



A Combined Gas- and Particle-phase Analysis of Highly Oxygenated Organic Molecules (HOM) from α -pinene Ozonolysis

Jian Zhao¹, Ella Häkkinen¹, Frans Graeffe¹, Jordan E. Krechmer^{2,a}, Manjula R. Canagaratna², Douglas R. Worsnop², Juha Kangasluoma¹, and Mikael Ehn¹

¹Institute for Atmospheric and Earth System Research/Physics, Faculty of Science, University of Helsinki, Helsinki, 00014, Finland.

²Aerodyne Research Inc., Billerica, Massachusetts, 01821, United States.

^aNow at: Bruker Daltonics Inc., Billerica, Massachusetts, 01821, United States

Correspondence to: Jian Zhao (jian.zhao@helsinki.fi) and Mikael Ehn (mikael.ehn@helsinki.fi)

Abstract. Highly oxygenated organic molecules (HOM) are important for the formation of secondary organic aerosol (SOA), which poses serious health risks and exerts great influence on Earth's climate. However, the speciation of particle-phase HOM and its relationship with gas-phase HOM formation has been limited by the lack of suitable analytical techniques. Here, combining a novel particle evaporation inlet VIA (Volatilization Inlet for Aerosols) with a nitrate chemical ionization mass spectrometer (NO₃-CIMS), gas- and particle-phase HOM products of α -pinene ozonolysis were studied under different conditions. Within the 50-min residence time of our Teflon chamber, we observed enhancement of C₁₆-C₁₉ HOM dimers in particles compared to the HOM that were condensing. In particular, gas-phase dimer formation was considerably suppressed in experiments with the addition of CO or NO, but dimers still made up a considerable fraction of the observed SOA. In addition to the generally shorter carbon skeletons of the particle phase dimers (i.e. C₁₆-C₁₉) compared to the gas phase (C₁₉-C₂₀), average O/C ratios of the HOM (especially in the dimer range) also decreased slightly in the particle phase. C₁₇H₂₆O₂ compounds, which have often been reported by previous offline measurements, dominate the particle-phase HOM mass spectra in α -pinene ozonolysis experiments. Our results indicate that these C₁₇ compounds might be related to particle-phase processes within one hour after HOM condensation. However, the new VIA-NO₃-CIMS system used in this work will require more detailed characterization to better understand how the thermal desorption and wall effects may modify the measured particle-phase HOM distributions. Nevertheless, for example organic nitrate measured by this novel VIA-NO₃-CIMS system was consistent with the measurements of an Aerodyne Aerosol Mass Spectrometer (AMS), showing the capability of this system as a promising technique for particle-phase HOM measurements. Taken together, we believe that this system is a promising technique for combined online gas- and particle-phase HOM measurements.

1 Introduction

Secondary organic aerosol (SOA), formed from condensable oxidation products of volatile organic compounds (VOCs), contributes a large fraction to tropospheric fine particles, which greatly influence human health and global climate (Mauderly and Chow, 2008; Jimenez et al., 2009; Boucher et al., 2013). Recently, highly oxygenated organic molecules (HOM) formed through rapid autoxidation (i.e. consecutive intramolecular H-shifts followed by O₂ addition) of peroxy radicals (RO₂) were found to play an important role in the growth and/or formation of new particles in the atmosphere (Crounse et al., 2013; Ehn et al., 2014; Kulmala et al., 2014; Jokinen et al., 2014; Kirkby et al., 2016; Bianchi et al., 2019). RO₂ are ubiquitous intermediates formed from the attack of various oxidants on VOC precursors, and therefore the fate of RO₂ and related formation pathways toward HOM is crucial to understanding the gas-phase chemistry and corresponding SOA formation in Earth's lower atmosphere (Orlando and Tyndall, 2012; Ehn et al., 2017).

Monoterpenes (C₁₀H₁₆), emitted from various terrestrial vegetation systems, account for ~15% of biogenic VOCs emissions (by mass) at a global scale (Guenther et al., 2012), and may contribute largely to fine organic aerosols (Ding et al., 2014; Zhang et al., 2018; McFiggans et al., 2019). α -pinene (AP) is one of the most important monoterpenes, which has been shown to form C₁₀ monomers and C₂₀ dimers as HOM products (Ehn et al., 2014). An endocyclic double bond and relatively large carbon skeleton increases the yield of these low volatile HOM products compared to many other biogenic VOCs (e.g.



isoprene C_5H_8 , the largest biogenic gas precursor), making it important in terms of SOA formation (Donahue et al., 2012; Donahue et al., 2013; Jokinen et al., 2015; Kurten et al., 2016). Furthermore, the ozonolysis of α -pinene produces aldehydic H atoms, which greatly promotes H-shifts in the RO_2 , and consequently increases HOM yields (Mentel et al., 2015; Rissanen et al., 2014; Bianchi et al., 2019). Consistently, HOM yields from other oxidants (e.g. OH, NO_3 , and Cl) reacting with α -pinene are typically lower (Jokinen et al., 2015; Berndt et al., 2016; Kurtén et al., 2017; Wang et al., 2020). These recent HOM yield findings are also consistent with earlier SOA mass yields, which have showed that O_3 forms more SOA than the other oxidants (Hoffmann et al., 1997; Griffin et al., 1999; Presto et al., 2005).

The understanding of HOM formation from ozonolysis of α -pinene has gradually improved during the past decade, owing to the developments/improvements of state-of-the-art mass spectrometers and quantum chemical calculations (Ehn et al., 2012; Vereecken et al., 2015; Riva et al., 2019; Bianchi et al., 2019; Iyer et al., 2021). The nitrate chemical ionization inlet combined with an atmospheric pressure interface time-of-flight mass spectrometer (NO_3 -CIMS for short hereafter) has been widely used to detect gas phase HOM as adduct clusters with NO_3^- with high sensitivity and selectivity (Junninen et al., 2010; Jokinen et al., 2012; Ehn et al., 2014; Hyttinen et al., 2017). Other reagent ions (e.g. CH_3COO^- , $C_3H_7NH_3^+$, and NH_4^+) have also been used with different sensitivities/selectivities to increase our understanding of HOM species and autoxidation (Berndt et al., 2016; Berndt et al., 2018; Riva et al., 2019). However, the characterization of HOM in the particle phase as well as the physicochemical links between gas- and particle-phase HOM are still not clear. One of the difficulties is the lack of suitable analytical techniques to identify the labile peroxides in α -pinene SOA (Docherty et al., 2005; Reinnig et al., 2009; Krapf et al., 2016), which are formed literally after each H-shift of RO_2 during HOM formation in the gas phase (Crounse et al., 2013; Mentel et al., 2015; Wang et al., 2017). The lifetime of organic peroxides in α -pinene SOA was reported to be around one hour through dark decomposition (Krapf et al., 2016). Formation of high molecular weight ester dimers with aldehydes/ketones through the Baeyer-Villiger reactions was proposed as another particle-phase loss of these hydroperoxides (Claflin et al., 2018). The concomitant formation of carboxylic acids through the Baeyer-Villiger reactions is consistent with the fact that a large portion of organic acids was observed in ambient SOA (Hallquist et al., 2009; Yatavelli et al., 2015). However, the experiments conducted by Claflin et al. (2018) were under conditions of high VOC concentrations (corresponding 3 mg m^{-3} SOA), and whether the proposed reactions can happen under ambient conditions is still unknown. Another study suggested that some HOM species (assuming being keto-hydroperoxides) may undergo decomposition through the Korcek mechanism in the particle phase, which forms C_7 - C_9 carboxylic acids and C_1 - C_4 carbonyls (Mutzel et al., 2015).

Here, a novel particle evaporation inlet named VIA (Volatilization Inlet for Aerosols, Aerodyne Research, Inc.) combined with a NO_3 -CIMS was used to detect HOM in the particle phase at trace concentrations. There are several special features of this system compared to currently used/developed techniques (Häkkinen et al., 2022). Traditional molecular level SOA measurements were usually conducted in an offline or semi-online way, during which filter/metal surfaces are necessary to collect particles (Hallquist et al., 2009; Yatavelli et al., 2012; Mutzel et al., 2015). We tried to minimize unwanted surface reactions of SOA by heating it directly within the sampling flow after removing the gases, which is important for the detection of labile organic compounds. Another online technique that was recently developed is the Extractive Electrospray Ionization (EESI) inlet combined with a mass spectrometer, aim at providing rapid online molecular detection of SOA (Lopez-Hilfiker et al., 2019; Pospisilova et al., 2020) and updated to a dual-phase inlet by Lee et al. (2022). The EESI has proven very useful for SOA detection, but it cannot provide comparable gas-phase measurements. Both gas- and particle-phase organic molecules could be detected by a Filter Inlet for Gases and Aerosols (FIGAERO) coupled with an iodide/acetate CIMS (Lopez-Hilfiker et al., 2014; Lopez-Hilfiker et al., 2015) or a Chemical Analysis of Aerosol Online (CHARON) inlet coupled with a Proton-Transfer-Reaction (PTR) mass spectrometer (Eichler et al., 2015; Müller et al.,



2017). These are robust and powerful systems for online (CHARON) or semi-online (FIGAERO) SOA measurements, though each reagent ion and setup provides its own selectivity and sensitivity, with neither being specifically selective towards HOM (Riva et al., 2019). Recently, a Thermal Desorption-Differential Mobility Analyzer (TD-DMA) coupled with a NO_3 -CIMS was deployed to detect nano (~ 10 -30 nm) SOA particles in a semi-online approach (Wagner et al., 2018; Caudillo et al., 2021). Similar to our system, nitrate was used as the reagent ion for both gas- and particle-phase measurements, reducing ambiguities when comparing HOM species measured in the gas versus particle phase. The advantage of the TD-DMA system is that it uses a DMA to size-selected particles, but the drawback is the lowered (bulk) sensitivity due to the need for charging in the DMA. Thus, the novel VIA inlet coupled with a NO_3 -CIMS will be a promising system to identify HOM species in the particle phase, using a fully online approach.

In this work, in order to provide molecular speciation of HOM in the particle phase and deduce its dependence on HOM formed in the gas phase, we systematically studied HOM products in both phases using the novel VIA- NO_3 -CIMS system. From our earlier experiments, higher dimer than monomer concentrations were observed in the particle phase for four different VOCs using this system (Häkkinen et al., 2022), in contrast to the measurements in the gas phase. However, those experiments were mainly focused on the particle phase and were conducted with high precursor conditions using a Potential Aerosol Mass (PAM) reactor (Kang et al., 2007; Lambe et al., 2011). Here, starting with relatively lower concentrations of α -pinene and O_3 (dark reactions) as our base case in a Teflon chamber, CO or NO_2/NO was introduced into this reaction system to change the distribution of HOM products, meanwhile investigate the impacts in both the gas and particle phases. HOM products were grouped according to different carbon numbers, along with dimer to monomer ratios (D/M), to compare their distribution in different phases. In addition, bulk chemical compositions (O/C, H/C, and N/C) of HOM in the particle phase were estimated and compared to a widely used Aerodyne Aerosol Mass Spectrometer (Decarlo et al., 2006).

2 Experimental section

2.1 Experimental setup

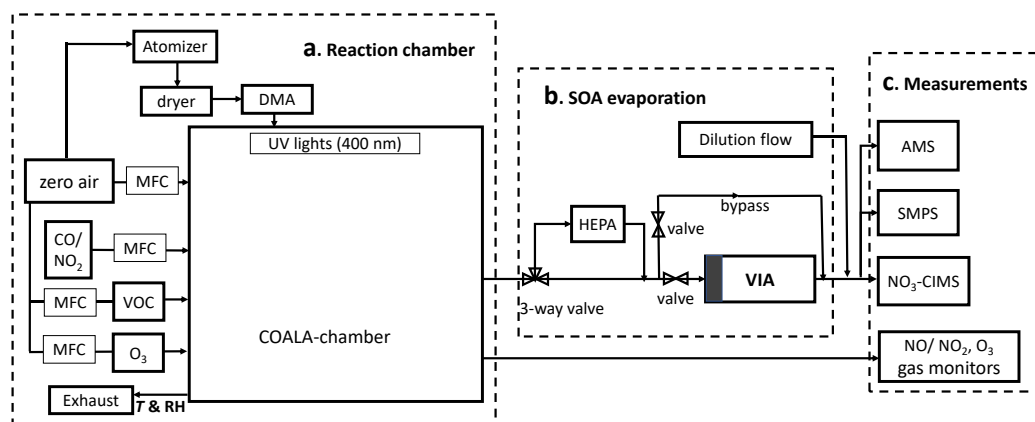


Figure 1. Setup for the thermal evaporation experiments used in this work. Note that all three parts (a, b, and c) are necessary for the particle-phase HOM measurements, whereas only part a and c are used for the gas-phase HOM measurements (i.e. all instruments sampling directly from the chamber).

The setup for experiments conducted in this work includes three parts as shown in Figure 1. a) Reaction chamber (2 m^3 , Teflon), in which HOM and SOA were formed from α -pinene ozonolysis. The experiments were operated in a continuous



flow mode (residence time ~50 min) under dry conditions ($RH < 2\%$) and at room temperature ($26 \pm 1^\circ\text{C}$). The total inflow of 40-L min^{-1} cleaned air was purified by a clean air generator (AADCO model 737-14, Ohio, USA). α -pinene was injected using a syringe (Hamilton) and a syringe pump (Cole Pamer, IL, USA). O_3 was generated by an ozone generator (Dasibi 1008-PC). Sodium chloride solution (NaCl in Milli-Q water) was used to generate seed particles (80 nm) to make the HOM
 115 condense onto larger particles rather than forming new (small) particles. NaCl was used instead of the more common ammonium sulfate (AS) to prevent the potential depleting of reagent ions by large amounts of sulfuric acid, and to avoid its nucleation with evaporated HOM vapors. CO or NO_2 was injected into the chamber from a gas cylinder during some of the experiments to change the gas-phase HOM chemistry. NO could be generated from photolysis of NO_2 with UV lights (400
 120 nm, LEDLightmake.inc). The entire flow system was controlled by a series of mass flow controllers (MKS, G-Series, $0.05\text{-}50\text{ L min}^{-1}$, Andover, MA, USA). For more detailed descriptions of the “COALA” chamber system and facilities at the University of Helsinki, please refer to Riva et al. (2019) and Peräkylä et al. (2020).

b) SOA evaporation system, a novel VIA inlet was deployed to evaporate particles for analysis. First, the gaseous vapors were removed from the sample flow ($1.5\text{-}2\text{ L min}^{-1}$) by a honeycomb-activated charcoal denuder (shaded region in Fig. 1,
 125 part b) in front of the heating region within the VIA. The remaining particles were then heated to a chosen temperature (25 to 300°C , measured from the surface of the heating tube by a thermocouple) causing evaporation of molecules in the particles. After the VIA, a dilution flow (10 L min^{-1}) was introduced, in part to cool down the sample before entering the NO_3 -CIMS, and to provide enough flow for all instruments.

Based on our tests, the removal efficiency of this charcoal denuder depends on the flow rate through VIA as well as the
 130 concentrations and types of the target gas. The best removal performance was obtained with a small flow rate ($1\text{-}2\text{ L min}^{-1}$). We observed that $>95\%$ of $\sim 4\text{ ppm O}_3$ and $>80\%$ of $500\text{-}800\text{ ppb } \alpha\text{-pinene}$ were removed at 1.5 L min^{-1} . Thus, the gas denuder worked efficiently under low concentrations used in this work (Table 1) and gas-phase reactions within the VIA were assumed to be negligible. The particle transmission efficiency of VIA was $\sim 90\%$ tested with room air ($10\text{-}200\text{ nm}$ particles). The denuder was replaced with a new/cleaned one before the start of each experiment to prevent potential
 135 contaminations, and it can be regenerated by heating to 100°C in a clean air flow for about four hours. Using valves, we could either pass the sampled air through a high efficiency particulate air (HEPA) filter to remove particles and through the VIA to remove the gaseous vapors for a background measurement, or bypass the HEPA and VIA completely to measure the unperturbed sample. However, to avoid the additional losses and dilution, direct gas-phase HOM measurements were done by completely bypassing part b) of the setup. A more detailed characterization of this VIA- NO_3 -CIMS system was
 140 investigated in a previous work (Häkkinen et al., 2022).

c) Measurements, a suite of gas and particle phase instruments were deployed. Gas phase and evaporated HOM from SOA were measured by a NO_3 -CIMS (Tofwerk AG/Aerodyne Research, Inc.), which is highly sensitive to detect HOM as adduct clusters with NO_3^- (Junninen et al., 2010; Jokinen et al., 2012). Nitrate ions were generated from nitric acid (HNO_3) through an X-ray source. The concentrations of HOM were estimated by applying a calibration factor $C_x (= 2 \times 10^{10}\text{ cm}^{-3})$ to the raw
 145 signals after being normalized by the sum of reagent ions (i.e. nitrate monomer, dimer, and trimer). This calibration factor was estimated using the sulfuric acid signals obtained by the VIA- NO_3 -CIMS system from several nanograms of AS assuming 70% evaporation at 300°C (Häkkinen et al., 2022). Thus, a large uncertainty comes with this estimation, along with other factors including the total losses of HOM during evaporation and sampling. We expect the uncertainty be at least a factor 3 ($+200\%\text{-}67\%$), and therefore the absolute concentrations are only reported in this work to give rough
 150 approximations of the sampled concentrations. SOA mass was measured by a long time-of-flight Aerosol Mass Spectrometer (LTOF-AMS, Aerodyne Research, Inc.), using thermal evaporation (600°C) and electron impact ionization (70 eV) techniques (Decarlo et al., 2006; Canagaratna et al., 2007). The (relative) ionization efficiency of AMS towards inorganic



species was calibrated with 300 nm ammonium nitrate (AN) and AS particles (Jayne et al., 2000). The particle size distribution (10-500 nm) was measured with a scanning mobility particle sizer system (SMPS, consisting of one 3081 long differential mobility analyzer and one 3750 condensation particle counter, TSI). O_3 , NO , and NO_2 were monitored by a UV photometric analyzer (Model 49P, Thermo Environmental Instruments), an Eco Physics gas analyzer (Model CLD 780 TR), and a chemiluminescence $NO-NO_2-NO_x$ analyzer (Model 42i, Thermo Fisher Scientific), respectively. In addition, a proton transfer reaction time-of-flight mass spectrometer (PTR-TOF 8000, Ionicon Analytik GmbH) was used to measure α -pinene. Unfortunately, it worked intermittently and was only connected to the chamber twice during the first experiments. The concentrations of α -pinene in the first experiment were evaluated based on 5-min averaged signals from the data acquisition panel applying the sensitivity obtained from the calibration conducted that morning (two brown cross markers in Fig. 2b). For the rest experiments, we used the syringe pump injection rates to estimate the injected concentrations of α -pinene.

