

1 **Air pollution reductions caused by the COVID-19 lockdown open up a way to preserve**
2 **the Himalayan Glaciers**

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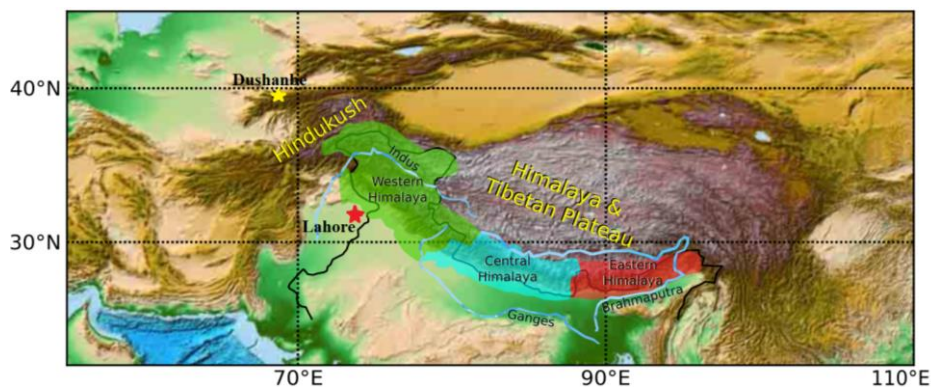
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13 Abstract

14 The rapid melting of glaciers in the Hindu Kush Himalayas (HKH) during recent decades poses
15 an alarming threat to water security for larger parts of Asia. If this melting persists, the entire
16 Himalayan glaciers are estimated to disappear by end of the 21st century. Here, we assess the
17 influence of the spring 2020 COVID-19 lockdown on the HKH, demonstrating the potential
18 benefits of a strict emission reduction roadmap. Chemistry-climate model simulations,
19 supported by satellite and ground measurements, show that lower levels of gas and aerosol
20 pollution during lockdown led to changes in meteorology, and to a reduction in black carbon
21 in snow (2-14%) and thus in snow melting (10-40%). This caused increases in snow cover (6-
22 12%) and mass (2-20%) and a decrease in runoff (5-55%) over the HKH and Tibetan Plateau,
23 ultimately leading to an enhanced snow-equivalent-water (2–55%). We emphasize the
24 necessity for immediate anthropogenic pollution reductions to address the hydro-climatic threat
25 to billions of people in South Asia.

27 **1. Introduction**

28 The Hindu Kush Himalayan (HKH) mountains and Tibetan plateau is the largest snow-cladded
29 region outside the Poles (Fig. 1). This region is also referred to as High Mountain Asia,
30 although that includes the Tien Shan and some other northern ranges. The HKH meltwater
31 feeds rivers in India and China that drive the agriculture, hydropower generation, and economy
32 of these countries (Hussain et al., 2019; Sabin et al., 2020; Lee et al. 2021a). The Himalayan
33 snowmelt in spring provides ~50% of the annual freshwater to ~4 billion people of South Asia
34 and East Asia (Sarangi et al 2019, Sabin et al., 2020). Rapid Himalayan snowmelt caused a
35 loss of ~40 % of the Himalayan glacier area compared to the Little Ice Age, 400 to 700 years
36 ago, i.e. ~0.92 to 1.38 mm sea-level equivalent (Lee et al., 2021b). The snow mass over the
37 Himalayas has generally decreased during the last 30 years (except for a few Karakoram
38 glaciers that show an increasing trend in snow mass) (Hussain et al., 2019. The alarming rate
39 of snow melting of 0.02 to 0.6 cm °C⁻¹ day⁻¹ raised concerns about the sustainability of water
40 supply (Tiwari et al., 2015) and loss of glaciers in the region (Hussain et al., 2019, Lee et al.,
41 2021b). Model simulations for extreme scenarios show that Himalaya snow melting could
42 cause the glaciers to disappear by the end of the 21st century (Cruz et al. 2007, Hock et al.,
43 2019).



48 **Figure 1:** Map of the Hindu Kush Himalayas region with the Western (70 - 80° E, 30° - 35° N),
49 Central (80° - 87° E, 28° - 30° N), and Eastern Himalayas (88° - 95° E, 26° - 30° N). A yellow
50 and red star indicates the location of the AERONET sun photometer stations Dushanbe
51 (68.858° E, 38.553° N) and Lahore (74.264° E, 31.480° N), respectively.

52 The accelerated thinning of Himalayan glaciers is attributed to climate change causing
53 shifts in air temperature and precipitation, as well as the atmospheric distribution and
54 deposition of light-absorbing particles i.e., dust and black carbon (BC) (IPCC Climate Change
55 2013, Krishnan et al., 2019). Among the aforementioned factors, snow darkening due to the
56 deposition of absorbing aerosols is an integral component of Himalayan snowmelt and runoff
57 (Lau et al., 2010). The snow-melting efficacy of BC is higher than that of greenhouse gases
58 (Qian et al., 2011; Nair et al. 2013; Ma et al., 2019; Sarangi et al., 2019). The increasing energy
59 demand of the densely populated South Asian region has increased the emission of greenhouse
60 gases and BC aerosol in the last few decades (Fadnavis et al., 2017, Krishnan et al., 2020),
61 leading to enhanced darkening and snow melting (Usha et al., 2021).

