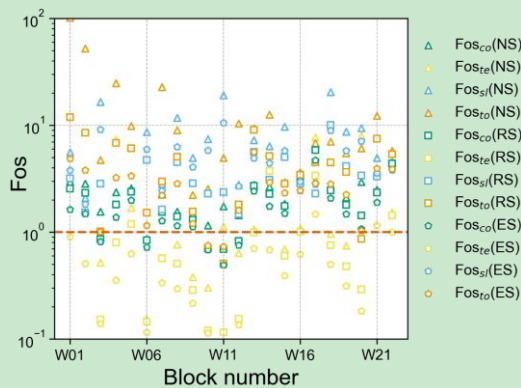
Block number	Free faces	h [m]	a [m]	b [m]	d_1 [m]	d_2 [m]	d_3 [m]	α [°]	Dip direction [°]			NS (Natural scenario)						RS (Rainfall scenario)						ES (Earthquake scenario)					
									BD	J1	J2	Fos_{te}	Fos_{co}	Fos_{sl}	Fos_{to}	Fos_{min}	Fos_{te}	Fos_{co}	Fos_{sl}	Fos_{to}	Fos_{te}	Fos_{co}	Fos_{sl}	Fos_{to}	Fos_{te}	Fos_{co}	Fos_{sl}	Fos_{to}	
W01	3	23	7.2	6.1	0.65	0.25	0.17	6	78	7	97	-	2.99	5.61	101.54	2.99	-	2.56	3.18	11.91	0.90	1.63	3.81	4.88					
W02	3	23	6.42	5.25	0.78	0.4	0.31	16	148	51	141	-	2.84	2.10	52.28	2.10	-	2.33	1.54	8.49	0.51	1.48	1.82	3.79					
W03	2	20	3.5	2.6	0.84	0.55	-	7	341	53	143	0.52	1.56	16.53	4.72	0.52	0.15	0.86	2.83	1.02	0.14	0.81	9.12	1.01					
W04	2	19	4.6	4.6	0.62	0.77	-	7	273	65	155	7.35	2.37	-	24.74	2.37	0.80	1.81	-	6.83	0.35	1.38	-	3.23					
W05	2	15	16.7	5.6	2.13	1.36	-	5	283	50	140	1.70	2.57	-	9.86	1.70	1.19	2.39	-	6.10	0.63	1.99	-	3.36					
W06	3	20	16.7	9.7	7.5	4.2	3.9	5	302	226	316	0.15	0.87	8.67	1.53	0.15	0.15	0.84	4.73	1.52	0.12	0.72	5.96	1.16					
W07	2	22	9.2	3.7	0.64	0.8	-	12	324	315	405	-	2.27	2.82	22.86	2.27	0.57	1.55	1.62	2.97	0.34	1.28	2.44	2.21					
W08	2	23	12	7.9	2	1.9	-	3	317	332	422	0.76	1.55	11.75	8.99	0.76	0.51	1.40	4.51	5.09	0.29	1.14	6.29	2.84					
W09	2	18	8.4	6	0.9	2.5	-	8	60	335	425	0.38	1.48	4.98	2.23	0.38	0.29	1.30	2.87	1.56	0.22	1.12	4.08	1.20					
W10	2	23	5.7	3.3	1.3	0.85	-	5	329	313	403	0.30	1.16	7.41	2.53	0.30	0.12	0.71	2.30	0.71	0.11	0.68	5.84	0.75					
W11	3	22	1.1	2	0.1	0.64	0.1	4	327	120	210	1.13	1.74	19.08	4.97	1.13	0.12	0.69	2.37	0.51	0.07	0.49	10.57	0.73					
W12	2	25	3.9	4	0.74	0.96	-	12	355	297	387	0.64	1.44	2.78	10.36	0.64	0.15	0.82	1.48	1.81	0.14	0.75	2.70	1.61					
W13	2	12	11.9	10.9	3	2.28	-	7	36	73	163	1.06	2.77	7.28	9.39	1.06	0.99	2.71	5.63	9.02	0.70	2.41	4.93	5.65					
W14	3	19	13	5	0	1.1	0	8	296	73	163	-	2.67	6.40	12.57	2.67	3.75	2.28	3.09	5.15	0.68	1.75	4.41	2.94					
W15	2	18	22	6	8.3	0	-	8	351	200	290	0.70	1.84	9.74	2.93	0.70	0.60	1.75	5.03	2.83	0.39	1.50	5.79	2.34					
W16	3	11	5.2	7.6	0	2.9	0	13	42	144	234	1.09	3.04	3.46	3.65	1.09	1.01	2.96	2.84	3.45	0.62	2.45	2.98	2.45					
W17	3	7	8	2	0	0.56	0	20	30	156	246	7.71	6.72	3.07	6.83	3.07	3.40	5.87	2.29	4.49	1.48	4.70	2.81	2.86					
W18	2	12	8.5	4.5	1.61	1.27	-	2	252	253	343	0.97	2.66	20.49	7.05	0.97	0.75	2.46	10.06	4.50	0.50	2.08	8.90	2.82					
W19	2	15	4.2	5.2	1.6	0.68	-	5	28	56	146	0.75	2.12	8.71	5.49	0.75	0.48	1.80	4.17	3.66	0.31	1.48	5.79	2.24					
W20	3	15	1.8	1.7	0.23	0.5	0.3	4	20	63	153	7.96	2.95	9.44	6.08	2.95	0.29	1.43	3.39	0.87	0.18	1.07	7.12	1.03					
W21	3	20	18.9	9	0	2	0	7	348	71	161	-	2.51	4.96	12.25	2.51	-	2.36	3.31	7.48	1.15	1.90	3.58	3.95					
W22	2	7	5.4	5.7	1	1.65	-	6	294	53	143	1.53	4.48	-	5.78	1.53	1.44	4.38	-	5.37	1.00	3.81	-	3.88					

