



The effect of Static Chamber's Base on N₂O Flux in Drip Irrigation

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Abstract

10 Static chambers are commonly used to provide *in-situ* quantification of N₂O fluxes. Despite their
benefits, when left in the field, the physicochemical conditions inside the chamber's base may differ from
the ambient, especially in drip-irrigated systems. This research aimed to study the effects of static
chambers' bases on water and N-forms distribution and the impact it has on N₂O measurements in drip
irrigation. N₂O emissions were measured in a drip-irrigated avocado orchard for two years, using bases
15 with a dripper at their center (In) and bases installed adjacent to the dripper (adjacent). During the
irrigation/fertigation season, the measured N₂O_{In} fluxes were greater than the N₂O_{Adjacent} fluxes ($0.82 \pm$
 0.15 vs. 0.36 ± 0.05 ng cm⁻² sec⁻¹). In contrast, during the winter, when the orchard is not irrigated or
fertilized, insignificant differences were observed between the measured N₂O_{Adjacent} and N₂O_{In} fluxes.
Three dimensional simulations of water flow and N-forms transport and transformations showed two
20 opposing phenomena (a) increased water contents, N concentrations, and downward flushing when the
dripper is placed inside the base, and (b) hampering of the lateral distribution of water and solutes into
the most bio-active part of the soil inside the base when the base is placed adjacent to the dripper. It also
showed that both "In" and "adjacent" practices underestimate the "true" cumulative flux from a dripper
with no base by ~25% and ~50%, respectively. A nomogram in a non-dimensional form corresponding
25 to all soil textures, emitter spacings and discharge rates, was developed to determine the optimal diameter
of an equivalent cylindrical base to be used along a single dripline.

1. Introduction

Static chambers are commonly used to provide *in-situ* quantification of N₂O fluxes from soil-plant
systems (Clough et al., 2020). Ideally, such chambers should be as large as feasibly possible in order to
30 capture spatial variation, where most chambers cover a surface area of 0.03 – 0.25 m². Commonly, static
chambers are built from two separate parts: (i) bases (also known as collars or anchors) that are pushed
into the ground, and (ii) chambers that are placed and sealed onto the bases during flux measurements.

It is acknowledged that static chambers suppress the gas concentration gradient at the soil-atmosphere
interface. The magnitude of the chamber-induced errors is known to increase with increased chamber
35 deployment time, decreased chamber height, greater soil air-filled porosity (Venterea, 2010). Different



methods have been developed to quantify/correct these errors (Venterea, 2010; Venterea et al., 2020). Proper use of static chambers calls for minimal disturbance of the soil surface and the prevention of lateral diffusion of N_2O beneath the bases' wall. To reduce perturbation of the soil structure, which can release pulses of gases, and to minimize the influence on the carbon and nitrogen turnover in the bases
40 (mainly the decaying process of cut roots) (Clough et al., 2020), bases are installed prior to the beginning of the experiment and left in the same location for the duration of the project. During installation, bases need to be inserted to at least the depth where N_2O concentrations are not being perturbed by feedback effects from the chamber (Rochette and Eriksen-Hamel, 2008; Healy et al., 1996; Hutchinson and Livingston, 2001). Despite their benefits, when bases are left in the field, atypical soil water contents,
45 aeration, temperature, and microbial processes may develop in them (Clough et al., 2020). Such perturbations must be minimized or avoided, especially in drip-fertigation systems where the localized concentrations of N and water content may increase the potential for N_2O emissions (Smart et al., 2011; Baram et al., 2018; Burton et al., 2008; Zebarth et al., 2008).

A review of the published literature in which static chambers were used at drip-irrigated fields shows
50 inconsistency regarding the chamber's location relative to the emitter (drinker). In many studies, chamber's bases (i.e., metal or PVC frame) are installed adjacent to the drip lines to depths of 5 to 13 cm, prior to the beginning of the experiment and left in the same location for the duration of the project (Garland et al., 2014; Baram et al., 2018; Alsina et al., 2013; Verhoeven and Six, 2014; Tian et al., 2017; Scheer et al., 2008; Fentabil et al., 2016). An additional strategy is to install the bases to similar depths,
55 a day to few hours before sampling (Sanchez-Martín et al., 2010; Sánchez-Martín et al., 2008). Both strategies are applied to reduce perturbation of the soil structure following the base's insertion, as discussed above. In such practices, the base serves as a hydraulic barrier that prevents lateral movement of water from the dripper into the soil volume confined by it, especially close to the soil surface. Accordingly, the water-filled pore-space (WFPS; the ratio between the volumetric water content and the
60 porosity, also known as "water saturation degree", S) and the N-species concentrations in the base will most likely differ from the ambient ones. Sanchez-Martín et al. (2010), acknowledged these biases (especially in the water content) and, therefore, removed the bases after sampling, allowing lateral redistribution between sampling days. Wolff et al. (2017) and Vallejo et al. (2014) tried to overcome these limitations by installing the bases and sampling the fluxes directly after the fertigation event.
65 However, as mentioned before, such a sampling method could be affected by perturbation of the soil structure and root damage (Clough et al., 2020). Placement of drippers above the bases was also suggested as a way to minimize the differences between the soil water content and N concentrations inside the bases and that of the surroundings (Heller et al., 2010; Garland et al., 2014). Another suggestion was to split the dripper discharge into two separate drippers, one in the bases and the other outside
70 (Fentabil et al., 2016). Though these latter methods seem to overcome the base's disturbance of water and solute distribution in the topsoil, their actual impact on water and N-species distribution in the soil inside the base and their effect on the N_2O fluxes were never tested.

Many models have been developed to simulate N_2O emissions from soils. In all models, N_2O emission results mainly from nitrification and denitrification reactions, which in turn are affected by the following



75 parameters in the soil: (a) mineral N concentration [namely, nitrate (NO_3^-) and ammonium (NH_4^+)], (b)
WFPS as a proxy for soil aeration and gas diffusion coefficient, (c) temperature, (d) pH, (e) redox
potential, and (f) carbon availability (Rabot et al., 2015; Hénault et al., 2019; Wu and Zhang, 2014). This
research aimed to study the effects of static chambers' bases on water and N-forms distribution inside it
and the impact it has on N_2O emission measurements in drip irrigation. We used both field measurements
80 and three-dimensional (3-D) simulations of flow and transport in order to test the effect of the base
diameter and its location relative to the dripper lateral on N_2O emissions.

