Validating the Nernst–Planck transport model under reaction-driven flow conditions using RetroPy v1.0

Po-Wei Huang[1], Bernd Flemisch[2], Chao-Zhong Qin[3], Martin O. Saar[1,4], and Anozie Ebigbo[5]

[1]Geothermal Energy and Geofluids Group, Institute of Geophysics, Department of Earth Sciences, ETH Zurich, Zurich, Switzerland
[2]Institute for Modelling Hydraulic and Environmental Systems, University of Stuttgart, Stuttgart, Germany
[3]State Key Laboratory of Coal Mine Disaster Dynamics and Control, Chongqing University, Chongqing, China
[4]Department of Earth and Environmental Sciences, University of Minnesota, Minneapolis, USA
[5]Chair of Hydromechanics, Helmut Schmidt University, Hamburg, Germany

Correspondence: Po-Wei Huang(powei.huang@erdw.ethz.ch), Anozie Ebigbo(ebigbo@hsu-hh.de)

Abstract. Reactive transport processes in natural environments often involve many ionic species. The diffusivities of ionic species vary. Since assigning different diffusivities in the advection-diffusion equation leads to charge imbalance, a single diffusivity is usually used for all species. In this work, we apply the Nernst–Planck equation, which resolves unequal diffusivities of the species in an electroneutral manner, to model reactive transport. To demonstrate the advantages of the Nernst–Planck model, we compare the simulation results of transport under reaction-driven flow conditions using the Nernst–Planck model with those of the commonly used single-diffusivity model. All simulations are also compared to well-defined experiments on the scale of centimeters. Our results show that the Nernst–Planck model is valid and particularly relevant for modeling reactive transport processes with an intricate interplay among diffusion, reaction, electromigration, and density-driven convection.

1 Introduction

Reactive transport processes are fundamental in large-scale subsurface applications such as: geological hydrogen storage (Aftab et al., 2022; Hassanpouryouzband et al., 2022), geothermal energy extraction (Randolph and Saar, 2011; Fleming et al., 2020; Ezekiel et al., 2022), biogeochemical subsurface treatments (Carrera et al., 2022), contaminant remediation (Frizon et al., 2003; Reddy and Cameselle, 2009), recovery of minerals (Chang et al., 2021), and carbon sequestration and storage (Luhmann et al., 2013; Tutolo et al., 2014; Pogge von Strandmann et al., 2019; Grimm Lima et al., 2020; Ling et al., 2021). On a smaller scale, reactive transport governs the processes of steel corrosion in reinforced concrete (Yu et al., 2020), cementation processes (Samson and Marchand, 2007; Cochepin et al., 2008), dendrite growth in microelectronics (Illés et al., 2022), and electrochemical reactors (Rivera et al., 2021). In all aforementioned settings, the reacting fluid consists of multiple ionic species and may be subject to an external electric field. In some applications, the electromigration of ionic species can play a significant role in the overall transport process. A typical model, which accounts for electromigration, is the Nernst–Planck equation.

Modeling electrochemical processes using the Nernst–Planck equation is usually coupled with the Poisson equation for the electrical potential, also known as the Poisson–Nernst–Planck (PNP) model. The PNP model has been applied in vari-

ous electrochemical applications (Jasielec, 2021). In geoscientific applications, benchmarks of the Nernst–Planck model have been proposed and verified considering charged clay surfaces (Tournassat and Steefel, 2021). Charged surfaces affect the water permeability of a porous material (in contrast to air permeability, which is not affected by electric charges) (Revil and Leroy, 2004; Kwon et al., 2004b, a; Cheng and Milsch, 2020). The PNP model is utilized at the pore scale to determine such water permeabilities (Priya et al., 2021). Pore-scale experiments with external charge effects have also been conducted to investigate reactive transport processes (Sprocati et al., 2019; Sprocati and Rolle, 2022; Rolle et al., 2022; Izumoto et al., 2022). Rasouli et al. (2015) verified numerical software using electromigration benchmarks (Lichtner, 1994; Glaus et al., 2013) and multicomponent-electromigration experiments considering dispersion effects (Rolle et al., 2013). Finally, multicomponent transport experiments in porous media confirm the validity of the Nernst–Planck model (Muniruzzaman et al., 2014; Muniruzzaman and Rolle, 2015, 2017; Rolle et al., 2018).

The above benchmark and validation studies focused on reactions and the mixing of fluids, subject to flow introduced by fluid pressure or electric fields. In this work, we propose another set of benchmark problems for numerical models, based on chemically driven convection experiments (Avnir and Kagan, 1984; Eckert and Grahn, 1999; Eckert et al., 2004), which focus on how reaction and mixing of fluids influence fluid flow. The concept of reaction-driven flow is as follows: Consider reaction $A + B \rightarrow C$, where A and B are the reactants and C is the product. If A, B, and C have different material properties, for example, viscosity, diffusivity, and/or density, the reaction can induce fluid flow and hydrodynamic instabilities in such systems (De Wit, 2016, 2020). We are interested in aqueous systems, where ionic species exist, such that the Nernst–Planck model is applicable.

Zalts et al. (2008) performed experiments using an acid and a base to study chemically driven convection. Related experiments and theories of the underlying mechanism have emerged: unequal species diffusivities (Citri et al., 1990; Almarcha et al., 2010, 2011; Trevelyan et al., 2011; Lemaigre et al., 2013; Kim, 2019; Jotkar et al., 2021), the same species diffusivity but with varying viscosity (Hejazi and Azaiez, 2013), concentration-dependent species diffusivities (Bratsun et al., 2015, 2017, 2021, 2022). Chemically driven convection has also been investigated using a weak acid and a weak base (Cherezov et al., 2018). Since the Nernst–Planck model is expected to be able to account for unequal species diffusivities and has the characteristics of concentration-dependent diffusive fluxes, we examine the validity of the Nernst–Planck model by comparing our Nernst–Planck simulations with chemically driven experiments, conducted by others.

In addition, we present simulations of convective dissolution of CO_2, a process which is important in the underground storage of CO_2 (Singh et al., 2019). Thomas et al. (2016) performed experiments of convective dissolution of CO_2 into alkaline brines, including $LiOH$ and $NaOH$, and observed that different cations lead to varying onset times and fingering patterns. Thomas et al. (2016) proposed that the inert cation, which does not participate in the acid-base reaction between the dissolved CO_2 and the brine, plays a major role in the development of convective instabilities. Since dissolved CO_2 forms charged and uncharged states ($CO_2(aq)$, HCO_3^-, CO_3^{2-}) when the pH changes as a consequence of neutralization reactions, it is natural to apply the Nernst–Planck model in such cases.

In the following sections, we first establish the fundamental equations of the conservation laws and the accompanying numerical methods. We then provide details concerning the two reaction-driven flow experiments and a CO_2 dissolution experiment—all conducted by others. We compare the differences between the simulations of the single-diffusivity transport

model and the Nernst–Planck model, and elucidate the necessity of considering electromigration effects caused by differing diffusivities during reactive transport modeling.

2 Materials and Methods

2.1 Fundamental equations for reactive transport processes

In this section, we delineate the fundamental equations describing the reactive transport of aqueous species.

2.1.1 Mass balance of the fluid components

We consider a single-phase fluid composed of N components, which can be ionic species, dissolved gas, or minerals. The single-phase fluid occupies a certain volume V_f, and the ith component of the fluid has the mass m_i. The mass concentration of the ith component is defined as

$$\rho_i = \frac{m_i}{V_f}. \tag{1}$$

We use the term mass concentration to clarify that ρ_i is not the density of the component. The fluid density, ρ, can be obtained by summing up the mass concentrations of the components, i.e.,

$$\rho = \sum_{i=1}^{N} \rho_i = \sum_{i=1}^{N} \frac{m_i}{V_f}. \tag{2}$$

In general, the fluid density is a function of fluid pressure, p, temperature, T, and fluid composition (i.e., the mass concentrations, ρ_i). The equation of state of the fluid density can be written as

$$\rho = \rho(p, T, \rho_i; \ i = 1, 2, ..., N). \tag{3}$$

The mass balance of each fluid component is described by the continuity equation,

$$\frac{\partial \rho_i}{\partial t} + \nabla \cdot (\rho_i \boldsymbol{v}_i) = Q_i, \tag{4}$$

where $\rho_i \boldsymbol{v}_i$ is the mass flux and Q_i is the source/sink term introduced by chemical reactions. Since chemical reactions preserve fluid mass, we have $\sum_i Q_i = 0$. We express the mass flux of the ith component as

$$\rho_i \boldsymbol{v}_i = \rho_i \boldsymbol{u} + \rho_i \boldsymbol{u}_{EP,i} + \boldsymbol{J}_{diff,i}, \tag{5}$$

where $\boldsymbol{u}$ is the barycentric velocity, $\boldsymbol{u}_{EP,i}$ is the electrophoretic velocity, and $\boldsymbol{J}_{diff,i}$ is the diffusive flux. In the following sections, we define the models of the diffusive flux, the electrophoretic flux, and the barycentric flux.

2.1.2 Modeling of the diffusive flux

We use standard Fickian diffusion to model the diffusive flux,

$$\boldsymbol{J}_{\text{diff},i} = -D_i M_i \nabla C_i \quad i = 1, 2, ..., N-1, \tag{6}$$

where D is the diffusivity, M is the molar mass, and C is the molar concentration. The contribution of the diffusing solvent can be neglected, considering $\rho_i/\rho_N \ll 1$. Such an assumption leads to a particular choice of the Nth component, which contributes the most mass to the fluid. In our experiments of interest, the Nth component is $H_2O(l)$. Fickian diffusion utilizes the diagonal-diffusivity-matrix assumption, which is valid for many electrolytes (Dreyer et al., 2013). A diagonal diffusivity matrix means that short-range interactions between ions are negligible, which restricts our model to dilute solutions (Kontturi et al., 2008).

2.1.3 Modeling of the barycentric flux

The type of experiments of interest for this work is performed in a Hele-Shaw cell, for which fluid flow can be described by Darcy's law,

$$\boldsymbol{u} = -\frac{k}{\eta} \left(\nabla p - \rho \boldsymbol{g} \right), \tag{7}$$

where k is the permeability, η is the dynamic viscosity of the fluid, and $\boldsymbol{g}$ is the gravitational acceleration. When combined with the incompressibility assumption,

$$\nabla \cdot (\rho \boldsymbol{u}) = 0, \tag{8}$$

$\boldsymbol{u}$ and p can be obtained. The scalar permeability of a Hele-Shaw cell, with a gap width of w, is given by

$$k = \frac{w^2}{12}. \tag{9}$$

2.1.4 Modeling of the electrophoretic flux

When the components in the fluids are electrically charged, for example, the ions Na^+ or Cl^-, we have to consider the effect of Coulombic forces among the ions. Such Coulombic forces are typically modeled by incorporating the electric field $\boldsymbol{E}$ into the mass fluxes of the fluid components (Boudreau et al., 2004). The electrophoretic flux is described by

$$\rho_i \boldsymbol{u}_{\text{EP},i} = \frac{z_i \rho_i D_i F}{RT} \boldsymbol{E}, \tag{10}$$

where z_i is the charge number of the fluid component, F is the Faraday constant, R is the ideal gas constant, and T is the fluid temperature. We require the charge conservation equation to define the electric field in a multicomponent fluid. To obtain the charge conservation equation of the ionic species, it is useful to use the flux defined in molar concentration. Using $\rho_i = M_i C_i$, we have the mass conservation equation (without considering the source/sink terms introduced by reactions)

$$\frac{\partial C_i}{\partial t} = -\nabla \cdot (C_i \boldsymbol{u} + C_i \boldsymbol{u}_{\text{EP},i} + \boldsymbol{J}_{\text{diff},i}/M_i). \tag{11}$$

Following the derivations of Kirby (2010), the charge density is defined as

$$\rho_{\mathrm{E}} = F \sum_{i=1}^{N-1} z_i C_i. \tag{12}$$

The charge conservation equation is the sum of Eq. (11) over all ionic species multiplied by the charge number and the Faraday constant,

$$\frac{\partial \rho_{\mathrm{E}}}{\partial t} = -F \, \nabla \cdot \left(\sum_{i=1}^{N-1} z_i C_i \boldsymbol{u} + z_i C_i \boldsymbol{u}_{\mathrm{EP},i} + z_i \boldsymbol{J}_{\mathrm{diff},i}/M_i \right). \tag{13}$$

We assume the aqueous solution is locally electroneutral,

$$\sum_{i=1}^{N-1} z_i C_i = 0, \tag{14}$$

which is an approximation considering the Debye length (typically on the order of nanometers) vanishes at the length scale of the considered system (Dickinson et al., 2011). The electroneutrality assumption eliminates the barycentric flux term in the charge conservation equation. We consider Fickian diffusive flux, Eq. (6), and the electrophoretic flux, Eq. (10), so that the charge conservation is given by

$$\frac{\partial \rho_{\mathrm{E}}}{\partial t} = -\nabla \cdot \left(\sum_{i=1}^{N-1} \frac{D_i C_i (z_i F)^2}{RT} \boldsymbol{E} - \sum_{i=1}^{N-1} D_i z_i F \nabla C_i \right). \tag{15}$$

The multiplier of the electric field, $\boldsymbol{E}$, is the electrical conductivity ($\mathrm{C^2 m^{-3} kg^{-1} s}$),

$$\sigma = \sum_{i=1}^{N-1} \frac{D_i C_i (z_i F)^2}{RT}. \tag{16}$$

When the diffusivities of all species are equal, $D_i = D$, the diffusive flux in the charge conservation equation is

$$D \sum_{i=1}^{N-1} z_i F \nabla C_i = D \nabla \left(F \sum_{i=1}^{N-1} z_i C_i \right) = 0. \tag{17}$$

Then, the charge conservation equation yields Laplace's equation of the electric potential, i.e., $\nabla \cdot (\sigma \nabla \phi) = 0$, where $\boldsymbol{E} = -\nabla \phi$. However, the diffusivity of the chemical species varies with molecular size and can differ by up to an order of magnitude. Such variations in molecular diffusivity lead to diffusive fluxes that produce an electric field. The model that combines mass conservation, Eq. (11), and Poisson's equation of electrostatics, is known as the Poisson–Nernst–Planck (PNP) model (Pamukcu, 2009).

The PNP model aims at resolving both the electric potential and the molar concentrations, subject to the boundary conditions of the electric potential. The mathematical properties of the PNP model, with no external flow (Filipek et al., 2017), coupled with Darcy flow (Herz et al., 2012; Ignatova and Shu, 2022), and coupled with Navier–Stokes flow (Lee, 2021; Constantin et al., 2021, 2022) are still being explored. Rigorous upscaling and homogenization of the PNP–Stokes system has been performed

(Ray et al., 2012a; Kovtunenko and Zubkova, 2021), and the upscaled formulations are examined by numerical modeling (Frank et al., 2011; Ray et al., 2012b). Numerical methods and solution techniques of the PNP equations are an ongoing research field (Flavell et al., 2014; Song et al., 2018; Shen and Xu, 2021; Liu and Maimaitiyiming, 2021; Yan et al., 2021; Liu et al., 2022; Zhang et al., 2022). Models similar to the PNP model also arise in modeling multicomponent interdiffusion of solids (Bożek et al., 2019), where the experimental comparison and the development of numerical methods are studied by Sapa et al. (2020) and Bożek et al. (2021).

In this work, we enforce the null current condition

$$\sum_{i=1}^{N-1} \frac{D_i C_i (z_i F)^2}{RT} \boldsymbol{E} - \sum_{i=1}^{N-1} D_i z_i F \nabla C_i = 0 \tag{18}$$

to ensure that no charge accumulates at the continuum scale of interest,

$$\frac{\partial \rho_{\mathrm{E}}}{\partial t} = 0 \,. \tag{19}$$

This result is a simplification of the PNP model. The null current condition has been utilized in modeling multicomponent ionic transport (Lichtner, 1985; Cappellen and Gaillard, 1996; Giambalvo et al., 2002; Muniruzzaman and Rolle, 2019; Cogorno et al., 2022; López-Vizcaíno et al., 2022). The combined assumptions of local electroneutrality and null current are referred to as strict electroneutrality by Lees et al. (2017). Furthermore, if there are no sources of the electric field (Tabrizinejadas et al., 2021), then the electric field can be represented by

$$\boldsymbol{E} = \frac{RT}{F} \frac{\sum_{j=1}^{N-1} D_j z_j \nabla C_j}{\sum_{k=1}^{N-1} (z_k)^2 D_k C_k} \,, \tag{20}$$

where we use j and k as summation indices. Hence, the molar electrophoretic flux of individual species can be formulated as

$$C_i \boldsymbol{u}_{\mathrm{EP},i} = \frac{z_i D_i C_i F}{RT} \boldsymbol{E} = z_i D_i C_i \frac{\sum_{j=1}^{N-1} z_j D_j \nabla C_j}{\sum_{k=1}^{N-1} (z_k)^2 D_k C_k} \,. \tag{21}$$

There are many physical interpretations of such a flux. For example, one is that the dissolved species is advecting due to the electric field. Another point of view can be obtained by combining Eq. (21) with the Fickian diffusion flux,

$$\boldsymbol{J}_{\mathrm{diff},i}/M_i + C_i \boldsymbol{u}_{\mathrm{EP},i} = -D_i \nabla C_i + \frac{z_i D_i C_i}{\sum_{k=1}^{N-1} (z_k)^2 D_k C_k} \sum_{j=1}^{N-1} z_j D_j \nabla C_j \tag{22}$$

$$= -\sum_{j=1}^{N-1} D_{ij} \nabla C_j \,, \tag{23}$$

which corresponds to the "natural generalization" of Fickian diffusion, denoted by Onsager (1945). Following Boudreau et al. (2004), the cross-coupling diffusivities are given by

$$D_{ij} = \delta_{ij} D_i - \frac{z_i z_j D_i D_j C_i}{\sum_{k=1}^{N-1} (z_k)^2 D_k C_k} = \delta_{ij} D_i - \frac{z_i D_i C_i \otimes z_j D_j}{\sum_{k=1}^{N-1} (z_k)^2 D_k C_k} \,, \tag{24}$$

where Einstein's summation convention is not applied. Hence, locally electroneutral diffusion is equivalent to nonlinear multicomponent diffusion with a diffusivity tensor being a rational function of molar concentrations of the charged species (Rubinstein, 1990). Similar forms of the cross-coupling diffusivities can also be found in Lichtner (1985).

Combining Eq. (11) and Eq. (22) yields

$$\frac{\partial C_i}{\partial t} = -\nabla \cdot \left(-D_i \nabla C_i + z_i D_i C_i \frac{\sum_{j=1}^{N-1} z_j D_j \nabla C_j}{\sum_{k=1}^{N-1} (z_k)^2 D_k C_k} + \boldsymbol{u} C_i \right), \tag{25}$$

which is the Nernst–Planck equation for single-phase-multicomponent systems. If all species have the same diffusivity, $D_i = D$, and the aqueous solution is electroneutral, Eq. (14). Then, the single-phase-multicomponent Nernst–Planck equation reduces to

$$\frac{\partial C_i}{\partial t} = -\nabla \cdot (-D \nabla C_i + \boldsymbol{u} C_i), \tag{26}$$

which we refer to as the single-diffusivity model. Both the Nernst–Planck model and the single diffusivity model satisfy the electroneutrality condition, Eq. (14), and the zero-charge accumulation assumptions, Eq. (19). However, of the two models, only the Nernst–Planck model is capable of handling the physics of varying species diffusivities. To communicate the effectiveness of the Nernst–Planck equation in reactive transport scenarios, we compare the modeling results with the single diffusivity model, which is a common approach. We set the diffusivity of the single diffusivity model as $3 \times 10^{-3}\,\mathrm{mm^2\,s^{-1}}$, which is within the common values for species diffusivities in water at $25\,^\circ\mathrm{C}$.

