Long-range transported pollution from the Middle East and its impact on carbonaceous aerosol sources over Cyprus.

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Abstract. The geographical origin and source apportionment of submicron carbonaceous aerosols (organic aerosols, OA, and black carbon, BC) have been investigated here for the first time by means of high time resolution measurements at an urban background site of Nicosia, the capital city of Cyprus, in the Eastern Mediterranean. This study covers a half-year period, encompassing both the cold and warm periods with continuous observations of the physical and chemical properties of PM1 performed with an Aerosol Chemical Speciation monitor (ACSM), an Aethalometer, accompanied by a suite of various ancillary off and on-line measurements. Carbonaceous aerosols were dominant during both seasons (cold and warm periods), with a respective contribution of 57% and 48% to PM1, respectively, and exhibited recurrent intense night-time peaks (>20–30 µg m⁻³) during the cold period associated with local domestic heating. Findings of this study show that high concentrations of sulfate (close to 3 µg m⁻³) were continuously recorded, standing among the highest ever reported for Europe and originating from the Middle East region.

Source apportionment of the OA and BC fractions was performed using the Positive Matrix Factorization (PMF) approach and the combination of two models (aethalometer model and multilinear regression), respectively. Our study revealed elevated hydrocarbon-like organic aerosol (HOA) concentrations in Nicosia (among the highest reported for a European urban background site), originating from a mixture of local and regional fossil-fuel combustion sources. Although air masses from the Middle East had a low occurrence and were observed mostly during the cold period, they were shown to strongly affect the mean concentrations levels of BC and OA in Nicosia during both seasons. Overall, the present study brings to our attention the need to further characterize primary and secondary carbonaceous aerosols in the Middle East; an undersampled region characterized by continuously increasing fossil fuel (oil and gas) emissions and extreme environmental conditions, which can contribute to photochemical aging.

1. Introduction

At the crossroads of three continents (Europe, Africa, Asia), the Eastern Mediterranean and Middle East (EMME) region faces many challenges, such as rapid population growth – with its currently 400 million inhabitants – as well as political and socio-economic instabilities. Environmental conditions in the region are exceptional, with the two largest deserts worldwide (Sahara,
Arabian) being among the most water scarce ecosystems on the planet (Terink et al., 2013). Climate change in this region is extraordinarily rapid; summer temperatures, in particular, are increasing by more than twice the global mean rate (Lelieveld et al., 2014), with significant impacts especially in urban areas (Mouzourides et al., 2015). While aerosol mass loadings over the EMME are dominated by desert dust, concentrations of fine particles due to anthropogenic emissions are also high (Basart et al., 2009) and will likely increase with continued population growth (Pozzer et al., 2012), making anthropogenic pollution in the area a leading health risk and an important climate forcing factor (Osipov et al., 2022).

Based on modelling studies, it has been also concluded that the EMME is characterized by highly favourable conditions for photochemical smog and ozone (O$_3$) formation leading to air quality standards being drastically exceeded (Lelieveld et al., 2014; Zanis et al., 2014). These enhanced concentrations of fine particulates and ozone have major human health implications, contributing to premature mortality (Giannadaki, et al., 2014; Lelieveld et al. 2015), which may be further exacerbated by the effects of heatwaves occurring during summer within the EMME region (Zittis et al., 2022).

Although data derived from satellite observations of NO$_2$ and SO$_2$ has revealed strong air pollution trends in the Middle East since 2010 (Lelieveld et al., 2015a), many pollution sources are still missing in emission inventories (Mclinden et al. 2016).

Thus, there is a current lack of a regional approach to characterize air pollution, with in-situ observation being insufficient, unavailable, or of low quality (Kadygrov et al., 2015; Ricaud et al., 2018; Paris et al., 2021), limiting the possibility to reduce uncertainties in regional emission inventories and implement efficient abatement strategies.

Significant efforts have been put forward in recent years to characterize the atmospheric composition in-situ over Cyprus, a central location of the EMME region (e.g. Kleanthous et al., 2014; Debevec et al., 2017 and 2018; Pikridas et al., 2018; Dada et al., 2020; Baalbaki et al., 2021; Vrekoussis et al., 2022). In-situ ground-based PM observations have clearly shown that contributions of dust to PM$_{10}$ over Cyprus are among the highest for the entire Mediterranean basin (Querol et al. 2009; Pey et al. 2013; Kleanthous et al., 2014; Pikridas et al., 2018; Achilleos et al., 2020), during dust storm events, leading to increased hospitalization, particularly attributed to cardiovascular-related diseases (Middleton et al., 2008; Tsangari et al., 2016) and short-term effects associated with daily mortality (Neophytou et al., 2013). These high levels of regional particulate matter are responsible for exceedances in PM$_{10}$ EU limits in major Cypriot cities (Querol et al., 2009). Past studies on PM trends and sources highlighted the important contribution of local (urban) emissions to PM$_{10}$ (Achilleos et al. 2014; Pikridas et al., 2018) but also showed a predominant regional pattern for PM$_{2.5}$ with major contribution of sulphur-rich sources (Achilleos et al., 2016). Based on 17 years of continuous observations of reactive gases in Cyprus, Vrekoussis et al., (2022) further confirmed the major contribution of long-range transport (incl. Middle East) in the observed concentration levels of carbon monoxide (CO) and sulphur dioxide (SO$_2$), two tracers of combustion sources.

Those studies have highlighted the unique location of Cyprus as a receptor site of major regional pollution hotspots, making the island one of the most polluted EU member states in terms PM and O$_3$ concentrations; the only one impacted by long-range transport of poorly-regulated air pollutants originating from Middle East countries. However, still few studies are currently available to assess the contribution of regional anthropogenic emissions to PM levels in Cyprus. The filter-based chemical speciation study reported by Achilleos et al., (2016) is currently the most exhaustive one and was based on 24-h integrated (PM$_{2.5}$ and PM$_{10}$) filter samples collected every 3 days for a period of one year (2012) in four cities in Cyprus. This study concluded that Cypriot cities, like many others in Europe, are characterized by a major contribution of regional sulphate and local (urban) emissions from traffic and domestic heating biomass burning.

Herewith, a detailed description of submicron (<1μm, PM$_1$) chemical composition and the further source apportionment of BC and OA is presented for the first time in Cyprus. State-of-the-art on-line instrumentation (e.g. Q-ACSM, Aethalometer) were deployed for the investigation of the temporal variability of aerosol composition at a location representative of the urban background pollution in the capital city of Nicosia. Source apportionment of submicron organic aerosols was performed using the organic fragments of the ACSM and Positive Matrix Factorization (PMF). The consistency of these results was assessed against the chemical analysis of parallel filter samples and on-line measurements of external tracers. This study was extended to
a 6-month duration in order to cover the two main seasons of the semi-arid Eastern Mediterranean climate (short mild and wet winter vs. long hot and dry summer), offering a comprehensive understanding of the daily and monthly variability of local and regional sources of carbonaceous aerosols. Cold and warm periods were compared to highlight the complexity of local (combustion) sources and the importance of regional ones. These results were further processed to apportion Black Carbon sources in Nicosia with emphasis on local versus regional contribution.

2. Material and Methods

2.1 Sampling site

Cyprus: Cyprus is the third largest island in the Mediterranean Sea, extending approximately 240km long from east-to-west and 100 km wide. The closest countries and their distance from the capital city of Nicosia are respectively Turkey (110km), Syria (250km), Lebanon (250km), Israel (300km), Egypt (400km), Jordan (430 km), and Greece (900 km from the Greek mainland), (Fig. 1a).

The population of Cyprus (ca. 1 million inhabitants) is rather small compared to its neighbouring countries and the rapidly growing (overall 400 million) population of the region (Lelieveld et al., 2013). The main urban areas of the island shown in Fig. 1b, are those of Nicosia (ca. 245,000 inhabitants), Limassol (ca. 150,000 inhabitants), Larnaca (ca. 50,000 inhabitants) and Paphos (ca. 35,000 inhabitants). Cyprus has a Mediterranean and semi-arid climate with two main seasons: a mild cold season (from December to March) and a hot warm season lasting about eight months (from April till November). Rain occurs mainly in the cold season, with the warm one being extremely dry (i.e., almost no rain between May and September)(Michaelides et al., 2018).

Nicosia: Nicosia is the largest city of the island and the southeasternmost of European Union Member States’ capitals. Nicosia is currently partitioned in two, with a buffer zone in-between under the control of the United Nations; the southern part being the capital of the Republic of Cyprus. The northern part of Nicosia (and northern part of the island) is not controlled by the government of the republic of Cyprus (Resolution 550, UN security council, 1984) (Fig. 1c). Geographically, Nicosia is located in the centre of the island, within the Mesaoria plain, 150 m above sea level (asl), which is delimited on its northern and southern edges by two mountain ranges; the Kyrenia Range culminating at 1,024 m asl, and the Troodos Mountains culminating at 1,952 m asl, respectively. This topography channels winds within a more or less west-east corridor (Fig. S5) feeding the city of Nicosia with long-range transported air masses from Europe, Africa, or the Middle East.

Measurements were performed at the Cyprus Atmospheric Observatory’s Nicosia station (CAO-NIC) located at the Cyprus Institute premises (Athalassa Campus; 174 m asl; 35.14N, 33.38E; Fig. 1c). The measurement site is considered as an urban background site, located within a low population density residential area with no significant local pollution hotspots in its vicinity (i.e., no dense road traffic, industry, commercial centers, restaurants, etc) and next to the Athalassa Forestry Park.

