1 Ambient carbonaceous aerosol levels in Cyprus and the role of 2 pollution transport from the Middle East.

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21 Abstract. The geographical origin and source apportionment of submicron carbonaceous aerosols (organic aerosols, OA, and

22 black carbon, BC) have been investigated here for the first time deploying high-time resolution measurements at an urban

23 background site of Nicosia, the capital city of Cyprus, in the Eastern Mediterranean. This study covers a half-year period,

24 encompassing both the cold and warm periods with continuous observations of the physical and chemical properties of PM₁

25 performed with an Aerosol Chemical Speciation monitor (ACSM), an Aethalometer, accompanied by a suite of various

26 ancillary off and on-line measurements. Carbonaceous aerosols were dominant during both seasons (cold and warm periods),

27 with a contribution of 57% and 48% to PM₁, respectively, and exhibited recurrent intense night-time peaks (>20-30 μg m⁻³)

28 during the cold period associated with local domestic heating. The findings of this study show that high concentrations of

29 sulfate (close to 3 µg m⁻³) were continuously recorded, standing among the highest ever reported for Europe and originating

30 from the Middle East region.

31 Source apportionment of the OA and BC fractions was performed using the Positive Matrix Factorization (PMF) approach and

32 the combination of two models (aethalometer model and multilinear regression), respectively. Our study revealed elevated

33 hydrocarbon-like organic aerosol (HOA) concentrations in Nicosia (among the highest reported for a European urban

34 background site), originating from a mixture of local and regional fossil-fuel combustion sources. Although air masses from

35 the Middle East had a low occurrence and were observed mostly during the cold period, they were shown to strongly affect

36 the mean concentrations levels of BC and OA in Nicosia during both seasons. Overall, the present study brings to our attention

37 the need to further characterize primary and secondary carbonaceous aerosols in the Middle East; an undersampled region

38 characterized by continuously increasing fossil fuel (oil and gas) emissions and extreme environmental conditions, which can

39 contribute to photochemical aging.

40 1. Introduction

41 At the crossroads of three continents (Europe, Africa, Asia), the Eastern Mediterranean and Middle East (EMME) region faces 42 many challenges, such as rapid population growth – with its currently 400 million inhabitants – as well as political and socio-

43 economic instabilities. Environmental conditions in the region are exceptional, with the two largest deserts worldwide (Sahara

- 44 and Arabian) being among the most water scarce ecosystems on the planet (Terink et al., 2013). Climate change in this region
- 45 is extraordinarily rapid; summer temperatures, in particular, are increasing by more than twice the global mean rate (Lelieveld
- 46 et al., 2014), with significant impacts, especially in urban areas (Mouzourides et al., 2015). While aerosol mass loadings over
- 47 the EMME are dominated by desert dust, concentrations of fine particles due to anthropogenic emissions are also high (Basart
- 48 et al., 2009) and will likely increase with continued population growth (Pozzer et al., 2012), making anthropogenic pollution
- 49 in the area a leading health risk and an important climate forcer (Osipov et al., 2022).
- 50 Based on modelling studies, it has been also concluded that the EMME is characterized by highly favourable conditions for
- 51 photochemical smog and ozone (O₃) formation leading to air quality standards being drastically exceeded (Lelieveld et al.,
- 52 2014; Zanis et al., 2014). These enhanced concentrations of fine particulates and ozone have major human health implications,
- 53 contributing to premature mortality (Giannadaki et al., 2014; Lelieveld et al., 2015), which may be further exacerbated by the
- 54 effects of heatwaves occurring during summer within the EMME region (Zittis et al., 2022).
- Although data from satellite observations of NO_2 and SO_2 has revealed strong air pollution trends in the Middle East since 2010 (Lelieveld et al., 2015a), many pollution sources are still missing in emission inventories (Mclinden et al., 2016). Thus, there is a current lack of a regional approach to characterize air pollution, with in-situ observation being insufficient, unavailable, or of low quality (Kadygrov et al., 2015; Ricaud et al., 2018; Paris et al., 2021), limiting the possibility to reduce uncertainties in regional emission inventories and implement efficient abatement strategies.
- 60 Significant efforts have been put forward in recent years to characterize the atmospheric composition in-situ over Cyprus, a central location of the EMME region (e.g., Kleanthous et al., 2014; Debevec et al., 2017 and 2018; Pikridas et al., 2018; Dada 61 62 et al., 2020; Baalbaki et al., 2021; Vrekoussis et al., 2022). In-situ ground-based PM observations have clearly shown that 63 contributions of dust to PM₁₀ over Cyprus are among the highest for the entire Mediterranean basin (Querol et al., 2009; Pey 64 et al., 2013; Kleanthous et al., 2014; Pikridas et al., 2018; Achilleos et al., 2020), during dust storm events, leading to increased 65 hospitalization, particularly attributed to cardiovascular-related diseases (Middleton et al., 2008; Tsangari et al., 2016) and short-term effects associated with daily mortality (Neophytou et al., 2013). These high levels of regional particulate matter are 66 67 responsible for exceedances in PM10 EU limits in major Cypriot cities (Querol et al., 2009). Past studies on PM trends and sources highlighted the important contribution of local (urban) emissions to PM₁₀ (Achilleos et al., 2014; Pikridas et al., 2018) 68 69 but also showed a predominant regional pattern for $PM_{2.5}$ with a major contribution of sulfur-rich sources (Achilleos et al., 70 2016). Based on 17 years of continuous observations of reactive gases in Cyprus, Vrekoussis et al. (2022) further confirmed 71 the major contribution of long-range transport (incl. Middle East) in the observed concentration levels of carbon monoxide
- 72 (CO) and sulfur dioxide (SO₂), two tracers of combustion sources.
- Those studies have highlighted the unique location of Cyprus as a receptor site of major regional pollution hotspots, making the island one of the most polluted EU member states in terms of PM and O₃ concentrations, the only one impacted by longrange transport of poorly-regulated air pollutants originating from Middle East countries. However, few studies are currently available to assess the contribution of regional anthropogenic emissions to PM levels in Cyprus. The filter-based chemical speciation study reported by Achilleos et al., (2016) is currently the most exhaustive one and was based on 24-h integrated ($PM_{2.5}$ and PM_{10}) filter samples collected every 3 days for a period of one year (2012) in four cities in Cyprus. This study concluded that Cypriot cities, like many others in Europe, are characterized by a major contribution of regional sulfate and
- 80 local (urban) emissions from traffic and domestic heating biomass burning.
- 81 Herewith, a detailed description of submicron (<1µm, PM₁) chemical composition and the further source apportionment of BC
- and OA is presented for the first time in Cyprus. State-of-the-art on-line instrumentation (e.g., Q-ACSM, Aethalometer) was
- 83 deployed to investigate the temporal variability of aerosol composition at a location representative of the urban background
- 84 pollution in the capital city of Nicosia. Source apportionment of submicron organic aerosols was performed using the organic
- 85 fragments of the ACSM and Positive Matrix Factorization (PMF). The consistency of these results was assessed against the
- 86 chemical analysis of parallel filter samples and on-line measurements of external tracers. This study was extended to a 6-month

87 duration in order to cover the two main seasons of the semi-arid Eastern Mediterranean climate (short, mild and wet winter vs.

88 long, hot and dry summer), offering a comprehensive understanding of the daily and monthly variability of local and regional

89 sources of carbonaceous aerosols. Cold and warm periods were compared to highlight the complexity of local (combustion)

90 sources and the importance of regional ones. These results were further processed to apportion Black Carbon sources in Nicosia

91 with emphasis on local versus regional contribution.

92 2. Material and Methods

93 2.1 Sampling site

94 *Cyprus*: Cyprus is the third largest island in the Mediterranean Sea, extending approximately 240km long from east-to-west 95 and 100km wide. The closest countries and their distance from the capital city of Nicosia are respectively Turkey (110km), 96 Syria (250km), Lebanon (250km), Israel (300km), Egypt (400km), Jordan (430 km), and Greece (900 km from the Greek 97 mainland), (Fig. 1a).

The population of Cyprus (approximately 1 million inhabitants) is rather small compared to its neighbouring countries and the rapidly growing (overall 400 million) population of the region (Lelieveld et al., 2013). The main urban areas of the island shown in Fig. 1b, are those of Nicosia (c.a. 245,000 inhabitants), Limassol (c.a. 150,000 inhabitants), Larnaca (c.a. 50,000 inhabitants) and Paphos (c.a. 35,000 inhabitants). Cyprus has a Mediterranean and semi-arid climate with two main seasons: a mild cold season (from December to March) and a hot warm season lasting about eight months (from April to November).

Rain occurs mainly in the cold season, with the warm one being extremely dry (i.e., almost no rain between May andSeptember)(Michaelides et al., 2018).

- *Nicosia*: Nicosia is the largest city on the island and the southeasternmost of the European Union Member States' capitals.
 Nicosia is currently partitioned in two, with a buffer zone in-between under the control of the United Nations; the southern part being the capital of the Republic of Cyprus. The northern part of Nicosia (and the northern part of the island) is not controlled by the government of the republic of Cyprus (Resolution 550, UN security council, 1984) (Fig. 1c). Geographically, Nicosia is located in the centre of the island, within the Mesaoria plain, 150 m above sea level (asl), which is delimited on its northern and southern edges by two mountain ranges; the Kyrenia Range culminating at 1,024 m asl, and the Troodos
- 111 Mountains culminating at 1,952 m asl, respectively. This topography channels winds within a more or less west-east corridor
- 112 (Fig. S5), feeding the city of Nicosia with long-range transported air masses from Europe, Africa, or the Middle East.
- Measurements were performed at the Cyprus Atmospheric Observatory's Nicosia station (CAO-NIC) located at the Cyprus Institute premises (Athalassa Campus; 174 m asl; 35.14N, 33.38E; Fig. 1c). The measurement site is considered an urban
- 115 background site, located within a low population density residential area with no significant local pollution hotspots in its
- 116 vicinity (i.e., no dense road traffic, industry, commercial centers, restaurants, etc.) and next to the Athalassa Forestry Park.
- 117 The period and duration of measurements presented here (07 December 2018 31 May 2019) were chosen to i) capture weather
- 118 conditions, atmospheric dynamics, and long-range pattern of the two main seasons, ii) investigate the contribution of domestic
- 119 heating emissions in winter, and iii) assess the potential increasing contribution of photochemical produced secondary aerosols
- 120 during the start of the dry and warm season. Local time (LT) in Cyprus is given as Eastern European Standard Time (EET)
- 121 (UTC+02:00 in winter and UTC+03:00 during the summer).



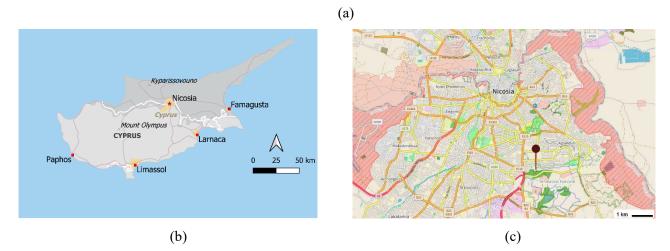


Figure 1: (a) Geographic location of the island of Cyprus and its closest Northern African and Middle Eastern neighbouring countries. (b) Location of the main cities of the Republic of Cyprus. Maps a,b were created by QGIS software v.3.26.3 utilizing the Natural Earth data (https://qgis.org). (c) Satellite view of the Nicosia agglomeration (grey area). The buffer zone dividing the island and the city is marked with red stripes; the location of the measurement site (CAO-NIC; The Cyprus Institute, Athalassa campus) is noted in red. (© OpenStreetMap contributors 2022. Distributed under the Open Data Commons Open Database License (ODbL) v1.0)

128 2.2 On-line Aerosol Instrumentation

129 On-line aerosol instrumentation has been operated following the Standard Operating Procedures defined by the European

130 Research Infrastructure on Aerosols, Clouds, and Trace Gases ACTRIS (<u>https://www.actris.eu</u>), and Cost COLOSSAL

131 (CA16109, 2021).

