

1      **Reference Soil Groups Map of Ethiopia Based on Legacy Data and Machine**  
2      **Learning Technique: EthioSoilGrids 1.0**

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25     **Abstract.** Up-to-date digital soil resource information and its comprehensive understanding are  
26     crucial to supporting crop production and sustainable agricultural development. Generating such  
27     information through conventional approaches consumes time and resources, and is difficult for  
28     developing countries. In Ethiopia, the soil resource map that was in use is qualitative, dated (since  
29     1984), and small-scaled (1:2 M) which limit its practical applicability. Yet, a large legacy soil profile  
30     data accumulated over time and the emerging machine learning modelling approaches can help in  
31     generating a high-quality quantitative digital soil map that can provide better soil information. Thus,

32 a group of researchers formed a coalition of the willing for soil and agronomy data sharing and collated  
33 about 20,000 soil profile data and stored them in a central database. The data were cleaned and  
34 harmonised using the latest soil profile data template and 14,681 profile data were prepared for  
35 modelling. Random Forest was used to develop a continuous quantitative digital map of 18 World  
36 Reference Base (WRB) soil groups at 250 m resolution by integrating environmental covariates  
37 representing major soil-forming factors. The map was validated by experts through a rigorous process  
38 involving senior soil specialists/pedologists checking the map based on purposely-selected district  
39 level geographic windows across Ethiopia. The map is expected to have tremendous value in soil  
40 management and other land-based development planning, given its improved spatial resolution and  
41 quantitative digital representation.

42 **Keywords:** soil profiles, environmental covariates, modelling, expert validation, Reference Soil  
43 Group

## 44 1 Introduction

45 Soils are important resources that support the development and production of various economic, social,  
46 and ecosystem services, and are useful in climate change mitigation and adaptation (Baveye et al.,  
47 2016). Data on soils' physical and chemical characteristics and their spatial distribution are needed to  
48 define and plan their functions over time and space, which are important steps towards sustainable use  
49 and management of soils (Elias, 2016; Hengl et al., 2017).

50 In Ethiopia, soil surveys and mapping have been conducted at various scales with varying scopes,  
51 approaches, methodologies, qualities, and levels of detail (Abayneh, 2001; Abayneh and Berhanu,  
52 2007; Berhanu, 1994; Elias, 2016; Zewdie, 2013). The most recent country-wide digital soil mapping  
53 efforts focused primarily on soil characteristics (Ali et al., 2020; Iticha and Chalsissa, 2019; Tamene  
54 et al., 2017), although soil class maps are equally important for allocating a particular soil unit for  
55 specific use (Leenaars et al., 2020a; Wadoux et al., 2020). Many attempts have been made to improve  
56 digital soil information systems (Hengl et al., 2021, 2017, 2015; Poggio et al., 2020). However, the  
57 initiatives were based on limited and unevenly distributed soil profile data (e.g., 1.15 soil profiles per  
58 1,000 km<sup>2</sup> for Ethiopia) which restricts the accuracy and applicability of the products.

59 In Ethiopia, thousands of soil profile data have been collected since the 1960s (Erkossa et al., 2022),  
60 but these data were scattered across different institutions and individuals (Ali et al., 2020).  
61 Furthermore, country-wide quantitative and gridded spatial soil type information does not exist (Elias,  
62 2016). The Ethiopian Soil Information System (EthioSIS) project attempted to develop a countrywide

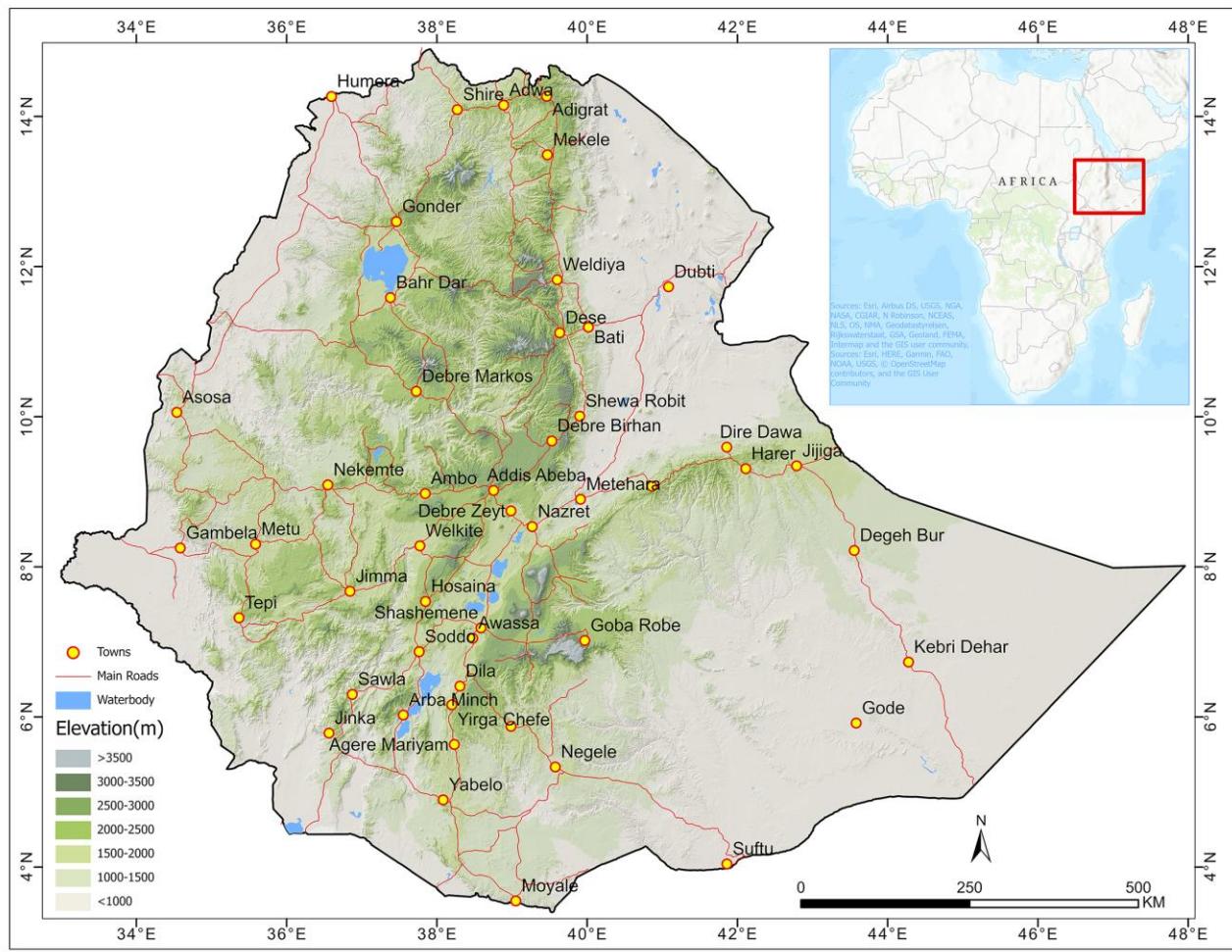
63 digital soil map focusing on topsoil characteristics, including plant nutrient content, but overlooked  
64 soil resource mapping (Ali et al., 2020; Elias, 2016), despite a strong need for a high-resolution soil  
65 resource map (Mulualem et al., 2018).

66 Ethiopia has an area of about 1.14 mill. km<sup>2</sup> consisting of varied environments, making its soils  
67 extremely heterogeneous. Capturing the heterogeneity using conventional soil survey and mapping  
68 approaches is an expensive and time-consuming endeavour (Hounkpatin et al., 2018). This can be  
69 circumvented using available legacy soil profile data accumulated over decades and tapping into the  
70 potential of advanced analytical techniques to develop high-resolution digital soil maps (Hounkpatin  
71 et al., 2018; Kempen, 2012, 2009). Therefore, the objectives of this study were to (1) develop a national  
72 legacy soil profile dataset that can be used as an input for various digital soil mapping exercises, and  
73 (2) generate an improved 250 m digital Reference Soil Groups (RSGs) map of Ethiopia.

## 74 **2 Methods**

### 75 **2.1 The study area**

76 The study area covered the entire area of Ethiopia (1.14 mill. km<sup>2</sup>) located between 3°N and 15° N,  
77 and between 33° E and 48° E (Figure 1). The topography of the country is marked by a large altitudinal  
78 variation, ranging from 126 meter below sea level at Dalol in the northeast to 4,620 m at Ras Dashen  
79 Mountain in the northwest (Billi, 2015; Enyew and Steeneveld, 2014). Ethiopia's wide range of  
80 topography, climate, parent material, and land use types created conditions for the formation of  
81 different soil types (Abayneh, 2005; Berhanu and Ochtman, 1974; Donahue, 1972; Mesfin, 1998;  
82 Nyssen et al., 2019; Virgo and Munro, 1978; Zewdie, 2013, 1999). More than 33% of the country is  
83 covered by the central, upper and highland complex (Abegaz et al., 2022), which embraces Africa's  
84 most prominent mountain system (Hurni, 1998).



86 **Figure 1.** Location map of Ethiopia, overview map © Esri World Topographic Map.

87 The country's complex topography strongly determines both rainfall and temperature patterns, by  
 88 modifying the influence of the large-scale ocean-land-atmosphere pattern, thus creating diverse  
 89 localised climates. Spatially, rainfall is characterised by a general decreasing trend in the direction  
 90 from the west- to east, north, northeast, south and southeast. The lowlands in the southeast and  
 91 northeast, covering approximately 55% of the country's land area, are characterised by arid and semi-  
 92 arid climates. Annual rainfall ranges from less than 300 mm in the south-eastern and north-western  
 93 lowlands to over 2,000 mm in the southwestern (southern portion of the western highlands). The  
 94 eastern lowlands get rain twice a year, in April–May and October–November, with two dry periods in  
 95 between. The total annual precipitation in this region varies from less than 500 to 1,000 mm. The driest  
 96 of all regions is the Denakil Plain, which receives less than 500 mm and sometimes none (Fazzini et

97 al., 2015). Temperatures are also greatly influenced by the rapidly changing altitude and the mean  
98 monthly values vary from ~35°C in the northeast lowlands to less than 7.5°C over the north and central  
99 highlands.

100 The country is characterised by a wide variety of geological formations (Abyneh, 2005; Alemayehu  
101 et al., 2014; Elias., 2016; Jarvis et al., 2011; Zewdie, 2013). These include (i) recent and old volcanic  
102 activities; (ii) the highlands consisting of igneous rocks (mainly basalts); (iii) steep-sided valleys  
103 characterised by strong colluvial and alluvial deposits; (iv) metamorphic rocks exposed by denudation  
104 process; and (v) various sedimentary rocks like limestone and sandstone in the relatively lower areas.

105 Diverse biophysical factors affecting the spatial distribution of vegetated land cover which in turn both  
106 as single and combined factors result in diverse soil types and properties across Ethiopia's landscapes  
107 (Hurni, 1998; Nyssen et al., 2019; WLRC, 2018). The spatio-temporal vegetation cover of the country  
108 has been characterised by a long history of landuse-landcover changes(WLRC, 2018). In terms of the  
109 type and spatial coverage of major landuse/landcover classes, woody vegetation (forest, woodland,  
110 and shrub and bush lands) covers about 57% of the country in accordance with the national 2016 map  
111 (WLRC, 2018). This is followed by cultivated land (20%) and grasslands (12%). Barren lands are  
112 estimated to cover about one-tenth of the area of the country while other minor lands with ecological  
113 significance (i.e., wetlands, water bodies and sub-afro-alpine and afro-alpine ) cover about 1.2% of  
114 the country's land mass.