The period and duration of measurements presented here (07 December 2018 - 31 May 2019) were chosen to i) capture weather conditions, atmospheric dynamic, and long-range pattern of the two main seasons, ii) investigate the contribution of domestic heating emissions in winter, and iii) assess the potential increasing contribution of photochemical produced secondary aerosols during the start of the dry and warm season. Local time (LT) in Cyprus is given as Eastern European Standard Time (EET) (UTC+02:00 in winter and UTC+03:00 during the summer).
2.2 On-line Aerosol Instrumentation

On-line aerosol instrumentation has been operated following the Standard Operating Procedures defined by ACTRIS (https://www.actris.eu), the European Research Infrastructure on Aerosols, Clouds, and Trace Gases. Non-refractory submicron (NR-PM$_1$) aerosol chemical composition, i.e. organics, sulphate, nitrate, ammonium and chloride, was continuously monitored using a Quadrupole ACSM (Aerosol Chemical Speciation Monitor; Aerodyne Research Inc.) at a 30-min time resolution (Ng et al. 2011a). The instrument, along with a scanning mobility particle sizer (SMPS, described below), sampled through a sharp cut cyclone operated at 4 L min$^{-1}$ (SCC 1.197, BGI Inc., USA), yielding an aerosol cut-off diameter of approximately 1.3μm. A nafion dryer was installed upstream, keeping sample RH below 30%. Data were retrieved using ACSM local v.1.6.0.3, implemented within Igor Pro (v. 6.37, Wavemetrics Inc., USA). The ACSM is designed and built around similar technology as the aerosol mass spectrometer (Jayne et al., 2000), where an aerodynamic particle focusing lens is combined with particle flash vaporization in high vacuum at 600 °C, electron impact ionization, separation and final detection of the resulting ions using a quadrupole mass spectrometer. Mass concentrations are calculated using a chemical composition-dependent collection efficiency (Middlebrook et al., 2012).

Black carbon (BC) measurements were conducted using an AE-33 7-wavelength aethalometer (Magee Scientific, US) at a 1-min time resolution. The aethalometer sampled ambient aerosol through a PM$_{2.5}$ aerosol inlet (SCC 1.829, BGI Inc., USA) at a flow rate of 5 L min$^{-1}$, after passing through a nafion dryer. The instrument internally corrected the filter loading effect in real-time, while a fixed value ($C_0$=1.39) was applied to compensate for the multi-scattering effect (Drinovec et al., 2015). BC was apportioned to source specific components, namely BC$_{FF}$ related to fossil fuel combustion and BC$_{WB}$ related to wood burning, by applying the "aethalometer model" (Sandradewi et al., 2008) on the 470 – 950 nm wavelength pair, using the instrument’s default
values for fossil fuel combustion and wood-burning aerosol Absorption Ångström Exponent, $\text{AAE}_{\text{ff}}=1$ and $\text{AAE}_{\text{wb}}=2$, respectively.

2.3 Ancillary measurements

SMPS: Particle number size distributions were monitored using a scanning mobility particle sizer (SMPS) consisting of an electrostatic classifier (model 3080, TSI Inc., USA) coupled with a condensation particle counter (CPC; model 3070, TSI Inc. USA) operating at a 5-min time resolution and at a 1 L min$^{-1}$ sample flow rate, measuring particles with diameter ranging from 9 to 700 nm. Ambient aerosols were drawn through a nafion dryer, placed upstream, keeping sample RH below 30 %. Volume concentrations of assumed spherical particles derived by the SMPS were converted into mass concentrations using a variable density calculated by the methodology described in Bougiatioti et al. (2014). The respective mass fractions time series of chemical species were calculated based on the ACSM measurements. A density value of 1.77 g cm$^{-3}$ was used for ammonium sulphate, and 1.35 g cm$^{-3}$ for organics (Florou et al., 2017; Lee et al., 2010), the two dominant compounds of PM$_{10}$ in Nicosia as detailed further below.

Filter sampling: Co-located 24h PM$_{2.5}$ samples were collected on quartz fiber filters (Tissuquartz, 47mm diameter, Pall) using a low volume sampler (Leckel SEQ47/50) operating at a flowrate of 2.3 m$^3$ h$^{-1}$. The filter samples were analysed for i) organic and elemental carbon using an OC/EC Lab Instrument (Sunset Laboratory Inc., OR, USA) implementing the EUSAAR II protocol (Cavalli and Putaud, 2008), ii) carbohydrates, including levoglucosan, mannosan, galactosan, using an Ion Chromatography Pulsed Amperometric Detection method (Thermo - Model ICS-3000) and iii) anions (Cl$^-$, NO$_3^-$, SO$_4^{2-}$, MSA, Oxalate) and cations (K$^+$, Na$^+$, NH$_4^+$, Mg$^{2+}$, Ca$^{2+}$) using ion chromatography (Thermo - Model ICS-5000).

Proton Transfer Reaction - Mass Spectrometry (PTR-MS): Air was sampled through a 20m long, 3/8” o.d. (1/4” i.d.) sheathed Teflon line that ran from the roof of the building to the instrument. A Teflon filter (0.2µm diameter porosity) was installed at the inlet to prevent large aerosol particles and insects from entering the sampling line. The resulting residence time of air in the line was estimated to ca. 0.5 min. Temporal resolution of Volatile Organic Compounds (VOCs) measured by the PTR-MS (Ionicon Analytik, Austria) was approximately two minutes (the time required to measure 55 different ions at 2 seconds per ion). The basic operation principles of the PTR-MS instrument have been described in detail by Lindinger et al. (2011). Briefly, a stable flow of air and high concentrations of H$_2$O$^+$ ions are continuously sampled into a drift tube held at 2.2 mbar pressure. There, compounds with a proton affinity greater than water, including a large selection of Oxygenated Volatile Compounds (OVOCs), undergo efficient proton-transfer reactions with the H$_2$O$^+$ ions to produce protonated organic product ions, which can be detected by a mass spectrometer.

Meteorological Parameters: Standard meteorological parameters (temperature, relative humidity, wind speed and direction) were obtained 10 m above ground at the Athalassa Forestry Park Meteorological station of the Cyprus Department of Meteorology, located at c.a. 1.3 km east from CAO-NIC station. Wind speed and direction data were further used in this study for component-specific non-parametric wind regression analysis (NWR) performed using the ZeFir toolbox (Petit et al., 2017) developed within the Igor Pro software (Wavemetrics Inc.). A co-located automatic CIMEL CE370 micro-LIDAR was operated continuously to retrieve the Planetary Boundary Layer Height (PBLH) and better assess the influence of atmospheric dynamic on in-situ ground-based observations.

Air masses backtrajectory analysis: Five-day air mass backtrajectories arriving at 1000m altitude above the sampling site were computed every 6 hours, using the Hybrid Single-Particle Lagrangian Integrated Trajectory model (HYPLIT4; Stein et al., 2015) using the Global Data Assimilation System (GDAS 1) meteorological data fields (with 1$^\circ$ spatial resolution). Backtrajectories were coupled to measured concentrations, assessing origins and source contributions to specific chemical components, by applying the Potential Source Contribution Function (PSCF) technique as implemented in the ZeFir toolbox described above.
2.4. Source Apportionment analysis

Positive Matrix Factorization (PMF) is an advanced multivariate factor analysis tool that attempts to identify the contributing factors, or sources, of atmospheric pollutants at a sampling site. For this study, source apportionment was performed on the organic mass spectra dataset collected by the ACSM. The (PMF) method (Paatero and Tapper, 1994) using the multilinear engine (ME-2) model developed by Paatero (1999) was implemented using the SoFi (Source Finder) toolkit (SoFi 6D; Canonaco et al., 2013). PMF allows the decomposition of the OA mass spectra matrix X into two matrices, G and F and a remaining residual matrix E:

$$X = G * F + E \quad (1)$$

Where X is the input dataset matrix (measured quantity), F is the resulting source profile matrix, G is the source contribution matrix (temporal variability of each source), and E represents the model residual matrix. Based on a number of criteria the optimal solution is selected, aiming at being physically meaningful that can be supported by external indicators (ancillary measurements), and trying to minimize values in the residual matrix E. Model input data and error matrices (in μg m⁻³), were exported using the ACSM software. Data points with a signal-to-noise (S/N) ratio smaller than 0.2 were removed and points with S/N between 0.2 and 2 were downweighted by increasing their estimated error values (Ulbrich et al., 2009; Paatero and Hopke, 2003). m/z values ranging from 10 to 120 were used in the analysis. CO₂ related variables were excluded from the PMF and finally reinserted in the solution.

Source apportionment of OA was performed following the general steps described by Crippa et al. (2014) and the recently updated harmonised standard operating procedures for seasonal OA PMF (Chen et al., 2022). As a first step, unconstrained PMF analyses were performed with a number of factors ranging from 2 to 8 in order to identify the most relevant number of factors and potential sources. If primary organic aerosol factor profiles such as Hydrocarbon-like OA (HOA), or biomass burning-like OA (BBOA) were found, then the corresponding site-specific primary OA (POA) mass spectra (see discussion below) or spectra found in the literature (e.g. Ng et al., 2011) and Crippa et al., (2014)) were set as constrains in the PMF, using the “a-value” approach (Paatero and Hopke, 2009; Canonaco et al., 2013). A sensitivity analysis was then performed with different a-values to assess the level of constrain introduced in each factor with i) a constrained HOA using, as an anchor the HOA spectrum found in Ng et al. (2011) with the a-values ranging between 0.05 and 2.0, ii) a constrained BBOA factor with the a-values from 0 to 0.5, and iii) a constrained cooking OA (COA) factor from Mohr et al. (2012) with a-values from 0.2 to 0.5. Once this sensitivity analysis was completed, the evaluation of the PMF results showed that the BBOA factor could not account for the entire m/z 60 mass fragment, which was then distributed within 2 factors. Additionally, the correlation of BBOA with BCₘₖ showed to be unsatisfactory (section S1). On the other hand, given the BBOA factor’s sensitivity to the type of solid fuel used, different biomass burning factor profiles have been reported in various regions around the world (Mohr et al., 2012). Consequently, a site-specific BBOA factor profile (BBOAₛ) was selected. The BBOAₛ spectrum was calculated as an average of 20 PMF runs from the initial unconstrained PMF for the cold period, validated by it’s timeseries correlation to BCₑ. Since aged OA (i.e. Oxygen-like OA, OOA) factors show more variability between measurement sites in terms of their mass spectra, no constrain was introduced for these factors (Canonaco et al., 2015).

