1 Ultradian rhythms in shell composition of photosymbiotic and non-photosymbiotic 2 mollusks Niels J. de Winter<sup>1,2</sup>, Daniel Killam<sup>3</sup>, Lukas Fröhlich<sup>4</sup>, Lennart de Nooijer<sup>5</sup>, Wim Boer<sup>5</sup>, Bernd R. 3 Schöne<sup>4</sup>, Julien Thébault<sup>6</sup>, Gert-Jan Reichart<sup>2,5,7</sup> 4 5 Affiliations 6 7 <sup>4</sup>Analytical, Environmental and Geochemistry group (AMGC), Vrije Universiteit Brussel, Brussels, **Belgium** 8 9 <sup>2</sup>Dept<sup>1</sup>Dept of Earth Sciences, Utrecht UniversityVrije Universiteit Amsterdam, 10 UtrechtAmsterdam, the Netherlands <sup>2</sup>Analytical, Environmental and Geochemistry group (AMGC), Vrije Universiteit Brussel, Brussels, 11 12 **Belgium** <sup>3</sup> San Francisco Estuary Institute, Richmond, CA, USA 13 Department of Paleontology, Institute of Geosciences, 14 4 Johannes Gutenberg UniversitätUniversity of , Mainz, Germany 15 <sup>5</sup> Dept. of Ocean Systems, Royal Netherlands Institute for Sea Research (NIOZ), Texel, the 16 17 Netherlands 18 <sup>6</sup>Univ Brest, CNRS, IRD, Ifremer, LEMAR, 29280 Plouzané, France, (ORCID: 0000-0002-3111-4428) 19 20 <sup>7</sup>Dept. of Earth Sciences, Utrecht University, Utrecht, the Netherlands 21 Corresponding author: N.J. de Winter, niels.de.winter@vub.ben.j.de.winter@vu.nl 22

#### 24 Abstract

The chemical composition of mollusk shells is a useful tool in (paleo)climatology since it captures 25 inter- and intra-annual variability in environmental conditions. Trace element and stable isotope 26 analyses analysis with improved sampling resolution now allows in-situ determination of the 27 28 composition of mollusk shell volumes precipitated at daily to sub-daily time intervalsnow enable 29 the use of mollusk shells for paleoenvironmental reconstructions at a daily to sub-daily resolution. Here, we discuss hourly resolved Mg/Ca, Mn/Ca, Sr/Ca and Ba/Ca profiles measured by laser 30 31 ablation ICP-MS through shells of photosymbiotic giant clams (Tridacna maxima, Tridacna T. squamosa and Tridacna-T. squamosina) and the non-photosymbiotic scallop Pecten maximus. 32 33 Precise sclerochronological age models and spectral analysis allowed us to extract daily and tidal 34 rhythms in the trace element composition of these shells. We find weak but statistically significant 35 expression of these periods and concludeWe find significant expression of these periodicities but 36 conclude that this cyclicity explains less than 10\_% of the sub-annual variance in trace element profiles. Tidal and diurnal rhythms explain variability of at most 0.2 mmol/mol (~10\_% of mean 37 value) in Mg/Ca and Sr/Ca, while ultradian Mn/Ca and Ba/Ca cyclicity has a median amplitude of 38 less than 2 µmol/mol (~\_40\_% and 80\_% of the mean of Mn/Ca and Ba/Ca, respectively). Daily 39 40 periodicity in Sr/Ca and Ba/Ca is stronger in Tridacna than in Pecten, with Pecten showing stronger tidal periodicity. One T. squamosa specimen which grew under a sunshade exhibits 41 some of among the strongest diurnal cyclicity. Daily cycles in trace element composition of giant 42 clams are therefore unlikely to be driven by variations in direct insolation-itself, but rather reflect 43 44 an inherent biological rhythmic process affecting element incorporation. Finally, the large amount of short-term trace element variability unexplained by periodic variability tidal and daily rhythms 45 highlights the dominance of aperiodic processes in mollusk physiology and/or environmental 46 conditions on shell composition at the sub-daily scale. Future studies should aim to investigate 47

whether part of this aperiodic remaining variability in shell chemistry reliably records weather
patterns or circulation changes in the paleoenvironment.

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#### 51 **1. Introduction**

Patterns in growth increments, microstructure, and chemical composition of accretionary 52 53 carbonate bioarchives yield detailed information about the environmental conditions and 54 biological rhythm of carbonate producing animals ((Dunbar and Wellington, 1981; Jones, 1983; Witbaard et al., 1994; Klein et al., 1996; Surge et al., 2001; Schöne et al., 2005a; Ivany, 2012; 55 Schöne and Gillikin, 2013; DeCarlo and Cohen, 2017; Killam and Clapham, 2018a) Dunbar and 56 57 Wellington, 1981; Jones, 1983; Witbaard et al., 1994; Klein et al., 1996; Surge et al., 2001; Schöne 58 et al., 2005a; Ivany, 2012; Schöne and Gillikin, 2013; DeCarlo and Cohen, 2017; Killam and Clapham, 2018). These characteristics have spurred the development of a multitude of 59 techniques for extracting information about life history (Jones and Quitmyer, 1996; Schöne et al., 60 61 2005b; Eggins et al., 2003; Anand and Elderfield, 2005; Goodwin et al., 2009; Mahé et al., 2010; 62 Comboul et al., 2014; DeCarlo and Cohen, 2017; Judd et al., 2018; Winter et al., 2022)-(Jones and Quitmyer, 1996; Schöne et al., 2005b; Goodwin et al., 2009; Mahé et al., 2010; Comboul et 63 al., 2014; DeCarlo and Cohen, 2017; Judd et al., 2018; de Winter, 2022), carbonate chemistry 64 65 (Sinclair et al., 1998; Lazareth et al., 2003; Schöne et al., 2010; de Winter and Claeys, 2017; Warter and Müller, 2017; Huyghe et al., 2021; de Winter et al., 2021b) (Sinclair et al., 1996; 66 Lazareth et al., 2003; Schöne et al., 2010; de Winter and Claeys, 2017; Warter and Müller, 2017; 67 Huyghe et al., 2021; de Winter et al., 2021a) and microstructure (Lazier et al., 1999; Checa et al., 68 69 2007; Popov, 2014; Gilbert et al., 2017; Crippa et al., 2020; Höche et al., 2020, 2021; Wichern et al., 2022) (Lazier et al., 1999; Checa et al., 2007; Popov, 2014; Gilbert et al., 2017; Crippa et al., 70 2020; Höche et al., 2020; 2021; Wichern et al., 2022) from carbonate shells and skeletons. As a 71 result, (fossil) carbonate skeletons have gained much attention as archives of past environmental 72

and climate change (Lough, 2010; Schöne and Gillikin, 2013; Ivany and Judd, 2022) (e.g., Lough,
 2010; Schöne and Gillikin, 2013; Ivany and Judd, 2022 and references therein).

75 Three characteristics make the shells of marine mollusks especially valuable as climate archives: (1) Nearly all marine mollusks precipitate their shells in isotopic equilibrium with the oxygen 76 77 isotope value of ambient sea water, except for juvenile oysters-and, some mollusks growing near 78 hydrothermal vents and some deep-burrowing species (Schöne et al., 2004; Hallmann et al., 2008; Wisshak et al., 2009; Huyghe et al., 2021; de Winter et al., 2022) (Schöne et al., 2004; 79 Wisshak et al., 2009; Huyghe et al., 2021; de Winter et al., 2022), (2) mollusk shells have a high 80 fossilization potential and long geological history, dating back to the beginning of the Phanerozoic 81 (Al-Aasm and Veizer, 1986a, b; Jablonski et al., 2003; Cochran et al., 2010; Jablonski et al., 2017; 82 83 de Winter et al., 2017, 2018; Coimbra et al., 2020) (Al-Aasm and Veizer, 1986a; b; Jablonski et 84 al., 2003; Cochran et al., 2010; Jablonski et al., 2017; de Winter et al., 2017; 2018; Coimbra et 85 al., 2020), (3) the incremental growth of mollusk shells allows for internal dating within the shell, 86 yielding chronologies of shell growth with sub-annual precision (Richardson et al., 1980; Jones, 1983; Schöne et al., 2005b; Goodwin et al., 2009; Huyghe et al., 2019) (Richardson et al., 1980; 87 Jones, 1983; Schöne et al., 2005c; Goodwin et al., 2009; Huyghe et al., 2019). These advantages 88 89 enable mollusk shells to record important information about climate and ambient water chemistry on the seasonal scale. Thereby, reconstructions from mollusk shells are highly complementary to 90 other, less highly resolved but longer-term, climate and environmental reconstructions like 91 sedimentary records, tree rings and ice cores (Black, 2009; Bougeois et al., 2014; Petersen et al., 92 93 2016; Tierney et al., 2020; de Winter et al., 2021a) (Black, 2009; Bougeois et al., 2014; Petersen et al., 2016; Tierney et al., 2020; de Winter et al., 2021b). 94

The resolution of the mollusk shell archive is not limited to seasonal variability. Studies monitoring
the behavior of mollusks during growth experiments show that their activity varies as a function
of environmental conditions (e.g., temperature and food availability) and follows ultradian rhythms

98 which may contain daily to hourly periodicities, probably linked to diurnal and tidal cycles, or lack 99 periodic behavior altogether (Rodland et al., 2006; García-March et al., 2008; Tran et al., 2011; Ballesta-Artero et al., 2017; Xing et al., 2019; Tran et al., 2020) (Rodland et al., 2006; García-100 101 March et al., 2008; Tran et al., 2011; Ballesta-Artero et al., 2017; Xing et al., 2019; Tran et al., 102 2020). Analyses of growth patterns and, more recently, composition of shell carbonate deposited 103 at these short time intervals show that these rhythms can be recorded in mollusk shells (Pannella, 104 1976; Richardson et al., 1980; Sano et al., 2012; Warter et al., 2018; de Winter et al., 2020) (Pannella, 1976; Richardson et al., 1980; Sano et al., 2012; Warter et al., 2018; de Winter et al., 105 106 2020). This raises the question whether mollusk shells reliably record behavioral changes, high frequency (paleo-)-weather or circulation patterns (e.g. (Komagoe et al., 2018; Yan et al., 2020; 107 108 Poitevin et al., 2020) (e.g., Komagoe et al., 2018; Yan et al., 2020; Poitevin et al., 2020). 109 Alternatively, the presence of daily cyclicity in shell chemistry may yield information about the 110 paleobiology of extinct mollusks, such as the presence-use of photosymbiosis (e.g. (Sano et al., 111 2012; Warter et al., 2018; de Winter et al., 2020) (e.g., Sano et al., 2012; Warter et al., 2018; de 112 Winter et al., 2020). The latter seems plausible given the effect of photosymbiosis on shell 113 mineralization in modern tridacnids (Ip and Chew, 2021) (Ip and Chew, 2021) and on the trace 114 element composition of aragonite in modern photosymbiotic scleractinian corals (Cohen et al., 2002; Meibom et al., 2003; Inoue et al., 2018) (Cohen et al., 2002; Meibom et al., 2003; Inoue et 115 116 al., 2018). If proven true, daily variability in bivalve shells may serve as a proxy for photosymbiosis 117 in the fossil record (e.g. (de Winter et al., 2020) (e.g., de Winter et al., 2020). This is of interest 118 because photosymbiosis is a derived adaptation of some tropical bivalve species (e.g., tridacnids) and its prevalence in the fossil record has important implications for the ecological niche of fossil 119 120 mollusks (e.g., (Vermeij, 2013) Vermeij, 2013). In addition, photosymbiosis can affect mollusk shell composition, and understanding it is therefore critical for the interpretation of chemical 121 122 proxies in mollusk shells for environmental reconstructions (Killam et al., 2020) (Killam et al., 2020). Finally, improving our understanding of photosymbiosis in tropical ecosystems sheds light 123

on the resilience of photosymbiotic organisms to environmental change, now and in the geological
 past. The latter is of special interest in light of the ongoing climate and biodiversity crises, which
 are profoundly affecting these sensitive ecosystems\_(Pandolfi and Kiessling, 2014)-(Pandolfi and
 Kiessling, 2014).