62

63 The economic slowdown caused by the COVID-19 pandemic measures led to a drastic
64 reduction in public and freight transportation, industrial emissions, and energy use (Fadnavis
65 et al., 2021a). This resulted in a substantial decline in emissions of several atmospheric
66 pollutants including greenhouse gases and black carbon aerosol (Forster et al. 2020; Kanniah
67 et al., 2020; Le Quéré et al 2020), and potentially reduced deposition of dark aerosols on snow
68 and ice (Bair et al., 2021). Remote sensing approaches show cleaner snow with ~30% less
69 light-absorbing impurities in snow during the lockdown period over Asia between March and
70 May 2020 (Bair et al 2021). This led to decreased snowmelt by 25 – 70 mm in 2020 compared
71 to the last 20-year mean for March-May over Western Himalayas due to decreased radiative
72 forcing induced by BC and dust deposition on snow/ice surfaces and related changes in snow
73 absorption and surface albedo (Bair et al., 2021). Bair et al. (2021) also found that 6.6 km⁻³ of
74 melt water stayed in the Indus Basin. Gauge and reservoir data for this part of the world,
75 however, are not freely available. Impacts of reduced levels of air pollution on changes in the
76 snow mass, surface water runoff, and water reservoir over the HKH are not reported hitherto.

77 Here, we provide a detailed analysis of the impact of reduced pollution over HKH and Tibetan
78 plateau region during the COVID-19 lockdown period between March and May 2020. We used
79 global simulations with the chemistry-climate model ECHAM6-HAMMOZ (Schultz et al.,
80 2018, Tegen et al., 2019), updated with an improved BC-in-snow parameterization (Huang
81 2018), in order to contrast the 2020 COVID-19 (COVID) with the typical, unchanged (control,
82 CTL) air pollution conditions. The COVID simulations are performed using a COVID-19
83 emission inventory where emissions are reduced based on Google and Apple mobility data
84 (Forster et al., 2020; details in section 2.2).

85 **2. Methods**

86 **2.1 Observational data**

87 We used monthly snow cover fraction from NASA's Moderate Resolution Imaging
88 Spectroradiometer (MODIS) satellite product on a $0.5 \times 0.5^\circ$ resolution (version 6, level 3; Hall
89 et al., 2006) for the years 2000 – 2020 (<https://nsidc.org/data/MOD10CM/versions/6>). For
90 aerosol information we used monthly mean satellite AOD at $1 \times 1^\circ$ resolution from the MODIS
91 Terra level-3 dark target and deep blue retrievals at 550 nm wavelength for 2001-2020
92 (<https://giovanni.gsfc.nasa.gov>). Uncertainty in MODIS AOD data over snow are documented
93 by Huang et al (2020). We also used ground-based sun photometer observations of AOD from
94 the Aerosol Robotic Network (AERONET) (Martonchik et al., 2004) at the stations Dushanbe
95 (68.858° E, 38.553° N) for the period 2010-2020 and Lahore (74.264° E, 31.480° N) for the
96 period 2006 – 2020, situated in HKH region (<https://aeronet.gsfc.nasa.gov>).

97

98 **2.2 The ECHAM6-HAMMOZ model description and Experimental set-up**

99 We performed 10-member ensemble experiments using the state-of-the-art aerosol-
100 chemistry-climate model ECHAM6-HAMMOZ (version echam6.3-ham2.3-moz1.0; Schultz et

101 al., 2018, Tegen et al., 2019). The model comprises the atmospheric general circulation model
102 ECHAM6 (Stevens et al., 2013), the atmospheric chemistry module MOZ (Schultz et al, 2018),
103 and the Hamburg Aerosol Model (HAM; Stier et al., 2005; Zhang et al., 2012). The HAM
104 component predicts the nucleation, growth, evolution, and sinks of sulphate (SO_4^{2-}), black
105 carbon (BC), particulate organic matter (POM), sea salt (SS), and mineral dust (DU) aerosols.
106 Seven log-normal modes describe the size distribution of the aerosol population with a
107 prescribed variance in the aerosol module. The MOZ submodule describes the trace gas
108 chemistry from the troposphere to the lower thermosphere. The chemical mechanism includes
109 the O_x , NO_x , HO_x , ClO_x and BrO_x chemical families, along with CH_4 and its degradation
110 products. Several primary non-methane hydrocarbons (NMHCs) and related oxygenated
111 organic compounds are also described. It contains 108 species, 71 photolytic processes, 218
112 gas-phase reactions and 18 heterogeneous reactions with aerosol (Schultz et al., 2018). Details
113 of emissions (anthropogenic, biomass burning, biogenic, fossil fuel etc.) and model
114 parametrisation and other details are reported in the past Fadnavis et al. (2017, 2019a,b, 2021b).
115 Anthropogenic and biomass burning emissions of sulphate, and black carbon (BC) and organic
116 carbon (OC) are based on the AEROCOM-ACCMIP-II emission inventory for year 2020
117 (Lamarque et al., 2010; Textor et al., 2006). Additional consideration for the reduction of snow
118 albedo due to BC in snow is implemented but extended for the MOZ module. The model also
119 accounts for changes in snow albedo due to airborne BC deposited in the top layer of snow,
120 while the effect of other in-snow aerosol particles (e.g. dust, OC) is not simulated in the model.
121 Influxes of BC in snow include below-cloud and in-cloud wet scavenging, as well as dry
122 deposition and sedimentation. Snowmelt and glacier runoff remove the in-snow BC at a
123 reduced efficiency, leading to enhanced concentration, while fresh and pristine snowfall leads
124 to reductions in BC concentration.