2.2 Chamber experiments

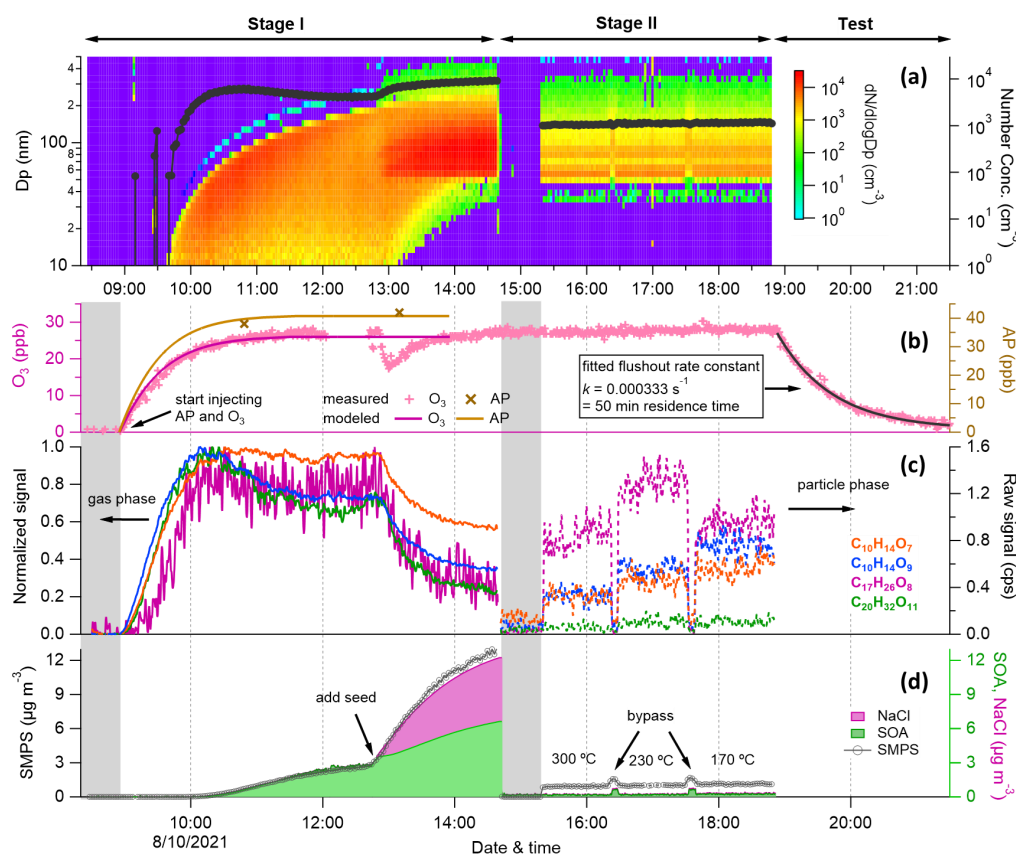


Figure 2. Overview of experiment No. 1 (input flow with 53 ppb α -pinene and 33 ppb O_3). (a) Particle number concentration and size distribution sampled by SMPS as a function of time. Time series of (b) measured and modeled ozone and α -pinene concentration (1-min averaged) in the chamber, (c) gas-phase (solid lines, normalized to their maximums) and particle-phase (dashed lines, raw signals) HOM species (10-s averaged), (d) total aerosol, organics, and sodium chloride mass concentrations from SMPS (2.2-min averaged, black circles) and AMS measurements (20 s averaged). The first and second shaded areas are gas phase and particle phase background measurements, respectively. Note that the generally lower particle phase measurements in stage II compared to stage I are owing to the addition of a dilution flow (Fig. 1), and no measurements given above were corrected for chamber wall loss.



A typical experiment was conducted as shown in Figure 2, including two stages. In stage I, all instruments sampled directly from the chamber (i.e. part b in Fig. 1 was removed from the system). The COALA chamber was flushed with purified air for at least 12 hours before each experiment (chamber background was determined during this period, first shaded area in Fig. 2). HOM formed rapidly once α -pinene and O_3 were injected into the chamber. SOA particles took ~40 minutes to form and grow larger than 10 nm to be measured by an SMPS, whereas it took about 1.5 hours for the AMS to capture the SOA formation. Then, NaCl seed particles were added into the chamber after HOM and SOA concentrations stabilized, which can be seen as a dramatic increase of the particle mass in Figure 2d. However, the thermal desorption temperature of 600 °C used by the AMS was not enough for effective NaCl evaporation/measurements, resulting in discrepancies in aerosol mass concentration compared to the SMPS (Fig. S1). The mass concentrations shown in Fig. 2d were modified based on AMS and SMPS measurements and the method to reconstruct SOA and NaCl mass is explained in the Supplementary (section S1). In contrast to the aerosol mass, gas-phase HOM decreased upon seed addition owing to enhanced condensation onto the newly added particles that increased the condensation sink. By calculating the difference in gas-phase HOM spectra before and after seed particle addition, we can deduce the condensing HOM products. We will use the term “condensed gas phase” hereafter to directly compare the condensing HOM to the particle-phase HOM measurements. In stage II, the VIA was mounted manually between the chamber and instruments for particle-phase HOM measurements as shown in Figure 1. After background measurements for the particle phase (second shaded area in Fig. 2), three different temperatures (170, 230, and 300 °C) were set for the VIA heater, each one kept for about one hour for data collection. Note that the analysis and following discussions mainly focused on the results of 230 °C, which was found to give the highest sensitivity/signals for total particle-phase HOM measurements with less losses or thermal decomposition compared to other temperatures. However, as shown in Fig. 2c, larger molecules (C_{17} and C_{20} compounds) were found to evaporate at higher temperatures, while signals of smaller or more volatile molecules (C_{10} compounds) started to decrease at the highest temperatures. This effect is consistent with the previous results reported by Häkkinen et al. (2022) and needs to be kept in mind when interpreting the particle-phase HOM mass spectra. Between the change of temperatures, the sampling line bypassed VIA to obtain SOA mass concentration without heating. Then, the evaporated SOA mass could be estimated by taking the difference of measurements between bypass and VIA modes. At the end of the experiment, a test stage was included for determining kinetic parameters for use in a box model (described in section 2.4), such as the flush-out rate (as shown in Fig. 2b) and the photolysis rate of NO_2 . All experiments conducted in this work are summarized in Table 1 and will be explained in section 2.4.

Table 1. Experimental Conditions for the COALA Chamber Experiments. The steady state conditions refer to the stabilized concentrations before adding seed particles.

No.	Input flow (in) conditions					steady state (ss) conditions				
	[AP] _{in} (ppb)	[O ₃] _{in} (ppb)	[NO ₂] _{in} (ppb)	[CO] _{in} (ppm)	UVlights (400 nm)	[AP] _{ss} (ppb)	[O ₃] _{ss} (ppb)	[NO ₂] _{ss} (ppb)	[NO] _{ss} (ppb)	[SOA] _{ss} ($\mu g m^{-3}$)
1	53	33	0	0	off	41	26.5	0	0	2.84
2	80	33	0	0	off	63.6	24.0	0	0	7.01
3	120	33	0	0	off	98.3	20.5	0	0	11.83
4	120	33	0	100	off	101.2	20.2	0	0	9.30
5	120	33	7.6	0	off	97	21.9	5.3	0	8.40
6	120	33	7.7	0	on	96.2	33.1	3.3	0	10.9
7	120	33	17.3	0	on	95.3	46.6	8.1	0	13.95
8	120	33	46.6	0	on	92.9	66.0	28.5	1.4	7.50
9	53	33	46.3	0	on	38.9	51.6	35.4	2.5	1.34



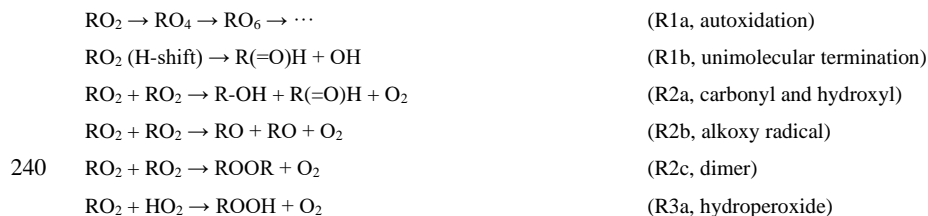
2.3 Data analysis

The mass spectrometry data sets were analyzed using Igor Pro (WaveMetrics, Inc., USA) based packages (Tofware_3_2_0 and ToF_AMS_HRAnalysis_v1_25A). Based on the mechanisms of HOM formation, “HOM” was suggested by Bianchi et al. (2019) to describe compounds directly related to gas phase autoxidation (Crounse et al., 2013) in atmospheric conditions and with at least six oxygen atoms, so that could be effectively detected by a NO₃-CIMS as used in this work. Although organic compounds with O numbers less than six (e.g. norpinic acid C₈H₁₂O₄ and pinic acid C₉H₁₄O₄) were largely observed in both the gas and particle phase in α-pinene ozonolysis (Lopez-Hilfiker et al., 2015; Zhang et al., 2015), these compounds would not be regarded as HOM, nor would they be effectively detected by a NO₃-CIMS (Hyytinen et al., 2015; Hyytinen et al., 2017). For experiments with NO_x, measured HOM signals were separated into two groups as HOM_{CHO} (C_xH_yO_z compounds) and N-containing HOM_{ON} (C_xH_yN₁₋₂O_z compounds), based on the number of N atom in the fitted molecular formula excluding the reagent ion (NO₃⁻). The peak positions located at odd or even nominal mass also helped to separate these compounds. Note that the contribution of N-containing peaks was only considered for NO_x experiments, owing to the relatively low contribution of nitrate dimer (HNO₃·NO₃⁻) charged signals to the total signals (4.3±1.5%) for α-pinene ozonolysis experiments. In addition, HOM products were grouped according to different carbon numbers to investigate their distribution. C₈-C₁₀ and C₁₆-C₂₀ compounds were regarded as HOM monomer and dimer, respectively, to be consistent with previous research works (Zhao et al., 2018b; Molteni et al., 2019). Thus, HOM species discussed in this study were limited to molecules with a formula as C_{x>7}H_{y>7}N₀₋₂O_{z>5} (excluding the NO₃⁻ reagent ions), which contributed 83-91% of the total measured signals (shaded areas in Fig. 3).

Finally, elemental ratios were calculated in order to compare the bulk chemical composition of SOA measured by the AMS and the VIA-NO₃-CIMS system. Molar O/C, H/C, and N/C ratios of AMS were calculated using the Improved-Aiken method (IA), with uncertainties of 28%, 13%, and 22%, respectively (Aiken et al., 2008; Canagaratna et al., 2015). Nitrate signals measured by the AMS are almost all from organic nitrates, which are consistent with the estimation of NO_x⁺ ratio method (Farmer et al., 2010; Day et al., 2022). Molar O/C, H/C, and N/C ratios of the NO₃-CIMS measurements (signal weighted) were calculated based on the fitted chemical formulas. The uncertainty mainly comes from the accuracy of high-resolution fitting results, relative calibration factors among different HOM compounds (instead of the absolute calibration factors), and the signal-to-noise ratios of each compound.

2.4 Kinetic model and RO₂ chemistry

A simple 0-D kinetic box model was used to model α-pinene concentrations and support the interpretation of the experimental results. The chemical reactions and physical parameters are described in Section S2. As shown in Figure 2b, the modeled α-pinene is consistent with the PTRTOF measurements. Based on the only two α-pinene measurements during the first experiment, modeled α-pinene concentrations were estimated with an accuracy of 10%, though this will probably be less accurate for the NO_x experiments because more reactions need to be considered in the model. The main reactions that will be discussed in the following sections are listed below:





245 Peroxy radicals are the key intermediate in α -pinene ozonolysis leading to HOM formation (Ehn et al., 2017; Iyer et al., 2018; Berndt et al., 2018). They can undergo unimolecular autoxidation forming more oxidized RO_2 (Eq. R1a) or termination forming monomer products (Eq. R1b). They can also undergo bimolecular reactions (Eq. R2,3,4) with $\text{RO}_2/\text{HO}_2/\text{NO}$ to form closed-shell monomer/dimer products or alkoxy radicals (RO), which may or may not form another RO_2 to restart this cycle (Mentel et al., 2015; Berndt et al., 2018; Bianchi et al., 2019). The branching ratios depend on the concentration of reaction

250 partners.
 In α -pinene ozonolysis, the fate of RO_2 is mainly through unimolecular (Eq. R1a,b) and bimolecular (Eq. R2a,b,c) reactions. Different α -pinene concentrations will change the relative importance of these two kinds of pathway. Thus, the concentration-dependent distribution of HOM products in both phases were investigated in experiment No. 1-3 (α -pinene, 53-120 ppb). In experiment No. 4 with CO as an OH scavenger, ozonolysis was largely isolated so that we could focus on
 255 O_3 -derived RO_2 and corresponding HOM products. Note that added CO will turn OH to HO_2 , thus making HO_2 a competitive partner to react with RO_2 (Eq. R3a,b). In experiment No. 5, NO_2 was added under dark conditions to introduce NO_3 (formed from the reaction of NO_2 and O_3) initiated oxidation forming nitrooxy peroxy radicals (nRO_2), which has been proposed to not undergo autoxidation very efficiently (Kurtén et al., 2017). Nevertheless, the fate of nRO_2 through bimolecular (Eq. R2a,b,c) reactions with another RO_2/nRO_2 would still be expected. Finally, NO was formed after turning on the UV lights
 260 (experiment NO. 6-9, inflow $\text{VOC}/\text{NO}_x = 15.6\text{-}1.1$), which partly converted RO_2 to organic nitrates as N-containing HOM monomers (Eq. R4a) instead of through bimolecular (Eq. R2a,b,c) reactions with another RO_2 . Relatively lower O_3 to α -pinene ratios (0.28-0.62) than their distribution in the real atmosphere were used in this work to reduce reactions between NO and O_3 , which maximizes bimolecular reactions between NO and RO_2 . Note that the gas-phase HOM were measured in experiments No. 6-9 with both lights off and on, whereas the particle phase measurements were only obtained with the lights
 265 on. Experiment No.5 was conducted with lights off during the entire experiment, which is the only one containing the information about particle-phase HOM products from NO_3 and α -pinene reactions without NO chemistry. In the following sections, both the identification of gas and particle phase HOM products and their relationship will be investigated in detail for each reaction system.

3. Results

270 3.1 HOM distribution under different conditions

3.1.1 α -pinene and O_3 dark reactions

Mass defect plots of particle-phase signals measured by the VIA- NO_3 -CIMS system from α -pinene ozonolysis are given in Fig. 3 a-c, which showed large differences compared to gas phase measurements (Fig. S2). In the gas phase, $\text{C}_{10}\text{H}_{14,16}\text{O}_z$ dominates HOM monomers followed by $\text{C}_8\text{H}_{12}\text{O}_z$, while the dimers (much lower than monomers) were almost evenly spread
 275 among $\text{C}_{17}\text{-C}_{20}$ compounds (Fig. 3d). However, in the particle phase, dimers (dominated by $\text{C}_{17}\text{H}_{26}\text{O}_z$ followed by $\text{C}_{18,19}\text{H}_{26,28}\text{O}_z$) showed comparable or even higher contribution to total HOM signals than monomers (mainly C_{10} compounds). Similarly, higher dimer than monomer signals of α -pinene SOA was measured in sodiated ion adducted ($[\text{M} + \text{Na}]^+$, electrospray ionization) mass spectra (Zhang et al., 2017). In addition, $\text{C}_{16}\text{H}_{24,26}\text{O}_z$ was almost exclusively observed in the particle phase, indicating particle-phase related sources of these compounds. Oxygen numbers also changed, peaking



around O₆-O₇ in the gas phase but shifted to O₇-O₁₀ in the particle phase (Fig. 3f). Nevertheless, the signal-weighted O/C of the particle-phase HOM is lower than that of the gas phase (details in section 3.2), owing to large shift of C numbers in the particle phase. Last, H numbers (Fig. 3d) are highly correlated with C numbers for HOM products both in the gas and particle phases, as expected. And a small amount of C₁₀H₁₅O₂ and C₉H₁₃O₂ peroxy radicals were measured in the gas phase but were negligible in the particle phase.