Non-refractory submicron (NR-PM₁) aerosol chemical composition, i.e. organics, sulfate, nitrate, ammonium and chloride, was continuously monitored using a Quadrupole ACSM (Aerosol Chemical Speciation Monitor; Aerodyne Research Inc.) at a 30-min time resolution (Ng et al. 2011a). The instrument, along with a scanning mobility particle sizer (SMPS, described below), sampled through a sharp cut cyclone operated at 4 L min⁻¹ (SCC 1.197, BGI Inc., USA), and was equipped with a PM₁ aerodynamic lens, yielding an aerosol cut-off diameter of approximately 1.3µm. Data were retrieved using ACSM local

- 137 v.1.6.0.3, implemented within Igor Pro (v. 6.37, Wavemetrics Inc., USA). The ACSM is designed and built around similar
- 138 technology as the aerosol mass spectrometer (Jayne et al., 2000), where an aerodynamic particle focusing lens is combined
- 139 with particle flash vaporization in high vacuum on the surface of a standard tungsten vaporizer heated at 600 °C, followed by
- 140 electron impact ionization, separation and final detection of the resulting ions using a quadrupole mass spectrometer. Mass
- 141 concentrations are corrected for incomplete detection due to particle bounce using the chemical composition-dependent
- 142 collection efficiency (CDCE) (Middlebrook et al., 2012). The determined parameters, response factor (RF) and relative
 - 143 ionization efficiency (RIE) are reported in table S2.
 - 144 Black carbon (BC) measurements were conducted using a 7-wavelength aethalometer (AE-33 Magee Scientific, US) at a 1-
 - 145 min time resolution. The aethalometer sampled ambient aerosol through a PM_{2.5} aerosol inlet (SCC 1.829, BGI Inc., USA) at

- 146 a flow rate of 5 L min⁻¹ after passing through a nafion dryer. The instrument internally corrected the filter loading effect in
- 147 real-time, while a fixed value ($C_0=1.39$) was applied to compensate for the multi-scattering effect (Drinovec et al., 2015). BC
- 148 was apportioned to source specific components, namely BC_{ff} related to fossil fuel combustion and BC_{wb} related to wood
- 149 burning, by applying the "aethalometer model" (Sandradewi et al., 2008) on the 470 950 nm wavelength pair. The
- 150 instrument's default values for fossil fuel combustion and wood-burning aerosol Absorption Ångström Exponent, AAE_{ff}=1
- 151 and AAE_{wb}=2, respectively were selected after performing a sensitivity analysis on the AAE values (Supplement Section 3).

152 2.3 Ancillary measurements

SMPS: Particle number size distributions were monitored using a scanning mobility particle sizer (SMPS) consisting of an 153 electrostatic classifier (model 3080, TSI Inc., USA) coupled with a condensation particle counter (CPC; model 3070, TSI Inc. 154 155 USA) operating at a 5-min time resolution and at a 1 L min⁻¹ sample flow rate, measuring particles with a diameter ranging 156 from 9 to 700 nm. Ambient aerosols were drawn through a nafion dryer, and placed upstream, keeping sample RH below 30 157 %. Volume concentrations of assumed spherical particles derived by the SMPS were converted into mass concentrations using 158 a variable density calculated by the methodology described in Bougiatioti et al. (2014). The respective mass fractions time 159 series of chemical species were calculated based on the ACSM measurements. A density value of 1.77 g cm⁻³ was used for 160 ammonium sulfate, and 1.35 g cm⁻³ for organics (Florou et al., 2017; Lee et al., 2010), the two dominant compounds of PM_1

- 161 in Nicosia as detailed further below.
- 162 <u>Filter sampling:</u> Co-located 24h PM_{2.5} samples were collected on quartz fiber filters (Tissuquartz, 47mm diameter, Pall) using 163 a low volume sampler (Leckel SEQ47/50) operating at a flowrate of 2.3 m³ h⁻¹. The filter samples were analysed for i) organic 164 and elemental carbon using an OC/EC Lab Instrument (Sunset Laboratory Inc., OR, USA) implementing the EUSAAR II 165 protocol (Cavalli and Putaud, 2008), ii) carbohydrates, including levoglucosan, mannosan, galactosan, using an Ion 166 Chromatography Pulsed Amperometric Detection method (Thermo - Model ICS-3000) and iii) anions (Cl⁻, NO₃⁻, SO₄²⁻, MSA, 167 Oxalate) and cations (K⁺, Na⁺, NH₄⁺, Mg²⁺, Ca²⁺) using ion chromatography (Thermo - Model ICS-5000).
- 168 Proton Transfer Reaction - Mass Spectrometry (PTR-MS): Air was sampled through a 20m long, 3/8" o.d. (1/4" i.d.) 169 sheathed Teflon line that ran from the roof of the building to the instrument. A Teflon filter (0.2µm diameter porosity) was 170 installed at the inlet to prevent large aerosol particles and insects from entering the sampling line. The resulting residence time 171 of air in the line was estimated to be approximately 0.5 min. The temporal resolution of Volatile Organic Compounds (VOCs) 172 measured by the PTR-MS (Ionicon Analytik, Austria) was approximately two minutes (the time required to measure 55 different ions at 2 seconds per ion). The basic operation principles of the PTR-MS instrument have been described in detail by 173 174 Lindinger et al. (2011). Briefly, a stable flow of air and high concentrations of H_3O^+ ions are continuously sampled into a drift 175 tube held at 2.2 mbar pressure. There, compounds with a proton affinity greater than water, including a large selection of 176 Oxygenated Volatile Organic Compounds (OVOCs), undergo efficient proton-transfer reactions with the H₃O⁺ ions to produce 177 protonated organic product ions, which can be detected by a mass spectrometer.
- Meteorological Parameters: Standard meteorological parameters (temperature, relative humidity, wind speed and direction) were obtained from the meteorological station of the Cyprus Department of Meteorology, installed 10 m above ground, located at the Athalassa Forestry Park (164 m asl) lying approximately 1.3 km east of the CAO-NIC station. Wind speed and direction data were further used in this study for component-specific non-parametric wind regression analysis (NWR) performed using the ZeFir toolbox (Petit et al., 2017) developed within the Igor Pro software (Wavemetrics Inc.). A co-located automatic CIMEL CE370 micro-LIDAR was operated continuously to retrieve the Planetary Boundary Layer Height (PBLH) and better assess the influence of atmospheric dynamics on in-situ ground-based observations.
- Air masses back trajectory analysis: Five-day air mass back trajectories arriving at 1000m altitude above the sampling site
 were computed every 6 hours, using the Hybrid Single-Particle Lagrangian Integrated Trajectory model (HYPLIT4; Stein et
- 187 al., 2015) using the Global Data Assimilation System (GDAS 1) meteorological data fields (with 1° spatial resolution). Back

- 188 trajectories were coupled to measured concentrations, assessing origins and source contributions to specific chemical
- 189 components, by applying the Potential Source Contribution Function (PSCF) technique as implemented in the ZeFir toolbox
- 190 described above.

191 **2.4. Source Apportionment analysis**

Positive Matrix Factorization (PMF) is an advanced multivariate factor analysis tool that attempts to identify the contributing
 factors, or sources, of atmospheric pollutants at a sampling site. For this study, source apportionment was performed on the

- 194 organic mass spectra dataset collected by the ACSM. The (PMF) method (Paatero and Tapper, 1994) using the multilinear
- 195 engine (ME-2) model developed by Paatero (1999) was implemented using the SoFi (Source Finder) toolkit (SoFi 6D;
- 196 Canonaco et al., 2013). PMF allows the decomposition of the OA mass spectra matrix X into two matrices, G and F and a
- 197 remaining residual matrix, E:
- 198 X = G * F + E (1)
- 199 Where X is the input dataset matrix (measured quantity), F is the resulting source profile matrix, G is the source contribution 200 matrix (temporal variability of each source), and E represents the model residual matrix. Based on a number of criteria, the 201 optimal solution is selected, aiming at being physically meaningful that can be supported by external indicators (ancillary 202 measurements), and trying to minimize values in the residual matrix E. Model input data and error matrices (in µg m⁻³), were 203 exported using the ACSM software. Data points with a signal-to-noise (S/N) ratio smaller than 0.2 were removed. Points with 204 S/N between 0.2 and 2 were down-weighted by increasing their estimated error values (Ulbrich et al., 2009; Paatero and Hopke, 205 2003). M/z (mass-to-charge ratio) values ranging from 10 to 120 were used in the analysis. CO₂-related variables were excluded from the PMF and finally reinserted into the solution. 206
- 207 Source apportionment of OA was performed following the general steps described by Crippa et al. (2014) and the recently 208 updated harmonised standard operating procedures for seasonal OA PMF (Chen et al., 2022). As a first step, unconstrained 209 PMF analyses were performed with a number of factors ranging from 2 to 8 in order to identify the most relevant number of 210 factors and potential sources. If primary organic aerosol factor profiles such as Hydrocarbon-like OA (HOA) or biomass burning-like OA (BBOA) were found, then the corresponding site-specific primary OA (POA) mass spectra (see discussion 211 212 below) or spectra found in the literature (e.g., Ng et al., 2011 and Crippa et al., 2014) were set as constrains in the PMF, using 213 the "a-value" approach (Paatero and Hopke, 2009; Canonaco et al., 2013). A sensitivity analysis was then performed with 214 different a-values to assess the level of constrain introduced in each factor with i) a constrained HOA using, as an anchor the 215 HOA spectrum found in Ng et al. (2011) with the a-values ranging between 0.05 and 2.0, ii) a constrained BBOA factor with 216 the a-values from 0.2 to 0.5 from Ng et al. (2011), and iii) a constrained cooking OA (COA) factor from Crippa et al. (2014) 217 with a-values from 0.2 to 0.5. Once this sensitivity analysis was completed, the evaluation of the PMF results showed that the 218 BBOA factor could not account for the entire m/z 60 mass fragment, which fragment was distributed within 2 factors. 219 Additionally, the correlation of BBOA with BC_{wb} showed to be unsatisfactory (section S1). On the other hand, given the 220 BBOA factor's sensitivity to the type of solid fuel used, different biomass-burning factor profiles have been reported in various 221 regions around the world (Xu et al., 2020; Trubetskaya et al., 2021). Consequently, a site-specific BBOA factor profile 222 (BBOA_{cv}) was selected. The BBOA_{cv} spectrum was calculated as an average of 20 PMF runs from the initial unconstrained 223 PMF for the cold period, validated by it's time-series correlation to BCwb. Since aged OA (i.e. Oxygen-like OA, OOA) factors 224 show more variability between measurement sites in terms of their mass spectra, no constrain was introduced for these factors 225 (Canonaco et al., 2015).
- 226 In this study, the BBOA factor a major contributor of OA during winter could not be properly resolved when performing
- 227 the PMF analysis on the entire period dataset. A seasonal approach was followed instead, separating the OA dataset into two
- 228 periods that were then used to describe both the two periods (cold and warm, respectively). The criteria used to delineate those
- two periods are presented and discussed in section 3.2.