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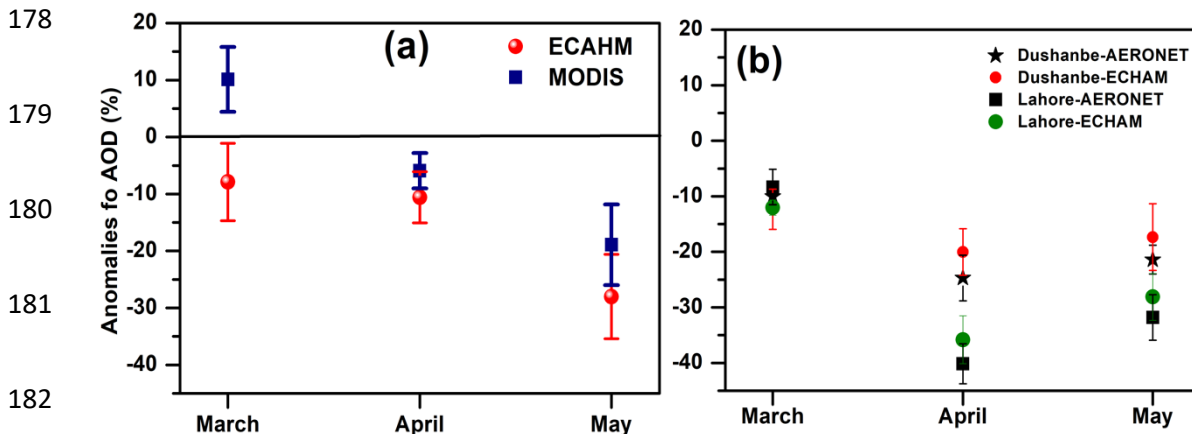
126 The model simulations were performed at T63 horizontal resolution ($1.875^\circ \times 1.875^\circ$) with 47
127 levels in the vertical from the surface to 0.01 hPa (corresponding to approx. 80 km), and with
128 a time step of 20 minutes. To understand the effect of the COVID-19 restrictions on snow over
129 Himalayas and Tibetan plateau region we conducted a control (CTL) and a COVID-19
130 (COVID) simulation. We adopted an ensemble approach (with 10 ensemble members) for the
131 above two experiments. Ten spin-up simulations were performed from 1 to 31 December 2019
132 to generate stabilised initial fields for the 10 ensemble members. Emissions were the same in
133 each of the 10 members during the spin-up period. Control simulations were extended with the
134 same setup until 1 June 2020. While for the COVID simulations (10 ensemble members each),
135 the anthropogenic emission of all gases and aerosols were changed since 1 January 2020
136 according to Google and Apple mobility data as in Forster et al. (2020). The COVID-19
137 emissions were prepared by deriving scaling factors between the input4MIPS SSP245 baseline
138 and the version5 of the Forster et al. (2020) 2-year blip scenario, separately for each species
139 and each grid point (see Fig. S5a). Subsequently, these scaling factors have been applied to the
140 AeroCom-II ACCMIP emissions. This ensures consistency of the drop in emissions
141 independent of the absolute emission values in the AeroCom-II ACCMIP and the input4MIPS
142 SSP245 data sets. The global mean emission changes in carbon monoxide (CO, 2-24%), black
143 carbon (BC, 3-23%), organic carbon (OC, 2-17%), sulfur dioxide (SO₂, 3-23%), nitrogen o
144 xides (NO_x, 2-30%), methane (CH₄, 2-5%), and ammonia (NH₃, 0-3%) during the period
145 January to 1 July 2020 (COVID - CTL) are in agreement with previous studies Forster et al.
146 (2020) and Le Quéré et al., (2020) (Fig. S5b). Our model experiments follow the CovidMIP
147 protocol (Jones et al., 2021). The COVID and CTL simulations ended on 1 June 2020. To
148 investigate the effects of COVID-19 emissions in spring (i.e., since 1 March 2020), we
149 analysed the difference between COVID and CTL simulations for the spring season in 2020.
150 The same dust parametrisation was employed in the CTL and COVID simulations.

151 A limitation of our simulation is the relatively coarse spatial resolution in the ECHAM6-
152 HAMMOZ model (1.875°x1.875°). Other studies used a finer spatially resolved regional
153 model; for example Sarangi et al. (2020) use a 12 x 12 km (~ 0.10°) grid in the regional WRF-
154 Chem-SNICAR model over the same region. In our model grid of 1.875°, many of the
155 Himalayan sub ranges are smaller than a pixel, and, hence, the topographic influences, which
156 are substantial in the mountains are limited. One effect may be that snowfall and snow on the
157 ground are underestimated (e.g., Liu et al., 2022). The coarse grid size can impact the anomalies
158 found here as the changes in snow mass are small, at most +16 mm, and the bias in the likely
159 underestimated snow mass may change between the control and COVID simulations. Biases
160 are, however, the same in the control and COVID simulations and, thus, their effects will be
161 diluted when we compute the anomalies.

162 **3 Comparison of AOD over Western, Central, Eastern Himalayas and Tibetan Plateau** 163 **regions**

164 We elaborate on the comparison of MODIS AOD with our model simulations over
165 Western, Central, Eastern Himalayas and Tibetan Plateau regions (Fig. S6). Both MODIS and
166 the model show a reduction in AOD during spring 2020 over the aforementioned regions of
167 HKH. The estimated differences in AOD during March to May 2020 vary between 0.8 – 11%
168 over Western and Central Himalayas, and 8 – 16% over Eastern Himalayas. Over the Tibetan
169 plateau region, in contrast to the model simulations, MODIS shows an enhancement (2 – 16 %) in
170 AOD (Fig. S6). This may be due to dust aerosols, which are transported during spring from
171 western Asia and locally, generating dust piles over the Tibetan Plateau (Fadnavis et al., 2017,
172 2021a). The simulated dust aerosol concentration in spring 2020 over the Tibetan Plateau
173 region is smaller in the COVID than in the non-COVID (i.e. CTL) situation (Fig. S1c). The
174 changes in simulated dust are a response to meteorology differences between the COVID and
175 CTL simulations (Fig. S7).