2. Materials and methods

2.1. Study site and N_2O measurements

N_2O fluxes were measured over a two-year period in a drip-irrigated avocado orchard. The orchard is
85 located near Kibbutz-Yasur in Western Galilee, Israel. The soil at the site is a Vertisol (58% clay
dominated by montmorillonite) (Nemera et al., 2020). The climate at the site is Mediterranean,
characterized by a relatively long dry season (April – October) requiring irrigation and a distinct rainy
period during the winter (November – March). The trees are planted 3.5m apart on ridges (1.6m wide,
0.4m high), with 6m between rows. Each row of trees was irrigated with a set of two driplines (laterals),
90 located 0.9m apart along both sides of the trees, with 0.5m spaced 1.6 L h⁻¹ drippers (UNIRAM, Netafim).
From April through November, the orchard was fertigated every other day, using treated wastewater
enriched with an ammonium sulfate nitrate solution ($\text{NH}_4:\text{NO}_3 = 3:1$), maintaining 50 – 70 mg-N L⁻¹ in
the fertigation solution.

From June 2018 through June 2020, N_2O fluxes were measured at mid-morning using
95 accumulation static chambers that were installed at 12 random locations in the avocado orchard. At each
location, two chamber bases were inserted, one with a dripper at its center (In) and one adjacent to the
dripper (adjacent) (Fig. 1). The bases were made from opaque polyvinylchloride (PVC) rings, 10 cm high
and 19 cm internal diameter (ID, surface area of 283.5 cm²). The rings were inserted to a depth of 6-8
cm two weeks prior to the start of the sampling campaign and remained in the soil for the duration of the
100 experiment. The chambers were built from a 20 cm sewer PVC cup (volume of 3119 cm³), equipped with
a vent (3 mm Swagelok bulkhead union with a 12 cm long coiled copper tube, 1.5 mm i.d.), covered with
a bubble reflective foil and a rubber skirt to ensure sealing with the base. Fluxes were measured in real-
time by circulating the headspace in the static gas chamber via Teflon tubes into a Fourier-transform
infrared spectrometer (FTIR; Gasmeter DX4000, Gasmeter Technologies, Finland). During the enclosure
105 period, N_2O concentrations were recorded every five seconds, as an average average of 50 reads. N_2O
fluxes (q) [g cm⁻² sec⁻¹] were calculated based on the linear slope, representing the increase in N_2O
concentration throughout 4 to 8 min enclosure time (Eq. 1). The Pearson's correlation coefficient (r^2) was
calculated for the linearity of the slope, and readings were accepted when r^2 was >0.70.

$$[1] \quad q = \frac{dC_{gas}}{dt} * \frac{V_{chamber}}{A_{chamber}} * \frac{P}{RT} * M_w$$



where C_{gas} is the measured gas concentration [$\mu\text{L L}^{-1}$], t is the time [sec], $V_{chamber}$ [cm^3] and $A_{chamber}$ [cm^2]
110 are the chamber volume and surface area respectively, P is the ambient pressure [atm], R is the gas law
constant [$0.08206 \text{ L atm mol}^{-1} \text{ }^\circ\text{K}^{-1}$], T is the temperature [$^\circ\text{K}$], and M_w is the molecular weight of the gas
[g mol^{-1}]. Daily values were obtained following the method of Parkin and Kaspar (2006), using a Q_{10}
value of 3.72.

2.2. Modelling water flow, nutrients fate, and N_2O emissions

115 2.2.1. The physical domain and its parametrization

Employing a Cartesian coordinate system (x_1, x_2, x_3), where x_1 directed downwards, a subplot of the orchard
consisting of a 3-D, spatially heterogeneous, variably saturated flow domain which extends over $L_1=2\text{m}$,
 $L_2=15\text{m}$, and $L_3=10\text{m}$ along the x_1, x_2 , and x_3 axes, respectively, is considered here. The subplot includes
two adjacent tree rows, located 6m apart, with four trees, located 3.5m apart, along each row (Fig. SI-1)

120 Following Russo et al. (2020), the van Genuchten (1980) (VG) five-parameter model (*i.e.*,
saturated conductivity, K_s , shape parameters, α and n , the saturated, θ_s , and residual, θ_r , values of water
content, θ), was implemented here for the local description of the constitutive relationships for
unsaturated flow. Based on previous studies (e.g., Russo et al., 1997; Russo and Bouton, 1992), it is
assumed here that each of the VG parameters is a second-order stationary, statistically anisotropic,
125 random space function, characterized by a constant mean, and a two-point covariance. Parameters of the
latter, the variance and the correlation length-scales, were adopted from Russo and Bouton (1992). Grain-
size distribution data was obtained by the laser diffraction method (Eshel et al., 2004) from 0.3m-segments
of five soil cores extending to a depth of 1.2m. The data was used to estimate the local-scale VG parameters
by an optimization procedure. For more details see Russo et al. (2020). Mean values of the VG parameters
130 were estimated using the soil texture-based procedure suggested by Mishra et al. (1989). Details of the
generation of the 3-D, cross-correlated realizations of the spatially heterogeneous VG parameters are given
in Russo et al. (2006). Mean values and coefficients of variation (CV) of the resultant VG parameters are
given in Table 1 of Russo et al. (2020). The numerical grid used for the generation of the 3-D VG
parameters' field, was modified in order to account for the application of water by the drip irrigation system
135 and for the geometry of the ridges. For more details, see Russo et al. (2020).

In addition, deterministic molecular diffusion coefficient for chloride (Cl^-), nitrate (NO_3^-) and
ammonium (NH_4^+) in water, $D_0=5.4 \cdot 10^{-5} \text{ m}^2\text{d}^{-1}$, dimensionless Henry's constant for N_2O , $K_H=0.2$, and pore-
scale dispersion tensor (with longitudinal dispersivity, $\lambda_L=2 \times 10^{-3}\text{m}$ and transverse dispersivity, $\lambda_T=1 \times 10^{-4}\text{m}$
(Perkins and Johnston, 1963), were considered in the simulations. First-order rate constants for
140 nitrification and denitrification, K_1 and K_2 , respectively, and liquid-solid partitioning coefficient for
ammonium, K_{d1} , were taken into account as depth-dependent, implementing values within the range
suggested by Lotse et al. (1992). Estimates of the root uptake coefficients for ammonium and nitrate, K_{u1}
and K_{u2} , respectively, were calculated by extending the method of Nye and Tinker (1977); for more
details see Russo et al. (2013). Root distribution data, adopted from Salgado and Cautin (2008), were
145 employed in order to construct a time-invariant, normalized root depth-distribution function for the
avocado trees.



2.2.2. Quantification of the flow and the transport

150 Considering water and nitrogen extraction by plant roots, water flow and solute (ammonium, nitrate, and chloride) transport in the 3-D, unsaturated, spatially heterogeneous flow system were simulated employing numerical solutions of the 3-D Richards equation and the 3-D single-region, advection-dispersion equation (ADE), respectively. Following Russo et al. (2015), the flow model was modified to account for irrigation by drippers. The iterative procedure described in Russo et al. (2006) was employed in order to determine the size of the time-dependent ponding area that may develop around the drippers at the soil surface during an irrigation event. Furthermore, following Russo et al. (2020), the sink term representing water uptake by the plant roots, which appears on the right-hand side of the Richards equation, was modified to account for the effect of the oxygen availability on water uptake. The maximization iterative (MI) approach proposed by Neuman et al. (1975) was adopted here in order to calculate water uptake by the plant roots and, concurrently, actual transpiration rate, $\tau_a(t)$.

