

1 Winter brown carbon over six China's megacities: Light 2 absorption, molecular characterization, and improved 3 source apportionment revealed by multilayer perceptron 4 neural network

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16 **Abstract.** Brown carbon (BrC) constitutes a large fraction of organic carbon and exhibits strong light
17 absorption properties, thus affecting the global radiation budget. In this study, we investigated the light
18 absorption properties, chemical functional bonds, and sources of BrC in six megacities in China, namely
19 Beijing, Harbin, Xi'an, Chengdu, Guangzhou, and Wuhan. The average values of the BrC light
20 absorption coefficient and the mass absorption efficiency at 365 nm in northern cities were higher than
21 those in southern cities by 2.5 and 1.8 times, respectively, demonstrating the occurrence of abundance of
22 BrC in northern China's megacities. Fourier transform-infrared (FT-IR) spectra revealed sharp and
23 intense peaks at 1640, 1458–1385, and 1090–1030 cm⁻¹, which were ascribed to aromatic phenols,
24 confirming the contribution of primary emission sources (e.g., biomass burning and coal combustion) to
25 BrC. In addition, we noted peaks at 860, 1280–1260, and 1640 cm⁻¹, which were attributed to
26 organonitrate and oxygenated phenolic groups, indicating that secondary BrC also existed in six
27 megacities. Positive matrix factorization (PMF) coupled with multilayer perceptron (MLP) neural
28 network analysis were used to apportion the sources of BrC light absorption. The results showed that
29 primary emissions (e.g., biomass burning, tailpipe emissions, and coal combustion) made a major
30 contribution to BrC in six megacities. However, secondary formation processes made a greater

31 contribution to light absorption in the southern cities (17.9%–21.2%) than in the northern cities (2.1%–
32 10.2%). These results can provide a basis for the more effective control of BrC to reduce its impacts on
33 regional climates and human health.

34 **1 Introduction**

35 Brown carbon (BrC) constitutes a vital fraction of carbonaceous aerosols and exhibits strong light
36 absorption properties in near-ultraviolet (UV) and visible wavelength regions (Laskin et al., 2015; Wu et
37 al., 2021; Zhang et al., 2022). Therefore, it has received extensive attention in recent years (Laskin et al.,
38 2015; Yan et al., 2018; Yuan et al., 2020). BrC has substantial effects on radiative forcing, cloud
39 condensation, ice cores, and climate (Ma et al., 2020; Sreekanth et al., 2007). On the basis of remote
40 sensing observations and chemical transport model results, studies have detected a BrC-induced
41 nonnegligible positive radiative forcing ranging from 0.1 to 0.6 W m⁻² on a global scale (Jo et al., 2016;
42 Wu et al., 2020).

43 BrC in urban atmospheres can originate from numerous sources, including incomplete combustion of
44 fossil fuels (Soleimanian et al., 2020), biomass burning (Shen et al., 2017; Soleimanian et al., 2020),
45 forest fires, and residential coal combustion (Kirchstetter et al., 2004; Soleimanian et al., 2020). In
46 addition, both primary BrC and gaseous pollutants emitted from anthropogenic and biological activities
47 can be converted into secondary BrC through a series of atmospheric chemical reactions (Kumar et al.,
48 2018; Laskin et al., 2015). Studies have determined that the absorption properties of BrC exhibited
49 distinct temporal and spatial variations in different regions and cities, and these properties were closely
50 related to diverse emissions sources and complex atmospheric aging processes (Chung et al., 2012; Wu
51 et al., 2021). For example, Devi et al. (2016) observed that BrC contributed differently to light absorption
52 in the rural and urban southeast United States. Mo et al. (2021) studied the light absorption coefficient
53 of BrC at 365 nm (BrC $b_{\text{abs}365}$) in ten Chinese cities, which found that the BrC $b_{\text{abs}365}$ value displayed
54 obvious spatial (northern China > southern China) variations. Furthermore, a stronger light absorption
55 ability in cold seasons (fall and winter) in Beijing (Cheng et al., 2016), Xi'an (Shen et al., 2017), Seoul
56 (Kim et al., 2016), Taiyuan and other cities (Mo et al., 2021) has been found to be strongly associated
57 with increased biomass burning emissions for heating. The mass absorption efficiency at 365 nm
58 (MAE₃₆₅) of BrC has been widely used to evaluate the light-absorbing ability of BrC (Bao et al., 2022).

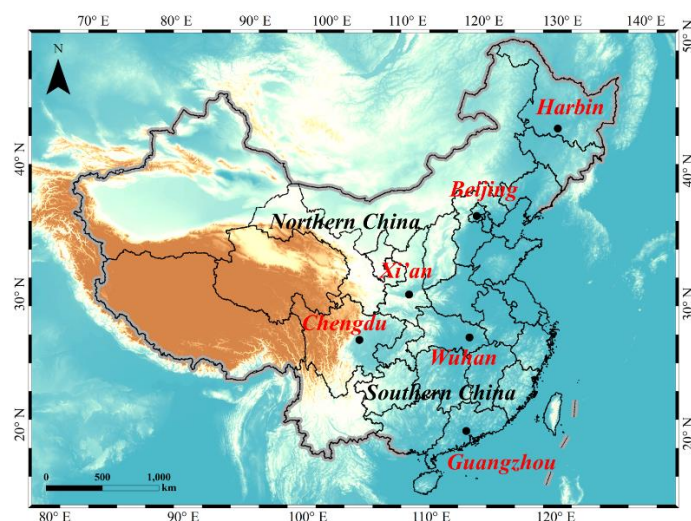
59 Xie et al. (2017) found that the BrC MAE₃₆₅ values from biomass burning ($1.28 \pm 0.12 \text{ m}^2 \text{ g}^{-1}$) were
60 higher than those from vehicle emissions ($0.62 \pm 0.76 \text{ m}^2 \text{ g}^{-1}$). Ni et al. (2021) noted that BrC MAE₃₆₅
61 values can be decreased from $1.43 \text{ m}^2 \text{ g}^{-1}$ to $0.11 \text{ m}^2 \text{ g}^{-1}$ with the BrC aerosol aged. Another study noted
62 that secondary organic aerosol (SOA) formation processes constituted a major source of BrC in Atlanta
63 and Los Angeles; moreover, the optical properties of BrC differed considerably between the two cities
64 due to differences in secondary BrC precursors (Zhang et al., 2011).