177 4.1 Reduction of airborne aerosols and in-snow BC concentration over the Himalayas



183 **Figure 2:** (a) Changes in monthly mean AOD (%) during March - May 2020 from MODIS in
 184 comparison to mean of 2001-2019 and ECHAM-HAMMOZ (COVID minus CTL) averaged
 185 over the Hindu Kush Himalayas (HKH) and Tibetan Plateau region (75° - 95° E, 30° - 35° N),
 186 (b) same as (a) but for AOD from AERONET observations and ECHAM-HAMMOZ model
 187 results at Dushanbe (68.858° E, 38.553° N, climatology 2010-2019) and Lahore (74.264° E,
 188 31.480° N, climatology 2006-2019). Vertical bars in Fig (a)-(b) indicate the standard deviation
 189 within ten members of model simulations, and within monthly mean anomalies from MODIS
 190 for years 2001-2019.

191

192 The COVID-19 lockdown restrictions in spring 2020 decreased the anthropogenic
 193 aerosol amounts over the HKH ranges (Western, Central, and Eastern Himalayas), and the
 194 Tibetan Plateau region. The ECHAM6-HAMMOZ model simulations show that COVID
 195 lockdown resulted in a cleaner atmosphere during March - May 2020 over the HKH ranges and
 196 Tibetan Plateau region. There is a reduced level of Aerosol Optical Depth (AOD) over the
 197 region throughout spring 2020 by -8.1 ± 6.2 % in March, -10.2 ± 4.7 % in April, -27 ± 6.9 % in
 198 May compared to the CTL (non COVID) simulation (Fig. 2a). This is supported by NASA's
 199 Moderate Resolution Imaging Spectroradiometer (MODIS) measurements also showing a
 200 reduction in AOD in April (-5.6 ± 3.3 %) and May (-18.8 ± 7.2 %) 2020 compared to the mean
 201 over the last 20 years (Fig. 2a). Thus, both model simulations and MODIS AOD show a
 202 reduction in aerosol pollution in April - May 2020. For March 2020, MODIS measurements

203 show AOD enhancement by 10.2 ± 4.8 %, which is due to increased dustiness over the HKH
204 region (see section 3 for a detailed discussion). AOD measurements at two Aerosol Robotic
205 Network (AERONET) sun photometer stations in Dushanbe (68.858° E, 38.553° N)
206 and Lahore (74.264° E, 31.480° N) show an AOD reduction in agreement with our model
207 simulations (Fig. 2b). There are differences among MODIS, AERONET and the model. The
208 changes in AOD during COVID compared to no-COVID period is less in the model than the
209 MODIS observations by 4.2 - 9.8 % and higher than the AERONET observations by 1.8 - 4.2
210 %. These differences are due to the fact that the simulated AOD change is in response to the
211 reduction of anthropogenic aerosols and associated circulation responses, while MODIS and
212 AERONET measurements show the effect of all atmospheric processes. Also, note that the
213 MODIS AOD values are spatial averages representative for a relatively large area while the
214 AERONET values are point measurements. Importantly, changes in simulated AOD in 2020
215 fall within the standard deviation of satellite and ground-based measurements indicating
216 reliability of our simulations (except for March 2020 with respect to MODIS). Our model
217 simulations also show a reduction in BC burden by 15 - 55% (Fig. S1a), and sulfate burden by
218 22 - 24 % over the HKH and Tibetan Plateau regions in spring 2020 (Fig. S1b). Interestingly,
219 dust burden also shows a reduction over these regions (Fig. S1c, Fig. S2a-c), except over central
220 Himalaya in March and April 2020. The lower dust load is related to the interactive change in
221 atmospheric dynamics in the model, which also leads to changes in the wet and dry deposition
222 rates of dust (Fig. S2d-i) (details in section 3). A drop in BC is also observed in Aerosol
223 Radiative Forcing Over India Network (ARFINET) ground-based measurements over the Indo-
224 Gangetic Plain (> 50 %), north-eastern India ($>30\%$), Himalaya regions (16 - 60%), and Tibet
225 (70%) during spring 2020 (Gogoi et al., 2021; Liu et al., 2021). A similar impact of the
226 reduction of energy consumptions on decrease in AOD during the COVID-19 lockdown period,
227 i.e., in spring 2020 compared to the 2010-2019 climatology is also seen over South and East

228 Asia (40 %) and the Indo-Gangetic Plain (IGP) by 30 – 40 % in satellite measurements
229 (Fadnavis et al., 2021a; Srivastava et al., 2021; Pandey et al., 2021; Shafeeque et al , 2021).