## 115 **2.2 Legacy soil profile data collation and preparation**

116 The soil profile data generated over decades through various soil survey missions were kept in a  
117 variety of formats with limited accessibility. There has been no institution with a mandate to coordinate  
118 the generation, collation, harmonisation, and sharing of soil profile data. This led to the formation of  
119 a group of individuals and institutions who were willing to exchange soil and agronomy data.  
120 Established in 2018, the group known as the Coalition of the Willing (CoW) was committed to  
121 addressing the challenges posed by the lack of the soil and agronomy data access and sharing in the  
122 country (Tamene et al., 2021).

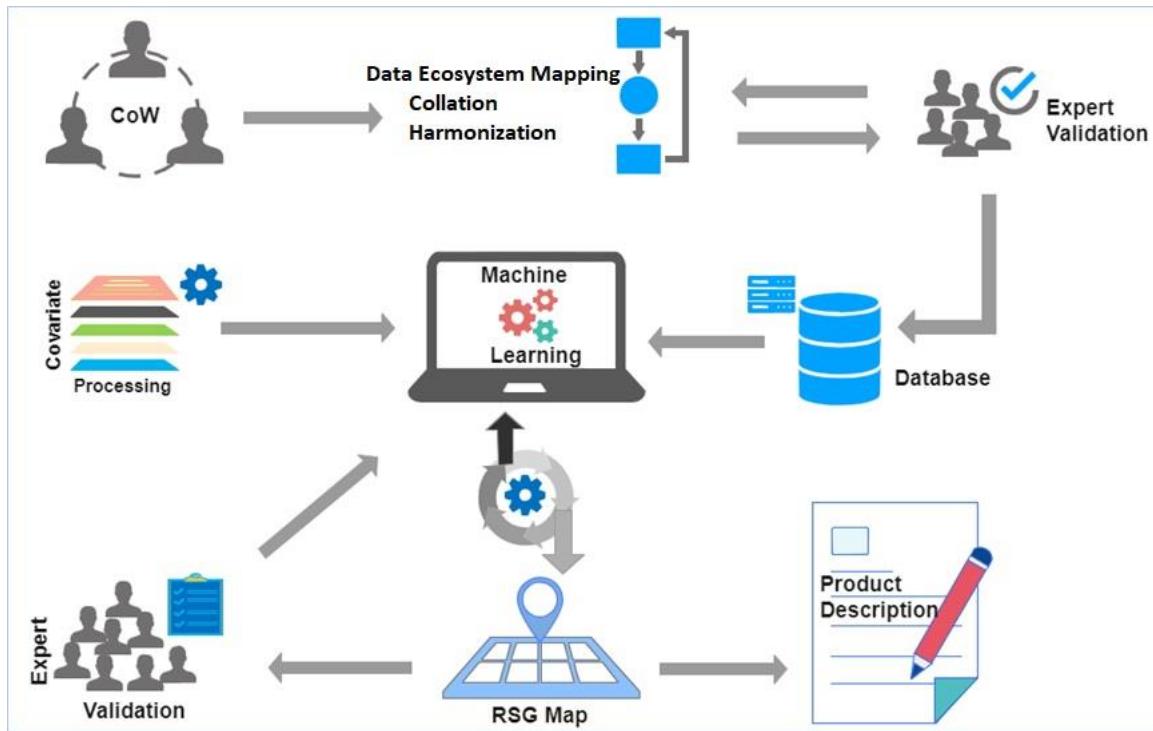
123 The CoW conducted a national soil and agronomy data ecosystem mapping which revealed that a  
124 plethora of legacy soil resource data sets do exist across different institutions and individuals (Ali et  
125 al., 2020). The assessment also revealed that a sizable proportion of the data holders were willing to

126 share the data in their custody, provided that some regulations are put in place to administer the data.  
127 The CoW developed and approved internal data sharing guidelines (CoW, 2020), and facilitated data  
128 collation campaigns, which involved both formal and informal approaches to data holders.

129 Through a data collation campaign, soil profile data collected between the 1970s and 2021 were  
130 acquired from over 88 diverse sources (Ali et al., 2020; Tamene et al., 2021). Initially, 8,000 profile  
131 data points were collated and subjected to improved modelling techniques to create a provisional WRB  
132 reference soil group map of Ethiopia. This was presented to various partners and data-holding  
133 institutions to demonstrate the power of data sharing. This created awareness and enabled us to  
134 mobilise and collate over 20,000 legacy soil profile data. These data were then added to the national  
135 data repository.

136 The data had varying levels of completeness in terms of soil field and environmental descriptions and  
137 laboratory analysis. These required a rigorous expert-based quality assessment and standardisation  
138 before compiling into a harmonised format. The expanded version of the Africa Soil Profile (AfSP)  
139 database (Leenaars et al., 2014) template was used for standardising and harmonising the data. Out of  
140 the collated soil profile data, 14,681 georeferenced data points were extracted based on completeness  
141 and cleanliness for the purposes of modelling. The cleaned soil profile data set contained, at least, the  
142 reference soil group (RSG) nomenclature as outlined in the WRB legend. While the original soil  
143 profile records were set in different coordinate systems, all were projected into the adopted standard  
144 georeferencing system, namely WGS84, decimal degrees in the QGIS (3.20.2) environment (QGIS  
145 Development Team, 2021). To verify their position, soil profile locations were plotted using a standard  
146 WGS84 coordinate system to verify that points are matching with the site description,  
147 geomorphological settings, and at the very least the source project boundary outline.

148 The accuracy of the data depends on the quality and reliability of the survey data itself which in turn  
149 requires expert knowledge and experience in soil description and classification (Leenaars et al.,  
150 2020a). In this study, data cleaning, validation, reclassification, and verification were carried out by a  
151 team of prominent national pedologists and soil surveyors, including those involved in the generation  
152 of some of the soil profile data themselves (Figure 2).



153 **Figure 2.** Schematic presentation of data acquisition and workflow.

154 In addition, the Ministry of Agriculture (MoA) soil survey and mapping experts and other volunteers  
 155 have validated the legacy soil profile observations. This led to the reclassification of the soil types as  
 156 deemed necessary. Such validation and reclassification involved re-examining the geomorphological  
 157 setup of the soil profile locations using Google Earth as well as reviewing the site and soil descriptions  
 158 and the corresponding laboratory data, and reviewing the proposed soil type. The harmonised data sets  
 159 in the database were used as input soil profile data for modelling and mapping IUSS WRB reference  
 160 soil groups.

162 **2.3 Preparation and selection of environmental covariates**

163 **2.3.1 Covariates acquisition and preparation**

164 In order to develop spatially continuous soil class/type maps, data on environmental covariates that  
 165 represent directly or indirectly the soil-forming factors have to be integrated with soil profile data  
 166 (Hengl and MacMillan, 2019). Environmental covariates are spatially explicit proxies of soil-forming  
 167 factors based on the soil-environment relationship (McBratney et al., 2003, Shi et al., 2018).  
 168 Acquisition and preparation of covariates is a crucial step in digital soil mapping using machine

169 learning algorithms (McBratney et al., 2003; Miller et al., 2021). In this study, 68 potential candidate  
170 environmental variables representing soil-forming factors (climate, organisms, relief, parent material,  
171 and time) were derived from diverse remote sensing products and thematic maps (Hengl and  
172 MacMillan, 2019; McBratney et al., 2003).

173 Relief and topography-related covariates were derived from 90-meter Shuttle Radar Topography  
174 Mission (SRTM) digital elevation model (DEM) (Vågen, 2010). Climate-related variables including  
175 long-term mean, minimum, maximum, and standard deviation temperature, and precipitation data for  
176 the period between 1983 and 2016 (Dinku et al., 2014) were acquired from Enhancing National  
177 Climate Services (ENACTS-NMA) initiatives with 4 km resolutions (Dinku et al., 2014). Moderate  
178 Resolution Imaging Spectroradiometer (MODIS) imagery raw bands and derived indices (Vågen,  
179 2010), were downloaded from USGS EarthExplorer (<https://earthexplorer.usgs.gov/>) to represent  
180 vegetation-related factors. National geological (Tefera et al., 1996), and land use and land cover  
181 (WLRC-AAU, 2018) thematic maps of Ethiopia were gathered to represent parent material and  
182 organisms, respectively.

183 Downscaling (disaggregating) or upscaling (aggregating) of rasters were also performed to match the  
184 target resolution. A 250 m spatial resolution was chosen to accommodate both the spatial resolution  
185 of the major covariates inputs and make it applicable for large-scale analysis. All layers were masked  
186 for buildings and water bodies by the national boundary of Ethiopia and a stacked layer was created  
187 using raster package (R Core Team, 2020) to extract covariate values at the locations of soil profiles.  
188 One-hot encoding technique using dummyVars function available in Caret package (Kuhn, 2008) was  
189 used to pre-process and convert categorical covariates into a binary vector. Each element of the binary  
190 vector represents the presence or absence of that category. One-hot encoding is beneficial because it  
191 allows machine-learning algorithms to interpret categorical variables as numerical features. The  
192 covariate pre-processing, visual inspection for inconsistencies, and resampling to a target grid of 250  
193 m were conducted in QGIS [3.20.2] (QGIS Development Team, 2021), SAGA GIS [7.8.2] (Conrad et  
194 al., 2015) and R [version 4.05] (R Core Team, 2020) software packages. All input data were projected  
195 to a common Lambert azimuthal equal-area projection with the latitude of origin 8.65 and centre of  
196 meridian 39.64 which is the centre point for Ethiopia. This projection was selected since it is effective  
197 in minimising area distortions over land. Each covariate was adjusted to have an identical spatial  
198 resolution, extent and projection using two resampling methods. Continuous covariates were

199 resampled using the bilinear spline method, whereas categorical covariates were resampled using the  
200 nearest neighbour method.

201 **2.3.2 Covariates' selection**

202 Selecting an optimal set of covariates for effectively represent the soil–environment relationship is a  
203 key step in Digital Soil Mapping (DSM) since improper selection of covariates will affect the quality  
204 of model outputs (Shi et al., 2018; Huang et al. 2020). In this study, near-zero variance assessment  
205 was conducted using `nearZeroVar` function available in R *caret* package (Kuhn, 2008) to identify and  
206 remove environmental variables that have little or no variance. In addition, preliminary Random Forest  
207 model training was performed to assess and identify covariates having high variable importance. After  
208 expert judgement, a total of 27 environmental variables (24 continuous and 3 categorical) were  
209 selected for modelling and predicting Reference Soil Groups.