65 China has a high concentration of atmospheric water-soluble organic carbon, which has a major impact
66 on regional air quality, visibility, and the climate (Mo et al., 2021). However, to our knowledge, limited
67 study was conducted to insight to the optical profiles, molecular composition, and sources apportionment
68 of BrC in a large scale in China. Accurately understanding the spatial variations of the sources and light
69 absorption properties of BrC in China is essential for reducing uncertainty about the effects of BrC on
70 the climate. Many studies have used receptor modelling techniques such as positive matrix factorization
71 (PMF) coupled with multiple linear regression analysis to assign the sources of BrC (Bao et al., 2022;
72 Lei et al., 2019; Soleimanian et al., 2020). For example, Bao et al. (2022) obtained specific source
73 contributions to BrC $b_{\text{abs}365}$ in Nanjing based on PMF and MLR method, confirming that the key
74 contributors to BrC $b_{\text{abs}365}$ were mainly derived from biomass burning, primary industrial, and traffic
75 emissions. Lei et al. (2018) investigated the source apportionment of BrC $b_{\text{abs}365}$ in Yulin and showed that
76 the residential coal combustion was the highest contributor to BrC $b_{\text{abs}365}$ in winter. Soleimanian et al.
77 (2020) used the principal component analysis (PCA) coupled with MLR source apportionment model,
78 which identified fossil fuel combustion was the dominant source of BrC $b_{\text{abs}365}$ in central Los Angeles
79 during summer (38%), followed by SOA (30%) and biomass burning (12%). However, atmospheric
80 processes are generally non-linear in nature, thus traditional deterministic models could be limited. The
81 artificial neural network (ANN) based models, such as multilayer perceptron (MLP), have been shown
82 to provide meaningful results closer to realistic estimates than most linear models (Borlaza et al., 2021a;
83 Elangasinghe et al., 2014). Therefore, in this study, a winter campaign for PM_{2.5} sampling was conducted
84 over six China's megacities. The purposes of this study were to 1) investigate the spatial variations of the
85 carbonaceous matter concentrations and optical properties of BrC across six representative urban areas
86 in China, 2) determine the molecular composition of BrC, and 3) insight the relationship between light
87 absorption and BrC sources by using PMF coupled with ANN-MLP.

88 2 Methods

89 2.1 Samples collection

90 PM_{2.5} samples were collected in six cities in China (Figure 1): three cities in northern China (Beijing
91 [BJ], Harbin [HrB], and Xi'an [XA]) and three cities in southern China (Chengdu [CD], Guangzhou
92 [GZ], and Wuhan [WH]). We classified the cities as being in northern or southern cities according to their
93 geographic location, such as “north or south of the Huaihe River”. Owing to geographical factors, these
94 cities exhibit considerable differences in terms of energy structure and climate. The average annual
95 temperature in northern cities is generally below 15°C, while in southern cities it is usually above 15°C
96 (Mo et al., 2021). Information about the six cities and the sampling sites is summarized in Table S1
97 (Supporting Information).



98
99 **Figure 1. PM_{2.5} samples were taken in six Chinese cities.**

100 For sample collection, filter samplers were mounted on rooftops between 8 and 30 m above the ground,
101 and samples were collected from November 20 to December 22, 2019. In BJ, HrB, and GZ, a mini-
102 volume sampler operating at 5 L min⁻¹ (Airmetrics, Springfield, OR, USA) was used to collect PM_{2.5}
103 samples on 47-mm quartz-fiber filters (Whatman, Maidstone, UK) for 24 h. In CD, a medium-volume
104 PM_{2.5} sampler operating at 100 L min⁻¹ (HY-100SFB, Hengyuan, Qingdao, China) was used to collect
105 PM_{2.5} samples on 90-mm quartz-fiber filters (Whatman). Moreover, in XA and WH, a high-volume
106 sampler (HVS-PM_{2.5}, Thermo-Anderson Inc. Cleves, OH, USA) with a flow rate of 1.13 m³ min⁻¹ was
107 used to collect PM_{2.5} samples on quartz-fiber filters (203 mm × 254 mm, Whatman, QMA). Before
108 sample collection, all quartz filters were prebaked at 780 °C for 7 h to eliminate any residual carbon. A

109 detailed description of the quality control procedures for the filters before and after the sampling
110 processes can be found in the article by Shen et al (2017). After the sampling processes, the samples were
111 sealed and stored below 0 °C to avoid evaporative losses before analysis.

112 **2.2 Chemical analysis**

113 The organic carbon (OC) and elemental carbon (EC) of the PM_{2.5} samples were analyzed using a
114 Thermal and Optical Carbon Analyzer (DRI Model 2001A, Atmoslytic, Inc., USA) in accordance with
115 the improved Interagency Monitoring of Protected Visual Environment (IMPROVE) thermal/optical
116 reflectance protocol. Detailed descriptions of the OC and EC measurement methods can be found in the
117 article by Cao et al (2004). A portion of each filter (about 2.84 cm²) was extracted using 10 mL of
118 ultrapure water to analyze water-soluble inorganic ions (Na⁺, NH₄⁺, K⁺, Mg²⁺, Ca²⁺, Cl⁻, NO₃⁻, and SO₄²⁻)
119 through ion chromatography (Dionex 500, Dionex Corp, USA). A detailed description of the ion analysis
120 method used in this study can be found in the article by Shen et al (2008).