- 230 One factor was consequentially constrained with the resulting BBOAcy spectrum (with an a-value in the 0-0.5 range, using
- 231 steps of 0.02), obtaining the optimal solution using an a-value equal to 0.46. A widely referred-to standard mass spectrum (Sun
- et al., 2016; Duan et al., 2020) derived from Ng et al. (2011) was used to constrain the HOA factor, with an a-value of 0.2,
- 233 thus obtaining the best correlation with BC_{ff}, a tracer for traffic-related emissions. A detailed description of the OA source
- apportionment analysis can be found in section S1 in the supplementary material.

235 3. Results and Discussion

236 **3.1. On-line aerosol data quality check**

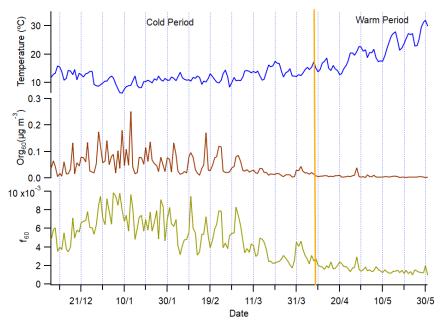
- 237 A chemical mass closure exercise for PM_1 was performed at a temporal resolution of 1h to check the quality of the on-line aerosol measurements. Chemically reconstructed PM1 was calculated as the sum of the mass concentration of all non-refractory 238 species measured by the ACSM (OA, NO_3^- , SO_4^{2-} , NH_4^+ , Cl^-) plus the BC concentrations measured by the Aethalometer AE-239 33 (Putaud et al., 2004). The contribution of other chemical constituents to submicron aerosols, such as sea salt and dust 240 241 (measured by co-located filter sampling), was found to be low and therefore neglected here. A scatter plot of the ACSM + AE-242 33 measurements vs. the SMPS-derived PM₁ concentrations is shown in Figure S4a. The results indicate a very good correlation ($r^2 = 0.88$; N=1823) and a slope of 1.2 (Fig. S4a). This 20% discrepancy lies within the uncertainty of the on-line 243 244 instruments. It could be attributed to the cut-off size of the SMPS at 700nm, which is slightly lower compared than the ACSM. 245 In addition, ACSM individual chemical species were compared with co-located off-line analyses performed on daily PM_{2.5} 246 filters. As shown in Fig. S4b-e, very good agreement was obtained between on-line and off-line measurements with $r^2 \ge 0.80$ (N=165-175) for all species. The discrepancy between ACSM and filter measurements for nitrate (slope of 1.3) could be 247 248 attributed to the volatilization of HNO₃ from the filter surface due to the presence of semi-volatile ammonium nitrate. The 249 obtained slopes for ammonium and sulfate below 1:1 (0.81 and 0.85, respectively) are consistent with the fact that fine 250 (NH₄)₂SO₄ aerosols, mainly originating from secondary processes and long-range transport (Sciare et al., 2010; Freutel et al., 251 2013), can be found at a large size mode possibly exceeding 1 μ m, consequently not being sampled by the ACSM.
- The study investigated the aerosol ion balance using both online and offline inorganic measurements. The ratio of the measured concentration of $NH_4^+_{Measured}$ and the estimated concentration of $NH_4^+_{Predicted}$, as calculated in Jiang et al., (2019), was used for this purpose. The results showed a slope of 0.80 for online measurements and 0.96 for offline measurements. These findings, suggest that the atmospheric aerosol observed during the study period was predominantly neutral, taking into account the uncertainties of ammonium concentrations reported in Q-ACSM intercomparison studies (Crenn et al., 2015), as well as the species' relatively high detection limit (Ng et al., 2011).
- An interesting result obtained from the comparison of OA (ACSM) with OC (from filters) is an OM-to-OC ratio of 1.42 which is at the lower end of ratios reported for urban environments, which usually exhibit typical values of 1.6 ± 0.2 (Petit et al., 2015; Theodosi et al., 2011; Brown et al., 2013). Without neglecting the fact that two different size fractions were compared (PM₁ for the ACSM and PM_{2.5} for the filter sampling), this low ratio probably point to long-chain hydrocarbon OA that often are related to primary combustion (poorly oxidized) OA (Aiken et al., 2008). As such, this ratio could represent an independent
- 263 means of verification of the consistency of our source apportionment between primary and secondary OA.
- 264 Finally, black carbon concentrations derived from light absorption measurements (Aethalometer AE-33) were compared
- against filter-based EC measurements (see Fig. S4f). Data from the two techniques correlate very well ($r^2=0.83$), with a BC/EC
- 266 ratio of 1.67 being similar to studies in other urban areas (Rigler et al., 2020; Liu et al., 2022), highlighting the existence of a
- 267 BC absorption enhancement (E_{abs}) attributable to a lensing effect induced by other chemical species, among which secondary
- 268 OA may play an important role (Zhang et al., 2018).

269 3.2 Meteorological conditions

270 Delineation of cold vs. warm seasons: The ACSM organic mass at m/z 60 is characteristic of the fragmentation of

271 levoglucosan, a product of cellulose pyrolysis and well-established biomass burning marker (Alfarra et al., 2007). Its respective contribution to total OA (f_{60}) was used in this study as an indicator of biomass burning for domestic heating to delineate cold 272 273 vs. warm seasons, comparing with the 0.3% threshold proposed by Cubison et al. (2011) for air masses influenced by biomass burning. Except for a single small peak in early May, corresponding to open fires for the celebration of the Greek Orthodox 274 275 Easter, the last instance when f_{60} was above the threshold was recorded during the first week of April (Fig. 2). From then 276 onwards, daily air temperature started rising constantly, from about 15°C at the beginning of April up to to 30°C at the end of May. These two features dictated the division of the dataset into two periods: a cold period of four months (07/12/2018-277 08/04/2019), with an average temperature of $12 \pm 4^{\circ}$ C, and a warm period of two months (09/04/2019 - 31/05/2019), with an 278

279 average temperature of $20 \pm 7^{\circ}$ C.



280

Figure 2: Time series of air temperature (blue), m/z 60 organic concentration (org60, brown) and f_{60} fragment (green) for the cold and warm periods. The vertical line is used to delineate the measurements within the two seasons.

Wind sectors: During these two periods, a distinct pattern in the wind sectors and the air masses arriving at the sampling site was observed. As seen in Fig. S5, the dominant wind direction for the cold period was the NW-SW [225 $^{\circ}$ - 315 $^{\circ}$] sector encompassing 48% of the total wind directions, while the NE-SE [45 $^{\circ}$ - 135 $^{\circ}$] sector covered 26%. During the warm period, the weight of this proportion is shifting even more towards the NW-SW [225 $^{\circ}$ - 315 $^{\circ}$] sector, having a 62% of total air masses while only 17% are arriving from the NE-SE [45 $^{\circ}$ - 135 $^{\circ}$] sector.

Air mass origin: A cluster analysis was performed (Fig. S6a,b) for both periods in order to better assess the main upwind 288 289 regions responsible for long-range transported air pollution over Cyprus and their change relative to the period of the year. The number of clusters used in each season was determined by considering the percentage change in Total Spatial Variance 290 291 (TSV) as a function of the number of clusters of merged trajectories (Fig. S6c,d) and the mean trajectory paths of each cluster 292 (Fig. S6e,f). The first large drops observed in TSV from the two - to - three and the three - to - four cluster transition could 293 not represent all the recorded trajectories and especially the ones describing air masses arriving in Nicosia from the east. The 294 next remarkable decrease in TSV was recorded when moving to seven clusters. Thus, for both periods, seven clusters were 295 chosen to better represent all air masses arriving in Nicosia.

A significant part of most of the calculated mean trajectory path representing clusters arriving in Cyprus was found to be related to the wider western sector, with many of them though, passing over Turkey before reaching Cyprus. Interestingly, this analysis showed one cluster (Cluster 1) arriving from the Middle East (close to Lebanon and Syria) and another four (Cluster

299 1, 2, 5, 6) passing over the western part of Turkey for the cold period. For the warm period, the only clusters arriving from the

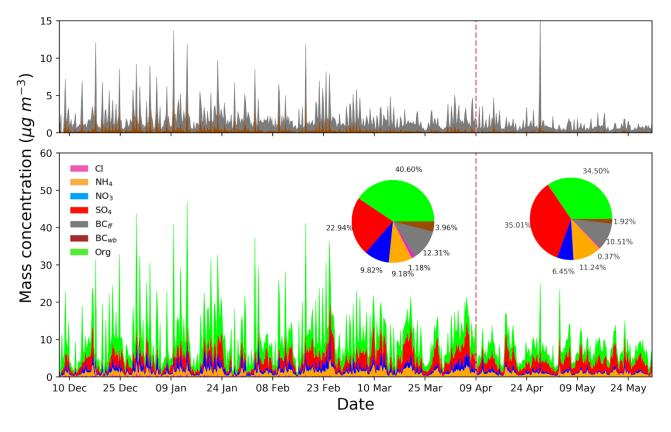
300 Middle East were the ones related to Turkey (Clusters 1, 5, 6). Plotting all individual 72h back trajectories (Fig S6e,f) showed

301 that a clear portion (almost 25% of the calculated trajectories) are being influenced by the Middle East, especially for the cold

302 period (Fig S6e).

303 3.3. Chemical composition of PM₁

- 304 Seasonal perspective of PM₁: The time series of PM₁ chemical composition derived from the ACSM (OA, SO₄²⁻, NH₄⁺, NO₃⁻
- 305, Cl⁻) and the Aethalometer (BC_{ff}, BC_{wb}) are depicted for the entire measuring period in Figure 3. Averaged data (6h averaging 306 period) are shown here for clarity. Furthermore, the relative average contribution of each chemical constituent to total PM₁
- 306 period) are shown here for clarity. Furthermore, the relative average contribution of each chemical constituer
- 307 concentrations is depicted in the respective inner pie charts for both periods.



308

Figure 3: Stacked area plots of the chemical composition time series for PM₁ in Nicosia derived from 6-hour averages of ACSM and AE-33 measurements. The vertical dashed red line separates the cold from the warm season. The average relative contribution of each species is shown in the respective pie charts (inner panels) for each season.