2.1.5 Chemical equilibrium using Gibbs-energy minimization

Another aspect of mass balance amongst the fluid components, which we now denote as the chemical species, is the reaction that changes the amount of the chemical species while conserving the mass balance of chemical elements. The reaction source/sink term in the mass balance of each fluid component, Q_i, is implicitly formulated as the Gibbs-energy minimization problem,

$$\min_n G = \sum_{i=1}^{N} \mathrm{n}_i \mu_i \quad \text{subject to} \quad \begin{cases} \text{mass balance of chemical elements} \\ \mathrm{n} \geq 0 \end{cases}, \tag{27}$$

where G is the Gibbs energy and n_i is the molar amount of the ith chemical species. The chemical potential is defined as

$$\mu_i = \mu_i^\circ + RT \log a_i, \tag{28}$$

where μ_i° is the standard chemical potential of the ith species, and a_i is the activity of the ith species. For the standard chemical potential, we employ the SUPCRT07 database (Johnson et al., 1992). To evaluate the activities, we use the Helgeson-Kirkham-Flowers (HKF) activity model (Helgeson and Kirkham, 1974a, b, 1976; Helgeson et al., 1981), which is based on the extended Debye–Hückel activity model. For the activity model of $CO_2(\mathrm{aq})$, we use the Drummond model (Drummond, 1981). The combination of the SUPCRT database and the HKF activity model is utilized in many geochemical modeling packages,

such as EQ3/6 (Wolery, 1992), CHILLER (Reed, 1998), ChemApp (Petersen and Hack, 2007), Perplex (Connolly, 2009), GEM-Selektor (Kulik et al., 2012), PHREEQC (Parkhurst and Appelo, 2013), Geochemist's Workbench (Bethke, 2007), and Reaktoro (Leal, 2022) (Miron et al., 2019). General reviews of the activity models can be found in Marini (2007); Hingerl (2012).

190   We use Reaktoro to solve the chemical equilibrium problem. The numerical methods and implementations of the Gibbs-energy minimization problem are detailed in Leal et al. (2014, 2016, 2017). The equilibrium calculations are performed at every time step and for every discretized volume of the simulation domain, and calculations of similar chemical states can be cached for efficiency. Efficient lookup tables and machine learning approaches of chemical-equilibrium calculations are in active research (Huang et al., 2018; De Lucia and Kühn, 2021; De Lucia et al., 2021; Savino et al., 2022; Laloy and Jacques, 195   2022), and significant speedups in solving reactive transport problems have been achieved (Kyas et al., 2022; Bordeaux-Rego et al., 2022). In particular, Reaktoro employs on-demand machine learning and physics-based interpolation, which is able to conserve mass and control interpolation accuracy (Leal et al., 2020). Surrogate models of the full reactive transport process, using machine learning, have also been developed (Sprocati and Rolle, 2021). However, to avoid approximation errors, machine learning approaches and surrogate models are not employed in this work.

2.2   Numerical schemes

This section elaborates on the numerical schemes we use for solving the fundamental equations of the reactive transport processes.

2.2.1   The barycentric flux

We utilize the mixed finite element formulation to solve for fluid velocity and fluid pressure under the constraints of mass 205   balance, Eq. (8) and momentum balance, Eq. (7). To avoid spurious solutions, the finite element spaces have to satisfy the Ladyzhenskaya-Babuška-Brezzi (LBB) condition. We refer the reader to other works for further discussions regarding the LBB condition in the context of fluid flow problems (Donea and Huerta, 2003), geodynamics (Thieulot and Bangerth, 2022), and elasticity and electromagnetic applications (Boffi et al., 2013). We select LBB-compatible elements, namely the zeroth-order Raviart-Thomas (RT) element for the barycentric flux and the piecewise constant element for the pressure (Raviart and 210   Thomas, 1977). Another popular LBB-compatible element is the Brezzi-Douglas-Marini element (Brezzi et al., 1985). The mixed finite element formulation can be written in matrix form as

$$
\begin{pmatrix} A & B^T \\ B & 0 \end{pmatrix} \begin{bmatrix} \boldsymbol{u} \\ p \end{bmatrix} = \begin{bmatrix} f \\ 0 \end{bmatrix},
\tag{29}
$$

which leads to a saddle point problem. Following Benzi et al. (2005)'s review of saddle point problems, one can apply Uzawa's method (Uzawa, 1958) to solve such problems. Given an initial condition of p^n at step n, Uzawa's method solves $\boldsymbol{u}$ and p

iteratively,

$$\boldsymbol{u}^{n+1} = A^{-1}(f - B^T p^n), \tag{30}$$

$$p^{n+1} = p^n + \omega B \boldsymbol{u}^{n+1}, \tag{31}$$

where $\omega > 0$ is a relaxation parameter. Normally, Uzawa's method converges slowly and requires many iterations, so that improving Uzawa's method is an ongoing research effort (Bacuta, 2006). We use one of the improved methods, the augmented Lagrangian Uzawa's method, introduced by Fortin and Glowinski (1983):

$$\begin{pmatrix} A + rB^T B & B^T \\ B & 0 \end{pmatrix} \begin{bmatrix} \boldsymbol{u} \\ p \end{bmatrix} = \begin{bmatrix} f \\ 0 \end{bmatrix}, \tag{32}$$

where $r > 0$ is a tuning parameter. We solve the matrix problem iteratively,

$$\boldsymbol{u}^{n+1} = (A + rB^T B)^{-1}(f - B^T p^n), \tag{33}$$

$$p^{n+1} = p^n + \omega B \boldsymbol{u}^{n+1}. \tag{34}$$

When $\omega = r$, we can choose sufficiently large r to accelerate the convergence of Uzawa's method. However, the matrix $A + rB^T B$ becomes more ill-conditioned as r increases, which may lead to a large number of iterations to solve when applying an iterative method (Fortin and Glowinski, 1983). Hence, we use the MUMPS (Amestoy et al., 2001, 2019) direct solver to update the velocity, Eq. (33). The degrees of freedom of $\boldsymbol{u}$ in our simulations range from $17\,000$ to $140\,000$.

2.2.2 Transport of fluid components

In the previous section, we selected an LBB-compatible velocity and pressure pair. The basis function of pressure is a piecewise constant, which can be described as the average pressure of a certain cell volume in a given mesh. Therefore, it is straightforward to define the molar concentrations as piecewise constants. We use the finite volume method to discretize the transport of the fluid components, Eq. (25):

$$\int_{\Omega} q_i \frac{\partial C_i}{\partial t} d\Omega = -\int_{\Gamma} q_i \left(-D_i \nabla C_i + z_i D_i C_i \frac{\sum_{j=1}^{N-1} z_j D_j \nabla C_j}{\sum_{k=1}^{N-1} (z_k)^2 D_k C_k} + \boldsymbol{u} C_i \right) \cdot \boldsymbol{n} \, d\Gamma, \tag{35}$$

where the fluxes are yet to be defined. Considering no-flux boundary conditions of the fluid components, we use the two-point flux approximation scheme (TPFA) for the diffusive flux,

$$\int_{\Gamma} q_i D_i \nabla C_i \cdot \boldsymbol{n} \, d\Gamma = -\int_{\Gamma_{\text{int}}} [q_i] D_i \frac{[C_i]}{h} \, d\Gamma, \tag{36}$$

where Γ_{int} denotes the interior cell interfaces, $[\bullet]$ is the jump operator that evaluates the difference of a function across a common interface of two cells, and h is the distance between the cell centers. We assumed that the diffusivity of each species is constant. For diffusivities varying in space, Tournassat et al. (2020) discussed and compared various averaging approaches

of the diffusivities at the interfaces between cells. For the advective flux, we utilize the full upwinding method,

$$-\int_{\Gamma} q_i C_i \, \boldsymbol{u} \cdot \boldsymbol{n} \, d\Gamma = -\int_{\Gamma_{\text{int}}} [q_i] C_i^{\text{up}} \, \boldsymbol{u} \cdot \boldsymbol{n} \, d\Gamma, \tag{37}$$

where C_i^{up} is the molar concentration in the upwind direction. For the Nernst–Planck term, we follow Tournassat et al. (2020)'s approach:

$$-\int_{\Gamma} q_i z_i D_i C_i \frac{\sum_{j=1}^{N-1} z_j D_j \nabla C_j}{\sum_{k=1}^{N-1} (z_k)^2 D_k C_k} \cdot \boldsymbol{n} \, d\Gamma = \int_{\Gamma_{\text{int}}} [q_i] z_i D_i \{C_i\} \frac{\sum_{j=1}^{N-1} z_j D_j [C_j]}{\sum_{k=1}^{N-1} (z_k)^2 D_k \{C_k\} h} \, d\Gamma, \tag{38}$$

where $\{\bullet\}$ is the averaging operator. We use the arithmetic average to represent the molar concentration at the cell interface, which is suitable when all charged components are mobile (Gimmi and Alt-Epping, 2018), or when there are no membrane effects or large concentration gradients (Tournassat et al., 2020). UFL (Alnæs, 2012) is utilized for defining the variational forms in our code implementation.

With regard to the time-stepping schemes, we use an explicit scheme for the upwind advection term and the Crank–Nicolson scheme for the diffusion and Nernst–Planck terms. The timestep size Δt is determined in an adaptive manner, where the minimum and maximum timestep sizes are prescribed for each simulation. To avoid negative concentration values, we perform a transformation of variables $C_i = \exp(c_i)$, which leads to the logarithmically transformed Nernst–Planck equations (Kirby, 2010). The transport equations are assembled using FEniCS (Alnæs et al., 2015) and solved using PETSc (Balay et al., 2022).

2.2.3 Coupling flow, transport, and chemical equilibrium

Of many coupling approaches in reactive transport modeling (Steefel and MacQuarrie, 2018; Carrayrou et al., 2004; Abd and Abushaikha, 2021), we apply the sequential non-iterative approach (SNIA) to couple flow, transport, and chemical equilibrium. Figure 1 illustrates the simulation procedure, where the fluid flow, the species transport, and the equilibrium steps are solved consecutively.

Due to the simplicity and effectiveness in applying the SNIA coupling between the transport solvers and the chemical equilibrium codes, this coupling method is widely adopted in the following software: OpenGeoSys (Shao et al., 2009; Naumov et al., 2022), poreReact (coupling of OpenFOAM (Weller et al., 1998) and Reaktoro) (Oliveira et al., 2019), CSMP++GEM (Yapparova et al., 2019), FEniCS–Reaktoro (Damiani et al., 2020), Osures (Moortgat et al., 2020), FEniCS-based Hydro-Mechanical-Chemical solver (Kadeethum et al., 2021), PorousFlow (based on the MOOSE Framework (Permann et al., 2020)) (Wilkins et al., 2021), IC-FERST-REACT (Yekta et al., 2021), COMSOL and PHREEQC (Jyoti and Haese, 2021), GeoChemFoam (coupling of OpenFOAM and PHREEQC) (Maes and Menke, 2021), coupling of Reaktoro and Firedrake (Rathgeber et al., 2016) (Kyas et al., 2022), and P3D-BRNS (Golparvar et al., 2022). For reviews of reactive transport codes and the underlying coupling approaches, we refer the reader to the publications by Gamazo et al. (2015); Damiani et al. (2020).
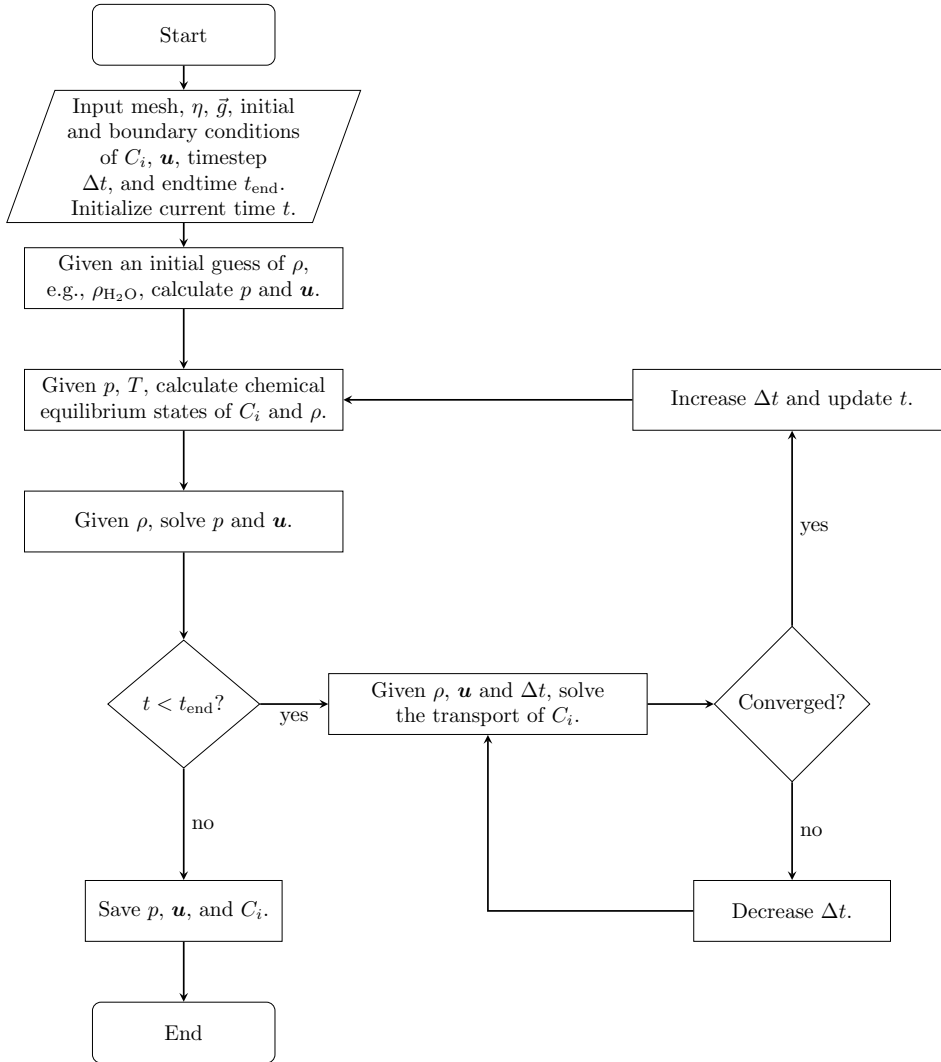
Figure 1. Flow chart of the simulation procedures that couple flow, transport, and chemical equilibrium, using the sequential non-iterative approach (SNIA).

2.3 Selected experiments for model evaluation

270   We select three experiments, performed (by others) in a Hele-Shaw cell, to evaluate the effectiveness of the Nernst–Planck model in reactive transport situations. The diffusivities of the aqueous species are listed in Table 1. The density and dynamic viscosity of the relevant aqueous solutions are listed in Table 2. Our models assume constant dynamic viscosity, set as the average viscosity of the aqueous solutions. All experiments are performed in a lab environment, and we assume a constant temperature of $25\,°C$ and a constant background pressure of 100 kPa. No-flow boundary conditions are prescribed on all

off
275   sides of the simulation domain. Table 3 specifies the chemical compositions of the aqueous solutions, which defines the initial conditions of the mass transport equations.

Table 1. Diffusivities of aqueous species in $25\,°C$ water at infinite dilution

	H^+	OH^-	Na^+	Cl^-	Li^+	NO_3^-	HCO_3^-	CO_3^{2-}	$CO_2(aq)$
$D\,(10^{-3}\,\mathrm{mm^2\,s^{-1}})$	9.311	5.273	1.334	2.032	1.029	1.902	1.101	0.804	2.045

The diffusivities of aqueous species other than HCO_3^-, CO_3^{2-}, and $CO_2(aq)$ are taken from Vanýsek (2021). The diffusivities of HCO_3^-, CO_3^{2-}, and $CO_2(aq)$ are calculated using Zeebe (2011).

Table 2. Dynamic viscosity and density of aqueous solutions and water at $25\,°C$

		1M HCl	1M NaOH	1.4M NaOH	1.5M HNO₃	0.01 M LiOH	H2O(l)
Literature	$\eta\,(\mathrm{mPa\,s})$	0.959[a]	1.129[b]	1.232[b]	0.933[c]	0.893[b]	0.89
Literature	$\rho\,(\mathrm{g/cm^3})$	1.014[d]	1.039[b]	1.055[b]	1.046[c]	0.997[b]	0.997
Reaktoro	$\rho\,(\mathrm{g/cm^3})$	1.012	1.037	1.053	1.044	0.997	0.997

Linearly interpolated using [a]Nishikata et al. (1981), [b]Sipos et al. (2000), [c]Zaytsev and Aseyev (1992), and [d]Åkerlöf and Teare (1938). The viscosity and density of water are taken from Lemmon and Harvey (2021).

Table 3. The chemical compositions of the aqueous solutions. The unit of the chemical compositions is in $\mathrm{mol\,L^{-1}}$. In the table, dashes denote "not applicable" because the chemical species are not considered in the simulation cases.

	H^+	OH^-	Na^+	Cl^-	Li^+	NO_3^-	HCO_3^-	CO_3^{2-}	$CO_2(aq)$	H2O(l)
1M HCl	1.0	2.109×10^{-14}	10^{-20}	1.0	-	-	-	-	-	54.170
1M NaOH	2.199×10^{-14}	1.0	1.0	10^{-20}	-	-	-	-	-	55.360
1.4M NaOH	1.537×10^{-14}	1.4	1.4	-	-	10^{-20}	-	-	-	55.361
1.5M HNO₃	1.5	1.287×10^{-14}	10^{-20}	-	-	1.5	-	-	-	52.712
0.01 M LiOH	1.235×10^{-12}	0.01	-	-	0.01	-	10^{-20}	10^{-20}	10^{-20}	55.345

2.3.1   Chemically driven convection of acid–base systems

Almarcha et al. (2010) experimentally showed that simple acid–base reactions, such as $A + B \rightarrow C$, can lead to hydrodynamic instabilities due to differences in diffusivities and densities of the chemical species. For the first case in our numerical modeling

280   study, we follow the experimental setup described in Almarcha et al. (2010). The acid (A) is $1\,M$ HCl and the base (B) is $1\,M$ NaOH, where M is the molar concentration $(\mathrm{mol\,L^{-1}})$. The experiment was performed by Almarcha et al. (2010) in a Hele-Shaw cell, which was $3.1\,cm$ wide, $5\,cm$ high, and had a $0.5\,mm$ gap width. To prevent Rayleigh-Taylor instability, the

less-dense acid was placed on top of the base. We consider the following species as the main fluid components in this case: H^+, OH^-, Na^+, Cl^-, and $H_2O(l)$. Please refer to the left panel of Figure 2 for the setup. Table 3 lists the chemical composition of 1 M HCl and 1 M NaOH. The prescribed minimum timestep size is 3×10^{-2} seconds, and the maximum timestep size is 2.0 seconds.

For the second modeling case, we choose a similar experiment, but with a strikingly different instability process. In their chemically driven convection experiments, Bratsun et al. (2017) reported a shock-wave-like structure of the acid–base interface. In further studies, this phenomenon has subsequently been reproduced using many different acid–base pairs (Mizev et al., 2021; Bratsun et al., 2021). For our numerical comparison study, we use the acid-base pair from Bratsun et al. (2021), i.e., HNO_3– NaOH. The acid consists of 1.5 M HNO_3 and is placed on top of 1.4 M NaOH. The Hele-Shaw cell of this experiment has a width of 2.5 cm, a height of 9.0 cm, and a gap width of 1.2 mm. We consider the following species as the main fluid components in this case: H^+, OH^-, Na^+, NO_3^-, and $H_2O(l)$. The setup is shown in Figure 2, and Table 3 lists the initial chemical composition of 1.5 M HNO_3 and 1.4 M NaOH. The prescribed minimum timestep size is 1×10^{-3} seconds, and the maximum timestep size is 1.0 seconds. We perform the simulation in a 2D geometry with the same width, height, and species initial conditions as in the experiment, and compare the simulation results with the time series of interferometry snapshots taken during the experiments.