128 In this study, we investigate shell growth patterns and shell chemistry of the photosymbiotic 129 bivalves Tridacna maxima, T. squamosa and T. squamosina as well as the non-photosymbiotic scallop Pecten maximus. P. maximus was chosen as a non-photosymbiotic counterpart in 130 131 comparison with the tridacnids because of its comparatively high growth rate and the presence of 132 daily striae visible on the outside-outer surface of its shell, which make it possible to construct 133 accurate shell chronologies (Chauvaud et al., 2005)-(Chauvaud et al., 2005). We combine ultrahigh-resolution (hourly resolved) Mg/Ca, Sr/Ca, Mn/Ca and Ba/Ca measurements in the shells 134 135 with detailed sclerochronology to investigate the variability in these trace elements over time in 136 all four species. The aim of this study is to investigate (1) whether the shells record high-frequency variability in shell chemistry that can be linked to environmental and/or circadian rhythms and (2) 137 whether the presence of photosymbiosis influences the expression of this variability in the shells' 138 composition. 139

#### 141 **2. Materials and methods**

#### 142 2.1 Preparation of *P. maximus* specimens

Three specimens of the King scallop P. maximus (labeled "PM2", "PM3" and "PM4") were 143 collected alive on 15/11/2019 on the southern coast of the Bay of Brest near Lanvéoc, France 144 145 (48°17'N 4°30'W) by SCUBA divers at a depth of approximately 8 m (see\_(Fröhlich et al., 2022) 146 Fröhlich et al., 2022; Figure 1). Note that water depth in the Bay of Brest varies significantly due 147 to the macrotidal regime with a mean tidal range of 2.8 – 5.9 m with extreme ranges up to 7.2 m (Guillaume Olivier et al., 2021; Service Hydrographique et Océanographique de la Marine -148 149 Géoportail, 2022) (Guillaume-Olivier et al., 2021; Service Hydrographique et Océanographique 150 de la Marine; 2022). Collected specimens contained at least one full year of growth based on the 151 visibility of one winter growth line on the outside of the shell (age class 1; see (Thébault et al., 152 2022) Thébault et al., 2022-; Fig 1F and S1). Specimens were frozen at -20°C immediately after 153 collection. Soft body parts and epibionts were removed from the shells before further treatment. 154 Shells were superficially cleaned using a plastic brush and adhering sediment was removed by ultrasonication in deionized water. The flat, left valves were used for elemental and 155 sclerochronological analysis following previous studies on P. maximus (Thébault et al., 2022; 156 Fröhlich et al., 2022) (Thébault et al., 2022; Fröhlich et al., 2022). 157

158 High-resolution color photos were made of the outside of the left valve of the shell using a mirrorreflex camera (Canon EOS 600 DSLR camera connected to a Wild Heerbrugg binocular 159 160 microscope equipped with a Schott VisiLED MC 1000 sectoral dark-field light source) aimed downward perpendicular to the working surface. Overlapping images of the shells were stitched 161 162 together using Image Composite Editor v2.0.3.0 (Microsoft Research Computational Photography Group, Redmond, WA, USA). The stitched images were used to count and measure daily striae 163 on the shell surface (see Fig. 1 and S1). To obtain a fully focused composite of the complete 164 165 shell, dynamic focusing was applied to allow all parts of the slightly curved surface of the shell to

come into focus. Dynamic focus images where later stitched together using focus stacking in
 Helicon Focus (Helicon Focus 7.7.5; HeliconSoft, Kharkiv, Ukraine; see S1).

168 Cross sections were cut through all three P. maximus shells perpendicular to the daily growth 169 lines (striae) from the ventral margin of the shell to the shell hinge (see Fig. 1B-, Fig. 1E-FD and 170 **S1**) along the axis of maximal growth. Shells were fortified with a protective layer of metal epoxy 171 (Gluetec Wiko Epofix 05) before sectioning using a Buehler Isomet 1000 low-speed precision saw (Buehler Inc, Lake Bluff, IL, USA) equipped with a diamond-coated wafering thin blade (0.4 mm 172 173 thickness; number 15LC 11–4255) at 200 rpm. Parallel cuts were made to allow shell sections to 174 be glued to glass plates for grinding high-grade polishing (down to F1200 grit SiC powder) and 175 high-grade polishing and (1 µm Al<sub>2</sub>O<sub>3</sub> suspension on Buhler polishing cloth, MasterTex). Two cross sections were made through specimens PM2 and PM3: One through a "rib" of the shell (i.e., 176 177 radial segment that protrudes away from the interior, named PM2\_1 and PM3\_1) and one through 178 a "valley" (i.e., radial segment between two "ribs" that lies deeper towards the interior, named PM2\_2 and PM3\_2; see Fig. 1 and S1). The dual sections were cut to compare shell chemistry 179 180 between the "ribs" and "valleys" of the shell-. Specimen PM4 was only sectioned once, through a "valley" in the shell, making a total of five cross sections through the *P. maximus* specimens. 181



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Figure 1: Overview of sample locations and preparation steps. A) Location of the Bay of 184 Brest, with the red star indicating the sampling location. B) Schematic cross section through P. 185 maximus showing how the LA-ICP-MS line scan (red line) was positioned within the outer shell 186 187 layer (OSL). EC) Schematic representation of a segment through the shell of P. maximus showing the striae which are deposited daily and which were counted to establish age models (see also 188 B). DF) Left valve of P. maximus (PM2) used in this study, with dashed lines showing the position 189 of cross sections through ribs (red) and valleys (blue) in the shell. Black arrow indicates growth 190 191 direction away from the shell hinge. The black dotted line highlights a winter growth stop. CE)

192 Schematic cross section through a tridacnid, illustrating the positions of parrallel parallel LA-ICP-193 MS line scans (red lines) through these shells within the OSL. **PF**) Position of the Gulf of Agaba, 194 with the red star indicating the sample location for tridacnids. E) Schematic representation of a 195 segment through the shell of P. maximus showing the striae which are deposited daily and which 196 were counted to establish age models (see also B). F) Left valve of P. maximus (PM2) used in 197 this study, with dashed lines showing the position of cross sections through ribs (red) and valleys 198 (blue) in the shell. Black arrow indicates growth direction away from the shell hinge. The black dotted line highlights a winter growth stop. G-I) Pictures of (from left to right) T. squamosa 199 200 (specimen TSFRS1), T. maxima (specimen TM29) and T. squamosina (specimen SQSA1) with 201 dashed red lines indicating the positions of the cross sections used for LA-ICP-MS analysis (see 202 C) and black arrows indicating the direction of growth. KJ) Example of Mutvei-stained cross 203 section through a T. maxima specimen used to visualize and count growth lines, with the insert 204 showing part of the OSL where growth lines were counted (red lines) to establish age models for 205 the tridacnids. Black arrows indicate the direction of growth. JK) Schematic representation of the 206 LA-ICP-MS line scanning setup with the rectangular spot size (100 \* 20 µm; see S11) that was 207 positioned parallel to the growth layers in the shell. K) Example of Mutvei-stained cross section through a T. maxima specimen used to visualize and count growth lines, with the insert showing 208 209 part of the OSL where growth lines were counted (red lines) to establish age models for the 210 tridacnids. Black arrows indicate the direction of growth.

#### 212 2.2 Preparation of *Tridacna* specimens

A total of 5 tridacnid specimens, two *T. maxima* (named **TM29** and **TM84**), two *T. squamosa* (named **TS85** and **TSFRS1**) and one *T. squamosina* (**SQSA1**) specimen, were collected in the summer of 2016 from beach death assemblages on the coast of the Gulf of Aqaba with permit from the Israeli National Parks Authority (**Figure 1**; see details in\_(Killam et al., 2020)-Killam et al., 2020). One cultured *Tridacna squamosa* shell (**TSM1**) was obtained from the National Center for Mariculture, Eilat. Species were determined following shell characteristics of the local population as cited in (Roa-Quiaoit, 2005)-Roa-Quiaoit (2005).

220 All shells were sectioned along the axis of maximum growth after removing epibionts using a 221 metal brush (see Fig. 1G-I). Original microstructure and preservation of the original aragonite 222 mineralogy of all specimens was confirmed using Scanning Electron Microscopy and X-ray 223 Diffraction Spectroscopy following (Gannon et al., 2017) Gannon et al. (2017) and (Kontoyannis 224 and Vagenas, 2000) Kontovannis and Vagenas (2000; (see details in (Killam et al., 2020)). Shell 225 segments were partially embedded in Araldite 2020 epoxy resin (Huntsman Corp., Woodlands, TX, USA) before being sectioned in direction of maximum growth using a slow-rotating saw 226 equipped with a thin wafered saw blade ( $\alpha < 1$  mm). Parallel cross sections produced 5-10 mm 227 thick sections that were polished using progressively finer SiC polishing disks. 228

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## 230 2.3 Microscopy and photography

Polished surfaces of all 11 cross sections (5 *Pecten*, 6 *Tridacna*) were imaged using an Epson®
1850 flatbed scanner (Seiko Epson Corp., Nagano, Japan) at a pixel resolution of 6400 dpi (±4
µm pixel size) as well as by stitching micrographs made using a KEYENCE VHX-5000 digital
microscope using x250 magnification together into composite images (see **S1**). Cross sections

were imaged both before and after trace element analyses to allow the trace element profiles tobe referenced relative to the cross sections.