121 **2.3 Optical properties of methanol extracts**

122 A 0.526-cm² punch was ultrasonically extracted from each filter sample by using 5 mL of methanol
123 (HPLC Grade, Fisher Scientific, NH, USA) for 30 min. Subsequently, all extracts were filtered through
124 a microporous membrane with a diameter of 25 mm and pore size of 0.22 μm (Puradisc 25 TF, PTFE
125 membrane) to remove insoluble components. The UV–visible absorption spectra of the BrC samples
126 were determined using a liquid waveguide capillary cell–total OC spectrophotometer (LWCC-2100,
127 World Precision, Sarasota, FL, USA) between the wavelengths of 200 and 700 nm. The BrC optical
128 properties such as $b_{abs365, methanol}$ (The absorption coefficient for methanol extracts at 365 nm) and $MAE_{365, methanol}$
129 (normalized by $b_{abs365, methanol}$ to organic carbon, OC) were calculated as showed in previous study
130 (Lei et al., 2019) and details was listed in Text S1.

131 **2.4 Fourier transform infrared spectroscopy spectra**

132 Functional groups in the samples collected in six megacities were characterized using a Fourier
133 transform infrared (FT-IR) spectrometer (Bruker Optics, Billerica, MA, USA). The method described in
134 section 2.3 was used to extract the BrC filtrates, then the BrC extracts were concentrated to 0.5 mL under
135 a gentle nitrogen flow, after which they were mixed with 0.2 g of KBr (FT-IR grade, Sigma-Aldrich) and

136 then blown with nitrogen to complete dryness. The resulting extract–potassium bromide mixture was
137 ground in an agate mortar and examined through FT-IR spectroscopy. The FT-IR spectrum of each sample
138 was recorded in transmission mode by averaging 64 scans using a standard optical system with KBr
139 windows. The spectra were recorded in the wavelength range of 4000–400 cm^{-1} at a resolution of 4 cm^{-1} .
140 Before analyzing the aerosol extract samples, we obtained the baseline spectrum by analyzing pure KBr.

141 **2.5 Source apportionment of BrC light absorption coefficient at 365 nm**

142 In this study, the source apportionment of BrC was conducted using the PMF coupled with ANN-MLP
143 methods by following the steps: 1) identification and quantification of the major sources of $\text{PM}_{2.5}$ for the
144 six cities using PMF (The United States Environmental Protection Agency, PMF 5.0); 2) produces a
145 predictive model by ANN-MLP for one variable ($\text{BrC } b_{\text{abs}365}$) based on the values of the input variables
146 ($\text{PM}_{2.5}$ sources daily contributions). PMF is a bilinear factor model that has been widely used in source
147 apportionment studies (Cao et al., 2012; Lei et al., 2018; Li et al., 2021; Shen et al., 2010; Tao et al.,
148 2017). In the present study, water-soluble inorganic ions (Na^+ , NH_4^+ , K^+ , Mg^{2+} , Ca^{2+} , NO_3^- , SO_4^{2-} and
149 Cl^-) and carbon fractions (OC1, OC2, OC3, OC4, EC1, and EC2) were used as data inputs for PMF. The
150 PMF model was run multiple times, extracting four to six factors. A more detailed description of these
151 items can be found in the article by Lei et al (2019). Subsequently, an MLP model was constructed. The
152 model was developed using IBM SPSS Statistics for Windows, version 23 (IBM Corp., Armonk, NY,
153 USA). The detail information of the ANN-MLP model construction and training was described in Text
154 S2. After ANN-MLP model training, the obtained MLP model was applied to a set of virtual datasets.
155 Each virtual dataset consists of each source with the same mass contribution (from PMF analysis) as the
156 original dataset, but with one source set to zero. The $\text{BrC } b_{\text{abs}365}$ contribution for a specific source was
157 obtained by subtracting the $\text{BrC } b_{\text{abs}365}$ simulation value obtained using the virtual dataset from the BrC
158 $b_{\text{abs}365}$ simulation value obtained using the original MLP model, which contains all the source
159 contributions (Borlaza et al., 2021a).

160 **3 Results and discussion**

161 **3.1 General description of $\text{PM}_{2.5}$ and its chemical species in six megacities**

162 As presented in Table S2, the $\text{PM}_{2.5}$ concentrations in six cities ranged from 9.9 to 241.9 $\mu\text{g m}^{-3}$ and
163 exhibited a significant spatial variation ($p < 0.01$), indicating the complexity of air pollution and spatial

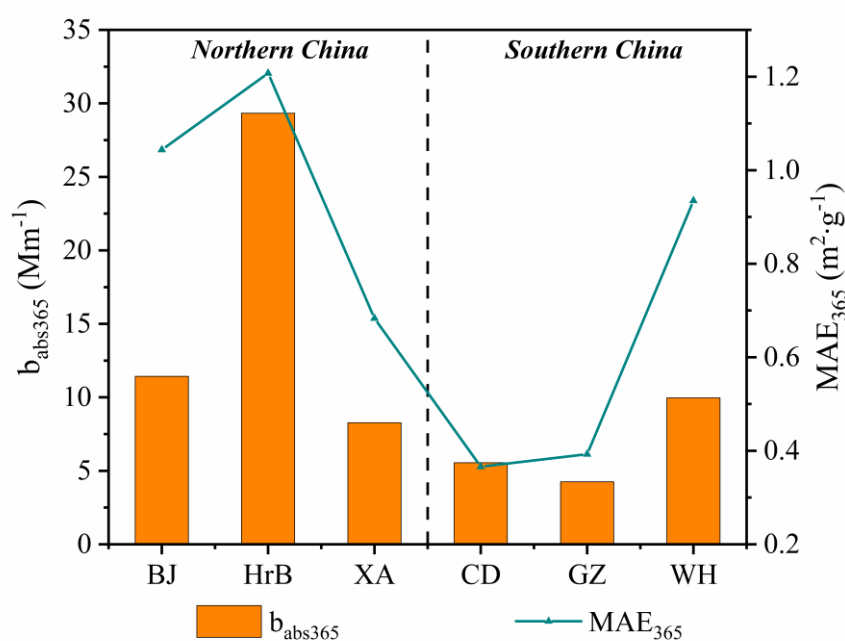
164 differences in air pollution levels in China. HrB had the highest average PM_{2.5} concentration (85.5 ± 43.9
165 $\mu\text{g m}^{-3}$), which exceeded National Air Quality Standard grade-II (24-h average: $75 \mu\text{g m}^{-3}$) and was 1.5,
166 1.1, 1.2, 2.0 and 1.3 times higher than those recorded in BJ, XA, CD, GZ, and WH, respectively. This
167 phenomenon indicates that PM_{2.5} pollution is still a major challenge in China, particularly in northern
168 China.