312 Although intense and short-duration peaks are observed for carbonaceous aerosols (OA, BC_{ff}, BC_{wb}), background NR-PM₁ concentration levels (between peak values) remain well below 10 µg m⁻³ for the 6-h average in both seasons. In other words, 313 314 no PM₁ pollution episodes (with e.g., concentrations above 10 μ g m⁻³) lasting for consecutive days were observed. Such lack 315 of intense and persistent PM₁ pollution episodes differs from what is reported in central and northern Europe, where stagnant 316 (anticyclonic) conditions occur together with continental (polluted) air masses, mainly in winter and springtime (e.g., Petit et al., 2015). This suggests that the relatively low emissions from Cyprus (compared to the neighboring countries) and its remote 317 318 marine location (i.e., far from densely populated areas) may prevent the build-up of high PM₁ pollution events over Nicosia. 319 On the other hand, clear differences can be observed between both periods, with significantly higher PM₁ concentrations during the cold period, associated with repeated, intense peaks of OA and BC - not observed during the warm season - and suggesting 320 local combustion emissions. The highest PM₁ concentrations were observed between December 28th 2018 and January 13th 321 322 2019 (Fig. 3) and were associated with low temperatures and Christmas holidays, both likely to promote domestic heating use. 323 During the warm period, the higher contribution of sulfate, and lower contribution of OA, are clearly noticeable. The

- 324 contribution of nitrate during the warm period, most probably in the form of semi-volatile NH_4NO_3 , remains marginal, possibly
- 325 due to non-favourable thermodynamic conditions preventing its formation and accumulation.
- 326 **PM**₁ **chemical composition**: For the cold period, the average calculated mass concentration of PM₁ (calculated as the sum of
- 327 chemical components measured by AE-33 and ACSM) was $12.35 \pm 9.77 \ \mu g \ m^{-3}$, with $10.34 \pm 7.92 \ \mu g \ m^{-3}$ being the average
- 328 concentration of the non-refractory species (Table 1). OA constitutes the larger fraction of PM_1 mass, with an average
- 329 concentration of $5.03 \pm 5.48 \ \mu g \ m^{-3}$ (41 %), followed by sulfate (23 %), black carbon (16 %), nitrate (10 %), ammonium (9
- 330 %), and chloride (1 %). These concentrations and the overall distribution of chemical components in NR-PM₁ are similar to
- 331 those measured by ACSM in other European cities (Bressi et al., 2021). Concentrations appear to decline during the warm
- 332 period, with an average calculated PM₁ concentration of $8.18 \pm 4.65 \ \mu g \ m^{-3}$, including $7.18 \pm 3.81 \ \mu g \ m^{-3}$ from the non-
- refractory components. The dominant species during the warm period were sulfate and OA, each representing 35 % of PM₁,
- followed by black carbon (12 %), ammonium (11 %) and nitrate (6%). During that period, chloride concentrations were
- negligible, contributing less than 1 % (Table 1).
- Table 1: Species mean, standard deviation, median concentrations and respective contribution to PM₁ during cold and warm periods
 in Nicosia.

Cold Period				Warm Period				
μg m ⁻³	Mean	Std	Median Contribution (%)		Mean	Std	Median	Contribution (%)
OA	5.03	5.48	3.35	41	2.83	1.91	2.51	35
SO 4 ²⁻	2.84	1.89	2.60	23	2.87	1.50	2.61	35
NO₃ ⁻	1.22	1.25	0.75	10	0.53	0.56	0.34	6
NH_4^+	1.14	0.77	1.01	9	0.92	0.55	0.84	11
Cl⁻	0.14	0.21	0.07	1	0.03	0.08	0.01	<1
BC	2.01	2.31	1.26	16	1.01	1.46	0.66	12
PM ₁	12.35	9.77	10.01	100	8.18	4.65	7.53	100

338

339 Interestingly, sulfate concentrations recorded in Nicosia are higher compared to what is commonly observed in other European 340 countries and Mediterranean cities (Table 2) and likely reflect a regional pattern of sulfur-rich emissions compared to Europe, where SO_2 emissions have strongly decreased during the last decades (Smith et al., 2011; Chin et al., 2014) thanks to the 341 342 implementation of specific abatement measures on reducing sulfur emissions (European NEC Directive (EU, 2016) and United 343 Nation Gothenburg (1999) protocol). More specifically, the importance of sulfur emissions in Turkey (2 455 Gg, EEA 2021), which were 50% higher compared to the total SO_x emissions of the EU 28 in 2019, together with the fact that half of the air 344 345 masses reaching Cyprus are passing over Turkey (see Fig. S6) are key contributors to the high concentrations of sulfate in our 346 study.

347 Shipping emissions appear to have a relatively minor impact on the concentration of sulfate. To more accurately determine the 348 contribution of shipping emissions to SO₄⁻², SO₂, and total PM_{2.5} a supplementary analysis was conducted using the WRF-349 Chem model, which simulates both physical and chemical processes occurring in the atmosphere. This model has been 350 extensively evaluated in several studies for the Eastern Mediterranean (Kushta et al., 2018) and Europe (Berger et al., 2016; Tuccella et al., 2012). Following the set-up used in Giannakis et al., (2019) and driven by the EDGAR v.5 anthropogenic 351 352 emission inventories (Crippa et al., 2019), two annual-long simulations were performed: firstly, including all sectoral emissions 353 in the model (baseline simulation So) and a second simulation where shipping emissions have been omitted (scenario 354 simulation, S₁) to identify the impact of shipping on gaseous and aerosol sulfur-related species concentrations (SO₂ and SO₄⁻ 355 ²) and total PM_{2.5} over the Central and Eastern Mediterranean. The figures (S final) describe the contribution of shipping in absolute terms (Fig. S7 a,c,e) and as a percentage (Fig. S7 b,d,f) for the SO_4^{-2} , SO_2 and total $PM_{2.5}$ calculated for each species. 356 357 According to these results, the highest impact of shipping on near-ground modelled concentrations of the three species (SO4-

358 ², SO₂ and PM_{2.5}) was estimated along the central Mediterranean region (yellow grids, west of the Balkans and Greece), as

- 359 well as a small section south of Greece. The Levantine basin, where Cyprus is located, experiences significantly lower
- 360 influence under the no-shipping emissions sensitivity test. More specifically, over the East Mediterranean, SO₄-² concentrations
- 361 represent a relative change of only about 6-8% when including shipping emissions.
- 362

363	Table 2: Comparison of concentration, and percentage contribution to PM ₁ , between the main submicron chemical species derived
	by ACSM.

	PM_1	OA	SO4 ²⁻	$\mathrm{NH_4}^+$	NO ₃ -	Cl	Reference	
	(µg m ⁻³)	$(\mu g \ m^{-3})$	(µg m ⁻³)	Reference				
Nicosia Cold (DJFM)	12.32	5.03	2.81	1.14	1.22	0.31	This study	
Nicosia Warm (AM)	8.18	2.83	2.87	0.92	0.53	0.05	This study	
Cyprus RB* (Annual)	7.6	3.26	2.66	0.98	0.23	-	Chen et al. (2022)	
European UB** (Annual)	10.6	5.3	2.0	-	1.9	-	Bressi et al. (2021)	
S. Europe RB*** (Annual)	6.3	3.5	1.3	-	0.8	-	Bressi et al. (2021)	
Athens Winter	18.7	13.13	2.4	-	1.8	0.14	Stavroulas et al. (2019)	
Athens Spring	6.42	3.3	2.1	0.6	0.4	0.02	Stavroulas et al. (2019)	
Marseille Winter	11.9	6.17	1.12	0.86	1.58	0.09	Chazeau et al. (2021)	
Marseille Spring	8.09	3.86	1.06	0.70	1.13	0.04	Chazeau et al. (2021)	
Barcelona (Annual)	9.85	4.10	1.70	1.05	1.35	0.06	Via et al. (2021)	

365 * Cyprus Regional background

366 ** European urban background = Barcelona (Spain) + London (UK) + Prague (Czech) + Tartu (Estonia) + Zurich (Switzerland)

367 ***Southern European regional background = Ersa (Corsica, France) + Finokalia (Crete, Greece)

368 The main difference between the cold and warm periods lies in the decrease in the concentration of carbonaceous aerosols

369 (OA, BC) and NO₃⁻ by almost a factor of two. Several phenomena can explain this significant seasonal variation: the absence

of a domestic heating source (mainly biomass burning as explained in Fig. 2); the absence of Middle East air masses during

371 the warm period (see discussion later on); the increase in the Planetary Boundary Layer Height (PBLH) above Nicosia (Fig.

372 S8) enhancing vertical dilution of local emissions during the warm period and therefore lowering ground-based concentrations;

373 less favourable thermodynamic conditions, with warmer and dryer air, also preventing the condensation of semi-volatile

374 species (e.g., ammonium nitrate). Sulfate concentrations do not exhibit a similar seasonal pattern and therefore seem to be less

375 affected by the above factors. On the contrary, the increase in photochemistry enhances the formation of sulfate aerosols, and

376 the decrease in precipitation enhances aerosol lifetime, strengthening the impact of long-range transport.

377 3.4. Diurnal variability of PM1 chemical constituents

378 Figure 4 shows the diurnal variability of the PM1 species derived from the ACSM and AE-33 for both the cold (Fig. 4a) and

379 warm (Fig. 4b) periods. The diurnal variability of the apportioned BC related to fossil fuel combustion (BCff) and wood-

380 burning (BC_{wb}) are also depicted here.

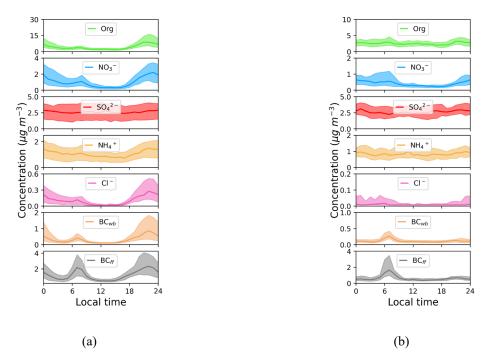


Figure 4: Median diurnal trends of the main submicron chemical constituents (OA, SO₄²⁻, NO₃⁻, NH₄⁺, Cl⁻ and BC) during the a) cold and
b) warm periods. The shaded area represents the 25th and 75th percentiles of the diurnals.

383 **Organic aerosols**: Organic aerosols dominate the cold period PM₁ concentration levels, exhibiting a night-time maximum 384 above 12 µg m⁻³ and a second smaller maximum at 4 µg m⁻³, coinciding with local traffic rush hour (06:00-09:00 LT). Elevated 385 OA concentrations in the cold period during the night (max at 22:00 LT) are a common, well-documented feature in many 386 urban environments across Europe and the Mediterranean (e.g., Florou et al., 2017; Stavroulas et al., 2019; Chazeau et al., 387 2021). They can be attributed to higher emissions from domestic heating, evening traffic peak and cooking activities. The 388 strong correlation between OA and BC_{wb} ($R^2=0.81$; N=2934; Fig. S9) suggests that residential wood burning is an important contributor to this nighttime peak. Interestingly, this peak is not significantly amplified by a lower PBLH during night-time, 389 390 which seems to remain relatively stable with no significant diurnal variability during the cold period (Fig. S8). It is also worth 391 noting that background OA concentrations observed both at the end of the night and middle of the day, when local emissions 392 are minimal, remain relatively high at around 3 µg m⁻³. The diurnal variability of OA is much less pronounced during the warm 393 period, suggesting a more important contribution of regional sources to OA compared to the strong dynamic of local emissions. 394 The assumption of a more important contribution from regional OA during the warm period is further supported by a mean 395 OA concentration of 2.83 µg m⁻³ (Table 2) that is close to the averaged OA concentrations of 3.2 µg m⁻³ reported for a 2-year 396 period continuous observations with Q-ACSM (2015-2016) at the rural background site of the Cyprus Atmospheric 397 Observatory at Agia Marian Xyliatou (CAO-AMX), at a distance roughly 40 km from Nicosia (Chen et al., 2022). During the 398 warm period, a small OA peak remains visible in the morning, with a similar amplitude to the cold season, likely to be related to traffic emissions. A second peak can be observed at 21:00 LT (not observed in BC), potentially originating from cooking 399 400 activities. Heavy oil combustion from shipping could possibly contribute to OA. Further to the poor contribution of shipping 401 emission on OA, a model study of sources of organic aerosols in Europe using CAMx (Jiang et al., 2019) showed that the contribution of "other anthropogenic sources" (gathering shipping, industry, and energy production) on OA (POA+SOA) was, 402 403 typically, of the order of 10% during summer and winter in the Eastern Mediterranean region close to Cyprus. Based on a simple receptor model, PM2.5 source apportionment performed in Nicosia, Achilleos et al. (2016) showed that the contribution 404 from shipping is approximately 8% to PM_{2.5}. Most of the transported mass is attributed to $SO_4^{2^2}$ with a minor contribution from 405 carbonaceous aerosols. In conclusion, shipping emissions are likely to play a minor role in OA concentrations. 406