2.3.2 Convective dissolution of CO_2 in reactive alkaline solutions

Thomas et al. (2016) performed experiments of convective dissolution of CO_2 in reactive alkaline solutions. The experiment setup consisted of a Hele-Shaw cell, saturated with alkaline solutions, and gaseous CO_2 flows over the top of the Hele-Shaw cell. Eventually, gaseous CO_2 dissolves into the alkaline solution, leading to density changes in the solution and, as a result, the formation of convection cells. Furthermore, Thomas et al. (2018) studied the convective dissolution of CO_2 in various molar concentrations of NaCl solutions using the aforementioned experimental setup. Likewise, Mahmoodpour et al. (2019) studied the effect of various salt solutions on the onset of convective instabilities by conducting experiments within porous media. Similar experimental studies in a sand box have been performed using dense miscible fluids (Neufeld et al., 2010; Agartan et al., 2015; Guo et al., 2021; Wang et al., 2021; Tsinober et al., 2022). Although such experiments do capture the effects of density-driven flow, electromigration effects are not involved. Other experimental studies of convective dissolution of gaseous CO_2 in aqueous solutions have been conducted in pressure-volume-temperature (PVT) cells (Yang and Gu, 2006), Hele-Shaw cells (Kneafsey and Pruess, 2010; Class et al., 2020; Zhang et al., 2020; Jiang et al., 2020; Teng et al., 2021), and 3D granular porous media (Mahmoodpour et al., 2020; Brouzet et al., 2022). Amarasinghe et al. (2020) reviewed CO_2 dissolution experiments and performed experiments of convective dissolution of CO_2 in a Hele-Shaw cell and porous media under supercritical CO_2 conditions.

Numerical studies of CO_2 convective dissolution have been performed considering the reactive transport of: a single component (Azin et al., 2013), a single component with mineral reactions (Babaei and Islam, 2018; Shafabakhsh et al., 2021), a single component in two fluid phases with a resolved gas–fluid boundary (Martinez and Hesse, 2016), multiple components with mineral reactions (Zhang et al., 2011), and multiple components in two fluid phases with mineral reactions (Audigane

et al., 2007; Soltanian et al., 2019; Sin and Corvisier, 2019). However, the Nernst–Planck model was not invoked in any of these works.

Using Lattice-Boltzmann methods with consideration of electrostatic forces, Fu et al. (2020, 2021) investigated how ionic species of various salt solutions with varying concentrations affect the onset time and mass transfer rate of CO_2 convective dissolution. So far, these are the only works we found that account for electromigration effects when modeling the convective dissolution of CO_2.

In the third case of our numerical study, we model the experiment of Thomas et al. (2016) and investigate the reactive transport of CO_2 in a Hele-Shaw cell including the electromigration of species using the Nernst–Planck model. Instead of comparing our results to the experiments at the full scale (165 mm by 210 mm) (Thomas et al., 2016), a simulation domain with a width of 100 mm, a height of 40 mm, and a gap width of 0.5 mm is utilized, see the right panel of Figure 2. The smaller simulation domain was chosen, as the relevant processes remained within this smaller domain. On the top boundary of the simulation domain, we assume that the gaseous CO_2 and the alkaline brine are always in equilibrium. The alkaline brine is 0.01 M LiOH. We consider the following species as the main fluid components in the case of convective dissolution of CO_2: H^+, OH^-, Li^+, $CO_2(aq)$, HCO_3^-, CO_3^{2-}, and $H_2O(l)$, and the initial conditions are presented in Table 3. The prescribed minimum and maximum timestep sizes are 5×10^{-3} seconds and 1.5 seconds, respectively. In the results section, we present the development of the fingers and the movement of the "shock wave" induced by density, reaction, and electromigration effects.

3 Results

This section compares the simulations and the experimental results of the chemically driven convection problems and the convective dissolution of CO_2 into a reactive alkaline solution. All simulations are performed using the Nernst–Planck and the single-diffusivity models.

3.1 Comparing the chemically driven convection experiments with the single-diffusivity model and the Nernst–Planck model

We present the results of the chemically driven convection between 1 M HCl and 1 M NaOH solutions. Figure 3 compares the digital interferometry images obtained from experiments with the simulated fluid densities. The interferometry images have a size of 13 mm by 23.075 mm, from which we select a sub-region of the same size from the simulations. This region is indicated by the red box in the right panel of Figure 3. Note that the interferometry images show the density contrast of the fluids (induced by changes in refractive index), not the density itself. Hence, the simulated density plots cannot be compared with the interferometry images pixel by pixel. Nonetheless, in the left panel of Figure 3, only the simulations with the Nernst–Planck model replicate the experimentally observed fingering effect with a similar number of fingers. At the simulation time of 60 seconds, a low-density layer (shown in yellow) emerges in the Nernst–Planck model, an effect which cannot be observed in the single-diffusivity model. At 115 seconds, the low-density layer develops instabilities, causing density-driven flow. However,
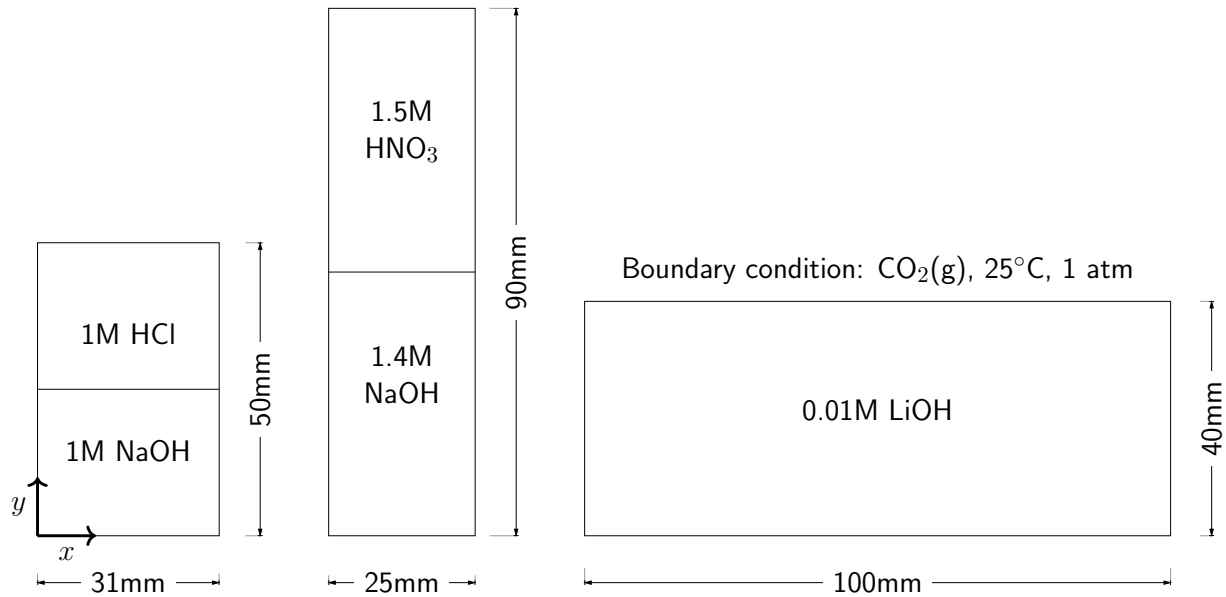
Figure 2. The selected experimental setups: The left panel shows the chemically driven convection experiments with $1\,M$ HCl–$1\,M$ NaOH (Almarcha et al., 2010). The middle panel shows the chemically driven convection setup with $1.5\,M$ HNO_3–$1.4\,M$ NaOH (Mizev et al., 2021). The right panel shows a smaller setup of CO_2 dissolution with $0.01\,M$ LiOH solutions, compared to the original setup ($165\,mm$ by $210\,mm$), designed by Thomas et al. (2016). From left to right, the gap widths of the Hele-Shaw cells are $0.5\,mm$, $1.2\,mm$, and $0.5\,mm$. All Hele-Shaw cells are closed, so that no fluid leaves the domain.

the time it takes for fingering to commence is longer in the simulations than in the experiments. This delay can be attributed
350   to the finite representation by the numerical scheme of the diffusion and electromigration fluxes at the initially sharp interface between the acid and the base, which could be improved by adaptive mesh refinement or a higher mesh resolution.

     Figure 4 compares the images from the physical experiment with our numerical simulations that utilize the Nernst–Planck model with a lag time of 40 seconds. We further present the filled contours of the Na^+ molar concentration and the molar concentration of Na^+ and Cl^- combined. Na^+ is transported upwards, closely following the low-density fingers. The rightmost
355   plots of Figure 4 show that the low-density region is caused by lower concentrations of both Na^+ and Cl^-. The dark red region indicates lower salinity and corresponds to the yellow-turquoise, low-density region. More interestingly, a stable region of higher salinity (shown in blue) forms close to the initial contact line of the acid and the base.

     Figure 5 shows a comparison of the experiments of $1.5\,M$ HNO_3 and $1.4\,M$ NaOH solutions and the simulations. Both the Nernst–Planck model and the single-diffusivity model reproduce the shock-wave-like structures seen in the experimental
360   interferometry images, i.e., the rapid downward movement of the acid–base interface. In the interferometry images at the 3 seconds snapshot, fingers appear, which the simulations cannot reproduce at the 3 seconds simulation time. As discussed before, during the HCl–NaOH experiments, the mismatch is due to the poor approximation of diffusive fluxes at the initially
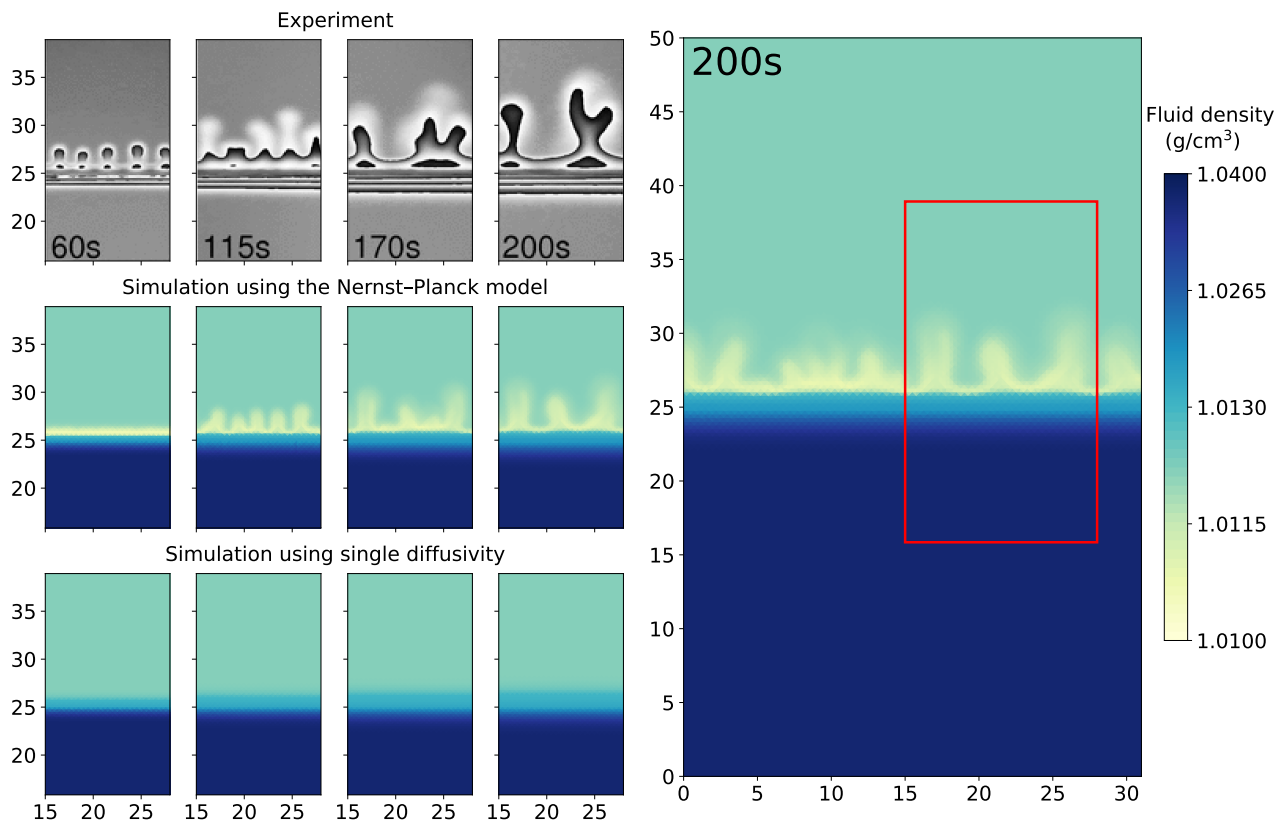
Figure 3. Comparing the simulations using the Nernst–Planck model and those using the single-diffusivity model with the experiments of chemically driven convection in HCl–NaOH solutions. The right panel shows the full simulation domain ($31\,\mathrm{mm}$ by $50\,\mathrm{mm}$), and the red box indicates the comparison window. On the left panel, the time increases from left to right, and the simulation time corresponds to the time shown in the experiment snapshots. The experimental figures are based on digital interferometry, where we adapted the figures from Almarcha et al. (2010).

sharp interface. For the later snapshots, 150 and 700 seconds, we observe the sparsening of the dark lines in the interferometry images. This indicates that the fluid densities become more homogeneous, and the simulations replicate such mixing physics.