Note: When there is no tensile stress in the mudstone foundation, Fos_{te} has no value. For the case of an anacinal slope, blocks do not slide and Fos_{sl} has no value. Both parameters are replaced by "-".

374 5.6 Discussion

375 5.16.1 Characteristics of rock block stability

376 There are up to 12 results of Fos per potential unstable block with the consideration of three scenarios and four failure modes
377 (i.e., partial damage and overall failure). Most Fos_{te} values are less than 1 in all scenarios (yellow points in Fig. 4012), except
378 for two blocks (i.e., W17 and W20), whose Fos_{te} values are also close to 1 under rainfall or earthquake scenarios. Although
379 most of Fos_{co} values (green points in Fig. 4012) are greater than 1, they are closer to the critical state of $Fos = 1$ than Fos_{sl}
380 and Fos_{to} (represented by blue and orange points in Fig. 4012, respectively). The compression damage of the exposed
381 mudstone can be investigated in the field survey (Fig. 4d). However, it is difficult to observe the phenomenon of tensile damage
382 inside the mudstone base. In the case of weak tensile strength, the mudstone base suffers from tensile failure, and compression
383 failure usually occurs before tension failure. According to the results, their Fos_{te} and Fos_{co} are less than 1 or close to 1, which
384 means that the underlying mudstone has been partially damaged due to slight compressive or tensile failure, and the blocks are
385 potentially unstable with the current depth of the basal cavity. However, most of the blocks do not exhibit overall failure, and
386 they still exist on the slope. Moreover, their Fos_{sl} and Fos_{to} values are greater than 1 in different scenarios, which is consistent
387 with this actuality. The results indicate that most of the blocks are close to a critical state, in which they are partially damaged
388 but the whole block is still stable.

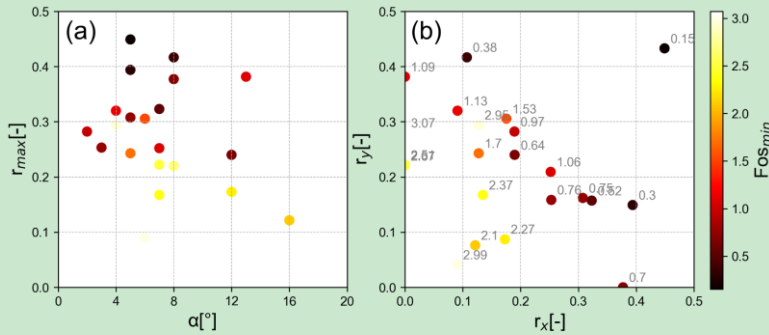


389
390 **Figure 40-12** Distribution of Fos in different scenarios. Shapes represent different scenarios and colours represent different failure modes.