160 Following Russo et al. (2013), the ADE was modified to account for nitrogen transformations and uptake by plants' roots in the soil-water-plant-atmosphere system. In addition, the competition between chloride and nitrate and its effect on the extraction of nitrogen by the plant roots, and the inhibition of nitrification induced by chloride were taken into account. For more details, see Russo and Kurtzman (2019). The uptake of nitrate and ammonium by the plants' roots was also calculated by a MI approach described by Eq. 6 in Russo et al. (2013).

165 N_2O emissions were calculated based on Hénault et al. (2005) and Hénault et al. (2019), accounting for nitrification and denitrification driven emissions. N_2O flux during denitrification (N_2O_{denit} ; $mg-N m^{-2} d^{-1}$) was calculated as a combination of the potential denitrification rate (D_p ; $mg-N m^{-2} d^{-1}$) and response functions to several environmental factors (Hénault et al. 2005):

$$\text{Eq. 2} \quad N_2O_{denit} = D_p \cdot F_w \cdot F_N \cdot F_T \cdot r_{max}$$

170 Where F_w is the denitrification response factor to the soil WFPS, assuming the WFPS parameter serves as a proxy of the oxygen availability for microorganisms (Supporting information (SI) Eq-SI1). F_N is the denitrification response factor to soil nitrate content (dimensionless) (based on Michaelis–Menten saturation curve) (Eq-SI2), F_T is the denitrification response factor to soil temperature (Eq-SI3), r_{max} is the maximum ratio of N_2O to denitrified nitrate under anaerobic incubations (in this study $r_{max} = 0.3$). N_2O production during nitrification ($mg-N m^{-2} d^{-1}$) was defined by Eq. 3 (Hénault et al. 2005):

$$\text{Eq. 3} \quad \begin{aligned} N_2O_{nit} &= z N_A, & WFPS < 0.62 \\ N_2O_{nit} &= r_{max} z N_A, & WFPS \geq 0.62 \end{aligned}$$

175 Where z is the proportion of nitrified nitrogen emitted as N_2O (in this study $z = 0.006$), and N_A is the actual areal nitrification rate ($mg-N m^{-2} d^{-1}$).

Details of the flow and the transport equations and the numerical schemes employed to solve them are given elsewhere (Russo and Kurtzman, 2019; Russo et al., 2013).

2.2.3. Implementation

180 Meteorological data collected in the Yasur orchard were used to estimate the reference evapotranspiration, $ET_0(t)$, using Penman-Monteith method. Potential evapotranspiration rates, $\varepsilon_p(t) = \varepsilon_p(t) + \tau_p(t)$ (where ε : evaporation, τ - transpiration), were estimated from the $ET_0(t)$ data using the time-



dependent crop coefficients actually used in the Yasur site. Assuming that the wetted soil surface area of the ridge is completely covered by the trees' canopy, a negligibly low soil evaporation rate was adopted for the surface area of the ridges, *i.e.*, $\tau_p(t) = \varepsilon \tau_p(t)$. For the soil surface area between the ridges outside the rooted zone, a negligibly small transpiration rate was assumed, *i.e.*, $\varepsilon_p(t) = \varepsilon \tau_p(t)$. Actual rates of water loss by evaporation, $\varepsilon_a(t)$, were implemented by a MI approach described in Russo et al. (2006).

The chamber base was modeled as a cuboid whose axes coincide with the coordinates of the flow system. The center of a given chamber base is located at a given, user-controlled point, $p = p(x_2, x_3)$, in the x_2x_3 -horizontal plane; it extends vertically from the soil surface, $x_1 = 0$, to the depth of $x_1 = Z_{\text{bot}}$, and horizontally from $x_{c21} = p - \delta x_2$ to $x_{c22} = p + \delta x_2$ and from $x_{c31} = p - \delta x_3$ to $x_{c32} = p + \delta x_3$, where $Z_{\text{bot}} = 0.10\text{m}$ and $\delta x_2 = \delta x_3$ vary between 0.1m to 0.2m. Unit-head-gradient is specified at $x_1 = Z_{\text{bot}}$, and no-flow is specified at $x_1 = 0$ and at the vertical planes of the chamber located at x_{c21} and x_{c22} and at x_{c31} and x_{c32} .

Appropriate initial conditions for the present analyses were created by considering measured water content and solutes' concentrations profiles obtained prior to the irrigation season. For the flow, a second-type upper boundary condition was imposed on the top boundary ($x_1 = 0$) with flux that is determined by the drippers' discharge and by the time-dependent potential soil evaporation flux. A unit-head-gradient-boundary was specified at the bottom boundary ($x_1 = L_1$). For the transport, a first-type upper boundary condition was imposed on the top boundary with inlet concentrations corresponding to the irrigation water concentrations. A zero-gradient-boundary was specified at the bottom boundary. No-flow conditions are assumed for the vertical boundaries located at $x_2 = 0$, $x_2 = L_2$, $x_3 = 0$, and $x_3 = L_3$ (Fig. SI-1)

For a given location at the horizontal x_2x_3 -plane and a given horizontal extent of the chamber base, starting at the beginning of the irrigation season (May 1st), flow and transport simulations proceeded for an irrigation period of 180d. Actual concentrations of nitrate and ammonium in the irrigation water (including amounts added as fertilizers) and concentration of chloride in the irrigation water used in the field experiments were implemented in the simulations.

2.3. Statistics

Data were analyzed using JMP® Pro Statistical Software version 15.0 (SAS Institute Inc., USA). For each base size or location, we used a t-test to analyze the effects of the base on the different variables. The data met the assumption of homogeneity of variances. Presented data are means \pm standard error (SE), with p values (p) representing the level of statistical significance.

3. Results

3.1. Measured N₂O fluxes in the field

Results from two years of measuring showed that the N₂O fluxes were higher during the fertigation season (April through October) than during the winter period (November through March). During the winter, no significant differences were observed between the measured N₂O_{In} and N₂O_{Adjacent} fluxes. In contrast, during the irrigation/fertigation season, the N₂O fluxes from the chambers with a dripper at their bases (N₂O_{In}) were on average 2.3 ± 0.56 times greater than the fluxes from the chambers adjacent to the dripper (N₂O_{Adjacent}) (0.82 ± 0.15 vs. 0.36 ± 0.05) (Fig. 2). In 4.3% of the samples, N₂O_{Adjacent} fluxes were



220 higher than the N_2O_{in} fluxes. This phenomenon only occurred when the measured fluxes were very low
($<0.01 \text{ ng cm}^{-2} \text{ sec}^{-1}$) in both locations and only following irrigation events and not fertigation events
(Fig. 2). In 17% of the measurements, the N_2O_{in} and $N_2O_{Adjacent}$ fluxes differed by 20% or less (Fig. 2).