In this study, the BBOA factor - a major contributor of OA during winter - could not be properly resolved when performing the PMF analysis on the entire period dataset. A seasonal approach was followed instead, separating the OA dataset into two periods that were then used to describe both the two periods (cold and warm respectively). The criteria used to delineate those two periods are presented and discussed in the below section 3.2.

One factor was consequently constrained with the resulting BBOAₛ spectrum (with an a-value in the 0-0.5 range, using steps of 0.02), obtaining the optimal solution using an a-value equal to 0.46. A widely referred-to standard mass spectrum (Sun et al., 2016; Duan et al., 2020) derived from Ng et al. (2011) was used to constrain the HOA factor, with an a-value of 0.2, thus obtaining the best correlation with BCₑ, a tracer for traffic related emissions. A detailed description of the OA source apportionment analysis can be found in section S1 in the supplementary material.
3. Results and Discussion

3.1. On-line aerosol data quality check

A chemical mass closure exercise for PM\textsubscript{10} was performed at a temporal resolution of 1 h to check the quality of the on-line aerosol measurements. Chemically reconstructed PM\textsubscript{10} was calculated as the sum of the mass concentration of all non-refractory species measured by the ACSM (OA, NO\textsubscript{3}\textsuperscript{-}, SO\textsubscript{4}\textsuperscript{2-}, NH\textsubscript{4}\textsuperscript{+}, Cl\textsuperscript{-}) plus the BC concentrations measured by the Aethalometer AE-33 (Pataud et al., 2004). The contribution of other chemical constituents to submicron aerosol, such as sea salt and dust (measured by co-located filter sampling), was found to be low and therefore neglected here. A scatter plot of the ACSM + AE-33 measurements vs the SMPS-derived PM\textsubscript{10} concentrations is shown in Figure S4b. The results obtained indicate a very good correlation ($r^2 = 0.88$; N=1823) and a slope of 1.2 (Fig. S4b). This 20% discrepancy lies within the uncertainty of the on-line instruments and could be attributed to the cut-off size of the SMPS at 700 nm that is slightly lower compared to the ACSM. In addition, ACSM individual chemical species were compared with co-located off-line analyses performed on daily PM\textsubscript{10} filters. As shown in Fig. S4c-f, very good agreement was obtained between on-line and off-line measurements with $r^2 \geq 0.80$ (N=165-175) for all species. The discrepancy between ACSM and filter measurements for nitrate (slope of 1.3) could potentially be attributed to the volatilization of HNO\textsubscript{3} from the filter surface due to the presence of semi-volatile ammonium nitrate. The obtained slopes for ammonium and sulfate below 1:1 (0.81 and 0.85, respectively) is consistent with the fact that fine (NH\textsubscript{4})\textsubscript{2}SO\textsubscript{4} aerosols, mainly originating from secondary processes and long-range transport (Sciare et al., 2010; Freutel et al., 2013), can be found at a large size mode possibly exceeding 1 \(\mu\)m, consequently not being sampled by the ACSM. A very striking result obtained from the comparison of OA (ACSM) with OC (from filters) is an OM-to-OC ratio of 1.42 which is at the lower end of ratios reported for urban environments, which usually exhibit typical values of 1.6 $\pm$ 0.2 (Petit et al., 2015; Theodosi et al., 2011; Brown et al., 2013). This low ratio clearly denotes a major contribution of long-chain hydrocarbon OA that often refer to primary combustion (poorly oxidized) OA. As such, this ratio will represent a valuable and independent means of verification of the consistency of our source apportionment between primary and secondary OA.

Finally, black carbon concentrations derived from light absorption measurements (Aethalometer AE-33) were compared against filter-based EC measurements (see Fig. S4a). Data from the two techniques correlate very well ($r^2=0.83$) with a BC/EC ratio of 1.67 being similar to studies in other urban areas (Rigler et al., 2020; Liu et al., 2022), highlighting the existence of a BC absorption enhancement ($E_{abs}$) attributable to a lensing effect induced by other chemical species, among which secondary OA may play an important role (Zhang et al., 2018).

3.2 Meteorological conditions

Delineation of cold vs warm seasons: The ACSM organic mass at m/z (mass-to-charge ratio) 60 is characteristic of the fragmentation of levoglucosan, a product of cellulose pyrolysis and well-established biomass burning marker (Alfarra et al., 2007). Its respective contribution to total OA ($f_{60}$) was used in this study as an indicator of biomass burning for domestic heating to delineate cold vs warm seasons, comparing with the 0.3% threshold proposed by Cubison et al., (2011) for air masses influenced by biomass burning. Except for a single small peak in early May, corresponding to open fires for the celebration of the Greek Orthodox Easter, the last instance when $f_{60}$ was above the threshold was recorded during the first week of April (Fig. 2). From then onwards, daily air temperature started rising constantly, from ca. 15°C in the beginning of April up to 30°C at the end of May. These two features dictated the division of the dataset into two periods: a cold period of four months (07/12/2018-08/04/2019), with an average temperature of 12 $\pm$ 4°C, and a warm period of two months (09/04/2019 – 31/05/2019), with an average temperature of 20 $\pm$ 7°C.
Figure 2: Time series of air temperature (blue), m/z 60 organic concentration (org60, brown) and f60 fragment (green) for the cold and warm periods. The vertical line is used to delineate the measurements within the two seasons.

**Wind sectors:** During these two periods a distinct pattern in the wind sectors and the air masses arriving at the sampling site was observed. As seen in Fig. S5, the dominant wind direction for the cold period was the NW-SW [225° - 315°] sector encompassing 48% of the total wind directions while the NE-SE [45° - 135°] sector covered 26%. During the warm period the weight of this proportion is shifting even more towards the NW-SW [225° - 315°] sector having a 62% of total air masses while only 17% are arriving from the NE-SE [45° - 135°] sector.

**Air mass origin:** A cluster analysis was performed (Fig. S6) for both periods in order to better assess the main upwind regions responsible for long-range transported air pollution over Cyprus and their change relative to the period of the year. Most of the air masses arriving in Cyprus were found to originate from Europe; many of them passing over Turkey before reaching Cyprus. Interestingly, this analysis showed one cluster arriving from Middle East for the cold period, whereas there were not enough trajectories passing over the Middle East to calculate a cluster for this area during the warm period.

### 3.3. Chemical composition of PM$_1$

**Seasonal perspective of PM$_1$:** Time series of PM$_1$ chemical composition derived from the ACSM (OA, SO$_4^{2-}$, NH$_4^+$, NO$_3^-$, Cl$^-$) and the Aethalometer ($BC_{ff}$, $BC_{wb}$) are depicted for the entire measuring period in Figure 3. Averaged data (6h averaging period) are shown here for clarity. Furthermore, the relative average contribution of each chemical constituent to total PM$_1$ concentrations, is depicted in the respective inserted pie charts for both periods.
Figure 3: Time series of the chemical composition of PM$_1$ in Nicosia derived from 6-hour averages of ACSM and AE33 measurements. The vertical dashed red line separates the cold from the warm season. The average relative contribution of each species is shown in the respective pie charts (inner panels) for each season.

Although intense and short duration peaks are observed for carbonaceous aerosols (OA, BC$_{ff}$, BC$_{wb}$), background NR-PM$_1$ concentration levels (between peak values) remain well below 10 μg m$^{-3}$ for the 6-h average in both seasons. In other words, no PM$_1$ pollution episodes (with e.g. concentrations above 10 μg m$^{-3}$) lasting for consecutive days were observed. Such lack of intense and persistent PM$_1$ pollution episodes differs from what is reported in central and northern Europe, where stagnant (anticyclonic) conditions occur together with continental (polluted) air masses, mainly in winter and springtime (e.g. Petit et al., 2015). This suggests that the relatively low emissions from Cyprus (compared to the neighboring countries) together with its remote marine location (i.e. far from densely populated areas) may prevent the build up of high PM$_1$ pollution events over Nicosia. On the other hand, clear differences can be observed between both periods, with significantly higher PM$_1$ concentrations during the cold period, associated with repeated intense peaks of OA and BC - not observed during the warm season – and suggesting local combustion emissions. The highest PM$_1$ concentrations were observed between December 28$^{th}$ 2018 and January 13$^{th}$ 2019 (Fig. 3) and were associated with low temperatures and Christmas holidays, both likely to promote the use of domestic heating. During the warm period, higher contribution of sulfate, and lower contribution of OA, are clearly noticeable. The contribution of nitrate during the warm period, most probably in the form of semi-volatile NH$_4$NO$_3$, remains marginal, possibly due to non-favourable thermodynamic conditions preventing its formation and accumulation.