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238 2.4 LA-ICP-MS analyses

Elemental ratios were based on measuring ratios of the isotopes <sup>25</sup>Mg, <sup>87</sup>Sr, <sup>55</sup>Mn and <sup>137</sup>Ba to <sup>43</sup>Ca along profiles through all shell cross sections using Laser Ablation – Inductively Coupled Plasma – Mass Spectrometry (LA-ICP-MS). Measurements were carried out on a laser ablation system (ESI NWR193UC; Elemental Scientific, Omaha, NE, USA) coupled to a quadrupole ICP-MS (iCap-Q, Thermo Fisher Scientific, Waltham, MA, USA) at the Royal Netherlands Institute for Sea Research (NIOZ). Operation parameters are provided in **S11**.<sup>-</sup>

247 Scan lines were programmed on the polished shell cross--sections in direction of growth as close 248 as possible to the outer edge of the shell, with the circular LA-ICP-MS spot oriented parallel to the growth lines (with a width of 20 µm in scanning direction, see Fig. 1J; S11). For the pectinids, 249 250 care was taken to target the outer portion of the outermost shell layer (oOSL) and avoid sampling 251 of the inner portion of the outer shell layer (iOSL) or inner shell layer (ISL), which was demonstrated to have a different chemical composition (see Freitas et al., 2009). For the 252 253 tridacnids, profiles were placed within the OSL close to (within 100 µm of) the outer edge of the shell in a first analytical session. However, since spikes of high Mg/Ca and Mn/Ca ratios were 254 255 observed in these results, parallel transects placed ~\_100 µm further towards the inside of the 256 shell were measured through all tridacnid shells to verify whether these spikes in Mg and Mn were 257 reproducible further inward (see S2). The combination of laser scan speed (4 µm \* s<sup>-1</sup>) and ICP-258 MS run cycle time ("sweep time"; 109 ms) yielded a mean spatial sampling resolution of 0.44 µm, determined as the arithmetic mean without taking into account the round shape of the laser spot. 259 However, note that the width of the LA-ICP-MS spot size in scanning direction (20 µm) caused 260 smoothing of the LA-ICP-MS signal, reducing the effective spatial sampling resolution. As a result, 261 262 each 0.44 µm wide segment of the shells in growth direction is sampled 45 times while the spot 263 moves across the shell. We tested the effect of this smoothing on a virtual record with characteristics similar to the Sr/Ca record obtained by (Sano et al., 2012) (see section 3.3 and 264 265 S13). All scan lines in pectinids and tridacnids were repeated a second time at the exact same location using a faster scan rate of 10  $\mu$ m<sub>-</sub> \* s<sup>-1</sup> to assess repeatability of the elemental signals 266 267 (see **S2**).

Data reduction was performed using an adapted version of the data reduction software SILLS
 (Signal Integration for Laboratory Laser Systems; (Guillong et al., 2008)) in Matlab. Raw LA-ICP MS data were calibrated using NIST610, (National Institute of Standards and Technologies,

Gaithersburg, MD, USA) using the reference values reported in the GeoReM database ((Jochum et al., 2005, 2011)). Quality control <u>reference</u> materials BAS<u>CRM</u> 752-393 (Bureau of Analyzed Samples, Middlesbrough, UK), RS3 and one matrix-matched carbonate standard (MACS-3; United States Geological Survey, Reston, VA, USA; (Wilson et al., 2008)) were used to monitor
the quality of the measurement. <u>Details on the accuracy of the LA-ICP-MS trace element results</u>
<u>relative to preferred values for the check standards is provided in S12</u>.

To increase the stability of the ICP-MS signal and to correct for drift, <sup>43</sup>Ca was used as internal standard. <u>Calcium concentrations were assumed to be constant (38 wt%) throughout the</u> <u>LAICPMS profiles within the same shell layer</u>. External drift-correction using repeated measurements on the JCp1 standard was applied. if tThe element/Ca drift was <u>less than</u> >5\_% during the analytical sequence. Drift during a single transect was found to be negligible.

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## 283 2.5 Age models

284 Trace element profiles in *P. maximus* shells were internally dated using daily striae visible on the 285 outer shell surface (Fig. 1=1D). Daily increment widths (perpendicular distances between 286 successive striae) were counted and measured multiple times, both on the outside of the shell using the focus-stacked images (see section 2.3) and by counting and measuring the distance 287 between growth layers in cross sections through the "valleys" of the shells (PM2 2 and PM3 2; 288 see S3) by different persons. Positions of daily striae on the outside of the shells were plotted 289 relative to distance along the LA-ICP-MS scan line using manual alignment of striae and the LA-290 291 ICP-MS path on microscope composites of cross sections through the shells, taking into account 292 the curvature of growth lines with distance away from the outer shell surface (see S3). The timing 293 of shell formation was determined by backdating the daily striae from the ventral margin (last visible stria mineralized on the date of shell collection, i.e., November 15, 2019), and by linearly 294

interpolating the timing of measurements located between daily growth lines based on theirdistance from daily striae positions (**S5**).

Trace element profiles from Tridacna shells were also dated using layer counting. Polished cross 297 298 sections through all tridacnids were imaged using UV luminescence (see Fig. 1J and S4) and 299 stained with Mutvei solution (Schöne et al., 2005c; Killam et al., 2021) to facilitate this counting. 300 However, despite staining and luminescence techniques, the expression of daily and semi-diurnal 301 (half-daily, or 12 h rhythmic) growth markings wereas insufficiently clear to count individual growth 302 lines along the full (multi-year) growth period recorded in all the shells. Therefore, we opted for a 303 hybrid method in which we measured the width of daily and semi-diurnal increments in parts of cross sections of each shell where they were well developed and used these measurements in 304 305 combination with the annual growth breaks (which are easily recognized on the shell) to create 306 age models for all specimens. age models were constructed based on parts of the shell where 307 daily and tidal layers could be identified with confidence. The median widths of daily or semi-308 diurnal increments were determined on these cross sections and compared to the width of annual increments identified based on growth breaks visible on the outer margin of the shell. The 309 distinction between diurnal (24 h) and tidal (~12 h) pacing of growth increments was made based 310 on the width of small-scale increments relative to the width of annual increments in the shell. A 311 312 von Bertalanffy growth model (Von Bertalanffy, 1957) was constructed for each specimen based 313 on the annual growth ( $\Delta L$ ) inferred from (semi-)diurnal growth line counting and the maximum 314 shell height (Linf) known for these species in the Red Sea from the literature (Roa-Quiaoit, 2005; 315 Mohammed et al., 2019):

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$$L_t = L_{inf} * \left(1 - e^{-kt}\right)$$
, with  $k = -ln\left(rac{\Delta L}{L_{inf}}
ight)$ 

In this formula, Lt is the shell height at time t and k is the growth constant (Brody growth coefficient;
(Munro, 1982)). Since cross sections through the tridacnids were made through the shell hinge

319 (in direction of the shell height) and literature values for L<sub>inf</sub> are reported with reference to shell 320 length (measured parallel to the shell hinge), allometric data on T. maxima, T. squamosa and T. squamosina from the literature was used to convert Linf values (which are commonly reported as 321 shell length) to shell height and make them relevant for the direction in which the trace element 322 323 profiles were measured on the cross sections (Roa-Quiaoit, 2005; Richter et al., 2008; 324 Mohammed et al., 2019). Uncertainties on the annual growth increment widths ( $\Delta L$ ) were calculated from the standard error of the mean width of daily and semi-diurnal growth increments 325 on which  $\Delta L$  is based, and uncertainties on the values for L<sub>inf</sub> were taken from variability in the 326 327 values in the literature. Both sources of uncertainty were propagated through the growth model using the variance formula (Ku, 1966) to obtain error envelopes on age-distance relationships 328 (growth curves) of tridacnids (see S5). All data processing steps described in this manuscript are 329 330 carried out using the open-source computational software package R (R Core Team, 2013), and 331 scripts detailing these calculations are provided in S6 and deposited on the open-access software repository GitHub (https://zenodo.org/record/6603175) 332

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334 2.6 Spectral analysis

335 Spectral analysis on the LA-ICP-MS data was used to isolate trace element variability at the sub-336 annual scale. All trace element profiles were first detrended using a LOESS filter with a span of 0.2 times the length of the record to remove longer term (i.e., seasonal to multi-annual) trends. 337 The detrended series were linearly resampled in the time domain before applying the Multi-Taper 338 Method (MTM; (Thomson, 1982)) to extract dominant frequencies from the data. Spectral analysis 339 340 was carried out using the "astrochron" package (Meyers, 2014) in R ((R Core Team, 2013); see script in S6). The significance of relevant periodicities was tested using a combination of "red 341 342 noise" estimation and a harmonic F-test (see (Meyers, 2012)). To visualize the evolution of periodic behavior across the shells, wavelet analysis was applied on all trace element profiles
using the "dpIR" package in R (see S6).

345

346 2.7 Extracting high-resolution variability

347 After detrending and spectral analysis, all trace element profiles were smoothed using a Savitzky-348 Golay filter with a width of 21 datapoints (8.4  $\mu$ m; equivalent to a timespan of ~1-5 h; S6) to remove high-frequency measurement noise within our LA-ICP-MS spot size. The Savitzky-Golay 349 350 filter was used because it preserves the tendency of the trace element profile and therefore retains 351 a larger fraction of the amplitude of periodicity close to the width of the filter compared to simpler 352 smoothing techniques (e.g. moving average; (Savitzky and Golay, 1964)). The 21-point window 353 size was chosen to smooth out the maximum amount of instrumental noise without losing periodicity larger than half the width of the laser spot (0.22 µm; equivalent to 23 datapoints). 354 355 Statistically significant (see section 2.6) variability in daily (~22-36 h; centered on the 24 h diurnal 356 cycle) and tidal-semi-diurnal (~8-14 h; centered on the 12.4 h tidal cycle) frequency bands was extracted from the trace element records using a combination of bandpass filtering (using the 357 "bandpass" function in the "astrochron" R package) and stacking (see S6). Trace element data 358 359 was stacked along bandpass filters using the following procedure: Maxima and minima in the 360 bandpass filter were used as tie points to reference each datapoint of the smoothed dataset relative to its position within the cycle on a scale from 0 to 1. These relative positions were then 361 used to divide the data into 10 bins (bin 1 contains positions 0 - 0.1, bin 2 contains data from 362 363 positions 0.1 - 0.2, etc.), giving the providing stacked data a resolution of 0.1 times the length of 364 the cycle under investigation. The full breakdown of variability within and between bins created in the stacking routine is provided in S7. Different sources of variance in the trace element records 365 were isolated by sequentially determining the variance left in the trace element records after each 366 367 of the data treatment steps explained above (see example in S7). This procedure allowed us to

- 368 quantify the amount of variance in each trace element profile explained by either diurnal or semi-
- 369 diurnal variability.