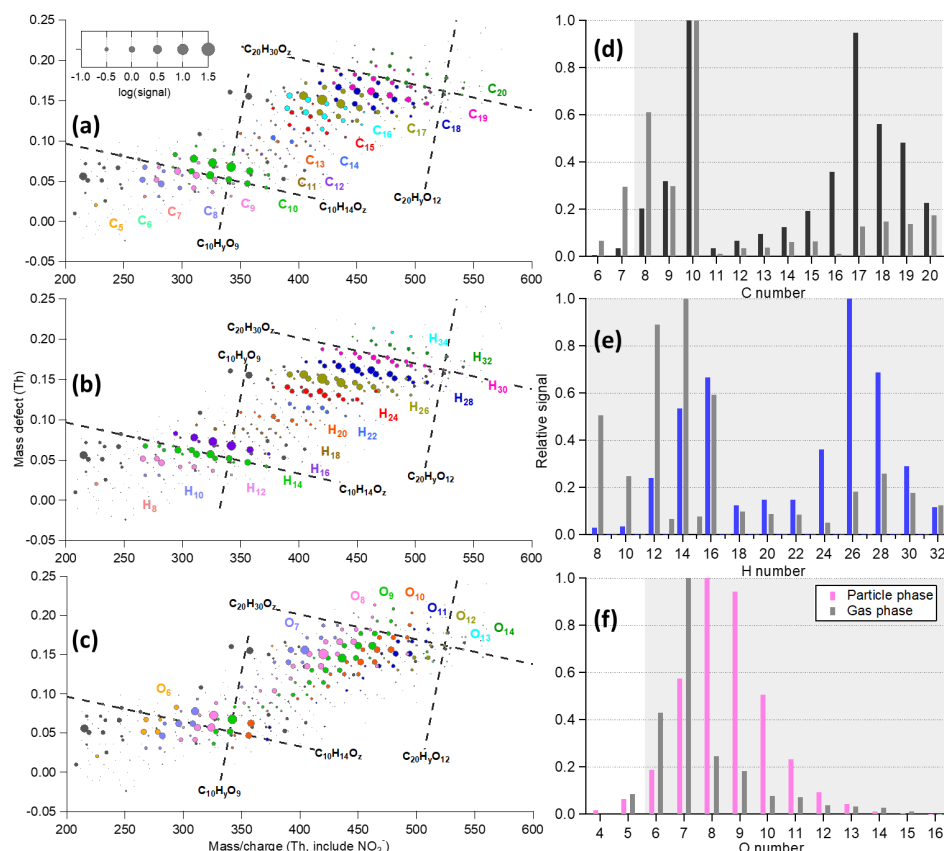


Figure 3. (left) Mass defect plot of particle phase signals (60-min average, 230 °C) measured in experiment No. 3 (input flow concentration as 120 ppb α -pinene and 33 ppb O₃). The data points are identical in panels a-c, but colored by the (a) carbon number, (b) hydrogen number, and (c) oxygen number. The area of the markers is proportional to the logarithm of measured raw signals. The grey markers are either unknown peaks (unidentified) or not HOM compounds according to the nomenclature of “HOM” (Bianchi et al., 2019). (right) Distribution of signals (normalized to the largest peak) grouped by (d) carbon number, (e) hydrogen number, and (f) oxygen number. The particle phase measurements are shown as colored bars, while the gas phase measurements are shown in grey. The shaded areas in panels d-f are the focus of this study.

Many elemental compositions identical to the monomers and dimers identified in this work were also reported by other mass spectrometry techniques (summarized in Table S2), though in general, species reported in this work contained more oxygen than in previous publications (Müller et al., 2009; Zhang et al., 2015; Mutzel et al., 2015; Kristensen et al., 2020; Pospisilova et al., 2020). For example, ester dimers with molecular formulas as C₁₇H₂₆O₇ and C₁₇H₂₆O₈ were identified by an Ultra-High Performance Liquid Chromatography/Electrospray Ionization quadrupole Time-of Flight Mass Spectrometer (UHPLC/ESI-qToF-MS) (Kristensen et al., 2020; Kahnt et al., 2018). However, the structure of many ester dimers with 5-7 oxygen atoms were proposed to contain carboxylic functional groups at both ends (Kristensen et al., 2016; Zhang et al., 2015), which would



possibly be “invisible” by a NO_3 -CIMS owing to large difficulties to form two hydrogen bonds with the reagent ions (Hyttinen et al., 2015). It remains unclear to what extent the identical elemental compositions reflect identical isomeric structures, given the differences between the methods. Nevertheless, the similarities in detected compositions are encouraging, even though one needs to be careful to not draw direct links to previous observations using different methods.

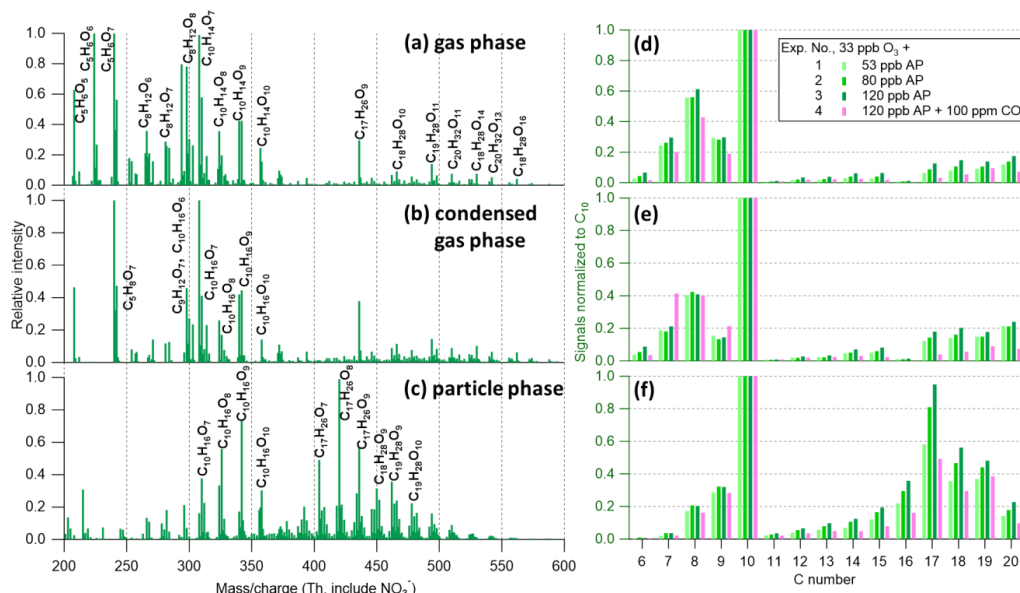


Figure 4. (left) Mass spectra of (a) gas phase (without seed particles), (b) condensed gas phase (from the difference between gas-phase mass spectra without and with seed particles, see section 2.2), and (c) particle phase (at 230 °C) measurements from experiment No.3 (input flow concentration as 120 ppb α -pinene and 33 ppb O_3). Gas phase and condensed gas phase mass spectra (30-min average) are normalized to $(\text{C}_{10}\text{H}_{14}\text{O}_7)\text{NO}_3$ at m/z 308 with m/z 224 and m/z 240 off the scale. Particle phase mass spectra (60-min average) is normalized to $(\text{C}_{17}\text{H}_{26}\text{O}_8)\text{NO}_3$ at m/z 420. (right) The distribution of HOM products for experiments No. 1-4, grouped by C numbers, are summarized in (d) for gas phase, (e) for condensed gas phase, and (f) for particle phase. Signals are normalized to the C_{10} family.

Different initial concentrations of α -pinene were used to investigate the concentration-dependent distribution of HOM products, summarized in Fig. 4d-f. In the gas phase, as more α -pinene reacted, more RO_2 was formed in the system, thus making self/cross reactions between two RO_2 more important, while decreasing the autoxidation (Eq. R1a) and unimolecular termination (Eq. R1b). We indeed see a lower fraction of compounds with large O numbers in the gas phase when increasing the α -pinene concentrations (Fig. S3), while C_8 monomers (likely from Eq. R2b followed by fragmentation) and all dimers (from reaction Eq. R2c) increased (Fig. 4d). The concentration-dependent trend of HOM distributions in the gas phase is consistent to that observed at relatively low α -pinene conditions (17-1692 ppt) with ~ 30 ppb O_3 (Molteni et al., 2019), despite clearly higher α -pinene (53-120 ppb) and O_3 (~ 33 ppb) concentrations used in this work.

When adding NaCl seed particles to convert more HOM vapors into SOA, mass spectra of the condensable gas phase could be obtained (i.e. the difference of gas-phase measurements between without and with seed particles, Fig. 4b). Low volatile monomers and almost all dimers (with roughly $m/z > 350$ Th including the reagent ion NO_3^-) showed comparable, largely kinetically limited, and irreversible condensation loss on the seed particles, similar to results from an earlier study (Peräkylä et al., 2020). The mass spectra of condensed gas-phase HOM were in general comparable to those of the gas phase (Fig. 4b vs. 4a and 4e vs. 4d), though with slightly higher dimer contribution. However, the particle-phase HOM distribution was dramatically different from the (condensed) gas phase measurements, and measured dimer concentrations were even higher



than monomers (Fig. 4f). With more α -pinene reacted, the enhancement of C_{16} - C_{19} dimers in the particle phase was much larger than that observed in the gas phase, while the particle-phase O number distribution did not show obvious concentration-dependent changings among different experiments (Fig. S3). In addition, $C_{17}H_{26}O_9$ was identified in both gas and particle phases, whereas $C_{17}H_{26}O_{7,8}$ were only detected in the particle phase. These compounds have also often been identified from filter measurements (Müller et al., 2009; Zhang et al., 2015; Mutzel et al., 2015; Kristensen et al., 2020). All these particle and gas phase comparisons raise an interesting question-- what is the source of those C_{16} - C_{19} dimers? The evenly distributed C_{17} - C_{20} HOM dimers in the gas phase and their extremely low volatilities (Peräkylä et al., 2020; Hyttinen et al., 2021) would have us expecting a similar distribution in the particle phase. However, the very different HOM distribution measured in the two phases implies that particle phase reactions are taking place. We cannot rule out that our VIA sampling setup is perturbing the true distribution, and these limitations are discussed in more detail in section 3.4. Earlier studies have speculated on different processes forming dimers in the particle phase. For example, $C_{17}H_{26}O_8$ (nominal mass at 358 amu, m/z 420 in Fig.4), which showed the highest signal in the dimer range (negligible in the gas phase), was also reported as the most abundant oligomer in α -pinene SOA measured by a High-Performance Liquid Chromatography and Fourier Transform Ion Cyclotron Resonance-Mass Spectrometry (HPLC/FTICR-MS) (Müller et al., 2009). In that study, liquid chromatography followed by mass spectrometry showed that $C_8H_{12}O_4$ and $C_9H_{14}O_4$ (cis-pinic acid and isobaric compounds detected at m/z 185) are monomeric units that make up this C_{17} dimer. It was later proposed that $C_8H_{12}O_4$ could be terpenylic acid (and isobaric compounds detected at m/z 171) (Yasmeen et al., 2010; Beck and Hoffmann, 2016). However, the proposed esterification process needs aqueous and acid conditions to turn the lactone group of terpenylic acid into two alcohol moieties (Beck and Hoffmann, 2016), which is not available in this study (dry NaCl seed particles). In addition, terpenylic acid was proposed to be formed either through OH initiated oxidation of α -pinene or an O_3 initiated pathway of α -pinene oxide (yield ~2%) (Claeys et al., 2009). However, in the CO (used as an OH scavenger) experiment (No. 4, discussed in the next section), when both these pathways were very limited, $C_{17}H_{26}O_8$ was still observed as the highest dimer peak in Figure 5b. In addition, it has been shown in other studies that a dominate role of dimers (including $C_{17}H_{26}O_8$) in the particle phase were primarily measured from ozonolysis of α -pinene (Kristensen et al., 2014; Zhang et al., 2017).

Recently, Claflin et al. (2018) used derivatization-spectrophotometric methods to quantify the distribution of different functional groups in very high-concentration α -pinene SOA (~3 mg m^{-3}). They compared the measurements to the modeled results of the Master Chemical Mechanism and proposed that hydroperoxides/peroxycarboxylic acids along with aldehydes/ketones could be converted to carboxyl monomers and ester dimers through the Baeyer-Villiger reaction in SOA. This would be an explanation of those “newly formed” dimers measured in the particle phase, which may also be important under atmospheric conditions (i.e. low SOA concentrations). Furthermore, particle phase decomposition of more carbon-containing HOM compounds (e.g. C_{10} in the monomer range and C_{20} in the dimer range) was proposed as an important decay pathway of α -pinene SOA (Pospisilova et al., 2020), which could largely explain the relative enhancement of C_{16} - C_{19} comparing to C_{20} compounds. However, from their results, C_{16} - C_{17} signals peak after ~2 hours of the α -pinene ozonolysis in batch-mode experiments, whereas the residence time is ~1 hour for the continuous-mode experiments conducted in this work. Nevertheless, about half of the compounds were proposed to have half-lives of less than one hour (Pospisilova et al., 2020). If the thermal energy supplied during the evaporation of SOA within the VIA can promote this decay process, then the observed particle-phase HOM distribution would be expected to shift to lower C numbers.

Although the above discussion mainly focused on C_{17} compounds, C_{18} - C_{20} HOM dimers observed in the gas phase were still identified in the particle phase. Thus, the particle phase dimers measured in this work may be partly intact HOM dimers formed and condensed from the gas phase (Berndt et al., 2018; Jokinen et al., 2014), and partly particle-phase formed dimers



which are multifunctional organic compounds being detected by a NO_3 -CIMS (thus were regarded as “HOM”). However, as discussed in section 2.2, some part of the enhancement of C_{16} - C_{19} compounds observed in the particle phase is likely due to a higher detection efficiency for HOM dimers than monomers with the VIA setup at 230 °C. In order to understand this “dimer enhancement” better, we conducted experiments where we tried to modify the dimer/monomer (D/M) ratios of HOM in the gas phase, by adding CO or NO to the system (as discussed in the following sections), to investigate the effects of the corresponding D/M ratios in the particles.

3.1.2 α -pinene and O_3 dark reactions with CO

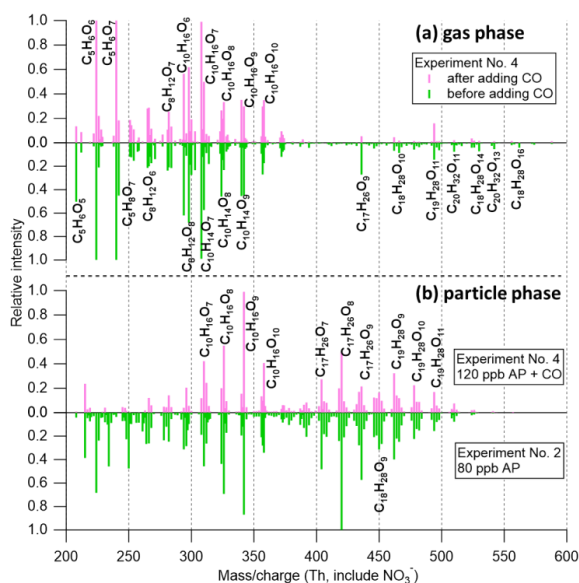


Figure 5. Mass spectra of (a) gas phase (30-min average) measurements before (green) and after (pink) adding 100 ppm CO in experiment No.4 (input flow concentration as 120 ppb α -pinene and 33 ppb O_3), normalized to $(\text{C}_{10}\text{H}_{14}\text{O}_7)\text{NO}_3$ at m/z 308 with m/z 224 and m/z 240 off the scale, and (b) particle phase (60-min average) measurements with (pink) and without (green) CO, normalized to $(\text{C}_{10}\text{H}_{16}\text{O}_9)\text{NO}_3$ at m/z 342 and $(\text{C}_{17}\text{H}_{26}\text{O}_8)\text{NO}_3$ at m/z 420, respectively. Note that oxidation rate of α -pinene decreased to about half due to the decreased OH oxidation after adding CO (assuming OH yield being ~0.8 from ozonolysis of α -pinene). Therefore, experiment No. 2 (80 ppb α -pinene) was chosen for the particle-phase comparison because of comparable SOA concentrations.