PM$_1$ chemical composition: For the cold period, the average calculated mass concentration of PM$_1$ (calculated as the sum of chemical components measured by AE33 and ACSM) was 12.32 ± 9.77 μg m$^{-3}$, with 10.30 ± 7.92 μg m$^{-3}$ being the average concentration of the non-refractory species (Table 1). OA constitutes the larger fraction of PM$_1$ mass, with an average concentration of 5.03 ± 5.48 μg m$^{-3}$ (41%), followed by sulfate (23%), black carbon (16%), nitrate (10%), ammonium (9%), and chloride (3%). These concentrations and the overall distribution of chemical components in NR-PM$_1$ are quite similar to those measured by ACSM in other European cities (Bressi et al., 2021). Concentrations appear to decline during the warm period, with an average calculated PM$_1$ concentration of 8.18 ± 4.65 μg m$^{-3}$, including 7.15 ± 3.80 μg m$^{-3}$ from the non-refractory
components. The dominant species during the warm period were sulfate and OA, each representing 35% of PM$_1$, followed by black carbon (12%), ammonium (11%) and nitrate (6%). During that period, chloride concentrations were negligible, contributing less than 1% (Table 1).

Table 1: Species mean, standard deviation, median concentrations and respective contribution to PM$_1$ during cold and warm periods in Nicosia.

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<th>Cold Period</th>
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<td>Mean (μg m$^{-3}$)</td>
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<tr>
<td>OA</td>
<td>5.03</td>
<td>5.48</td>
</tr>
<tr>
<td>SO$_4^{2-}$</td>
<td>2.81</td>
<td>1.89</td>
</tr>
<tr>
<td>NO$_3^-$</td>
<td>1.22</td>
<td>1.25</td>
</tr>
<tr>
<td>NH$_4^+$</td>
<td>1.14</td>
<td>0.77</td>
</tr>
<tr>
<td>Cl$^-$</td>
<td>0.31</td>
<td>0.25</td>
</tr>
<tr>
<td>BC</td>
<td>2.01</td>
<td>2.31</td>
</tr>
<tr>
<td>PM$_1$</td>
<td>12.32</td>
<td>9.77</td>
</tr>
</tbody>
</table>

Interestingly, sulfate concentrations recorded in Nicosia are much higher compared to what is commonly observed in other European countries and Mediterranean cities (Table 2) and likely reflect a regional pattern of sulfur-rich emissions compared to Europe where SO$_2$ emissions have strongly decreased during the last decades (Smith et al., 2011; Chin et al., 2014) thanks to the implementation of specific abatement measures on reducing sulfur emissions (European NEC Directive (EU, 2016) and United Nation Gothenburg (1999) protocol). More specifically, the importance of sulfur emissions in Turkey (2 455 Gg, EEA 2021) which were 50% higher compared to the total SO$_2$ emissions of the EU 28 in 2019, together with the fact that half of air masses reaching Cyprus are passing over Turkey (see Fig. S6) are key contributors to the high concentrations of sulfate in our study.

Table 2: Comparison of concentration, and percentage contribution to PM$_1$, between the main submicron chemical species derived by ACSM.

<table>
<thead>
<tr>
<th></th>
<th>PM$_1$ (μg m$^{-3}$)</th>
<th>OA (μg m$^{-3}$)</th>
<th>SO$_4^{2-}$ (μg m$^{-3}$)</th>
<th>NH$_4^+$ (μg m$^{-3}$)</th>
<th>NO$_3^-$ (μg m$^{-3}$)</th>
<th>Cl$^-$ (μg m$^{-3}$)</th>
<th>Reference</th>
</tr>
</thead>
<tbody>
<tr>
<td>Nicosia Cold (DJFM)</td>
<td>12.32</td>
<td>5.03</td>
<td>2.81</td>
<td>1.14</td>
<td>1.22</td>
<td>0.31</td>
<td>This study</td>
</tr>
<tr>
<td>Nicosia Warm (AM)</td>
<td>8.18</td>
<td>2.83</td>
<td>2.87</td>
<td>0.92</td>
<td>0.53</td>
<td>0.05</td>
<td>This study</td>
</tr>
<tr>
<td>Cyprus RB* (Annual)</td>
<td>7.6</td>
<td>3.26</td>
<td>2.66</td>
<td>0.98</td>
<td>0.23</td>
<td>-</td>
<td>Chen et al. (2022)</td>
</tr>
<tr>
<td>European UB** (Annual)</td>
<td>10.6</td>
<td>5.3</td>
<td>2.0</td>
<td>1.9</td>
<td>-</td>
<td>Bressi et al. (2021)</td>
<td></td>
</tr>
<tr>
<td>S. Europe RB*** (Annual)</td>
<td>6.3</td>
<td>3.5</td>
<td>1.3</td>
<td>0.8</td>
<td>-</td>
<td>Bressi et al. (2021)</td>
<td></td>
</tr>
<tr>
<td>Athens Winter</td>
<td>18.7</td>
<td>13.13</td>
<td>2.4</td>
<td>1.8</td>
<td>0.14</td>
<td>Stavroulas et al. (2019)</td>
<td></td>
</tr>
<tr>
<td>Athens Spring</td>
<td>6.42</td>
<td>3.3</td>
<td>2.1</td>
<td>0.6</td>
<td>0.4</td>
<td>0.02</td>
<td>Stavroulas et al. (2019)</td>
</tr>
<tr>
<td>Marseille Winter</td>
<td>11.9</td>
<td>6.17</td>
<td>1.12</td>
<td>0.86</td>
<td>1.58</td>
<td>0.09</td>
<td>Chazeau et al. (2021)</td>
</tr>
<tr>
<td>Marseille Spring</td>
<td>8.09</td>
<td>3.86</td>
<td>1.06</td>
<td>0.70</td>
<td>1.13</td>
<td>0.04</td>
<td>Chazeau et al. (2021)</td>
</tr>
<tr>
<td>Barcelona (Annual)</td>
<td>9.85</td>
<td>4.10</td>
<td>1.70</td>
<td>1.05</td>
<td>1.35</td>
<td>0.06</td>
<td>Via et al. (2021)</td>
</tr>
</tbody>
</table>

* Cyprus Regional background
** European urban background = Barcelona (Spain) + London (UK) + Prague (Czech) + Tartu (Estonia) + Zurich (Switzerland)
*** Southern European regional background = Ersa (Corsica, France) + Finokalia (Crete, Greece)

The main difference between the cold and warm periods lies in the decrease of concentration of carbonaceous aerosols (OA, BC) and NO$_3^-$ by almost a factor of two. Several phenomena can explain this significant seasonal variation: the absence of a domestic heating source (mainly biomass burning as explained in Fig. 2); the absence of Middle East air masses during the warm period (see discussion later on); the increase in the Planetary Boundary Layer Height (PBLH) above Nicosia (Fig. S7) enhancing vertical dilution of local emissions during the warm period and therefore lowering ground-based concentrations; less favourable
thermodynamic conditions, with warmer and dryer air, also preventing the condensation of semi-volatile species (e.g. ammonium nitrate). Sulfate concentrations do not exhibit a similar seasonal pattern and therefore seem to be less affected by the above factors. On the contrary, the increase in photochemistry enhances the formation of sulfate aerosols, and the decrease in precipitation enhances aerosol lifetime, strengthening the impact of long-range transport.

### 3.4. Diurnal variability of PM$_1$ chemical constituents

Figure 4 shows the diurnal variability of the PM$_1$ species derived from the ACSM and AE-33 for both the cold (Fig. 4a) and warm (Fig. 4b) periods. The diurnal variability of the apportioned BC related to fossil fuel combustion (BC$_{ff}$) and wood-burning (BC$_{wb}$) are also depicted here.

**Figure 4**: Diurnal variability of the main submicron chemical constituents (OA, SO$_4^{2-}$, NO$_3^-$, NH$_4^+$, Cl$^-$ and BC) during the a) cold and b) warm periods. Diurnal profiles of BC$_{ff}$ and BC$_{wb}$ are embedded.

**Organic aerosols**: Organic aerosols clearly dominate the cold period PM$_1$ concentration levels, exhibiting a night-time maximum above 12 µg m$^{-3}$, and a second smaller maximum at 4 µg m$^{-3}$ coinciding with local traffic rush hour (06:00-09:00 LT). Elevated OA concentrations in the cold period during the night (max at 22:00 LT) are a common well-documented feature in many urban environments across Europe and the Mediterranean (e.g. Florou et al., 2017; Stavroulas et al., 2019; Chazeau et al., 2021). They can be attributed to higher emissions from domestic heating, evening traffic peak and cooking activities. The strong correlation between OA and BC$_{wb}$ ($R^2=0.81; N=2934$; Fig. S8) suggests that residential wood burning is an important contributor to this nighttime peak. Interestingly, this peak is not significantly amplified by a lower PBLH during night-time, which seems to remain relatively stable with no significant diurnal variability during the cold period (Fig. S7). It is also worth noting that background OA concentrations observed both at the end of the night and middle of the day, when local emissions are minimal, remain relatively high at ca. 3 µg m$^{-3}$. The diurnal variability of OA is much less pronounced during the warm period, suggesting a more important contribution of regional sources to OA compared to the strong dynamic of local emissions. The assumption of a more important contribution from regional OA during the warm period is further supported by a mean OA concentration of 2.83 µg m$^{-3}$ (Table 2) that is close to the averaged OA concentrations of 3.2 µg m$^{-3}$ reported for a 2-year period continuous observations with Q-ACSM (2015-2016) at the rural background site of the Cyprus Atmospheric Observatory at Agia Marian Xyliatou (CAO-AMX), ca. 40km distance from Nicosia (Chen et al., 2022). During the warm period, a small OA peak remains visible in the morning, with similar amplitude to the cold season, likely to be related to traffic emissions. A second peak can be observed at 21:00 LT (not observed in BC) which may possibly originate from cooking activities.