169 The average concentration of OC, a major chemical component of PM_{2.5}, ranged from 5.6 to 19.4 μg
170 m^{-3} in six megacities; these cities can be arranged (in descending order) as follows in terms of the average
171 OC concentration: HrB > XA > BJ > WH > GZ > CD (Table S2). Similar to the PM_{2.5} trend, the average
172 OC concentration in the northern cities ($15.5 \pm 7.9 \mu\text{g m}^{-3}$) was higher than that in the southern cities
173 ($9.2 \pm 4.6 \mu\text{g m}^{-3}$), which can be attributed to substantial emissions from residential heating (i.e., coal
174 and biomass combustion) in winter in northern China (Zhang et al., 2021). In addition, these residential
175 fuels can emit an abundant OC emission (Lei et al., 2018; Sun et al., 2017). To assess the sources of
176 atmospheric BrC, we estimated the concentrations of primary OC (POC) and secondary OC (SOC) by
177 using the EC tracer method (Ram and Sarin, 2011). Detailed calculation method was described in Text
178 S3. As presented in Table S2, the average SOC concentrations throughout the measurement period ranged
179 from 1.0 (CD) to 9.2 $\mu\text{g m}^{-3}$ (HrB), and the fractional contributions of SOC to OC varied from 22.6% to
180 66.6%. The average POC concentrations ranged from 4.0 (GZ) to 10.2 $\mu\text{g m}^{-3}$ (HrB), and POC
181 constituted 34.4%–77.4% of the total OC mass in the six cities. Accordingly, the SOC and POC
182 concentrations exhibited typical spatial fluctuations, which were consistent with the fluctuations of the
183 PM_{2.5} and total OC concentrations. These results reveal that primary emissions usually dominated
184 secondary formation processes, especially in the northern cities.

185 3.2 Light absorption properties of BrC

186 As plotted in Figure 2, the light absorption coefficient ($b_{\text{abs}}, \text{Mm}^{-1}$) values for BrC exhibited significant
187 spatial variations across the six cities ($1.7\text{--}64.1 \text{Mm}^{-1}$; $p < 0.01$). We executed Student *t* test at the 95%
188 confidence level and observed that HrB had the highest average BrC $b_{\text{abs}365}$ value ($29.3 \pm 14.2 \text{Mm}^{-1}$),
189 followed by BJ ($11.4 \pm 3.9 \text{Mm}^{-1}$), WH ($10.0 \pm 3.2 \text{Mm}^{-1}$), XA ($8.3 \pm 2.4 \text{Mm}^{-1}$), CD ($5.6 \pm 2.7 \text{Mm}^{-1}$),
190 and GZ ($4.3 \pm 1.4 \text{Mm}^{-1}$). The average BrC $b_{\text{abs}365}$ value in the northern cities was $15.7 \pm 12.3 \text{Mm}^{-1}$,
191 which was 2.5 times higher than that in the southern cities ($p < 0.01$). The large variation in the measured
192 BrC $b_{\text{abs}365}$ values in these megacities was observed, which reflected that the light absorption of BrC was

193 heavily affected by chromophore sources (Huang et al., 2018; Soleimanian et al., 2020), aging during
 194 atmospheric transportation (Lambe et al., 2013), and meteorological conditions (Li et al., 2021). Light-
 195 absorbing carbonaceous aerosols were believed to be responsible for the considerable absorption of light
 196 in the atmosphere (Xie et al., 2020). As presented in Figure S2, we observed positive correlations between
 197 BrC $b_{\text{abs}365}$ and POC in the six cities (r range: 0.61–0.92). Similar correlations were observed between
 198 BrC $b_{\text{abs}365}$ and SOC (r range: 0.51–0.80), indicating that the sources of atmospheric BrC in the six cities
 199 were quite complex. Apart from primary emissions, secondary formation processes also seemed to have
 200 a considerable contribution to BrC in these cities. Biomass burning was revealed to be the dominant
 201 source of BrC in these cities during winter (Cheng et al., 2016; Shen et al., 2017; Sun et al., 2017; Cheng
 202 et al., 2022). Furthermore, we observed high correlations (r range: 0.69–0.92) between BrC $b_{\text{abs}365}$ and
 203 K^+ , which is commonly regarded as a tracer of biomass burning (Shen et al., 2010), in HrB, BJ, XA, and
 204 WH (Figure S3). This evidence supports the aforementioned findings that emissions from biomass
 205 burning might be the major BrC source in winter in these cities. For the southern cities CD and GZ, the
 206 low BrC $b_{\text{abs}365}$ values ($1.7\text{--}11.5 \text{ Mm}^{-1}$) are of the same order of magnitude as those reported previously
 207 in Nanjing ($3.3\text{--}13 \text{ Mm}^{-1}$; Chen et al., 2019; Chen et al., 2018), Seoul ($0.9\text{--}7.3 \text{ Mm}^{-1}$; Kim et al., 2016),
 208 and Hong Kong ($4.8\text{--}10.6 \text{ Mm}^{-1}$; Zhang et al., 2020). The aging or oxidation of aerosols were confirmed
 209 to be the major source of BrC in these regions, indicating that secondary aerosols are likely a major
 210 source of winter BrC in CD and GZ.