391 5.26.2 Relationship between Fos and geometric parameters

392 Fig. 41-13 presents the relationship between Fos_{min} and two main geometric parameters, the dip of the contact surface and
393 the retreat ratio. In general, the dip angle of the contact surface (α) is the key factor influencing the sliding failure mode. The

394 horizontal axis in Fig. 4a-13a is α between the rock blocks and underlying mudstone. Most of the points in Fig. 4a-13a are
 395 in the interval $[0, 8^\circ]$, which is consistent with the features of sub-horizontal strata in the study area. The shade of the points
 396 does not change significantly in the x -axis direction, as Fig. 4a-13a shows. Therefore, compared with the maximum retreat
 397 ratio (r_{max}), the dip of the contact surface has less influence on rockfall stability in the study area. There was a significant
 398 positive correlation between the retreat ratio (r_{max}) and Fos_{min} . In Fig. 4b-13b, as the retreat ratios increase in the positive
 399 direction of the x -axis and y -axis, the rock blocks show a notable tendency to be unstable.



400
 401 **Figure 4-13** Correlation between Fos and the dip of contact surface and retreat ratio. Here, α is the dip angle of the contact surface between
 402 rock block and underlying mudstone, r_x and r_y are the retreat ratio along x direction and y direction, respectively, equal to d_1/a and d_2/b ,
 403 and r_{max} is the larger of r_x and r_y .

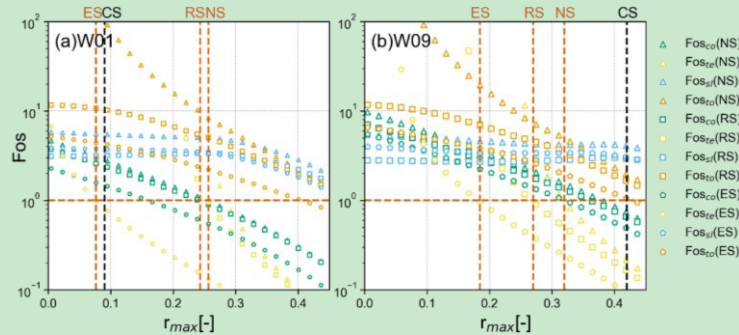
404 5.3.6.3 Definition of rockfall susceptibility

405 To explore the variation in Fos with the progressive erosion process of the cavity on the blocks, the cavity retreat velocities in
 406 different directions are assumed to be equal (5 mm/year, Zhang et al. (2016)). Fig. 42-14 shows the variations in Fos of two
 407 specific blocks during the evolution process of the mudstone cavity. In the initial stage, the cavity is small, and the
 408 overlyingoverhanging block is stable; all Fos values are greater than 1.0. The cavity expands over time as the mudstone
 409 weathers; then, the contact area decreases, and non-uniform distributed stress arises. When the stress exceeds the ultimate
 410 strength of mudstone in a partial area, Fos_{co} and Fos_{te} decrease significantly, as shown in Fig. 42-14. The instability of the
 411 blocks starts from the failure (or damage) of the foundation. Fos_{te} and Fos_{co} reach the critical state much earlier than Fos_{sl}
 412 and Fos_{to} . For these two specific blocks, when r_{max} increases to 0.4, Fos_{sl} and Fos_{to} are still higher than 1.0. This means
 413 that the rock blocks can remain globally stable in this condition.

414 These results further elucidate the stability analysis model proposed in this study. Fos_{co} and Fos_{te} introduced in this model
 415 present the damage state of basal mudstone caused by compressive and tensile stresses, which do not provide global instability
 416 of the overlyingoverhanging block as sliding and toppling. However, Fos_{co} and Fos_{te} are important preliminary signs of

417 subsequent global failure of the rock block, as presented through the numerical simulation in Section 4. The damage in the
 418 basal mudstone can significantly accelerate weathering and prompt expansion of the cavity, which will lead to global failure.
 419 The lower Fos_{co} and Fos_{te} are, the lesser the safety margin of the blocks. Therefore, the four Fos used in this study can
 420 provide a more comprehensive quantification of rockfall stability.