210 **2.4 Modelling and mapping soil types/reference soil groups**

211 **2.4.1 Model tuning and quantitative evaluation**

212 In digital soil mapping, machine-learning techniques have been extensively used to determine the  
213 relationship between soil types and environmental variables (McBratney et al., 2003). Many machine-  
214 learning models were developed in the past decades for digital soil mapping to spatially predict soil  
215 classes based on existing soil data and soil-forming environmental covariates (Heung et al., 2016).  
216 Random Forest (RF), a tree-based ensemble method, is one of the most promising machine learning  
217 techniques available for digital soil mapping (Breiman, 2001; Heung et al., 2016), which has gained  
218 popularity due to its high overall accuracy and has been widely used in predictive soil mapping  
219 (Brungard, 2015; Hengl et al., 2018). Examples of the main strengths of the RF model are its ability  
220 to handle numerical and categorical data without any assumption of the probability distribution; and  
221 its robustness against nonlinearity and overfitting (Breiman, 2001; Svetnik et al., 2003). While  
222 building the RF model, data was split into training (80 %) and testing (20 %) components using  
223 random sampling for training the model and evaluating its performance, respectively (Kuhn, 2008).  
224 Hyper-parameter optimization and repeated cross-validation on the training dataset were performed  
225 for optimal model application using the `ranger` method of *Caret* package. The three tuning parameters  
226 for `ranger` method are `mtry`, `splitrule`, and `.min.node.size`. Generally this function is used to tune the  
227 parameters in modelling in an automated fashion, as this will automatically check all the possible

228 tuning parameters and return the optimised parameters on which the model gives the best accuracy.  
229 Model tuning was performed with a repeated 10-fold cross-validation procedure applying multiple  
230 combinations of hyper-parameters for the ranger method. This is a fast implementation of RF  
231 particularly suited for high-dimensional data (Wright and Ziegler, 2017). Then the number of  
232 covariates used for the splits (mtry), splitting rules (splitrule) and minimum node size (min.node.size)  
233 were optimised. The parameter ntree was adjusted to 1,000 in the model, and mtry values (10, 15, 20),  
234 min.node.size values (5, 10, 15), and splitrule values (“variance”, “extratrees”, and “maxstat”) were  
235 fed for the optimization procedure. The accuracy of the testing dataset was related to the model  
236 performance for the new dataset, indicating the capacity of the model to predict at the unsampled  
237 location. A confusion matrix was also used to calculate a cross-tabulation of observed and predicted  
238 classes with associated statistics i.e., producer’s accuracy and user’s accuracy.

#### 239 **2.4.2 Software and computational framework**

240 In this study, various open-source software packages that provide a comprehensive set of tools and  
241 diverse capabilities were used for data preparation, analysis and visualisation. Data pre-processing and  
242 preparation were performed using QGIS (QGIS Development Team, 2021) and SAGA GIS (Conrad  
243 et al., 2015). For statistical analysis and machine learning modelling, R (R Core Team, 2020) and  
244 relevant libraries were installed on a Windows server 2016 standard with 250 GB of working memory  
245 to handle the challenges associated with large-scale data processing and analysis.

#### 246 **2.4.3 Expert evaluation of spatial patterns of the beta-version soil map**

247 Visual inspection of the DSM output over the terrain was used to identify abnormalities and assess  
248 how effectively it depicts landscape components (Rossiter et al., 2022). For this, we employed an  
249 expert-based qualitative assessment of the model output. This technique was used to complement  
250 model-based accuracy assessment and confirm agreement or indicate areas of concern. This was  
251 implemented by a panel of senior soil specialists/pedologists checking the map based on purposely  
252 selected district level geographic windows across Ethiopia, representing different agro-ecological  
253 zones known to have diverse soil occurrences, and familiar to the panel of experts. Accordingly, an  
254 expert validation workshop was conducted using the first version of the reference soil groups (RSGs)  
255 map. About 45 multi-disciplinary scientists including soil surveyors, pedologists, geologists, and  
256 geomorphologists were drawn from national and international research, development, and higher

257 learning institutions to review the draft RSG map in plenary. This was followed by breakout sessions  
258 where groups of experts evaluated the map based on their experience and knowledge of soil-landscape  
259 relations of the country and examined geographic windows.

260 Most importantly, disagreements regarding RSGs occurrence and patterns of the modelling outputs  
261 across topo-sequences and contrasting soil-forming factor sequences were identified and discussed.  
262 Further, inferences on parts of the DSM framework that require improvement were recommended.  
263 After finalising the evaluation at the group's level assessment, each group presented the results in the  
264 plenary followed by a discussion to get feedback from other participants. Following the plenary  
265 discussions, the participants created a group of six senior pedologists to work on the recommendations  
266 including changing the quality mask layer, validation of the additional data obtained during the event,  
267 and assessment of re-modelling outputs.

268 After the second model was re-run, the group of senior pedologists together with geospatial experts  
269 re-evaluated the output using the selected districts based on the feedback from the first review, which  
270 was mainly on areas where there were “minor” and “major” concerns. Consequently, some  
271 improvements were made e.g., in the areas where Vertisols, Fluvisols, and Leptosols were  
272 overestimated. Further, underestimated RSGs (Alisols, Solonetz, Planosols, Acrisols, Lixisols,  
273 Phaeozems, and Gleysols) showed a slight increase in area coverage and pattern improvements.  
274 However, the total area of Leptosols and Cambisols increased from the first run due to the partial  
275 exclusion of the mask layer used in the first round of modelling. The mask layer used in the first run  
276 was criticised for quality issues as it excluded significant soil areas and due to its weakness in  
277 capturing non-soil areas such as rock outcrops, salt flats, swamps and sand dunes. Nevertheless, the  
278 spatial patterns of these soils occurring across previously considered “non-soil areas” were examined  
279 by the panel of experts. In parallel, geospatial and soil experts checked the raster map of the RSGs in  
280 the GIS environment to ensure areas with ‘no concern’ before re-running the model are kept the same  
281 or changes are accepted by the panel of experts. The map from the second run is presented in this  
282 paper as EthioSoilGrids version 1.0 product.

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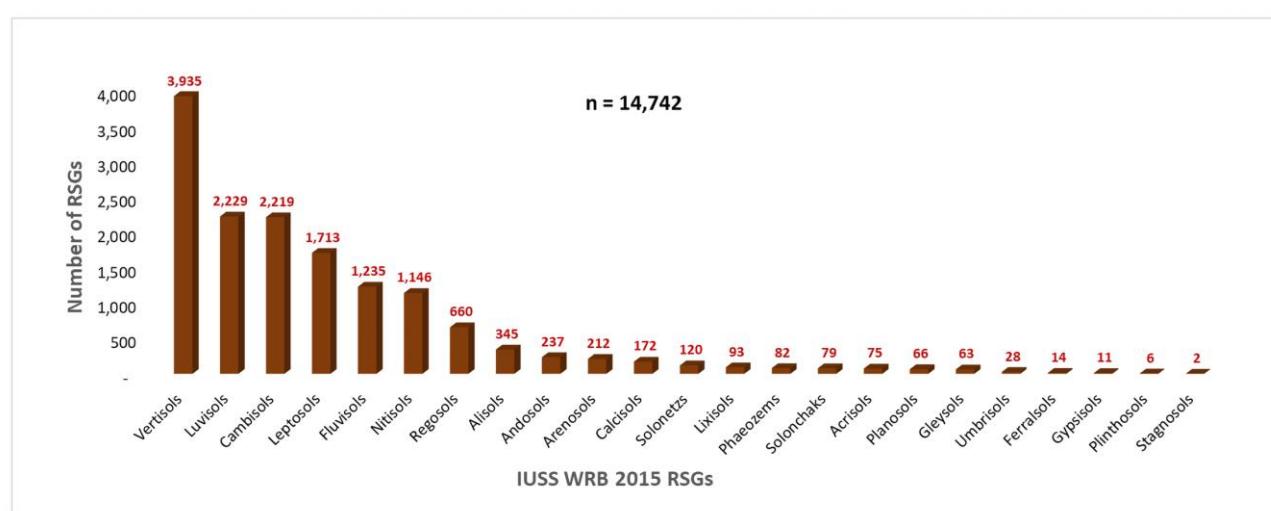
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285 **3 Results and Discussion**

286 **3.1 Soil profile datasets**

287 Using the IUSS WRB, 2015, the preliminary identified 14,742 georeferenced legacy soil profiles were  
288 classified/reclassified into twenty-three reference soil groups (RSGs). Nearly 90% of the soil profile  
289 points represented Vertisols, followed by Luvisols, Cambisols, Leptosols, Fluvisols, and Nitisols,  
290 which were found to be the dominant soil types in Ethiopia (Figure 3). The remaining 10% represented  
291 the Regosols, Alisols, Andosols, Arenosols, Calcisols, Solonetz, Lixisols, Phaeozems, Solonchaks,  
292 Acrisols, Planosols, Gleysols, Umbrisols, Ferralsols, Gypsisols, Plinthosols, and Stagnosols.

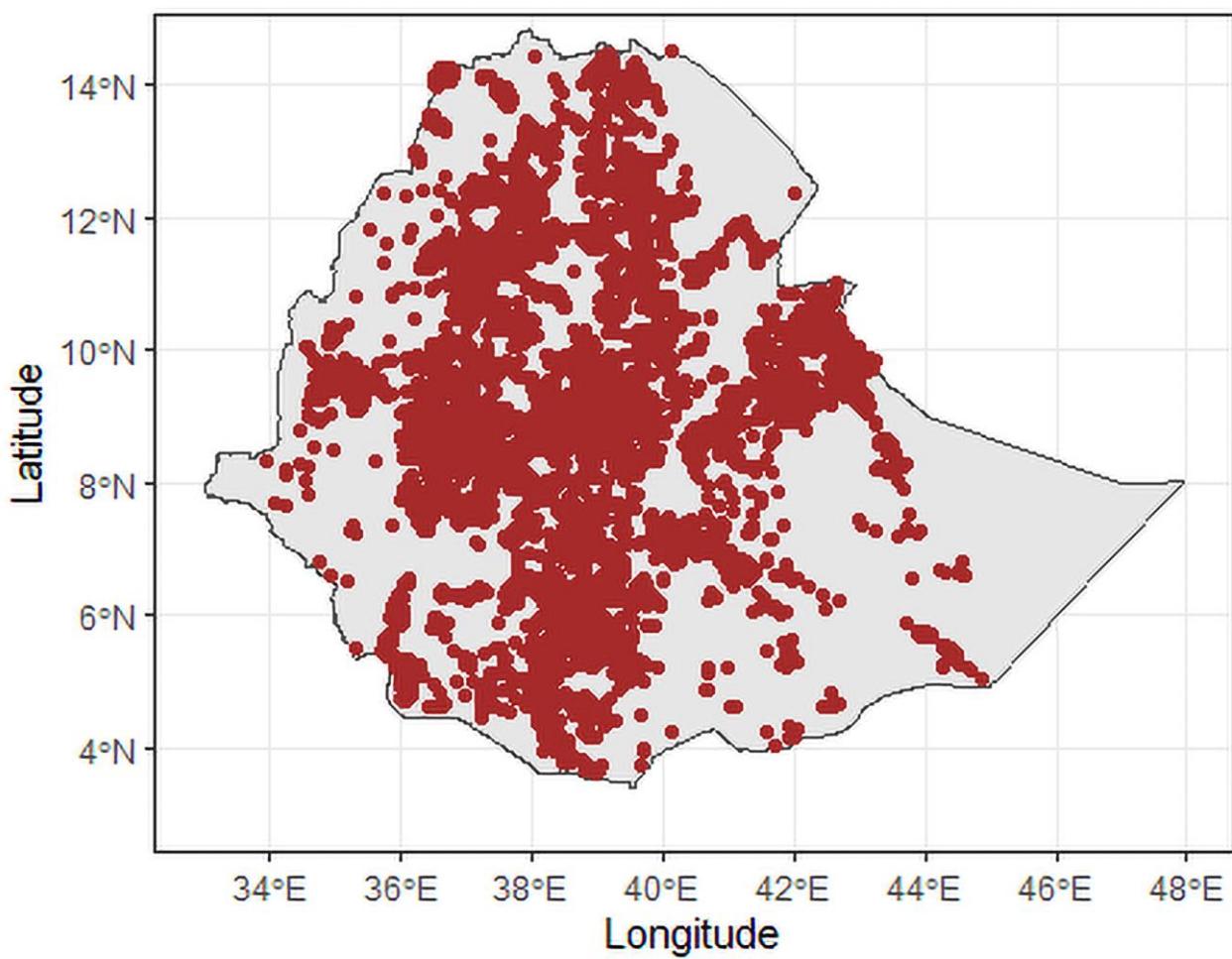
293 According to this study, about 72% of the IUSS WRB (2015) RSGs were confirmed to occur in  
294 Ethiopia. This reconfirms the characterization of Ethiopia as a land of soil diversity having endowed  
295 with a diverse range of soil types (Elias, 2016; Mishra et al., 2004). One of the limitations with legacy  
296 soil data in categorical mapping is the imbalanced soil samples, in that all classes were not equally  
297 represented (Wadoux et al., 2020). For this study, soil profiles with less than 30 observations were  
298 objectively excluded from the model after examining the accuracy and spatial distribution of each  
299 reference soil group. Five reference soil groups (Umbrisols, Ferralsols, Gypsisols, Plinthosols, and  
300 Stagnosols) were excluded from the model and the EthioSoilGrids version 1.0 map.