407 Black carbon: During the cold season, BC follows a bimodal diurnal pattern, which can be further apportioned by focusing 408 on its source-specific components BC_{ff} and BC_{wb}. The fossil fuel component exhibits two maxima, one in the early morning, 409 coinciding with traffic rush hour, and one in the late afternoon, most probably related to both traffic and an increase in energy 410 demand due to domestic heating (see discussion later on). On the other hand, BC_{wb} diurnal variability is dominated by a night-411 time maximum (20:00 - 01:00 LT), peaking one hour after BCff and linked to wintertime residential wood-burning emissions, 412 contributing up to 33 % of total BC. During the warm season, the BC diurnal pattern is characterised by the absence of a night-413 time maximum while still exhibiting a significant peak in the morning, dominated by BCff. The very low contribution of 414 biomass-related combustion particles during the warm period, as previously noted from m/z 60 in Fig. 2, is further supported 415 here, with BC_{wb} exhibiting a nearly flat diurnal variability with close-to-zero mass concentrations. The contribution of shipping in the Mediterranean on Black Carbon (BC) concentrations was investigated from model estimates by Marmer et al. (2009) 416 based on three (3) most commonly used ship emissions inventories: 1) EDGAR FT by Olivier et al. (2005), 2) Evring et al. 417 418 (2005), and 3) EMEP by Vestreng et al. (2007). Results showed that shipping emissions were contributing to typically 15-25% 419 of BC in the E. Mediterranean, far from the shipping routes (which is the case for Cyprus). A similar result was found from a 420 more detailed (Positive Matrix Factorization) PM_{2.5} source apportionment analysis performed in Nicosia in 2018, with heavy 421 oil combustion contributing 7% to PM2.5 (Bimenyimana et al., 2023 under review), and the relevant factor containing less than $0.1 \mu g \text{ m}^{-3} \text{ of EC}.$ 422

423 Secondary inorganic aerosols: During the cold season, non-refractory nitrate and chloride detected by the Q-ACSM are 424 mostly present in the form of semi-volatile NH₄NO₃ and NH₄Cl (Guo et al., 2017; Theodosi et al., 2018). They show a night-425 time maximum (Fig. 4-a), reflecting the presence of gas precursors (NH₃, HNO₃, HCl) and the more favourable thermodynamic 426 conditions with lower temperatures, higher relative humidity, and condensation sink due to high PM concentrations of 427 combustion aerosols (traffic, domestic heating). Additionally, there is a smaller morning NO₃-peak, most probably linked to 428 traffic (Foret et al., 2022). This is not observed for chloride, suggesting that HCl may not be as abundant in the morning 429 compared to the evening. The less favourable thermodynamic conditions during the warm period lead to very small 430 concentrations of semi-volatile NO3⁻ and Cl⁻ (Fig. 4b). As expected, sulfate does not show a pronounced diurnal pattern, 431 irrespective of the period, and pointing to regionally-processed aerosols (Fig. 4a,b).

432 **3.5. OA Source Apportionment**

433 **3.5.1 OA source apportionment during the cold period**

For the cold period, the optimal PMF result has been found using a 5-factor solution following the approach detailed in section 434 435 2.4. The identification of OA sources related to these 5 factors was then performed following the typical combination of information from i) OA mass spectra (Fig. 5a), ii) the correlation of each factor with source-specific tracers (see Fig. 5b), iii) 436 437 their diurnal variability (Fig. 6a), and iv) their daily (week days vs. week-end) pattern (also Fig. 6b). The five factors were 438 then assigned to the following sources: A primary BBOA (Biomass Burning Organic Aerosol), two primary HOA 439 (Hydrocarbon-like Organic Aerosol; HOA-1 and HOA-2) and two secondary OA sources, namely low-volatile MO-OOA (More-Oxidized Oxygenated Organic Aerosol) and semi-volatile LO-OOA (Less-Oxidized Oxygenated Organic Aerosol). 440 This source apportionment is presented and justified below for each factor: 441

442 **HOA-1 (Hydrocarbon-Like OA Type 1):** The mass spectrum of HOA-1 (Fig. 5a) is consistent with a fossil fuel (traffic) 443 combustion source that can be identified by the prevailing contributions of the ion series representing C_nH_{2n-1} (m/z = 27, 41, 444 55, 69, 83, 97, typical fragments of cycloalkanes or unsaturated hydrocarbon chains) and C_nH_{2n+1} (m/z = 29, 43, 57, 71, 85, 445 99, typical fragments of alkane chains). Hence, this factor mass spectrum is well correlated to eight selected HOA factors 446 related to vehicular traffic found in the literature (Fig. S10a) and relevant to European and Mediterranean environments. The 447 traffic-related origin of the HOA-1 factor can be further confirmed by the good correlation with BC_{ff} (R²=0.65; N=2934; Fig. S11a), benzene (R²=0.72; N=1165; Fig. S11b). The diurnal variability of HOA-1 shows a bimodal cycle with a sharp maximum during the morning rush hour with an amplitude similar to BC_{ff} (Fig. 6a), and a broader maximum in the evening possibly encompassing emissions from traffic and diesel-fired residential heating systems. In the weekly cycle, as depicted in Fig. 6b, the morning peak decreases on Saturday. It is nearly absent on Sunday mornings, aligned with the de-escalation of traffic emissions usually observed during weekend mornings.

453

454 **BBOA** (Biomass Burning OA): The mass spectrum of the site-specific BBOA factor (reported as BBOA_{ev} in section 2.4) 455 exhibits characteristic peaks at m/z 29, 60, and 73 (Fig. 5a), which are indicative of biomass burning (Crippa et al., 2014). The 456 mass spectrum is quite similar to other BBOA spectra found in the Mediterranean and Europe (Fig. S10c), with a key difference 457 here being the rather low contribution of a signal at m/z=43. The biomass burning-related origin of the factor is further confirmed by the strong correlation with BC_{wb} (R²=0.81; N=2934; Fig. S11c), benzene (R²=0.61; N=1162; Fig. S11d) and 458 459 levoglucosan (R²=0.94; N=125; Fig, S11e) a typical tracer of biomass burning (Fourtziou et al., 2017). The BBOA diurnal 460 pattern exhibits an expected well-marked night-time maximum around 22:00 LT, consistent with residential wood-burning 461 activities. This night-time maximum is observed throughout the week (Fig. 6a), confirming the important role of wood burning 462 for heating in the city. Interestingly, the higher concentrations of BBOA as well as BCwb (Fig. 6b) were observed on Sunday 463 evenings, pointing to the recreational use of fireplaces, leading to enhanced residential wood-burning emissions during the 464 weekend, a feature also reported in other sites in Europe and the US (Bressi et al., 2016; Rattanavaraha et al., 2017; Zhang et al., 2019). 465

466 **HOA-2 (Hydrocarbon-Like OA Type 2):** The mass spectrum obtained for this factor (Fig. 5a) is similar to the HOA-1 factor, 467 with high signals for the ion series $C_nH_{2n+1}^+$ and $C_nH_{2n-1}^+$. The main differences between these two factors occur in the relative 468 contribution of m/z 41 compared to m/z 43 and the relative contribution of m/z 55 compared to m/z 57, which are both much 469 higher for HOA-2 than for HOA-1. Furthermore, the contribution of signal to m/z 44 is more significant in HOA-2, which can 470 imply a mix of various sources and/or a possibly higher degree of atmospheric processing. Other discrepancies with HOA-1 471 concern its diurnal variability, with an intense maximum at night (Fig. 6a), and its average concentration levels, which are 472 almost three times higher than HOA-1.

473 Influence of cooking activities: The HOA-2 diurnal profile has a small peak at 13:00 LT and a significantly higher one at 21:00 474 LT, effectively coinciding with typical meal times in Cyprus as well as those reported in the literature for Greece (Siouti et al., 2021), therefore indicating the influence of cooking activities to this factor. When plotting f_{55} vs f_{57} (Mohr et al., 2012) and 475 476 colouring the data points by the corresponding time of day, a distinct pattern appears with data of higher f_{55} over f_{57} being 477 clustered to the top left of the triangle, close to the fitted lines representing cooking (Fig. S12) and coinciding with midday and 478 evening hours. The night-time maxima pattern is consistent throughout the week, with the higher concentrations being recorded 479 on Friday and Saturday evenings (Fig. 6b), in line with an expected food service sector activity increase as part of Nicosia 480 inhabitants' leisure in the weekend. The mass spectrum of HOA-2, even though left unconstrained, is highly correlated to COA 481 found in other studies (Fig. S10b) in both Mediterranean and continental European urban environments. Additionally, the non-482 negligible signal at m/z=60 points to the widely spread habit of meat charbroiling (Kaltsonoudis et al., 2017).

483 Influence of power plant emissions: A closer look at the diurnal variability of the HOA-2 factor shows a certain persistence of 484 this factor throughout the day, even when cooking activities are more or less absent (Fig. 6a). Such pattern could imply the 485 influence of other combustion sources, not necessarily of local origin. The influence of other combustion sources would also help to explain why HOA-2 average concentrations are roughly 3 times higher than OA related to traffic (HOA-1), as it is very 486 487 unlikely that cooking activities can contribute solely to the observed HOA-2 concentrations. A possible contributing source 488 could be related to the energy production sector on the island, which relies exclusively on heavy fuel oil. In a recent study, 489 Vrekoussis et al. (2022), utilizing satellite observations, have identified that power plants located to the North (Teknecik 490 powerplant, PP4, 362MW), North-East (Kalecik powerplant, PP5, 153MW) and South-East (Dhekelia power station, PP3,