370 3. Results

## 371 3.1 Trace element data

372 LA-ICP-MS line scans yielded profiles of Sr/Ca, Mg/Ca, Mn/Ca and Ba/Ca in growth direction on 11 cross--sections through shells of P. maximus, T. maxima, T. squamosa and T. squamosina 373 374 (Fig. 2; see S0 for raw data of all scans). Trace element profiles of consecutive line scans on the same transect show high repeatability: sSub-millimeter scale patterns in Sr/Ca, Mg/Ca, Mn/Ca, 375 376 and Ba/Ca are highly repeated repeatable reproducible between consecutive line scans on the same LA-ICP-MS transect:, R<sup>2</sup> values between trace element results of time-equivalent shell 377 samples typically exceed 0.8, and the mean element ratio differences between time equivalent 378 379 samples in different line scans are typically <0.05 mmol/mol for Mg/Ca (<5 % relative to mean value), <0.02 mmol/mol for Sr/Ca (<2 % relative to mean value), <0.5 µmol/mol for Mn/Ca (<10 380 381 % relative to mean value) and <0.2 µmol/mol for Ba/Ca (<8 % relative to mean value; is less than 0.05 mmol/mol for the most variable profiles (Mg/Ca, with lower differences for the lower-382 concentration Mn/Ca and Ba/Ca records; see S2). Remeasured transects further away from the 383 outer shell surface in tridacnids (see section 2.4) differ more from the original transects than 384 those measured on the exact same locality in the shell: R<sup>2</sup> values between parallel lines in different 385 386 localities are 0.3 – 0.5 for Mg/Ca and Sr/Ca and <0.3 for Mn/Ca and Ba/Ca, reflecting intra-shell 387 variability in trace element composition in the tridacnids (S2). Overall, sub-millimeter scale patterns in trace element composition are reproduced in parallel line scans, and the mean offset 388 between the lines was always less than 0.2 mmol/mol. 389

Pectinid and tridacnid shells contain similar mean Sr/Ca and Ba/Ca ratios (Sr/Ca of  $1.3 \pm 0.3$  and  $1.5 \pm 0.6$  mmol/mol respectively; Ba/Ca of  $2.8 \pm 2.5$  and  $3.0 \pm 5.1$  µmol/mol respectively; uncertainty is calculated as  $1\sigma$ ). Mean Mg/Ca and Mn/Ca ratios are higher in *P. maximus* than in *Tridacna* species (Mg/Ca =  $3.1 \pm 0.9$  and  $0.7 \pm 0.9$  mmol/mol; Mn/Ca =  $7.8 \pm 4.7$  and  $2.7 \pm 7.8$ µmol/mol;  $1\sigma$ ; **Figure 2**; **S4**). Differences between tridacnid specimens generally exceed the 395 differences between tridacnids and pectinids (1o of Ba/Ca among all tridacnid specimens = 2.1 396 µmol/mol). Individual records like those in TM84 and PM3\_1 show large variability (especially in Ba/Ca and Mn/Ca) compared to other specimens of the same species. Inter-specimen variability 397 398 is higher in tridacnid shells than in pectinids (inter-specimen relative standard deviations as a 399 fraction of mean ratio for Ba/Ca: 0.74 vs 0.64, Mg/Ca: 0.37 vs 0.20, Sr/Ca: 0.19 vs 0.03 and Mn/Ca: 0.78 vs 0.33 for tridacnids and pectinids, respectively). Figure 2 shows that this variability 400 between tridacnids is not readily explained by differences between species, but mostly reflects 401 differences in the trends within the records, with some specimens (e.g., TM84, TSM1 and TS85) 402 403 showing trends in composition towards the end of the record (see also S8). Trace element compositions in tridacnid shells are significantly more skewed towards higher values than in 404 405 pectinids (mean skewness per element and per specimen is 9.7 for tridacnids and 0.9 for 406 pectinids), reflecting the high peaks in trace element composition observed in tridacnid profiles, 407 especially near the ventral margin (e.g., specimens TM84, TSM1 and TS85; see section 2.4; Fig. 408 2; S8). Finally, "rib" and "valley" segments through the same specimen of *P. maximus* show similar 409 patterns in trace elements, but absolute concentrationselement-to-calciumal ratios (especially of 410 Ba/Ca and Mn/Ca) can be quite different, highlighting heterogeneity within the shells of P. 411 maximus (Fig. 2).

412 Plots of trace element variability reveal dominant high-frequency variability superimposed on seasonal-scale patterns (Figure 2). Trace element profiles in pectinids, reflecting only one 413 414 growing season, show a typical seasonal pattern in Sr/Ca and Mg/Ca with maxima in the 415 elemental ratio in the middle of the profile (corresponding to the summer)., while Mn/Ca and Ba/Ca 416 in pectinids are more variable, showing multiple peaks in the same growth year. Peaks in Mn/Ca 417 and Ba/Ca are synchronous between profiles through the same specimen, but not between 418 specimens, possibly showing that growth resumed on different days for different specimens after 419 the winter stop. Like in the pectinid profiles, Tridacnid Mg/Ca, and Sr/Ca ratios in tridacnids show similar patterns, with one or two distinct cycles per growth year. <u>Mn/Ca and Ba/Ca ratios in</u>
 tridacnids show more regular annual or biannual variability than pectinids (most notably specimen
 **SQSA1**). It must be noted, however, that *P. maximus* shells only recorded one growth season.
 limiting the interpretation of seasonal growth patterns.

424 Plots of trace element variability zoomed in on smaller sections of the record reveal dominant 425 high-frequency variability superimposed on seasonal-scale patterns (Figure 3). Both tridacnid 426 and pectinid LA-ICP-MS records show distinct high-frequency cyclicity in all analyzed studied trace element-to-Ca ratios with amplitudes on nearly the same order of magnitude as the longer-term 427 428 variability described by the records on a seasonal to multi-annual scale (Fig. 2). However, higher frequency variability in tridacnid ratio profiles is characterized by more extreme peaks, especially 429 in Mg/Ca, skewing the distribution of trace element values. Mn/Ca and Ba/Ca appear to be less 430 431 variable in tridacnid shells than in pectinids, except for specimen TM84, which shows a sharp 432 increase in Mn and Ba towards the end of its lifetime. Mn/Ca and Ba/Ca ratios in tridacnids show more regular annual or biannual variability than pectinids (most notably specimen SQSA1). It 433 must be noted, however, that P. maximus shells only recorded one growth season, limiting the 434 interpretation of seasonal growth patterns. 435



Figure 2: Overview of LA-ICP-MS results of Sr/Ca (blue), Mg/Ca (green), Mn/Ca (purple) and Ba/Ca (orange) in pectinid (upper panel) and tridacnid (lower panel) specimens. Vertical axes are equal for plots positioned next to each other (but different for the two groups of tridacnid and pectinid plots). Shaded lines show raw LA-ICP-MS data while solid lines indicate 0.2 span LOESS fits through the data highlighting monthly-scale variability. A direct comparison of trace elemental ratios between specimens is provided in **S8**.



## 450 3.2 Age models

451 Growth line counting in the *P. maximus* shells was repeated multiple times on both the outer shell surface and in cross sections through the shell by different persons (Table 1; S3). The variability 452 453 in counting results shows that the growth lines were not always equally easy to distinguish. In 454 PM2 and PM3, the most likely number of increments (228 and 220 respectively) was counted 455 confirmed in both cross sections and on the outside of the shell, with other analyses yielding both 456 higher and lower numbers. In PM4, counts on the outside and on the one available cross section 457 were very close (Table 1). In this case, the counting in the cross section (278 increments) was 458 chosen over the slightly lower count on the outside of the shell (272-273 increments) as reference since the LA-ICP-MS profile was measured on the same cross section and could be directly linked 459 460 to the counted increments. Note that, while the consistency in daily counts between specimens from the same year and environment is encouraging, inter-specimen variability in the onset of 461 462 spring growth or the onset of the winter stop (potentially prior to the collection date of November 463 15<sup>th</sup>, 2019) places some uncertainty on the absolute date assigned to *P. maximus* shell sections 464 (Chauvaud et al., 2005). The fact that the mean increment width between the P. maximus specimens which grew in the same year in the same environment is highly consistent lends 465 466 confidence to the layer counting result (Table 1). The difference in number of days of growth between specimens can be caused either by variability in the day on which seasonal growth 467 commenced (in spring) or the day on which the winter growth stop commenced (in autumn; 468 469 Chauvaud et al., 1998). The sampling date (November 15<sup>th</sup>, 2019) does not preclude the onset of 470 winter growth cessation before the time of sampling. The age-distance relationships (growth curves) resulting from the sclerochronology are shown in S5. 471

# **Table 1: Growth increments counting in** *P. maximus*

Specimen	Increments counted	Increments counted in	Mean <u>daily</u> increment
	on outer surface	cross sections	width [μm ± 1σ]
PM2	226, <b>228</b> , 234, 241	227, <b>228</b> , 233	249 ± 19
PM3	<b>220</b> , 226, 243	213, <b>220</b> , <b>220</b>	249 ± 22
PM4	272, 273	278	247 ± 4

Layer counting in tridacnid shells yielded estimates of semi-diurnal, daily and annual growth
(Table 2; S4). Annual growth rates calculated from layer counting are highly consistent between
specimens from the same species from the same environment (within 1 mm/yr; Table 2), lending
confidence to the growth line counting results. The von Bertalanffy growth models based on these
growth line countings are plotted in S5. Statistics of the parameters (L<sub>inf</sub> and k) of these growth
models and their uncertainty are provided in S4.

Table 2: Growth line counting in Tridacna shells. Column 3 shows the total number of 483 484 increments counted in the specimen, column 4 shows their median width and column 5 shows the width of an annual increment in the specimen. Note that increments could not be counted over 485 486 the entire growth period of the shells, so the numbers in column 3 represent representative 487 numbers of increments counted in those parts of the shells where they were distinct enough for 488 counting (see section 2.5). Increment timing (semi-diurnal vs diurnal) was established based on 489 the relative difference between small-increment width and annual increment widththe distance between annual growth breaks measured on the outside of the shells. 490

Specimen	Species	# counted	Median	Annual	Increment
		increments	increment	growth	timing
			width [µm]	[mm]	
TM29	T. maxima	274	26.5	27.9	Semi-diurnal
TM84	T. maxima	109	39.1	26.6	Diurnal
TS85	T. squamosa	310	40.3	20.2	Diurnal
TSFRS1	T. squamosa	225	23.3	20.1	Semi-diurnal
TSM1	T. squamosa	180	33.3	20.6	Diurnal
SQSA1	T. squamosina	153	22.3	14.9	Diurnal

491

## 493 <u>3.3 Growth rates and temporal sampling resolution</u>

Growth rates are highly similar between specimens of the same species (Table 1 and Table 2; 494 495 **S3-5**), with *P. maximus* achieving the highest growth rates (~\_220 growth days \* ~\_250  $\mu$ m/d ≈ 55 mm/yr; Table 1), followed by T. maxima (~27 mm/yr; Table 2), T. squamosa (~20 mm/yr; Table 496 497 2) and T. squamosina (15 mm/yr; Table 2). The-Dividing the width of the daily increments (e.g., 498 250 µm for P. maximus) by spatial LA-ICP-MS sampling resolution of 0.44 µm (see section 2.4), 499 these growth rates would seem to translate age models reveal that the average to a temporal 500 resolution of the LA-ICP-MS line scans was 0.04 h, 0.24 h, 0.44 h and 0.27 h for P. maximus, T. 501 maxima, T. squamosa and T. squamosina, respectively. These estimates were calculated by dividing the width of the daily increments (e.g., 250 µm for P. maximus) by the resolution of the 502 LAICPMS data (0.4 µm) to obtain the number of LAICPMS measurements per day (e.g., 625 503 504 pts/day for *P. maximus*, yielding a mean sampling resolution of 0.04h). Note that However, the our 505 LA-ICP-MS slit-spot size is wider (20 µm in sampling direction) than the spatial sample resolution, causing causes some significant smoothing, reducing the frequency that can be resolved on the 506 scale of this very fine temporal resolution. Nevertheless, our test applying the LA-ICP-MS settings 507 on a virtual Sr/Ca dataset with the characteristics of the diurnal record by (Sano et al., 2012) and 508 509 growth rates equivalent to our specimens (see S13) reveals that our analytical methodology can 510 easily resolve daily variability even in our slowest-growing specimen (T. sSquamosina SQSA1), conserving at least 75 % of the signal amplitude. The LA-ICP-MS profiles record trace element 511 512 variability during growth periods ranging between 220 days (for **PM3**) and 1041 days (for **SQSA1**).