302 **Figure 3.** Number of soil profile points per WRB reference soil groups.

303 After excluding the built-up) and water surface areas the average soil profile density was 13.1 per  
304 1,000 km<sup>2</sup> (Figure 4), but the actual density varied across the different parts of the country. The  
305 variation tends to follow river basins, sub-basins, and agricultural land-use types-based studies from  
306 which most of the legacy data were pulled.. For instance, in 30 intervention districts of the Capacity  
307 Building for Scaling up of Evidence-Based Best Practices in Agricultural Production in Ethiopia  
308 (CASCAPE) project, the average profile density was about 87 profiles per 1,000 km<sup>2</sup> for a total area  
309 of about 26,830 km<sup>2</sup> (Leenars et al., 2020a). Similarly, semi-detailed soil mapping missions in 15  
310 districts conducted through the Bilateral Ethiopia-Netherlands Effort for Food, Income and Trade  
311 (BENEFIT)-REALISE project generated about 217 observations per 1,000 km<sup>2</sup> (Leenars et al.,  
312 2020b).

313 A soil type and depth map compilation and updating mission at a 1:250,000 scale by the Water Land  
314 Resource Centre (WLRC) of Addis Ababa University collated and used about 3,949 legacy soil  
315 profiles for the entire country (Ali et al., 2020), which is about 3.5 profiles per 1,000 km<sup>2</sup>. Although  
316 the distribution is not even and the eastern lowlands are sparsely represented, the number of data used  
317 in this study is 8.5 times higher than the 1,712 legacy soil profiles data currently existing in the Africa  
318 soil profile database (Batjas et al., 2020; Leenaars et al., 2014).



320 **Figure 4.** Spatial distribution of collated legacy soil profile data.

321 The soil profiles distribution across the 32 agro-ecological zones (AEZ) of Ethiopia revealed that all,  
 322 except two—tepid per-humid mid-highland (0.13% landmass) and very cold sub-humid sub-afro alpine  
 323 to afro-alpine (0.03% landmass)—were represented by soil profile observations. Furthermore, about  
 324 95% of the profile observations represented 91% of the AEZs aerial coverage (Appendix A). The  
 325 distribution of legacy soil profiles varied across AEZs. In general, the top-ranked lowland AEZs with  
 326 roughly 56% area coverage were represented by 23% of the total profile observations, whereas top-  
 327 ranked highland AEZs with 20% area coverage received 47% of profile observations. For instance,  
 328 warm desert, warm moist, hot arid, and warm sub-moist lowlands with area coverage of around 20%,  
 329 15%, 11%, and 10%, were represented roughly by 3%, 11%, 2%, and 7% of the total profiles,  
 330 respectively. Tepid moist mid highlands (8% area coverage), tepid sub-humid mid highlands (7% area

331 coverage), and tepid sub-moist mid highlands (5% area coverage) each were represented by 20%,  
332 15%, and 12% of the profiles, respectively.

333 **3.2 Modelling and Mapping**

334 **3.2.1 Variable importance**

335 The reference soil group spatial pattern is primarily influenced by long-term average surface  
336 reflectance, flow-based DEM indices, and precipitation. Figure 5 shows variables of importance for  
337 determining RSGs spatial prediction. The top-ranked variables were (i) long-term MODIS Near-  
338 Infrared (NIR) reflectance; (ii) multiresolution index of valley bottom flatness, (iii) long-term mean  
339 day-land surface temperature; (iv) long-term mean soil moisture; (v) standard deviation of long-term  
340 precipitation; (vi) long-term mean precipitation; and (vii) topographic wetness index.

341 MODIS long-term mean spectral signatures showed high relative importance. According to Hengl et  
342 al. (2017), accounting for seasonal vegetation fluctuation and inter-annual variations in surface  
343 reflectance, long-term temporal signatures of the soil surface, derived as monthly averages from long-  
344 term MODIS imagery, were more effective. Furthermore, Hengl and MacMillan (2019) explained that  
345 long-term average seasonal signatures of surface reflectance provide a better indication of soil  
346 characteristics than only a single snapshot of surface reflectance.

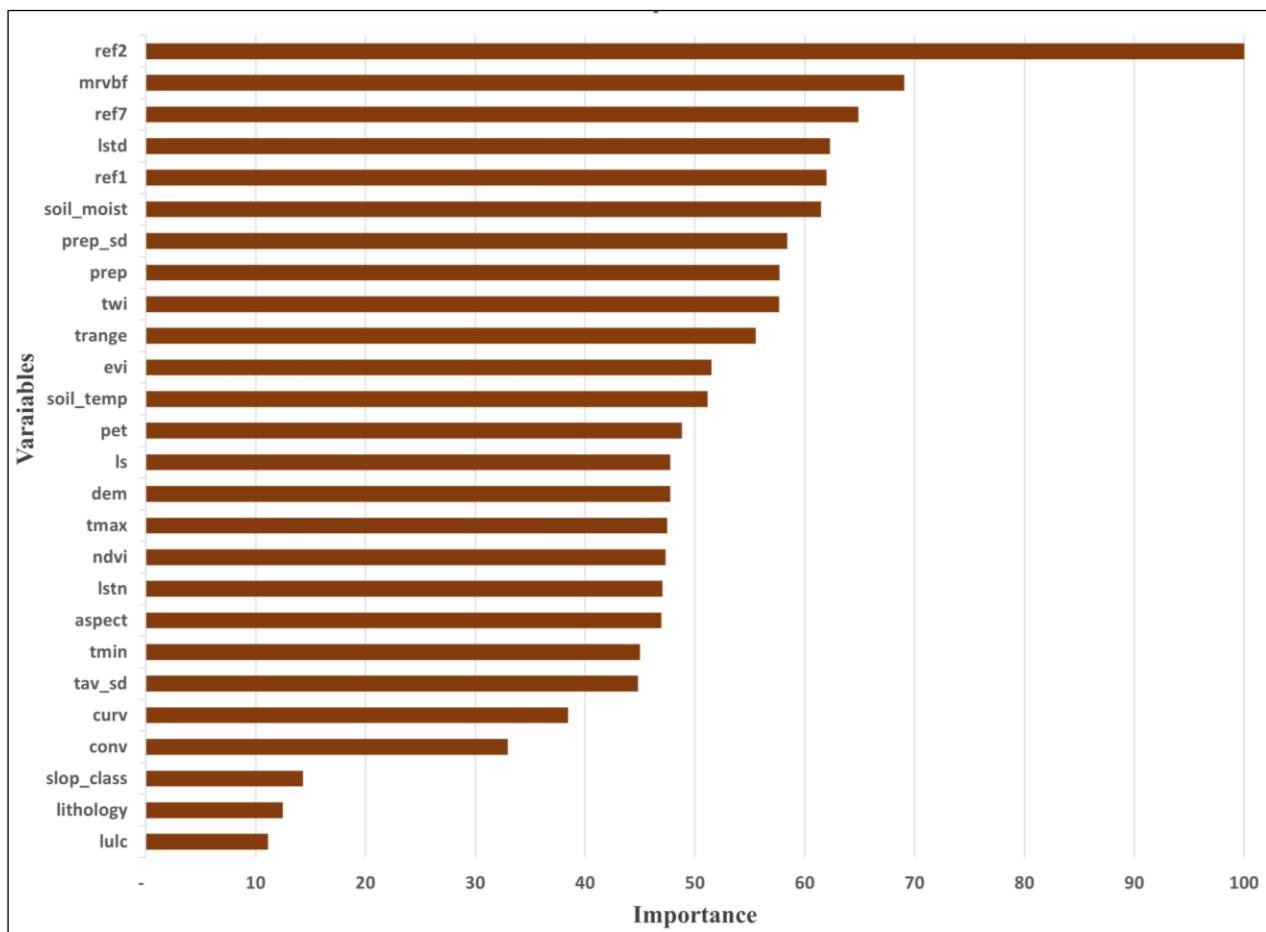
347 The Multi-Resolution Valley Bottom Flatness Index, a DEM-derived topography index, is the second  
348 top-ranked covariate driving soil variability across Ethiopia. This hydrological/soil removal and  
349 accumulation/deposition index is used to distinguish valley floor and ridgeline landscape positions  
350 (Soil Science Division Staff, 2017) highly responsible for multiple soil-forming processes to operate  
351 over a particular landscape, resulting in a wide range of soil development. The influence of topography  
352 on spatial soil variation is manifested in every landscape of Ethiopia (Belay, 1997; Mesfin, 1998;  
353 Nyssen et al., 2019; Zewdie, 2013).

354 Long-term daily mean land surface temperature, mean soil moisture, rainfall standard deviation and  
355 mean annual rainfall were among the top-ranked covariates for predicting reference soil groups' spatial  
356 variation across the country. In Ethiopia, different soil genesis studies revealed that climate has a  
357 significant influence on soil development and properties and is, therefore, responsible for having  
358 widely varying soils in the country (Abayneh, 2006, 2005; Fikru, 1988, 1980; Zewdie, 2013).

359

360 Among the most important covariates for predicting reference soil groups in the Ethiopian highlands,  
 361 are monthly average soil moisture for January (ranked 3<sup>rd</sup>), long-term average soil moisture (ranked  
 362 4<sup>th</sup>), and monthly average soil moisture for August (ranked 5<sup>th</sup>) (Leenars et al., 2020a). In the current  
 363 study, soil moisture was among the ten top ranked covariates in modelling and explaining long-  
 364 distance soil type variability across the country.

365



366 **Figure 5.** Random forest covariate relative importance for modelling RSGs.

367 Note: prep=Precipitation; prep\_sd=The standard deviation of precipitation; tmax=Maximum  
 368 Temperature; tmin=Minimum Temperature; trange=Temperature range; tav\_sd=Standard deviation of  
 369 average temperature; pet=Potential evapotranspiration; lstd=Land surface temperature- Day;  
 370 lstdn=Land surface temperature-Night; soil\_moist=Soil Moisture ; soil\_temp=Soil temperature; DEM  
 371 =Digital elevation model (Elevation); twi =Topographic wetness Index; aspect=Topographic Aspect;  
 372 curv=Topographic Curvature; conv=Topographic convergence index; ls=Slope Length and Steepness

373 factor (ls\_factor); morph=Terrain Morphometry; mrvbf=Multiresolution index of valley bottom  
374 flatness; slope=Slope class (%); ndvi=Normalised Difference Vegetation Index (NDVI);  
375 evi=Enhanced Vegetation Index (EVI); lulc=Land use/ landcover; lithology=Geology; ref1=Red band  
376 ;ref2=Near-Infrared; ref7=Mid-Infrared.

377 In this study, lithology showed a relatively low influence on soil variability may be due to the use of  
378 a coarse-scale and less detailed lithology map, which may not sufficiently capture the spatial  
379 variability of the parent materials.