365   In Figure 6, we show the simulations of the molar concentrations of Na^+ and the streamlines of the fluid velocity field at 25, 50, 75, 100, and 125 seconds. One can clearly see that the Na^+ plumes follow the streamlines. The maximum velocity in the y direction of the Nernst–Planck simulations and the single-diffusivity simulations during the five presented snapshots are $4.086\,\mathrm{mm/s}$ and $4.184\,\mathrm{mm/s}$, respectively. Both models replicate the experiments, and no particular difference can be observed. In such experimental settings, where density-driven convection is dominant, a distinctive description of the diffusive 370   fluxes is less crucial.
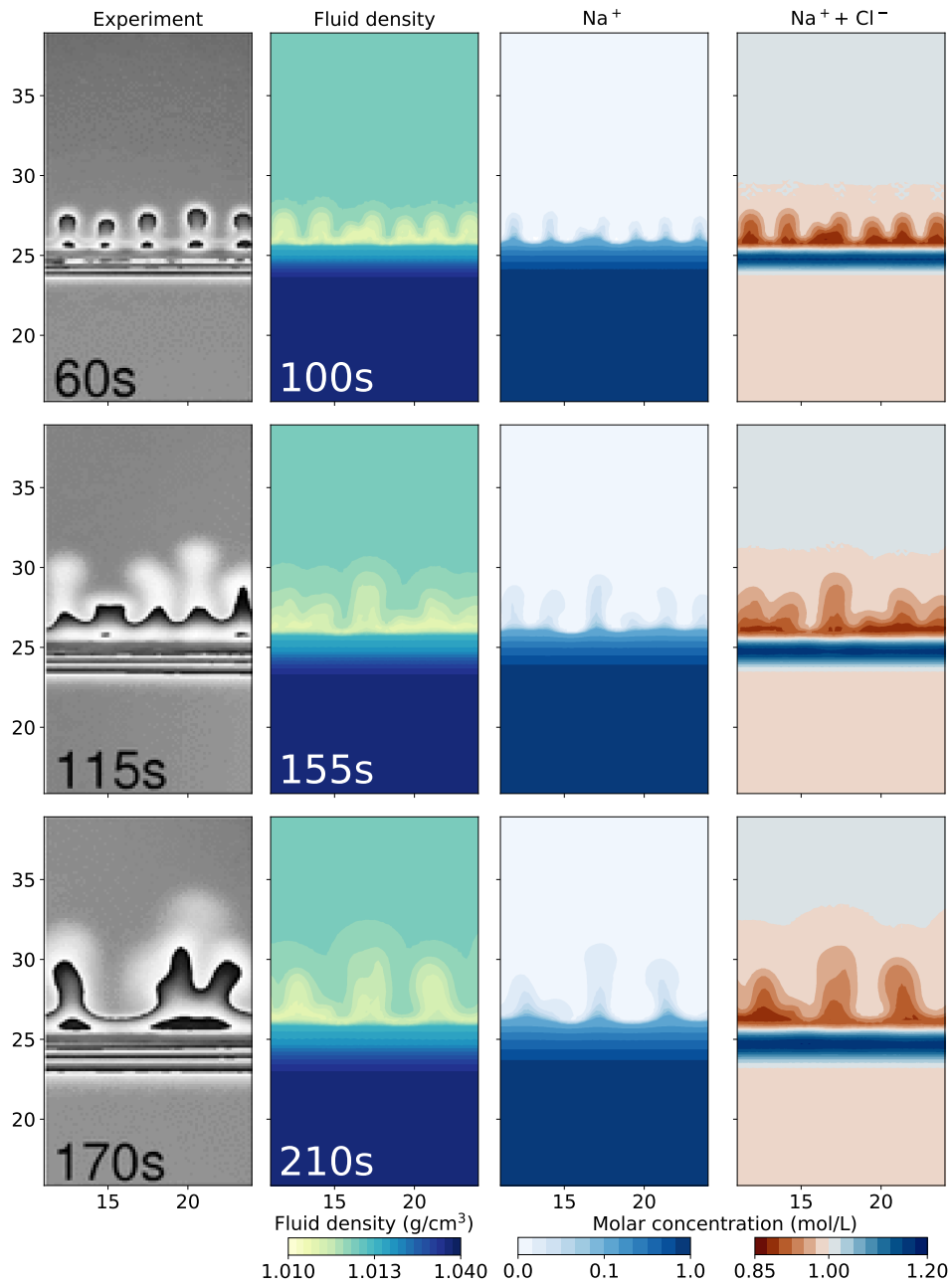
Figure 4. Comparing the simulation results using the Nernst–Planck model with the experimental results of chemically driven convection in HCl–NaOH solutions. The leftmost column shows the experimental figures, adapted from Almarcha et al. (2010). The filled contour plots present the simulation results, showing the fluid density, molar concentration of Na^+, and the sum of the molar concentrations of Na^+ and Cl^- with a lag time of 40 seconds.
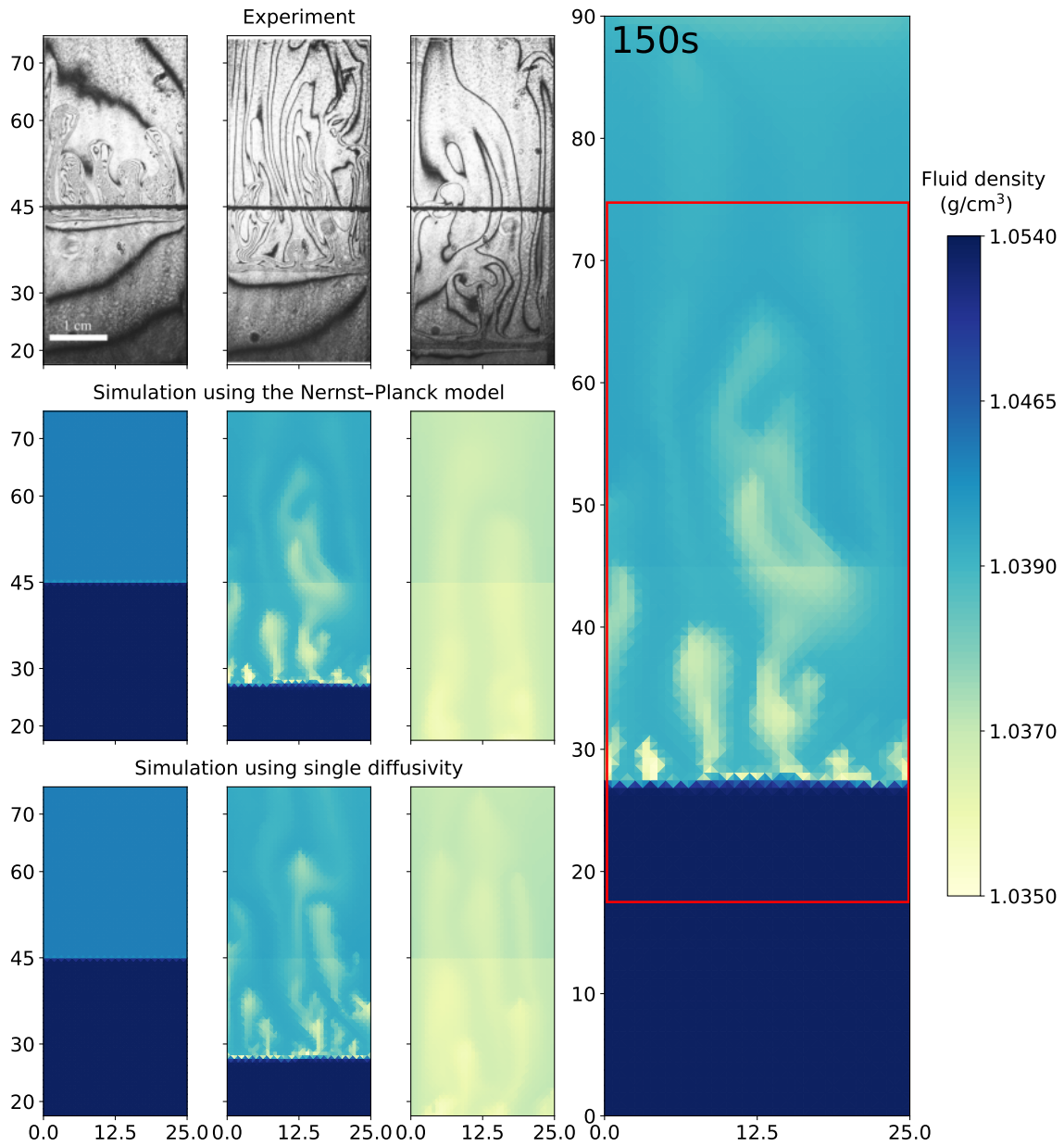
Figure 5. Comparing the simulations using the Nernst–Planck model and those using the single-diffusivity model with the experiments of chemically driven convection in HNO_3–$NaOH$ solutions. The right panel shows the full simulation domain ($25\,mm$ by $90\,mm$), and the red box indicates the comparison window. In the left panel, the time increases from left to right, where the snapshots correspond to 3, 150, and 700 seconds. The figures from the experiments are based on digital interferometry, and were adapted from Mizev et al. (2021).
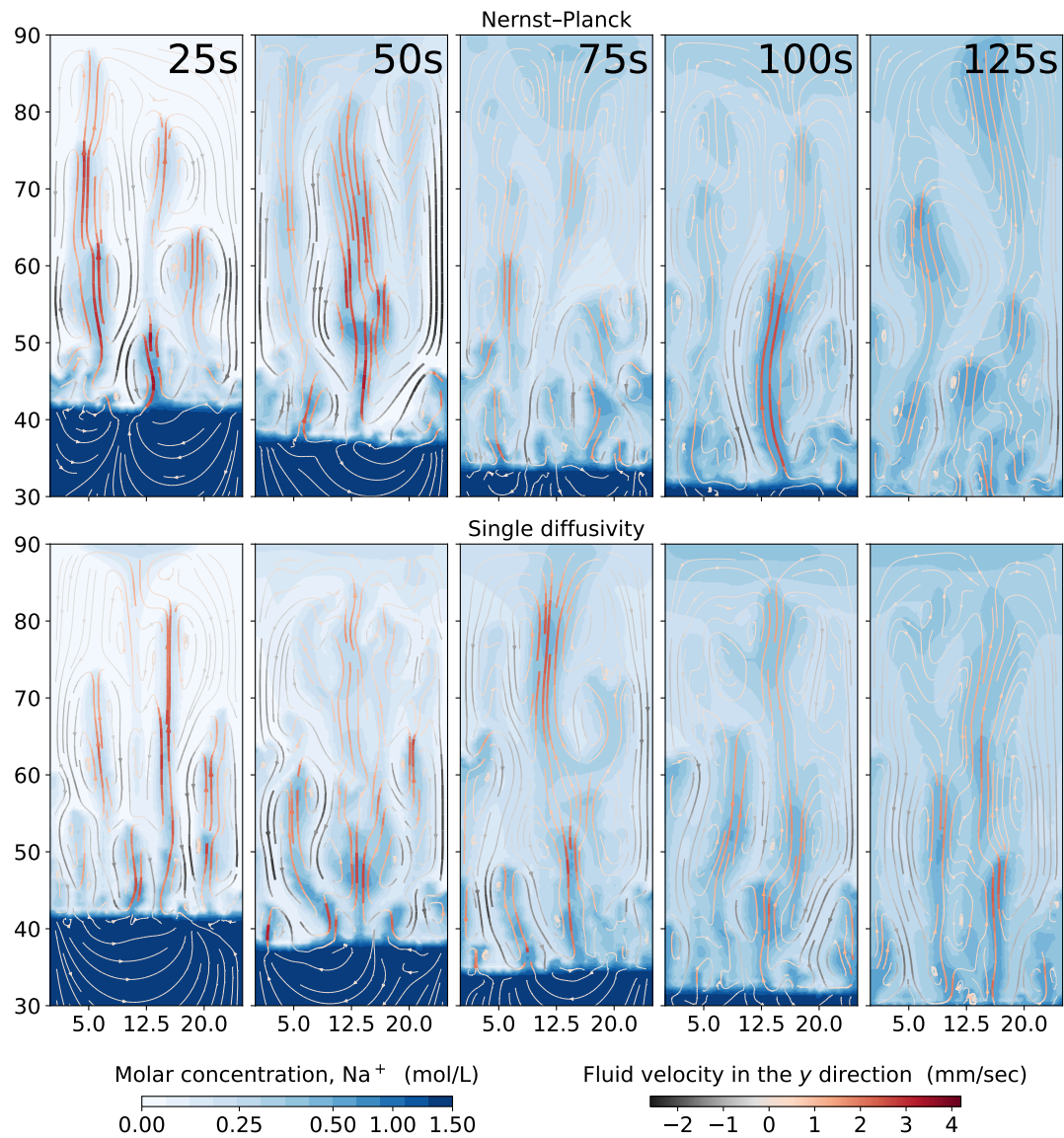
Figure 6. Comparing the simulations using the Nernst–Planck model and those using the single-diffusivity model in modeling chemically driven convection of HNO_3–$NaOH$ solutions. The top and the bottom panel show the simulations employing the Nernst–Planck model and the simulations using one single diffusivity, respectively. The time increases from left to right, and the snapshots correspond to 25, 50, 75, 100, and 125 seconds.

3.2 Simulation of convective dissolution of CO_2 in reactive alkaline solutions

Figure 7 compares the schlieren images of convective dissolution of CO_2 in 0.01 M LiOH solution from the experiments with the density plots using the Nernst–Planck and the single-diffusivity models. The onset time of the Nernst–Planck model and the single-diffusivity model are approximately 600 and 1400 seconds, respectively. Therefore, the densities of the two simulation results are shown for the onset time plus the experimental snapshot, 780 seconds. At these time steps, the convective fingers in both models reach similar depths of $\sim 17\,\mathrm{mm}$, as shown in the schlieren image from the experiment. The simulation result of the Nernst–Planck model is comparable to the experiment of CO_2 dissolving into the LiOH brine and shows a broader range of finger sizes compared to those of the single-diffusivity model. In contrast, the convective fingers in the single-diffusivity model resemble the fingers seen in the schlieren image, where CO_2 dissolves in H_2O, where the electromigration effects are weaker, compared to the LiOH solution.

Figure 8 shows the density plots of the simulation results using the Nernst–Planck model and the single-diffusivity model. In the top panel, we show the onset of the convective instability. The single-diffusivity model develops instabilities roughly 800 seconds later than the Nernst–Planck model. The second panel from the top shows the convective fingers after their onset; the Nernst–Planck model yields twice the number of fingers, compared to the single-diffusivity model. In these two panels, we observe that the Nernst–Planck model generates a low-density layer, colored in yellow, at the dissolution front, which causes a shorter onset time and more fingers due to a larger density contrast. In the two bottom panels, we show the evolution of the simulations at later times. There are differences in the fingering pattern, such as more secondary convective fingers being generated by the Nernst–Planck model. Moreover, the low-density layer in the Nernst–Planck model persists even at later time.

To further quantify the simulation results, Figure 9 visualizes the pH plots as well as the molar concentrations of Li^+, $CO_2(aq)$, and CO_3^{2-}. We use a diverging color palette to present the acid regions (orange) and the base regions (purple). The molar concentrations of $CO_2(aq)$ and CO_3^{2-} follow the pH color palette, where most $CO_2(aq)$ exists in the acidic region and most CO_3^{2-} exists in the basic region. Such changes in the CO_2 states lead to changes in the charge states $z = 0, 1, 2$ for $CO_2(aq)$, HCO_3^-, and CO_3^{2-}, respectively. In the Li^+ plot of the Nernst–Planck simulations, we observe up to 30% differences in the molar concentration compared to the initial condition ($0.01\,\mathrm{mol/L}$). The red regions are where Li^+ accumulates, and we can clearly see the interaction between Li^+ and CO_3^{2-}, whereby Li^+ migrates towards the double-charged CO_3^{2-}. Meanwhile, the single-diffusivity simulation shows almost no contrast in the filled contours, other than due to numerical errors. The difference between the maximum and minimum molar concentrations of Li^+ is $\sim 10^{-8}\,\mathrm{mol\,L^{-1}}$. Using the Nernst–Planck model, we confirm the influence of the inert (spectator) ions in the reactive-transport situation discussed in Thomas et al. (2016).

4 Discussion

In this section, we discuss whether the Nernst–Planck model is valid, and its differences are compared to the commonly used single-diffusivity model.
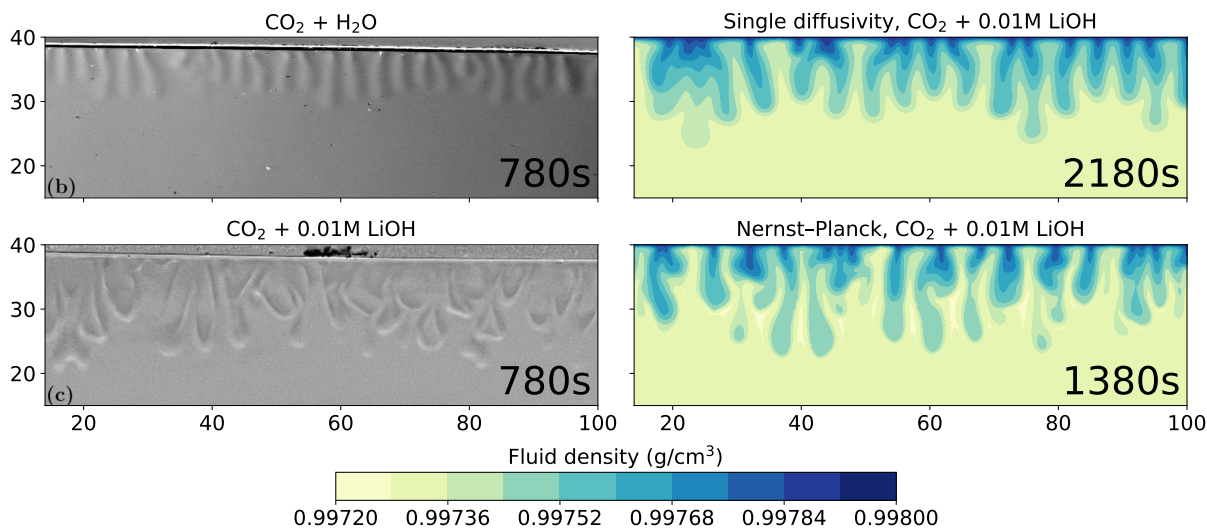
Figure 7. Comparing the simulations using the Nernst–Planck model and those using the single-diffusivity model with the experiments of convective dissolution of CO_2 in 0.01 M LiOH solution. We plot the experimental figure of CO_2 dissolving in pure water in the top-left to demonstrate the differences in fingering patterns, compared with CO_2 dissolving in LiOH solution. The filled contours show the fluid density. The experimental figures are adapted from Thomas et al. (2016).

4.1 The validity of the Nernst–Planck model under reactive flow conditions

Comparing to two chemically driven convection experiments conducted by Almarcha et al. (2010) and Mizev et al. (2021), we show that the Nernst–Planck numerical model is able to reproduce the convective fingering observed in the experiments, especially in the 1 M HCl–1 M NaOH case. The mechanism that leads to convection is known as the reactive diffusive layer convection (DLC) instability, which arises when a less-dense solution of a fast-diffusing component overlies a denser solution of a slow-diffusing component (Lemaigre et al., 2013). In our Nernst–Planck model, the fast-diffusing component is H^+. Additionally, the convective plumes of the reactive DLC instability are only observed above the initial contact line between the acid and the base (Almarcha et al., 2010; Lemaigre et al., 2013). This is the asymmetric feature of the reactive DLC instability, which can also be observed in the Nernst–Planck simulations (Figure 3). Directly quoting Lemaigre et al. (2013): "The asymmetry of the DLC pattern is related to the chemical reaction which consumes the acid before it can accumulate in the lower layer and replaces it by a third species of intermediate contribution to density". This third species is salt, a product of the acid–base neutralization reaction. Using the Nernst–Planck model, we not only replicate this stabilizing lower layer, but also find that the accumulation of salt can exceed the salt concentration at chemical equilibrium (in this case, approximately $0.5\,\mathrm{mol\,L^{-1}}$), as shown in Figure 4.

In the 1.5 M HNO_3–1.4 M NaOH case, both the Nernst–Planck model and the single-diffusivity model reproduce the shock-wave-like structure observed in the experiments performed by Mizev et al. (2021). The instability in this case is convection-controlled (CC) (Bratsun et al., 2017). The underlying mechanism of this CC instability has previously been attributed to
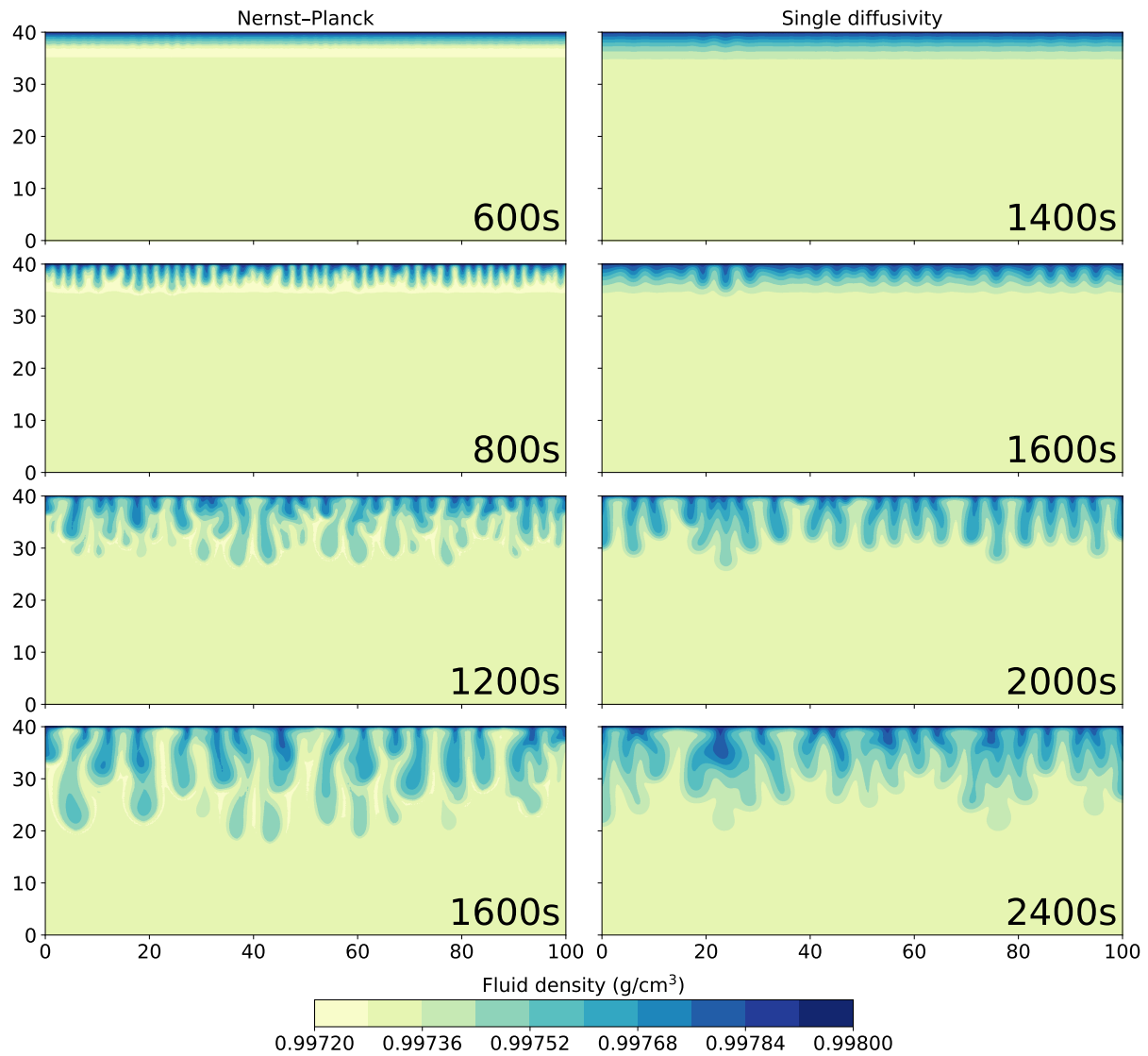
Figure 8. Comparing the simulations, using the Nernst–Planck model and those using the single-diffusivity model, when numerically modeling convective dissolution of CO_2 in 0.01 M LiOH solution. The filled contours show the fluid density. Time increases from top to bottom and is indicated in the plots.
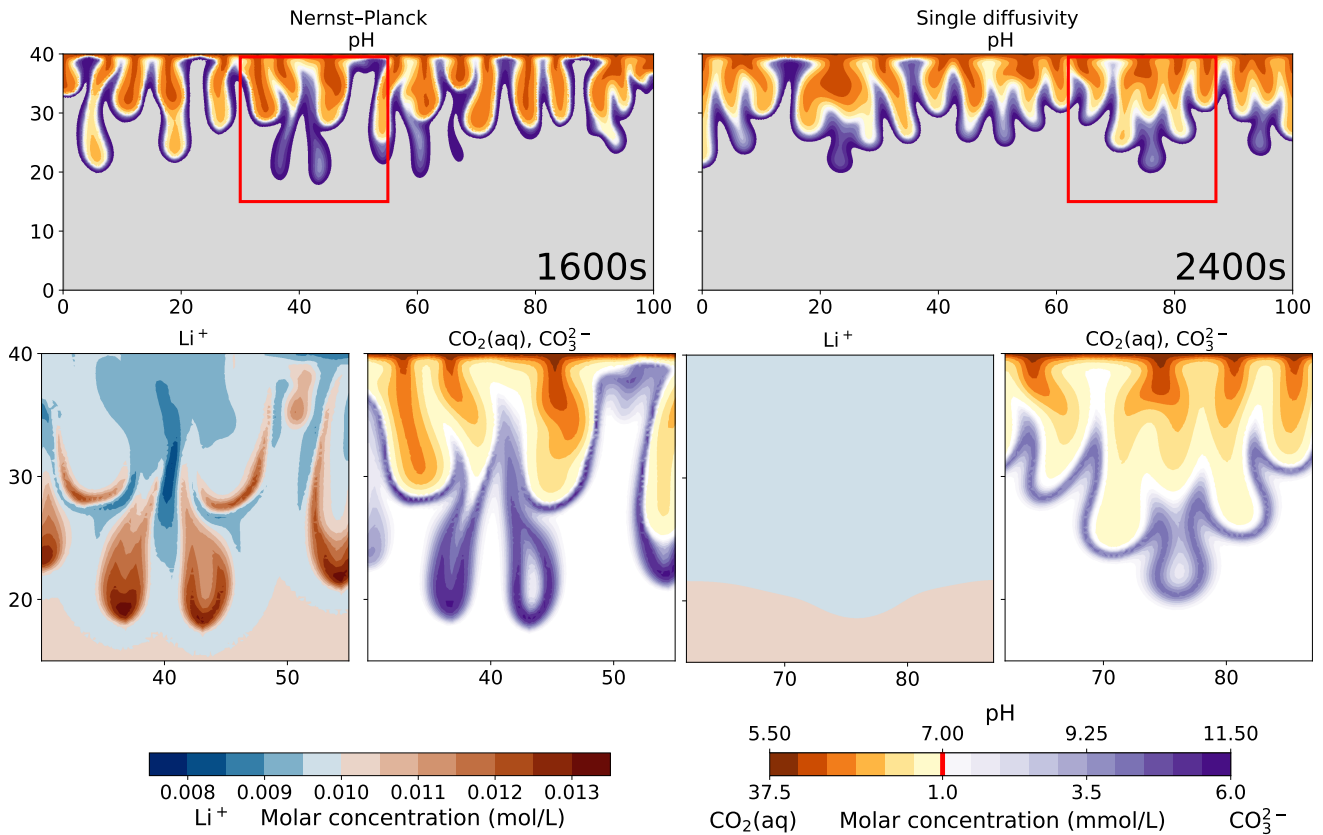
Figure 9. Comparing the simulations, using the Nernst–Planck model and those using the single-diffusivity model, when numerically modeling convective dissolution of CO_2 in $0.01\,M$ LiOH solution. The top panel shows the pH values. The orange palette represents acidic regions (pH$<$ 7) and the purple palette represents basic regions (pH$>$ 7). The gray region shows the unperturbed part of the LiOH solution (pH$>$ 11.5). In the bottom panel, we show the zoomed-in view of the red boxes in the top panel. We plot the molar concentrations of Li^+, $CO_2(aq)$, and CO_3^{2-}. Notice that the molar concentrations of $CO_2(aq)$ and CO_3^{2-} share the same color bar with the pH values. The red vertical line in the color bar indicates the minimum molar concentrations of $CO_2(aq)$ and CO_3^{2-} shown in the plots.