**Black carbon**: During the cold season, BC follows a bimodal diurnal pattern, which can be further apportioned by focusing on its source-specific components BC$_{ff}$ and BC$_{wb}$. The fossil fuel component exhibits two maxima, one in the early morning,
coinciding with traffic rush hour, and one in the late afternoon, most probably related to both traffic and increase in energy demand due to domestic heating (see discussion later on). On the other hand, BC_{wb} diurnal variability is dominated by a night-time maximum (20:00 - 01:00 LT), peaking one hour after BC_{f} and linked to wintertime residential wood-burning emissions, contributing up to 33% of total BC. During the warm season, the BC diurnal pattern is characterised by the absence of a night-time maximum, while still exhibiting a significant peak in the morning, dominated by BC_{f}. The very low contribution of biomass-related combustion particles during the warm period, as previously noted from m/z 60 in Fig. 2, is further supported here with BC_{wb} exhibiting a nearly flat diurnal variability with close-to-zero mass concentrations.

Secondary inorganic aerosols: During the cold season, non-refractory nitrate and chloride detected by the Q-ACSM are mostly present in the form of semi-volatile NH\textsubscript{4}NO\textsubscript{3} and NH\textsubscript{4}Cl (Guo et al., 2017; Theodosi et al., 2018). They show a night-time maximum (Fig. 4-a), reflecting the presence of gas precursors (NH\textsubscript{3}, HNO\textsubscript{3}, HCl) and the more favourable thermodynamic conditions with lower temperatures, higher relative humidity, and condensation sink due to high PM concentrations of combustion aerosols (traffic, domestic heating). Additionally, there is a smaller morning NO\textsubscript{3} peak most probably linked to traffic (Foret et al., 2022). This is not observed for chloride, suggesting that HCl may not be as abundant in the morning compared to the evening. The less favourable thermodynamic conditions during the warm period leads to very small concentrations of semi-volatile NO\textsubscript{3} and Cl\textsuperscript{-} (Fig. 4b). As expected, sulfate does not show a pronounced diurnal pattern, irrespective of the period, and pointing to regionally-processed aerosols (Fig. 4a,b).

3.5. OA Source Apportionment

3.5.1 OA source apportionment during the cold period

For the cold period, the optimal PMF result has been found using a 5-factor solution following the approach detailed in section 2.4. The identification of OA sources related to these 5 factors was then performed following the typical combination of information from i) OA mass spectra (Fig. S9a), ii) the correlation of each factor with source-specific tracers (see Fig. 5), iii) their diurnal variability (Fig. 6a), and iv) their daily (week days vs week-end) pattern (also Fig. 6b). The five factors were then assigned to the following sources: A primary BBOA (Biomass Burning Organic Aerosol), two primary HOA (Hydrocarbon-like Organic Aerosol; HOA-1 and HOA-2) and two secondary OA sources, namely MO-OOA (Low-volatile Oxygenated Organic Aerosol) and LO-OOA (Semi-volatile Oxygenated Organic Aerosol). This source apportionment is presented and justified below for each factor:

HOA-1 (Hydrocarbon-Like OA Type 1): The mass spectrum of HOA-1 (Fig. S9a) is consistent with a fossil fuel (traffic) combustion source that can be identified by the prevailing contributions of the ion series representing C\textsubscript{n}H\textsubscript{2n+1} (m/z = 27, 41, 55, 69, 83, 97, typical fragments of cycloalkanes or unsaturated hydrocarbon chains) and C\textsubscript{n}H\textsubscript{2n+1} (m/z = 29, 43, 57, 71, 85, 99, typical fragments of alkane chains). Hence, this factor mass spectrum is well correlated to eight selected HOA factors related to vehicular traffic found in the literature (Fig. S10a) and relevant to European and Mediterranean environments. The traffic-related origin of the HOA-1 factor can be further confirmed by the good correlation with BC_{f} (R\textsuperscript{2}=0.65; N=2934; Fig. S11a), benzene (R\textsuperscript{2}=0.72; N=1165; Fig. S11b). The diurnal variability of HOA-1 shows a bimodal cycle with a sharp maximum during the morning rush hour with an amplitude similar to BC_{f} (Fig. 6a), and a broader maximum in the evening possibly encompassing emissions from traffic and diesel fired residential heating systems. In the weekly cycle as depicted in Fig. 6b, the morning peak decreases in Saturday and is nearly absent on Sunday mornings, aligned with the de-escalation of traffic emissions usually observed during weekend mornings.

BBOA (Biomass Burning OA): The mass spectrum of the site-specific BBOA factor (reported as BBOA\textsubscript{aer}, in section 2.4) exhibits characteristic peaks at m/z 29, 60, and 73 (Fig. S9a), which are indicative of biomass burning (Crippa et al., 2014). The
mass spectrum is quite similar to other BBOA spectra found in the Mediterranean and Europe (Fig. S10c) with a key difference here being the rather low contribution of signal at m/z=43. The biomass burning related origin of the factor is further confirmed by the strong correlation with BC_{wb} \left(R^2=0.81; N=2934; \text{Fig. S11c}\right), benzene \left(R^2=0.61; N=1162; \text{Fig. S11d}\right) and levoglucosan \left(R^2=0.94; N=125; \text{Fig. S11e}\right), a typical tracer of biomass burning (Fourtziou et al., 2017). The BBOA diurnal pattern exhibits an expected well-marked night-time maximum around 22:00 LT, consistent with residential wood burning activities. This night-time maximum is observed throughout the week (Fig. 6a) confirming the important role of wood burning for heating in the city. Interestingly, the higher concentrations of BBOA as well as BC_{wb} (Fig. 6b) were observed on Sunday evenings, pointing to a recreational use of fireplaces, leading to enhanced residential wood-burning emissions during the weekend, a feature also reported in other sites in Europe and the US (Bressi et al., 2016; Rattanavaraha et al., 2017; Zhang et al., 2019).

**HOA-2 (Hydrocarbon-Like OA Type 2):** The mass spectrum obtained for this factor (Fig. S9a) is similar to the HOA-1 factor, with high signals for the ion series C_{n}H_{2n+1}^{+} and C_{n}H_{2n+2}^{+}. The main differences between these two factors occur in the relative contribution of m/z 41 compared to m/z 43 and the relative contribution of m/z 55 compared to m/z 57 which are both much higher for HOA-2, than for HOA-1. Furthermore, the contribution of signal to m/z 44 is more significant in HOA-2, which can imply a mix of various sources and/or a possible higher degree of atmospheric processing. Other discrepancies with HOA-1 concern its diurnal variability with an intense maximum at night (Fig. 6a) and even more its average concentration levels, which are almost three times higher than HOA-1.

**Influence of cooking activities:** The HOA-2 diurnal profile has a small peak at 13:00 LT and a significantly higher one at 21:00 LT effectively coinciding with typical meal times in Cyprus as well as those reported in literature for Greece (Siouti et al., 2021), therefore indicating the influence of cooking activities to this factor. When plotting f_{55} vs f_{57} (Mohr et al., 2012) and colouring the data points by the corresponding time of day, a distinct pattern appears with data of higher f_{55} over f_{57} being clustered to the top left of the triangle, close to the fitted lines representing cooking (Fig. S12) and coinciding with midday and evening hours. The night-time maxima pattern is consistent throughout the week, with the higher concentrations being recorded on Friday and Saturday evenings (Fig. 6b), in line with an expected food service sector activity increase as part of Nicosia inhabitants’ leisure in the weekend. The mass spectrum of HOA-2, even though left unconstrained, is highly correlated to COA found in other studies (Fig. S10b) in both Mediterranean and continental European urban environments. Additionally, the non-negligible signal at m/z=60 points to the widely spread habit of meat charbroiling (Kaltsounidis et al., 2017).

**Influence of power plant emissions:** A closer look at the diurnal variability of the HOA-2 factor shows a certain persistence of this factor throughout the day, even when cooking activities are more or less absent (Fig. 6a). Such pattern could imply the influence of other combustion sources, not necessarily of local origin. The influence of other combustion sources would also help to explain why HOA-2 average concentrations are roughly 3 times higher than OA related to traffic (HOA-1), as it is very unlikely that cooking activities can contribute solely to the observed HOA-2 concentrations. A possible contributing source could be related to the energy production sector in the island which relies exclusively on heavy fuel oil. In a recent study, Vrekoussis et al. (2022), utilizing satellite observations, have identified that power plants located to the North (Tekneckoik powerplant, PP4, 362MW), North-East (Kalecik powerplant, PP5, 153MW) and South-East (Dhekelia power station, PP3, 460MW) of Nicosia at 22 km, 60 km and 38 km, respectively, are significantly contributing to columnar NO2 concentrations over the island. The importance of these emission hotspots, along with their location on the island, during both the cold and warm periods is illustrated in Fig. S13 and show, in particular for the Northern power plants (PP4, PP5), emissions as high as the traffic-related NO2 over Nicosia. Interestingly, in a source apportioning study on VOCs performed at the Cyprus Atmospheric Observatory – Agia Marina Xyliatou (CAO-AMX), a rural remote site 32 km southwest of Nicosia, Debevec et al. (2017) have resolved a factor related to industrial activity/powered generation, exhibiting a connection with winds arriving from the wider eastern sector.