3.2. Simulation results

Simulation results show that during irrigation, the WFPS down the soil profile under the chamber's base
225 (20 cm ID), with a dripper at its center (WFPS_{In}), were higher than the WFPS under a normal
representative dripper with no base (WFPS_{No}) (Fig. 3A and Fig. 4D). A day after the irrigation, the
WFPS_{In} decreased faster, leaving the soil profile under the WFPS_{No} treatments wetter. Throughout the
simulation period (i.e., from day five onwards), the WFPS at depths of 10, 20 and 30 cm hardly differ,
and the WFPS_{In} ranged from 0.77 to 0.63, while the WFPS_{No} ranged from 0.74 to 0.65.

230 Ammonium-N (NH_4^+-N) concentrations in and under the chamber's base ($NH_4^+-N_{In}$) increased
at all depths from $\sim 5 \text{ mgL}^{-1}$ at the start of the simulation to $\sim 25 \text{ mgL}^{-1}$ following 20 days. After 20 days,
the concentrations remained high with minor changes. In contrast, the concentrations under a normal
representative dripper with no base ($NH_4^+-N_{No}$) slightly increased during the 60 days of simulation (from
5 to 8 mgL^{-1}) (Fig. 3B). Unlike with the NH_4^+-N concentration, nitrate-N ($NO_3^- -N$) concentrations
235 showed a clear oscillating trend over time that corresponded to the N concentration in the fertigation
solution (Fig. 3C). The amplitude of change was higher when the dripper was placed inside the base.

Simulations with bases of variable sizes (i.e., no-base, 20, 30, and 40 cm ID) showed a clear
size impact on the average NH_4^+-N and $NO_3^- -N$ concentrations in the top 10 cm of the soil (Fig. 5). For
 $NO_3^- -N$, the smaller the base ID is, the higher the deviation from ambient (no-base) concentrations.
240 During fertigation events, $NO_3^- -N$ concentrations inside the 20, 30, and 40 cm bases increased by up to
212%, 159%, and 137%, respectively, relative to a dripper with no base around it. In contrast, between
fertigation events, the concentrations decreased down to 10%, 32%, and 61% of the concentrations under
a dripper with no base. Unlike the oscillating trend of $NO_3^- -N$ concentrations, NH_4^+-N concentrations
during the simulated 60 days of fertigation gradually accumulated in the soil until they stabilized at
245 around 20 mg L^{-1} . The sharpest buildup was observed under the 20 and 30 cm ID bases where the NH_4^+-N
concentrations were up to 300 times greater than the concentrations under a dripper with no-base (Figs.
5).

Simulated N_2O emissions showed a clear oscillating trend over time, which was affected by the
irrigation and fertigation regime (Fig. 6). During fertigation events, N_2O fluxes from the 20 cm ID base
250 were on average higher by $14 \pm 6\%$ ($p = 0.0345$) than the fluxes from a dripper with no base, with higher
fluxes from deeper parts ($< 10 \text{ cm}$) of the soil (Fig. 4C). However, the fluxes from the 30 and 40 cm ID
bases were lower by $10 \pm 5\%$ and $26 \pm 4\%$, respectively ($p < 0.001$, both). A day after a fertigation event,
the fluxes from the 20, 30, and 40 cm ID chambers were $-4 \pm 10\%$, $-14 \pm 5\%$, and $-26 \pm 3\%$ lower than the
fluxes from a dripper with no base ($p < 0.008$, all). More than one day after fertigation, the fluxes from
255 the 20, 30, and 40 cm ID chambers were significantly lower than from a dripper with no base, with the
greatest reduction in the 20 cm ID base ($-69 \pm 3\%$, $-42 \pm 4\%$, and $-24 \pm 3\%$). Irrigation events a day or two
days following fertigations drastically reduced the N_2O fluxes, leading to fluxes that equaled $33 \pm 6\%$,
 $67 \pm 8\%$, and $85 \pm 5\%$ of the fluxes measured from a dripper with no base.



Under a dripper with no-base and a 40 cm ID base, simulated N₂O emission was significantly
260 ($p < 0.05$) affected by the simulated WFPS, NH₄⁺-N, and NO₃⁻-N concentrations at depths of 10, 20, and
30 cm below the surface, with R² correlations of 0.10, 0.20 and 0.99, respectively (Table SI-2 and Table
SI-3). In contrast, under bases with a 20 and 30 cm ID, simulated N₂O emissions were significantly
affected only by the simulated NO₃⁻-N concentrations (Table SI-1, and Table SI-2, Table SI-3).

4. Discussion

265 Our field measurements show that N₂O_{Adjacent} fluxes may be higher than the N₂O_{In} fluxes following
irrigation events or several days after fertigation events (Fig. 2). These values suggest either higher WFPS
or higher concentration of N in the base adjacent to the dripper. Simulation results show that the
differences in the WFPS between a base with a dripper at its center and between an undisturbed dipper
are in the range of ±4% following irrigations every other day (Fig 3A). It also shows that these differences
270 can be greater (±8%), especially at a depth of 30 cm, following sporadic irrigation events or irrigation
events at the start of the growing season (Fig 3A and Fig. 4D). Analysis of the simulations results shows
that the WFPS had a significant impact on the N₂O fluxes only when no base was used (R²<0.004,
 $p > 0.598$ vs. R²<0.094, $p < 0.009$) (Table SI-3). Published literature shows that the WFPS-N₂O relation is
not always clear. N₂O emissions were thought to have their optimum in the range of 70–80% water-filled
275 pore space (WFPS) depending on soil type (Davidson et al., 2000). At higher soil water content, the
major end product of denitrification is N₂. Nevertheless, a comprehensive study of 51 soils across Europe
showed that most soils had their optimum N₂O emissions under WFPS wetter than 80% (Zechmeister-
Boltenstern et al., 2007). Similarly, Keller and Reinert (1994) found that N₂O emissions increased
exponentially with WFPS, even at WFPS greater than 80%. Bateman and Baggs (2005) showed that the
280 relative contribution of nitrification and denitrification to N₂O emission depended on the WFPS. Whereas
at 70% and 20% WFPS, most N₂O originated from denitrification, at 35–60% WFPS, nitrification was
the main process producing N₂O. Overall, the simulated WFPS indicated that the micropores in the soil
remained water-filled, which permits microbial activity without water stress, while the soil macropores
are air-filled, which permits relatively good aeration of the bulk of the soil, although anaerobic microsites
285 may exist. At such WFPS, both oxidative (nitrification) and reductive (denitrification) processes are
active in the soil.