unequal and concentration-dependent diffusivities (Bratsun et al., 2017, 2021, 2022). However, using the single-diffusivity model, our simulations can reproduce the experimental observations. This indicates that modeling CC instabilities is less sensitive to the specific diffusivity model used, as the CC instability is mostly caused by the density-driven convective flow of the reactants. Nevertheless, there are no drawbacks (other than being computationally more intensive) in employing the Nernst–Planck equations to model the CC instability cases.

In the case of CO_2 dissolving in 0.01 M LiOH solution, we observe a much later convective onset time in the numerical simulations than in the experiment by Thomas et al. (2016). This discrepancy is much larger than in the 1 M HCl–1 M NaOH case. It cannot only be explained by the ineffectiveness of the diffusive-flux approximations at the sharp interface of dissolving $CO_2(aq)$. Convective instability can be introduced in a variety of ways in laboratory experiments. For example, interfacial instabilities between the gas and the liquid can occur due to the gas flowing over the interface. Furthermore, if the injected CO_2 is not saturated with water, evaporation of water can induce gravitational instability (Bringedal et al., 2022). These localized physical processes at the gas-liquid boundary, which are not included in our numerical models, could be partly responsible for the later onset times we observe in our simulations.

Although the comparison between the experiment and the simulations of the two models in Figure 7 can be ambiguous, i.e., the Nernst–Planck model is not *the* model that describes the schlieren images, we emphasize that the difference between the experiments with Li^+ and those without it is apparent in the figure. When comparing the Nernst–Planck and the single-diffusivitiy models, the former exhibits an earlier fingering onset time. This is caused by electromigration effects, which generate the low-density layer shown in Figure 8. The electromigration of Li^+, an inert ion in this case, and subsequent effects on the overall transport processes shown in Figure 9, is clearly a motivation for more studies on the electromigration effects of inert (spectator) ions.

To summarize, we use the experiments conducted by Almarcha et al. (2010); Thomas et al. (2016); Mizev et al. (2021) and our numerical simulation results to show that the Nernst–Planck model is valid in the considered reactive transport scenarios. Particularly in the 1 M HCl–1 M NaOH experiments, where the convective fingers are controlled by diffusion. In the convection-controlled regime, electromigration effects are less crucial, and both the Nernst–Planck model and the single-diffusivity model replicate the 1.5 M HNO_3–1.4 M NaOH experiments well. In the case of CO_2 dissolving in an alkaline solution, the use of the Nernst–Planck model uncovers intricate details on how inert ionic species can affect the overall reactive transport processes.

4.2 Electromigration and the relevance of inert ionic species

Using the simulations of chemically driven convection in 1 M HCl–1 M NaOH solutions, we elaborate the main differences between the Nernst–Planck model and the single-diffusivity model. The top panel of Figure 10 shows the molar concentrations of salt, plotted against elevation in the Hele-Shaw cell. As expected, the single-diffusivity model shows the concentration profile of two diffusing Heaviside step functions of Na^+ and Cl^-. We plot the sum of Na^+ and Cl^- molar concentrations to demonstrate that in single-diffusivity models, Na^+ and Cl^- exhibit an ideal behavior, known as equimolar counterdiffusion. The Nernst–Planck model shows unexpected results, such as the accumulation of Na^+ and Cl^- close to the acid–base bound-

ary. The accumulation can be attributed to a larger diffusivity of the acid components, H^+ ($9.311 \times 10^{-3}\,\mathrm{mm^2s^{-1}}$) and Cl^- ($2.032 \times 10^{-3}\,\mathrm{mm^2s^{-1}}$), compared to the base components, OH^- ($5.273 \times 10^{-3}\,\mathrm{mm^2s^{-1}}$) and Na^+ ($1.334 \times 10^{-3}\,\mathrm{mm^2s^{-1}}$), which leads to more Cl^- migrating into the base. One might get the impression that this observation can be modeled by employing multiple diffusivities in the single-diffusivity model. However, we emphasize that the Nernst–Planck model has to be utilized to maintain electroneutrality at the continuum scale. Closely inspecting the fluid-density profile, the deficit in salt concentration leads to a lower density than the acid, which causes the reactive DLC instability. The single-diffusivity model is unable to depict the accumulation or deficit in the salt concentrations.

Moreover, the only reaction that occurs in this experiment is the neutralization reaction

$$H^+ + OH^- \rightleftharpoons H_2O(l). \tag{39}$$

No reaction occurs for Na^+ and Cl^-. Hence, they are inert (spectator) species. We elaborate on the transport of H^+, given by

$$\frac{\partial C_{H+}}{\partial t} = -\nabla \cdot \left(-D_{H+H+}\nabla C_{H+} - D_{H+OH-}\nabla C_{OH-} - D_{H+Na+}\nabla C_{Na+} - D_{H+Cl-}\nabla C_{Cl-} + \boldsymbol{u} C_{H+}\right), \tag{40}$$

where the generalized-diffusion interpretation of the Nernst–Planck equation is applied, Eq. (23), and D_{ij}, are defined by Eq. (24). In the Nernst–Planck model, the inert species contribute to the transport processes in the form of the cross-coupling diffusivities, D_{H+Na+}, D_{H+Cl-}, and the concentration gradients, ∇C_{Na+} and ∇C_{Cl-}. It has also been shown that the cross-coupling diffusivities (without reactions) can lead to hydrodynamic instabilities (Budroni, 2015). Such tight coupling between ionic species does not exist in the single-diffusivity model. Hence, the Nernst–Planck model is more expressive and capable of modeling the intricacies of reactive transport.

To further demonstrate the electromigration effects of the inert species, Figure 11 shows the species concentration profile of simulations of CO_2 dissolution in $0.01\,M$ LiOH solution. We plot the 600-seconds snapshot, which is the onset time of the Nernst–Planck model (see Figure 8). Figure 11 focuses on the concentration profile close to the top boundary, which is the elevation at $40\,mm$. When CO_2 dissolves into the LiOH solution, the following chemical reactions occur

$$CO_2(aq) + H_2O(l) \rightleftharpoons H_2CO_3(aq) \tag{41}$$

$$H_2CO_3(aq) \rightleftharpoons H^+ + HCO_3^- \tag{42}$$

$$H^+ + OH^- \rightleftharpoons H_2O(l), \tag{43}$$

where $H_2CO_3(aq)$ is a reaction intermediate that we have not considered in our numerical models. These reactions represent the entire process of $CO_2(aq)$ acidifying aqueous solutions. If we add Eqs. (41)–(43), this yields a simplified reaction,

$$CO_2(aq) + OH^- \rightleftharpoons HCO_3^-. \tag{44}$$

In Figure 11, at elevations above $38\,mm$, excessive amounts of $CO_2(aq)$ react with OH^-, producing HCO_3^-. However, due to the electroneutrality constraints, the concentration of HCO_3^- is limited by the concentrations of the cations, particularly Li^+. When $CO_2(aq)$ migrates deeper into the more basic regions, i.e., elevations between 36–$38\,mm$, another reaction happens,

$$HCO_3^- \rightleftharpoons H^+ + CO_3^{2-}. \tag{45}$$

In this basic region, HCO_3^- donates H^+ and forms CO_3^{2-}. We can still observe Li^+ constraining the amount of CO_3^- owing to electroneutrality. Another cation that can contribute to electroneutrality is H^+. However, in this basic region, the amount of H^+ is negligible. Although Li^+ is inert and does not take part in the reactions, Eqs. (41)–(45), it is still relevant in the overall reactive transport process. The transport of Li^+ in the top-left panel of Figure 11 is evidence of its involvement in the reactive transport process.

Focusing on the density plots in Figure 11, not only does the Nernst–Planck model produce lower densities than the single-diffusivity model, the density gradient between 37–$38\,\mathrm{mm}$ is also higher in the Nernst–Planck model. This can explain the earlier onset time of the Nernst–Planck model, compared to the onset time of the single-diffusivity model. We end the discussion with Figure 12, which presents Li^+ migrating as CO_2 is dissolving and reacting in the solution. The tight coupling between the equilibrium reactions, electromigration effects, and density-driven flow elicits further investigation, both numerically and experimentally.
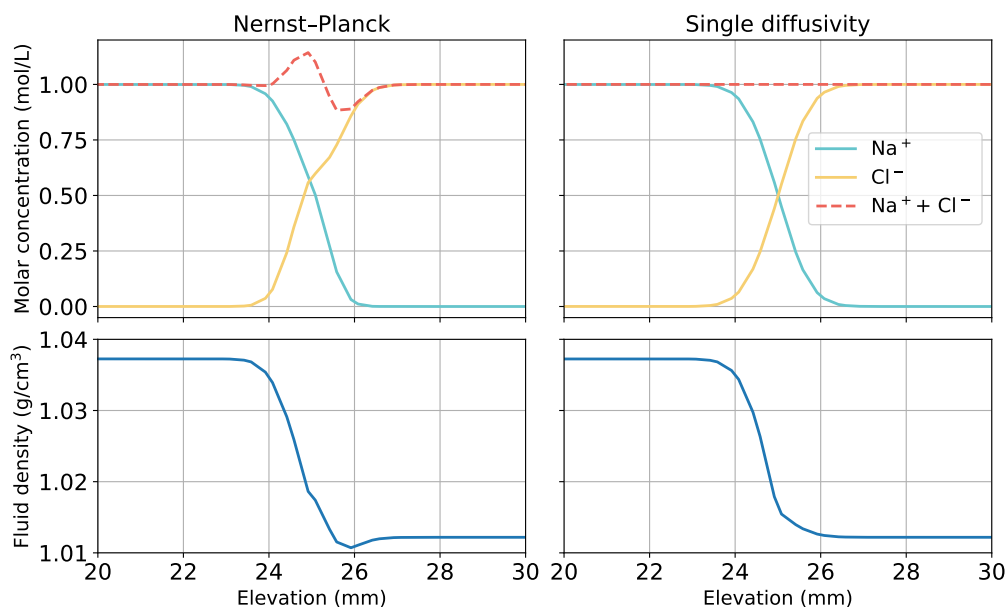
Figure 10. Comparing the simulations using the Nernst–Planck model and those using the single-diffusivity model in modeling chemically driven convection in $1\,\mathrm{M}$ HCl–$1\,\mathrm{M}$ $NaOH$ solutions. The snapshot is given at $t = 60$ seconds, when fingers have not yet appeared (see Figure 3). We plot the molar concentrations of salt and the fluid density against elevation in the Hele-Shaw cell, focusing on the contact region between the acid and the base, $25\,\mathrm{mm}$.

5 Conclusions

By comparing the single-diffusivity model and the Nernst–Planck model with reaction-driven flow experiments, we have demonstrated the importance and necessity under certain conditions of modeling the electromigration effects between charged
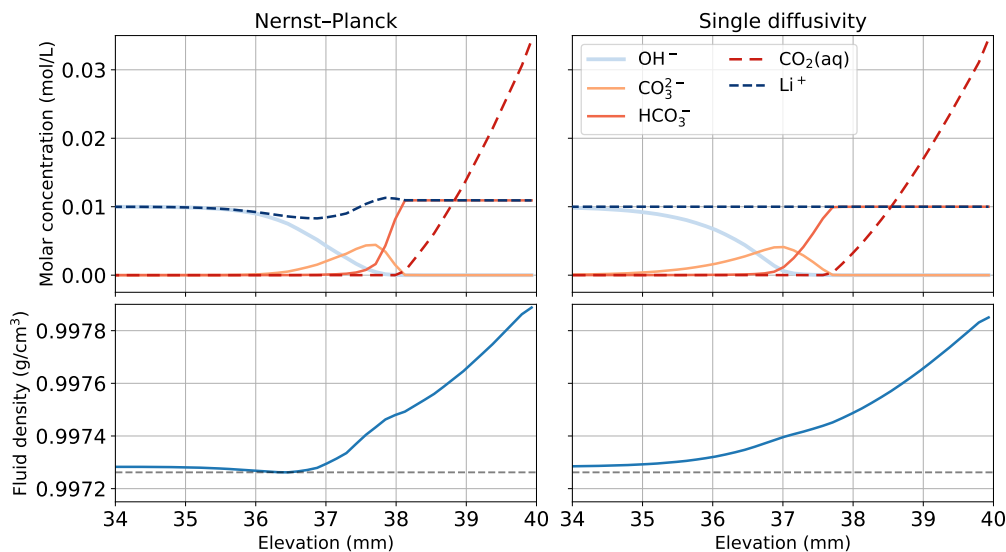
Figure 11. Comparing the simulations using the Nernst–Planck model and those using the single-diffusivity model in modeling convective dissolution of CO_2 in 0.01 M LiOH solution. The snapshot is given at $t = 600$ seconds, when fingers have not yet appeared (see Figure 8). We plot the molar concentrations of Li^+, OH^-, CO_3^{2-}, HCO_3^-, and $CO_2(aq)$ and the fluid density against elevation in the Hele-Shaw cell, focusing on the regions close to the top boundary of the Hele-Shaw cell. The gray-horizontal-dashed lines in the bottom panel show the minimum density observed in the Nernst–Planck model at $t = 600$ seconds.

species in aqueous environments using the Nernst–Planck model. Our results of simulating the reaction-driven flow experiments and convective dissolution of CO_2 show that non-reacting species (Na^+, Cl^-, Li^+) influence the overall reactive transport process via electromigration effects, which couple the charged species. Such electromigration effects cannot be modeled using the single-diffusivity model, and further studies on quantifying the valid regimes of the single-diffusivity model are recommended. We conclude that:

– By comparison of our numerical modeling results to previously published flow experiments, we have shown that the Nernst–Planck model is valid for modeling reactive transport processes. The processes in the experiments considered here are characterised by an intricate interplay between diffusion, reaction, electromigration, and density-driven convection.

– Compared to the often-used single-diffusivity model, the Nernst–Planck model enables the numerical modeling of the electromigration of ionic species, introduced by differing species diffusivities, resulting in more and improved physical insights.

– These reaction-driven flow experiments from literature can be further utilized for benchmarking reactive transport codes.
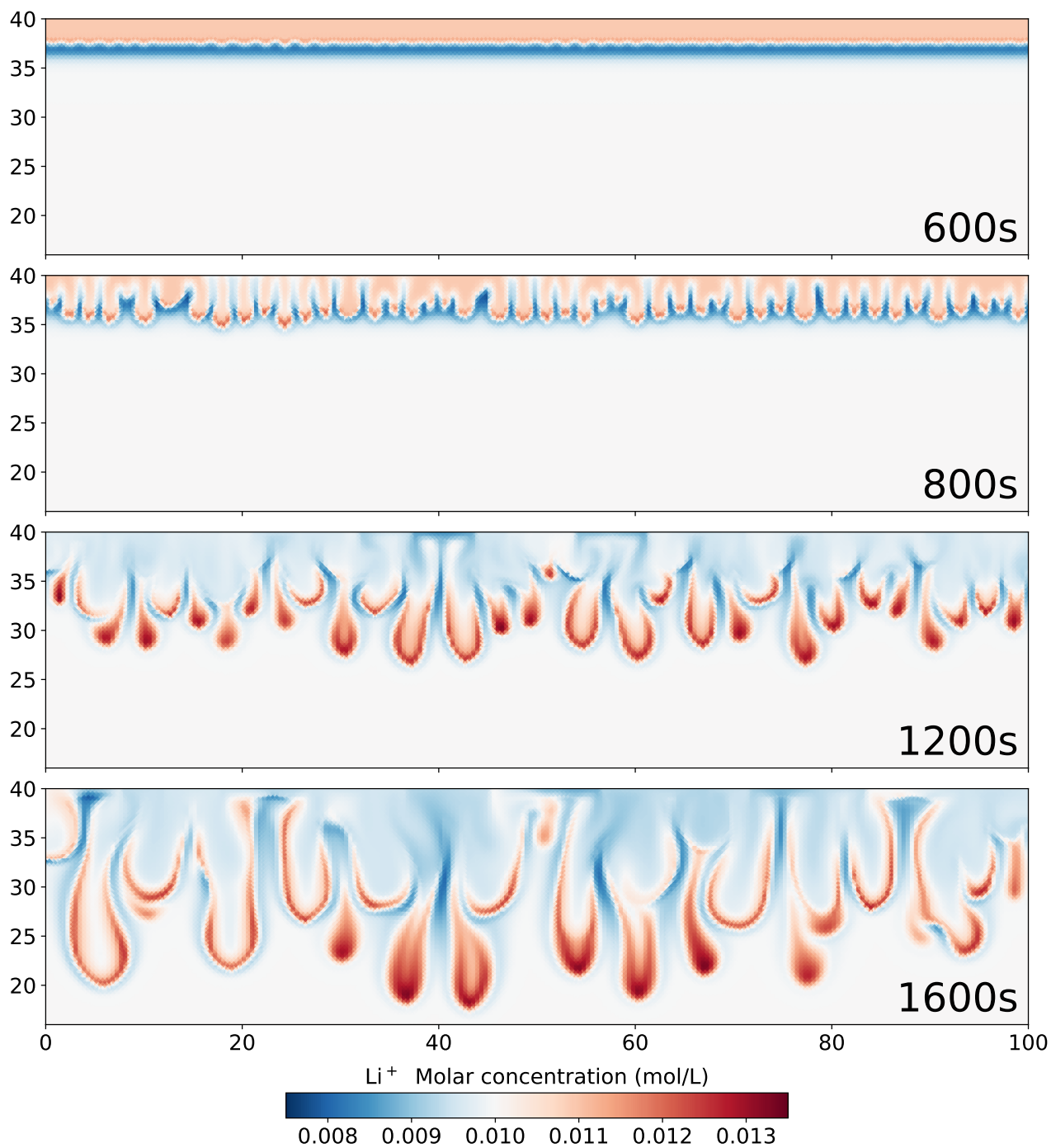
Figure 12. Simulation results using the Nernst–Planck model, showing the molar concentration of Li^+ during the convective dissolution of CO_2 in 0.01 M LiOH solution.