In order to assess the possible influence of Cypriot power plant emissions, the coupling of wind velocity, wind direction with the HOA-2 timeseries was performed through NWR analysis (Fig S14b). This analysis highlights the association of stagnant conditions (low wind speed / low dispersion) with high HOA-2 concentrations (i.e. night-time peaks) pointing to a more local
origin for this OA source. On the other hand, different features appear when wind velocities are higher, showing emissions originating from the NW and the E-NE sectors; i.e. downwind of power plants PP4, and PP5, although long-range transport influence cannot be ruled out. This is illustrated by the NWR of sulfate (Fig. S14f) which shows a dominant E sector likely to originate from regional emissions. Given the positioning of the sampling site, close to the edge of Nicosia urban fabric, with the Athalassa park lying to the east, such an observation can suggest transport of plumes from the operating powerplants, namely PP5 and PP3 to the city. Interestingly, a similar, yet even clearer image stands for SO2 concentrations – only half of which are considered to be of urban origin (Vrekoussis et al., 2022) – measured at a suburban background site (NicRes) and a traffic site (NicTra) in the city (Fig. S14g-h), with elevated SO2 concentrations being related to eastern winds of higher velocity, further corroborating that power generation related polluted plumes, traveling through the Mesaoria plain arriving to Nicosia can contribute to the HOA-2 factor.

Other combustion sources: Interestingly, chloride shows a good correlation with HOA-2 ($r^2=0.61$; N=2945; see Fig. S11f, Fig. 5). Chloride detected by the ACSM is in the form of NH4Cl (a secondary highly-volatile species). The source of this chloride is still widely debated and may originate from industrial activity or municipal (plastic-containing) waste burning (Gunthe et al., 2021). Another possible explanation of the good agreement between HOA-2 and chloride would be the use of CI-rich coal as a means for outdoor cooking in Nicosia could therefore reflect the influence of cooking activities that comprises a fraction of the HOA-2 factor.

**Less-Oxidized Oxygenated OA (LO-OOA):** With elevated contribution of m/z 44, the mass spectrum of this factor is consistent with a secondary OOA source. A higher m/z 43, and a lower m/z 44 (Fig. S9a) compared to MO-OOA, implies a less oxygenated (less-processed) component (Mohr et al., 2012). Finally, the time series of this factor is quite similar to NO2, with an overall good correlation value ($R^2 = 0.67$, N=2943; Fig. S11h) highlighting its semi-volatile character. This is further corroborated by the very good correlation of LO-OOA with chloride ($R^2 = 0.73$, N=2943; Fig S11i), another semi-volatile compound measured by the Q-ACSM. The diurnal variation of LO-OOA displays 1.5 times higher concentrations during the night compared to daytime (maximum of 1.84 ± 0.31 µg m$^{-3}$ at 22:00 LT; Fig. 6a); a pattern that is much more pronounced than the variability observed for MO-OOA. This feature highlights the fact that the presence of LO-OOA, is not exclusively controlled by photochemical processes. Instead, changes in thermodynamic equilibrium (due to lower T and increased RH), favouring the condensation of gas-phase semi-volatile material on one hand, and intense night-time chemistry (gas phase or heterogeneous) on the other hand, are among the processes that may account for the rapid night-time formation of LO-OOA. Atmospheric processing of biomass burning OA during periods of low photochemical activity (such as in winter or at night), known also as “dark” aging, have been reported recently (Kodros et al., 2020; Jorga et al., 2021) and could have contributed to the observed night-time formation of LO-OOA. Notably, the weekly cycle of LO-OOA, and its night-time maxima, appears to have the same pattern and intensity as those observed for BBOA (e.g., low peaks on Tuesday/Thursday, maximum on Sunday) (Fig. 6b). On the other hand the factor is correlated with both BBOA ($R^2=0.81$; Fig. S11k) and BCmeq ($R^2=0.66$; Fig. S11j). This observation could indicate a biomass burning contribution to LO-OOA, through fast oxidation of primary emissions, supported in several studies showing biomass burning linked OOA sources at night (Stavroulas et al., 2019; Kodros et al., 2020; Chen et al., 2021).

**More-Oxidized Oxygenated OA (MO-OOA):** The MO-OOA factor typically accounts for secondary organic aerosol formed in the atmosphere from gas-to-particle conversion processes of VOCs and their products, as well as atmospheric aging of primary OA (Petit et al., 2015; Stavroulas et al., 2019). Numerous VOC sources can contribute to OOA but lose their mass spectrum fingerprint owing to extended oxidation due to photochemical aging, which leads to enhanced signal at the m/z 44 fragment (CO2$^+$), a dominant tracer for OOA (Ng et al., 2011). The predominance of m/z 44 and the near absence of m/z 43 in the mass spectrum of the resolved MO-OOA factor (Fig. S9a) points to highly oxidized/aged secondary OA (i.e., originating from long-range transport). This is further supported by the relatively good agreement ($R^2=0.55$; N=2943; Fig. S11l) between concentrations of MO-OOA and sulphate (Fig. 5), a species of regional origin (Sciare et al., 2003). Nevertheless, the diurnal variability of MO-OOA does not closely follow sulfate showing a small increase of 20-30% every evening (Fig. 6a,b) which...
furthermore cannot be explained by atmospheric dynamics (c.f. the negligible PBLH diurnal variability for the cold period shown in Fig. S7). Alternatively, this would suggest that a fraction of MO-OOA is produced locally through night-time oxidation mechanisms as previously observed for LO-OOA.

Figure 5: Time series of the five OA factors resolved along with corresponding tracer compounds for the cold period.

Figure 6: Diurnal variability (left) and weekly cycles (right) of the five OA factors averaged over the cold period.

3.5.2. OA source apportionment during the warm period

For the warm period the optimal PMF solution was obtained using a 4-factor solution (HOA-1, HOA-2, MO-OOA, LO-OOA). As expected, the BBOA factor could not be resolved as previously highlighted by the low concentrations at m/z 60 reported during this period (Fig.2). Again, the identification of OA sources related to the 4 OA factors was performed following the typical combination of information from i) OA mass spectra (Fig. S9b), ii) the correlation of each factor with external source-specific tracers (Fig. 7 and Fig. S15), iii) their diurnal variability (Fig. 8a), and iv) their daily (week days vs week-end) pattern (also Fig. 8b). The mass spectra profiles for the 4-factor PMF solution during the warm period (Fig. S9b) were quite similar to the ones from the cold period (Fig. S9a).
HOA-1: For the warm period an a-value of 0.2 was selected for constraining the HOA-1 factor, again using the Ng et al. (2011b) factor’s diurnal pattern exhibiting much less variability, highlighting its dominant regional character. The time series of MO-545-1 time series follows the same pattern as the corresponding traffic-related HOA-1 factor reported for the cold period, showing a good correlation with BC$_{10}$ (R$^2$=0.62, N=1259; Fig. S15a). Its diurnal variability exhibits a bimodal pattern, with a typical sharp maximum in the morning (07:00 LT) and a smaller peak during the evening (Fig. 8a). On a weekly basis, this diurnal variability tends to be less pronounced on Saturdays and nearly absent on Sundays (Fig. 8b), reflecting reduced commuting during the weekend.

HOA-2: The HOA-2 factor still shows elevated concentrations during the warm period, close to 3 times higher compared to HOA-1 (Table S2). The profile remains quite unchanged between the cold and warm periods (R$^2$= 0.92; Fig. S10b) pointing to similar sources. No correlation was observed with chloride which may be expected due to unfavourable thermodynamic conditions hindering NH$_4$Cl formation as well as the lack of significant chloride sources during this period. A night-time maximum of HOA-2 is still observed when investigating the factor’s diurnal variability (Fig. 8a). Furthermore, a somewhat broader, compared to the cold period, maximum in the middle of the day (Fig. 8a) can also be observed. When going through the weekly variability, this midday maximum is particularly well defined on Sundays (Fig. 8b), while the evening peaks of Sundays and Mondays are the lowest. The above observations remain consistent with the cold period assessment, that HOA-2 is on one hand linked to cooking activities. For households activities are expected at noon and evenings while for restaurants, activity peaks on Sunday noon and is lower on Sunday evening and Monday reflecting the fact that such businesses remain closed on the first day of the week (Fig. 8b). On the other hand, the overall offset of HOA-2 observed against the HOA-1 diurnal profile persists, suggesting somewhat permanent background HOA-2 concentrations that cannot be explained by cooking activities alone. A contribution to this source by continuous emissions from power plants (see space-based (SP5-TROPOMI) vertical columns of NO$_2$ during the warm period in Fig. S13d), should be sought. In addition the HOA-2 NWR plot for the warm period reveals an even more significant enhancement of concentrations when moderate winds blow from the E-SE (Fig. S16b), a trend also observed for SO$_2$ during the same period (Fig S16c,f).

The above observations remain consistent with our assessment for the cold period: the HOA-2 factor consists of a mixed OA source that contains cooking activities (inc. coal combustion) and emissions from the powerplants located on the eastern part of the island. Indeed, the HOA-2 midday maximum can be linked to an increase in electricity demand at that time of day during the warm period due to an increase in air conditioning usage (Cyprus’ NECP 2021-2030, 2019).

LO-OOA: The LO-OOA factor profile exhibits some differences with the one resolved for the cold period (R$^2$ = 0.66) as illustrated in the correlation matrix of comparison to selected factor profiles found in the literature (Fig. S10d) while being very similar to those obtained in Athens/Piraeus during summer (Bougiatioti et al., 2014; Stavroulas et al., 2021). The LO-OOA timeseries shows a low agreement with NO$_2$ (R$^2$ = 0.31; N=1259; Fig. S15c) poorer than the observed correlation during the cold period (Fig. S11h). The diurnal pattern of the factor (Fig. 8a) shows maximum concentrations persisting throughout the night and early morning while a secondary maximum during midday can be observed. But overall, the diurnal pattern of LO-OOA is rather flat compared to the cold period, suggesting that local production may not be so important at that time compared to less variable regional background. Interestingly a midday hump similar to the one observed for HOA-2 is present, suggesting a common origin.