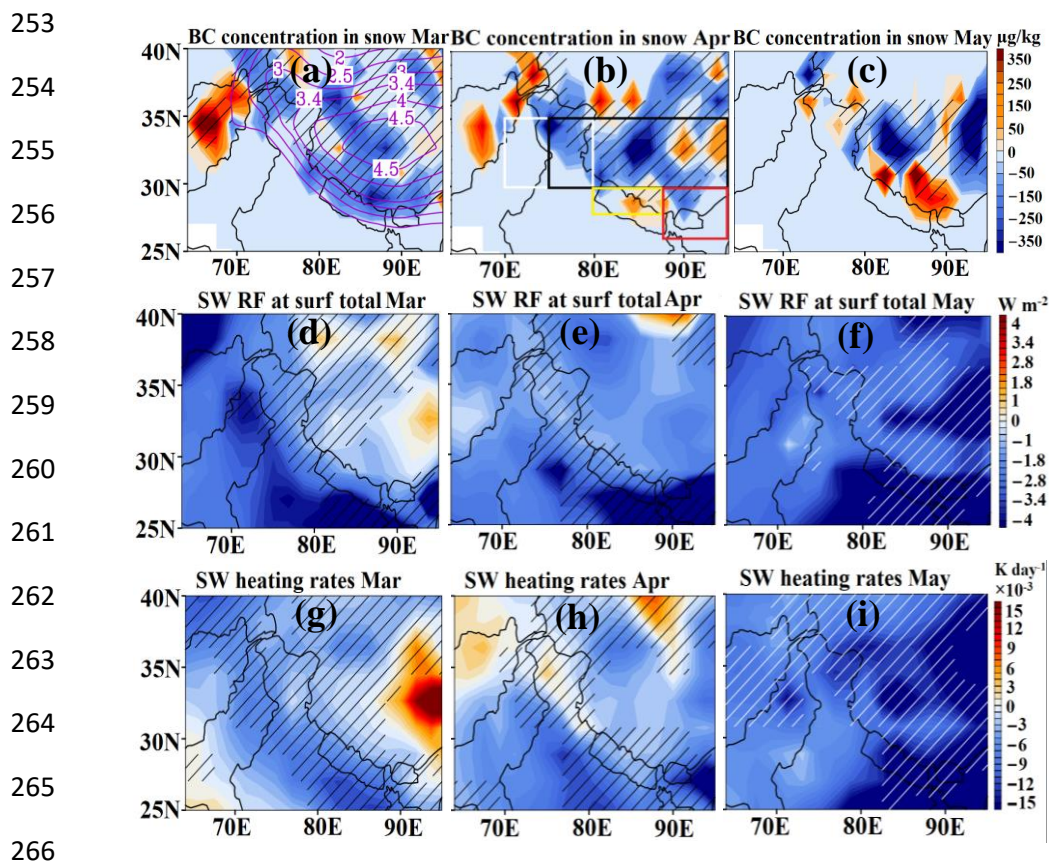
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231 The reduction in anthropogenic air pollution leads widely to a reduction in BC
232 concentration in the snow of approximately 25 - 350 $\mu\text{g kg}^{-1}$ (by 12 – 35 %) during spring 2020
233 (Fig. 3a-c) that reduce the snow darkening effect by embedded aerosol impurities. At the most
234 this amounts to about a 1.6% increase in visible snow albedo. Sporadically, however, the BC-
235 in snow concentrations have also increased in some areas of the Hindukush, Eastern Himalayas
236 and Kunlun Mountains. There are many factors at play that may lead to an increase in BC
237 concentration in snow in some locations. For instance, this includes increases in deposition of
238 BC following shifts in the atmospheric circulation (Fig.S3), accumulation of BC on surface
239 snow following partial snowmelt and minimal fresh snowfall, and less frequent occurrences of
240 complete snowmelt which would otherwise remove all accumulated BC in snow. Our
241 simulations reveal that the decrease in BC-in snow concentration and the overall reduction in
242 atmospheric pollution, as well as associated radiative effects, have decreased the shortwave
243 radiative forcing at the surface by 0.2 – 2 W m^{-2} in March – May 2020 (Fig. 3 d-f), leading to
244 a decrease in tropospheric heating by solar radiation of 0.001 to 0.015 K day^{-1} (Fig. 3 g-i). The
245 reduced anthropogenic BC over the HKH and Tibetan Plateau region resulted in less absorption
246 and re-emission of longwave radiation and, as a consequence, there is a reduction in longwave
247 radiative forcing in the atmosphere leading to a lower atmospheric heating (Fig. S4). This
248 decreased heating of the snowpack and tropospheric column is the combined effect of the reductions of
249 BC in snow, as well as changes in atmospheric concentrations of sulfate, OC and BC.

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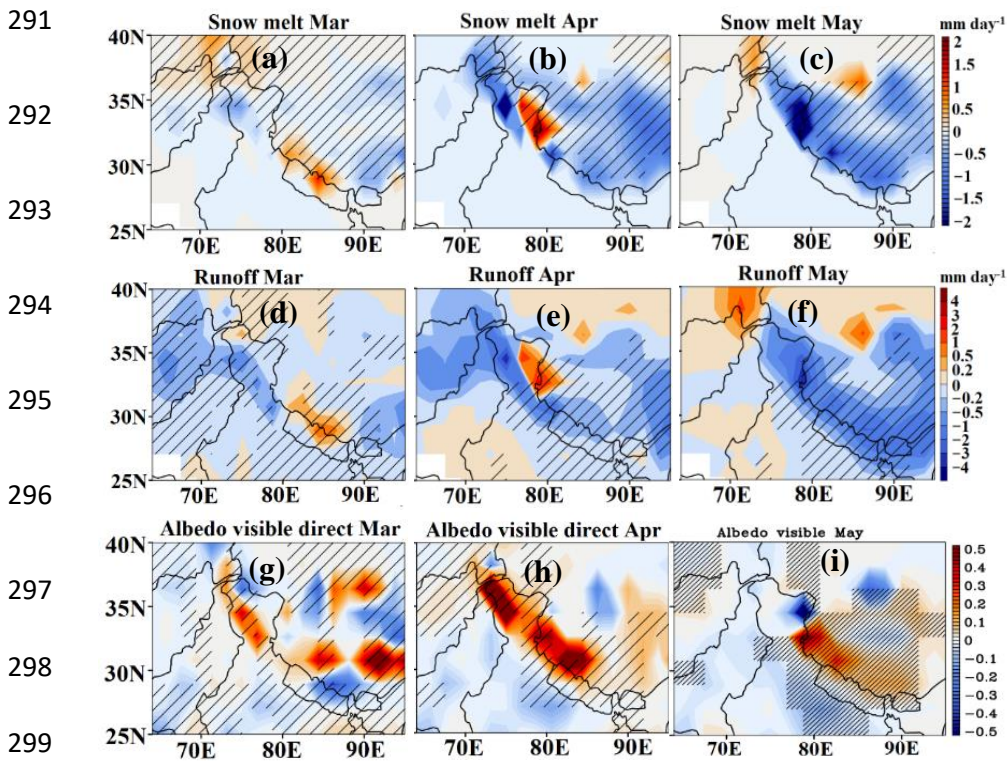