### 380 **3.2.2 Model performance**

381 The parameter optimization process resulted in mtry = 20, split rule= extra trees and minimum node  
382 size= 5. The overall accuracy of the model was 56.24% which ranged between 54.43% and 58.1%  
383 with a 95% confidence interval. The kappa values based on the internal cross-validation and testing  
384 dataset showed that the overall model performance produced using 10-fold cross-validation with the  
385 repeated fitting was 48%. Considering similar area-based digital soil class mapping efforts, the overall  
386 accuracy was in line with the accuracies that were typically reported for soil class maps developed  
387 with random forest model (Leenaars et al., 2020a) and statistical methods (Heung et al., 2016; Holmes  
388 et al., 2015). Table 1 shows the confusion matrix at validation/testing points i.e., 20 % of the  
389 observation. Further, the matrix indicates the producer's accuracy (class representation of observed  
390 versus predicted) and user's accuracy were not similar for all RSGs. The map purity is in the order of  
391 Lixisols, Calcisols, Alisols, Phaeozems, Vertisols, Andosols, Solonchaks, Fluvisols, Arenosols,  
392 Leptosols, Luvisols, Nitisols, and Cambisols. However, Vertisols, Calcisols, and Andosols are the  
393 observed classes that are best represented by the map followed by Fluvisols, Alisols, Nitisols,  
394 Leptosols, Luvisols and Cambisols.

395 Global Soil Grids at 250 m resolution used machine learning algorithms to map the global WRB  
396 reference soil groups with map purity and weighted kappa of 28% and 42%, respectively (Hengl et al.,  
397 2017). The Soil Grids 250 m WRB soil groups/classes prediction output-spatial soil patterns were not  
398 evaluated based on expert knowledge while in this study we did an extensive back and forth qualitative  
399 assessment by a panel of pedologists. The quantitative accuracy in the present study (about 56%)  
400 coupled with an expert-based qualitative evaluation of the predicted maps indicated the development  
401 and achievement of a substantially enhanced national product for users of spatial soil resource

402 information. This finding is a step forward and acceptable considering that Soil Grids are not expected  
403 to be as accurate as locally produced maps and models that use much more local point data and finer  
404 local variables (Mulder et al., 2016). Further, the data and findings in this study can help improve the  
405 soil maps of Africa as it partially addresses the concern by Hengl et al. (2017) who recognised that  
406 WRB RSGs modelling in the global Soil Grids 250 m is critically uncertain for parts of Africa. This  
407 is mainly attributed to limited access to more local point data by regional and global modelling  
408 initiatives, unlike the present study which accessed a large number of legacy soil profile datasets.

409

410

411 **Table 1.** Confusion matrix of random forest RSG prediction (at validation/testing observations).  
412

Prediction	Reference																	User Accuracy		
	Acrisols	Alisols	Andosols	Arenosols	Calcisols	Cambisols	Fluvisols	Gleysols	Leptosols	Lixisols	Luvisols	Nitisols	Phaeozems	Planosols	Regosols	Solonchaks	Solonetzs	Vertisols		
Acrisols	1	0	0	0	0	0	0	0	0	0	1	0	0	0	0	0	1	0	0.33	
Alisols	0	40	0	0	0	0	1	1	0	0	9	4	0	0	2	0	0	2	59	
Andosols	0	0	28	1	1	3	5	0	2	0	2	0	0	0	0	0	1	1	0.68	
Arenosols	0	0	0	11	0	2	1	0	0	0	5	0	0	0	0	0	0	1	0.64	
Calcisols	0	0	0	0	21	0	1	0	0	0	2	0	0	0	0	0	0	5	0.55	
Cambisols	2	3	6	9	1	197	28	2	35	2	47	16	5	1	16	3	3	28	0.72	
Fluvisols	1	0	3	5	1	34	144	0	9	0	15	7	0	0	1	5	5	17	0.49	
Gleysols	0	0	0	0	0	0	1	2	0	0	1	0	0	1	0	0	0	0	0.58	
Leptosols	0	1	4	3	3	47	11	0	176	0	27	7	1	0	32	0	0	24	0.40	
Lixisols	0	0	0	0	0	0	0	0	0	1	0	0	0	0	0	0	0	0	1	
Luvisols	2	16	3	8	0	34	13	2	33	3	216	30	3	0	25	1	0	41	1.00	
Nitisols	6	8	0	0	1	23	8	3	18	8	29	132	0	1	8	0	1	21	0.50	
Phaeozems	0	0	0	0	0	0	0	0	0	0	1	0	2	0	0	0	0	0	0.49	
Planosols	0	0	0	0	0	0	0	0	0	0	1	1	0	5	1	0	0	1	0.67	
Regosols	0	0	0	0	0	7	1	0	7	1	8	1	0	0	22	0	0	5	0.55	
Solonchaks	0	0	0	0	0	0	0	0	1	0	0	0	0	0	0	3	1	0	0.42	
Solonetzs	0	0	0	0	1	4	1	0	0	0	0	0	0	0	0	1	6	0	0.60	
Vertisols	3	1	3	5	5	92	32	2	61	3	81	31	5	5	25	2	6	641	0.46	
Producer Accuracy	0.07	0.58	0.60	0.26	0.62	0.44	0.58	0.17	0.51	0.06	0.49	0.58	0.13	0.38	0.17	0.20	0.25	0.81	0.64	
Total	15	69	47	42	34	443	247	12	342	18	445	229	16	13	132	15	24	787	-	2,930

413 **3.2.3 Modelling and Mapping: EthioSoilGrids Version 1.0**

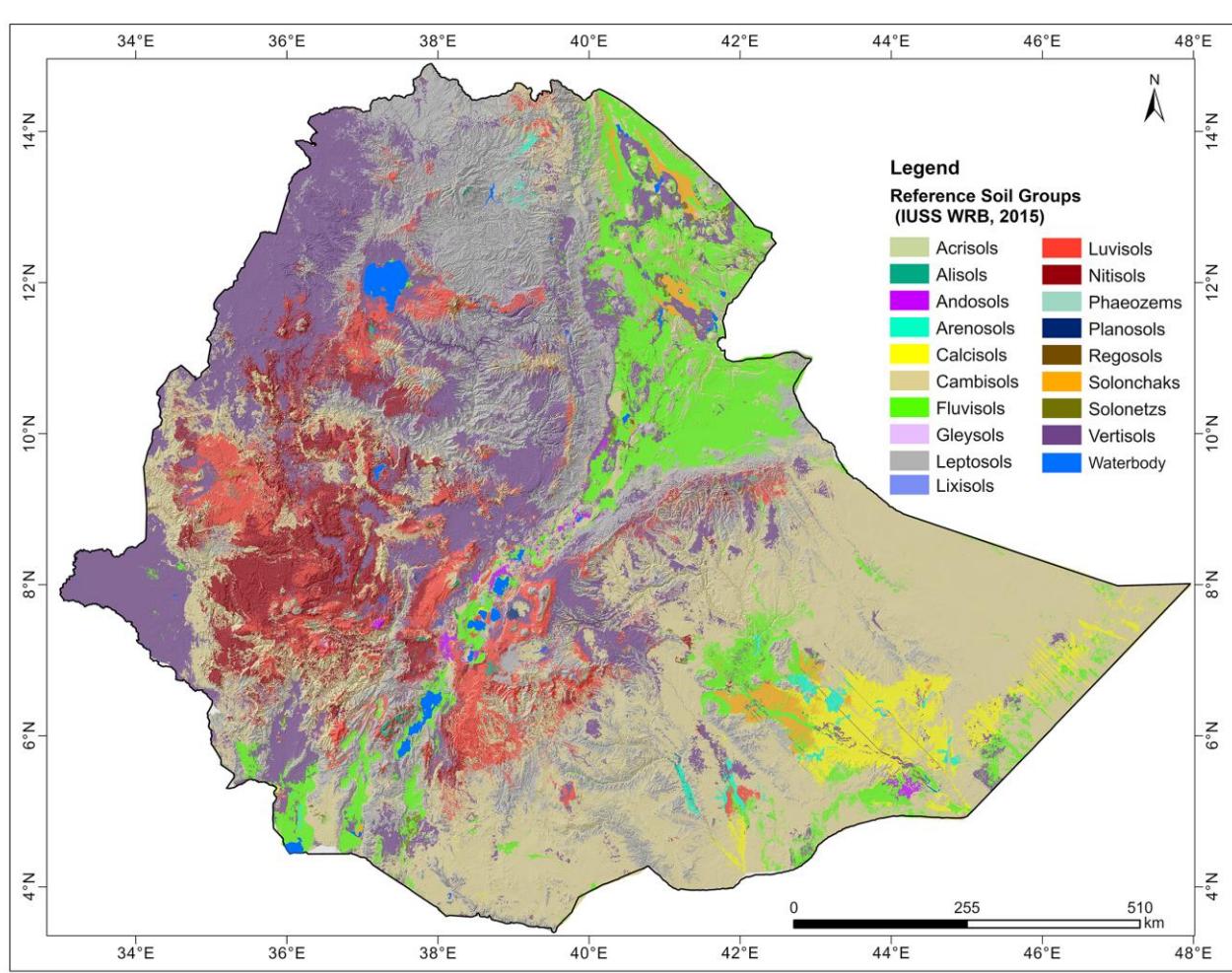
414 The study identified eighteen reference soil groups in Ethiopia, mapped at 250 m resolution (Figure  
415 6). The model prediction showed that seven soil reference groups including Cambisols, Leptosols,  
416 Vertisols, Fluvisols, Nitisols, Luvisols, and Calcisols covered nearly 98% of the total land area of the  
417 country (Figure 7). Five soil reference groups (Solonchaks, Arenosols, Regosols, Andosols, and

418 Alisols) were estimated to cover about 2% of the land area, while trace coverages of Solonetz,  
419 Planosols, Acrisols, Lixisols, Phaeozems, and Gleysols were also found in some pocket areas.

420 In terms of spatial distribution, Nitisols and Luvisols dominated the northwestern and southwestern  
421 highlands while the southeastern lowlands were dominantly covered by Cambisols, Calcisols, and  
422 Fluvisols with some Solonchaks. The Vertisols extensively cover the north and south-western  
423 lowlands along with the Ethio-Sudan border areas and central highland plateaus. The probability of  
424 occurrence of each RSG was mapped (Appendix C) in each modelling spatial window (i.e., the cell  
425 size of 250-meter X 250 m). The dominant RSGs were aggregated based on the most probable RSGs  
426 in each spatial modelling window. There was high correspondence between the top seven ranked  
427 prediction probabilities and observed soil types as confirmed visually by overlaying observed classes  
428 and prediction probabilities.