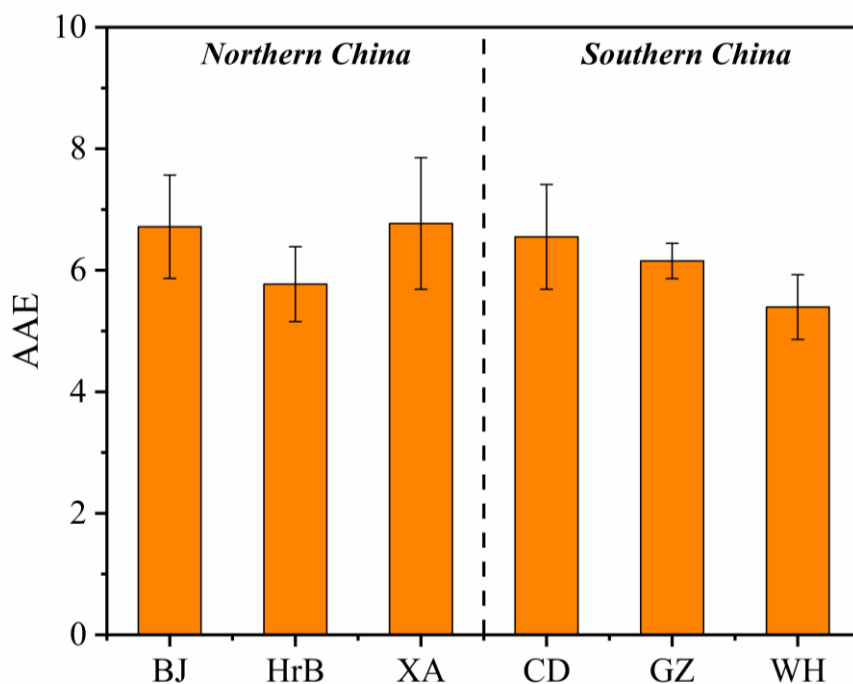
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212 **Figure 2. Spatial variations of BrC light absorption properties from six Chinese cities. The bars represent the**
213 **light absorption coefficient at 365 nm ($b_{\text{abs}365}$, left axis), and the lines represent the mass absorption efficiency**
214 **at 365 nm (MAE_{365} , right axis).**

215 The mass absorption efficiency (MAE , $\text{m}^2 \text{g}^{-1}$) is a key parameter for describing the light absorption
216 ability of atmospheric BrC (Li et al., 2021; Peng et al., 2020). Figure 2 illustrated the average MAE
217 values measured at 365 nm (BrC MAE_{365}) in the six cities; compared with the value measured in CD
218 ($0.37 \pm 0.18 \text{ m}^2 \text{g}^{-1}$), those measured in the other five cities were higher by 1.1–3.3 times. These cities
219 can be arranged as follows (in descending order) in terms of the measured BrC MAE_{365} values: HrB >
220 BJ > WH > XA > GZ > CD. These differences in BrC MAE_{365} values can be attributed to the variance
221 of the light absorption capacity of BrC in different megacities. The average BrC MAE_{365} values measured
222 in BJ, HrB, XA, and WH (range: $0.68\text{--}1.21 \text{ m}^2 \text{g}^{-1}$) were within the MAE ranges of biomass burning,
223 such as, the average MAE_{365} measured for BrC were $0.97 \pm 0.26 \text{ m}^2 \text{g}^{-1}$ for wood burning (Du et al.,
224 2014), $1.05 \pm 0.08 \text{ m}^2 \text{g}^{-1}$ for corn stalk combustion (Du et al., 2014), and $1.28 \pm 0.12 \text{ m}^2 \text{g}^{-1}$ for wheat
225 stubble burning (Xie et al., 2017; Lei et al., 2018), indicating that biomass burning may be a major source
226 of winter BrC in these cities. Biomass burning is commonly regarded as the main emission source for
227 BrC, which has a high absorption capacity, as indicated by field observations and model predictions
228 (Desyaterik et al., 2013; Feng et al., 2013; Lei et al., 2018). Notably, the MAE_{365} values derived for BrC
229 emitted from primary fossil fuel combustion are similar to those derived for biomass burning (Yan et al.,
230 2017); for example, former studies have revealed that the BrC MAE_{365} values produced by primary
231 emissions from residential coal combustion were in the range of $0.30\text{--}1.51 \text{ m}^2 \text{g}^{-1}$ (Ni et al., 2021; Yan
232 et al., 2017). Therefore, coal combustion may also be a potential source of BrC in these cities. By contrast,
233 we observed lower average BrC MAE_{365} values in GZ and CD (range: $0.37\text{--}0.39 \text{ m}^2 \text{g}^{-1}$). Previous studies
234 have revealed relatively low BrC MAE values from motor vehicle emissions, including gasoline vehicle
235 emissions ($0.62 \pm 0.76 \text{ m}^2 \text{g}^{-1}$; Xie et al., 2017) and motorcycle emissions ($0.20 \pm 0.08 \text{ m}^2 \text{g}^{-1}$; Du et al.,
236 2014). These findings suggest that the BrC sampled in GZ and CD mainly originated from traffic
237 emissions. In addition, laboratory experiments in a previous study revealed that MAE_{365} values decreased
238 from 1.43 to $0.11 \text{ m}^2 \text{g}^{-1}$ with aerosol aging, which suggests the production of SOA (Ni et al., 2021). This
239 finding demonstrates that secondary formation processes are among the main sources of BrC in CD and
240 GZ.