In experiment No. 4, CO was used as an OH scavenger (i.e. to isolate O_3 initiated reactions) to test the hypothesis that O_3 derived RO_2 and corresponding monomers are important for dimer formation in the particle phase (Kristensen et al., 2014; Zhang et al., 2017) and to change the gas-phase HOM distribution. α -pinene and O_3 were first added into the chamber similar to experiment No. 3, whereafter CO was added to largely shut down OH initiated formation of HOM. The added CO increases HO_2/RO_2 ratios compared to the pure α -pinene ozonolysis systems (roughly by a factor of 10^4 estimated by the kinetic model), making HO_2 a competitive partner to react with RO_2 (Eq. R3a,b). Thus, lower relative concentrations of dimers were observed (Fig. 5a). Total HOM concentrations also decreased after adding CO, likely due to the decreased oxidation rate of α -pinene (Fig. S4a). However, enhancement of C_{16} - C_{19} dimers compared to monomers was still observed in the particle phase (Fig. 5b and Fig. 4f). Furthermore, the peaks identified in the particle phase were almost the same as the ones without CO, strongly suggesting that ozonolysis products of α -pinene are important and perhaps necessary for the formation of these dimers (Kristensen et al., 2014; Zhang et al., 2017). In fact, the mass spectra of the CO experiment (No.4)



correlated well with the ones of experiment No. 2 without CO, with $r^2 = 0.82, 0.75$, and 0.80 for gas, condensed gas, and particle phase measurements, respectively.

In addition, although the majority of HOM products decreased largely in the gas phase after adding CO, $C_{10}H_{16}O_{8,10}$ increased significantly (Fig. S4a), which was also obvious from their time series (Fig. S5). This rapid increase with CO suggested that $C_{10}H_{15}O_{8,10}$ existed as dominant RO₂ in the gas phase, and were terminated rapidly by HO₂ forming hydroperoxides (Eq. R3a). On the other hand, $C_{10}H_{18}O_z$ and $C_{20}H_{34}O_z$ compounds were only formed through reactions of OH-derived RO₂ and therefor decreased after adding CO. Although $C_{10}H_{14}O_z$ and $C_{20}H_{30}O_z$ compounds were formed only through ozonolysis of α -pinene (Molteni et al., 2019), a general decrease was also observed, likely owing to the RO₂ autoxidation in some cases being outcompeted by HO₂.

3.1.3 α -pinene and O₃ reactions with NO (UV lights on)

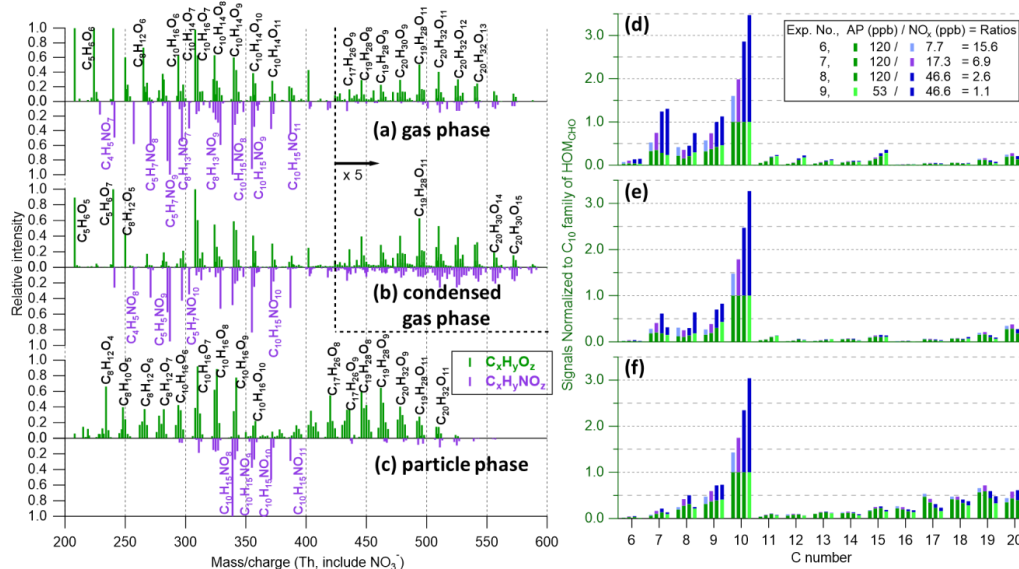


Figure 6. (left) Mass spectra of (a) gas phase (without seed), (b) condensed gas phase (from the difference between spectra without and with seed particles), and (c) particle phase (at 230 °C) measurements from experiment No.7 (input flow condition of 120 ppb α -pinene, 33 ppb O₃, and 17.3 ppb NO₂ with UV lights on). Gas phase and condensed gas phase mass spectra (30-min average) are normalized to (C₁₀H₁₄O₇)NO₃ at m/z 308 with m/z 240, and/or m/z 208 and m/z 224 off the scale, with peaks larger than 420 Th being multiplied by a factor of 5. Particle phase mass spectra (60-min average) are normalized to (C₁₀H₁₆O₇)NO₃ at m/z 310 with m/z 339 off the scale. Nitrogen-containing ions (C_xH_yNO_z, i.e. HOM_{ON}) are plotted downward while the other ions are plotted upwards. (right) The distribution of HOM products (varying initial α -pinene to NO_x ratios from 15.6 to 1.1), grouped by carbon numbers, in (d) gas phase, (e) condensed gas phase, and (f) particle phase. Signals are normalized in each experiment to the C₁₀ family of HOM_{CHO} compounds (green bars) and the blue bars (HOM_{ON}) stack on the top of the HOM_{CHO}.

NO (generated from the photolysis of NO₂) was introduced into the α -pinene ozonolysis system in experiments No. 6-9, and it acted similarly to HO₂ in the CO experiment to partly terminate RO₂ into monomers instead of forming dimers in the gas phase. Although a relatively smaller branching ratio towards the formation of HOM_{ON} monomers (Eq. R4a) than RO (Eq. R4b) would be expected, one more N atom makes the HOM_{ON} monomers (dominated by C₁₀H_{15,17}NO_z and C₉H_{13,15}NO_z compounds, in Fig. 6) easier to track, compared to the enhanced hydroperoxides in the CO experiment. As shown in Figure 6d, with lower initial AP/NO_x ratios, the contribution of HOM_{ON} monomers increased and HOM_{CHO} dimers decreased in the gas phase. Similar effects of NO-- terminating RO₂ and forming organic nitrates instead of dimers, thus suppressing SOA



formation-- were reported from previous studies (Ehn et al., 2014; Zhao et al., 2018a; Pullinen et al., 2020). The condensed gas phase followed the patterns of the gas-phase measurements.

Particle-phase HOM again showed relatively higher contributions of dimers (C_{16} - C_{20} compounds) compared to the gas phase (Fig. 6f vs. 6d). However, different from experiments No. 1-4, NO was used in experiments No. 6-9 to terminate RO_2 to RONO₂, which are not as effective particle-phase precursors as HOM_{CHO} monomers to form dimers through the Baeyer-Villiger reaction in SOA (Claflin et al., 2018). Consequently, less “enhancement” of C_{16} - C_{19} HOM_{CHO} compounds in the particle phase was observed compared to experiments without NO_x (Fig. 6f vs. 4f). In addition, highly functionalized organic nitrates observed in the particle phase have been found to decompose within 2-4 hours based on ambient measurements (Lee et al., 2016), suggesting that the formation of HOM_{ON} might be larger than we detected.

3.1.4 α -pinene and NO₃ dark reactions (UV lights off)

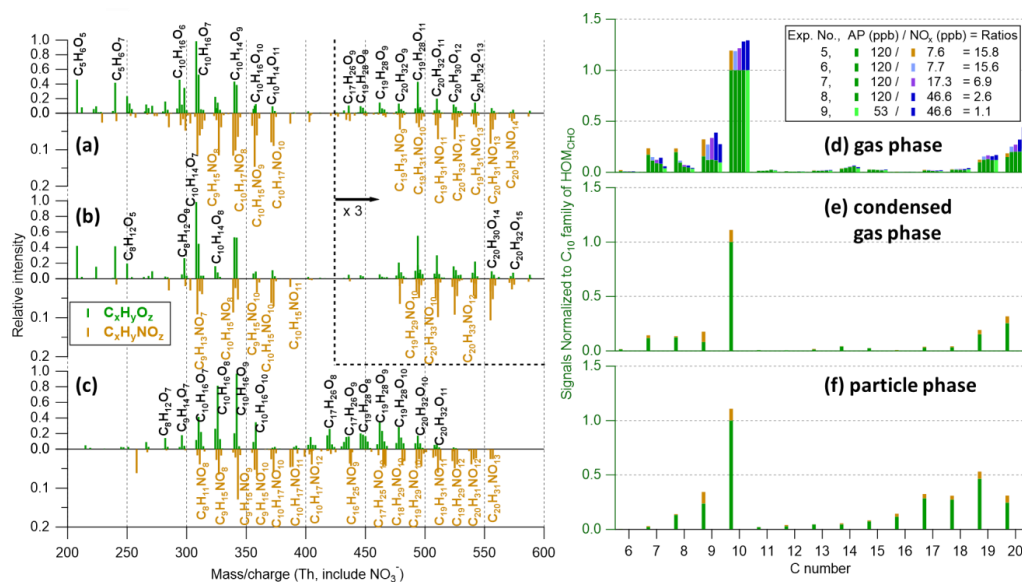


Figure 7. (left) Mass spectra of (a) gas phase (without seed), (b) condensed gas phase (from the difference between spectra without and with seed particles), and (c) particle phase (at 230 °C) measurements from experiment No.5 (input flow condition of 120 ppb α -pinene, 33 ppb O₃, and 7.6 ppb NO₂ with UV lights off). Note that the y-axis of HOM_{CHO} (green spectrum, upward) and HOM_{ON} (brown spectrum, downward) are on different scales. Gas phase and condensed gas phase mass spectra (30-min average) are normalized to (C₁₀H₁₄O₇)NO₃ at m/z 308, with peaks larger than 420 Th being multiplied by a factor of 3. Particle phase mass spectra (60-min average) are normalized to (C₁₀H₁₆O₉)NO₃ at m/z 342. (right) The distribution of HOM products (varying initial α -pinene to NO_x ratios from 15.8 to 1.1), grouped by carbon numbers, in (d) gas phase, (e) condensed gas phase, and (f) particle phase. Signals are normalized to HOM_{CHO} C₁₀ family. The UV lights were off during the entire experiment, making experiment No.5 (HOM_{ON} as brown bars) the only one with particle-phase HOM measurements from α -pinene and NO₃ reactions.

Particle-phase HOM from NO₃ initiated oxidation was investigated by adding NO₂ into the α -pinene ozonolysis system under dark conditions (i.e. UV lights off) in experiment No. 5. NO₃ oxidation of alkenes almost exclusively undergoes electrophilic addition to the double bond instead of H abstraction (Kerdouci et al., 2010), thus multigeneration oxidation by NO₃/O₃ should be negligible. In experiments No. 6-9, α -pinene, O₃, and NO₂ (7.6-46.6 ppb) were added into the chamber until a steady state was reached, whereafter the UV lights were turned on to photolyze NO₂ and NO₃ to NO (results discussed in section 3.1.3). Thus, only gas-phase measurements from NO₃ initiated oxidation were obtained from those experiments



(Fig. 7d), whereas experiment No. 5 was conducted with UV lights off during the entire period, making it the only with particle-phase HOM measurements of the α -pinene and NO_3 reactions (Fig. 7f).

For both NO and NO_3 experiments, HOM_{ON} were dominated by compounds with only one nitrogen atom ($94.5 \pm 0.9\%$) in both gas and particle phases. This result indicates that NO-termination of NO_3 -derived nitrooxy- RO_2 was negligible in both systems, and likewise that HOM dimer formation through cross-reactions of two nRO_2 were unlikely. Note that many previous studies used N_2O_5 as the NO_3 source (and often with concentrations higher than VOCs) to avoid O_3 initiated oxidation, thus $\text{C}_x\text{H}_y\text{N}_2\text{O}_z$ compounds (e.g. $\text{C}_{20}\text{H}_{32}\text{N}_2\text{O}_{9,10}$) were the largest signals in the dimer range, formed by self/cross reactions of two nRO_2 (Claffin and Ziemann, 2018; Takeuchi and Ng, 2019; Bell et al., 2021). Here, O_3 derived RO_2 still contributed largely to HOM in this work, thus cross reactions of nRO_2 with RO_2 was the main process forming dimers (with only one N-atom, e.g. $\text{C}_{19}\text{H}_{31}\text{NO}_z$ and $\text{C}_{20}\text{H}_{31}\text{NO}_z$). In addition, since HOM_{ON} signals generally increased after turning on the UV lights (Fig. S6), peroxy acyl nitrates (PANs) formed through NO_2 -termination of RO_2 (Rissanen, 2018) were likely less important than NO-termination due to lower reaction rates and lower thermal stability of the products (Atkinson, 2000; Orlando and Tyndall, 2012; Berndt et al., 2015).

In the NO_3 experiment (experiment No.5, Fig. 7), C_{10} compounds dominated HOM_{CHO} in the monomer range in both gas and particle phases, while C_9 and C_{10} compounds were the largest HOM_{ON} monomers. Similarly, $\text{C}_{10}\text{H}_{15,17}\text{NO}_z$ and $\text{C}_9\text{H}_{13,15}\text{NO}_z$ were identified as the main HOM_{ON} products during NO_3 initiated α -/ β -pinene oxidation using a FIGAERO-iodide-CIMS for gas- and particle-phase measurements (Boyd et al., 2015; Nah et al., 2016; Takeuchi and Ng, 2019) and when using an EESI-TOF for SOA measurements at the Hyytiälä boreal forest site (Pospisilova et al., 2020). A detailed comparison of HOM_{ON} peaks identified in the particle phase from this work with other techniques is summarized in Table S3. Again, C_{16} - C_{18} compounds were nearly exclusively measured in the particle phase along with a relative enhancement of C_{19} compounds compared to the gas phase, indicating potential particle-phase related processes.

3.1.5 Summary of dimer/monomer distributions

In this section, the dimer to monomer (D/M) ratio was calculated for each experiment to relate the distribution of HOM products in both phases. Overall, D/M ratios observed in the particle phase followed the trend of gas phase for all experimental conditions. As dimers condensed more readily than monomers, the gas phase D/M ratios decreased after adding seed particles (Fig. S7d) and the condensed gas phase D/M ratios (grey markers in Fig. 8b) was roughly 1.5-fold compared to the gas phase. However, the D/M ratios measured in the particle phase were 2-9 times as high as those observed in the (condensed) gas phase (Fig. S8). Interestingly, the experiments where we observed the lowest gas-phase D/M ratios (experiment No. 4, 8, 9) were the ones where the particle-phase D/M enhancement was the largest. In case there was dimerization taking place in the particles, this process should indeed be seen most clearly when the amount of condensing dimers was the lowest. Overall, these highest discrepancies between different phases are larger than the expected underestimation of monomer signals (by a factor of ~ 2 estimated by comparing the measurements between 230°C and 170°C , as shown in Fig. 2c for C_{10} compounds). Nevertheless, further investigations will be needed to study to what extent could this enhancement be attributed to particle-phase chemistry.

One striking feature of Fig. 8b is that despite the D/M ratio of gas-phase HOM varying by a factor of 3 between different experiments (i.e., range of y-values for red markers in Fig. 8b), the relative changes of D/M ratios between different experiments (for the same phase) were clearly lower. For example, between experiments No. 1 and No. 3, the gas phase D/M ratio increased by $\sim 50\%$ (from 0.2 to 0.3) and the particle phase D/M ratio increased by a very similar amount, $\sim 54\%$ (from 1.1 to 1.7). We can only speculate about the underlying reasons behind this behavior. One possibility is that there is dimer formation in the particles that scales suitably to produce the observed results. Another possibility is that our sensitivity



to HOM monomers in the particle phase is lower than for HOM monomers in the gas phase. As indicated earlier, the average C-atom content in the HOM dimers was lower in the particles than in the gas phase under all conditions. The loss of C atoms is unlikely to be limited to the dimers, and similar changes in the monomers can be expected (whether through particle-phase chemistry or thermal decomposition). The loss of a few functional groups from a HOM monomer can potentially make it undetectable by the NO_3 -CIMS, which in turn would cause an increase in the observed D/M ratio. Future studies will hopefully be able to shed more lights on this matter. The potential role of thermal decomposition upon heating cannot be ruled out and will be discussed further in section 3.4.