513

#### 514 3.3 <u>3.4</u> Spectral analysis

515 Normalized power spectra and significance level of daily and tidal periodicities in pectinid and 516 tridacnid records are shown in **Figure 3**-<u>4</u> and **Figure 45**, respectively. Full spectral analysis

517 results for all trace element records in all specimens are provided in S9. All P. maximus power spectra (Fig. 43) reveal semi-diurnal (12\_h) periodicity in Sr/Ca and Ba/Ca with >\_86\_% statistical 518 significance. Only sections through the ribs of the shells (PM2\_1 and PM3\_1) show semi-diurnal 519 520 periodicity in Mg/Ca and Mn/Ca (> 90 % significance). Daily periodicity is present in some, but not 521 all, pectinid profiles, but there seems to be no consistent pattern in the presence of diurnal 522 variability between specimens, between sections through ribs or valleys in the shell or between 523 trace element records. Most power spectra of trace element profiles in pMost pectinid trace element recordss also contain -show peaks associated with multi-day tidal-periodicities, the most 524 525 dominant being a 7-day period, with weaker expression of cyclicity associated with the fortnightly 526 (14d) cycle or lunar month (28d) including strong 7-day periodicities and weaker cyclicity 527 associated with 14-day and 28-day cycles. The latter is partly suppressed by the 0.2 span LOESS 528 filter (equivalent to a 44-56 day period depending on the length of the record) applied on the 529 records to remove the seasonal trend from the records. However, these lower frequency cycles 530 are clearly visible in the wavelets (see S9).

A much more consistent expression of diurnal periodicity is found in the t<u>T</u>ridacnid trace element profiles compared to those inshow more consistent daily periodicity than the pectinid recordss (**Fig. <u>54</u>**). Especially Sr/Ca and Ba/Ca records through nearly all tridacnid specimens exhibit strong (>\_90\_% confidence level) power in the daily period, while <u>daily cyclicity is weaker in</u> Mn/Ca and Mg/Ca records <u>exhibit much less periodicity</u>. Sr/Ca records in the tridacnids also contain a significant (>\_96\_%) semi-diurnal component, whose tidal origin seems clear in most specimens by peaks in power in the longer (7\_d, 14\_d and 28\_d) tidal components.



**Figure 34**: Multi-taper method spectrograms of Sr/Ca (**A**), Mg/Ca (**B**), Mn/Ca (**C**) and Ba/Ca (**D**) records through the five pectinid cross sections after detrending (see **section 2.6**). All spectra are normalized by dividing by the highest power peak and plotted on the same horizontal axis. Grey shaded lines show raw data while solid black lines plot 21-point moving average smoothed curves. Red vertical dashed lines highlight the expected periods of tidal variability while black vertical dashed lines indicate 1-day periodicities. Significance levels of peaks on these periods (see **section 2.6** and (Meyers, 2012)) are rounded to the nearest whole percentage point.



**Figure 45**: Multi-taper method spectrograms of Sr/Ca (**A**), Mg/Ca (**B**), Mn/Ca (**C**) and Ba/Ca (**D**) records through the six tridacnid cross sections after detrending (see **section 2.6**). All spectra are normalized by dividing by the highest power peak and plotted on the same horizontal axis. Grey shaded lines show raw data while solid black lines plot 21-point moving average smoothed curves. Red vertical dashed lines highlight the expected periods of tidal variability while black vertical dashed lines indicate 1-day periodicities. Significance levels of peaks on these periods (see **section 2.6** and (Meyers, 2012)) are rounded to the nearest whole percentage point.

# 555 <u>3.4</u> <u>3.5</u> Variance decomposition

556 Variability at the daily (24 h) and tidal-semi-diurnal (12 h) scale in all trace element records through 557 all specimens was extracted using bandpass filtering (section 2.7; see S9 and S10). The median amplitude of variability within these stacks was plotted against the median period of the variability 558 559 per element and per specimen to highlight dominant periodicities in the trace element data 560 (Figure 65). As noted in the spectral analysis results (section 3.34), trace element composition 561 in tridacnid shells is more strongly controlled by daily variability than in pectinid shells (Fig. 56; S10). The difference is especially noticeable in Sr/Ca and Ba/Ca ratios, which show a clear divide 562 563 between daily periodicity in tridacnid shells and tidal periodicity in pectinids (see Fig. 56). The 564 differences in Mg/Ca and Mn/Ca ratios are less clearobyious.



**Figure 56:** Cross plot showing the amplitude of variability of dominant spectral periods in Sr/Ca (A), Mg/Ca (B), Mn/Ca (C) and Ba/Ca (D) against the period (duration) of the cycle. Round symbols indicate the median amplitude of the cycle, while vertical bars and lines show interquartile differences and ranges in the amplitude over the record. Horizontal bars indicate the width of the bandpass filter used to extract periodic variability. Colors highlight different specimens (see legend).

574 An example of the distribution of normalized variability within the trace element records after each 575 data processing step is shown in **S7**. From tThis example it is clearshows that a large fraction of the variance in the records (73 % in this record after trimming outliers) is explained by low-576 577 frequency (seasonal scale) variability (S7). Of the remaining smoothed and detrended dataset, at 578 most 20 % of the variance is explained by daily and tidal (semi-diurnal) periodicity (see Figure 76 579 and Table 3). A full decomposition of variance in all trace element records through all specimens 580 is provided in S7. Figure 65 and Figure 6A-7A shows that, overall, the variance explained by daily periodicity is higher in tridacnids than in pectinids (Wilcoxon signed rank test; W = 44; p = 581 0.009). The difference between species is smaller for tidal variability (Fig. 76B). There is no clear 582 583 difference in relative dominance of tidal-semi-diurnal variability between trace element records, but daily variability is more strongly expressed in Ba/Ca and Mn/Ca records, especially in tridacnid 584 585 shells. Finally, T. squamosa specimen TSM1, which grew under a sunshade, does not exhibit 586 significantly lower daily periodicity compared to the other tridacnid specimens.



**Figure 67**: Summary of relative variance (in %) of significant daily (left) and tidal (right) variability extracted from trace element records. Colors highlight different specimens (see legend). Note that the *T. squamosa* specimen **TSM1** which grew under a sunshade is highlighted with a black arrow.
**Table 3**: Overview of the relative (in %) variance associated with daily and tidal variability in all
trace element records through all specimens. Empty cells represent records for which no
significant tidal or daily periodicity was found (see Fig. <u>4-5</u>3-4).

	Daily variance [% relative to detrended record]				Tidal variance			
					[% relative to detrended record]			
	Ba/Ca	Mg/Ca	Mn/Ca	Sr/Ca	Ba/Ca	Mg/Ca	Mn/Ca	Sr/Ca
PM2_1	3.8 %			1.0 %	6.6 %	2.8 %	5.1 %	3.2 %
PM2_2		0.9 %	4.1 %		4.9 %			6.3 %
PM3_1	1.7 %	1.3 %	3.6 %		4.3 %	2.2 %	3.7 %	2.7 %
PM3_2		2.0 %		1.3 %	4.5 %			2.4 %
PM4	3.1 %		1.0 %	1.7 %	6.2 %			9.5 %
TM29			2.9 %	3.5 %		9.0 %		7.8 %
TM84	3.6 %			2.8 %				7.6 %
TS85	8.1 %	0.9 %		2.7 %			5.5 %	7.8 %
TSFRS1	2.1 %		3.2 %	3.5 %				8.5 %
TSM1	12 %		8.5 %	6.7 %			10 %	20 %
SQSA1	6.9 %	3.8 %	11 %	2.1 %			7.1 %	2.6 %



597

**Figure 78**: Schematic overview of environmental parameters interpreted to affect shell growth and composition of pectinids in the Bay of Brest and tridacnids in the Gulf of Aqaba. The table at the bottom provides a schematic qualitative overview of the amount of variance in the trace element records of the taxa is explained by daily (sun symbol), tidal (moon symbol) or aperiodic (cloud symbol) variability in the environment.

# 604 4. Discussion

# 605 <u>4.1 Compositional differences between pectinids and tridacnids</u>

- 606 Pectinid and tridacnid shells are characterized by similar mean Sr/Ca and Ba/Ca ratios (Sr/Ca of
- 607 <u>1.3 ± 0.3 and 1.5 ± 0.6 mmol/mol respectively; Ba/Ca of 2.8 ± 2.5 and 3.0 ± 5.1 µmol/mol</u>
- 608 respectively; uncertainty is calculated as 1 σ). Mean Mg/Ca and Mn/Ca ratios are higher in P.
- 609 maximus than in Tridacna species (Mg/Ca =  $3.1 \pm 0.9$  and  $0.7 \pm 0.9$  mmol/mol; Mn/Ca =  $7.8 \pm 4.7$
- 610 and 2.7  $\pm$  7.8  $\mu$ mol/mol; 1  $\sigma$ ; Figure 2; S4). The incorporation of higher concentrations of the
- 611 elements Mg and Mn, which have a lower ionic radius than Sr and Ba, in calcitic pectinid shells
- 612 <u>compared to aragonitic tridacnid shells makes sense given the higher partition coefficient of these</u>
- 613 elements in calcite than in aragonite (Day and Henderson, 2013; Wassenburg et al., 2016).
- 614 Differences between tridacnid specimens generally exceed the differences between tridacnids and pectinids (10 of Ba/Ca among all tridacnid specimens = 2.1 µmol/mol), suggesting large inter-615 616 species differences. However, individual records such as those in TM84 and PM3\_1 show large variability (especially in Ba/Ca and Mn/Ca) compared to other specimens of the same species. 617 618 Inter-specimen variability is higher in tridacnid shells than in pectinids (inter-specimen relative standard deviations as a fraction of mean ratio for Ba/Ca: 0.74 vs 0.64 µmol/mol, Mg/Ca: 0.37 vs 619 0.20 mmol/mol, Sr/Ca: 0.19 vs 0.03 mmol/mol and Mn/Ca: 0.78 vs 0.33 µmol/mol for tridacnids 620 621 and pectinids, respectively). Figure 2 shows that this variability between tridacnids is not readily 622 explained by differences between species, but mostly reflects differences in the trends within the 623 records, with some specimens (e.g., TM84, TSM1 and TS85) showing trends in composition towards the end of the record (see also S8). It thus seems that inter-specimen differences, 624 perhaps driven by growth stress, are dominant drivers of trace element variability in giant clams. 625 626 This interpretation is corroborated by the observation that trace element compositions in tridacnid shells are significantly more skewed towards higher values than in pectinids (mean skewness per 627 628 element and per specimen is 9.7 for tridacnids and 0.9 for pectinids). This skewness reflects the

high peaks in trace element composition observed in tridacnid profiles, especially near the ventral
 margin (e.g., specimens TM84, TSM1 and TS85; see section 2.4; Fig. 2; S8), which hint towards
 growth disturbances at the individual level.