241 The absorption Ångström exponent (AAE) measurements at 330–550 nm represents the wavelength

242 dependence of light absorption by BrC (Cheng et al., 2017). We observed that the average AAE values
243 for BrC varied from 5.4 to 6.8 in the six cities (Figure 3). In general, the AAE values obtained in this
244 study were higher than those obtained at the Nepal Climate Observatory-Pyramid (3.7–4.0; 330–500 nm)
245 (Kirillova et al., 2016) and in the Los Angeles Basin (4.82 ± 0.49 ; 300–600 nm) (Zhang et al., 2013) and
246 lower than those obtained at the Tibetan Plateau (8.2 ± 1.4 ; 365–550 nm) (Zhu et al., 2018). Nevertheless,
247 the values obtained in this study were comparable to those obtained in Beijing (5.3–7.3; 310–450 nm)
248 (Cheng et al., 2016; Wu et al., 2021), Nanjing (6.7; 300–600 nm) (Chen et al., 2018), the Indo-Gangetic
249 Plain (5.3; 300–700 nm) (Srinivas et al., 2016), New Delhi (5.1; 330–400 nm) (Kirillova et al., 2014),
250 Seoul (5.5–5.8; 300–700 nm) (Kim et al., 2016), and Xi'an (5.3–6.1; 330–550 nm) (Huang et al., 2018).
251 These similarities can primarily be attributed to the consistent solubility of chromophores, which are
252 sensitive to the type of fuel used, the combustion conditions, and the solvents used (Cao et al., 2021; Huo
253 et al., 2018). Furthermore, the AAE values obtained in this study were within the range of those reported
254 by previous studies for coal combustion (5.5–6.4; 300–500nm) (Ni et al., 2021), biomass burning (4.4–
255 8.7; 300–550nm) (Xie et al., 2017), and gasoline vehicle emissions (6.2–6.9; 300–550 nm) (Xie et al.,
256 2017). This suggested that BrC in our study may have multiple sources. Additionally, in contrast to the
257 trends observed for the BrC $b_{\text{abs}365}$ and BrC MAE₃₆₅ values in the various cities, the AAE values observed
258 in CD and GZ were higher than those observed in the other cities. A previous study reported that the
259 AAE values for SOA were higher than those for primary organic aerosols (Saleh et al., 2013), and
260 previous laboratory combustion experiments revealed that the aging of biomass burning aerosols
261 generally engenders an increase in AAE values (from 6.93 to 15.59; Sengupta et al., 2018). These
262 findings suggested that BrC in the cities in this study was also affected by secondary formation processes.



263

264 **Figure 3. AAE values of BrC in six cities. AAE is calculated between 330 and 550 nm.**

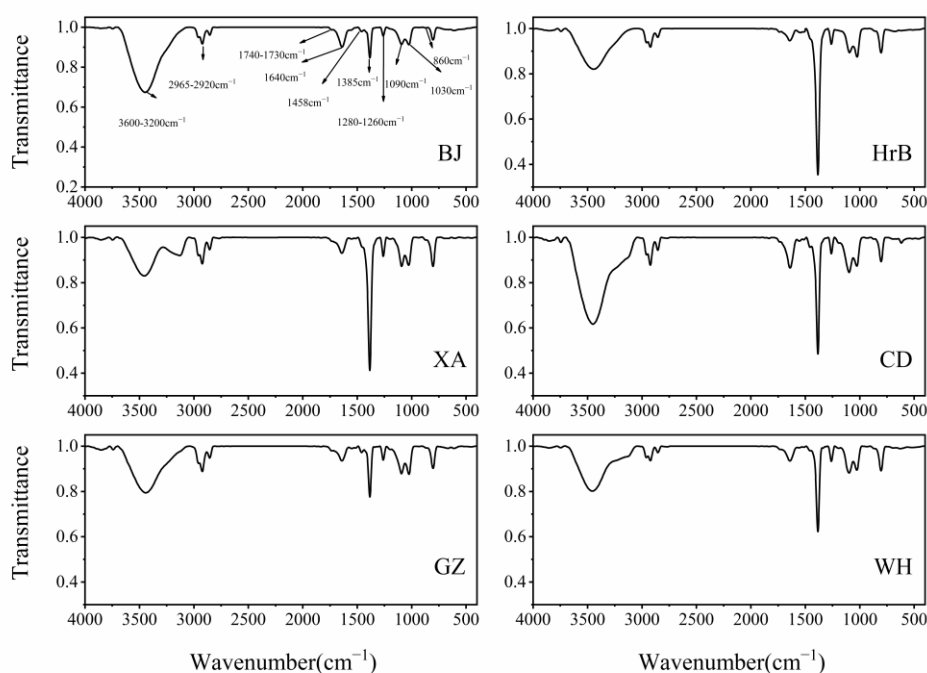
265 **3.3 Molecular structure of BrC**

266 In order to further explore the reasons for the differences in the optical properties of BrC among these
 267 cities, the functional groups of BrC were measured using FT-IR spectroscopy. Figure 4 illustrates the FT-
 268 IR spectra of BrC fractions within the region of 4000–400 cm^{-1} in the six cities. The band in the region
 269 of 400–800 cm^{-1} resulted from the interference from water vapor inside the instrument and thus can be
 270 ignored (Zhang et al., 2020). The broad and strong peak at 3450 cm^{-1} was contributed to the O-H stretch
 271 of H-bonded hydroxyl groups, phenols and carboxylic (Fan et al., 2016; Mukherjee et al., 2020). The
 272 sharp band near 1740 cm^{-1} was usually assigned to the C=O bonds of ketones, quinones, and amides
 273 (Duarte et al., 2005; Kristensen et al., 2015). We also attributed the sharp and intense absorption peaks
 274 at 2850–2990 cm^{-1} to aliphatic asymmetric and symmetric C–H stretching vibrations (Coury and Dillner,
 275 2008). Some bands were also displayed near 1640, 1458 and 1030 cm^{-1} , previous studies confirmed that
 276 these bands were generally ascribed to the C=C and C–H stretching of aromatic rings (Fan et al., 2016;
 277 Zhao et al., 2022), indicating the presence of aromatic groups. These results demonstrate the complexity
 278 of the chemical composition of BrC in the six cities, mainly containing aliphatic chains, carboxylic
 279 groups, and aromatic groups.