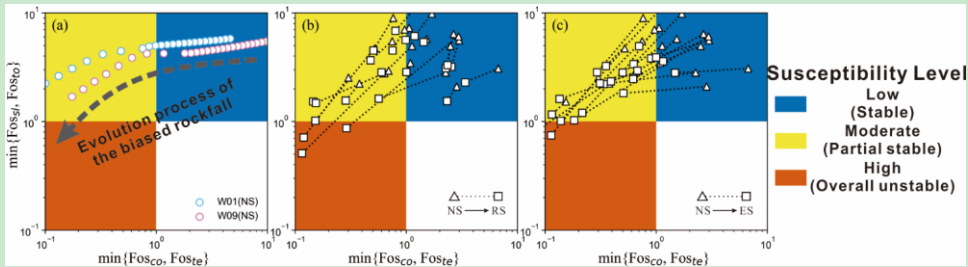
421 This result is consistent with Fig. 4912, in which 63.7% of the yellow and green points (Fos_{te} and Fos_{co}) are located between
 422 $Fos = 0.7$ and $Fos = 2.0$. This result can be validated by the field phenomena. In the study area, rock damage (e.g., micro-
 423 fractures and cleavages) can be observed in the underlying mudstone. However, most overlyingoverhanging rock blocks are
 424 stable at the present time. This means that even if Fos_{sl} or Fos_{to} is higher than 1, its foundation has begun to be damaged. In
 425 the case of heavy rain or earthquakes, Fos_{sl} and Fos_{to} may be reduced to less than 1, and the rockfall occurs.



426
 427 **Figure 12-14** Variation in Fos with r_{max} . (a) and (b) are the results for W01 and W09, respectively, which represent the situation of the
 428 blocks with two and three free faces. The black dotted line (CS) approximately represents the current state of the unstable blocks. The red
 429 dotted lines correspond to the critical values of r in different scenarios.

430 Based on the meaning of four Fos , rockfall susceptibility can be divided into three levels. When both Fos_{co} and Fos_{te} are
 431 greater than 1, the overall rock block is stable, and the mudstone base is not damaged, which is defined as “low susceptibility”
 432 and represented by the blue area in Fig. 4315. With the development of cavity erosion, when Fos_{co} or Fos_{te} is less than 1 and
 433 Fos_{sl} and Fos_{to} are higher than 1, the base undergoes be damaged, and the overlyingoverhanging sandstone blocks remain
 434 relatively stable. This state is defined as “moderate susceptibility” and represented by the yellow area. When Fos_{sl} or Fos_{to}
 435 is less than 1 in some scenarios, the rock blocks are in a “high susceptibility” state, which means that rockfalls are highly likely
 436 to occur. Fig. 43a-15a indicates that along with the increase in the cavity retreat ratio, the susceptibility of W01 and W09
 437 changes from low susceptibility to moderate susceptibility in the natural scenario. As Fig. 43b-15b and c show, when rainfall
 438 or earthquake occurs, Fos_{sl} or Fos_{to} of some blocks is less than 1, which means that some blocks have evolved to the state of
 439 high susceptibility and the overall sandstone blocks are unstable.

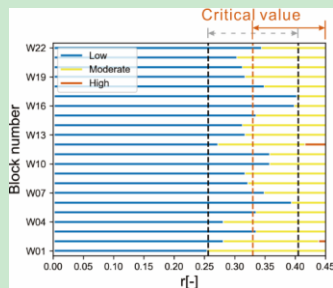
440



441
 442 **Figure 13-15** Rockfall susceptibility based on the combination of four Fos s. The susceptibility is defined as three levels, represented by red,
 443 yellow and blue. (a) shows the progressive failure process of the rock block changing from low susceptibility to moderate susceptibility as
 444 the cavity retreat ratio increases (illustrated by W01 and W09 in the natural scenario). (b) and (c) show the change in susceptibility of biased
 445 rock blocks, when the scenario changes from natural conditions to rainfall and earthquake conditions.

446 **5.46.4 Critical retreat ratio in the study area**

447 The cavity plays an important role in the progressive failure process of biased rockfall. To analyse the effect of the retreat ratio
 448 on the stability of rock blocks, all blocks in the study area were selected to calculate their Fos s and susceptibility level with the
 449 increasing r , whose retreat velocities in different directions are assumed to be equal. Fig. 14-16 shows that along with the
 450 increase in the retreat ratio, the susceptibility level of rock blocks changes from low to moderate susceptibility. Corresponding
 451 to the critical state of $\min\{Fos_{co}, Fos_{te}\} = 1$ of all blocks, the minimum retreat ratio is 0.26, and the maximum retreat ratio
 452 is 0.41, as marked by the vertical black dotted line in Fig. 1416. According to the statistical analysis of critical retreat ratios,
 453 both mean and median are 0.33. Therefore, the critical retreat ratio of the rock blocks in the study area can be determined as
 454 0.33, which is marked by the vertical red dotted line in the Fig. 1416. The critical retreat ratio calculated by this method can
 455 be used for the preliminary identification of potential unstable rock blocks in a specific area, which can help concentrate limited
 456 risk treatment resources on these priorities. It should be emphasized that the mechanical parameters and analysis scenarios
 457 significantly affect the critical value. Therefore, the elaborative risk control of a given rockfall should be arranged based on its
 458 specific parameters and analysis scenarios.