267 **Figure 3:** Spatial distribution of anomalies (COVID minus CTL) of BC concentration in snow
 268 ($\mu\text{g kg}^{-1}$) for (a) March, (b) April, and (c) May 2020; (d-f) shortwave radiative forcing (W m^{-2})
 269 at the surface and (e-g) tropospheric heating rates (K day^{-1}) due to changes in BC
 270 concentration in snow (COVID minus CTL). Hatched areas indicate the 95%-significance
 271 level. Contours in panel (a) indicate topography in km. Boxes in panel (b) indicate boundaries
 272 of Western Himalayas (WH, white), Central Himalayas (CH, yellow), Eastern Himalayas (EH,
 273 red) and Tibetan Plateau (black).

274 4.2 Impacts on snow melting, surface water runoff, and snow cover

275 Further we show that the decrease in aerosol pollution reduced the snow melting in
 276 spring 2020 by 0.2 to 2.5 mm day^{-1} corresponding to 10 – 50 % (Fig. 4 a-c). The amount of
 277 reduction of snow melting is pronounced over the western Himalayas in May. As a result of a
 278 reduction in snowmelt, surface water runoff has been drastically reduced by 2-4 mm -ay^{-1} (5 -
 279 55 %) (Fig. 4 d-f). The reduction in the runoff is most pronounced in May over the entire
 280 Himalayas and central Tibetan Plateau region. Estimates from remote sensing measurements
 281 also show the reduction of runoff by 6.5 km^3 of melted water in the Indus River Basin (Bair et
 282 al. 2020). In the past, studies have shown that elevated levels of light-absorbing aerosols

283 (elemental carbon: 13 to 75 ng g⁻¹ and dust: 32 to 217 μg g⁻¹) can contribute to about 3 to
 284 10 mm day⁻¹ of snowmelt over western Himalayas (Thind et al. 2019). A sensitivity analysis
 285 by (Santra et al., 2019) using a glacier mass balance model shows that BC-induced snow albedo
 286 reduction leads to an increase in annual runoff of 4 – 18%. In contrast to impacts of rising
 287 anthropogenic emissions during the past decades, emission reductions during the 2020
 288 COVID-19 lockdown period caused a brighter snow albedo and therefore an enhanced surface
 289 reflection with albedo increases of 0.2 – 0.5 (see Fig. 4g-i), leading to less atmospheric heating
 290 as well as associated reduced snowmelt and surface water runoff in spring 2020.