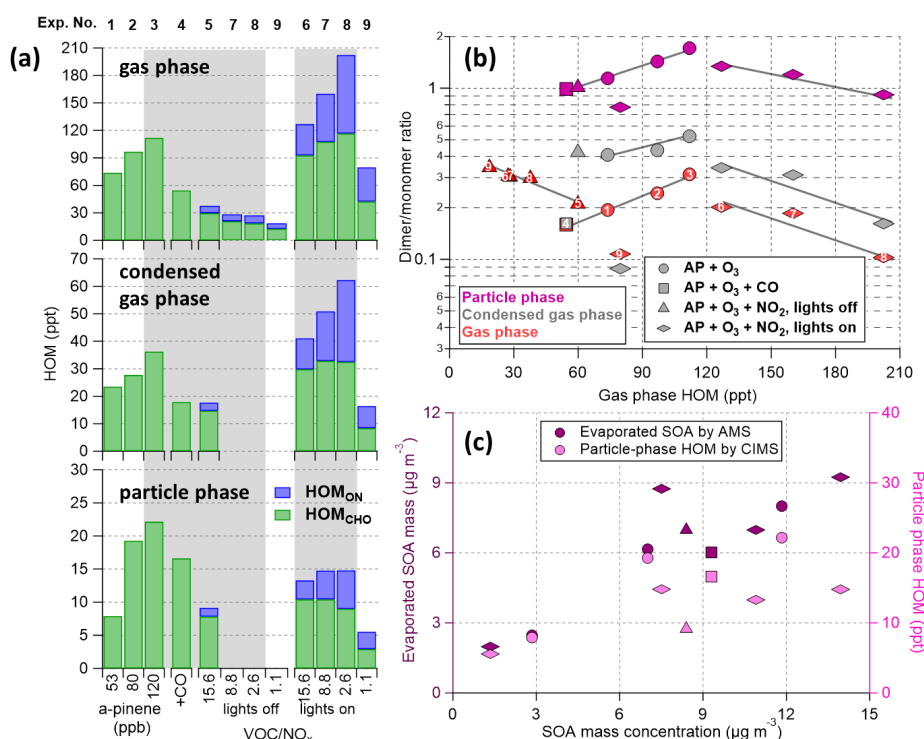


Figure 8. Summary of (a) total HOM concentrations estimated in the gas, condensed gas, and particle phase (different scales were used for different phases), (b) dimer to monomer (D/M, i.e. $\text{C}_{16}\text{-C}_{20}/\text{C}_8\text{-C}_{10}$) ratios, and (c) particle phase HOM and SOA measurements. The evaporated SOA mass was calculated from AMS measurements (i.e. the mass difference between VIA and bypass modes). Experiment numbers are given in panel (a) and (b). Shaded areas in panel (a) represents experiments with initial α -pinene concentrations of 120 ppb. The HOM concentrations are given in absolute values here to be able to compare the different experiments more quantitatively, but the large uncertainties (at least a factor 3) need to be kept in mind. The gas-phase HOM concentration is used as the x-axis in (b) as a surrogate for the overall α -pinene reaction rate, which could not be plotted directly due to the lack of VOC measurement. Grey solid lines in panels (b) only serve as guides to connect experiments with similar conditions.

In the NO experiments (No. 6-9 with UV lights on), D/M ratios of the condensed gas phase HOM changed the most, nearly by a factor 4 (Fig. 8b). However, the concurrent decrease in particle phase D/M ratios was smaller, closer to a factor 2. This change is also clear from Fig. S8. If there was particle phase dimer formation taking place, the effect should be most visible when the gas phase dimer contributions are the lowest, as in these experiments. While the very high NO_x/VOC ratios in these experiments make the direct comparison to other experiments hard, this can be seen as an indication that particle phase dimer formation is indeed taking place, as has been suggested earlier (Kalberer et al., 2004; Clafin et al., 2018). It should also be noted that in these experiments, the gas and particle phase absolute concentrations changed clearly from other



experiments (Fig. 8a). HOM concentrations ($= \text{HOM}_{\text{CHO}} + \text{HOM}_{\text{ON}}$) increased 13-81% in the gas phase, whereas particle-phase HOM decreased 33-40% (and SOA decreased 11-50%, Fig. S7) compared to the reference α -pinene and O_3 experiments (No. 1 and 3). This indicates a higher average volatility of the observed gas phase HOM, largely explainable by suppressed formation of low volatile HOM_{CHO} dimers and increased formation of more volatile HOM_{ON} monomers.

Exactly how this may impact the D/M ratios in gas and particle phase HOM remains an open question.

As there are still uncertainties relating to the VIA response, e.g. in correctly accounting for wall losses of vapors in the current experimental setup (especially within the VIA inlet), only a qualitative comparison was given here. As shown in Fig. 8c, ~80% of SOA evaporated within the VIA system based on AMS measurements, which is consistent with the integrated mass based on SMPS datasets (Fig. S7b). However, particle-phase HOM concentrations measured by the VIA- NO_3 -CIMS was much lower (by a factor of ~5-6, accounting for the dilution and assuming an average molecular mass of HOM compounds as 300 g mol^{-1}) compared to the AMS measurements. This difference is large enough that even after accounting for our large uncertainties in quantification, we expect that a noticeable fraction of the particle-phase products are not detected using this system. Vapor losses to the walls after evaporation as well as the selectivity of our ionization are expected to be the two main reasons for this. More detailed characterization of this inlet system will be investigated in future studies in order to evaluate if the wall losses can either be avoided or accounted for. Nevertheless, the trend of particle-phase HOM is comparable to the SOA measurements (Fig. 8c), with the largest discrepancies observed under high SOA loadings with NO_x added.

3.2 Elemental composition of HOM

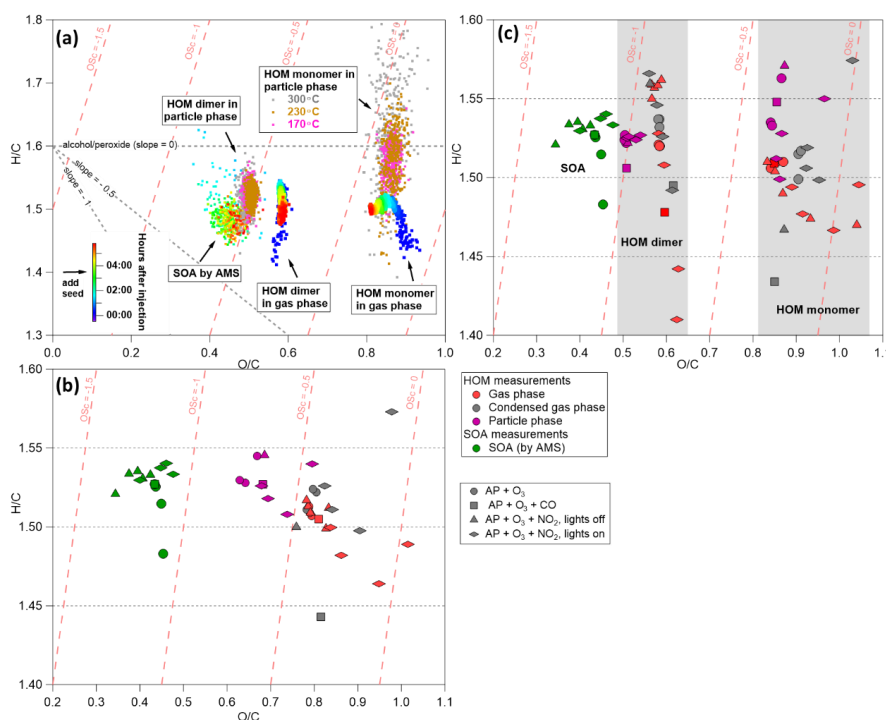


Figure 9. Van Krevelen diagrams of the VIA- NO_3 -CIMS and AMS data. (a) Gas-phase HOM (10-s averaged data points) and SOA measured by AMS (20-s averaged data points) from experiment No. 1. The SOA and the gas-phase HOM data are colored



by the time after α -pinene and O_3 injection, while the particle-phase HOM are colored by the temperature of the VIA. (b) Average “bulk” HOM and SOA elemental ratios for all experiments. (c) Average gas phase (red, 30-min), condensed gas phase (grey, 30-min), and particle phase (purple, 60-min) HOM elemental ratio separated into monomers (C_8 - C_{10}) and dimers (C_{16} - C_{20}).

535 In this section, Van Krevelen diagrams are used to investigate the bulk chemical properties of HOM measured in both gas and particle phases, and compared to the more routine elemental analysis from the AMS (Decarlo et al., 2006; Aiken et al., 2007; Aiken et al., 2008; Canagaratna et al., 2015).

As shown in Fig. 9a, gas-phase HOM monomers formed rapidly after the injection of α -pinene and O_3 , with bulk chemical composition stabilizing at elemental ratios resembling $C_{10}H_{15}O_8$. Dimers followed the H/C of monomers but showed lower
 540 O/C ratios, as has been observed also in earlier studies (Jokinen et al., 2014; Berndt et al., 2018). In the particle phase, the chemical composition of HOM measured at different temperatures were quite comparable, thus the datasets obtained at 230 °C were used to represent the particle-phase HOM measurements. Since the elemental ratios of HOM measurements are relatively stable 1.5 hours after injection of gas precursors, signal-weighted means were calculated for each phase after reaching a steady state for each experiment separately (Fig. 9b, c). In general, it is interesting that all data sets showed
 545 comparable H/C ratios (1.5 ± 0.1) but a relatively large range of O/C ratios. Especially, O/C ratios of HOM measured in the particle phase (purple markers in Fig. 9b) are lower compared to the gas (red markers) and condensed gas (grey markers) phase measurements (0.63-0.80 vs. 0.76-1.02). The AMS measured SOA showed the lowest O/C ratios (0.34-0.48), which are comparable with previous studies (Alfarra et al., 2006; Chhabra et al., 2010), summarized in Table S4. One possible explanation for the inconsistency of O/C ratios between AMS and CIMS measurements is that the VIA- NO_3 -CIMS measured
 550 only the most oxidized part of the SOA (as discussed in section 3.1.5).

HOM measurements were also separated into monomers and dimers for more detailed comparisons (Fig. 9c). In α -pinene ozonolysis experiments, the chemical compositions of gas-phase HOM are comparable under different initial α -pinene concentrations (red circles in Fig. 9c). The condensed gas phase monomers showed slightly higher O/C ratios, while dimers showed comparable O/C ratios compared to the gas phase (red vs. grey circles), which is consistent with the volatility
 555 distribution of HOM products (Peräkylä et al., 2020). However, a small O/C gap was observed between the particle phase and the (condensed) gas phase for dimers (0.50-0.51 vs. 0.58-0.59). In experiment No.4, adding CO largely shut down OH derived RO_2 , which lead to lower H/C ratios of gas-phase dimers compared to that before CO addition (Fig. S8a). On average, the H/C ratios of dimers decreased from 1.52 to 1.48 in the gas phase and decreased from 1.55 to 1.51 in the particle phase. The AMS measurements did not show detectable effects after adding CO. Again, the gas and particle phases showed a small
 560 gap in the dimer range (red vs. purple square in Fig. 9c) as in experiment No. 1-3.

In NO_x experiments, H/C ratios of both gas-phase HOM monomer and dimer decreased after turning on the UV lights (Fig. S8b), owing to the interruption of NO_3 initiated oxidation (i.e. addition of NO_3 formed nRO_2 radicals have one more H atom than O_3 derived RO_2). However, all particle-phase measurements (purple diamonds and triangles in Fig. 9c) were relatively comparable with other experiments (especially for dimers), indicating similar particle-phase transformations after the
 565 condensation of HOM from the gas phase. On the other hand, in both NO_3 (lights off) and NO (lights on) experiments, higher O/C ratios were observed in the monomer range compared to experiments without NO_x because of three more O atoms from the nitrate functional group. In addition, higher N/C ratios were observed in both the gas (0.037-0.07 vs. 0.018-0.02) and particle phase (0.01-0.04 vs. 0.01) with UV lights on compared to those without UV lights, confirming that more HOM_{ON} was formed through NO and RO_2 reactions than NO_3 oxidations. Correspondingly, higher O/C ratios of SOA were observed
 570 with UV lights on compared to the dark NO_3 experiment (0.40-0.48 vs. 0.34-0.41).



3.3 Organic nitrate (ON) measurements

Finally, organic nitrate measured by the AMS and the VIA-NO₃-CIMS system were compared to evaluate the performance of this novel system. Based on AMS measurements, nitrate mass tracked well with organics (Fig. 10a, b) for each experiment separately, indicating nitrate signals were from organic nitrates. The nitrate/organic ratio in the AMS also increase very clearly as the NO_x/VOC ratio in the chamber increased. The nitrate/organic mass ratios increased from 0.012-0.061 to 0.018-0.075 after turning on the UV lights, in agreement with earlier discussion that more intensive HOM_{ON} formation took place in NO experiments than in NO₃ experiments. The NO⁺/NO₂⁺ ratios, a marker for the contribution of organic nitrates to the measured nitrate signal, observed in experiment No. 6-9 did not show discrepancies between UV lights on and off (Fig. 10c), and were well separated from the data set of ammonium nitrate calibration (5.2-5.6 vs. 1.6).

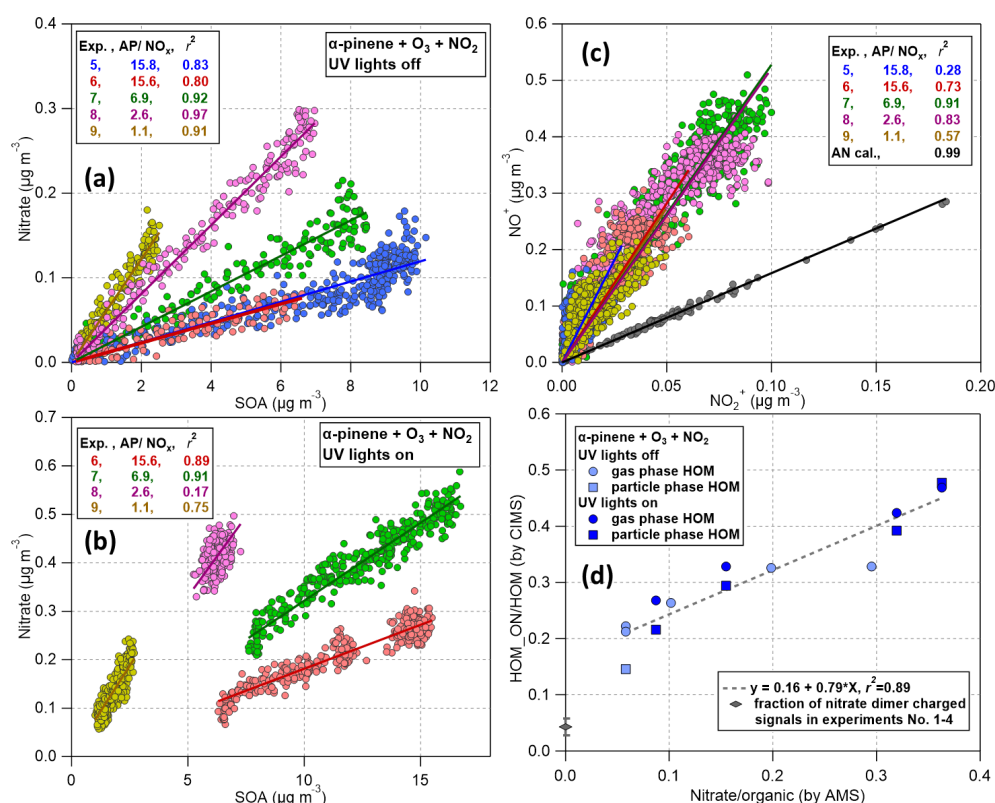


Figure 10. Scatter plots of organic nitrate observations with the VIA-NO₃-CIMS and AMS. (a) AMS-measured nitrate mass vs. organic mass during NO_x experiments (No. 5-9) with UV lights off. (b) Same as (a), but with UV lights on. (c) NO⁺ vs. NO₂⁺ ions measured by the AMS, including experiments with both UV lights off and on. (d) Molar fraction of organic nitrate (HOM_{ON}) to total organic signals (HOM = HOM_{CHO} + HOM_{ON}) measured by the VIA-NO₃-CIMS system vs. the ratio of nitrate to organics measured by the AMS ((nitrate/62)/(SOA/300)).

Assuming an average molar mass as 300 g mol⁻¹ for organic molecules, the molar fraction of ON to total organic aerosols (= (nitrate/62)/(SOA/300)) was estimated to be 5.8-30% for NO₃ experiments and 8.7-36% for NO experiments based on AMS measurements. Interestingly, the molar fraction of HOM_{ON} to total HOM signals measured in both the gas (circles in Fig. 10d) and particle phase (squares in Fig. 10d) by the VIA-NO₃-CIMS correlated well with the above AMS measurements, suggesting that chemical composition of gas- and particle-phase HOM (i.e. relative contribution of HOM_{CHO} and HOM_{ON})



were quite comparable. Note that the slope of the fitted line is sensitive to several factors, including different calibration factors between HOM_{CHO} and HOM_{ON} and the estimated average molar mass of SOA. Thus, the value should be only qualitatively interpreted. We also note that some part of the non-zero intercept in Figure 10d (16%) is likely contributed by charging in the NO_3 -CIMS by the nitrate dimer ($\text{HNO}_3 \cdot \text{NO}_3^-$), which is higher than that in α -pinene ozonolysis experiments (estimated to be $4.3 \pm 1.5\%$). It is possible that this ratio would increase further in NO_x experiments if there was HNO_3 formed at some amounts in the gas phase. Finally, neither the $\text{NO}^+/\text{NO}_2^+$ nor the ON molar fraction analysis showed any clear differences between the NO and NO_3 experiments based on both AMS and NO_3 -CIMS datasets. Thus, other parameters (e.g. UV lights on/off in the chamber or hour of the day in the atmosphere) is needed to distinguish HOM_{ON} formed by NO_3 oxidation from those by NO termination of RO_2 .

3.4 Discussion on the factors that may affect the HOM detection

Thermal decomposition is a well-known issue with techniques based on the evaporation of molecules from the particle phase to identify the chemical composition of vapors, especially for labile organic mixtures. It remains unclear to what extent the thermal decomposition of SOA happened under different temperatures and/or if the thermal desorption techniques could promote particle-phase reactions. However, if thermal decomposition did play an important role in this VIA- NO_3 -CIMS system, we would expect more decomposition of dimers (labile RO-OR compounds) than monomers, and in particular at the highest temperature. Yet, we observed much higher dimer/monomer ratios in the particle phase than the gas phase, suggesting that we might instead “enhance” dimers if the thermal decomposition would primarily remove monomers from our sensitivity window.