Code and data availability. The RetroPy code is available on GitHub (https://github.com/pwhuang/RetroPy), and version 1.0 is archived on Zenodo (https://doi.org/10.5281/zenodo.7371384). The RetroPy code is published under the GNU Lesser General Public License (LGPL). The codes that produced the simulations are located in the example folder of RetroPy, where the chemical_convection folder contains the chemically-driven convection examples, and the CO2_convection folder contains the CO_2 convective dissolution example. The data and scripts that recreate the figures are available on Zenodo (https://doi.org/10.5281/zenodo.7362225).

Video supplement. We provide the simulation results of Figure 4, 6, and 12 in the mp4 format. They are located in https://doi.org/10.3929/ethz-b-000579224.

Author contributions. Po-Wei Huang: Conceptualization, Methodology, Software, Formal analysis, Writing - Original Draft, Writing - Review & Editing, Visualization

Bernd Flemisch: Writing - Review & Editing

Chao-Zhong Qin: Writing - Review & Editing

Martin O. Saar: Supervision, Writing - Review & Editing

Anozie Ebigbo: Conceptualization, Writing - Review & Editing, Project administration, Funding acquisition

Competing interests. The authors declare that there is no known conflict of interest.

Acknowledgements. This work was supported by the Swiss National Science Foundation project entitled "Analysing spatial scaling effects in mineral reaction rates in porous media with a hybrid numerical model." We also thank the Werner Siemens-Stiftung (Werner Siemens Foundation) for its support of the Geothermal Energy and Geofluids (GEG.ethz.ch) Group at ETH Zurich, Switzerland.

We would like to thank our colleagues, Xiang-Zhao Kong and Isamu Naets, for all the helpful discussions. Furthermore, this work could not have been done without the visualization tool matplotlib (Hunter, 2007). We utilized the color palettes provided by Harrower and Brewer (2003); Davis (2019); Crameri (2021).

References

Abd, A. S. and Abushaikha, A. S.: Reactive transport in porous media: a review of recent mathematical efforts in modeling geochemical reactions in petroleum subsurface reservoirs, SN Applied Sciences, 3, 401, https://doi.org/10.1007/s42452-021-04396-9, 2021.

Aftab, A., Hassanpouryouzband, A., Xie, Q., Machuca, L. L., and Sarmadivaleh, M.: Toward a Fundamental Understanding of Geological Hydrogen Storage, Ind. Eng. Chem. Res., 61, 3233–3253, https://doi.org/10.1021/acs.iecr.1c04380, 2022.

Agartan, E., Trevisan, L., Cihan, A., Birkholzer, J., Zhou, Q., and Illangasekare, T. H.: Experimental study on effects of geologic heterogeneity in enhancing dissolution trapping of supercritical CO_2, Water Resour. Res., 51, 1635–1648, https://doi.org/10.1002/2014WR015778, 2015.

Åkerlöf, G. and Teare, J.: A Note on the Density of Aqueous Solutions of Hydrochloric Acid, J. Am. Chem. Soc., 60, 1226–1228, https://doi.org/10.1021/ja01272a063, 1938.

Almarcha, C., Trevelyan, P. M. J., Grosfils, P., and De Wit, A.: Chemically Driven Hydrodynamic Instabilities, Phys. Rev. Lett., 104, 1–4, https://doi.org/10.1103/PhysRevLett.104.044501, 2010.

Almarcha, C., R'Honi, Y., De Decker, Y., Trevelyan, P. M. J., Eckert, K., and De Wit, A.: Convective Mixing Induced by Acid–Base Reactions, The Journal of Physical Chemistry B, 115, 9739–9744, https://doi.org/10.1021/jp202201e, 2011.

Alnæs, M. S.: UFL: a finite element form language, in: Automated Solution of Differential Equations by the Finite Element Method: The FEniCS Book, edited by Logg, A., Mardal, K.-A., and Wells, G. N., pp. 303–338, Springer Berlin Heidelberg, Berlin, Heidelberg, https://doi.org/10.1007/978-3-642-23099-8_17, 2012.

Alnæs, M. S., Blechta, J., Hake, J., Johansson, A., Kehlet, B., Logg, A., Richardson, C., Ring, J., Rognes, M. E., and Wells, G. N.: The FEniCS Project Version 1.5, Archive of Numerical Software, 3, 9–23, https://doi.org/10.11588/ans.2015.100.20553, 2015.

Amarasinghe, W., Fjelde, I., Åge Rydland, J., and Guo, Y.: Effects of permeability on CO_2 dissolution and convection at reservoir temperature and pressure conditions: A visualization study, Int. J. Greenh. Gas Con., 99, 103 082, https://doi.org/10.1016/j.ijggc.2020.103082, 2020.

Amestoy, P. R., Duff, I. S., L'Excellent, J.-Y., and Koster, J.: A Fully Asynchronous Multifrontal Solver Using Distributed Dynamic Scheduling, SIAM J. Matrix Anal. A., 23, 15–41, https://doi.org/10.1137/S0895479899358194, 2001.

Amestoy, P. R., Buttari, A., L'Excellent, J.-Y., and Mary, T.: Performance and Scalability of the Block Low-Rank Multifrontal Factorization on Multicore Architectures, ACM T. Math. Software, 45, 1–26, https://doi.org/10.1145/3242094, 2019.

Audigane, P., Gaus, I., Czernichowski-Lauriol, I., Pruess, K., and Xu, T.: Two-dimensional reactive transport modeling of CO_2 injection in a saline aquifer at the Sleipner site, North Sea, Am. J. Sci., 307, 974–1008, https://doi.org/10.2475/07.2007.02, 2007.

Avnir, D. and Kagan, M.: Spatial structures generated by chemical reactions at interfaces, Nature, 307, 717–720, https://doi.org/10.1038/307717a0, 1984.

Azin, R., Raad, S. M. J., Osfouri, S., and Fatehi, R.: Onset of instability in CO_2 sequestration into saline aquifer: scaling relationship and the effect of perturbed boundary, Heat Mass Transfer, 49, 1603–1612, https://doi.org/10.1007/s00231-013-1199-7, 2013.

Babaei, M. and Islam, A.: Convective-Reactive CO_2 Dissolution in Aquifers With Mass Transfer With Immobile Water, Water Resour. Res., 54, 9585–9604, https://doi.org/10.1029/2018WR023150, 2018.

Bacuta, C.: A Unified Approach for Uzawa Algorithms, SIAM J. Numer. Anal., 44, 2633–2649, https://doi.org/10.1137/050630714, 2006.

Balay, S., Abhyankar, S., Adams, M. F., Benson, S., Brown, J., Brune, P., Buschelman, K., Constantinescu, E. M., Dalcin, L., Dener, A., Eijkhout, V., Gropp, W. D., Hapla, V., Isaac, T., Jolivet, P., Karpeev, D., Kaushik, D., Knepley, M. G., Kong, F., Kruger, S., May, D. A.,

McInnes, L. C., Mills, R. T., Mitchell, L., Munson, T., Roman, J. E., Rupp, K., Sanan, P., Sarich, J., Smith, B. F., Zampini, S., Zhang, H., Zhang, H., and Zhang, J.: PETSc Web page, https://petsc.org/, last access: 2022-11-03, 2022.

Benzi, M., Golub, G. H., and Liesen, J.: Numerical solution of saddle point problems, Acta Numer., 14, 1–137, https://doi.org/10.1017/S0962492904000212, 2005.

Bethke, C. M.: Geochemical and Biogeochemical Reaction Modeling, Cambridge University Press, 2 edn., https://doi.org/10.1017/CBO9780511619670, 2007.

Boffi, D., Brezzi, F., and Fortin, M.: Mixed Finite Element Methods and Applications, Springer, Berlin, Heidelberg, https://doi.org/10.1007/978-3-642-36519-5, 2013.

Bordeaux-Rego, F., Sanaei, A., and Sepehrnoori, K.: Enhancement of Simulation CPU Time of Reactive-Transport Flow in Porous Media: Adaptive Tolerance and Mixing Zone-Based Approach, Transport Porous Med., 143, 127–150, https://doi.org/10.1007/s11242-022-01789-1, 2022.

Boudreau, B. P., Meysman, F. J. R., and Middelburg, J. J.: Multicomponent ionic diffusion in porewaters: Coulombic effects revisited, Earth Planet. Sc. Lett., 22, 653–666, https://doi.org/10.1016/j.epsl.2004.02.034, 2004.

Bożek, B., Sapa, L., and Danielewski, M.: Difference Methods to One and Multidimensional Interdiffusion Models with Vegard Rule, Math. Model. Anal., 24, 276–296, https://doi.org/10.3846/mma.2019.018, 2019.

Bożek, B., Sapa, L., Tkacz-Śmiech, K., Zajusz, M., and Danielewski, M.: Compendium About Multicomponent Interdiffusion in Two Dimensions, Metall. Mater. Trans. A, 52, 3221–3231, https://doi.org/10.1007/s11661-021-06267-9, 2021.

Bratsun, D., Kostarev, K., Mizev, A., and Mosheva, E.: Concentration-dependent diffusion instability in reactive miscible fluids, Phys. Rev. E, 92, 011 003, https://doi.org/10.1103/PhysRevE.92.011003, 2015.

Bratsun, D., Mizev, A., Mosheva, E., and Kostarev, K.: Shock-wave-like structures induced by an exothermic neutralization reaction in miscible fluids, Phys. Rev. E, 96, 053 106, https://doi.org/10.1103/PhysRevE.96.053106, 2017.

Bratsun, D., Mizev, A., and Mosheva, E.: Extended classification of the buoyancy-driven flows induced by a neutralization reaction in miscible fluids. Part 2. Theoretical study, J. Fluid Mech., 916, A23, https://doi.org/10.1017/jfm.2021.202, 2021.

Bratsun, D. A., Oschepkov, V. O., Mosheva, E. A., and Siraev, R. R.: The effect of concentration-dependent diffusion on double-diffusive instability, Phys. Fluids, 34, 034 112, https://doi.org/10.1063/5.0079850, 2022.

Brezzi, F., Douglad, J., and Marini, L. D.: Two families of mixed finite elements for second order elliptic problems, Numerische Mathematik, 47, 217–235, https://doi.org/10.1007/BF01389710, 1985.

Bringedal, C., Schollenberger, T., Pieters, G. J. M., van Duijn, C. J., and Helmig, R.: Evaporation-Driven Density Instabilities in Saturated Porous Media, Transport Porous Med., 143, 297–341, https://doi.org/10.1007/s11242-022-01772-w, 2022.

Brouzet, C., Méheust, Y., and Meunier, P.: CO_2 convective dissolution in a three-dimensional granular porous medium: An experimental study, Phys. Rev. Fluids, 7, 033 802, https://doi.org/10.1103/PhysRevFluids.7.033802, 2022.

Budroni, M. A.: Cross-diffusion-driven hydrodynamic instabilities in a double-layer system: General classification and nonlinear simulations, Phys. Rev. E, 92, 063 007, https://doi.org/10.1103/PhysRevE.92.063007, 2015.

Cappellen, P. V. and Gaillard, J.-F.: Chapter 8. Biogeochemical Dynamics in Aquatic Sediments, in: Reactive Transport in Porous Media, edited by Lichtner, P. C., Steefel, C. I., and Oelkers, E. H., pp. 335–376, De Gruyter, https://doi.org/10.1515/9781501509797-011, 1996.

Carrayrou, J., Mosé, R., and Behra, P.: Operator-splitting procedures for reactive transport and comparison of mass balance errors, J. Contam. Hydrol., 68, 239–268, https://doi.org/10.1016/S0169-7722(03)00141-4, 2004.

Carrera, J., Saaltink, M. W., Soler-Sagarra, J., Wang, J., and Valhondo, C.: Reactive Transport: A Review of Basic Concepts with Emphasis on Biochemical Processes, Energies, 15, https://doi.org/10.3390/en15030925, 2022.

Chang, E., Brewer, A. W., Park, D. M., Jiao, Y., and Lammers, L. N.: Selective Biosorption of Valuable Rare Earth Elements Among Co-Occurring Lanthanides, Environ. Eng. Sci., 38, 154–164, https://doi.org/10.1089/ees.2020.0291, 2021.

Cheng, C. and Milsch, H.: Permeability Variations in Illite-Bearing Sandstone: Effects of Temperature and NaCl Fluid Salinity, J. Geophys. Res.-Sol. Ea., 125, e2020JB020 122, https://doi.org/10.1029/2020JB020122, 2020.

Cherezov, I., Cardoso, S. S., and Kim, M. C.: Acceleration of convective dissolution by an instantaneous chemical reaction: A comparison of experimental and numerical results, Chem. Eng. Sci., 181, 298–310, https://doi.org/10.1016/j.ces.2018.02.005, 2018.

Citri, O., Kagan, M. L., Kosloff, R., and Avnir, D.: Evolution of chemically induced unstable density gradients near horizontal reactive interfaces, Langmuir, 6, 559–564, https://doi.org/10.1021/la00093a007, 1990.

Class, H., Weishaupt, K., and Trötschler, O.: Experimental and Simulation Study on Validating a Numerical Model for CO_2 Density-Driven Dissolution in Water, Water, 12, https://doi.org/10.3390/w12030738, 2020.

Cochepin, B., Trotignon, L., Bildstein, O., Steefel, C., Lagneau, V., and Van der lee, J.: Approaches to modelling coupled flow and reaction in a 2D cementation experiment, Adv. Water Resour., 31, 1540–1551, https://doi.org/10.1016/j.advwatres.2008.05.007, 2008.

Cogorno, J., Stolze, L., Muniruzzaman, M., and Rolle, M.: Dimensionality effects on multicomponent ionic transport and surface complexation in porous media, Geochim. Cosmochim. Ac., 318, 230–246, https://doi.org/10.1016/j.gca.2021.11.037, 2022.

Connolly, J. A. D.: The geodynamic equation of state: What and how, Geochem. Geophy. Geosy., 10, https://doi.org/10.1029/2009GC002540, 2009.

Constantin, P., Ignatova, M., and Lee, F.-N.: Interior Electroneutrality in Nernst–Planck–Navier–Stokes Systems, Archive for Rational Mechanics and Analysis, 242, 1091–1118, https://doi.org/10.1007/s00205-021-01700-0, 2021.

Constantin, P., Ignatova, M., and Lee, F.-N.: Existence and stability of nonequilibrium steady states of Nernst–Planck–Navier–Stokes systems, Physica D: Nonlinear Phenomena, 442, 133 536, https://doi.org/10.1016/j.physd.2022.133536, 2022.

Crameri, F.: Scientific colour maps, https://doi.org/10.5281/zenodo.5501399, 2021.

Damiani, L. H., Kosakowski, G., Glaus, M. A., and Churakov, S. V.: "A framework for reactive transport modeling using FEniCS–Reaktoro: governing equations and benchmarking results", Computat. Geosci., 24, 1071–1085, https://doi.org/10.1007/s10596-019-09919-3, 2020.

Davis, M.: Palettable: Color palettes for Python, https://jiffyclub.github.io/palettable, last access: 2022-11-25, 2019.

De Lucia, M. and Kühn, M.: DecTree v1.0 – chemistry speedup in reactive transport simulations: purely data-driven and physics-based surrogates, Geosci. Model Dev., 14, 4713–4730, https://doi.org/10.5194/gmd-14-4713-2021, 2021.

De Lucia, M., Kühn, M., Lindemann, A., Lübke, M., and Schnor, B.: `POET` (v0.1): speedup of many-core parallel reactive transport simulations with fast DHT lookups, Geosci. Model Dev., 14, 7391–7409, https://doi.org/10.5194/gmd-14-7391-2021, 2021.

De Wit, A.: Chemo-hydrodynamic patterns in porous media, Philosophical Transactions of the Royal Society A: Mathematical, Physical and Engineering Sciences, 374, 20150 419, https://doi.org/10.1098/rsta.2015.0419, 2016.

De Wit, A.: Chemo-Hydrodynamic Patterns and Instabilities, Annu. Rev. Fluid Mech., 52, 531–555, https://doi.org/10.1146/annurev-fluid-010719-060349, 2020.

Dickinson, E. J. F., Limon-Peterson, J. G., and Compton, R. G.: The electroneutrality assumption in electrochemistry, J. Solid State Electr., 15, 1335–1345, https://doi.org/10.1007/s10008-011-1323-x, 2011.

Donea, J. and Huerta, A.: Finite Element Methods for Flow Problems, John Wiley & Sons, https://doi.org/10.1002/0470013826, 2003.

Dreyer, W., Guhlke, C., and Müller, R.: Overcoming the shortcomings of the Nernst–Planck model, Phys. Chem. Chem. Phys., 15, 7075–7086, https://doi.org/10.1039/C3CP44390F, 2013.

Drummond, S.: Boiling and mixing of hydrothermal fluids: chemical effects on mineral precipitation, Ph.D. thesis, Pennsylvania State University, https://www.proquest.com/docview/303157791, last access: 2022-11-21, 1981.

645 Eckert, K. and Grahn, A.: Plume and Finger Regimes Driven by an Exothermic Interfacial Reaction, Phys. Rev. Lett., 82, 4436–4439, https://doi.org/10.1103/PhysRevLett.82.4436, 1999.

Eckert, K., Acker, M., and Shi, Y.: Chemical pattern formation driven by a neutralization reaction. I. Mechanism and basic features, Phys. Fluids, 16, 385–399, https://doi.org/10.1063/1.1636160, 2004.

Ezekiel, J., Adams, B. M., Saar, M. O., and Ebigbo, A.: Numerical analysis and optimization of the performance of 650 CO_2-Plume Geothermal (CPG) production wells and implications for electric power generation, Geothermics, 98, 102 270, https://doi.org/10.1016/j.geothermics.2021.102270, 2022.

Filipek, R., Kalita, P., Sapa, L., and Szyszkiewicz, K.: On local weak solutions to Nernst–Planck–Poisson system, Appl. Anal., 96, 2316–2332, https://doi.org/10.1080/00036811.2016.1221941, 2017.

Flavell, A., Machen, M., Eisenberg, B., Kabre, J., Liu, C., and Li, X.: A conservative finite difference scheme for Poisson–Nernst–Planck 655 equations, J. Comput. Phys., 13, 235–249, https://doi.org/10.1007/s10825-013-0506-3, 2014.

Fleming, M. R., Adams, B. M., Kuehn, T. H., Bielicki, J. M., and Saar, M. O.: Increased Power Generation due to Exothermic Water Exsolution in CO_2 Plume Geothermal (CPG) Power Plants, Geothermics, 88, 101 865, https://doi.org/10.1016/j.geothermics.2020.101865, 2020.

Fortin, M. and Glowinski, R.: Augmented Lagrangian Methods in Quadratic Programming, in: Augmented Lagrangian Methods: Applica- 660 tions to the Numerical Solution of Boundary-Value Problems, edited by Fortin, M. and Glowinski, R., vol. 15 of *Studies in Mathematics and Its Applications*, chap. 1, pp. 1–46, Elsevier, https://doi.org/10.1016/S0168-2024(08)70026-2, 1983.