MO-OOA: The factor profile of MO-OOA resolved during the warm period is strikingly identical to the profile found in the cold period (their R$^2$ is almost 1; Fig. S10e), while being excellently correlated to other highly oxygenated OA factors resolved in both the urban and regional background in the Eastern Mediterranean (Bougiatioti et al., 2014; Stavroulas et al., 2019, 2021) as well as in continental Europe (Crippa et al., 2014). The winter night-time peaks are not observed anymore (Fig. 8a), with the factor’s diurnal pattern exhibiting much less variability, highlighting its dominant regional character. The time series of MO-
OOA correlates good to SO$_4^{2-}$ ($R^2=0.53$; $N=1259$; Fig.S15b), confirming this regional and highly processed origin. The concentration levels of MO-OOA during the warm period are lower than in the cold (Table S2). However, its relative contribution to total OA during the warm period remains similar (45%).

Figure 7: Time series of the four OA factors resolved along with corresponding tracer compounds for the warm period.

![Figure 7](image)

Figure 8: Diurnal variability (a) and weekly cycles (b) of the four OA factors resolved during the warm period.

3.6. Spatial and seasonal variability of OA sources

3.6.1. Seasonal variability of OA sources

Primary OA: The mass concentration of the three primary OA factors (HOA-1, HOA-2, BBOA) represents as much as 40% of total organic aerosols during the cold period (Fig. 9), with POA contribution significantly decreasing in the warm period (22% to total OA) due to the absence of the significant residential wood burning source which during the cold period accounted for 12% of total OA. The important contribution of primary sources in Nicosia has been also highlighted earlier, by the rather low OA/OC ratio of 1.42 (Section 3.1). In a recent publication covering several European sites, Chen et al. (2022) reported that in
urban sites, solid fuel combustion-related OA components were 21.4% of total OA during winter months, higher than what is found for BBOA in Nicosia, owing to the rather milder winters in the city.

The traffic related primary factor in Nicosia (HOA-1) was found to be rather stable in terms of contribution to total OA across this study’s two seasons averaging 7% and 6% respectively for the cold and warm periods, being lower than the figure reported in other European Urban sites (12.7%, Chen et al., 2022). On the other hand, the HOA-2 factor represents ca 2/3 of the total HOA in Nicosia with little variation from winter (72%) to summer (66%) to total HOA (Fig. 9). Comparing it with COA in urban locations resolved by Chen et al. (2022), during both winter (14.4% compared to 21% in the cold season in Nicosia) and spring (15% versus 16% in Nicosia during the warm season), the higher values reported in Nicosia further support the assumption that the HOA-2 represents a mixed combustion source.

Secondary OA: A higher degree of oxidation is observed for the LO-OOA factor during the warm period, given the much higher contribution of signal at m/z 44 compared to the respective cold period factor. This discrepancy, reported in several studies (Huang et al., 2019; Duan et al., 2020), is explained by higher photochemistry during the warm period which promotes the oxidation of OA, resulting in an LO-OOA profile with a higher m/z 44 fraction. This result is also consistent with a less-oxidized LO-OOA formed during the cold period from night-time chemistry. The range of LO-OOA concentration levels are different between cold and warm periods (0.05-7.74 μg m⁻³ and 0.05-4.00 μg m⁻³, respectively), while the mean concentrations for both periods are similar (0.86 and 0.95 μg m⁻³ for cold and warm periods respectively). The contribution of LO-OOA relative to total OA is double during the warm period compared to the cold, reflecting both the absence of the biomass burning source as well as the prevailing conditions favoring atmospheric processing of primary OA and SOA precursors. During the cold period, LO-OOA intense peaks suggest an influence from local emissions while during the warm period, the less-variable LO-OOA diurnal variability highlights the influence of more intense photochemical processing at medium-to-large geographical scale. MO-OOA is found to be the major contributor to total OA for both the cold (44%) and warm (45%) periods, higher in both cases than the average MO-OOA contributions reported for other European urban sites (Chen et al., 2022) underlining the importance of highly processed secondary OA over Nicosia (Fig. 9).

![Figure 9: Relative contribution of PMF resolved OA sources to total OA for the cold period (left) and the warm period (right) respectively.](image)

### 3.6.2. Geographic origin of OA sources

The geographic origin of OA sources (local vs regional) is further assessed here using both Non-parametric Wind Regression (NWR) analyses as well as the regional scale coupling concentrations to air mass back trajectories through PSCF.

**Cold period:** During this period, primary OA factors, especially HOA-1 and BBOA, have an expected strong local component that is characterized by high concentrations at low wind speeds (hourly average 1.4 m s⁻¹) when winds are originating from the W-SW sector (Fig. S14a,c), pointing to the busy highway connecting Nicosia to the other major cities in the island while integrating the highly populated residential areas of Strovolos and Lakatamia municipalities. (Fig. 1c). As discussed earlier the HOA-2 factor, apart from its local influence (also in the W-SW sector), exhibits significant concentrations related to higher wind
s from the NW and the E-NE sectors that could originate from power plants but also possibly from long-range transport. Interestingly, a small local contribution from the city, still within the W-SW sector, can be also observed for both LO-OOA and MO-OOA, consistent with the peaks observed that could originate from local night-time chemistry. Still high concentrations of MO-OOA (and to a lesser extent LO-OOA) are observed with high wind speeds and Eastern directions (Fig. S14e,d). Although the contribution of the power plant PP5 located in the East sector (Fig. S13c) cannot be excluded, PSCF analysis points out that the hotspots of MO-OOA can be traced in neighbouring countries (eg. Syria, Lebanon, South Turkey) in the middle East (Fig 10a). These areas also represent hotspots of SO\textsubscript{4}\textsuperscript{2-} according to PSCF analysis (Fig. S17a).

**Warm period:** Given the generally higher wind speeds recorded, in comparison to the cold season (average of 1.93 m s\textsuperscript{-1} vs 1.36 m s\textsuperscript{-1} in the cold period), all OA factors show elevated concentrations coupled with higher wind speeds. The most striking result is the major influence of the E-SE sector for all OA sources, although this sector is upwind of Nicosia and therefore poorly influenced by local city emissions. As noted previously for the cold period, long-range transported OA from the Middle East is expected to be the main driver here to explain the influence of the E-SE sector, at least for LO-OOA and MO-OOA (Fig. S16c,d). This is once again confirmed from the PSCF results reported in Fig. 10b for the warm period. The HOA-1 factor still shows maxima for low wind speeds (<5 km h\textsuperscript{-1}) characteristic of local emissions and the SW-S direction, but also exhibits significant contribution related to the E-SE sector. Although the influence of the power plant PP5 on HOA-2 is expected, contribution of this source can not be exclude for HOA-1 as well. On the other hand, quantification of the Middle Eastern contribution to the HOA-2 factor remains to be assessed, since the current dataset cannot provide sufficient information on separating the contribution of power plants on the island versus more regional Middle East emissions (Fig. S165b). Although this hypothesis needs further investigation, the presence of HOA-2 in the Middle East would be consistent with recent findings highlighting the importance of OC emissions from diesel generators used in Lebanon as a mean of complementary power generation (Fadel et al., 2022).

![Figure 10: PSCF plots for MO-OOA during the cold and warm period. Colorscale represents PSCF values while the sampling site is denoted with a white dot.](https://doi.org/10.5194/egusphere-2022-1081)

In conclusion, based on the relative contribution of OA factors (Fig. 9) and the NWR analysis (Fig. S14, S16), it can be reasonably assumed that significant amount of measured OA in Nicosia is originating from long-range transport with the Middle East being the major source region, during both cold and warm periods. This is the first time that such high contribution of OA from the Middle East is highlighted over Cyprus. Assuming that biomass combustion and biogenic emissions of OA in the desert regions of the Middle East are relatively limited, these results suggest that most of primary and secondary OA originating from the Middle East could be of fossil fuel origin, which is consistent with the previously reported large use of oil in this region.
3.7. Spatial and seasonal variability of BC sources

The above conclusion on the influence of primary and secondary OA sources from the Middle East region, and its strong fossil fuel origin, motivates a careful examination of the geographic origin and sources of BC concentrations recorded in Nicosia.

Baseline (i.e. lowest) BC_{ref} concentrations are typically observed in the middle of the night and in the middle of the day, when local emissions are at their minimum (See Fig. 4). As such, these background concentrations can be considered as a first qualitative indicator of background BC_{ref} concentrations of regional origin. Interestingly, these baseline BC_{ref} concentrations appear to be in phase with those of sulphate (Fig. 11), as well as the MO-OOA factor derived from the OA PMF analysis. This observation points to the possible use of MO-OOA as a tracer for regional BC_{ref}. Hence, it brings further evidence on the importance of regional emissions on carbonaceous aerosol concentrations in Nicosia.

![Figure 11: Temporal variability of BC_{ref} and SO_{4}{^2}\textsuperscript{-} concentrations during the entire measuring periods.](image)

The assumption that transported regional pollution can affect BC_{ref} concentrations in Nicosia can be further supported by investigating the BC_{ref} NWR polar plots for both the cold and warm seasons (Fig. S18a,b). Elevated concentrations related to local emissions were observed for calm conditions with low wind speeds (<5 km h\(^{-1}\)) in the SW sector, as previously observed for HOA-1. Interestingly, BC_{ref} NWR plots show a distinct contribution at higher wind speeds (~15 km h\(^{-1}\)) and the NE-SE (Middle East) sector, during both the cold and warm periods, with estimated concentrations of roughly 1.5 μg m\(^{-3}\), further support the major role of the Middle East in the observed BC concentration levels in Nicosia (Fig S18 a,b).