In section 3.2, lower O/C ratios of the particle-phase HOM were observed comparing to the gas-phase HOM, which could partly be attributed to the thermal decomposition of organic compounds within the VIA. The carbonyl, peroxide, esters, carboxyl, and hydroxyl were ubiquitous functional groups identified in α -pinene SOA under various conditions (Claflin et al., 2018). Decarboxylation of $\text{C}(=\text{O})\text{O}-$ functional groups may play a relatively more important role than the dehydration of hydroxyl (with H_2O being the main thermal decomposition product) owing to their expected higher contributions to SOA. Concurrent formation of CO, H_2O , and CO_2 from decarboxylation would decrease both the O/C and H/C ratios for the particle-phase measurements, which was observed at 200 °C on the vaporizer of the AMS (Canagaratna et al., 2015) and in the FIGAERO-CIMS (Stark et al., 2017). With proper combinations of $\text{C}(=\text{O})\text{OH}$ and $\text{C}(=\text{O})\text{OR}$ thermal decomposition, it is possible to decrease O/C as well as slightly increase H/C (e.g. on average losing more CO_2/CO than H_2O).

In addition, thermal decomposition of peroxide oligomers was observed in the thermogram of an acetate-FIGAERO measurements, with an estimated dissociation temperature around 80 °C for O-O bond (Lopez-Hilfiker et al., 2015).

However, large amounts of C_{16} - C_{19} dimers measured in this work suggested either these labile peroxides transferred to other more stable functional groups in SOA, or the much short residence time at elevated temperature in the VIA (less than 0.2 s), as well as sampling without surface contact, may lead to less thermal decomposition compared to the traditional collection-and-desorption methods (Häkkinen et al., 2022). Nevertheless, it is clear that we need further investigations to determine if particle-phase processing or thermal decomposition (or other hardware effects) are the larger causes for the changes observed between the particle and (condensed) gas phase in our work. In particular, the temperature-dependent sensitivity (including evaporation and transmission efficiencies) for different types of organic molecules will need to be characterized in future works.



4 Conclusion

Combined gas- and particle-phase HOM measurements were conducted using the novel VIA inlet combined with a NO₃-CIMS in this study. α -pinene ozonolysis was studied under different conditions, including using CO as an OH scavenger and varying NO_x concentrations with and without UV lights to perturb the oxidation chemistry and gas-phase dimer formation. Within the 50-min residence time of the chamber, we observed distinct differences between HOM in particles and the HOM that were condensing. In particular, we noted clearly enhanced C₁₆-C₁₉ HOM dimers in SOA that were not observed in the gas phase. This “enhancement” might be partly attributed to a slightly higher sensitivity towards dimers than monomers using the VIA inlet system. However, different chemical composition identified between two phases implies the existence of potential particle-phase chemistry. In addition, gas-phase dimer formation was considerably suppressed in both the CO and NO experiments, and this also led to decreases in particle-phase dimers, yet dimers still made up a considerable fraction of the observed SOA.

In addition to the generally shorter carbon skeletons of the particle phase dimers (mainly C₁₆-C₁₉) compared to the gas phase (mainly C₁₉-C₂₀), average O/C ratios of the HOM (especially in the dimer range) also decreased slightly in the particle phase. Earlier studies have proposed that e.g. Baeyer-Villiger reactions could convert peroxydicarboxylic acids and hydroperoxides to ester dimers in reactions with aldehydes/ketones (Claflin et al., 2018; Bakker-Arkema and Ziemann, 2020), which is not inconsistent with the results observed in this study. Our results indicate that particle-phase reactions are taking place under the time scales probed here, with some of the reactions potentially leading to the formation of new dimers, while some reactions may be causing loss of small fragments (e.g. C₂₀ dimers forming C₁₇ dimers). These predominant C₁₇H₂₆O_z compounds identified in the particle phase, which have often been reported by previous offline measurements, might be related to HOM dimer formation in both the gas and particle phases. However, the new VIA-NO₃-CIMS system used in this work will require more detailed characterization in order to better understand how the thermal desorption and wall effects may modify the HOM distributions from those existing in the particles. The NO₃-CIMS is known to be selective towards the most oxygenated molecules, which also means that we are likely not detecting all the evaporating compounds. This is also indicated by the measured O/C ratios being clearly higher than those measured by an Aerodyne AMS. However, the measured organic nitrate fraction measured by the VIA-NO₃-CIMS was found to track those measured by the AMS. Taken together, we believe that this system is a promising technique for combined online gas- and particle-phase HOM measurements.

Data availability. Data are available upon request from the corresponding authors.

Author contributions. ME and JZ designed the study. ME, JK, JEK, MC, and DR support the experiment setup and analysis methods. JEK built the first version of the VIA. JZ, EH, and FG conducted the experiments. JZ analyzed the data. JZ prepared the manuscript with contributions from all co-authors.

Competing interests. Jordan Krechmer, Manjula Canagaratna, and Douglas Worsnop work for Aerodyne Research, Inc., which developed and commercialized the VIA inlet used in this study.

Acknowledgements. This work was supported by funding from Academy of Finland (grants 317380, 320094, 325656, 345982, and 346370) and University of Helsinki 3-year grant (75284132). Jian Zhao thanks Lauri Franzon and Otso Peräkylä for helpful discussions about the experiments and results. Ella Häkkinen thanks the Vilho, Yrjö and Kalle Väisälä Foundation for financial support. Frans Graeffe acknowledge Svenska Kulturfonden (grants 167344 and 177923) for financial support.



References

- Aiken, A. C., DeCarlo, P. F., and Jimenez, J. L.: Elemental analysis of organic species with electron ionization high-resolution mass spectrometry, *Anal. chem.*, 79, 8350-8358, doi:10.1021/ac071150w, 2007.
- 670 Aiken, A. C., Decarlo, P. F., Kroll, J. H., Worsnop, D. R., Huffman, J. A., Docherty, K. S., Ulbrich, I. M., Mohr, C., Kimmel, J. R., and Sueper, D.: O/C and OM/OC ratios of primary, secondary, and ambient organic aerosols with high-resolution time-of-flight aerosol mass spectrometry, *Environ. Sci. Technol.*, 42, 4478-4485, <https://doi.org/10.1021/es703009q>, 2008.
- Alfarra, M. R., Paulsen, D., Gysel, M., Garforth, A. A., Dommen, J., Prévôt, A. S. H., Worsnop, D. R., Baltensperger, U., and Coe, H.: A mass spectrometric study of secondary organic aerosols formed from the photooxidation of anthropogenic and biogenic precursors in a reaction chamber, *Atmos. Chem. Phys.*, 6, 5279-5293, doi:10.5194/acp-6-5279-2006, 2006.
- 675 Atkinson, R.: Atmospheric chemistry of VOCs and NO_x, *Atmos. Environ.*, 34, 2063-2101, doi:10.1016/s1352-2310(99)00460-4, 2000.
- Bakker-Arkema, J. G. and Ziemann, P. J.: Measurements of Kinetics and Equilibria for the Condensed Phase Reactions of Hydroperoxides with Carbonyls to Form Peroxyhemiacetals, *ACS Earth and Space Chemistry*, 4, 467-475, doi:10.1021/acsearthspacechem.0c00008, 2020.
- 680 Beck, M. and Hoffmann, T.: A detailed MS_n study for the molecular identification of a dimer formed from oxidation of pinene, *Atmos. Environ.*, 130, 120-126, <https://doi.org/10.1016/j.atmosenv.2015.09.012>, 2016.
- Bell, D. M., Wu, C., Bertrand, A., Graham, E., Schoonbaert, J., Giannoukos, S., Baltensperger, U., Prevot, A. S. H., Riipinen, I., El Haddad, I., and Mohr, C.: Particle-phase processing of α -pinene NO₃ secondary organic aerosol in the dark, *Atmos. Chem. Phys.*, 22, 13167-13182, doi:10.5194/acp-22-13167-2022, 2022.
- 685 Berndt, T., Mentler, B., Scholz, W., Fischer, L., Herrmann, H., Kulmala, M., and Hansel, A.: Accretion Product Formation from Ozonolysis and OH Radical Reaction of α -Pinene: Mechanistic Insight and the Influence of Isoprene and Ethylene, *Environ. Sci. Technol.*, 52, 11069-11077, doi:10.1021/acs.est.8b02210, 2018.
- Berndt, T., Richters, S., Kaethner, R., Voigtländer, J., Stratmann, F., Sipilä, M., Kulmala, M., and Herrmann, H.: Gas-Phase Ozonolysis of Cycloalkenes: Formation of Highly Oxidized RO₂ Radicals and Their Reactions with NO, NO₂, SO₂, and Other RO₂ Radicals, *The Journal of Physical Chemistry A*, 119, 10336-10348, doi:10.1021/acs.jpca.5b07295, 2015.
- Berndt, T., Richters, S., Jokinen, T., Hyttinen, N., Kurten, T., Otkjaer, R. V., Kjaergaard, H. G., Stratmann, F., Herrmann, H., Sipilä, M., Kulmala, M., and Ehn, M.: Hydroxyl radical-induced formation of highly oxidized organic compounds, *Nature communications*, 7, 13677, doi:10.1038/ncomms13677, 2016.
- 695 Bianchi, F., Kurtén, T., Riva, M., Mohr, C., Rissanen, M. P., Roldin, P., Berndt, T., Crounse, J. D., Wennberg, P. O., Mentel, T. F., Wildt, J., Junninen, H., Jokinen, T., Kulmala, M., Worsnop, D. R., Thornton, J. A., Donahue, N., Kjaergaard, H. G., and Ehn, M.: Highly Oxygenated Organic Molecules (HOM) from Gas-Phase Autoxidation Involving Peroxy Radicals: A Key Contributor to Atmospheric Aerosol, *Chem. Rev.*, 119, 3472-3509, doi:10.1021/acs.chemrev.8b00395, 2019.
- Boucher, O., Randall, D., Artaxo, P., Bretherton, C., Feingold, G., Forster, P., Kerminen, V.-M., Kondo, Y., Liao, H., and Lohmann, U.: Clouds and aerosols, in: *Climate change 2013: the physical science basis. Contribution of Working Group I to the Fifth Assessment Report of the Intergovernmental Panel on Climate Change*, Cambridge University Press, 571-657, 2013.
- 700 Boyd, C. M., Sanchez, J., Xu, L., Eugene, A. J., Nah, T., Tuet, W. Y., Guzman, M. I., and Ng, N. L.: Secondary organic aerosol formation from the beta-pinene+NO₃ system: effect of humidity and peroxy radical fate, *Atmos. Chem. Phys.*, 15, 7497-7522, doi:10.5194/acp-15-7497-2015, 2015.
- 705



- Canagaratna, M., Jayne, J., Jimenez, J., Allan, J., Alfarra, M., Zhang, Q., Onasch, T., Drewnick, F., Coe, H., and Middlebrook, A.: Chemical and microphysical characterization of ambient aerosols with the aerodyne aerosol mass spectrometer, *Mass Spectrom. Rev.*, 26, 185-222, <https://doi.org/10.1002/mas.20115>, 2007.
- Canagaratna, M. R., Jimenez, J. L., Kroll, J. H., Chen, Q., Kessler, S. H., Massoli, P., Hildebrandt Ruiz, L., Fortner, E., Williams, L. R., Wilson, K. R., Surratt, J. D., Donahue, N. M., Jayne, J. T., and Worsnop, D. R.: Elemental ratio measurements of organic compounds using aerosol mass spectrometry: characterization, improved calibration, and implications, *Atmos. Chem. Phys.*, 15, 253-272, [doi:10.5194/acp-15-253-2015](https://doi.org/10.5194/acp-15-253-2015), 2015.
- Caudillo, L., Rörup, B., Heinritzi, M., Marie, G., Simon, M., Wagner, A. C., Müller, T., Granzin, M., Amorim, A., Ataei, F., Baalbaki, R., Bertozzi, B., Brasseur, Z., Chiu, R., Chu, B., Dada, L., Duplissy, J., Finkenzeller, H., Gonzalez Carracedo, L., He, X. C., Hofbauer, V., Kong, W., Lamkaddam, H., Lee, C. P., Lopez, B., Mahfouz, N. G. A., Makhmutov, V., Manninen, H. E., Marten, R., Massabò, D., Mauldin, R. L., Mentler, B., Molteni, U., Onnela, A., Pfeifer, J., Philippov, M., Piedehierro, A. A., Schervish, M., Scholz, W., Schulze, B., Shen, J., Stolzenburg, D., Stozhkov, Y., Surdu, M., Tauber, C., Tham, Y. J., Tian, P., Tomé, A., Vogt, S., Wang, M., Wang, D. S., Weber, S. K., Welti, A., Yonghong, W., Yusheng, W., Zauner-Wieczorek, M., Baltensperger, U., El Haddad, I., Flagan, R. C., Hansel, A., Höhler, K., Kirkby, J., Kulmala, M., Lehtipalo, K., Möhler, O., Saathoff, H., Volkamer, R., Winkler, P. M., Donahue, N. M., Kürten, A., and Curtius, J.: Chemical composition of nanoparticles from α -pinene nucleation and the influence of isoprene and relative humidity at low temperature, *Atmos. Chem. Phys.*, 21, 17099-17114, [doi:10.5194/acp-21-17099-2021](https://doi.org/10.5194/acp-21-17099-2021), 2021.
- Chhabra, P. S., Flagan, R. C., and Seinfeld, J. H.: Elemental analysis of chamber organic aerosol using an aerodyne high-resolution aerosol mass spectrometer, *Atmos. Chem. Phys.*, 10, 4111-4131, [doi:10.5194/acp-10-4111-2010](https://doi.org/10.5194/acp-10-4111-2010), 2010.
- Claeys, M., Iinuma, Y., Szmigielski, R., Surratt, J. D., Blockhuys, F., Van Alsenoy, C., Böge, O., Sierau, B., Gómez-González, Y., Vermeylen, R., Van der Veken, P., Shahgholi, M., Chan, A. W. H., Herrmann, H., Seinfeld, J. H., and Maenhaut, W.: Terpenylic Acid and Related Compounds from the Oxidation of α -Pinene: Implications for New Particle Formation and Growth above Forests, *Environ. Sci. Technol.*, 43, 6976-6982, [doi:10.1021/es9007596](https://doi.org/10.1021/es9007596), 2009.
- Clafflin, M. S. and Ziemann, P. J.: Identification and Quantitation of Aerosol Products of the Reaction of β -Pinene with NO₃ Radicals and Implications for Gas- and Particle-Phase Reaction Mechanisms, *The Journal of Physical Chemistry A*, 122, 3640-3652, [doi:10.1021/acs.jpca.8b00692](https://doi.org/10.1021/acs.jpca.8b00692), 2018.
- Clafflin, M. S., Krechmer, J. E., Hu, W., Jimenez, J. L., and Ziemann, P. J.: Functional Group Composition of Secondary Organic Aerosol Formed from Ozonolysis of α -Pinene Under High VOC and Autoxidation Conditions, *ACS Earth and Space Chemistry*, 2, 1196-1210, [doi:10.1021/acsearthspacechem.8b00117](https://doi.org/10.1021/acsearthspacechem.8b00117), 2018.
- Crounse, J. D., Nielsen, L. B., Jørgensen, S., Kjaergaard, H. G., and Wennberg, P. O.: Autoxidation of Organic Compounds in the Atmosphere, *The Journal of Physical Chemistry Letters*, 4, 3513-3520, [10.1021/jz4019207](https://doi.org/10.1021/jz4019207), 2013.
- Day, D. A., Campuzano-Jost, P., Nault, B. A., Palm, B. B., Hu, W., Guo, H., Wooldridge, P. J., Cohen, R. C., Docherty, K. S., Huffman, J. A., de Sá, S. S., Martin, S. T., and Jimenez, J. L.: A systematic re-evaluation of methods for quantification of bulk particle-phase organic nitrates using real-time aerosol mass spectrometry, *Atmos. Meas. Tech.*, 15, 459-483, [doi:10.5194/amt-15-459-2022](https://doi.org/10.5194/amt-15-459-2022), 2022.
- DeCarlo, P. F., Kimmel, J. R., Trimborn, A., Northway, M. J., Jayne, J. T., Aiken, A. C., Gonin, M., Fuhrer, K., Horvath, T., and Docherty, K. S.: Field-deployable, high-resolution, time-of-flight aerosol mass spectrometer, *Anal. chem.*, 78, 8281-8289, <https://doi.org/10.1021/ac061249n>, 2006.
- Ding, X., He, Q.-F., Shen, R.-Q., Yu, Q.-Q., and Wang, X.-M.: Spatial distributions of secondary organic aerosols from isoprene, monoterpenes, β -caryophyllene, and aromatics over China during summer, *J. Geophys. Res.- Atmos.*, 119, 877-811, [doi:10.1002/2014JD021748](https://doi.org/10.1002/2014JD021748), 2014.