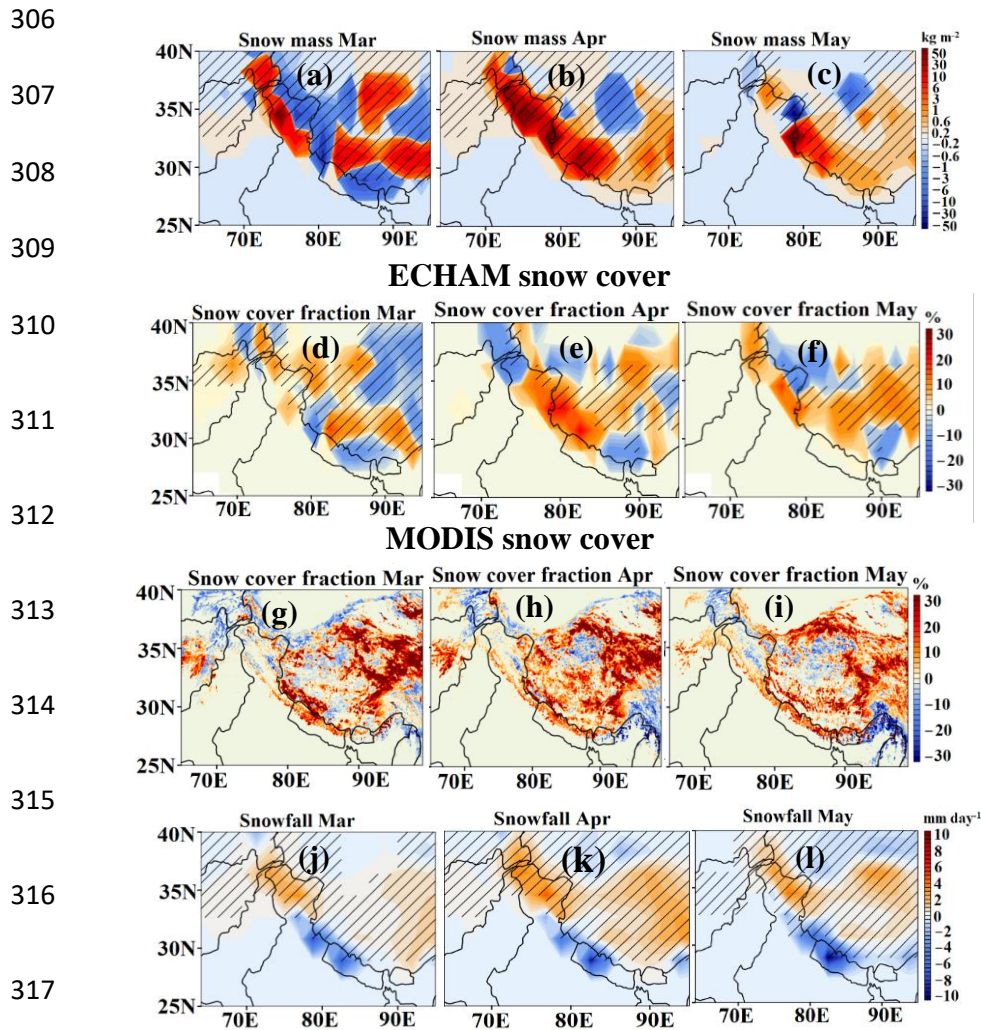


300 **Figure 4:** Spatial distribution of anomalies of (a-c) snow melt (mm day⁻¹), (d-f) surface water
 301 runoff (mm day⁻¹) for March to May 2020 (COVID minus CTL) and (g-i) surface albedo mean
 302 in the visible. Hatched areas indicate the 95%-significance level.

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305



318 **Figure 5:** Monthly mean anomalies (COVID minus CTL) for March to May 2020 of (a-c) the
 319 snow mass (kg m^{-2}), (d-f) snow cover fraction (%) as modelled by ECHAM6-HAMMOZ as
 320 well as (g-i) snow cover fraction from MODIS satellite measurements (%) with respect to the
 321 climatological average 2000-2019, and, (j-l) snowfall as modelled by ECHAM6-HAMMOZ.
 322 Hatched areas indicate the 95%-significance level.

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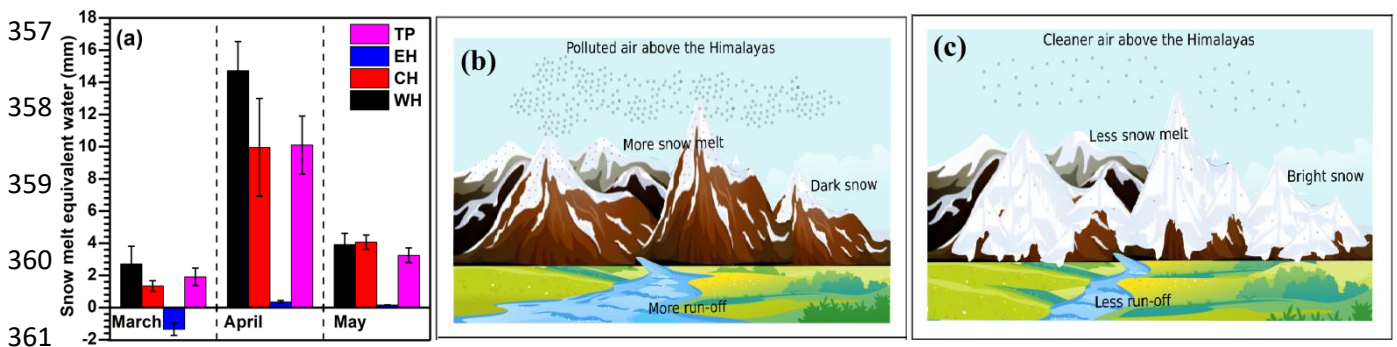
324 Our simulations also indicate that these changes lead to an increase in snow mass of
 325 $0.2\text{-}50 \text{ kg m}^{-2}$, i.e. 10-40% (Fig. 5a-c) and snow cover fraction of 2-30% during spring 2020
 326 (Fig. 5d-f). MODIS measurements also show a remarkable agreement with the model
 327 simulations (especially during April - May 2020), with increased snow cover of about 15-30%
 328 over the parts of Western Himalayas and Central Himalayas and the Tibetan Plateau region
 329 and decreased by 5-12 % over parts of North-East Himalayas especially in April and May 2020
 330 (Fig. 5 g-i). However, there are also some differences in terms of exact regions of snow cover

331 enhancement or reduction respectively, since the MODIS observations include the influence of
332 real-time meteorology, while meteorology in the model ensemble include internal variability
333 and do not replicate the exact conditions observed by MODIS. Our model simulations show
334 that air pollution reductions in the COVID-19 lockdown period and associated changes in
335 radiative forcing caused changes in the tropospheric circulation and thermodynamics (see
336 Fadnavis et al., 2020 for a detailed analysis). These changes in meteorology have increased
337 snowfall by 2-5 mm day⁻¹ (3-20 %) over the Western Himalayas and Tibetan Plateau region
338 (Fig. 5 j-l). The increase in snowfall over these regions will contribute to enhancement in snow
339 mass and snow cover (Fig. 5 a-f) and albedo (Fig. 4 g-i). In a few areas, however, this also
340 contributes to a more efficient BC deposition on snow, as described above (Fig. 3). This BC
341 enrichment in snow at a few places, however, has no influence on the fact that overall the
342 COVID-19 measures reduced the BC-in snow concentration and thus increased the visible
343 snow albedo (see Fig. 4g-i).