280 In contrast to these similar functional groups, the apparent differences of typical functional bands were

281 also found among these cities. The strong band near 3130 cm^{-1} denoting O–H band (Fan et al., 2016;
282 Mukherjee et al., 2020) were only detected in XA, CD and WH, and the same peak were observed in the
283 spectra from the corn straw burning (Fan et al., 2016) and coal combustion (Zhang et al., 2022), which
284 stressed the emissions of biomass burning and coal combustion with high abundance of oxygenated
285 phenolic compounds in these cities. Moreover, the peak at 1385 cm^{-1} was generally considered to be
286 derived from the O–H bond deformation and C–O stretching of phenolic groups (Fan et al., 2016;
287 Mukherjee et al., 2020; Zhang et al., 2020), and the same peak was observed in the FT-IR spectra of BrC
288 samples derived from the combustion of biomass materials (Fan et al., 2016). These observations
289 indicated the contribution of biomass burning to BrC; this was because that biomass burning can release
290 heat-modified lignin derivatives such as aromatic phenols (e.g., syringyl and guaiacyl) (Duarte et al.,
291 2007; Fan et al., 2016; Zhao et al., 2022). It was noted that the abundance of this peak was different
292 among six cities, and was significantly higher in HrB, XA and WH, which indicated biomass burning
293 contributed differently to BrC in six cities, and higher contribution was occur in HrB, XA and WH than
294 those in other cities. Previous studies have shown that BrC from biomass burning has a high light
295 absorption capacity (Cao et al., 2021; Desyaterik et al., 2013; Kumar et al., 2018), which supported that
296 these cities with higher abundance of aromatic phenol functional groups were consisted with higher BrC
297 $b_{\text{abs}365}$ (range: $8.3\text{--}29.3\text{ Mm}^{-1}$) and BrC MAE₃₆₅ (range: $0.68\text{--}1.21\text{ m}^2\text{ g}^{-1}$) values in section 3.2.

298 Furthermore, we observed three peaks at 860 , $1280\text{--}1260$, and 1640 cm^{-1} , demonstrating the presence
299 of organic-nitrate (C–ONO₂) and oxygenated phenolic groups (Day et al., 2010; Zhang et al., 2020).
300 Previous studies have shown that the anthropogenic volatile organic compounds, sulfates, nitrates and
301 other acidic particle components from coal and biomass combustion may enhance the contents of these
302 functional groups through aqueous-phase formation under high humidity conditions (Gilardoni et al.,
303 2016; Wang et al., 2019; Zhang et al., 2020). Therefore, the FT-IR spectra indicated that all the BrC
304 samples from six cities have the contribution of secondary generation. Besides, the abundance of
305 functional groups at these wavenumbers, especially at 1640 cm^{-1} , was higher in CD than that in other
306 cities. These results might indicate that the secondary source of BrC was relatively high in CD.



307
308 **Figure 4. FTIR spectra of BrC in six megacities.**

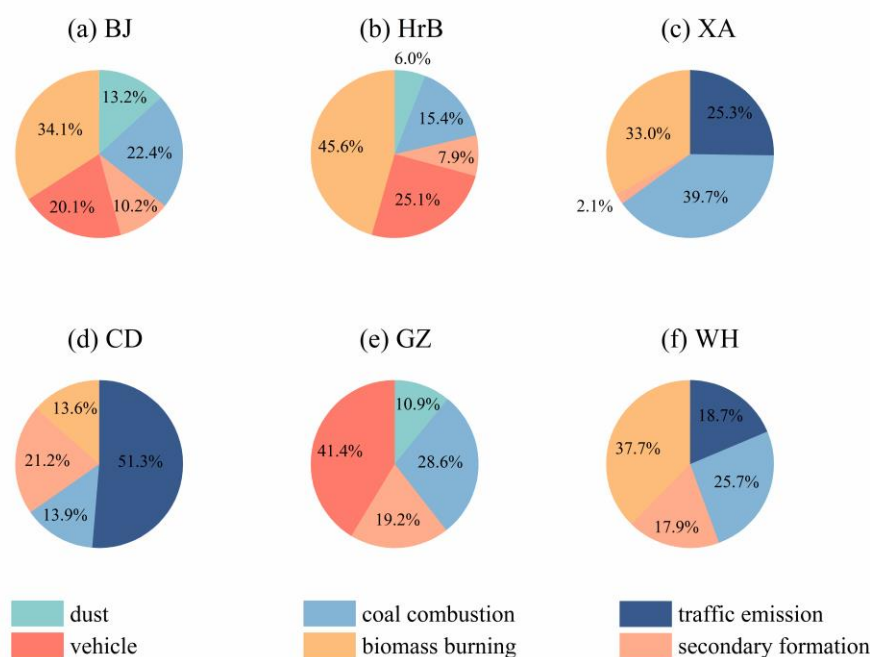
309 **3.4 Source apportionment of BrC**

310 Considering the complexity of atmospheric processes, and the correlation and/or nonlinear interaction
 311 between independent variables (i.e., multicomponent or multi-source interactions), we attempted to apply
 312 ANN techniques of nonlinear functions, such as MLP model, combined PMF analysis to predict the
 313 source contribution of allocated BrC from PM_{2.5} sources in this study. The PMF-apportioned source
 314 contributions to PM_{2.5} in the six cities are presented in Figures S4 and S5. A good correlation was
 315 observed between the measured and PMF-reconstructed PM_{2.5} mass concentrations in all sites (BJ: r =
 316 0.99; HrB: r = 0.90; XA: r = 0.97; CD: r = 0.97; GZ: r = 0.94; WH: r = 0.95), theoretical Q_{true} and Q_{robust}
 317 displayed a <5% difference, and scaled residuals of >95% data were in the range of -3 to 3. These
 318 evidences demonstrated the validity and robustness of our PMF solutions (Borlaza et al., 2021b; Tao et
 319 al., 2021). As illustrated in Figure S4, the first source was dominated by sulfate, OC, and EC and was
 320 considered to represent from coal combustion (Huang et al., 2014). The second source comprised high
 321 concentrations of NH_4^+ , NO_3^- , and SO_4^{2-} and was considered to represent secondary formation processes
 322 (Shen et al., 2010). Furthermore, the third source comprised high loadings of K^+ and was considered to
 323 represent biomass burning (Shen et al., 2010). The fourth source primarily comprised Na^+ , Mg^{2+} , and
 324 Ca^{2+} and was thus determined to represent fugitive dust (Shakeri et al., 2016; Shen et al., 2016; Sun et

325 al., 2019). The fifth source contained high concentrations of Mg^{2+} , Ca^{2+} , NO_3^- , OC, and EC and was thus
326 identified as representing traffic-related emissions (Shakeri et al., 2016). Finally, the sixth source
327 comprised high concentrations of OC, EC, and NO_3^- and was considered to represent vehicle emissions
328 (Shakeri et al., 2016).