It is well established that higher N concentrations lead to higher N₂O emissions (Wolff et al.,
2014, 2017; Baram et al., 2018; Schellenberg et al., 2012; Alsina et al., 2013; Butterbach-Bahl et al.,
2013). A clear indication of this phenomenon can be seen in Figures 3 and 4, the model's simulation of
290 water flow and NO₃⁻ and NH₄⁺ transport in the subsurface. When the dripper is placed in the center of
the base, NO₃⁻_{In} concentrations in the top 30 cm may be 50 – 64 % higher during fertigation events and
up to 67% lower following irrigation or in the days following fertigation events. Inspection of the trend
over time shows that, on average, NO₃⁻_{In} concentrations are decreased by -19%±5% ($p = 0.017$) relative
to a dripper with no base. This phenomenon results from the higher WFPS and the geometry of the base
295 that limits lateral flow. As such, the water flow in the base is essentially 1D in the vertical direction,
which expedites downward water flow and N transport into the subsurface. NH₄⁺, unlike NO₃⁻, is



positively charged, hence readily adsorbs to the clays in the soil. Accordingly, NH_4^+ concentrations in the top 30 cm increased by 280% relative to a dripper with no base and remained higher throughout the season ($p < 0.001$) (Fig. 3).

300 Inspection of the correlations between the simulated N_2O fluxes and the NH_4^+ and NO_3^- concentrations show that N_2O emissions were mainly influenced by the NH_4^+ and NO_3^- concentrations (Table SI-2 and Table SI-3). These, in turn, were affected by the base's ID, with higher ID leading to less biases. It is known that N_2O emission fluxes vary from one fertilizer event to another, even at the same site with the same fertilizer type under similar environmental conditions (Cowan et al., 2020). Here we
305 show that an additional factor that must be accounted for is the location of the chamber's base relative to the water source and the perturbation that the base has on water and N-species distribution. Simulation results show that placing the dripper inside the base may increase the N_2O flux during a fertigation event by up to 52% relative to a dripper without a base. In tandem, N_2O fluxes following irrigation events or in the days following fertigation may be up to 91% lower when the dripper is placed inside the base (Fig.
310 7A). A similar effect is observed when the chamber's base is positioned adjacent to the dripper (i.e., up to 23% increase during fertigation and up to 97% decrease following irrigation events or in the days following fertigation). One should note that the modeled N_2O fluxes resulted mainly from denitrification, as suggested by Eq. 3 and its relation to the WFPS (Fig. 3A, Table SI-2, and Table SI-3). It is evident that under a different WFPS- $\text{N}_2\text{O}_{\text{mit}}$ pattern, the base impact on the N_2O may increase in tandem with the
315 higher NH_4^+ concentrations. All of these results provide a good indication of the two opposing phenomena (a) increased WFPS, N concentrations and downward flushing when the dripper is placed inside the base, and (b) hampering of the lateral distribution of water and solutes into the most bio-active part of the soil inside the base when the base is placed adjacent to the dripper.

Comparison of cumulative N_2O emission measured in 2018, 2019, and 2020 showed the $\text{N}_2\text{O}_{\text{In}}$
320 flux to be 50% – 60% higher than the $\text{N}_2\text{O}_{\text{Adjacent}}$ flux (98 – 120 mg m^{-2} vs. 60 – 74 mg m^{-2}). Simulation results showed similar differences between the two ($\text{N}_2\text{O}_{\text{In}}$ ~68% higher, 9.53 vs. 6.48 $\text{mg N}_2\text{O-N m}^{-2}$ over 60 days) (Fig. 7B). It also shows that both methods underestimate the "true" cumulative flux from a dripper with no base by ~25% and ~50%, respectively. These values suggest that in addition to measurement errors due to suppression of the gas concentration gradient at the soil-atmosphere in static
325 chambers (e.g., Venterea et al., 2020; Venterea, 2010), the impact of the chamber's base on the water and N distribution provides an additional level of complexity, leading to an erroneous estimate of the true N_2O flux. Accordingly, the emission factors measured in such setups (e.g., Alsina et al., 2013; Baram et al., 2018; Fentabil et al., 2016; Garland et al., 2014; Scheer et al., 2008; Tian et al., 2017; Verhoeven and Six, 2014) are likely lower than the actual emission factors.

330 The degree to which the location of the chamber's base relative to the dripper affects the N_2O flux will depend on the soil properties and on the chamber's ID. Overall, an increase in the chamber's ID will decrease the abovementioned biases by reducing the lateral flow constraints posed by the chamber's base. An indication of this process can be seen in Figure 5 for the clayey soil used in this study. Such clayey soils have a large capillary length (i.e., tens of cm long), which supports high lateral capillary flow.
335 Accordingly, the use of a chamber with a larger ID (i.e., the simulated 40 cm or even larger) is required



to reduce the negative effects of the base on the water distribution near the surface and to provide a more reliable representation of the ambient fluxes around drippers.

A recommendation on the diameter of the chamber's base follows a simplified analysis of steady water flow from either a dripper (emitter) with no base (an undisturbed dripline) or a dripper surrounded by the cylindrical base (i.e., "In"). It is suggested to determine the diameter of the cylindrical base (d_{cyl} , cm), such that the water potential (or water content) at a chosen depth below the dripper (d_{ref} , cm) will be equal to that generated at the same depth below an undisturbed dripline with equally-spaced (d_{emit} , cm) drippers, depending on the soil capillary length (α^{-1} , cm). The matric flux potential in this location (ϕ_{ref} , $\text{cm}^2 \text{h}^{-1}$) is evaluated with analytical solutions to the linearized, steady water flow equation for a surface point source confined in a strip (its width equals d_{emit}), imitating an undisturbed single dripline, and for a surface point source in the center of a laterally-confined cylindrical domain, mimicking the dripper surrounded by the inserted-into-the-soil base wall (Eqs. [19] and [31] in Communar and Friedman (2011), respectively). The computations were performed using DIDAS (Friedman et al., (2016), <https://app.agri.gov.il/didas/>), firstly computing the ϕ_{ref} generated at depth d_{ref} under the undisturbed dripper and then determining d_{cyl} that will generate the same ϕ_{ref} at the same depth in an iterative, trial-and-error procedure.