- 491 460MW) of Nicosia at 22 km, 60 km and 38 km, respectively, are significantly contributing to columnar NO₂ concentrations
- 492 over the island. The importance of these emission hotspots, along with their location on the island, during both the cold and
- 493 warm periods is illustrated in Fig. S13 and shows, particularly for the Northern power plants (PP4, PP5), emissions as high as
- 494 the traffic-related NO₂ over Nicosia. Interestingly, in a source apportioning study on VOCs performed at the Cyprus
- 495 Atmospheric Observatory Agia Marina Xyliatou (CAO-AMX), a rural remote site 32 km southwest of Nicosia, Debevec et
- 496 al. (2017) have resolved a factor related to industrial activity/power generation, exhibiting a connection with winds arriving
- 497 from the wider eastern sector.
- 498 In order to assess the possible influence of Cypriot power plant emissions, the coupling of wind velocity, and wind direction 499 with the HOA-2 time-series was performed through NWR analysis (Fig S14b). This analysis highlights the association of 500 stagnant conditions (low wind speed / low dispersion) with high HOA-2 concentrations (i.e., night-time peaks), pointing to a more local origin for this OA source. On the other hand, different features appear when wind velocities are higher, showing 501 502 emissions originating from the NW and the E-NE sectors; i.e. downwind of power plants PP4 and PP5, although long-range 503 transport influence cannot be ruled out. This is illustrated by the NWR of sulfate (Fig. S14f), which shows a dominant E sector 504 likely to originate from regional emissions. Given the positioning of the sampling site, close to the edge of Nicosia's urban 505 fabric, with the Athalassa park lying to the east, such an observation can suggest the transport of plumes from the operating 506 powerplants, namely PP5 and PP3 to the city. Interestingly, a similar yet even clearer image stands for SO₂ concentrations – 507 only half of which are considered to be of urban origin (Vrekoussis et al., 2022) - measured at a suburban background site (NicRes) and a traffic site (NicTra) in the city (Fig. S14g-h), with elevated SO₂ concentrations being related to eastern winds 508 509 of higher velocity, further corroborating that power generation related polluted plumes, traveling through the Mesaoria plain 510 arriving to Nicosia can contribute to the HOA-2 factor.
- 511 Other combustion sources: Interestingly, chloride shows a good correlation with HOA-2 ($r^2=0.61$; N=2945; see Fig. S11f, Fig. 512 5b). Chloride detected by the ACSM is in the form of NH₄Cl (a secondary highly-volatile species). The source of this chloride 513 is still widely debated and may originate from industrial activity or municipal (plastic-containing) waste burning (Gunthe et 514 al., 2021). Another possible explanation of the good agreement between HOA-2 and chloride would be the use of Cl-rich coal 515 as a means for outdoor cooking in Nicosia could therefore reflect the influence of cooking activities that comprises a fraction 516 of the HOA-2 factor.
- 517 Less-Oxidized Oxygenated OA (LO-OOA): With elevated contribution of m/z 44, the mass spectrum of this factor is consistent with a secondary OOA source. A higher m/z 43 and a lower m/z 44 (Fig. 5a) compared to MO-OOA implies a less 518 519 oxygenated (less-processed) component (Mohr et al., 2012). Finally, the time series of this factor is quite similar to NO₃⁻, with an overall good correlation value ($R^2 = 0.67$, N=2943; Fig. S11h), highlighting its semi-volatile character. This is further 520 corroborated by the very good correlation of LO-OOA with chloride (R² = 0.73, N=2943; Fig S11i), another semi-volatile 521 522 compound measured by the Q-ACSM. The diurnal variation of LO-OOA displays 1.5 times higher concentrations during the 523 night compared to daytime (maximum of $1.84 \pm 0.31 \,\mu g \, \text{m}^{-3}$ at 22:00 LT; Fig. 6a); a pattern that is much more pronounced 524 than the variability observed for MO-OOA. This feature highlights that the presence of LO-OOA, is not exclusively controlled 525 by photochemical processes. Instead, changes in thermodynamic equilibrium (due to lower T and increased RH), favouring 526 the condensation of gas-phase semi-volatile material on the one hand, and intense night-time chemistry (gas phase or 527 heterogenous) on the other hand, are among the processes that may account for the rapid night-time formation of LO-OOA. 528 Atmospheric processing of biomass burning OA during periods of low photochemical activity (such as in winter or at night), 529 also known as "dark" aging, has been reported recently (Kodros et al., 2020; Jorga et al., 2021) and could have contributed to 530 the observed night-time formation of LO-OOA. Notably, the weekly cycle of LO-OOA, and its night-time maxima, appears 531 to have the same pattern and intensity as those observed for BBOA (e.g., low peaks on Tuesday/Thursday, maximum on 532 Sunday) (Fig. 6b). On the other hand, the factor is correlated with both BBOA ($R^2=0.81$; Fig. S11k) and BC_{wb} ($R^2=0.66$; Fig. 533 S11j). This observation could indicate a biomass-burning contribution to LO-OOA through fast oxidation of primary

534 emissions, supported by several studies showing biomass burning linked to OOA sources at night (Stavroulas et al., 2019;

535 Kodros et al., 2020; Chen et al., 2021).

More-Oxidized Oxygenated OA (MO-OOA): The MO-OOA factor typically accounts for secondary organic aerosol formed 536 537 in the atmosphere from gas-to-particle conversion processes of VOCs and their products, as well as atmospheric ageing of primary OA (Petit et al., 2015; Stavroulas et al., 2019). Numerous VOC sources can contribute to OOA but lose their mass 538 539 spectrum fingerprint owing to extended oxidation due to photochemical aging, which leads to enhanced signal at the m/z 44 fragment (CO_2^+), a dominant tracer for OOA (Ng et al., 2011). The predominance of m/z 44 and the near absence of m/z 43 in 540 541 the mass spectrum of the resolved MO-OOA factor (Fig. 5a) points to highly oxidized/aged secondary OA (i.e., originating 542 from long-range transport). This is further supported by the relatively good agreement (R²=0.55; N=2943; Fig. S111) between 543 concentrations of MO-OOA and sulfate (Fig. 5b), a species of regional origin (Sciare et al., 2003). Nevertheless, the diurnal 544 variability of MO-OOA does not closely follow sulfate showing a small increase of 20-30% every evening (Fig. 6a,b), which 545 furthermore cannot be explained by atmospheric dynamics (c.f. the negligible PBLH diurnal variability for the cold period 546 shown in Fig. S8). Alternatively, this would suggest that a fraction of MO-OOA is produced locally through night-time 547 oxidation mechanisms as previously observed for LO-OOA. Similar nighttime increases of high oxygenated OA factors, 548 related to local sources, have been reported in both northern European urban sites (Zhang et al., 2019; Lin et al., 2022) as well 549 as in the Eastern Mediterranean urban environment (Athens, Greece), where a link to oxidized primary residential wood 550 burning emissions as a potential driver of the low volatility OOA factor diurnal variability, was also suggested (Stavroulas et al., 2019). 551

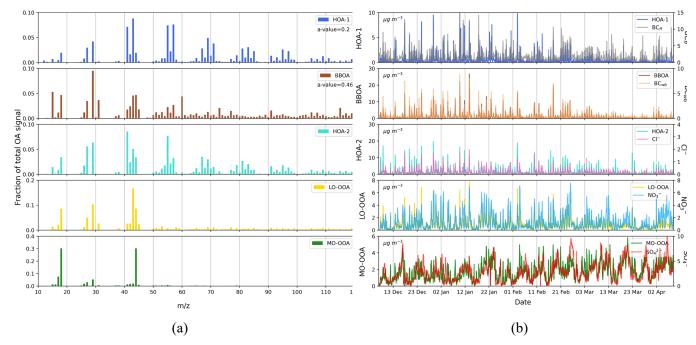
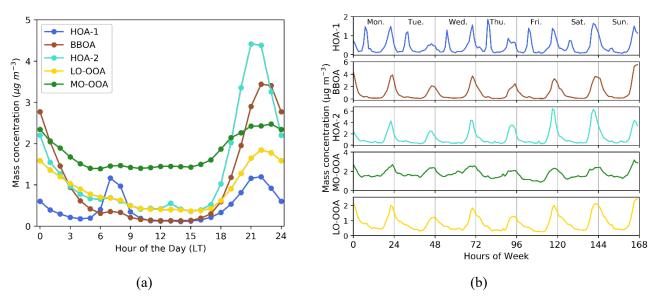
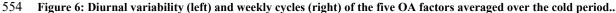


Figure 5: Mass spectra of the PMF (a) and time series of the five OA factors resolved along with corresponding tracer compounds(b) for the cold period.





555 **3.5.2.** OA source apportionment during the warm period

556 For the warm period, the optimal PMF solution was obtained using a 4-factor solution (HOA-1, HOA-2, MO-OOA, LO-OOA).

557 As expected, the BBOA factor could not be resolved, as previously highlighted by the low concentrations at m/z 60 reported

558 during this period (Fig.2). Again, the identification of OA sources related to the 4 OA factors was performed following the

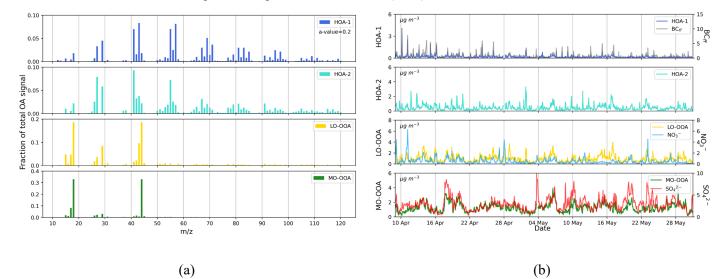
typical combination of information from i) OA mass spectra (Fig. 7a), ii) the correlation of each factor with external sourcespecific tracers (Fig. 7b and Fig. S15), iii) their diurnal variability (Fig. 8a), and iv) their daily (weekdays vs. weekend) pattern (also Fig. 8b). The mass spectra profiles for the 4-factor PMF solution during the warm period (Fig. 7a) were quite similar to

562 the ones from the cold period (Fig. 5a).

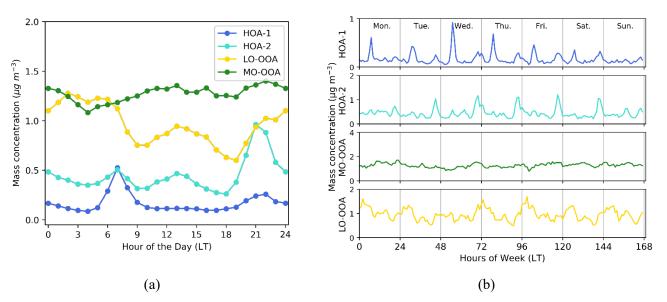
563 HOA-1: For the warm period, an a-value of 0.2 was selected for constraining the HOA-1 factor, again using the Ng et al. (2011b) HOA profile as a reference. The resolved factor profile is nearly identical to the one obtained for the cold season (R^2 564 565 = 0.99, Fig. S10a). It is also very well correlated to traffic-related HOA factor profiles found in other Mediterranean 566 (Kostenidou et al., 2015; Gilardoni et al., 2016; Florou et al., 2017; Stavroulas et al., 2019) and European cities (Lanz et al., 2010; Crippa et al., 2014) as depicted in detail in Fig S10a. The HOA-1 time series follows the same pattern as the 567 corresponding traffic-related HOA-1 factor reported for the cold period, showing a good correlation with BCff, (R²=0.62, 568 569 N=1259; Fig. S15a). Its diurnal variability exhibits a bimodal pattern, with a typical sharp maximum in the morning (07:00 570 LT) and a smaller peak during the evening (Fig. 8a). On a weekly basis, this diurnal variability tends to be less pronounced on 571 Saturdays and nearly absent on Sundays (Fig. 8b), reflecting reduced commuting during the weekend.