429 The overall occurrence and the relative position of each of the RSG along the topo-sequence and its  
430 association with other RSGs agree with previous works (Abayneh, 2006; Ali et al., 2010; Abdenna et  
431 al., 2018; Asmamaw and Mohammed, 2012; Belay, 2000, 1998, 1997, 1996; Driessen et al., 2001;  
432 Elias, 2016; FAO 1984a; Fikre, 2003; Mitku, 1987; Mohammed and Belay, 2008; Mohammed and  
433 Solomon, 2012; Mulugeta et al., 2021; Nyssen et al., 2019; Sheleme, 2017; Shimeles et al., 2007;  
434 Tolossa, 2015; Zewdie, 2013). However, in some cases, the RSGs' position along the topo-sequence  
435 and association with other RSGs require further investigation. The observed disparities might be  
436 attributed to the positional accuracy of legacy point observations, modelling approach, and most  
437 importantly the level of detail and scale/resolution of the environmental variables used in this study.  
438 We used the currently available coarse resolution national geological map and hence soil parent  
439 material might be inadequately represented in the model, which probably resulted in irregular RSGs  
440 sequences. For instance, the main driving factors to establish and explain soil-landscape variability in  
441 May-Leiba catchment of northern Ethiopia were geology (soil parent material) and different mass  
442 movements (Van de Wauw et al., 2008). These factors led to Cambisols–Vertisols catenas  
443 on basalt and Regosols–Cambisols–Vertisols catenas on limestone formations. Similar studies  
444 identified parent material strongly determines the soil type (e.g. Vertisol, Luvisol, Cambisol) (Nyssen  
445 et al., 2019). In general, in areas where there is complex soil diversity and distribution of soils, one of  
446 the most important parameters is to identify parent material including effective techniques to capture

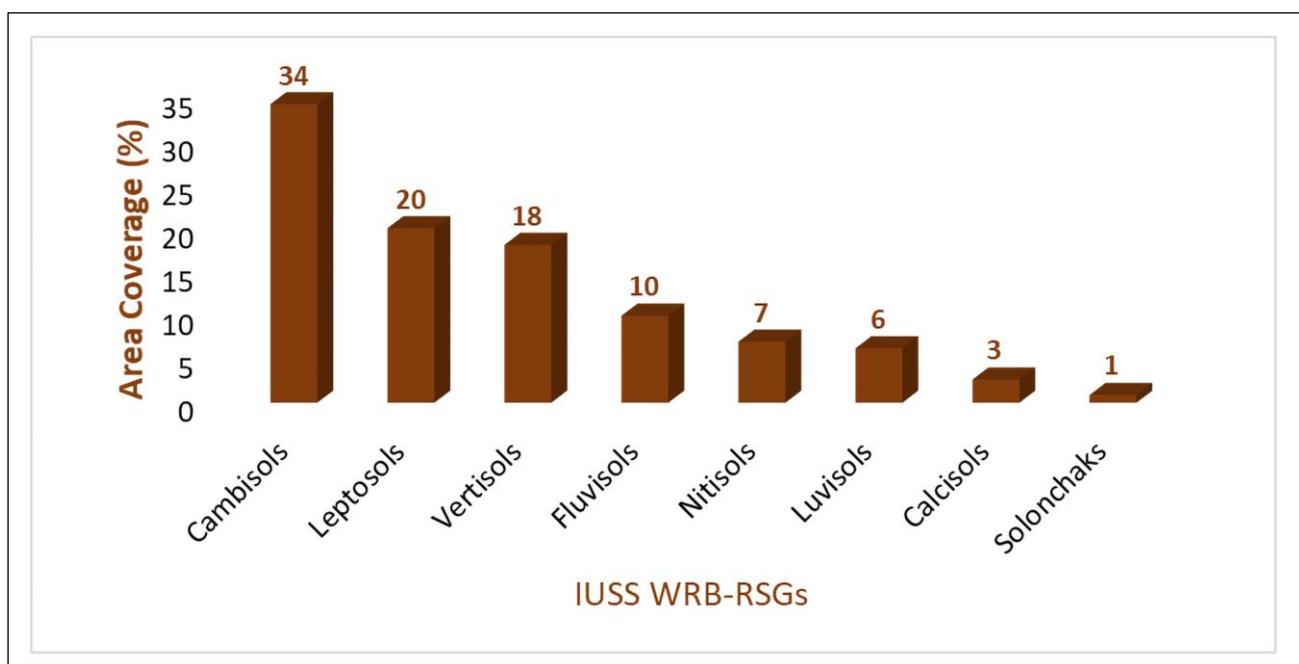
447 and delineate mass movement bodies, and human-induced soil erosion and deposition areas (Leenars  
448 et al., 2020a; Nyssen et al., 2019; Van de Wauw et al., 2008).



450 **Figure 6.** Major reference soil groups of Ethiopia (EthioSoilGrid V1.0).

451 Considering the third position of Cambisols in the order of frequency occurrence of RSGs per point  
452 observations (following Vertisols and Luvisols), these soils seem to be over-represented on the map  
453 (ranked 1<sup>st</sup>) apparently at the expense of Vertisols and Luvisols, and to some extent in places of  
454 Leptosols and other RSGs. This might be attributed to the fact that Cambisols create a geographical  
455 continuation with Vertisols and/or Luvisols at the lower slopes and Leptosols/ Regosols at the higher  
456 slopes, suggesting the presence of some bordering soil qualities in respective transitional zones (Ali et  
457 al., 2010; Asmamaw and Mohammed, 2012; Sheleme, 2017; Zewdie, 2013).

458 The proportion of area mapped as Cambisols (34 %) revealed new insights compared with the  
459 information from the most cited spatial soil maps: Cambisols ranked 2<sup>nd</sup> (21 %), 2<sup>nd</sup> (16 %), 4<sup>th</sup> (9 %),  
460 and 4<sup>th</sup> (8 %) as reported by Berhanu (1980), FAO (1984b), FAO (1998), and Soil Grids- Hengl et al  
461 (2017), respectively. This might be due to: (i) the number and distribution of profile observations,  
462 which is more extensive than the previous ones, (ii) the type and level of details of covariates  
463 considered; (iii) variations and rearrangements in the keys for classification of the RSGs among soil  
464 classification versions used in previous studies and misclassification/confusion of Vertisols with  
465 Vertic Cambisols, as legacy soil profile data coming from diverse sources.

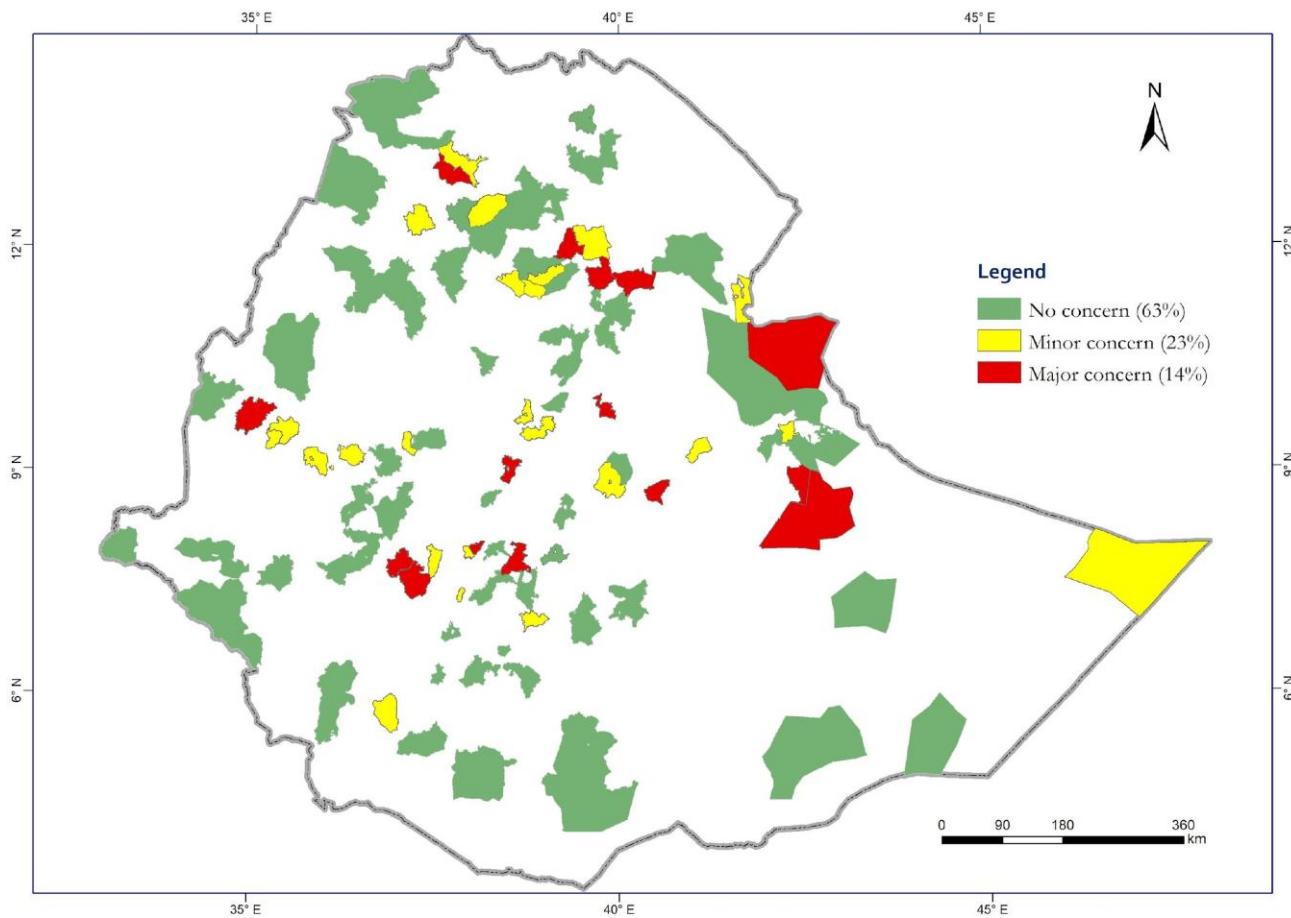


467 **Figure 7.** The area coverage (in %) for the major WRB RSGs (Note: the remaining 10 RSGs-  
468 Arenosols (0.44 %), Regosols (0.35 %), Andosols (0.31 %), Alisols (0.16 %), Solonetz (0.04 %),  
469 Planosols (0.04 %), Acrisols (0.02 %), Lixisols (0.02 %), Phaeozems (0.02 %), and Gleysols (0.01 %)  
470 were not plotted because of their relatively small area coverage).

471

472 **3.3 Expert validation of the soil map**

473 Expert knowledge of soil-landscape relations and soil distribution remains important to evaluate the  
474 predictive soil mapping results and assess if predicted spatial patterns make sense from a pedological  
475 viewpoint (Hengl et al., 2017; Poggio et al., 2020; Rossiter et al., 2022). An important step in  
476 qualitative model evaluation is, therefore, expert assessment whereby professionals with broad  
477 experience in soil survey and mapping can evaluate and improve the quality of the soil resource map.  
478 This can highlight areas of agreement or concern across the landscape (Rossiter et al., 2022). The  
479 expert validation workshop provided useful insights and tangible improvements to the development  
480 of the map. While the plenary discussion provided an overview of the approaches followed in  
481 developing the map, the group discussions helped to have an in-depth review of the selected polygons  
482 of the map assigned to them. Participants were split into five groups (with 8-10 members each) and  
483 have chosen up to 60 polygons representing areas with which at least one of the group members has  
484 sufficient information, including data sources. Overall, the groups have checked a total of 126  
485 polygons (Figure 8) which were fairly distributed across the country.



488 **Figure 8.** The spatial distribution of districts validated by stakeholders and feedback categories  
 489 according to the level of concerns raised.

490 The group members displayed the polygons one by one in a GIS environment and discussed the  
 491 predicted dominant and associated soil reference soil groups and labelled them in one of three  
 492 confirmation categories: 1. confirmed with '*no concern*', 2. confirmed with '*minor concern*', and 3.  
 493 confirmed with '*major concern*'. Confirmation with '*no concern*' was made when all members of a  
 494 group agreed on both the types, relative coverage and patterns of the predicted soils within the polygon.  
 495 Confirmation with '*minor concern*' was made when all or some of the team members agreed on the  
 496 predicted soil types within the polygons but did not agree on the order of abundance or the probability  
 497 occurrence of one or two soils including observed spatial patterns, while confirmation with '*major*  
 498 *concern*' was made when all members of the team did not agree on the predicted soil type, or when  
 499 the presence of another soil type, other than the predicted ones is noted.