329 The optimal neural network model for each site were explored by changing activation function types
330 (Tan H and Sigmoid), optimizing algorithms (scaled conjugate and gradient descent), and based on the
331 lowest root mean square error (RMSE) and the highest correlation coefficient (r) between observed and
332 MLP-modelled values (Borlaza et al., 2021a). Although there are other architectures that are more
333 complex for MLP models, a basic MLP architecture was considered sufficient for the input and output
334 data sets of this study.

335 Figure S6 shows the correlation between observed values and BrC $b_{\text{abs}365}$ predicted values from
336 selected MLP models. The good correlation indicated the reliability of the model results. On the basis of
337 the MLP results, we calculated the source-specific contributions to BrC in the six cities (Figure 5). The
338 primary sources including coal combustion, dust, vehicle, biomass burning and traffic emissions, and
339 their average contribution to BrC in the northern cities was 93.3%, which was 1.2 times higher than that
340 in the southern cities. Among these primary emissions, we noted that a higher contribution of biomass
341 burning to BrC in HrB, BJ, XA and WH compared to other cities, which is consistent with the higher
342 abundance of biomass burning products, such as aromatic phenol functional groups was founded in these
343 cities as discussed in section 3.3. As supported, the BrC from biomass burning have high MAE_{365} values
344 (Cao et al., 2021; Kumar et al., 2018), which can be also observed among these cities (range: 0.68–1.20
345 $\text{m}^2 \text{g}^{-1}$). In addition, we noted that the contribution of biomass burning to BrC in WH (37.7%) was higher
346 than that in CD (13.6%) and GZ (0%), which can explain the highest BrC MAE_{365} was observed in WH
347 among southern cities as shown in Figure 2. On average, the secondary formation source contribution to
348 BrC in southern cities was 19.4%, which was 2.9 times higher than that in northern cities. Besides, the
349 highest contribution was observed in CD with 21.2%, followed by $\text{GZ} > \text{WH} > \text{BJ} > \text{HrB} > \text{XA}$. This
350 result can be supported by the abundance of organic-nitrate functional groups, the relatively high AAE
351 value and low BrC MAE_{365} value in CD, which were closely related with the contribution of secondary
352 sources.



353

354 **Figure 5. The source contribution to BrC using multilayer perceptron neural network analysis in (a) BJ, (b)**
 355 **HrB, (c) XA, (d) CD, (e) GZ, (f) WH.**

356 4 Conclusions

357 We investigated the sources and light absorption properties of BrC in wintertime in six megacities
 358 across China. Both the b_{abs} and MAE_{365} of BrC at 365 nm in northern cities were approximately 2.5 and
 359 1.8 times higher than those in southern cities. The BrC MAE_{365} values measured in BJ, HrB, XA and
 360 WH were ranged from 0.68 to 1.21 $\text{m}^2 \text{g}^{-1}$, which were within the MAE ranges derived for biomass
 361 burning. Thus, these comparisons confirmed that emissions from biomass burning might be the major
 362 BrC source in winter in these cities. Previous studies have reported that MAE_{365} values decreased with
 363 aerosol aging while the AAE values of SOA were higher than those for POA. Besides, we noticed that
 364 the average BrC MAE_{365} and AAE values showed different trends in southern cities of CD and GZ, that
 365 is, the BrC MAE_{365} values of these two cities were lower than those of other cities, while the AAE values
 366 were relatively higher. These evidences supported the secondary formation process were among the main
 367 sources of BrC in CD and GZ.

368 The chemical functional groups of BrC in six cities mainly included aliphatic chains, carboxyl groups
 369 and aromatic groups. However, the apparent difference of typical functional bands revealed the important

370 contributions of primary biomass burning and coal combustion to BrC for high abundance of oxygenated
371 phenolic compounds in these cities, especially in HrB, XA and WH. In contrast, the presence of organic-
372 nitrate (C-ONO₂) and oxygenated phenolic groups in BrC molecular implied the contribution from
373 secondary formation in six megacities, especially in CD city.

374 Due to the complexity of atmospheric processes, which are usually non-linear in nature, and the
375 traditional linear-based source analytic models may be limited. Here, we used a multilayer perceptron
376 (MLP) model based on artificial neural network (ANN) to improve the source allocation of BrC in these
377 cities. Source apportionment of BrC based on PMF and ANN-MLP analysis revealed that primary
378 emissions (e.g., biomass burning, coal combustion, and vehicle emissions) were key contributors to BrC,
379 and their average contribution in northern cities was about 93.3%, which was 1.2 times higher than that
380 in southern cities. Secondary formation processes made a greater contribution to BrC in southern cities
381 (19.4%) than northern cities (6.7%). The results of our work can provide a basis for the development of
382 more effective practices to control BrC emissions at the regional level.

383

384 **Data availability.** The key data sets are publicly available on the Zendo data repository platform:
385 <https://zenodo.org/record/6790321>.

386

387 **Author contribution.** ZS: Conceptualization. DW, TZ, and SH: Data curation. ZS and HX: Funding
388 acquisition. DW and YL: Methodology. ZS and JC: Resources. DW: Writing - original draft. DW, ZS,
389 QZ, HX, JS, JC and YL: Writing - review & editing.

390

391 **Competing interests.** The authors declare that they have no conflict of interest.

392

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