The resulting reference water potential (Fig. 8A) and cylindrical base nomogram (Fig. 8B) are presented in a non-dimensional form corresponding to all soil textures and inter-emitter spacings and discharge rates (q , $\text{cm}^3 \text{h}^{-1}$), after normalizing all relevant lengths (d_{cyl} , d_{emit} , d_{ref} , cm), by the soil capillary length, (α^{-1} , cm), which describes the exponential decrease of the soil's hydraulic conductivity upon drying according to the Gardner's (1958) function, $K = K_s \exp(\alpha h)$ (K – hydraulic conductivity (cm h^{-1}), K_s – hydraulic conductivity at water-saturation (cm h^{-1}), h – pressure head (cm)), and reflects the dominance of capillarity over gravity forces in driving the water in partially-saturated soils. Coarse-textured, sandy soils are characterized with large α (cm^{-1}) values, and fine-textured, clayey soils with small ones.

The dimensionless matric flux potential ($\Phi_{ref} = 8\pi\phi_{ref}/\alpha q$, $\phi = K/\alpha$, or water content or pressure head (h) for a given soil (α) and dripper discharge rate (q)) decreases sharply with increasing distance between emitters (d_{emit}) or with coarsening (increasing α) of the soil texture, what counts is only their product αd_{emit} , and with increasing reference depth (αd_{ref}) below the emitter (Fig. 8A). At a dimensionless inter-emitter distance (αd_{emit}) of about 2, the $\Phi_{ref}(\alpha d_{emit})$ lines flatten as the effect of the neighboring emitters weakens and the potentials converge to those generated by a single emitter (not in a dripline, dash-dotted lines, Eq. [10] in Communar and Friedman (2010b)).

The upper limit of the dimensionless inter-emitter distance (αd_{emit}) in the depicted cylindrical base nomogram (Fig. 8B) is 2, as larger spacings are not recommended to assure overlap between the wetted bulbs (Communar and Friedman, 2010b), and a dimensionless reference depth of 1 (αd_{emit}) as the processes leading to N_2O emission are occurring at shallow depths. The diameter of the equivalent cylindrical base decreases sharply with increasing distance between emitters or for more sandy soils (αd_{emit}), the plotted ratio, d_{cyl}/d_{emit} , is approximately proportional to $(\alpha d_{emit})^{-1/3}$. This is so because the relative (dimensionless) effect of the parallel strip walls (i.e., of the neighboring emitters) increases with



375 d_{emit} . However, the dependence of d_{cyl}/d_{emit} on the reference depth (αd_{ref}) is mild, a slight increase with
depth of reference locations for small inter-emitter spacings (or for clayey soils), and virtually
independence for large inter-emitter spacings (or for sandy soils). This is good news, as it means that an
equivalent cylindrical base of chosen diameter can provide similar water contents at a range of depths
below the dripper. The diameter of the equivalent cylindrical base is larger than the inter-emitter spacing
380 ($d_{cyl}/d_{emit} > 1$) for smaller inter-emitter spacing, or for fine-texture soils, and slightly smaller than the
inter-emitter spacing ($d_{cyl}/d_{emit} < 1$) for larger inter-emitter spacing, or for coarse-texture soils (Fig. 8B).
As written above, intuitively, it is expected that since an infinitely-deep cylinder confines lateral water
flow in all directions, while the symmetry vertical planes between drippers along the dripline confines it
only in the direction of the dripline, d_{cyl} should be larger than d_{emit} . These results agree with the simulation
385 results discussed above, demonstrating mostly larger differences compared to undisturbed drippers for
bases of smaller diameters (Figs.5 and 6). As a sensible inter-emitter spacing is about one capillary length,
i.e., $\alpha d_{emit} = 1$, the recommended d_{cyl}/d_{emit} is about one (Fig. 8B), a base diameter equal to the inter-emitter
spacing. Notice that the d_{cyl}/d_{emit} (αd_{emit} , αd_{ref}) nomogram is independent of the dripper discharge rate
(q), since according to the linearized water flow equation used for the analysis (Eq. [5] in Communar and
390 Friedman (2010), the matric flux potential generated by the drippers (point sources) is simply
proportional to q , whatever is the geometry of the flow field.

The use of the nomogram is very simple. Suppose we want to determine the diameter of a
cylindrical base (d_{cyl}) that will optimally reproduce the wetting patterns under 50 cm-spaced drippers
(d_{emit}) along a single dripline in a clayey soil with a capillary length (α^{-1}) of 100 cm, by requiring that the
395 water potential (content) at a depth (d_{ref}) of 25 cm below the dripper will be the same. The dimensional
emitter spacing (αd_{emit}) is thus 0.5, and the vertical arrow stops at the dimensionless reference depth
(αd_{ref}) of 0.25 (black solid line in Fig. 8B), from which the horizontal arrow stretches to approximately
 $d_{cyl}/d_{emit} = 1.21$, i.e., the cylindrical base diameter should be larger by 21% compared to the inter-emitter
spacing, about 60 cm. The soil capillary lengths (α^{-1}) of most agricultural soils vary between
400 approximately 10 cm for sandy soils to 100 cm for structureless, clayey soils with common values of 20
to 40 cm for loams and fine sands (Friedman et al., 2016). If the value of the hydraulic conductivity at
saturation is known, the soil capillary length can be evaluated with the universal relationship $\alpha =$
 $0.04035K_s^{1/2}$ (in which α is measured in cm^{-1} and K_s in cm h^{-1}) (Fig. 12 in Communar and Friedman
(2010), also used in DIDAS).

405 The analysis used for constructing the $d_{cyl}/d_{emit}(\alpha d_{emit}$, $\alpha d_{ref})$ nomogram is based on addressing
only water flow and applying multiple simplifying assumptions of steady flow, assuming an infinitely
deep confining cylinder (as opposed to the just a few centimeters insertion of the chamber base, although
the effect of 10cm insetion seems to affect the wetting patterns at large depths, Fig. 4D) and disregarding
water uptake by plant roots and evaporation from the soil surface. Yet, it is believed that it provides
410 plausible guidelines for choosing the optimal inserted-into-the-soil base diameter. In parallel,
minimizing the depth to which bases are inserted to 1-2 cm will lower its impact on the water and N-
species distribution inside it. This is especially true when short closure times are needed, such as when
portable gas analyzers are used in the field. An alternative option is to develop a static chamber that does



not need a base. All of the above are crucial steps that are required to obtain accurate observations on
415 which small- and large-scale climate prediction models are built.

5. Conclusions

Based on two years of field measurements and numerical simulations of water flow, N-species transport
and reactions (i.e., nitrate and ammonium), and N₂O emissions, we conclude that the bases of static
chambers provide an underestimation of N₂O emissions when used in drip-irrigation. This is an outcome
420 of: (a) increased water contents, N concentrations, and downward flushing when the dripper is placed
inside the base, and (b) hampering of the lateral distribution of water and solutes into the most bio-active
part of the soil inside the base when the base is placed adjacent to the dripper. A nomogram is proposed
to determine the optimal diameter of a cylindrical base to be used along a single dripline. Further study
is suggested to determine the optimal insertion depth of bases based on the enclosure period. An alternative
425 option is to develop a static chamber that does not need a base.