632 4.24 Trace element variability in *P. maximus* 

633 4.<u>42</u>.1 Comparison with previous studies

634 Trace element concentrations in P. Maximus maximus analyzed in this study are in close 635 agreement with concentrations found in wild (live collected) pectinid shells in the literature (Lorrain et al., 2005; Barats et al., 2008; Poitevin et al., 2020; Fröhlich et al., 2022) (Lorrain et al., 2005; 636 Barats et al., 2008; Poitevin et al., 2020; Fröhlich et al., 2022). In these studies, Sr/Ca shows a 637 638 strong link with calcification rate (as measured by the width of daily shell increments; (Lorrain et 639 al., 2005) Lorrain et al., 2005), although previous studies did not assess variability on the (sub-) daily scale. The long-term trends in our Sr/Ca records seem to confirm this correlation, with 640 641 higher values being recorded in the middle of the growing season (day 50-150; Fig. 2) when 642 growth rates are highest (see S5). Previous studies demonstrate that Mg/Ca ratios in pectinid shells are at most partially related to temperature and/or salinity (Lorrain et al., 2005; 643 644 Poitevin et al., 2020). The fact that the studied P. maximus specimens, which all grew during 645 the same year in the same environment, do not show a synchronous Mg/Ca pattern (Fig. 2) corroborates this previous work and argues against a straightforward temperature 646 647 dependence for Mg/Ca in P. maximusThere is some discussion on the dependence of Mg/Ca ratios in pectinid shells to temperature and/or salinity (Lorrain et al., 2005; Poitevin et al., 2020). 648 This study's individuals that grew during the same year in the same environment do not show a 649 650 synchronous Mg/Ca pattern (Fig. 2), arguing against a simple temperature dependence for Mg/Ca 651 in P. maximus. In addition, the lack of strict coherence between profiles of Mg/Ca (and other elements) in parallel transects through P. maximus shells (e.g., PM2\_1 and PM2\_2; Fig. 2) hints 652 653 at compositional heterogeneity within the shells, in agreement with findings by (Lorrain et al.,

654 2005)hints at compositional heterogeneity within the shells. Low correlations between profiles 655 through the same shell at the daily scale are also partly driven by small misalignments of the 656 timing of shell formation between the transects at the sub-millimeter scale and variations in the 657 height of trace element peaks, especially in Mn/Ca and Ba/Ca, which are higher further towards 658 the outside of the shell (**S2**).

659 There is evidence suggesting that Mg content varies in mollusk shells in function of the amount of organic matter in the biomineral (Dauphin et al., 2003; Richard, 2009; Tanaka et al., 2019) 660 661 (Dauphin et al., 2003; Richard, 2009; Tanaka et al., 2019). Contrarily, Mn is known to be taken up in thermodynamic equilibrium in the mineral fraction of bivalve shells (Onuma et al., 1979; 662 Soldati et al., 2016) (Onuma et al., 1979; Soldati et al., 2016), and Mn/Ca ratios in P. maximus 663 664 have been shown to faithfully record fluctuations of dissolved Mn in the coastal environment 665 driven by riverine input and redox conditions (Barats et al., 2008) (Barats et al., 2008). Similarly, 666 there is strong evidence that Ba/Ca ratios in P. maximus (and other mollusks) record changes in 667 Ba available in the environment linked to primary productivity (e.g., (Gillikin et al., 2008; Thébault et al., 2009; Fröhlich et al., 2022) Gillikin et al., 2008; Thébault et al., 2009; Fröhlich et al., 2022). 668 Interestingly, our results (Fig. 2) show that background Ba/Ca values are not equal in the shells 669 670 of *P. maximus* and tridacnid specimens grown in the same environment. This contradicts the assessment by (Gillikin et al., 2005) that background Ba/Ca concentrations are a function of 671 environmental conditions and can be consistently subtracted from Ba/Ca records to separate 672 673 peak from background values. This The suggested relationship between Ba/Ca and primary 674 productivity would explain the skewed (skewness > 1; S8) character of the Ba/Ca records and the correlation between Ba/Ca and Mn/Ca in our the studied P. maximus specimens, as the reducing 675 676 conditions following peaks in primary productivity favor the remobilization of Mn into the water 677 column causing short-lived simultaneous increases in Ba/Ca and Mn/Ca in the shells (Dehairs et 678 al., 1989; Barats et al., 2008, 2009) (Dehairs et al., 1989; Barats et al., 2008; 2009).

679 4.42.2 -Short-term changes in shell composition in tridacnids P. maximus

In the context of the high-resolution trace element variability central to this study, it seems 680 plausible that short-term changes in the environment of the Bay of Brest were drivers of Mn/Ca 681 and Ba/Ca variability in *P. maximus* shells, while Mg/Ca and Sr/Ca composition is mostly driven 682 683 by changes in calcification rate. This would suggest that the significant tidal (12 h) component in 684 Ba/Ca and Mn/Ca records through *P. maximus* (Fig. 43) is driven directly by redox changes over the strong tidal cycle in the Bay of Brest (see (Polsenaere et al., 2021)-Polsenaere et al., 2021) 685 686 and resuspension of Ba and Mn due to tidal currents (Hily et al., 1992) (Hily et al., 1992), while tidal rhythms in Mg/Ca and Sr/Ca may be a consequence of the scallop's calcification response 687 to changes in its environment (e.g., temperature, salinity and light availability) through the large 688 689 (up to 7\_m range) tidal cycle (Fig. 87). The latter corroborates with previous studies in other calcitic 690 mollusk shells which demonstrated that Mg incorporation on short timescales is driven by the 691 metabolic response to subtle changes in the environment (Lazareth et al., 2007)-(Lazareth et al., 692 2007). Finally, care must be taken to interpret trace element variability in *P. maximus* shells, since large intra-shell gradients in Mg/Ca, Sr/Ca and Mn/Ca have previously been observed in this 693 species, making trace element composition highly dependent on the location of measurements 694 695 relative to the outer shell surface or positioning relative to striae on the shell surface\_(Freitas et al., 2009) (Freitas et al., 2009). Even though the LA-ICP-MS line scans in this study targeted 696 exclusively the oOSL of *P. maximus* specimens, variability in elemental ratios resulting from small 697 698 changes in the distance of the line scan from the outer edge of the shell cannot be fully excluded 699 (Richard, 2009) (Richard, 2009).

700

701 4.2-3 Trace element variability in tridacnids

702 4.23.1 Comparison with previous studies

703 Results for Sr/Ca, Mg/Ca and Ba/Ca in this study's tridacnid specimens broadly corroborate trace 704 element results in other tridacnid studies (e.g., (Elliot et al., 2009; Sano et al., 2012; Yan et al., 2013; Warter et al., 2018) Elliot et al., 2009; Sano et al., 2012; Yan et al., 2013; Warter et al., 705 706 2018). While data on Mn/Ca in the OSL of tridacnids is scarce, the Mn/Ca ratios in tridacnids in 707 this study (mean Mn/Ca =  $7.8 \pm 4.7 \mu mol/mol$ ) are similar to LAICPMSLA-ICP-MS Mn/Ca data available in the literature ((Warter et al., 2015), 4-10 µmol/mol), but significantly lower than Mn/Ca 708 values measured using total digestion Atomic Absorption Spectrometry ((Madkour, 2005), ~30 709 710 µmol/mol). The main difference between the techniques is that LA-ICP-MS (both in this study as 711 in (Warter et al., 2015)) sampled shell layers where growth lines were visible and did not include pre-treatment, while the total digestion study (Madkour, 2005) removed organic matter by roasting 712 the shells at 200°C prior to bulk shell analysis. Given that Mn in bivalve shells is typically 713 714 associated with the mineral fraction of the shell (Soldati et al., 2016), The the difference in results 715 may therefore hint at differences between shell layers within tridacnids, or differences in Mn 716 concentration between the organic and mineral fractions in the shells. Bivalve shells typically 717 contain between 1\_% and 5\_% organic matter (Marin and Luquet, 2004), with tridacnid shells being notable for their low organic matter content (< 1 %; (Taylor and Layman, 1972; Agbaje et al., 718 719 2017) Agbaje et al., 2017; Taylor and Layman, 1972). Given that most Mn in bivalve shells is typically associate with the mineral fraction of the shell (Soldati et al., 2016), it-lt thus seems 720 721 unlikely that such a largethe fraction of differences in Mn/Ca ratio between bulk analyses 722 (Madkour, 2005) and in situ analysis ((Warter et al., 2015); this study) could originate from 723 variations within the organic matrix, which only constitutes such a small fraction of the shell. 724 Therefore, we consider a difference in Mn concentration between shell layers in tridacnids more likely. The lack of consistent trace element offsets between the tridacnid species under study here 725 726 (T. maxima, T. squamosa and T. squamosina) confirms the chemical similarities of shells tridacnid 727 species found in previous studies (e.g., T. gigas; (Elliot et al., 2009; Yan et al., 2013)-Elliot et al., 2009; Yan et al., 2013; T. crocea; (Warter et al., 2018); T. derasa; (Sano et al., 2012)). 728

## 729 4.23.2 Short-term variability in Sr/Capaced to the day-night cycle

730 Sr/Ca in tridacnids is thought hypothesized to be strongly controlled by light intensity through a circadian rhythm linked to the day-night cycle (Sano et al., 2012; Warter et al., 2018). This would 731 732 explain the strong daily periodicity in Sr/Ca records through all tridacnids in this study. This daily 733 periodicity may be caused by the ctenidium in tridacnids working on a daily rhythm to keep the 734 acid-base balance in the hemolymph of the clams to offset the CO<sub>2</sub> depletion by photosymbionts (which is paced to the day-night cycle of light availability). In the process, Ca2+-channels and 735 736 Na<sup>+</sup>/H<sup>+</sup>-exchangers work to keep the charge balance in the internal fluid and provide nutrients and ions for shell mineralization, letting in compatible trace elements such as Sr<sup>2+</sup> (Ip and Chew, 737 2021). This mechanism follows the biomineralization model by (Carré et al., 2006) and is 738 supported by the high affinity of Sr<sup>2+</sup> with Ca-channels (Hagiwara and Byerly, 1981) and the high 739 740 ionic fluxes supported by this pathway, allowing enough membrane permeability to support the 741 fast shell formation in tridacnids (Coimbra et al., 1988; Sather and McCleskey, 2003). Following this line of reasoning, the preconcentration of Sr<sup>2+</sup> in the extrapallial fluid through Ca-channels 742 743 should have a larger effect on shell Sr/Ca ratios in tridacnids than the discrimination against Sr<sup>2+</sup> 744 (or other trace elements) through shell organic matrix during mineralization of the shell from this 745 fluid (as proposed in (Gillikin et al., 2005)). This model could explain the indirect link between 746 trace element uptake incorporation into the shell in of tridacnids and the day-night cycle without 747 a direct causal relationship between trace element concentration and light availability (as 748 demonstrated by the strong daily cycle in trace elements in the shaded **TSM1** specimen). It is 749 worth noting that experiments on freshwater bivalves (e.g., Corbicula fluminea; (Zhao et al., 2017)) revealed that a closure of the Ca<sup>2+</sup> channels did not influence Sr concentrations in the 750 shell, arguing against a kinetic effect on Sr partitioning into the shell. 751

752 4.23.3 Tidal vs. diurnal variability

753 Our spectral analysis does not allow us to distinguish between the expression of the solar day (24 754 h) and lunar day (~24.8 h) because the width of the bandpass filters used to extract periodicities (~22--36 h) encompass both frequencies. While we cannot exclude the possibility that some of 755 756 the daily (24 h frequency band) periodicity in tridacnid records is an expression of the lunar cycle, 757 it seems unlikely for most records except Sr/Ca, because the expression of the other lunar cycles 758 periodicities (most notably the ~12 h cycle) is much weaker in tridacnids compared to the pectinids 759 (see Fig. 6-7). Nevertheless, it remains possible that the diurnal cycle in Sr/Ca in tridacnids, 760 previously interpreted as a response to the day-night cycle, is in fact caused by a circadian rhythm 761 paced to the lunar day. Additionally, vertical mixing, a major driver of sea surface temperature 762 changes in the northern Gulf of Agaba, is shown to be driven by a combination of surface wind intensity (which has strong daily variability) and the presence of tidal currents (Carlson et al., 763 764 2014). It is therefore possible that changes in local surface water temperature partly control the 765 observed (semi-)diurnal variability in trace element concentrations in Red Sea tridacnids.