Frank, F., Ray, N., and Knabner, P.: Numerical investigation of homogenized Stokes–Nernst–Planck–Poisson systems, Computing and Visualization in Science, 14, 385–400, https://doi.org/10.1007/s00791-013-0189-0, 2011.

Frizon, F., Lorente, S., Ollivier, J., and Thouvenot, P.: Transport model for the nuclear decontamination of cementitious materials, Comp. 665 Mater. Sci., 27, 507–516, https://doi.org/10.1016/S0927-0256(03)00051-X, 2003.

Fu, B., Zhang, R., Liu, J., Cui, L., Zhu, X., and Hao, D.: Simulation of CO_2 Rayleigh Convection in Aqueous Solutions of NaCl, KCl, $MgCl_2$ and $CaCl_2$ using Lattice Boltzmann Method, Int. J. Greenh. Gas Con., 98, 103 066, https://doi.org/10.1016/j.ijggc.2020.103066, 2020.

Fu, B., Zhang, R., Xiao, R., Cui, L., Liu, J., Zhu, X., and Hao, D.: Simulation of interfacial mass transfer process accompanied by Rayleigh convection in NaCl solution, Int. J. Greenh. Gas Con., 106, 103 281, https://doi.org/10.1016/j.ijggc.2021.103281, 2021.

670 Gamazo, P., Slooten, L. J., Carrera, J., Saaltink, M. W., Bea, S., and Soler, J.: PROOST: object-oriented approach to multiphase reactive transport modeling in porous media, J. Hydroinform., 18, 310–328, https://doi.org/10.2166/hydro.2015.126, 2015.

Giambalvo, E. R., Steefel, C. I., Fisher, A. T., Rosenberg, N. D., and Wheat, C. G.: Effect of fluid-sediment reaction on hydrothermal fluxes of major elements, eastern flank of the Juan de Fuca Ridge, Geochim. Cosmochim. Ac., 66, 1739–1757, https://doi.org/10.1016/S0016-7037(01)00878-X, 2002.

675 Gimmi, T. and Alt-Epping, P.: Simulating Donnan equilibria based on the Nernst-Planck equation, Geochim. Cosmochim. Ac., 232, 1–13, https://doi.org/10.1016/j.gca.2018.04.003, 2018.

Glaus, M. A., Birgersson, M., Karnland, O., and Van Loon, L. R.: Seeming Steady-State Uphill Diffusion of 22Na+ in Compacted Montmorillonite, Environ. Sci. Technol., 47, 11 522–11 527, https://doi.org/10.1021/es401968c, 2013.

Golparvar, A., Kästner, M., and Thullner, M.: P3D-BRNS v1.0.0: A Three-dimensional, Multiphase, Multicomponent, Pore-scale Reactive Transport Modelling Package for Simulating Biogeochemical Processes in Subsurface Environments, Geoscientific Model Development Discussions, 2022, 1–28, https://doi.org/10.5194/gmd-2022-86, 2022.

Grimm Lima, M., Schädle, P., Green, C. P., Vogler, D., Saar, M. O., and Kong, X.-Z.: Permeability Impairment and Salt Precipitation Patterns During CO_2 Injection Into Single Natural Brine-Filled Fractures, Water Resour. Res., 56, e2020WR027213, https://doi.org/10.1029/2020WR027213, 2020.

Guo, R., Sun, H., Zhao, Q., Li, Z., Liu, Y., and Chen, C.: A Novel Experimental Study on Density-Driven Instability and Convective Dissolution in Porous Media, Geophys. Res. Lett., 48, e2021GL095619, https://doi.org/10.1029/2021GL095619, 2021.

Harrower, M. and Brewer, C. A.: ColorBrewer.org: An Online Tool for Selecting Colour Schemes for Maps, The Cartographic Journal, 40, 27–37, https://doi.org/10.1179/000870403235002042, 2003.

Hassanpouryouzband, A., Adie, K., Cowen, T., Thaysen, E. M., Heinemann, N., Butler, I. B., Wilkinson, M., and Edlmann, K.: Geological Hydrogen Storage: Geochemical Reactivity of Hydrogen with Sandstone Reservoirs, ACS Energy Letters, 7, 2203–2210, https://doi.org/10.1021/acsenergylett.2c01024, 2022.

Hejazi, S. H. and Azaiez, J.: Nonlinear simulation of transverse flow interactions with chemically driven convective mixing in porous media, Water Resour. Res., 49, 4607–4618, https://doi.org/10.1002/wrcr.20298, 2013.

Helgeson, H. C. and Kirkham, D. H.: Theoretical prediction of the thermodynamic behavior of aqueous electrolytes at high pressures and temperatures; I, Summary of the thermodynamic/electrostatic properties of the solvent, Am. J. Sci., 274, 1089–1198, https://doi.org/10.2475/ajs.274.10.1089, 1974a.

Helgeson, H. C. and Kirkham, D. H.: Theoretical prediction of the thermodynamic behavior of aqueous electrolytes at high pressures and temperatures; II, Debye-Huckel parameters for activity coefficients and relative partial molal properties, Am. J. Sci., 274, 1199–1261, https://doi.org/10.2475/ajs.274.10.1199, 1974b.

Helgeson, H. C. and Kirkham, D. H.: Theoretical prediction of the thermodynamic properties of aqueous electrolytes at high pressures and temperatures. III. Equation of state for aqueous species at infinite dilution, Am. J. Sci., 276, 97–240, https://doi.org/10.2475/ajs.276.2.97, 1976.

Helgeson, H. C., Kirkham, D. H., and Flowers, G. C.: Theoretical prediction of the thermodynamic behavior of aqueous electrolytes by high pressures and temperatures; IV, Calculation of activity coefficients, osmotic coefficients, and apparent molal and standard and relative partial molal properties to 600°C and 5 KB, Am. J. Sci., 281, 1249–1516, https://doi.org/10.2475/ajs.281.10.1249, 1981.

Herz, M., Ray, N., and Knabner, P.: Existence and uniqueness of a global weak solution of a Darcy-Nernst-Planck-Poisson system, GAMM-Mitteilungen, 35, 191–208, https://doi.org/10.1002/gamm.201210013, 2012.

Hingerl, F. F.: Geothermal electrolyte solutions: thermodynamic model and computational fitting framework development, Ph.D. thesis, ETH Zurich, https://doi.org/10.3929/ethz-a-009755479, 2012.

Huang, Y., Shao, H., Wieland, E., Kolditz, O., and Kosakowski, G.: A new approach to coupled two-phase reactive transport simulation for long-term degradation of concrete, Constr. Build. Mater., 190, 805–829, https://doi.org/10.1016/j.conbuildmat.2018.09.114, 2018.

Hunter, J. D.: Matplotlib: A 2D graphics environment, Comput. Sci. Eng., 9, 90–95, https://doi.org/10.1109/MCSE.2007.55, 2007.

Ignatova, M. and Shu, J.: Global Smooth Solutions of the Nernst–Planck–Darcy System, J. Math. Fluid Mech., 24, https://doi.org/10.1007/s00021-022-00666-7, 2022.

Illés, B., Medgyes, B., Dušek, K., Bušek, D., Skwarek, A., and Géczy, A.: Numerical simulation of electrochemical migration of Cu based on the Nernst-Plank equation, Int. J. Heat Mass Tran., 184, 122268, https://doi.org/10.1016/j.ijheatmasstransfer.2021.122268, 2022.

Izumoto, S., Huisman, J. A., Zimmermann, E., Heyman, J., Gomez, F., Tabuteau, H., Laniel, R., Vereecken, H., Méheust, Y., and Le Borgne, T.: Pore-Scale Mechanisms for Spectral Induced Polarization of Calcite Precipitation Inferred from Geo-Electrical Millifluidics, Environ. Sci. Technol., 56, 4998–5008, https://doi.org/10.1021/acs.est.1c07742, 2022.

720 Jasielec, J. J.: Electrodiffusion Phenomena in Neuroscience and the Nernst–Planck–Poisson Equations, Electrochem, 2, 197–215, https://doi.org/10.3390/electrochem2020014, 2021.

Jiang, L., Wang, S., Liu, D., Zhang, W., Lu, G., Liu, Y., and Zhao, J.: Change in Convection Mixing Properties with Salinity and Temperature: CO_2 Storage Application, Polymers, 12, https://doi.org/10.3390/polym12092084, 2020.

Johnson, J. W., Oelkers, E. H., and Helgeson, H. C.: SUPCRT92: A software package for calculating the standard molal thermodynamic

725 properties of minerals, gases, aqueous species, and reactions from 1 to 5000 bar and 0 to 1000°C, Comput. Geosci., 18, 899–947, https://doi.org/10.1016/0098-3004(92)90029-Q, 1992.

Jotkar, M., De Wit, A., and Rongy, L.: Control of chemically driven convective dissolution by differential diffusion effects, Phys. Rev. Fluids, 6, 053 504, https://doi.org/10.1103/PhysRevFluids.6.053504, 2021.

Jyoti, A. and Haese, R. R.: Validation of a multicomponent reactive-transport model at pore scale based on the coupling of COMSOL and

730 PhreeqC, Comput. Geosci., 156, 104 870, https://doi.org/10.1016/j.cageo.2021.104870, 2021.

Kadeethum, T., Lee, S., Ballarin, F., Choo, J., and Nick, H.: A locally conservative mixed finite element framework for coupled hydro-mechanical–chemical processes in heterogeneous porous media, Comput. Geosci., 152, 104 774, https://doi.org/10.1016/j.cageo.2021.104774, 2021.

Kim, M. C.: Effect of the irreversible $A+B \to C$ reaction on the onset and the growth of the buoyancy-driven instability in a porous medium:

735 Asymptotic, linear, and nonlinear stability analyses, Phys. Rev. Fluids, 4, 073 901, https://doi.org/10.1103/PhysRevFluids.4.073901, 2019.

Kirby, B. J.: Species and Charge Transport, in: Micro- and Nanoscale Fluid Mechanics: Transport in Microfluidic Devices, pp. 250–264, Cambridge University Press, https://doi.org/10.1017/CBO9780511760723.013, 2010.

Kneafsey, T. J. and Pruess, K.: Laboratory Flow Experiments for Visualizing Carbon Dioxide-Induced, Density-Driven Brine Convection, Transport porous med., 82, 123–139, https://doi.org/10.1007/s11242-009-9482-2, 2010.

740 Kontturi, K., Murtomäki, L., and A. Manzanares, J.: Ionic Transport Processes: in Electrochemistry and Membrane Science, Oxford University Press, https://doi.org/10.1093/acprof:oso/9780199533817.001.0001, 2008.

Kovtunenko, V. A. and Zubkova, A. V.: Existence and two-scale convergence of the generalised Poisson–Nernst–Planck problem with nonlinear interface conditions, Eur. J. Appl. Math., 32, 683–710, https://doi.org/10.1017/S095679252000025X, 2021.

Kulik, D. A., Wagner, T., Dmytrieva, S. V., Kosakowski, G., Hingerl, F. F., Chudnenko, K. V., and Berner, U. R.: Theoretical prediction of the

745 thermodynamic properties of aqueous electrolytes at high pressures and temperatures. III. Equation of state for aqueous species at infinite dilution, Computat. Geosci., 17, 1–24, https://doi.org/10.1007/s10596-012-9310-6, 2012.

Kwon, O., Herbert, B. E., and Kronenberg, A. K.: Permeability of illite-bearing shale: 2. Influence of fluid chemistry on flow and functionally connected pores, J. Geophys. Res.-Sol. Ea., 109, https://doi.org/10.1029/2004JB003055, 2004a.

Kwon, O., Kronenberg, A. K., Gangi, A. F., Johnson, B., and Herbert, B. E.: Permeability of illite-bearing shale: 1. Anisotropy and effects

750 of clay content and loading, J. Geophys. Res.-Sol. Ea., 109, https://doi.org/10.1029/2004JB003052, 2004b.

Kyas, S., Volpatto, D., Saar, M. O., and Leal, A. M. M.: Accelerated reactive transport simulations in heterogeneous porous media using Reaktoro and Firedrake, Computat. Geosci., 26, 295–327, https://doi.org/10.1007/s10596-021-10126-2, 2022.

Laloy, E. and Jacques, D.: Speeding Up Reactive Transport Simulations in Cement Systems by Surrogate Geochemical Modeling: Deep Neural Networks and k-Nearest Neighbors, Transport Porous Med., https://doi.org/10.1007/s11242-022-01779-3, 2022.

755   Leal, A. M., Blunt, M. J., and LaForce, T. C.: Efficient chemical equilibrium calculations for geochemical speciation and reactive transport modelling, Geochim. Cosmochim. Ac., 131, 301–322, https://doi.org/10.1016/j.gca.2014.01.038, 2014.

Leal, A. M., Kulik, D. A., and Kosakowski, G.: Computational methods for reactive transport modeling: A Gibbs energy minimization approach for multiphase equilibrium calculations, Adv. Water Resour., 88, 231–240, https://doi.org/10.1016/j.advwatres.2015.11.021, 2016.

760   Leal, A. M. M.: Reaktoro for Python and C++, https://reaktoro.org/, last access: 2022-11-21, 2022.

Leal, A. M. M., Kulik, D. A., Smith, W. R., and Saar, M. O.: An overview of computational methods for chemical equilibrium and kinetic calculations for geochemical and reactive transport modeling, Pure Appl. Chem., 89, 597–643, https://doi.org/doi:10.1515/pac-2016-1107, 2017.

Leal, A. M. M., Kyas, S., and Kulik, Dmitrii A.and Saar, M. O.: Accelerating Reactive Transport Modeling: On-Demand Machine Learning

765   Algorithm for Chemical Equilibrium Calculations, Transport Porous Med., 133, 161–204, https://doi.org/10.1007/s11242-020-01412-1, 2020.

Lee, F.-N.: Global Regularity for Nernst–Planck–Navier–Stokes Systems with mixed boundary conditions, Nonlinearity, 36, 255–286, https://doi.org/10.1088/1361-6544/aca50f, 2022.

Lees, E., Rokkam, S., Shanbhag, S., and Gunzburger, M.: The electroneutrality constraint in nonlocal models, J. Chem. Phys., 147, 124 102,

770   https://doi.org/10.1063/1.5003915, 2017.

Lemaigre, L., Budroni, M. A., Riolfo, L. A., Grosfils, P., and De Wit, A.: Asymmetric Rayleigh-Taylor and double-diffusive fingers in reactive systems, Phys. Fluids, 25, 014 103, https://doi.org/10.1063/1.4774321, 2013.

Lemmon, E. W. and Harvey, A. H.: Thermophysical Properties of Water and Steam, in: CRC Handbook of Chemistry and Physics, edited by Rumble, J. R., CRC Press/Taylor & Francis, Boca Raton, FL, 102nd edn., 2021.

775   Lichtner, P. C.: Continuum model for simultaneous chemical reactions and mass transport in hydrothermal systems, Geochim. Cosmochim. Ac., 49, 779–800, https://doi.org/10.1016/0016-7037(85)90172-3, 1985.

Lichtner, P. C.: Principles and Practice of Reactive Transport Modeling, MRS Proceedings, 353, 117–130, https://doi.org/10.1557/PROC-353-117, 1994.

Ling, F. T., Plattenberger, D. A., Peters, C. A., and Clarens, A. F.: Sealing Porous Media through Calcium Silicate Reactions with CO_2 to

780   Enhance the Security of Geologic Carbon Sequestration, Environ. Eng. Sci., 38, 127–142, https://doi.org/10.1089/ees.2020.0369, 2021.

Liu, H. and Maimaitiyiming, W.: Efficient, Positive, and Energy Stable Schemes for Multi-D Poisson–Nernst–Planck Systems, J. Sci. Comput., 148, 92, https://doi.org/10.1007/s10915-021-01503-1, 2021.

Liu, H., Wang, Z., Yin, P., and Yu, H.: Positivity-preserving third order DG schemes for Poisson–Nernst–Planck equations, J. Comput. Phys., 452, 110 777, https://doi.org/10.1016/j.jcp.2021.110777, 2022.

785   Luhmann, A. J., Kong, X.-Z., Tutolo, B. M., Ding, K., Saar, M. O., and Seyfried, W. E. J.: Permeability Reduction Produced by Grain Reorganization and Accumulation of Exsolved CO_2 during Geologic Carbon Sequestration: A New CO_2 Trapping Mechanism, Environ. Sci. Technol., 47, 242–251, https://doi.org/10.1021/es3031209, 2013.

López-Vizcaíno, R., Cabrera, V., Sprocati, R., Muniruzzaman, M., Rolle, M., Navarro, V., and Yustres, Á.: A modeling approach for elec­trokinetic transport in double-porosity media, Electrochim. Acta, 431, 141 139, https://doi.org/10.1016/j.electacta.2022.141139, 2022.

790   Maes, J. and Menke, H. P.: GeoChemFoam: Direct Modelling of Multiphase Reactive Transport in Real Pore Geometries with Equilibrium Reactions, Transport Porous Med., 139, 271–299, https://doi.org/10.1007/s11242-021-01661-8, 2021.

Mahmoodpour, S., Rostami, B., Soltanian, M. R., and Amooie, M. A.: Effect of brine composition on the onset of convection during CO_2 dissolution in brine, Comput. Geosci., 124, 1–13, https://doi.org/10.1016/j.cageo.2018.12.002, 2019.

Mahmoodpour, S., Amooie, M. A., Rostami, B., and Bahrami, F.: Effect of gas impurity on the convective dissolution of CO_2 in porous media, Energy, 199, 117 397, https://doi.org/10.1016/j.energy.2020.117397, 2020.

Marini, L.: Chapter 4 - The Aqueous Electrolyte Solution, in: Geological Sequestration of Carbon Dioxide, vol. 11 of *Developments in Geochemistry*, pp. 53–77, Elsevier, https://doi.org/10.1016/S0921-3198(06)80024-4, 2007.

Martinez, M. J. and Hesse, M. A.: Two-phase convective CO_2 dissolution in saline aquifers, Water Resour. Res., 52, 585–599, https://doi.org/10.1002/2015WR017085, 2016.

Miron, G. D., Leal, A. M. M., and Yapparova, A.: Thermodynamic Properties of Aqueous Species Calculated Using the HKF Model: How Do Different Thermodynamic and Electrostatic Models for Solvent Water Affect Calculated Aqueous Properties?, Geofluids, 2019, https://doi.org/10.1155/2019/5750390, 2019.

Mizev, A., Mosheva, E., and Bratsun, D.: Extended classification of the buoyancy-driven flows induced by a neutralization reaction in miscible fluids. Part 1. Experimental study, J. Fluid Mech., 916, A22, https://doi.org/10.1017/jfm.2021.201, 2021.

Moortgat, J., Li, M., Amooie, M. A., and Zhu, D.: A higher-order finite element reactive transport model for unstructured and fractured grids, Scientific Reports, 10, 15 572, https://doi.org/10.1038/s41598-020-72354-3, 2020.

Muniruzzaman, M. and Rolle, M.: Impact of multicomponent ionic transport on pH fronts propagation in saturated porous media, Water Resour. Res., 51, 6739–6755, https://doi.org/10.1002/2015WR017134, 2015.

Muniruzzaman, M. and Rolle, M.: Experimental investigation of the impact of compound-specific dispersion and electrostatic interactions on transient transport and solute breakthrough, Water Resour. Res., 53, 1189–1209, https://doi.org/10.1002/2016WR019727, 2017.

Muniruzzaman, M. and Rolle, M.: Multicomponent Ionic Transport Modeling in Physically and Electrostatically Hetero­geneous Porous Media With PhreeqcRM Coupling for Geochemical Reactions, Water Resour. Res., 55, 11 121–11 143, https://doi.org/10.1029/2019WR026373, 2019.