- Docherty, K. S., Wu, W., Lim, Y. B., and Ziemann, P. J.: Contributions of organic peroxides to secondary aerosol formed from reactions of monoterpenes with O₃, *Environ. Sci. Technol.*, 39, 4049-4059, doi:10.1021/es050228s, 2005.
- Donahue, N. M., Kroll, J. H., Pandis, S. N., and Robinson, A. L.: A two-dimensional volatility basis set – Part 2: Diagnostics
 750 of organic-aerosol evolution, *Atmos. Chem. Phys.*, 12, 615-634, doi:10.5194/acp-12-615-2012, 2012.
- Donahue, N. M., Ortega, I. K., Chuang, W., Riipinen, I., Riccobono, F., Schobesberger, S., Dommen, J., Baltensperger, U., Kulmala, M., Worsnop, D. R., and Vehkamäki, H.: How do organic vapors contribute to new-particle formation?, *Faraday Discuss.*, 165, 91, doi:10.1039/c3fd00046j, 2013.
- Ehn, M., Berndt, T., Wildt, J., and Mentel, T.: Highly Oxygenated Molecules from Atmospheric Autoxidation of Hydrocarbons:
 755 A Prominent Challenge for Chemical Kinetics Studies, *International Journal of Chemical Kinetics*, 49, 821-831, <https://doi.org/10.1002/kin.21130>, 2017.
- Ehn, M., Kleist, E., Junninen, H., Petäjä, T., Lönn, G., Schobesberger, S., Dal Maso, M., Trimborn, A., Kulmala, M., Worsnop, D. R., Wahner, A., Wildt, J., and Mentel, T. F.: Gas phase formation of extremely oxidized pinene reaction products in chamber and ambient air, *Atmos. Chem. Phys.*, 12, 5113-5127, doi:10.5194/acp-12-5113-2012, 2012.
- 760 Ehn, M., Thornton, J. A., Kleist, E., Sipila, M., Junninen, H., Pullinen, I., Springer, M., Rubach, F., Tillmann, R., Lee, B., Lopez-Hilfiker, F., Andres, S., Acir, I. H., Rissanen, M., Jokinen, T., Schobesberger, S., Kangasluoma, J., Kontkanen, J., Nieminen, T., Kurten, T., Nielsen, L. B., Jorgensen, S., Kjaergaard, H. G., Canagaratna, M., Maso, M. D., Berndt, T., Petaja, T., Wahner, A., Kerminen, V. M., Kulmala, M., Worsnop, D. R., Wildt, J., and Mentel, T. F.: A large source of low-volatility secondary organic aerosol, *Nature*, 506, 476-479, doi:10.1038/nature13032, 2014.
- 765 Eichler, P., Müller, M., D'Anna, B., and Wisthaler, A.: A novel inlet system for online chemical analysis of semi-volatile submicron particulate matter, *Atmos. Meas. Tech.*, 8, 1353-1360, 10.5194/amt-8-1353-2015, 2015.
- Farmer, D. K., Matsunaga, A., Docherty, K. S., Surratt, J. D., Seinfeld, J. H., Ziemann, P. J., and Jimenez, J. L.: Response of an aerosol mass spectrometer to organonitrates and organosulfates and implications for atmospheric chemistry, *Proc. Natl. Acad. Sci. USA*, 107, 6670-6675, doi:10.1073/pnas.0912340107, 2010.
- 770 Griffin, R. J., Cocker III, D. R., Flagan, R. C., and Seinfeld, J. H.: Organic aerosol formation from the oxidation of biogenic hydrocarbons, *J. Geophys. Res.- Atmos.*, 104, 3555-3567, <https://doi.org/10.1029/1998JD100049>, 1999.
- Guenther, A. B., Jiang, X., Heald, C. L., Sakulyanontvittaya, T., Duhl, T., Emmons, L. K., and Wang, X.: The Model of Emissions of Gases and Aerosols from Nature version 2.1 (MEGAN2.1): an extended and updated framework for modeling biogenic emissions, *Geosci. Model Dev.*, 5, 1471-1492, doi:10.5194/gmd-5-1471-2012, 2012.
- 775 Hallquist, M., Wenger, J. C., Baltensperger, U., Rudich, Y., Simpson, D., Claeys, M., Dommen, J., Donahue, N. M., George, C., Goldstein, A. H., Hamilton, J. F., Herrmann, H., Hoffmann, T., Iinuma, Y., Jang, M., Jenkin, M. E., Jimenez, J. L., Kiendler-Scharr, A., Maenhaut, W., McFiggans, G., Mentel, T. F., Monod, A., Prévôt, A. S. H., Seinfeld, J. H., Surratt, J. D., Szmigielski, R., and Wildt, J.: The formation, properties and impact of secondary organic aerosol: current and emerging issues, *Atmos. Chem. Phys.*, 9, 5155-5236, doi:10.5194/acp-9-5155-2009, 2009.
- 780 Hoffmann, T., Odum, J. R., Bowman, F., Collins, D., Klockow, D., Flagan, R. C., and Seinfeld, J. H.: Formation of Organic Aerosols from the Oxidation of Biogenic Hydrocarbons, *Journal of Atmospheric Chemistry*, 26, 189-222, doi:10.1023/A:1005734301837, 1997.
- Hyttinen, N., Rissanen, M. P., and Kurtén, T.: Computational Comparison of Acetate and Nitrate Chemical Ionization of Highly Oxidized Cyclohexene Ozonolysis Intermediates and Products, *The Journal of Physical Chemistry A*, 121, 2172-2179,
 785 doi:10.1021/acs.jpca.6b12654, 2017.



- Hyttinen, N., Kupiainen-Määttä, O., Rissanen, M. P., Muuronen, M., Ehn, M., and Kurtén, T.: Modeling the Charging of Highly Oxidized Cyclohexene Ozonolysis Products Using Nitrate-Based Chemical Ionization, *The Journal of Physical Chemistry A*, 119, 6339-6345, doi:10.1021/acs.jpca.5b01818, 2015.
- Hyttinen, N., Wolf, M., Rissanen, M. P., Ehn, M., Peräkylä, O., Kurtén, T., and Prisle, N. L.: Gas-to-Particle Partitioning of
790 Cyclohexene- and α -Pinene-Derived Highly Oxygenated Dimers Evaluated Using COSMOtherm, *The Journal of Physical Chemistry A*, 125, 3726-3738, doi:10.1021/acs.jpca.0c11328, 2021.
- Häkkinen, E., Zhao, J., Graeffe, F., Fauré, N., Krechmer, J., Worsnop, D., Timonen, H., Ehn, M., and Kangasluoma, J.: Online measurement of highly oxygenated compounds from organic aerosol, *EGUsphere*, 2022, 1-29, doi:10.5194/egusphere-2022-933, 2022.
- 795 Iyer, S., Reiman, H., Möller, K. H., Rissanen, M. P., Kjaergaard, H. G., and Kurtén, T.: Computational Investigation of RO₂ + HO₂ and RO₂ + RO₂ Reactions of Monoterpene Derived First-Generation Peroxy Radicals Leading to Radical Recycling, *The Journal of Physical Chemistry A*, 122, 9542-9552, doi:10.1021/acs.jpca.8b09241, 2018.
- Iyer, S., Rissanen, M. P., Valiev, R., Barua, S., Krechmer, J. E., Thornton, J., Ehn, M., and Kurtén, T.: Molecular mechanism for rapid autoxidation in α -pinene ozonolysis, *Nature communications*, 12, 878, doi:10.1038/s41467-021-21172-w, 2021.
- 800 Jayne, J. T., Leard, D. C., Zhang, X., Davidovits, P., Smith, K. A., Kolb, C. E., and Worsnop, D. R.: Development of an Aerosol Mass Spectrometer for Size and Composition Analysis of Submicron Particles, *Aerosol Sci. Technol.*, 33, 49-70, doi:10.1080/027868200410840, 2000.
- Jimenez, J., Canagaratna, M., Donahue, N., Prevot, A., Zhang, Q., Kroll, J. H., DeCarlo, P. F., Allan, J. D., Coe, H., and Ng, N.: Evolution of organic aerosols in the atmosphere, *Science*, 326, 1525-1529, doi:10.1126/science.1180353, 2009.
- 805 Jokinen, T., Sipilä, M., Junninen, H., Ehn, M., Lönn, G., Hakala, J., Petäjä, T., Mauldin III, R. L., Kulmala, M., and Worsnop, D. R.: Atmospheric sulphuric acid and neutral cluster measurements using CI-API-TOF, *Atmos. Chem. Phys.*, 12, 4117-4125, doi:10.5194/acp-12-4117-2012, 2012.
- Jokinen, T., Sipilä, M., Richters, S., Kerminen, V. M., Paasonen, P., Stratmann, F., Worsnop, D., Kulmala, M., Ehn, M., Herrmann, H., and Berndt, T.: Rapid autoxidation forms highly oxidized RO₂ radicals in the atmosphere, *Angewandte
810 Chemie*, 53, 14596-14600, doi:10.1002/anie.201408566, 2014.
- Jokinen, T., Berndt, T., Makkonen, R., Kerminen, V. M., Junninen, H., Paasonen, P., Stratmann, F., Herrmann, H., Guenther, A. B., Worsnop, D. R., Kulmala, M., Ehn, M., and Sipilä, M.: Production of extremely low volatile organic compounds from biogenic emissions: Measured yields and atmospheric implications, *Proc. Natl. Acad. Sci. USA*, 112, 7123-7128, doi:10.1073/pnas.1423977112, 2015.
- 815 Junninen, H., Ehn, M., Petäjä, T., Luosujärvi, L., Kotiaho, T., Kostianen, R., Rohner, U., Gonin, M., Fuhrer, K., Kulmala, M., and Worsnop, D. R.: A high-resolution mass spectrometer to measure atmospheric ion composition, *Atmos. Meas. Tech.*, 3, 1039-1053, doi:10.5194/amt-3-1039-2010, 2010.
- Kahnt, A., Vermeylen, R., Iinuma, Y., Safi Shalamzari, M., Maenhaut, W., and Claeys, M.: High-molecular-weight esters in α -pinene ozonolysis secondary organic aerosol: structural characterization and mechanistic proposal for their formation from
820 highly oxygenated molecules, *Atmos. Chem. Phys.*, 18, 8453-8467, doi:10.5194/acp-18-8453-2018, 2018.
- Kalberer, M., Paulsen, D., Sax, M., Steinbacher, M., Dommen, J., Prevot, A. S. H., Fisseha, R., Weingartner, E., Frankevich, V., Zenobi, R., and Baltensperger, U.: Identification of Polymers as Major Components of Atmospheric Organic Aerosols, *Science*, 303, 1659-1662, doi:10.1126/science.1092185, 2004.
- Kang, E., Root, M. J., Toohey, D. W., and Brune, W. H.: Introducing the concept of Potential Aerosol Mass (PAM), *Atmos.
825 Chem. Phys.*, 7, 5727-5744, doi:10.5194/acp-7-5727-2007, 2007.



- Kerdouci, J., Picquet-Varrault, B., and Doussin, J.-F.: Prediction of Rate Constants for Gas-Phase Reactions of Nitrate Radical with Organic Compounds: A New Structure–Activity Relationship, *ChemPhysChem*, 11, 3909-3920, <https://doi.org/10.1002/cphc.201000673>, 2010.
- Kirkby, J., Duplissy, J., Sengupta, K., Frege, C., Gordon, H., Williamson, C., Heinritzi, M., Simon, M., Yan, C., Almeida, J.,
 830 Trostl, J., Nieminen, T., Ortega, I. K., Wagner, R., Adamov, A., Amorim, A., Bernhammer, A. K., Bianchi, F., Breitenlechner, M., Brilke, S., Chen, X., Craven, J., Dias, A., Ehrhart, S., Flagan, R. C., Franchin, A., Fuchs, C., Guida, R., Hakala, J., Hoyle, C. R., Jokinen, T., Junninen, H., Kangasluoma, J., Kim, J., Krapf, M., Kurten, A., Laaksonen, A., Lehtipalo, K., Makhmutov, V., Mathot, S., Molteni, U., Onnela, A., Perakyla, O., Piel, F., Petaja, T., Praplan, A. P., Pringle, K., Rap, A., Richards, N. A., Riipinen, I., Rissanen, M. P., Rondo, L., Sarnela, N., Schobesberger, S., Scott, C. E., Seinfeld,
 835 J. H., Sipila, M., Steiner, G., Stozhkov, Y., Stratmann, F., Tome, A., Virtanen, A., Vogel, A. L., Wagner, A. C., Wagner, P. E., Weingartner, E., Wimmer, D., Winkler, P. M., Ye, P., Zhang, X., Hansel, A., Dommen, J., Donahue, N. M., Worsnop, D. R., Baltensperger, U., Kulmala, M., Carslaw, K. S., and Curtius, J.: Ion-induced nucleation of pure biogenic particles, *Nature*, 533, 521-526, doi:10.1038/nature17953, 2016.
- Krapf, M., El Haddad, I., Bruns, Emily A., Molteni, U., Daellenbach, Kaspar R., Prévôt, André S. H., Baltensperger, U., and
 840 Dommen, J.: Labile Peroxides in Secondary Organic Aerosol, *Chem*, 1, 603-616, <https://doi.org/10.1016/j.chempr.2016.09.007>, 2016.
- Kristensen, K., Cui, T., Zhang, H., Gold, A., Glasius, M., and Surratt, J. D.: Dimers in α -pinene secondary organic aerosol: effect of hydroxyl radical, ozone, relative humidity and aerosol acidity, *Atmos. Chem. Phys.*, 14, 4201-4218, doi:10.5194/acp-14-4201-2014, 2014.
- 845 Kristensen, K., Watne, Å. K., Hammes, J., Lutz, A., Petäjä, T., Hallquist, M., Bilde, M., and Glasius, M.: High-Molecular Weight Dimer Esters Are Major Products in Aerosols from α -Pinene Ozonolysis and the Boreal Forest, *Environ. Sci. Technol. Lett.*, 3, 280-285, doi:10.1021/acs.estlett.6b00152, 2016.
- Kristensen, K., Jensen, L. N., Quéléver, L. L. J., Christiansen, S., Rosati, B., Elm, J., Teiwes, R., Pedersen, H. B., Glasius, M., Ehn, M., and Bilde, M.: The Aarhus Chamber Campaign on Highly Oxygenated Organic Molecules and Aerosols (ACCHA):
 850 particle formation, organic acids, and dimer esters from α -pinene ozonolysis at different temperatures, *Atmos. Chem. Phys.*, 20, 12549-12567, doi:10.5194/acp-20-12549-2020, 2020.
- Kulmala, M., Petäjä, T., Ehn, M., Thornton, J., Sipilä, M., Worsnop, D. R., and Kerminen, V.-M.: Chemistry of Atmospheric Nucleation: On the Recent Advances on Precursor Characterization and Atmospheric Cluster Composition in Connection with Atmospheric New Particle Formation, *Annu. Rev. Phys. Chem.*, 65, 21-37, doi:10.1146/annurev-physchem-040412-
 855 110014, 2014.
- Kurten, T., Tiisanen, K., Roldin, P., Rissanen, M., Luy, J. N., Boy, M., Ehn, M., and Donahue, N.: alpha-Pinene Autoxidation Products May Not Have Extremely Low Saturation Vapor Pressures Despite High O:C Ratios, *The journal of physical chemistry. A*, 120, 2569-2582, doi:10.1021/acs.jpca.6b02196, 2016.
- Kurtén, T., Møller, K. H., Nguyen, T. B., Schwantes, R. H., Misztal, P. K., Su, L., Wennberg, P. O., Fry, J. L., and Kjaergaard,
 860 H. G.: Alkoxy Radical Bond Scissions Explain the Anomalously Low Secondary Organic Aerosol and Organonitrate Yields From α -Pinene + NO₃, *The Journal of Physical Chemistry Letters*, 8, 2826-2834, doi:10.1021/acs.jpclett.7b01038, 2017.
- Lambe, A. T., Ahern, A. T., Williams, L. R., Slowik, J. G., Wong, J. P. S., Abbatt, J. P. D., Brune, W. H., Ng, N. L., Wright, J. P., Croasdale, D. R., Worsnop, D. R., Davidovits, P., and Onasch, T. B.: Characterization of aerosol photooxidation flow reactors: heterogeneous oxidation, secondary organic aerosol formation and cloud condensation nuclei activity
 865 measurements, *Atmos. Meas. Tech.*, 4, 445-461, doi:10.5194/amt-4-445-2011, 2011.