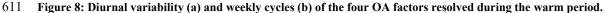
572 HOA-2: The HOA-2 factor still shows elevated concentrations during the warm period, close to 3 times higher than HOA-1 (Table S3). The profile remains quite unchanged between the cold and warm periods ($R^2 = 0.92$; Fig. S10b), pointing to similar 573 574 sources. No correlation was observed with chloride, which may be expected due to unfavourable thermodynamic conditions hindering NH₄Cl formation as well as the lack of significant chloride sources during this period. A night-time maximum of 575 576 HOA-2 is still observed when investigating the factor's diurnal variability (Fig. 8a). Furthermore, a somewhat broader, 577 compared to the cold period, maximum in the middle of the day (Fig. 8a) can also be observed. When going through the weekly variability, this midday maximum is particularly well defined on Sundays (Fig. 8b), while the evening peaks of Sundays and 578 579 Mondays are the lowest. The above observations remain consistent with the cold period assessment, that HOA-2 is on the one 580 hand linked to cooking activities. For household activities are expected at noon and evenings, while for restaurants, activity 581 peaks on Sunday noon and is lower on Sunday evening and Monday, reflecting the fact that such businesses remain closed on 582 the first day of the week (Fig. 8b). On the other hand, the overall offset of HOA-2 observed against the HOA-1 diurnal profile 583 persists, suggesting somewhat permanent background HOA-2 concentrations that cannot be explained by cooking activities

- alone. A contribution to this source by continuous emissions from power plants (see space-based (SP5-TROPOMI) vertical
- 585 columns of NO₂ during the warm period in Fig.S13d) should be sought. In addition, the HOA-2 NWR plot for the warm period
- 586 reveals an even more significant enhancement of concentrations when moderate winds blow from the E-SE (Fig. S16b), a
- 587 trend also observed for SO₂ during the same period (Fig S16e,f).
- 588 The above observations remain consistent with our assessment for the cold period: the HOA-2 factor consists of a mixed OA
- 589 source that contains cooking activities (inc. coal combustion) and emissions from the powerplants located on the eastern part
- 590 of the island. Indeed, the HOA-2 midday maximum can be linked to an increase in electricity demand at that time of day during
- the warm period due to an increase in air conditioning usage (Cyprus' NECP 2021-2030, 2019).
- 592 **LO-OOA:** The LO-OOA factor profile exhibits some differences from the one resolved for the cold period ($R^2 = 0.66$), as illustrated in the correlation matrix of comparison to selected factor profiles found in the literature (Fig. S10d) while being 593 594 very similar to those obtained in Athens/Piraeus during summer (Bougiatioti et al., 2014; Stavroulas et al., 2021). The LO-OOA time-series shows a low agreement with NO₃⁻ ($R^2 = 0.31$; N=1259; Fig. S15c) poorer than the observed correlation 595 596 during the cold period (Fig. S11h). The diurnal pattern of the factor (Fig. 8a) shows maximum concentrations persisting 597 throughout the night and early morning, while a secondary maximum during the midday can be observed. But overall, the 598 diurnal pattern of LO-OOA is rather flat compared to the cold period, suggesting that local production may not be so important 599 at that time compared to a less variable regional background. Interestingly a midday hump similar to the one observed for 600 HOA-2 is present, suggesting a common origin.
- MO-OOA: The factor profile of MO-OOA resolved during the warm period is strikingly identical to the profile found in the 601 602 cold period (their R² is almost 1; Fig. S10e), while being excellently correlated to other highly oxygenated OA factors resolved 603 in both the urban and regional background in the Eastern Mediterranean (Bougiatioti et al., 2014; Stavroulas et al., 2019, 2021) 604 as well as in continental Europe (Crippa et al., 2014). The winter night-time peaks are not observed anymore (Fig. 8a), with 605 the factor's diurnal pattern exhibiting much less variability, highlighting its dominant regional character. The time series of MO-OOA correlates well to SO₄²⁻ (R²=0.53; N=1259; Fig.S15b), confirming this regional and highly processed origin. The 606 607 concentration levels of MO-OOA during the warm period are lower than in the cold (Table S3). However, its relative 608 contribution to total OA during the warm period remains similar (45 %).



609 Figure 7: Mass spectra of the PMF (a) and the time series of the four OA factors resolved along with corresponding tracer compounds
610 (b) for the warm period.





612 **3.6. Spatial and seasonal variability of OA sources**

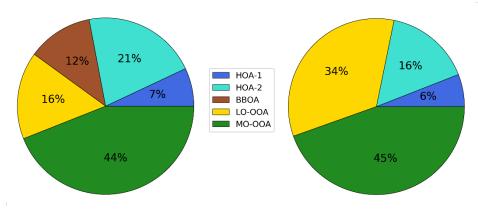
613 3.6.1. Seasonal variability of OA sources

Primary OA: The mass concentration of the three primary OA factors (HOA-1, HOA-2, BBOA) represents as much as 40 % of total organic aerosols during the cold period (Fig. 9), with POA contribution significantly decreasing in the warm period (22% to total OA) due to the absence of the significant residential wood burning source which during the cold period accounted for 12% of total OA. The important contribution of primary sources in Nicosia has also been highlighted earlier by the rather low OA/OC ratio of 1.42 (Section 3.1). In a recent publication covering several European sites, Chen et al. (2022) reported that in urban sites, solid fuel combustion-related OA components were 21.4 % of total OA during winter months, higher than what is found for BBOA in Nicosia, owing to the rather milder winters in the city.

The traffic-related primary factor in Nicosia (HOA-1) was found to be rather stable in terms of contribution to total OA across this study's two seasons, averaging 7% and 6%, respectively, for the cold and warm periods, being lower than the figure reported in other European Urban sites (12.7%, Chen et al., 2022). On the other hand, the HOA-2 factor represents ca 2/3 of the total HOA in Nicosia with little variation from winter (72 %) to summer (66 %) to total HOA (Fig. 9). Comparing it with COA in urban locations resolved by Chen et al. (2022), during both winter (14.4% compared to 21% in the cold season in Nicosia) and spring (15% versus 16% in Nicosia during the warm season), the higher values reported in Nicosia further support the assumption that the HOA-2 represents a mixed combustion source.

628 Secondary OA: A higher degree of oxidation is observed for the LO-OOA factor during the warm period, given the much higher signal contribution at m/z 44 than the respective cold period factor. This discrepancy, reported in several studies (Huang 629 630 et al., 2019; Duan et al., 2020), is explained by higher photochemistry during the warm period, which promotes the oxidation 631 of OA, resulting in a LO-OOA profile with a higher m/z 44 fraction. This result is also consistent with a less-oxidized LO-632 OOA formed during the cold period from night-time chemistry. The range of LO-OOA concentration levels is different between cold and warm periods (0.05-7.74 µg m⁻³ and 0.05-4.00 µg m⁻³, respectively), while the mean concentrations for both 633 periods are similar (0.86 and 0.95 µg m⁻³ for cold and warm periods respectively). The contribution of LO-OOA relative to 634 635 total OA is double during the warm period compared to the cold, reflecting both the absence of the biomass burning source as 636 well as the prevailing conditions favoring atmospheric processing of primary OA and SOA precursors. During the cold period, 637 LO-OOA intense peaks suggest an influence from local emissions, while during the warm period, the less-variable LO-OOA 638 diurnal variability highlights the influence of more intense photochemical processing at medium-to-large geographical scale. 640 cases than the average MO-OOA contributions reported for other European urban sites (Chen et al., 2022) underlining the

641 importance of highly processed secondary OA over Nicosia (Fig. 9).



642

Figure 9: Relative contribution of PMF resolved OA sources to total OA for the cold period (left) and the warm period (right), respectively.

645 3.6.2. Geographic origin of OA sources

The geographic origin of OA sources (local vs regional) is further assessed here using both Non-parametric Wind Regression
(NWR) analyses as well as the regional scale coupling concentrations to air mass back trajectories through PSCF.

648 Cold period: During this period, primary OA factors, especially HOA-1 and BBOA, have an expected strong local component 649 that is characterized by high concentrations at low wind speeds (hourly average 1.4 m s^{-1}) when winds are originating from the 650 W-SW sector (Fig. S14a,c), pointing to the busy highway connecting Nicosia to the other major cities in the island while 651 integrating the highly populated residential areas of Strovolos and Lakatamia municipalities. (Fig. 1c). As discussed earlier, 652 the HOA-2 factor, apart from its local influence (also in the W-SW sector), exhibits significant concentrations related to higher 653 wind speeds from the NW and the E-NE sectors that could originate from power plants but also possibly from long-range 654 transport. Interestingly, a small local contribution from the city, still within the W-SW sector, can also be observed for both 655 LO-OOA and MO-OOA, consistent with the peaks observed that could originate from local night-time chemistry. Still, high 656 concentrations of MO-OOA (and, to a lesser extent LO-OOA) are observed with high wind speeds and Eastern directions (Fig. S14e,d). Although the contribution of the power plant PP5 located in the East sector (Fig. S13c) cannot be excluded, PSCF 657 658 analysis points out that the hotspots of MO-OOA can be traced in neighbouring countries (eg. Syria, Lebanon and South Turkey) in the middle East (Fig 10a). These areas also represent hotspots of SO₄²⁻ according to PSCF analysis (Fig. S17a). 659 Warm period: Given the generally higher wind speeds recorded, in comparison to the cold season (average of 1.93 m s⁻¹ vs. 660 661 1.36 m s⁻¹ in the cold period), all OA factors show elevated concentrations coupled with higher wind speeds. The most striking 662 result is the major influence of the E-SE sector for all OA sources. However, this sector is upwind of Nicosia and, therefore,

663 poorly influenced by local city emissions. As noted previously, for the cold period, long-range transported OA from the Middle 664 East is expected to be the main driver to explain the influence of the E-SE sector, at least for LO-OOA and MO-OOA (Fig. S16c,d). This is again confirmed by the PSCF results reported in Fig. 10b for the warm period. The HOA-1 factor still shows 665 666 maxima for low wind speeds (<5 km h⁻¹) characteristic of local emissions and the SW-S direction but also exhibits significant 667 contribution related to the E-SE sector. Although the influence of the power plant PP5 on HOA-2 is expected, the contribution 668 of this source can not be excluded for HOA-1 as well. On the other hand, quantification of the Middle Eastern contribution to 669 the HOA-2 factor remains to be assessed since the current dataset cannot provide sufficient information on separating the 670 contribution of power plants on the island versus more regional Middle East emissions (Fig. S165b). Although this hypothesis 671 needs further investigation, the presence of HOA-2 in the Middle East would be consistent with recent findings highlighting

the importance of OC emissions from diesel generators used in Lebanon as a means of complementary power generation (Fadelet al., 2022).

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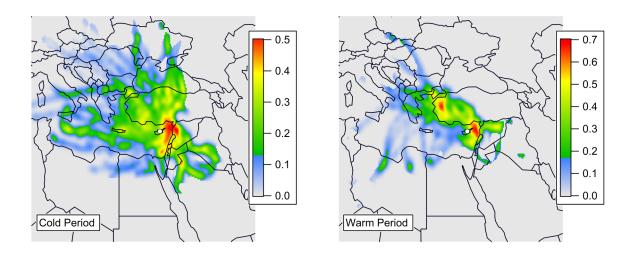




Figure 10: PSCF plots for MO-OOA during the cold and warm periods. The color scale represents the probability of air parcels arriving at the receptor site (white dot) for measured concentrations higher than the 75th percentile.

In conclusion, based on the relative contribution of OA factors (Fig. 9) and the NWR analysis (Fig. S14, S16), it can be reasonably assumed that a significant amount of measured OA in Nicosia originates from long-range transport with the Middle East being the major source region, during both cold and warm periods. This is the first time that such a high contribution of OA from the Middle East is highlighted over Cyprus. Assuming that biomass combustion and biogenic emissions of OA in the desert regions of the Middle East are relatively limited, these results suggest that most of the primary and secondary OA originating from the Middle East could be of fossil fuel origin, which is consistent with the previously reported extensive use of oil in this region.

685 3.7. Spatial and seasonal variability of BC sources

The above conclusion on the influence of primary and secondary OA sources from the Middle East region, and its strong fossil fuel origin, motivates a careful examination of the geographic origin and sources of BC concentrations recorded in Nicosia.

Baseline (i.e., lowest) BC_{ff} concentrations are typically observed in the middle of the night and in the middle of the day when local emissions are at their minimum (See Fig. 4). As such, these background concentrations can be considered as a first qualitative indicator of background BC_{ff} concentrations of regional origin. Interestingly, these baseline BC_{ff} concentrations appear to be in phase with those of sulfate (Fig. 11), as well as the MO-OOA factor derived from the OA PMF analysis. This observation points to the possible use of MO-OOA as a tracer for regional BC_{ff} . Hence, it brings further evidence of the importance of regional emissions on carbonaceous aerosol concentrations in Nicosia.