6. Author contribution

S. Baram and A. Bar-Tal supervised the project and acquired financial support. Baram designed the
experiments, analyzed the data, and wrote the manuscript with contributions from all co-authors. D.
Russo performed the detailed numerical simulations, and S. Friedman devised the equivalent cylinder
430 nomogram. A. Gal collected the data in the field.

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9. Figures

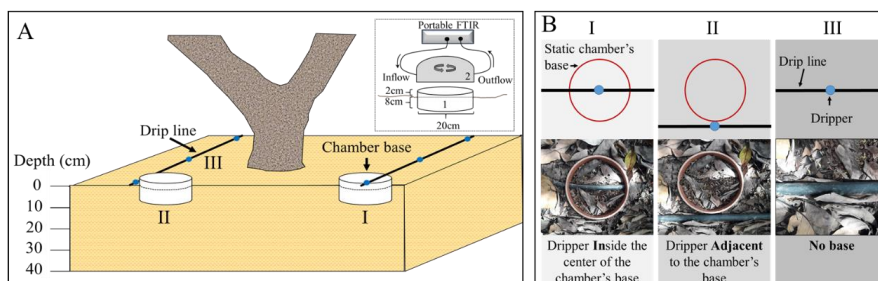


Figure 1. Schematic side view (A) and top view (B) representation of the static chamber setup used to measure N_2O fluxes in an Avocado orchard and their locations (I, II, III) relative to the drip line and the drippers.

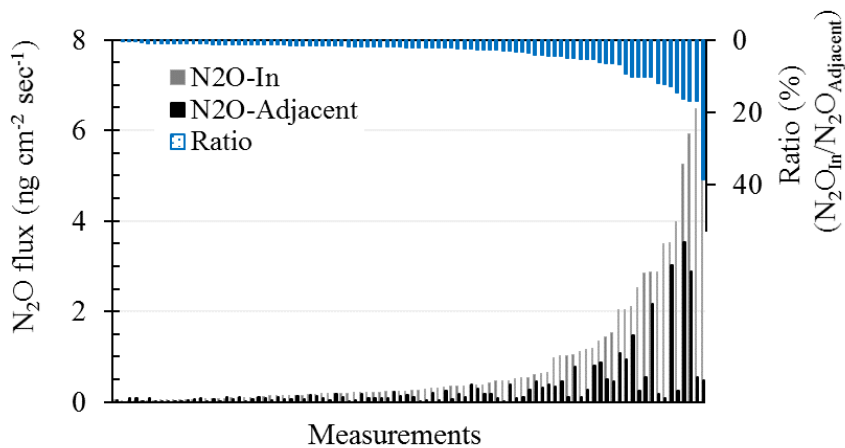


Figure 2. N_2O fluxes measured in static chambers with a drifter in their base (N_2O_{In}) and in static chambers with a drifter adjacent to their base ($N_2O_{Adjacent}$), and the ratio between N_2O_{In} and $N_2O_{Adjacent}$ at each one of the 12 sampling locations at all sampling days.

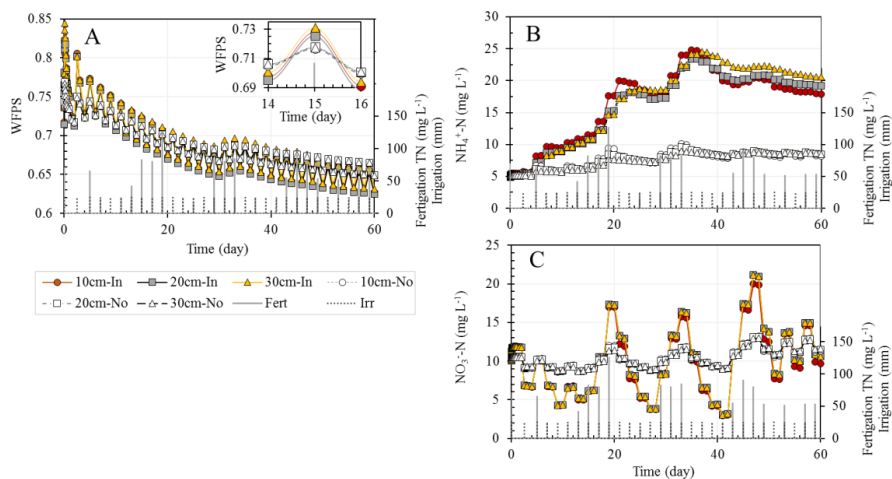
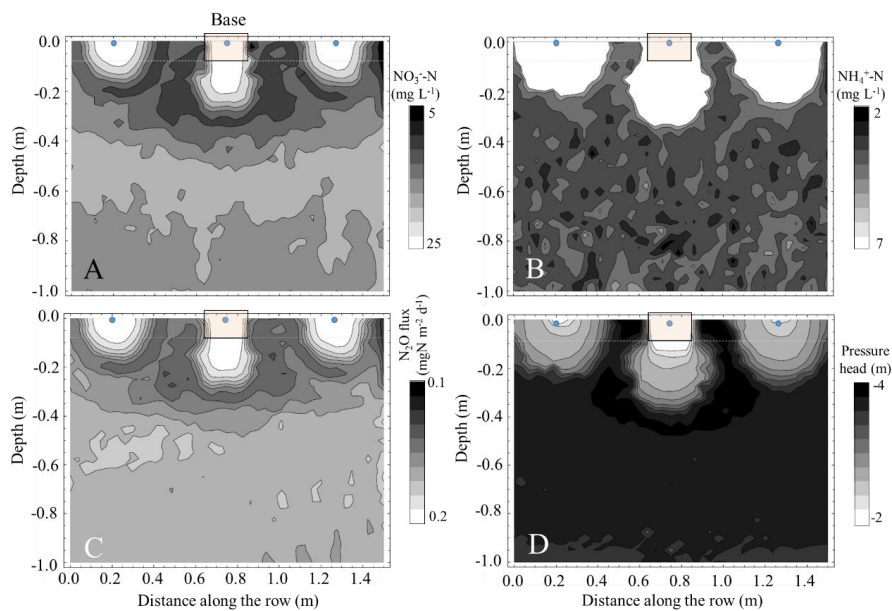


Figure 3. Simulation results of the change over time in (A) water filled pore space (WFPS), (B) NO₃⁻-N and (C) NH₄⁺-N concentrations at three depths (10, 20 and 30 cm) under a base of a static chambers, with a dripper at its center (filled shapes – In), and under a dripper without a base (light empty shapes – No) following irrigation every other day, and fertigation events. The top right corner of (A) shows a zoom in on the WFPS cycle during three consecutive days: the day after fertigation (day 14), the day of fertigation (day 15) and the day following the fertigation event (day 16). Total nitrogen (TN) concentration in the fertigation solution is the sum of NO₃⁻-N and NH₄⁺-N concentrations.