Muniruzzaman, M., Haberer, C. M., Grathwohl, P., and Rolle, M.: Multicomponent ionic dispersion during transport of elec­trolytes in heterogeneous porous media: Experiments and model-based interpretation, Geochim. Cosmochim. Ac., 141, 656–669, https://doi.org/10.1016/j.gca.2014.06.020, 2014.

Naumov, D. Y., Bilke, L., Fischer, T., Rink, K., Wang, W., Watanabe, N., Lu, R., Grunwald, N., Zill, F., Buchwald, J., Huang, Y., Bathmann, J., Chen, C., Chen, S., Meng, B., Shao, H., Kern, D., Yoshioka, K., Garibay Rodriguez, J., Miao, X., Parisio, F., Silbermann, C., Thiedau, J., Walther, M., Kaiser, S., Boog, J., Zheng, T., Meisel, T., and Ning, Z.: OpenGeoSys, https://doi.org/10.5281/zenodo.6405711, 2022.

Neufeld, J. A., Hesse, M. A., Riaz, A., Hallworth, M. A., Tchelepi, H. A., and Huppert, H. E.: Convective dissolution of carbon dioxide in saline aquifers, Geophys. Res. Lett., 37, https://doi.org/10.1029/2010GL044728, 2010.

Nishikata, E., Ishii, T., and Ohta, T.: Viscosities of aqueous hydrochloric acid solutions, and densities and viscosities of aqueous hydroiodic acid solutions, J. Chem. Eng. Data, 26, 254–256, https://doi.org/10.1021/je00025a008, 1981.

Oliveira, T. D. S., Blunt, M. J., and Bijeljic, B.: Modelling of multispecies reactive transport on pore-space images, Adv. Water Resour., 127, 192–208, https://doi.org/10.1016/j.advwatres.2019.03.012, 2019.

Onsager, L.: Theories and Problems of Liquid Diffusion, Ann. NY. Acad. Sci., 46, 241–265, https://doi.org/10.1111/j.1749-6632.1945.tb36170.x, 1945.

Pamukcu, S.: Electrochemical Transport and Transformations, in: Electrochemical Remediation Technologies for Polluted Soils, Sediments and Groundwater, edited by Reddy, K. R. and Cameselle, C., chap. 2, pp. 29–64, John Wiley & Sons, Ltd, https://doi.org/10.1002/9780470523650.ch2, 2009.

Parkhurst, D. L. and Appelo, C. A. J.: Description of input and examples for PHREEQC version 3–A computer program for speciation, batch-reaction, one-dimensional transport, and inverse geochemical calculations, http://pubs.usgs.gov/tm/06/a43, 2013.

Permann, C. J., Gaston, D. R., Andrš, D., Carlsen, R. W., Kong, F., Lindsay, A. D., Miller, J. M., Peterson, J. W., Slaughter, A. E., Stogner, R. H., and Martineau, R. C.: MOOSE: Enabling massively parallel multiphysics simulation, SoftwareX, 11, 100 430, https://doi.org/10.1016/j.softx.2020.100430, 2020.

Petersen, S. and Hack, K.: The thermochemistry library ChemApp and its applications, Int. J. Mater. Res., 98, 935–945, https://doi.org/10.3139/146.101551, 2007.

Pogge von Strandmann, P. A. E., Burton, K. W., Snæbjörnsdóttir, S. O., Sigfússon, B., Aradóttir, E. S., Gunnarsson, I., Alfredsson, H. A., Mesfin, K. G., Oelkers, E. H., and Gislason, S. R.: Rapid CO_2 mineralisation into calcite at the CarbFix storage site quantified using calcium isotopes, Nat. Commun., 10, https://doi.org/10.1038/s41467-019-10003-8, 2019.

Priya, P., Kuhlman, K. L., and Aluru, N. R.: Pore-Scale Modeling of Electrokinetics in Geomaterials, Transport Porous Med., 137, 651–666, https://doi.org/10.1007/s11242-021-01581-7, 2021.

Randolph, J. B. and Saar, M. O.: Combining geothermal energy capture with geologic carbon dioxide sequestration, Geophys. Res. Lett., 38, https://doi.org/10.1029/2011GL047265, 2011.

Rasouli, P., Steefel, C. I., Mayer, K. U., and Rolle, M.: Benchmarks for multicomponent diffusion and electrochemical migration, Computat. Geosci., 19, 523–533, https://doi.org/10.1007/s10596-015-9481-z, 2015.

Rathgeber, F., Ham, D. A., Mitchell, L., Lange, M., Luporini, F., Mcrae, A. T. T., Bercea, G.-T., Markall, G. R., and Kelly, P. H. J.: Firedrake: Automating the Finite Element Method by Composing Abstractions, ACM Trans. Math. Softw., 43, https://doi.org/10.1145/2998441, 2016.

Raviart, P. A. and Thomas, J. M.: A mixed finite element method for 2-nd order elliptic problems, in: Mathematical Aspects of Finite Element Methods, edited by Galligani, I. and Magenes, E., pp. 292–315, Springer Berlin Heidelberg, Berlin, Heidelberg, https://doi.org/10.1007/BFb0064470, 1977.

Ray, N., Muntean, A., and Knabner, P.: Rigorous homogenization of a Stokes–Nernst–Planck–Poisson system, J. Math. Anal. Appl., 390, 374–393, https://doi.org/10.1016/j.jmaa.2012.01.052, 2012a.

Ray, N., van Noorden, T., Frank, F., and Knabner, P.: Multiscale Modeling of Colloid and Fluid Dynamics in Porous Media Including an Evolving Microstructure, Transport Porous Med., 95, 669–696, https://doi.org/10.1007/s11242-012-0068-z, 2012b.

Reddy, K. R. and Cameselle, C.: Electrochemical Remediation Technologies for Polluted Soils, Sediments and Groundwater, John Wiley & Sons, https://doi.org/10.1002/9780470523650, 2009.

Reed, M. H.: Calculation of Simultaneous Chemical Equilibria in Aqueous-Mineral-Gas Systems and its Application to Modeling Hydrothermal Processes, in: Techniques in Hydrothermal Ore Deposits Geology, Society of Economic Geologists, https://doi.org/10.5382/Rev.10.05, 1998.

Revil, A. and Leroy, P.: Constitutive equations for ionic transport in porous shales, J. Geophys. Res.-Sol. Ea., 109, https://doi.org/10.1029/2003JB002755, 2004.

Rivera, F. F., Pérez, T., Castañeda, L. F., and Nava, J. L.: Mathematical modeling and simulation of electrochemical reactors: A critical review, Chem. Eng. Sci., 239, 116 622, https://doi.org/10.1016/j.ces.2021.116622, 2021.

Rolle, M., Muniruzzaman, M., Haberer, C. M., and Grathwohl, P.: Coulombic effects in advection-dominated transport of electrolytes in porous media: Multicomponent ionic dispersion, Geochim. Cosmochim. Ac., 120, 195–205, https://doi.org/10.1016/j.gca.2013.06.031, 2013.

Rolle, M., Sprocati, R., Masi, M., Jin, B., and Muniruzzaman, M.: Nernst-Planck-based Description of Transport, Coulombic Interactions, and Geochemical Reactions in Porous Media: Modeling Approach and Benchmark Experiments, Water Resour. Res., 54, 3176–3195, https://doi.org/10.1002/2017WR022344, 2018.

Rolle, M., Albrecht, M., and Sprocati, R.: Impact of solute charge and diffusion coefficient on electromigration and mixing in porous media, J. Contam. Hydrol., 244, 103 933, https://doi.org/10.1016/j.jconhyd.2021.103933, 2022.

Rubinstein, I.: Locally Electro-Neutral Electro-Diffusion without Electric Current, in: Electro-Diffusion of Ions, chap. 3, pp. 59–103, Society for Industrial and Applied Mathematics, https://doi.org/10.1137/1.9781611970814.ch3, 1990.

Samson, E. and Marchand, J.: Modeling the transport of ions in unsaturated cement-based materials, Comput. Struct., 85, 1740–1756, https://doi.org/10.1016/j.compstruc.2007.04.008, 2007.

Sapa, L., Bożek, B., Tkacz–Śmiech, K., Zajusz, M., and Danielewski, M.: Interdiffusion in many dimensions: mathematical models, numerical simulations and experiment, Math. Mech. Solids, 25, 2178–2198, https://doi.org/10.1177/1081286520923376, 2020.

Savino, M., Lévy-Leduc, C., Leconte, M., and Cochepin, B.: An active learning approach for improving the performance of equilibrium based chemical simulations, Computat. Geosci., 26, 365–380, https://doi.org/10.1007/s10596-022-10130-0, 2022.

Shafabakhsh, P., Ataie-Ashtiani, B., Simmons, C. T., Younes, A., and Fahs, M.: Convective-reactive transport of dissolved CO_2 in fractured-geological formations, Int. J. Greenh. Gas Con., 109, 103 365, https://doi.org/10.1016/j.ijggc.2021.103365, 2021.

Shao, H., Dmytrieva, S. V., Kolditz, O., Kulik, D. A., Pfingsten, W., and Kosakowski, G.: Modeling reactive transport in non-ideal aqueous–solid solution system, Appl. Geochem., 24, 1287–1300, https://doi.org/10.1016/j.apgeochem.2009.04.001, 2009.

Shen, J. and Xu, J.: Unconditionally positivity preserving and energy dissipative schemes for Poisson–Nernst–Planck equations, Numerische Mathematik, 148, 671–697, https://doi.org/10.1007/s00211-021-01203-w, 2021.

Sin, I. and Corvisier, J.: Multiphase Multicomponent Reactive Transport and Flow Modeling, Rev. Mineral. Geochem., 85, 143–195, https://pubs.geoscienceworld.org/msa/rimg/article/85/1/143/573304/Multiphase-Multicomponent-Reactive-Transport-and, last access: 2022-11-28, 2019.

Singh, M., Chaudhuri, A., Chu, S. P., Stauffer, P. H., and Pawar, R. J.: Analysis of evolving capillary transition, gravitational fingering, and dissolution trapping of CO_2 in deep saline aquifers during continuous injection of supercritical CO_2, Int. J. Greenh. Gas Con., 82, 281–297, https://doi.org/10.1016/j.ijggc.2019.01.014, 2019.

Sipos, P. M., Hefter, G., and May, P. M.: Viscosities and Densities of Highly Concentrated Aqueous MOH Solutions ($M^+ = Na^+$, K^+, Li^+, Cs^+, $(CH_3)_4N^+$) at 25.0 °C, J. Chem. Eng. Data, 45, 613–617, https://doi.org/10.1021/je000019h, 2000.

Soltanian, M. R., Hajirezaie, S., Hosseini, S. A., Dashtian, H., Amooie, M. A., Meyal, A., Ershadnia, R., Ampomah, W., Islam, A., and Zhang, X.: Multicomponent reactive transport of carbon dioxide in fluvial heterogeneous aquifers, J. Nat. Gas Sci. Eng., 65, 212–223, https://doi.org/10.1016/j.jngse.2019.03.011, 2019.

Song, Z., Cao, X., and Huang, H.: Electroneutral models for dynamic Poisson-Nernst-Planck systems, Phys. Rev. E, 97, 012 411, https://doi.org/10.1103/PhysRevE.97.012411, 2018.

Sprocati, R. and Rolle, M.: Integrating Process-Based Reactive Transport Modeling and Machine Learning for Electrokinetic Remediation of Contaminated Groundwater, Water Resour. Res., 57, e2021WR029 959, https://doi.org/10.1029/2021WR029959, 2021.

Sprocati, R. and Rolle, M.: On the interplay between electromigration and electroosmosis during electrokinetic transport in heterogeneous porous media, Water Res., 213, 118 161, https://doi.org/10.1016/j.watres.2022.118161, 2022.

905 Sprocati, R., Masi, M., Muniruzzaman, M., and Rolle, M.: Modeling electrokinetic transport and biogeochemical reactions in porous media: A multidimensional Nernst–Planck–Poisson approach with PHREEQC coupling, Adv. Water Resour., 127, 134–147, https://doi.org/10.1016/j.advwatres.2019.03.011, 2019.

Steefel, C. I. and MacQuarrie, K. T. B.: Chapter 2. Approaches to modeling of reactive transport in porous media, in: Reactive Transport in Porous Media, edited by Lichtner, P. C., Steefel, C. I., and Oelkers, E. H., pp. 83–130, De Gruyter, https://doi.org/10.1515/9781501509797-

910 005, 2018.

Tabrizinejadas, S., Carrayrou, J., Saaltink, M. W., Baalousha, H. M., and Fahs, M.: On the Validity of the Null Current Assumption for Modeling Sorptive Reactive Transport and Electro-Diffusion in Porous Media, Water, 13, https://doi.org/10.3390/w13162221, 2021.

Teng, Y., Wang, P., Jiang, L., Liu, Y., and Wei, Y.: New Spectrophotometric Method for Quantitative Characterization of Density-Driven Convective Instability, Polymers, 13, https://doi.org/10.3390/polym13040661, 2021.

915 Thieulot, C. and Bangerth, W.: On the choice of finite element for applications in geodynamics, Solid Earth, 13, 229–249, https://doi.org/10.5194/se-13-229-2022, 2022.

Thomas, C., Loodts, V., Rongy, L., and De Wit, A.: Convective dissolution of CO_2 in reactive alkaline solutions: Active role of spectator ions, Int. J. Greenh. Gas Con., 53, 230–242, https://doi.org/10.1016/j.ijggc.2016.07.034, 2016.

Thomas, C., Dehaeck, S., and De Wit, A.: Convective dissolution of CO_2 in water and salt solutions, Int. J. Greenh. Gas Con., 72, 105–116,

920 https://doi.org/10.1016/j.ijggc.2018.01.019, 2018.

Tournassat, C. and Steefel, C. I.: Modeling diffusion processes in the presence of a diffuse layer at charged mineral surfaces: a benchmark exercise, Computat. Geosci., 25, 1319–1336, https://doi.org/10.1007/s10596-019-09845-4, 2021.

Tournassat, C., Steefel, C. I., and Gimmi, T.: Solving the Nernst-Planck Equation in Heterogeneous Porous Media With Finite Volume Methods: Averaging Approaches at Interfaces, Water Resour. Res., 56, e2019WR026 832, https://doi.org/10.1029/2019WR026832, 2020.

925 Trevelyan, P. M. J., Almarcha, C., and De Wit, A.: Buoyancy-driven instabilities of miscible two-layer stratifications in porous media and Hele-Shaw cells, J. Fluid Mech., 670, 38–65, https://doi.org/10.1017/S0022112010005008, 2011.

Tsinober, A., Rosenzweig, R., Class, H., Helmig, R., and Shavit, U.: The Role of Mixed Convection and Hydrodynamic Dispersion During CO_2 Dissolution in Saline Aquifers: A Numerical Study, Water Resour. Res., 58, e2021WR030 494, https://doi.org/10.1029/2021WR030494, 2022.

930 Tutolo, B. M., Luhmann, A. J., Kong, X.-Z., Saar, M. O., and Seyfried, W. E. J.: Experimental Observation of Permeability Changes In Dolomite at CO_2 Sequestration Conditions, Environ. Sci. Technol., 48, 2445–2452, https://doi.org/10.1021/es4036946, 2014.

Uzawa, H.: Iterative methods for concave programming, in: Studies in linear and nonlinear progamming, edited by Arrow, K. J., Hurwicz, L., and Uzawa, H., vol. 15, pp. 1–46, Standford University Press, 1958.

Vanýsek, P.: Ionic Conductivity and Diffusion at Infinite Dilution, in: CRC Handbook of Chemistry and Physics, edited by Rumble, J. R.,

935 CRC Press/Taylor & Francis, Boca Raton, FL, 102nd edn., 2021.

Wang, S., Cheng, Z., Zhang, Y., Jiang, L., Liu, Y., and Song, Y.: Unstable Density-Driven Convection of CO_2 in Homogeneous and Heterogeneous Porous Media With Implications for Deep Saline Aquifers, Water Resour. Res., 57, e2020WR028 132, https://doi.org/10.1029/2020WR028132, 2021.

Weller, H. G., Tabor, G., Jasak, H., and Fureby, C.: A tensorial approach to computational continuum mechanics using object-oriented

940 techniques, Comput. Phys., 12, 620–631, https://doi.org/10.1063/1.168744, 1998.

Wilkins, A., Green, C. P., and Ennis-King, J.: An open-source multiphysics simulation code for coupled problems in porous media, Comput. Geosci., 154, 104 820, https://doi.org/10.1016/j.cageo.2021.104820, 2021.

Wolery, T. J.: EQ3/6, a software package for geochemical modeling of aqueous systems: Package overview and installation guide (Version 7.0), https://doi.org/10.2172/138894, 1992.

945 Yan, D., Pugh, M., and Dawson, F.: Adaptive time-stepping schemes for the solution of the Poisson-Nernst-Planck equations, Appl. Numer. Math., 163, 254–269, https://doi.org/10.1016/j.apnum.2021.01.018, 2021.

Yang, C. and Gu, Y.: Accelerated Mass Transfer of CO_2 in Reservoir Brine Due to Density-Driven Natural Convection at High Pressures and Elevated Temperatures, Ind. Eng. Chem. Res., 45, 2430–2436, https://doi.org/10.1021/ie050497r, 2006.

Yapparova, A., Miron, G. D., Kulik, D. A., Kosakowski, G., and Driesner, T.: An advanced reactive transport simulation scheme for hy-
950 drothermal systems modelling, Geothermics, 78, 138–153, https://doi.org/10.1016/j.geothermics.2018.12.003, 2019.

Yekta, A., Salinas, P., Hajirezaie, S., Amooie, M. A., Pain, C. C., Jackson, M. D., Jacquemyn, C., and Soltanian, M. R.: Re-active transport modeling in heterogeneous porous media with dynamic mesh optimization, Computat. Geosci., 25, 357–372, https://doi.org/10.1007/s10596-020-10009-y, 2021.

Yu, Y., Gao, W., Castel, A., Liu, A., Feng, Y., Chen, X., and Mukherjee, A.: Modelling steel corrosion under concrete non-uniformity and
955 structural defects, Cement Conrete Res., 135, 106 109, https://doi.org/10.1016/j.cemconres.2020.106109, 2020.

Zalts, A., El Hasi, C., Rubio, D., Ureña, A., and D'Onofrio, A.: Pattern formation driven by an acid-base neutralization reaction in aqueous media in a gravitational field, Phys. Rev. E, 77, 015 304, https://doi.org/10.1103/PhysRevE.77.015304, 2008.

Zaytsev, I. D. and Aseyev, G. G.: Properties of Aqueous Solutions of Electrolytes, CRC Press, Boca Raton, FL, 1992.

Zeebe, R. E.: On the molecular diffusion coefficients of dissolved CO_2, HCO_3^-, and CO_3^{2-} and their dependence on isotopic mass, Geochim.
960 Cosmochim. Ac., 75, 2483–2498, https://doi.org/10.1016/j.gca.2011.02.010, 2011.

Zhang, Q., Tu, B., Fang, Q., and Lu, B.: A structure-preserving finite element discretization for the time-dependent Nernst-Planck equation, Journal of Applied Mathematics and Computing, 68, 1545–1564, https://doi.org/10.1007/s12190-021-01571-4, 2022.

Zhang, W., Li, Y., and Omambia, A. N.: Reactive transport modeling of effects of convective mixing on long-term CO_2 geological storage in deep saline formations, Int. J. Greenh. Gas Con., 5, 241–256, https://doi.org/10.1016/j.ijggc.2010.10.007, 2011.

965 Zhang, Z., Fu, Q., Zhang, H., Yuan, X., and Yu, K.-T.: Experimental and Numerical Investigation on Interfacial Mass Transfer Mechanism for Rayleigh Convection in Hele-Shaw Cell, Ind. Eng. Chem. Res., 59, 10 195–10 209, https://doi.org/10.1021/acs.iecr.0c01345, 2020.