**BC source apportionment**: In order to better assess the relative contributions of the multiple primary OA sources (HOA-1, HOA-2) and to quantify the contribution of long-range transport from the Middle East to BC_{ref}, a multilinear regression (MLR) model was tentatively performed using the principle of co-emission of BC_{ref} and organic species by the different sources (Chirico et al., 2010; Laborde et al., 2013). This approach, used recently by Poulain et al. (2021), assumes that at any given time \(t\), BC_{ref} mass concentration is the sum of BC from traffic (traced by HOA-1), from a mixed combustion source (traced by HOA-2), and from long-range transport (traced by MO-OOA), as follows:

\[
[\text{BC}]_{\text{ref}} = [\text{BC}]_{\text{traffic}} + [\text{BC}]_{\text{mix combustion}} + [\text{BC}]_{\text{regional}} \tag{2}
\]

With:

\[
[\text{BC}]_{\text{traffic}} = a \times [\text{HOA-1}] \tag{3}
\]
\[
[\text{BC}]_{\text{mix combustion}} = b \times [\text{HOA-2}] \tag{4}
\]
\[
[\text{BC}]_{\text{regional}} = c \times [\text{MO-OOA}] \tag{5}
\]

Where \(a\), \(b\), and \(c\) are coefficients derived from the multi-linear regression model.

The above approach assumes that primary HOA-1 and HOA-2 can trace BC_{traffic} and BC_{mix combustion}, respectively. This is somewhat expected for traffic which has a typical HOA-1/BC_{traffic} ratio with little variations. For HOA-2, this assumption is valid for the fraction that is assumed to originate from power plant emissions, and for some of the cooking activities (e.g. when using charcoal combustion) but not necessarily all. As such, uncertainties of this approach is expected to be higher for HOA-2, compared to HOA-1. The use of MO-OOA to trace the regional source of BC would probably lead to even higher uncertainties due to the fact that MO-OOA is also sensitive to atmospheric photochemical processes and does integrate multiple sources.
Nevertheless, this latter assumption is believed to be acceptable given the good agreement reported above between baseline concentrations of BC_{ff} and MO-OOA (Fig. S19), and the above conclusions that carbonaceous aerosol originating from the Middle East are expected to be dominated by fossil fuel combustion. Note that MO-OOA was preferred here to LO-OOA to trace regional emissions due to the later’s somewhat more local character.

Combining equations 2-5 provides the multilinear regression model with the free regression parameters \(a, b, c\) which are fitted to the time-resolved BC_{ff} mass concentration measured by the Aethalometer and PMF results for the ACSM data:

\[
[\text{BC}]_{ff} = a \times [\text{HOA}-1] + b \times [\text{HOA}-2] + c \times [\text{MO-OOA}] \quad (6)
\]

Previous studies have shown that MLR models seem to have enhanced explanatory power when primary emissions are dominant (Laborde et al., 2013). To reduce this potential bias, the MLR model was applied distinctly for the two seasons separately.

During the cold period, a very good correlation between measured and modelled BC_{ff} was obtained (\(r^2 = 0.70; N = 2942\)), with the modelled BC_{ff} explaining 84% of the measured one (Fig. S20a). The regression coefficients \(a\) (HOA-1), \(b\) (HOA-2) and \(c\) (MO-OOA) were found to be 1.11 ± 0.03, 0.15 ± 0.01 and 0.41 ± 0.01, respectively. Regarding the warm period, it was not possible to obtain a positive value for \(b\) (HOA-2). A correlation between long-range transported HOA-2 and MO-OOA is, among other, a reason that can be proposed to explain why it has not been possible to extract a BC_{mix source} factor here. Therefore, BC_{ff} was only apportioned using HOA-1 and MO-OOA. A good correlation between measured and modelled BC_{ff} was obtained (\(r^2=0.62; N=1251\)), with the modelled BC_{ff} explaining 83% of observations (Fig S20b). The regression coefficients \(a\) (HOA-1) and \(c\) (MO-OOA) were found to be 3.05 ± 0.07 and 0.19 ± 0.01, respectively.

The combination of the Aethalometer model (apportioning BC_{ff} and BC_{wb}) and the MLR model (apportioning BC_{traffic}, BC_{mix sources}, and BC_{regional}) was performed to obtain an integrated picture of BC sources in Nicosia for both periods (see Fig. 12).

**Figure 12:** BC sources during the cold and the warm period in Nicosia

Spatial and seasonal variability of BC sources: During the cold period, BC was found to originate from four different sources denoting the complexity of combustion sources of different origin in Nicosia. BC_{regional} is the dominant source of BC (37%), while traffic, wood burning, and mix source are estimated to contribute to 26%, 26% and 11% of BC, respectively. From the perspective of BC_{ff} sources, long-range transport, traced by MO-OOA, remains the largest source of BC_{ff} during the cold period, contributing 63%, while BC_{ff} from local emissions constrained with HOA-1 and HOA-2 represents 24% and 13%, respectively (Fig S21). In other words, more than half of BC_{ff} in Nicosia was found to be regional and most probably originating from the Middle East during the cold period. This high contribution of regional BC_{ff} is quite unexpected for a medium-size European city...
4. Conclusions

Near-real-time chemical composition of submicron aerosols and source apportionment of carbonaceous aerosols was performed for the first time in Nicosia, a medium-size European capital city (circa 250,000 inhabitants) in Cyprus located in the Eastern Mediterranean and surrounded by Middle East countries with fast growing population and increasing emissions of air pollutants. Continuous observations were performed at an urban background site, for approximately 6 months (between 7 December 2018 and 31 May 2019), in order to obtain a large and representative dataset capturing specific features related to cold and warm periods, such as domestic heating and regional transport. Measurements of the major fractions of PM$_1$ were carried out with a Q-ACSM and an Aethalometer complemented by a comprehensive suite of collocated instruments (e.g., filter sampling, SMPS) to further assess the quality the acquired data.

Unlike many European cities, no clear PM$_{1}$ pollution episodes of several consecutive days could be observed over Nicosia. However, very intense peaks (above 40 $\mu$g m$^{-3}$, 1h averages) were recorded systematically every evening during the cold period. Carbonaceous aerosols (BC and OA) were identified as the main components of these peaks and were mostly attributed to local emissions from heating with only little contribution from local meteorology (PBL height did not show significant diurnal variability during the cold period).

Source apportionment of OA has been used to derive a local biomass burning OA (BBOA$_{loc}$) mass spectrum, in order to properly apportion the contribution of domestic wood burning. A total of five OA sources were identified during the cold period, among which four are typically reported within urban environments (HOA-1, BBOA, LO-OOA, MO-OOA). An additional one (HOA-2) was assigned as a mixture of several combustion sources such as cooking as well as a significant contribution from power plants located in the Northern part of the island. These power plants in addition represent major island-based hotspots of NO$_x$, as evidenced from satellite observations. Interestingly, a similar HOA-2 source was identified at our regional background site (40 km distance from Nicosia; Chen et al., 2022), pointing to a possible influence from these power plants to an extended part of the island. The impact of this specific source brings the OA contribution of primary sources up to 40 % over Nicosia during the cold period. Few additional features were noticed for the other OA sources with 1) a typical traffic-related (HOA-1) source observed during both seasons, 2) a biomass burning source (BBOA) related to domestic heating enhanced at night during the cold season and accounting for 12 % of the total OA, 3) a less oxidized secondary (LO-OOA) source of a semi-volatile character and influenced by local night-time chemistry that shows to be more oxidized (and of more regional origin) during the warm period, and 4) a secondary (MO-OOA) source mostly of regional origin but also influenced by night-time chemistry during the cold period.

The geographic origin of each OA source was assessed for both seasons. With the exception of MO-OOA, which shows systematically a strong regional component, HOA-1, HOA-2, and LO-OOA exhibit a clear local origin during both seasons, and a more pronounced influence from the Eastern wind sector during the warm period. The prevalence of this sector is systematically observed for MO-OOA highlighting the major role of Middle East emissions in contributing to ca. half of OA concentrations in Nicosia during both cold and warm seasons.
To further elucidate the influence of this complex mixture of OA sources on BC levels, a source apportionment of BC was performed combining i) the aethalometer model to separate BC into its fossil fuel (BC_{ff}) and wood burning components (BC_{wb}), and ii) a multi-linear regression model to apportion the contribution to BC_{ff} from traffic (constrained by HOA-1), mix combustion sources from cooking and power plants (constrained by HOA-2), and long-range transport from the Middle East (constrained by MO-OOA). Although several assumptions and uncertainties are associated with this approach, it has shown to provide an interesting tool to reconstruct the BC concentrations derived experimentally. Such BC apportionment performed for both cold and warm seasons solidified the conclusions reached through the OA source apportionment, with almost half of BC_{ff} being of regional origin with the Middle East playing an important role. This result is quite unexpected given that local traffic emissions are usually considered the dominant contributor to BC_{ff} in urban background environments. These conclusions have numerous implications related to PM regulation and the efficiency of local abatement strategies (in particular regarding traffic emissions), health (combustion aerosols being considered as particularly adverse for human health), and climate (major influence of light absorbing aerosols from Middle East fossil fuel emissions).

More accurate OA and BC source apportionment i) with more co-located high resolution measurements of specific trace metal and organic tracers, ii) better resolved OA mass spectra (e.g., from HR-ToF-AMS), iii) the use of various source specific mass spectra fingerprints (e.g., from cooking or power plants), and iv) multi-site measurements (incl. both urban and regional background) will enable a more accurate estimation of local vs regional fossil fuel emissions in Cyprus while better constraining the current regional efforts on air quality modelling and forecasting.

Data availability: All data used in this study can be accessed here: https://doi.org/10.5281/zenodo.7186341. More details on the analyses are available upon request to the contact author Aliki Christodoulou (a.christodoulou@cyi.ac.cy).

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