500 All three groups have rated the accuracy of the map at 60 +%; of the 126 polygons, they have expressed  
501 no concern for 63 %, minor concern for 23 % and a major concern for 14 % of the polygons.  
502 Furthermore, differences in the prevalence of RSGs and patterns of the modeling outputs across  
503 different soil forming factor sequences, as well as inferences about which areas of the DSM framework  
504 still need work, were identified and elaborated by the expert input, and presented in the subsequent  
505 sections.

### 506 **3.4 Evaluation of results, limitations and future direction**

507 Up-to-date soil resource spatial information is critically missing at a required scale and extent in  
508 Ethiopia. As a result, resource management strategies miss their targets. Furthermore, the absence of  
509 such data at a required resolution and extent, forced decision support tool developers to pick and use  
510 the data they can access and afford. As a result, model outputs appear more site-specific or  
511 representation becomes homogenous over the very heterogeneous landscapes that exist in reality. On  
512 the other hand, in large areas and complex landscapes such as Ethiopia, it is very difficult to address  
513 the demand for reasonably accurate and detailed soil-type maps using a conventional approach due  
514 to the costs involved, and resources and time it requires. For instance, given the vastness of the country  
515 and heterogeneous landscapes, a new conventional soil survey mission requires at least 170,000  
516 profile point observations to map the entire terrestrial land mass of Ethiopia at a scale of 1: 250,000  
517 with at least 1 observations per square centimetre. Moreover, the soil profile data requirement  
518 definitely could have been much higher as we increase the scale of mapping and density of  
519 observations. In the present study, machine-learning techniques combined with expert input were  
520 implemented to produce a countrywide soil resource map of Ethiopia at reasonably higher accuracy,  
521 less time and cost than that of conventional methods. In addition, rescue, compilations and  
522 standardization of about 14,681 geo-referenced legacy soil profiles that can be included in the National  
523 Soil Information System (NSIS) of Ethiopia and the World Soil Information Centre will support future  
524 national, regional and global DSM efforts. The approach used demonstrates the power of data and  
525 analytics to map the soil resources of Ethiopia and the output is an exemplary use case for similar  
526 digital content development efforts in Ethiopia and beyond.

527 Moreover, in this study the quality monitoring processes and methods were followed to filter dubious  
528 soil profiles, and soil classification and harmonization protocols. Then after, the study followed a  
529 robust modelling framework and generated new insights into the relative area coverage of WRB RSGs

530 of Ethiopia. In addition, the study provided coherent and up-to-date digital quantitative gridded spatial  
531 soil resource information to support the successful implementation of various digital agricultural  
532 solutions and decision support tools (DSTs).

533 The spatially explicit limitation of the present study is revealed by expert-based qualitative evaluation  
534 of spatial patterns across objectively selected geographic windows and prominent contrasting  
535 landscapes of Ethiopia. This qualitative assessment indicated areas of concern in terms of how well  
536 EthioSoilGrids version 1.0 represents soil geography across a mosaic of the country's landscapes. For  
537 instance, in the north-eastern lowlands of Ethiopia, mainly along the "Denakil" depression, Fluvisols,  
538 Cambisols and Vertisols were found on the map in areas where normally other soil types were expected  
539 to occur. In this area, the expected prediction and area coverage of Leptosols has been probably  
540 overshadowed by Fluvisols and Cambisols. Similarly, in some parts of western Ethiopia landscapes,  
541 the prediction of Vertisols overshadows other RSGs which resulted in area coverage underestimation  
542 of Fluvisols (along the "Akobo", "Gilo", and "Baro" rivers and their tributaries) and Alisols. Likewise,  
543 in the central parts of northwestern Ethiopia, the prediction of Nitisols has been overshadowed by  
544 Vertisols and Luvisols resulting in probable underestimation of the Nitisols area coverage.

545 The relatively low model performance and some classification errors in some of the examined  
546 geographic windows (e.g. the Denakil depression, along Akobo, Baro, and Gilo rivers and the Somali  
547 region) is, probably due to the paucity of samples from those areas (Figure 4), the inadequacy of the  
548 dataset by RSGs, and over-representation of the dataset by some RSGs such as Vertisols, Luvisols,  
549 and Cambisols. Balanced datasets are ideal to allow a decision tree algorithms to produce better  
550 classification but for datasets with uneven class size, the generated classification model might be  
551 biased towards the majority class (Hounkpatin et al., 2018; Wadoux et al., 2020). In addition,  
552 uncertainty around quality of included covariates, not considered covariates in the modelling process  
553 including management, use of validation methods that do not sufficiently control the effect of clustered  
554 samples, and small sample size for some RSGs could have possibly biased modelling results in some  
555 geographic areas.

556 To improve the modelling performance, future studies could explore (1) adding data for under-  
557 represented geographic areas, land uses and covariate spaces, (2) opportunities to include other  
558 covariates (parent material and management) that could capture the variability of the country  
559 heterogeneous landscapes, (3) dimension reduction of covariates (4) use of remedial measures for

560 imbalances in sample sizes, (5) comparing different cross-validation methods, (6) use of an ensemble  
561 modelling approach and/or robust modelling technique that accommodates neighbourhood size and  
562 connectivity analyses, (7) use of better resolution/quality mask layer to segregate non-soil areas (rock  
563 outcrops, salt flats, sand dunes and water bodies) from mapping areas, and (8) implementation of  
564 quantitative and qualitative comparison of national, regional, and global legacy soil maps/soil grids  
565 with new DSM products in terms of how well DSM products represent soil geography. In addition ,  
566 future digital soil mapping strategies in Ethiopia may require to consider new soil sampling missions  
567 in under-represented areas, adopt standard soil sampling, description guidelines and soil classification  
568 systems including soil physico-chemical and mineralogical analysis, and combine local soil  
569 nomenclature/classification systems with RSGs and develop a map of RSGs with qualifiers. At the  
570 moment the under-sampled and under-represented areas are the Somali region, the Denakil and the  
571 western and northwestern border areas of Ethiopia (Figure 4). Regardless of these limitations and to  
572 the best of our knowledge the EthioSoilGrids v1.0 product provides the most complete soil information  
573 available for Ethiopia.

## 574 **4 Conclusions**

575 Coherent and up-to-date country-wide digital soil information is essential to support digital  
576 agricultural transformation efforts. This study involved collation, cleaning, harmonization, and  
577 validation of the legacy soil profile data sets, involving soil scientists with different backgrounds  
578 individually and in groups. To develop the 250 m digital soil resource map, a machine learning  
579 modelling approach and expert validation were applied to the harmonised soil database and  
580 environmental covariates affecting soil-forming processes. Accordingly, about 20,000 soil profile data  
581 have been collated, out of which, about 14,681 were used for the modelling and mapping of eighteen  
582 RSGs out of the identified twenty-three RSGs. Although unevenly distributed, the legacy soil profile  
583 data used in the modelling covered most of the agro-ecologies of the country.

584 Among the mapped 18 RSGs, the highest number of observed (3,935) profiles represent Vertisols,  
585 followed by Luvisols, Cambisols and Leptosols, while Gleysols were represented with the lowest  
586 number (63) of profiles. The modelling revealed that MODIS long-term reflectance, multiresolution  
587 index of valley bottom flatness, land surface temperature, soil moisture, long-term mean annual

588 rainfall, and wetness index of the landscape are the most important covariates for predicting reference  
589 soil groups in Ethiopia.

590 Our ten-fold spatial cross-validation result showed an overall accuracy of about 56 % with varying  
591 accuracy levels among RSGs. The modelling result revealed that seven major soil reference groups  
592 including Cambisols (34 %), Leptosols (20 %), Vertisols (18 %), Fluvisols (10 %) Nitisols (7 %),  
593 Luvisols (6 %) and Calcisols (3 %) covered nearly 98 % of the total land area of the country, while  
594 minor coverage of other reference soil groups (Solonchaks, Arenosols, Regosols, Andosols, Alisols,  
595 Solonetz, Planosols, Acrisols, Lixisols, Phaeozems, and Gleysols) were also detected in some areas.  
596 Compared to the existing soil resource map, the coverage of the first three major soil groups has  
597 substantially increased which is related to the increased availability of soil profile data covering larger  
598 areas of the country, implying that these soils were previously underestimated. Cambisols and  
599 Vertisols which together represent nearly half of the total land area are relatively young with inherent  
600 fertility, implying the high agricultural potential for the country. However, given their limitations,  
601 these and the other soil types require the implementation of suitable land, water, and crop management  
602 techniques to sustainably exploit their potential.

603 The EthioSoilGrids version 1.0 product from this first countrywide RSGs modelling effort requires  
604 complementary activities. These include modelling and mapping that should go beyond RSGs and  
605 need to include 2<sup>nd</sup> level classifications including principal and supplementary qualifiers. Furthermore,  
606 soil atlas of Ethiopia with details of the soil physicochemical properties needs to be prepared together  
607 with the map, for which the authors and/or others responsible need to prioritize in their future research  
608 endeavours.

609

610

611

612 **Appendix A: Legacy soil profile data distribution**613 **Table A1.** Distribution of legacy soil profile data by agroecology zones.

Major agroecological zones	AEZ area coverage (%)*	Profiles observation (%)**
Warm arid lowland plains	19.76	3.40
Warm moist lowlands	15.12	10.74
Hot arid lowland plains	10.79	2.44
Warm sub-moist lowlands	9.63	6.94
Tepid moist mid highlands	8.05	20.21
Warm sub-humid lowlands	7.11	5.69
Tepid sub-humid mid highlands	6.63	15.26
Tepid sub-moist mid highlands	5.17	12.39
Warm semi-arid lowlands	2.75	3.23
Tepid humid mid highlands	2.65	2.48
Warm humid lowlands	2.29	0.45
Cool moist mid highlands	1.74	4.15
Hot sub-humid lowlands	1.67	0.07
Cool sub-moist mid highlands	1.16	3.00
Cool humid mid highlands	0.82	1.01
Warm per-humid lowlands	0.68	0.01

Major agroecological zones	AEZ area coverage (%)*	Profiles observation (%)**
Hot moist lowlands	0.59	3.56
Hot sub-moist lowlands	0.56	0.03
Cool sub-humid mid highlands	0.52	1.38
Tepid arid mid highlands	0.43	0.39
Hot semi-arid lowlands	0.40	2.05
Tepid semi-arid mid highlands	0.19	0.67
Cold moist sub-afro-alpine to afro-alpine	0.07	0.16
Cold sub-moist mid highlands	0.07	0.04
Cold sub-humid sub-afro-alpine to afro-alpine	0.06	0.03
Cold humid sub-afro-alpine to afro-alpine	0.06	0.01
Very cold humid sub-afro-alpine	0.04	0.02
Very cold sub-moist mid highlands	0.02	0.02
Very cold moist sub-afro-alpine to afro-alpine	0.01V	0.03
Hot per-humid lowlands	0.01	0.15
Tepid perhumid mid highland	0.13	0
Very cold sub-humid sub-afro alpine to afro-alpine	0.03	0

614 Note: \*= total area of Ethiopia 1.14mln km<sup>2</sup> ; \*\*=total number of profiles 14,681

615

616 **Appendix B: Environmental covariates**617 **Table B1.** List, description, spatial and temporal extent, and source of covariates used in modelling  
618 the reference soil groups.