459

460 **Figure 14-16** Effect of the retreat ratio (r) on the Fos of the rock block, which is illustrated by all blocks in the study area.

461 **5.5.6.5** Limitations

462 This study involves the development of an analytical model for the three-dimensional stability of biased rockfall, combining
 463 the basic LEM method and the consideration of the eccentricity effect. Due to the complexity of rock structure and force
 464 analysis, it is necessary to highlight the limitations of this model.

465 First, this study uses a three-dimensional coordinate system and bending theory. It is difficult to consider diverse shapes of
 466 rock blocks, and the rock block was simplified as a prismatic column. The assumption of fully persistent discontinuities may
 467 underestimate the stability of rock blocks, and ignores the stress transmission in joints or rock bridges. Then, following the
 468 basic framework of the general LEM method, this study assumed that the rock is not subjected to deformations. The complete
 469 stress-strain behaviour, such as the deformation in the mudstone layer, is not considered in this study. The mode of tension
 470 failure is very difficult to observe in the field, and it is currently verified by means of numerical simulation. Furthermore, the
 471 block stability is strongly influenced by the uncertainty of mechanical parameters. However, because of the difficulties in
 472 sampling strongly weathered mudstone, it is difficult to obtain adequate parameter values for uncertainty statistics. These
 473 limitations will be important considerations in future studies.

474 **6-7** Conclusion

475 Due to differential weathering in sub-horizontally interbedded of hard rock and soft rock, multi-layer biased rockfalls develop
 476 on steep slopes. In mountainous ranges, cut slopes, and coastal cliffs, rockfall may cause significant facility damage and
 477 casualties in residential areas and transport corridors. The aim of this study was to present a new three-dimensional analytical
 478 method for the stability of rock blocks with basal cavities. In this method, a non-uniform distributed stress due to the
 479 eccentricity effect is applied at the contact surface instead of a point force. The development of non-uniform distributed stress
 480 calculated by the proposed analytical methods was validated by numerical simulation, which presents the evolution process of
 481 biased rockfall from partial damage of the soft underlying layer, caused by non-uniform distributed stress, to toppling and

482 ~~sliding of overhanging hard rock block due to overall unbalanced force. In this method, a non-uniform distributed stress due to~~
483 ~~the eccentricity effect is applied at the contact surface instead of a point force.~~ The method considers four failure modes
484 according to the rockfall evolution process, including partial damage of the soft foundation (Fos_{co} and Fos_{te}) and overall
485 failure of the rock block (Fos_{st} and Fos_{to}).

486 Taking the northeast edge of the Sichuan Basin in Southwest China as the study area, the proposed method is used to calculate
487 the Fos of biased unstable rock blocks. The results show that in the natural scenario, the underlying mudstone of some rock
488 blocks has been partially damaged, and compression failure of the mudstone has been observed in the field. Some rock blocks
489 are expected to fail as a whole in rainfall or earthquake scenarios. The statistical analysis indicates that the retreat ratio is the
490 crucial factor influencing the Fos of biased rockfall. On the basis of different combinations of four Fos , rockfall susceptibility
491 was classified into three levels. As the retreat rate increases, the rock blocks undergo an evolution process from stability to
492 partial instability and then overall instability. Based on the current mechanical parameters of the eastern Sichuan Basin, the
493 critical retreat ratio from low to moderate rockfall susceptibility is 0.33.

494 The proposed method improves the three-dimensional mechanical model of a rock block with a basal cavity by considering
495 non-uniform distributed stress at the contact surface, which could promote the accuracy of rockfall stability analysis. Due to
496 the assumptions adopted and the complexity of the failure mechanism of biased rockfall, there are some limitations in this
497 method, mainly including the simplification of boundary conditions and rock deformation. These limitations will be important
498 considerations in future studies.

499 **Data availability**

500 All raw data can be provided by the corresponding authors upon request.

501 **Author contributions**

502 XS, BC and JD planned the campaign; XS and BC performed the field measurements; XS, BC, WW and BL designed and
503 developed the methodology. XS, BC and JD analysed the data; XS and BC wrote the manuscript draft; JD and WW reviewed
504 and edited the manuscript.

505 **Competing interests**

506 The authors declare that they have no conflicts of interest.

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