## 766 4.23.4 Seasonal variability and temperature relationships

On longer (seasonal) timescales, Sr/Ca in tridacnids has been suggested as a temperature proxy 767 similar to the well-known Sr/Ca-Sea Surface Temperature relationship in tropical corals (Lough, 768 769 2010; Yan et al., 2013). However, significantly lower Sr/Ca ratios in tridacnids compared to coral 770 aragonite (1.5 – 2.0 mmol/mol vs. 7.5 – 9.5 mmol/mol in corals; (Elliot et al., 2009); Fig. 2) suggest that tridacnids exert a large degree of biological control on the Sr concentration in their shells, 771 772 either possibly through the light-sensitive photosymbiosis-calcification relationship outlined above 773 (section 4.3.2) or otherwise through active Sr removal from the biomineralization site by Sr-774 binding organic molecules (following the model proposed by (Gillikin et al., 2005)). Similarly, 775 Mg/Ca ratios in tridacnids were previously thought to primarily record water temperature (e.g., 776 (Batenburg et al., 2011)) but our detailed investigation shows here large differences in Mg 777 concentration within tridacnid shells, and a strong anticorrelation of Mg with sulfur compounds

778 associated with the organic matrix in the shell (see section 4.2.14; (Dauphin et al., 2003)), has 779 been put forward as evidence for a strong control of calcification and microstructure on Mg composition in tridacnid shells (Yoshimura et al., 2014). However, evidence from studies on 780 foraminifera calcification demonstrate that the sulfur in biocarbonates is not organically bound 781 782 and that the covariation with Mg might instead be caused by lattice distortion due to incorporation 783 of Mg favoring simultaneous S incorporation (van Dijk et al., 2017). If this is also the case in tridacnids, the Mg/Ca-S/Ca covariance observed in previous studies might not preclude the use 784 of Mg/Ca as a temperature proxy. In that case, more detailed calibration studies (e.g. using 785 786 controlled growth experiments; (Warter et al., 2018)) are required to establish this proxy.

## 787 4.23.5 Ba/Ca and Mn/Ca in tridacnids

788 As in the pectinids, Ba/Ca ratios in tridacnids likely reflect changes in Ba concentration in the 789 environment, which can be caused by river input, upwelling of comparatively nutrient-rich waters 790 or blooms of Ba-rich phytoplankton (Vander Putten et al., 2000; Elliot et al., 2009). Given that Mn is mostly associated with the mineral fraction of bivalve shells and seems to fractionate into the 791 shell close to equilibrium with seawater (Onuma et al., 1979; Soldati et al., 2016), Mn/Ca ratios in 792 tridacnids likely reflect the availability of dissolved Mn in the water column, as in other mollusk 793 794 taxa (e.g., (Barats et al., 2009); see section 4.24). This assumption is supported by the correlation 795 between Mn/Ca and Ba/Ca measured in this study (Fig. 2), suggesting that both records are 796 influenced on seasonal timescales by variability in nutrient availability and redox conditions (sensu 797 (Dehairs et al., 1989)). Part of this correlation between Mn/Ca and Ba/Ca is driven by synchronous increases in both elements near the start and end of the profiles through tridacnid shells (Fig. 2). 798 799 These changes may reflect a decrease of control on shell composition during periods of stress, or alternatively reflect periods of slower growth which cause more primitive microstructures 800 (characterized by higher concentrations of trace elements) to be formed (Warter et al., 2018). 801

4.2.6 Environmental changes in the Gulf of Aqaba

803 Given that the Gulf of Agaba is oligotrophic, seasonally stratified, and lacks significant riverine 804 input (Nassar et al., 2014; Manasrah et al., 2019)(Nassar et al., 2014; Manasrah et al., 2018), the variability in nutrient concentrations and redox conditions driving Mn/Ca and Ba/Ca variability in 805 806 tridacnids are likely driven by convective overturning. The tidal amplitude is much smaller than in 807 the Bay of Brest (<\_1 m; (Manasrah et al., 2019)) and is unlikely to drive significant short-term 808 fluctuations in sea water chemistry. This may therefore explain the lack of tidal (12 h) periodicity 809 in Ba/Ca and Mn/Ca in tridacnids (Fig. 5-6). Nevertheless, tidal rhythms have been observed in 810 the behavior and growth of deep-sea bivalves living far below the direct influence of tides on the 811 environment, proving that such patterns can be recorded by the animals through their circadian rhythm (Schöne and Giere, 2005; Nedoncelle et al., 2013; Mat et al., 2020)(Schöne and Giere, 812 813 2005; Nedoncelle et al., 2013; Mat et al., 2020). In this case, the daily cycle seems to have been 814 more important for Ba/Ca in tridacnids, plausibly by driving diurnal changes in primary productivity 815 in the Gulf of Agaba. Alternatively, the daily periodicity found in tridacnid shell chemistry may in 816 fact be a response to the lunar day (~24.8 h) cycle, which is imprinted in the shell's chemical 817 composition through periodic exposure of the clams to extreme heat or air (subaerial exposure) in their shallow water environment during exceptionally low tides. The stress induced from this 818 819 exposure could have affected calcification and incorporation of trace elements (see 820 abovesection 4.3.5).

Interestingly, Sr/Ca ratios in tridacnids do exhibit tidal periodicity (**Table 2**; Fig. 3), perhaps driven
by a circadian rhythm in calcification linked to the tidal cycle, or by subtle changes in water
temperature driven by tidal currents (Carlson et al., 2014). This 12\_h periodic behavior is not
observed in previous studies of Sr/Ca ratios in tridacnids (Sano et al., 2012; Warter et al., 2018).
A recent valvometric study on tridacnids found a 12\_h period in activity, which supports the
hypothesis that a circadian rhythm paced to the tidal cycle could influence shell calcification
(Killam and Clapham, 2018b). Significant daily fluctuations in solar radiation (up to 1500 W m<sup>-2</sup>;

828 (Manasrah et al., 2019)) likely exerted a dominant control on the calcification of tridacnids in the 829 Red Sea, explaining the strong diurnal periodicity in Sr/Ca and Ba/Ca records in this study (see Fig. 67 and Fig. 98). As in the (non-symbiotic) pectinid data, it seems likely that the majority of 830 Mn/Ca and Ba/Ca variability in tridacnids directly reflects changes in the chemistry of the sea 831 832 water and its constituents (e.g., particulate organic matter) while Mg/Ca and Sr/Ca variations are driven by changes in calcification and microstructure. The latter may be indirectly influenced by 833 light intensity through photosynthesis by the symbionts, or by circadian rhythms paces to the 834 diurnal or tidal cycle. 835

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837 4.<u>3-4</u> Role of photosymbiosis on high-frequency chemical variability

### 838 4.34.1 Effect of symbiosis on calcificationshell formation

While the amplitude of diurnal variability in trace element concentrations does not vary much 839 840 between the symbiotic tridacnids and the non-symbiotic pectinids (Fig. 56), the amount of 841 variance in the trace element records explained by daily cyclicity is up to twice as high in tridacnids 842 (Fig. 76; Table 3). This suggests that the 24 h cycle has a much larger relative influence on trace 843 element composition (especially Sr/Ca and Ba/Ca) in tridacnids than in pectinids. This seems to point towards a role of the photosymbionts in calcification by tridacnids, such as through symbiont-844 mediated diurnal variation in the pH of the extrapallial fluid (Ip et al., 2006), as well as active 845 846 transport of HCO<sub>3</sub><sup>-</sup> for calcification (Chew et al., 2019) and as a C supply to the symbionts from 847 the host (Boo et al., 2021). Given the differences in absolute elemental ratios between these two groups of bivalves, comparing variance yields a more robust assessment of the relative 848 849 importance of tidal or diurnal variability on shell composition than looking at the absolute size 850 (amplitude) of the chemical cycle. While the difference in variance is clear, the importance of diurnal cyclicity on the photosymbiotic tridacnids is not as big as one might expect. Rarely more 851

than 10\_% of the variance is explained by day-night variability (**Table 3**). This seems to contradict the large daily Sr/Ca amplitudes found in (Warter et al., 2018) and the trace element fluctuations found in (de Winter et al., 2020), which rival the seasonal cycle in these trace element ratios in terms of amplitude. However, the percentages in **Table 3** relate to the amount of variation in the complete records through these individuals and therefore also contain areas of the shell where daily cyclicity is less pronounced, while values in previous studies often reflect maximum amplitudes recorded in parts of the shell with exceptionally clear daily increments.

## 859 4.<u>34</u>.2 Effect of differences in the environment

It seems plausible that part of the difference in diurnal variability between pectinids and tridacnids 860 861 is explained by a difference in the environment between the Gulf of Agaba and the Bay of Brest, rather than the presence of photosymbionts. The diurnal insolation cycle in the Gulf of Agaba is 862 larger than in the Bay of Brest (1500 vs 546 W\*m<sup>-2</sup> maximum summer irradiance; (Roberts et al., 863 2018; Manasrah et al., 2019)). If calcification in pectinids would be equally sensitive to sunlight, 864 this difference may explain much of the difference between the species. In this scenario, part of 865 the strong tidal component in the pectinid trace element data may be explained by the influence 866 of differences in water depth on the penetration of sunlight through the murky waters of the Bay 867 868 of Brest (Roberts et al., 2018). In fact, tidal movement can cause strong non-linear amplification 869 or reduction of the solar irradiance at the sea floor of the Bay of Brest by factors exceeding 10, especially outside the summer months, which in turn has a significant effect on primary 870 871 productivity in the water column (Roberts et al., 2018). This tidal effect is likely to be much weaker 872 in the Gulf of Agaba, given its comparatively low tidal amplitude, clear oligotrophic waters, and 873 much stronger and less seasonal day-night cycle (Manasrah et al., 2019). Indeed, even in nonphotosymbiotic bivalves, light and food availability are demonstrated to be major drivers of the 874 animal's behavior (e.g., (Ballesta-Artero et al., 2017)). The combination of the daily and tidal 875 cycles on solar irradiance at depth and photosynthesis in the Bay of Brest may therefore pose an 876

alternative pathway for strong tidal cyclicity in the trace element composition of pectinids in this
study and account for part of the twofold increase in daily variability in tridacnids compared to the
pectinids (Fig. 6-77-8; Table 3).