- Lee, B. H., Mohr, C., Lopez-Hilfiker, F. D., Lutz, A., Hallquist, M., Lee, L., Romer, P., Cohen, R. C., Iyer, S., Kurtén, T., Hu, W., Day, D. A., Campuzano-Jost, P., Jimenez, J. L., Xu, L., Ng, N. L., Guo, H., Weber, R. J., Wild, R. J., Brown, S. S., Koss, A., de Gouw, J., Olson, K., Goldstein, A. H., Seco, R., Kim, S., McAvey, K., Shepson, P. B., Starn, T., Baumann, K., Edgerton, E. S., Liu, J., Shilling, J. E., Miller, D. O., Brune, W., Schobesberger, S., D'Ambro, E. L., and Thornton, J. A.:
 870 Highly functionalized organic nitrates in the southeast United States: Contribution to secondary organic aerosol and reactive nitrogen budgets, *Proc. Natl. Acad. Sci. USA*, 113, 1516-1521, doi:10.1073/pnas.1508108113, 2016.
- Lee, C. P., Surdu, M., Bell, D. M., Dommen, J., Xiao, M., Zhou, X., Baccarini, A., Giannoukos, S., Wehrle, G., Schneider, P. A., Prevot, A. S. H., Slowik, J. G., Lamkaddam, H., Wang, D., Baltensperger, U., and El Haddad, I.: High-frequency gaseous and particulate chemical characterization using extractive electrospray ionization mass spectrometry (Dual-Phase-EESI-
 875 TOF), *Atmos. Meas. Tech.*, 15, 3747-3760, doi:10.5194/amt-15-3747-2022, 2022.
- Lopez-Hilfiker, F. D., Pospisilova, V., Huang, W., Kalberer, M., Mohr, C., Stefenelli, G., Thornton, J. A., Baltensperger, U., Prevot, A. S. H., and Slowik, J. G.: An extractive electrospray ionization time-of-flight mass spectrometer (EESI-TOF) for online measurement of atmospheric aerosol particles, *Atmos. Meas. Tech.*, 12, 4867-4886, doi:10.5194/amt-12-4867-2019, 2019.
- 880 Lopez-Hilfiker, F. D., Mohr, C., Ehn, M., Rubach, F., Kleist, E., Wildt, J., Mentel, T. F., Lutz, A., Hallquist, M., Worsnop, D., and Thornton, J. A.: A novel method for online analysis of gas and particle composition: description and evaluation of a Filter Inlet for Gases and AEROSols (FIGAERO), *Atmos. Meas. Tech.*, 7, 983-1001, doi:10.5194/amt-7-983-2014, 2014.
- Lopez-Hilfiker, F. D., Mohr, C., Ehn, M., Rubach, F., Kleist, E., Wildt, J., Mentel, T. F., Carrasquillo, A. J., Daumit, K. E., Hunter, J. F., Kroll, J. H., Worsnop, D. R., and Thornton, J. A.: Phase partitioning and volatility of secondary organic aerosol
 885 components formed from α -pinene ozonolysis and OH oxidation: the importance of accretion products and other low volatility compounds, *Atmos. Chem. Phys.*, 15, 7765-7776, doi:10.5194/acp-15-7765-2015, 2015.
- Mauderly, J. L. and Chow, J. C.: Health effects of organic aerosols, *Inhalation toxicology*, 20, 257-288, doi:10.1080/08958370701866008, 2008.
- McFiggans, G., Mentel, T. F., Wildt, J., Pullinen, I., Kang, S., Kleist, E., Schmitt, S., Springer, M., Tillmann, R., Wu, C., Zhao, D., Hallquist, M., Faxon, C., Le Breton, M., Hallquist, Å. M., Simpson, D., Bergström, R., Jenkin, M. E., Ehn, M., Thornton, J. A., Alfarra, M. R., Bannan, T. J., Percival, C. J., Priestley, M., Topping, D., and Kiendler-Scharr, A.: Secondary organic aerosol reduced by mixture of atmospheric vapours, *Nature*, 565, 587-593, doi:10.1038/s41586-018-0871-y, 2019.
- Mentel, T. F., Springer, M., Ehn, M., Kleist, E., Pullinen, I., Kurtén, T., Rissanen, M., Wahner, A., and Wildt, J.: Formation of highly oxidized multifunctional compounds: autoxidation of peroxy radicals formed in the ozonolysis of alkenes – deduced
 895 from structure–product relationships, *Atmos. Chem. Phys.*, 15, 6745-6765, doi:10.5194/acp-15-6745-2015, 2015.
- Molteni, U., Simon, M., Heinritzi, M., Hoyle, C. R., Bernhammer, A.-K., Bianchi, F., Breitenlechner, M., Brilke, S., Dias, A., Duplissy, J., Frege, C., Gordon, H., Heyn, C., Jokinen, T., Kürten, A., Lehtipalo, K., Makhmutov, V., Petäjä, T., Pieber, S. M., Praplan, A. P., Schobesberger, S., Steiner, G., Stozhkov, Y., Tomé, A., Tröstl, J., Wagner, A. C., Wagner, R., Williamson, C., Yan, C., Baltensperger, U., Curtius, J., Donahue, N. M., Hansel, A., Kirkby, J., Kulmala, M., Worsnop, D.
 900 R., and Dommen, J.: Formation of Highly Oxygenated Organic Molecules from α -Pinene Ozonolysis: Chemical Characteristics, Mechanism, and Kinetic Model Development, *ACS Earth and Space Chemistry*, 3, 873-883, doi:10.1021/acsearthspacechem.9b00035, 2019.
- Mutzel, A., Poulain, L., Berndt, T., Iinuma, Y., Rodigast, M., Böge, O., Richters, S., Spindler, G., Sipilä, M., Jokinen, T., Kulmala, M., and Herrmann, H.: Highly Oxidized Multifunctional Organic Compounds Observed in Tropospheric Particles:
 905 A Field and Laboratory Study, *Environ. Sci. Technol.*, 49, 7754-7761, doi:10.1021/acs.est.5b00885, 2015.



- Müller, L., Reinnig, M.-C., Hayen, H., and Hoffmann, T.: Characterization of oligomeric compounds in secondary organic aerosol using liquid chromatography coupled to electrospray ionization Fourier transform ion cyclotron resonance mass spectrometry, *Rapid Communications in Mass Spectrometry*, 23, 971-979, <https://doi.org/10.1002/rcm.3957>, 2009.
- Müller, M., Eichler, P., D'Anna, B., Tan, W., and Wisthaler, A.: Direct Sampling and Analysis of Atmospheric Particulate
 910 Organic Matter by Proton-Transfer-Reaction Mass Spectrometry, *Anal. chem.*, 89, 10889-10897, [doi:10.1021/acs.analchem.7b02582](https://doi.org/10.1021/acs.analchem.7b02582), 2017.
- Nah, T., Sanchez, J., Boyd, C. M., and Ng, N. L.: Photochemical Aging of α -pinene and β -pinene Secondary Organic Aerosol formed from Nitrate Radical Oxidation, *Environ. Sci. Technol.*, 50, 222-231, [doi:10.1021/acs.est.5b04594](https://doi.org/10.1021/acs.est.5b04594), 2016.
- Orlando, J. J. and Tyndall, G. S.: Laboratory studies of organic peroxy radical chemistry: an overview with emphasis on recent
 915 issues of atmospheric significance, *Chem. Soc. Rev.*, 41, 6294-6317, [doi:10.1039/C2CS35166H](https://doi.org/10.1039/C2CS35166H), 2012.
- Peräkylä, O., Riva, M., Heikkinen, L., Quéléver, L., Roldin, P., and Ehn, M.: Experimental investigation into the volatilities of highly oxygenated organic molecules (HOMs), *Atmos. Chem. Phys.*, 20, 649-669, [doi:10.5194/acp-20-649-2020](https://doi.org/10.5194/acp-20-649-2020), 2020.
- Pospisilova, V., Lopez-Hilfiker, F. D., Bell, D. M., El Haddad, I., Mohr, C., Huang, W., Heikkinen, L., Xiao, M., Dommen, J., Prevot, A. S. H., Baltensperger, U., and Slowik, J. G.: On the fate of oxygenated organic molecules in atmospheric aerosol
 920 particles, *Science Advances*, 6, eaax8922, [doi:10.1126/sciadv.aax8922](https://doi.org/10.1126/sciadv.aax8922), 2020.
- Presto, A. A., Huff Hartz, K. E., and Donahue, N. M.: Secondary Organic Aerosol Production from Terpene Ozonolysis. 2. Effect of NO_x Concentration, *Environ. Sci. Technol.*, 39, 7046-7054, [doi:10.1021/es050400s](https://doi.org/10.1021/es050400s), 2005.
- Pullinen, I., Schmitt, S., Kang, S., Sarrafzadeh, M., Schlag, P., Andres, S., Kleist, E., Mentel, T. F., Rohrer, F., Springer, M., Tillmann, R., Wildt, J., Wu, C., Zhao, D., Wahner, A., and Kiendler-Scharr, A.: Impact of NO_x on secondary organic
 925 aerosol (SOA) formation from α -pinene and β -pinene photooxidation: the role of highly oxygenated organic nitrates, *Atmos. Chem. Phys.*, 20, 10125-10147, [doi:10.5194/acp-20-10125-2020](https://doi.org/10.5194/acp-20-10125-2020), 2020.
- Reinnig, M. C., Warnke, J., and Hoffmann, T.: Identification of organic hydroperoxides and hydroperoxy acids in secondary organic aerosol formed during the ozonolysis of different monoterpenes and sesquiterpenes by on-line analysis using atmospheric pressure chemical ionization ion trap mass spectrometry, *Rapid Commun Mass Spectrom*, 23, 1735-1741,
 930 [doi:10.1002/rcm.4065](https://doi.org/10.1002/rcm.4065), 2009.
- Rissanen, M. P.: NO₂ Suppression of Autoxidation–Inhibition of Gas-Phase Highly Oxidized Dimer Product Formation, *ACS Earth and Space Chemistry*, 2, 1211-1219, [doi:10.1021/acsearthspacechem.8b00123](https://doi.org/10.1021/acsearthspacechem.8b00123), 2018.
- Rissanen, M. P., Kurten, T., Sipila, M., Thornton, J. A., Kangasluoma, J., Sarnela, N., Junninen, H., Jorgensen, S., Schallhart, S., Kajos, M. K., Taipale, R., Springer, M., Mentel, T. F., Ruuskanen, T., Petaja, T., Worsnop, D. R., Kjaergaard, H. G., and
 935 Ehn, M.: The formation of highly oxidized multifunctional products in the ozonolysis of cyclohexene, *J Am Chem Soc*, 136, 15596-15606, [doi:10.1021/ja507146s](https://doi.org/10.1021/ja507146s), 2014.
- Riva, M., Rantala, P., Krechmer, J. E., Peräkylä, O., Zhang, Y., Heikkinen, L., Garmash, O., Yan, C., Kulmala, M., Worsnop, D., and Ehn, M.: Evaluating the performance of five different chemical ionization techniques for detecting gaseous oxygenated organic species, *Atmos. Meas. Tech.*, 12, 2403-2421, [doi:10.5194/amt-12-2403-2019](https://doi.org/10.5194/amt-12-2403-2019), 2019.
- 940 Stark, H., Yatavelli, R. L. N., Thompson, S. L., Kang, H., Krechmer, J. E., Kimmel, J. R., Palm, B. B., Hu, W., Hayes, P. L., Day, D. A., Campuzano-Jost, P., Canagaratna, M. R., Jayne, J. T., Worsnop, D. R., and Jimenez, J. L.: Impact of Thermal Decomposition on Thermal Desorption Instruments: Advantage of Thermogram Analysis for Quantifying Volatility Distributions of Organic Species, *Environ. Sci. Technol.*, 51, 8491-8500, [doi:10.1021/acs.est.7b00160](https://doi.org/10.1021/acs.est.7b00160), 2017.
- Takeuchi, M. and Ng, N. L.: Chemical composition and hydrolysis of organic nitrate aerosol formed from hydroxyl and nitrate
 945 radical oxidation of α -pinene and β -pinene, *Atmos. Chem. Phys.*, 19, 12749-12766, [doi:10.5194/acp-19-12749-2019](https://doi.org/10.5194/acp-19-12749-2019), 2019.



- Vereecken, L., Glowacki, D. R., and Pilling, M. J.: Theoretical Chemical Kinetics in Tropospheric Chemistry: Methodologies and Applications, *Chem. Rev.*, 115, 4063-4114, doi:10.1021/cr500488p, 2015.
- Wagner, A. C., Bergen, A., Brilke, S., Fuchs, C., Ernst, M., Hoker, J., Heinritzi, M., Simon, M., Böhner, B., Curtius, J., and Kürten, A.: Size-resolved online chemical analysis of nanoaerosol particles: a thermal desorption differential mobility
950 analyzer coupled to a chemical ionization time-of-flight mass spectrometer, *Atmos. Meas. Tech.*, 11, 5489-5506, doi:10.5194/amt-11-5489-2018, 2018.
- Wang, Y., Riva, M., Xie, H., Heikkinen, L., Schallhart, S., Zha, Q., Yan, C., He, X.-C., Peräkylä, O., and Ehn, M.: Formation of highly oxygenated organic molecules from chlorine-atom-initiated oxidation of alpha-pinene, *Atmos. Chem. Phys.*, 20, 5145-5155, doi:10.5194/acp-20-5145-2020, 2020.
- 955 Wang, Z., Popolan-Vaida, D. M., Chen, B., Moshhammer, K., Mohamed, S. Y., Wang, H., Sioud, S., Raji, M. A., Kohse-Höinghaus, K., Hansen, N., Dagaut, P., Leone, S. R., and Sarathy, S. M.: Unraveling the structure and chemical mechanisms of highly oxygenated intermediates in oxidation of organic compounds, *Proc. Natl. Acad. Sci. USA*, 114, 13102-13107, doi:10.1073/pnas.1707564114, 2017.
- Yasmeen, F., Vermeylen, R., Szmigielski, R., Iinuma, Y., Böge, O., Herrmann, H., Maenhaut, W., and Claeys, M.: Terpenylic
960 acid and related compounds: precursors for dimers in secondary organic aerosol from the ozonolysis of α - and β -pinene, *Atmos. Chem. Phys.*, 10, 9383-9392, doi:10.5194/acp-10-9383-2010, 2010.
- Yatavelli, R. L. N., Lopez-Hilfiker, F., Wargo, J. D., Kimmel, J. R., Cubison, M. J., Bertram, T. H., Jimenez, J. L., Gonin, M., Worsnop, D. R., and Thornton, J. A.: A Chemical Ionization High-Resolution Time-of-Flight Mass Spectrometer Coupled to a Micro Orifice Volatilization Impactor (MOVI-HRToF-CIMS) for Analysis of Gas and Particle-Phase Organic Species,
965 *Aerosol Sci. Technol.*, 46, 1313-1327, doi:10.1080/02786826.2012.712236, 2012.
- Yatavelli, R. L. N., Mohr, C., Stark, H., Day, D. A., Thompson, S. L., Lopez-Hilfiker, F. D., Campuzano-Jost, P., Palm, B. B., Vogel, A. L., Hoffmann, T., Heikkinen, L., Aijala, M., Ng, N. L., Kimmel, J. R., Canagaratna, M. R., Ehn, M., Junninen, H., Cubison, M. J., Petaja, T., Kulmala, M., Jayne, J. T., Worsnop, D. R., and Jimenez, J. L.: Estimating the contribution of organic acids to northern hemispheric continental organic aerosol, *Geophys. Res. Lett.*, 42, 6084-6090,
970 doi:10.1002/2015gl064650, 2015.
- Zhang, H., Yee, L. D., Lee, B. H., Curtis, M. P., Worton, D. R., Isaacman-VanWertz, G., Offenberg, J. H., Lewandowski, M., Kleindienst, T. E., Beaver, M. R., Holder, A. L., Lonneman, W. A., Docherty, K. S., Jaoui, M., Pye, H. O. T., Hu, W., Day, D. A., Campuzano-Jost, P., Jimenez, J. L., Guo, H., Weber, R. J., de Gouw, J., Koss, A. R., Edgerton, E. S., Brune, W., Mohr, C., Lopez-Hilfiker, F. D., Lutz, A., Kreisberg, N. M., Spielman, S. R., Hering, S. V., Wilson, K. R., Thornton, J. A.,
975 and Goldstein, A. H.: Monoterpenes are the largest source of summertime organic aerosol in the southeastern United States, *Proc. Natl. Acad. Sci. USA*, 115, 2038-2043, doi:10.1073/pnas.1717513115, 2018.
- Zhang, X., McVay, R. C., Huang, D. D., Dalleska, N. F., Aumont, B., Flagan, R. C., and Seinfeld, J. H.: Formation and evolution of molecular products in α -pinene secondary organic aerosol, *Proc. Natl. Acad. Sci. USA*, 112, 14168-14173, doi:10.1073/pnas.1517742112, 2015.
- 980 Zhang, X., Lambe, A. T., Upshur, M. A., Brooks, W. A., Gray Bé, A., Thomson, R. J., Geiger, F. M., Surratt, J. D., Zhang, Z., Gold, A., Graf, S., Cubison, M. J., Groessl, M., Jayne, J. T., Worsnop, D. R., and Canagaratna, M. R.: Highly Oxygenated Multifunctional Compounds in α -Pinene Secondary Organic Aerosol, *Environ. Sci. Technol.*, 51, 5932-5940, doi:10.1021/acs.est.6b06588, 2017.
- Zhao, D., Schmitt, S. H., Wang, M., Acir, I. H., Tillmann, R., Tan, Z., Novelli, A., Fuchs, H., Pullinen, I., Wegener, R., Rohrer,
985 F., Wildt, J., Kiendler-Scharr, A., Wahner, A., and Mentel, T. F.: Effects of NO_x and SO₂ on the secondary organic aerosol



formation from photooxidation of α -pinene and limonene, Atmos. Chem. Phys., 18, 1611-1628, doi:10.5194/acp-18-1611-2018, 2018a.

Zhao, Y., Thornton, J. A., and Pye, H. O. T.: Quantitative constraints on autoxidation and dimer formation from direct probing of monoterpene-derived peroxy radical chemistry, Proc. Natl. Acad. Sci. USA, 115, 12142-12147, 990 doi:10.1073/pnas.1812147115, 2018b.