Categories	Covariates	Descriptions	Spatial resolution	Temporal resolution	Source
Climate	prep	Precipitation	4 km	1981 - 2016	ENACTS (Dinku et al.,2014)
	prep_sd	The standard deviation of precipitation	4 km	1981 - 2016	Derived from ENACTS (Dinku et al.,2014)
	tmax	Maximum Temperature	4 km	1983 - 2016	ENACTS (Dinku et al.,2014)
	tmin	Minimum Temperature	4 km	1983 - 2016	ENACTS (Dinku et al.,2014)
	trange	Temperature range	4 km	1983 - 2016	ENACTS (Dinku et al.,2014)
	tav_sd	Standard deviation of average temperature	4 km	1983 - 2016	Derived from ENACTS (Dinku et al.,2014)
	pet	Potential evapotranspiration	4 km	1981 - 2016	Derived from ENACTS (Dinku et al.,2014) using Modified Penman method
	lstd	Land surface temperature- Day (Aqua MODIS- MYD11A2 , time series monthly average)	1000 m	2002-2018	AfSIS <sup>a</sup>
	lstn	Land surface temperature-Night (Aqua MODIS- MYD11A2 , time series monthly average)	1000 m	2002-2018	AfSIS
	soil_moist	Soil Moisture (Derived from one-dimensional soil water balance)	4 km	1981 - 2016	Ethiopian Digital AgroClimate Advisory Platform (EDACaP)
	soil_temp	Soil temperature	30 km	1979 - 2019	ERA 5-Reanalysis ECMWF data <sup>b</sup>
Topography	DEM	Digital elevation model (Elevation)	90 m	-	SRTM- DEM (Vågen, 2010)

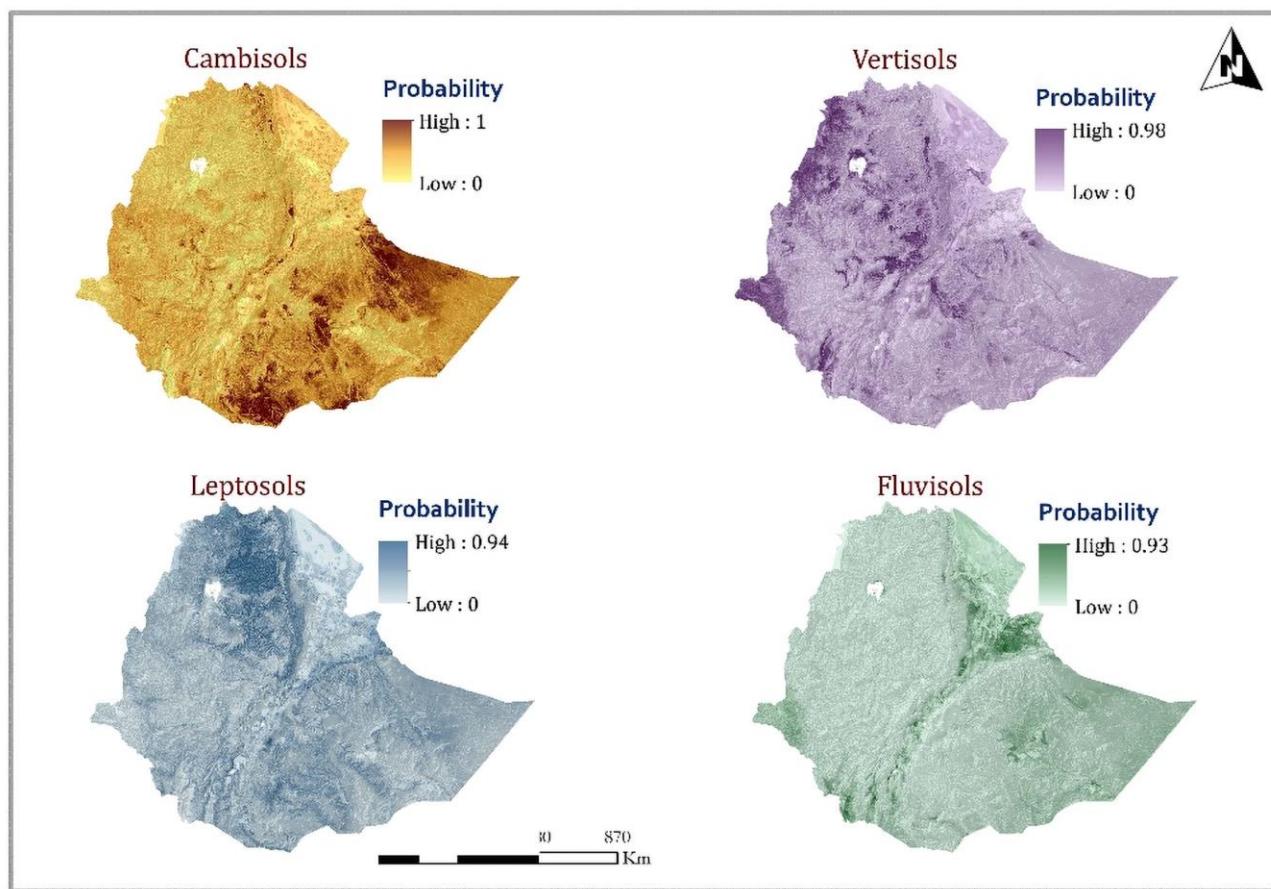
Categories	Covariates	Descriptions	Spatial resolution	Temporal resolution	Source
	twi	Topographic wetness Index	90 m	-	SAGA GIS-based SRTM-DEM derivative
	aspect	Topographic Aspect	90 m	-	SAGA GIS-based SRTM-DEM derivative
	curv	Topographic Curvature	90 m	-	SAGA GIS-based SRTM-DEM derivative
	conv	Topographic convergence index	90 m	-	SAGA GIS-based SRTM-DEM derivative
	ls	Slope Length and Steepness factor (ls_factor)	90 m	-	SAGA GIS-based SRTM-DEM derivative
	morph	Terrain Morphometry	90 m	-	SAGA GIS-based SRTM-DEM derivative
	mrvbf	Multiresolution index of valley bottom flatness	90 m	-	SAGA GIS-based SRTM-DEM derivative
	slope	Slope class (%)	90 m	-	SAGA GIS-based SRTM-DEM derivative
Vegetation	ndvi	Normalised Difference Vegetation Index (NDVI) (MODIS- MODIS MOD13Q1, time series monthly average)	250 m	2000-2021	AfSIS <sup>a</sup>
	evi	Enhanced Vegetation Index (EVI) (MODIS- MODIS MOD13Q1, time series monthly average)	250 m	2000-2021	AfSIS
	lulc	Land use/ landcover	30 m	2010	Water and Land Resource Centre-Addis Ababa University (WLRC-AAU, 2010)
parent material	lithology	Geology/parent material	1:2,000,000	1996	The Ethiopian Geological Survey (Tefera et al.,1996)
MODIS spectral refelectance	ref1	Red band (MODIS- MODIS MOD13Q1, time series monthly average)	250 m	2000 – 2018	AfSIS <sup>a</sup>
	ref2	Near-Infrared (MODIS- MODIS MOD13Q1, time series monthly average)	250 m	2000 – 2018	AfSIS

Categories	Covariates	Descriptions	Spatial resolution	Temporal resolution	Source
	ref7	Mid-Infrared (MODIS- MODIS MOD13Q1, time series monthly average)	250 m	2000 – 2018	AfSIS

619

620 **Appendix C: Probability of occurrence of reference soil groups**

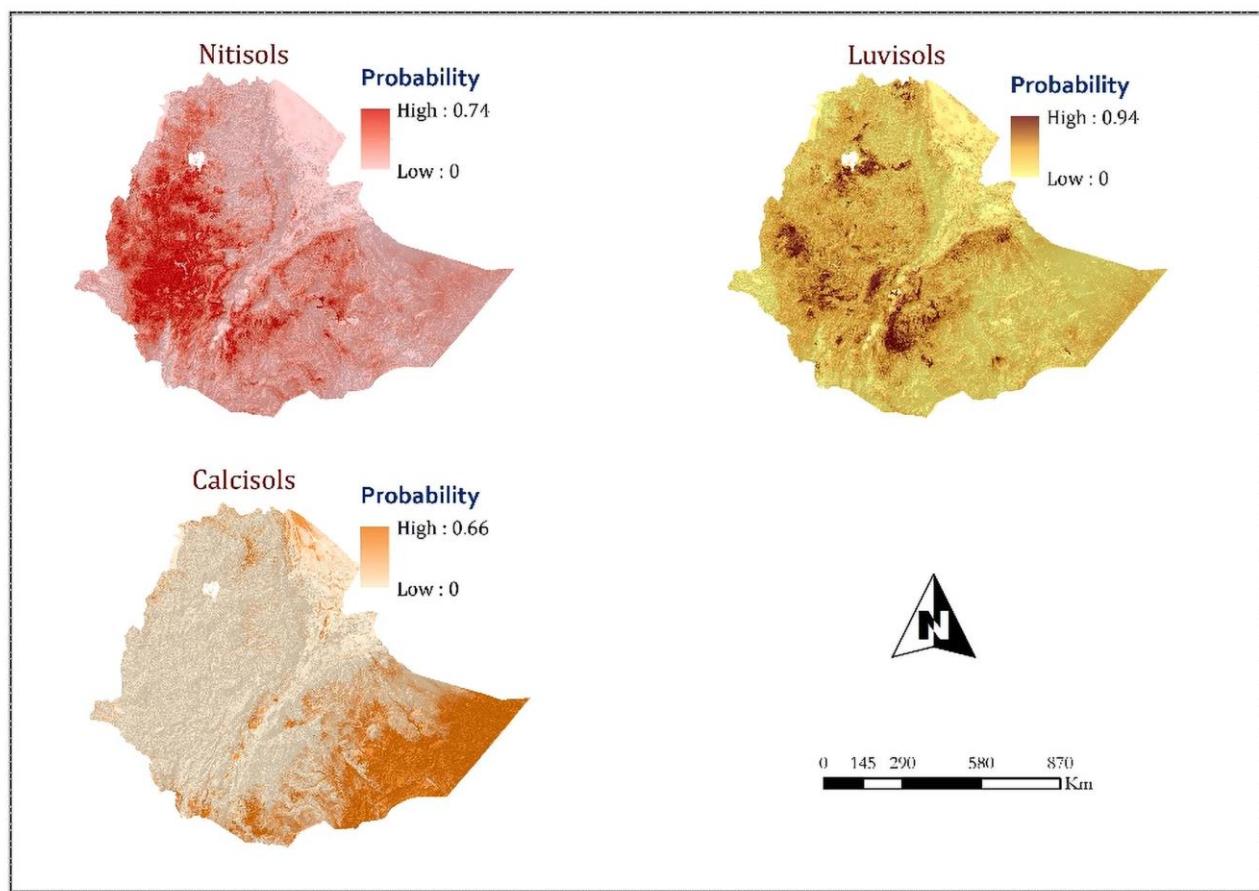
621



622

**Figure C1.** Occurrence probability maps of Cambisols, Leptosols, Vertisols, and Fluvisols.

623



**Figure C2.** Occurrence probability maps of Nitisols, Luvisols, and Calcisols.

627 **Data availability.** Full data will be available upon request based on the CoW guideline (CoW, 2020;  
628 <https://ethioagridata.com/>) and the MoA “Soil and Agronomy Data Management, Use and Sharing”  
629 directive No. 974/2023 Ethiopia (<https://nsis.moa.gov.et/>).

630 **Author contributions.** AA, TE, KG, WA, and LT conceived and designed the study, perform the  
631 analysis, and wrote the first draft, with substantial input and feedback from all authors. EM, TM, NH,  
632 AY, AM, TA, FW, AL, NT, AA, SG, YA, and BA, contributed to input data preparation, data  
633 encoding, and harmonization. Legacy data validation and review of subsequent versions of the paper  
634 were performed by MH, WH, AA, DT, GB, MG, SB, MA, AR, YGS, ST, DA, YW, DB, EZ, SC, and  
635 EE.

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