880 4.<u>34</u>.3 Effect of direct insolation

881 Specimen **TSM1** poses an interesting case study for investigating the link between sunlight and 882 calcification in tridacnids, since it grew under a sunshade and therefore experienced a dampened 883 diurnal variability in insolation compared to other giant clams in the area. The fact that this 884 specimen exhibits similar or even higher diurnal variability in shell chemistry (Fig. 67) argues against a direct influence of the rate of photosynthesis itself on calcification. Instead, it seems that 885 886 daily chemical variability is mostly an expression of circadian rhythm in tridacnids, which is 887 strongly (evolutionarily) coupled to the day-night cycle to optimize the symbiosis with primary 888 producers in its mantle, possibly through respiration rhythms carried out by the ctenidium (see 889 section 4.32; (Ip and Chew, 2021)). Symbionts have been shown to directly aid in calcification in 890 terms of through proton pumping (Armstrong et al., 2018), influencing internal acid-base chemistry (Ip et al., 2006), and valvometric studies show the clams bask in sunlight in daylight hours and 891 close partially at night when symbiosis is likely reduced (Schwartzmann et al., 2011). This 892 conclusion is further supported by the lack of a clear difference in diurnal cyclicity between trace 893 894 element records in T. maxima, T. squamosa and T. squamosina (Fig. 67; Table 3), even though 895 the degree of reliance on photosymbiosis is demonstrated to be highly variable between these 896 species (Killam et al., 2020). Therefore, it seems unlikely that sub-daily resolved trace element 897 records in tridacnids can be used as quantitative recorders of paleo-insolation, as was originally 898 suggested by (Sano et al., 2012). While the degree of symbiotic activity may not be clearly 899 recorded in the daily amplitude of trace element oscillations, the greater consistency of daily 900 periodic signal in the studied giant clams could relate to the direct biological control exerted by 901 the symbionts on the hosts' rhythms of calcification. Light exposure in giant clams promotes

expression of genes coding for proteins involved in Ca<sup>2+</sup>, H<sup>+</sup> and HCO<sub>3</sub><sup>-</sup> transport in the mantles 902 903 of giant clams (Ip et al., 2017; Chew et al., 2019), with the expression proposed to be at least partially mediated by photosensing on the part of the symbionts themselves (lp et al., 2017). 904 Differences between the daily consistency (spectral power) of photosymbiotic and non-905 906 photosymbiotic trace element profiles might still allow paleontologists to use the presence of 907 strong daily periodicity as a proxy for photosymbiosis in the fossil record (as suggested byin (de Winter et al., 2020)). However, the small differences found between pectinids and tridacnids in 908 909 this study and the comparatively large influence of environmental variability show that such 910 records should be interpreted with caution. Future studies could measure photosynthetic activity of the symbionts in tridacnids and attempt to relate this to the trace element composition of the 911 shell in an attempt to isolate the direct effect of photosymbiosis on shell composition. 912

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914 4.4-5 Aperiodic drivers of shell chemistry

### 915 4.4<u>5</u>.1 Circadian and behavioral changes

Even after controlling for instrumental noise, most (~\_90\_%) of the variance observed in our trace 916 917 element records is not directly related to the diurnal or tidal cycle. This suggests that aperiodic events at the scale of hours to days play an important role in the calcification of pectinids and 918 tridacnids. Given the large difference in ecological niche (e.g., photosymbiotic versus non-919 920 symbiotic) between these taxa, and the difference between the environment in which they grew, 921 this observation suggests that calcification of bivalves at the (sub-)daily scale may is generally be 922 dominated by aperiodic variability in calcification or in the environment. Part of this unaccounted 923 variability may be caused by variability in the animal's behavior, as documented by observations 924 of siphon and valve gape activity in cultured or monitored specimens of a variety of bivalve taxa (Rodland et al., 2006; Ballesta-Artero et al., 2017). While these experiments revealed quasi-925

926 periodic (3-7 minute and 60–90 minute periods) behavior unassociated with the tidal or daily cycle, records of activity of the individuals-bivalves also reveal less regular patterns on the scale of 2\_-927 24 h which may contribute to the aperiodic variance in trace element records (Rodland et al., 928 929 2006). Another example of aperiodic behavior potentially influencing shell chemistry is rapid valve 930 adduction or "coughing" observed in both pectinids and tridacnids, which serves as a mechanism 931 for expelling respiratory  $CO_2$  and faeces from the pallial cavity or to evade predation attempts (Robson et al., 2012; Soo and Todd, 2014). This behavior could resuspend sediment and produce 932 pulses of Mn and Ba at the sediment-water interface which are recorded as short-term, aperiodic 933 934 variability in these elements in the shell. The temporal sampling resolution (>1hseveral hours) of our trace element records after smoothing out measurement noise does not allow us to resolve 935 936 the types of periodic variability at the sub-hourly scale cited in these previous studies, meaning 937 that aperiodic variability in behavior and aliasing of these ultradian patterns likely contribute to the 938 aperiodic variability in our trace element records. On longer (sub-)seasonal timescales, activity in 939 bivalves is shown to be highly dependent on food and light availability (Ballesta-Artero et al., 940 2017), suggesting that aperiodic, short-term changes in these environmental factors could be a main driver of shell growth and composition and explain a large part of the variance in the trace 941 942 element records which is not explained by ultradian changes in the animal's behavior.

943 4.4<u>5</u>.2 Short-term environmental changes and paleoweather

Outside of regular fluctuations caused by tidal, daily, and seasonal cycles, changes in light and food availability at the hourly to daily scale are probably linked to circulation and weather phenomena. Previous studies show that enhanced vertical mixing during weather events such as storms, algal bloom events after wind-driven upwelling and pseudoperiodic dust deposition can temporarily increase the concentration of dissolved metals in surface waters, resuspend organic matter and temporarily increase primary productivity<del>,</del> (Lin et al., 2003; Al-Najjar et al., 2007; Iluz et al., 2009; Al-Taani et al., 2015; Komagoe et al., 2018)<del>(Lin et al., 2003; Al-Najjar et al., 2007;</del>

951 Iluz et al., 2009; AI-Taani et al., 2015; Komagoe et al., 2018). This will in turn lead to a shallowing 952 of the redoxcline through increased organic matter load at the sediment-water interface, which can be recorded in the composition of giant clam shells (Yan et al., 2020). Interestingly, data in 953 954 (Yan et al., 2020) suggest that recording an extreme weather event in Tridacna requires wind 955 speeds exceeding 20 km/h, a threshold which is almost never reached in the comparatively quiet 956 Gulf of Aqaba (Manasrah et al., 2019), while such events are common in the stormier Bay of Brest 957 (Hily et al., 1992; Chauvaud et al., 2005). This difference is also reflected in the periodicity of shell 958 composition, with the tridacnids having overall higher percentages of their variance explained by 959 daily and tidal variability than pectinids (Fig. 3 and 7), showing that aperiodic (potentially weathercontrolled) variability in shell composition has a stronger influence on the pectinids which grew in 960 961 the stormier Bay of Brest. highlights another difference between the environments of pectinid and 962 tridacnid specimens investigated in this study which could contribute to the variable expression 963 of periodicity in the trace element composition of their shells. A plausible scenario therefore emerges in which aperiodic weather events cause short-term variability in both the chemistry and 964 965 physical properties of the water column. These changes are subsequently recorded in bivalve shells, either directly because the weather events resuspend, remobilize or deliver trace elements 966 967 like Mn and Ba (e.g., (Dehairs et al., 1989; Gillikin et al., 2008; Mahé et al., 2010)), or indirectly because environmental stress associated with the event affects behavior and shell calcification, 968 resulting in a change in the incorporation of alkali-group cations (e.g., Mg and Sr) into the shell 969 970 biomineral ((Carré et al., 2006; Takesue et al., 2008); Fig. 87). Our results therefore highlight the 971 potential of high-resolution trace element records in bivalve shells to record short-term circulation changes and weather events, while prescribing caution in interpreting such records until the effect 972 973 of true environmental changes on the sub-daily scale can be separated from aperiodic ultradian 974 or behavioral patterns.

#### 976 5. Conclusions

Our high-resolution trace element records reveal that short-term variability on the tidal or daily 977 scale is recorded in the Mg, Sr, Mn, and Ba composition of shells of fast-growing mollusk species. 978 979 The application of spectral analysis and variance decomposition on these trace element records 980 is a useful tool to assess the influence of periodicity in the shallow marine environment on 981 calcification in mollusk shells. Our statistical analysis reveals that tidal and daily variability each 982 on average account for less than 10\_% of trace element variance in pectinids and tridacnids. In 983 photosymbiotic giant clam shells, the amount of variance in Sr and Ba paced to the daily cycle is two times higher than in the non-photosymbiotic pectinids, suggesting that photosymbiosis in 984 giant clams exerts some control on trace element composition in their shells. However, since only 985 ~10 % of the trace element variability in tridacnids is explained by diurnal variability, the 986 987 recognition of photosymbiosis in the fossil record from diurnal variability in fossil shell composition 988 will be complicated. In addition, differences between the mid-latitude environment of the pectinids 989 and the tropical environment of the tridacnids likely account for part of the difference in trace 990 element composition between the taxa.

991 -We propose that Ba and Mn composition in pectinids and tridacnids reflect short-term variability 992 in primary productivity and sea water chemistry which control the mobility of these elements. 993 Concentrations of Mg and Sr are likely controlled by short-term changes in growth and metabolic rate of the mollusks, which may be indirectly controlled by changes in their environment through 994 circadian rhythms or behavior, explaining the pacing of trace element composition to the tidal and 995 diurnal cycle. Most of the variance in trace element records in both taxa are not related to periodic 996 997 behavior at the 12 h or 24 h scale, likely recording aperiodic events in the environment related to weather-scale phenomena or circadian patterns. We thus conclude that mollusk shell carbonate 998 999 is a promising archive for recording weather-scale variability in shallow marine environments

across latitudes, potentially recording weather-scale phenomena in deep time, as long as these
 environmental effects can be separated by the influence of the <u>animal's</u> behavior<u>of the animal</u>.
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# 1003 Code availability

Scripts used for data processing and to create figures in this manuscript were uploaded to an open-access repository on GitHub (https://github.com/nielsjdewinter/TE\_circadian) and linked through Zenodo (https://zenodo.org/record/6603175).

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### 1008 Data availability

Supplementary data and figures referenced in this contribution were uploaded to the online openaccess repository Zenodo (https://doi.org/10.5281/zenodo.6602894).

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#### 1012 Author contribution

NJW designed the experiment after discussion with BRS, DK and LF. LF, DK, BRS and JT collected the samples. LF, DK and NJW together prepared samples for analyses and constructed shell chronologies using growth line counting. WB, LN, GJR and NJW carried out the LA-ICP-MS analyses and data processing. NJW designed and carried out the statistical analyses and wrote the R scripts guided by feedback from LN, BK, LN, WB and GJR. NJW wrote the first draft of the manuscript. All authors contributed to the writing process towards the final version of the manuscript.

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#### 1021 Competing interests

1022 The authors declare that they have no conflict of interest.

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