344

345 Himalaya snow is the largest source of freshwater for South Asia (Bolch et al., 2012).
346 The impact of reduced pollution on the surface water content in the Himalayas from our model
347 simulations is illustrated in Fig. 6a. The snow mass enhancement led to increase the snow
348 equivalent water by 2 to 14.7 mm (2.5 to 55 %). The western Himalayas show the highest
349 increase in snow equivalent water by 14.7 mm (55 %) followed by the Tibetan Plateau by 12
350 mm (by 22 %) and central Himalayas by 10 mm (by 18%) in April while the Eastern Himalayas
351 show a decrease in March (-1.3 mm; 10 %) and small enhancement in April by 1.1mm (2.3 %)
352 and May 2020 by 1.3 mm (2.7%) due to pollution reduction. Thus, human induced pollution
353 reduction during the COVID-19 lockdown benefitted the HKH in many ways. A schematic
354 shows the COVID-19 lockdown-induced effects in Figs. 6b-c: increased snow surface

355 reflectivity, reduced snowmelt and surface water runoff, as well as enhanced water content in
 356 the reservoir and snow.



362 **Figure 6:** (a) Change in water content (mm) of the Himalayan surface reservoirs (COVID
 363 minus CTL) from March to May 2020 over the Western Himalayas (WH), Central Himalayas
 364 (CH), Eastern Himalayas (EH) and Tibetan Plateau (TP). Vertical bars indicate the standard
 365 deviation within ten members of model simulations. Schematic illustrating the impacts of (b)
 366 air pollution on snow darkening in the Himalayas and surface water runoff for the usual
 367 polluted case and (c) the impacts of reduced pollution on snow brightening in the Himalayas
 368 and reduced surface water runoff, as observed during the 2020 COVID-19 lockdown period.

369

370 **5 Summary and conclusions:** A rising trend in Asian air pollution and associated climate
 371 change over the last few decades has had a detrimental impact on snow melting over the
 372 Hindu Kush Himalayas (HKH) and Tibetan Plateau region (Wester et al., 2019). Black
 373 carbon from increasing emissions of biomass burning, industrial and domestic combustion
 374 and transport is deposited on snow, reducing its albedo (i.e. darkening) (Bolch et al., 2019). A
 375 snow darkening effect along with pollution reduction, compounded with other climate change
 376 effects, accelerates the melting of snow and the disappearance of ice cover over the HKH and
 377 Tibetan Plateau region at an extraordinary rate (Usha et al., 2021). The drop in anthropogenic
 378 air pollution emissions, e.g. from energy production, during the COVID-19 lockdown period
 379 in spring 2020 reduced air pollutant levels worldwide (Forster et al., 2020). Our model
 380 simulations indicate that the associated reduction in anthropogenic aerosols and greenhouse
 381 gases in spring 2020 has benefited the HKH snow reservoirs. It caused an enhancement in the
 382 snow cover fraction by 6 - 12 % and snow mass by 2 - 20 %, corresponding to a decrease in
 383 snow melting by 10 - 40% and surface water runoff by 0.2 - 3 mm day⁻¹. As a consequence,
 384 the water content of the reservoir increased considerably by 4 to 59 %.

385 Our findings highlight that out of the two processes causing a retreat of Himalayan glaciers:
386 (1) a slow response to global climate change and (2) a fast response to local air pollution
387 (especially black carbon), a policy action on the latter is more likely to be within reach of
388 possible policy action on a shorter-term time scale and a more regional spatial scale. Even if
389 we stopped CO₂ emissions immediately, temperatures would not start decreasing. Our findings
390 confirm the importance of reducing short-lived climate forcers (black carbon) and their
391 complementary role to CO₂ mitigation (Rogelj et al., 2014). Reduction of air pollution to levels
392 similar with those recorded during the 2020 COVID-19 lockdown period, could safeguard
393 HKH glaciers, which are otherwise under the threat to disappear by the end of the 21st century.
394 Since 2000 Himalayan glaciers have been losing nearly half a meter of ice per year (Wester et
395 al., 2019). Our estimates indicate that air pollution reduction during COVID 19 lockdown in
396 spring 2020 caused a reduction in snow melt by 0.5 to 1.5 mm day⁻¹, indicating large benefits
397 to HKH glaciers. Even if global warming is kept below 1.5°C, one third of the glaciers in the
398 HKH region and more than half of those in the Eastern Himalaya will likely be lost by the end
399 of this century (Bolch et al., 2019). The speedily retreating glaciers and the snowpack loss are
400 already posing a threat to domestic sustainable water resources for billions of people in Asia
401 (Wood et al., 2021). However, if new economically and technically feasible policies would
402 reduce emissions of air pollutants (in particular black carbon) to at least lockdown period
403 levels, snowmelt could be reduced by 10 – 50%. Such policies will therefore bring substantial
404 benefits for sustained water supply, agriculture, and ecosystems in large parts of Asia.

405

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584

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593 **Author Contributions**

594 S.F. initiated the idea of the study. S.F. and B. H. performed model simulations. A.R. and A.
595 K. prepared Google based emission inventory. T.P.S., A.A., R.M. performed data analysis and
596 contributed in overall design. All authors contributed to discussions of the results and the
597 writing of the manuscript.

598 **Data and code availability**

599 The ECHAM-HAMMOZ model source code and all required input data are available to the
600 scientific community according to the HAMMOZ Software License Agreement through the
601 project website: <https://redmine.hammoz.ethz.ch/projects/hammoz>. The data that support the
602 findings of this study are openly available in zenodo at <http://doi.org/.../zenodo...>

603

604 **Competing Interests:** The authors declare no competing interests.