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Figure 4. Contours of the simulated (A) nitrate-N (NO_3^- -N) and (B) ammonium-N (NH_4^+ -N) concentrations, (C) vertical N_2O flux (positive upwards) and (D) pressure head distributions in the vertical x_1x_3 -plane of the flow domain in the vicinity of one of the dripper line laterals along one of the ridges whose center is located at $x_2=5.1\text{m}$ (Fig. SI-1). The figure shows three adjacent drippers, where the center dripper is confined by a static chamber's base. The presented data is a snapshot following 30 days of fertigation.

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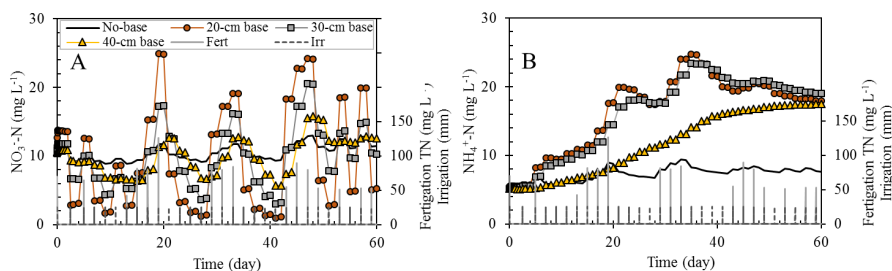


Figure 5. Simulation results of the change over time in $\text{NO}_3\text{-N}$ and $\text{NH}_4^+\text{-N}$ concentrations in the top soil (0–10cm) under bases of variable sizes (i.e., no-base, 20, 30, and 40 cm internal diameter) with a dripper at their centers. Total nitrogen (TN) concentration in the fertigation solution is the sum of $\text{NO}_3\text{-N}$ and $\text{NH}_4^+\text{-N}$ concentrations.

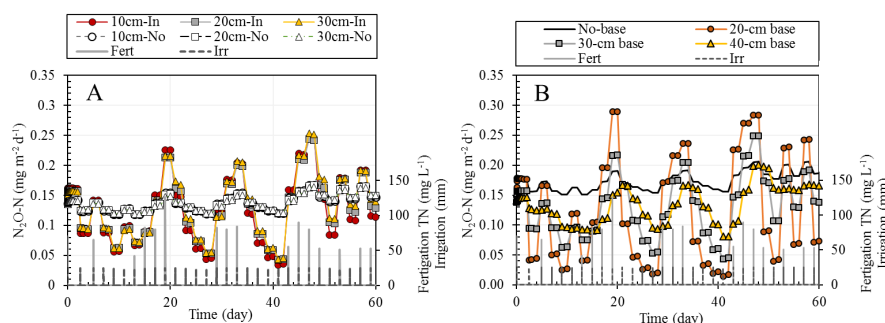


Figure 6. Simulation results of the change over time in $\text{N}_2\text{O-N}$ fluxes (A) at 10, 20 and 30cm below a base with a dripper at its center (In) and under a dripper without a base (No), and (B) in the topsoil (0–10cm) under bases of variable sizes (i.e., no-base, 20, 30, and 40 cm internal diameter) with a dripper at their centers. Total nitrogen (TN) concentration in the fertigation solution is the sum of $\text{NO}_3\text{-N}$ and $\text{NH}_4^+\text{-N}$ concentrations.

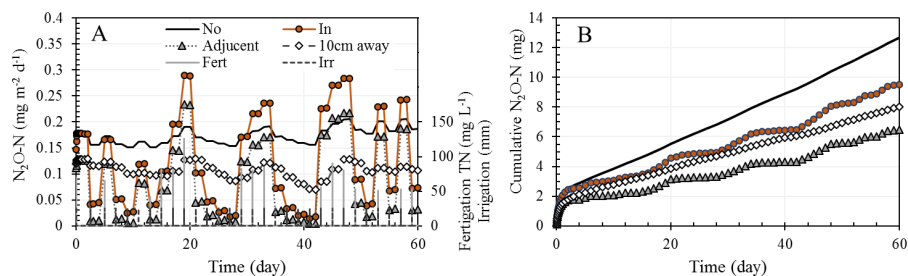


Figure 7. Simulation results of (A) daily and (B) cumulative $\text{N}_2\text{O-N}$ emissions from the soil surface under (No) and 10 cm away from a dripper without a chamber base, from a chamber base with a dripper at its center (In) and a base adjacent to the dripper (adjacent), during consecutive fertigation events over a period of 60 days.

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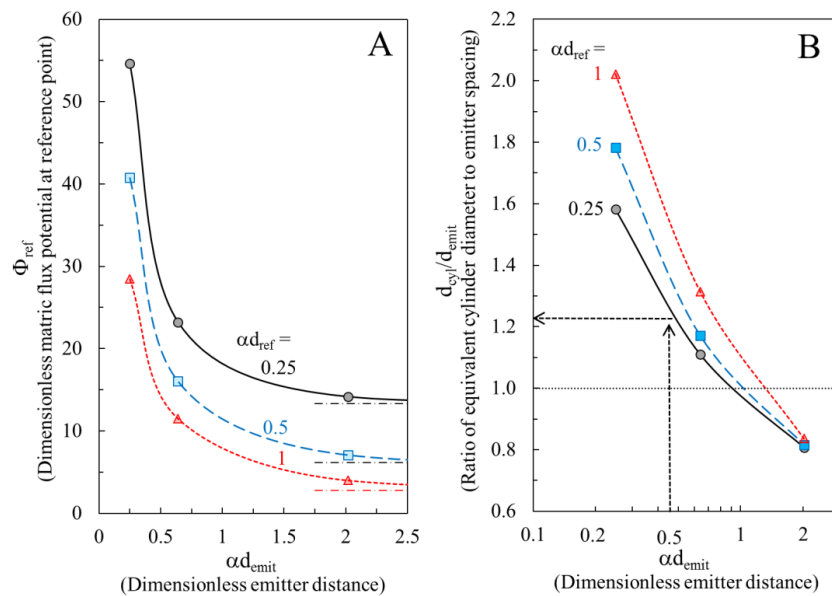


Figure 8. (A) Non-dimensional matric flux potential (Φ_{ref}), and (B) the ratio between the equivalent cylindrical base diameter (d_{cyl}) and the emitter spacing (d_{emit}), as a function of the non-dimensional distance between emitters (αd_{emit}), for different non-dimensional depths (αd_{ref}) below the emitter. The dash-dotted lines on (A) represent Φ_{ref} at the given depth below a single emitter. $\Phi_{ref} = 8\pi\phi_{ref}/\alpha q$, α is the soil's capillary lengths, q is the emitter discharge rate.

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