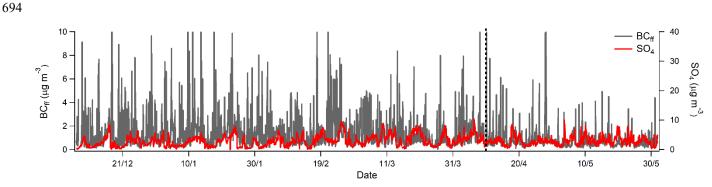


Figure 11: Temporal variability of BCff and SO4²⁻ concentrations during the entire measuring periods.

The assumption that transported regional pollution can affect BC_{ff} concentrations in Nicosia can be further supported by investigating the BC_{ff} NWR polar plots for both the cold and warm seasons (Fig. S18a,b). Elevated concentrations related to local emissions were observed for calm conditions with low wind speeds (<5 km h⁻¹) in the SW sector, as previously observed for HOA-1. Interestingly, BC_{ff} NWR plots show a distinct contribution at higher wind speeds (~15 km h⁻¹) and the NE-SE (Middle East) sector, during both the cold and warm periods, with estimated concentrations of roughly 1.5 μ g m⁻³, further support the major role of the Middle East in the observed BC concentration levels in Nicosia (Fig S18 a,b).

701 BC source apportionment: In order to better assess the relative contributions of the multiple primary OA sources (HOA-1,

- 702 HOA-2) and to quantify the contribution of long-range transport from the Middle East to BC_{ff}, a multilinear regression (MLR)
- 703 model was tentatively performed using the principle of co-emission of BC_{ff} and organic species by the different sources
- 704 (Chirico et al., 2010; Laborde et al., 2013). This approach, used recently by Poulain et al. (2021), assumes that at any given
- 705 time (t), BC_{ff} mass concentration is the sum of BC from traffic (traced by HOA-1), from a mixed combustion source (traced

 $[BC]_{regional} = c \times [MO-OOA] (5)$

- 706 by HOA-2), and from long-range transport (traced by MO-OOA), as follows:
 - $[BC]_{ff} = [BC]_{traffic} + [BC]_{mix \ combustion} + [BC]_{regional} (2)$
- 707 708

709

With:

- $[BC]_{traffic} = a \times [HOA-1] (3)$
- 710 $[BC]_{mix combustion} = b x [HOA-2] (4)$
- 711

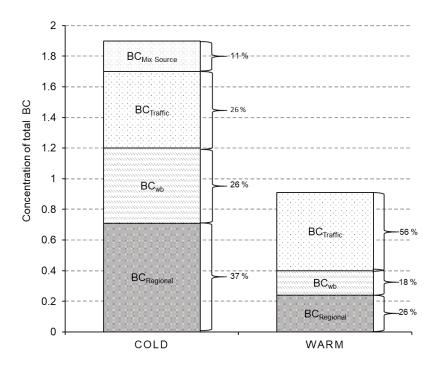
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712 Where a, b and c are coefficients derived from the multi-linear regression model.

713 The above approach assumes that primary HOA-1 and HOA-2 can trace BC_{traffic} and BC_{mix combustion}, respectively. This is 714 expected for traffic with a typical HOA-1/BCtraffic ratio with little variations. For HOA-2, this assumption is valid for the 715 fraction assumed to originate from power plant emissions and for some of the cooking activities (e.g., when using charcoal 716 combustion) but not necessarily all. As such, the uncertainties of this approach are expected to be higher for HOA-2 compared 717 to HOA-1. The use of MO-OOA to trace the regional source of BC would probably lead to even higher uncertainties because MO-OOA is also sensitive to atmospheric photochemical processes and does integrate multiple sources. Nevertheless, this 718 719 latter assumption is believed to be acceptable given the good agreement reported above between baseline concentrations of 720 BC_{ff} and MO-OOA (Fig. S19) and the above conclusions that carbonaceous aerosols originating from the Middle East are 721 expected to be dominated by fossil fuel combustion. Note that MO-OOA was preferred here to LO-OOA to trace regional 722 emissions due to the latter's somewhat more local character.

Combing equations 2-5 provides the multilinear regression model with the free regression parameters a, b, c, which are fitted to the time-resolved BC_{ff} mass concentration measured by the Aethalometer and PMF results for the ACSM data:

- $[BC]_{ff} = a x [HOA-1] + b x [HOA-2] + c x [MO-OOA] (6)$
- Previous studies have shown that MLR models have enhanced explanatory power when primary emissions dominate (Laborde et al., 2013). To reduce this potential bias, the MLR model was applied distinctly for the two seasons separately.
- During the cold period, a very good correlation between measured and modelled BC_{ff} was obtained ($r^2 = 0.70$; N = 2942), with 728 729 the modelled BC_{ff} explaining 84 % of the measured one (Fig. S20a). The regression coefficients a (HOA-1), b (HOA-2) and 730 c (MO-OOA) were found to be 1.11 ± 0.03 , 0.15 ± 0.01 and 0.41 ± 0.01 , respectively. Regarding the warm period, it was not 731 possible to obtain a positive value for b (HOA-2). A correlation between long-range transported HOA-2 and MO-OOA is, 732 among other, a reason that can be proposed to explain why it has not been possible to extract a BC_{mix source} factor here. Therefore, 733 BC_{ff} was only apportioned using HOA-1 and MO-OOA. A good correlation between measured and modelled BC_{ff} was 734 obtained (r²=0.62; N=1251), with the modelled BC_{ff}, explaining 83% of observations (Fig S20b). The regression coefficients a (HOA-1) and c (MO-OOA) were found to be 3.05 ± 0.07 and 0.19 ± 0.01 , respectively. 735
- 736 The combination of the Aethalometer model (apportioning BC_{ff} and BC_{wb}) and the MLR model (apportioning BC_{traffic}, BC_{mix}
- 737 source, and BCregional) was performed to obtain an integrated picture of BC sources in Nicosia for both periods (see Fig. 12).



738

739 Figure 12: BC sources during the cold and the warm period in Nicosia

740 Spatial and seasonal variability of BC sources: During the cold period, BC was found to originate from four different sources 741 denoting the complexity of combustion sources of different origins in Nicosia. BC_{regional} is the dominant source of BC (37%), 742 while traffic, wood burning, and mix source are estimated to contribute to 26 %, 26% and 11% of BC, respectively. From the 743 perspective of BC_{ff} sources, long-range transport, traced by MO-OOA, remains the largest source of BC_{ff} during the cold 744 period, contributing 63 %, while BCff from local emissions constrained with HOA-1 and HOA-2 represents 24% and 13%, 745 respectively (Fig S21). In other words, more than half of BC_{ff} in Nicosia was found to be regional and probably originated 746 from the Middle East during the cold period. This high contribution of regional BC_{ff} is quite unexpected for a medium-sized 747 European city like Nicosia, where local traffic is likely to be the main contributor to BC_{ff}. 748 Nevertheless, extra caution should be taken here. The obtained contribution of 63% for BC_{ff} regional should be seen as an 749 upper limit since a fraction of MO-OOA was shown to be of local origin during the cold period. During the warm period, the 750 picture remains similar, with traffic and wood burning representing two-thirds of BC (56 % & 18 %). Here, BC regional 751 contributed 26 % to total BC. From the perspective of BC_{ff} sources during the warm period, the long-range transport contributed 41 %, while BC_{ff} from local emissions constrained with HOA-1 represents 59 % (Fig S21). Although the two 752

753 models (Aethalometer and MLR) are associated with non-negligeable uncertainties, the BC source apportionment obtained

shows that local emissions cannot be considered only for BC, with demonstrated significant contribution of Middle East fossil
fuel emissions.

756 4. Conclusions

Near-real-time chemical composition of submicron aerosols and source apportionment of carbonaceous aerosols was 757 758 performed for the first time in Nicosia, a medium-sized European capital city (circa 250,000 inhabitants) in Cyprus located in 759 the Eastern Mediterranean and surrounded by Middle East countries with fast-growing population and increasing emissions of 760 air pollutants. Continuous observations were performed at an urban background site for approximately 6 months (between 7 761 December 2018 and 31 May 2019) in order to obtain a large and representative dataset capturing specific features - related to 762 both the cold and warm periods - such as domestic heating and regional transport. Measurements of the major fractions of PM_1 763 were carried out with a Q-ACSM and an Aethalometer complemented by a comprehensive suite of collocated instruments 764 (e.g., filter sampling, SMPS) to assess the quality of the acquired data further.

23

Unlike many European cities, no clear PM_1 pollution episodes of several consecutive days could be observed over Nicosia. However, very intense peaks (above 40 µg m⁻³, 1h averages) were recorded systematically every evening during the cold period. Carbonaceous aerosols (BC and OA) were identified as the main components of these peaks and were mostly attributed to local emissions from heating with little contribution from local meteorology (PBL height did not show significant diurnal variability during the cold period). Furthermore, a significant portion of PM₁ was found to be related to long range transported aerosol, while the influence of shipping emissions was estimated to be rather low (less than 8%).

- 771 Source apportionment of OA has been used to derive a local biomass burning OA (BBOA_{cv}) mass spectrum in order to 772 apportion the contribution of domestic wood burning properly. A total of five OA sources were identified during the cold 773 period, among which four are typically reported within urban environments (HOA-1, BBOA, LO-OOA, MO-OOA). An 774 additional one (HOA-2) was assigned as a mixture of several combustion sources, such as cooking as well as a significant 775 contribution from power plants located in the Northern part of the island. These power plants in addition, represent major 776 island-based hotspots of NO_x , as evidenced by satellite observations. Interestingly, a similar HOA-2 source was identified at 777 our regional background site (40 km distance from Nicosia; Chen et al., 2022), pointing to a possible influence from these 778 power plants to an extended part of the island. The impact of this specific source brings the OA contribution of primary sources 779 up to 40 % over Nicosia during the cold period. Few additional features were noticed for the other OA sources with 1) a typical 780 traffic-related (HOA-1) source observed during both seasons, 2) a biomass burning source (BBOA) related to domestic heating 781 enhanced at night during the cold season and accounting for 12 % of the total OA, 3) a less oxidized secondary (LO-OOA) 782 source of a semi-volatile character, influenced by local night-time chemistry, that was more oxidized (i.e., of a less local 783 character) during the warm period, and 4) a secondary (MO-OOA) source mostly of regional origin but also influenced by 784 night-time chemistry during the cold period.
- The geographic origin of each OA source was assessed for both seasons. Except for MO-OOA, which systematically shows a strong regional component, HOA-1, HOA-2, and LO-OOA exhibit a clear local origin during both seasons, and a more pronounced influence from the Eastern wind sector during the warm period. The prevalence of this sector is systematically observed for MO-OOA highlighting the major role of Middle East emissions in contributing to almost half of OA concentrations in Nicosia during both cold and warm seasons.
- 790 To further elucidate the influence of this complex mixture of OA sources on BC levels, source apportionment of BC was 791 performed by combining i) the aethalometer model to separate BC into its fossil fuel (BC_{ff}) and wood burning components 792 (BC_{wb}), and ii) a multi-linear regression model to apportion the contribution to BC_{ff} from traffic (constrained by HOA-1), mix 793 combustion sources from cooking and power plants (constrained by HOA-2), and long-range transport from the Middle East 794 (constrained by MO-OOA). Although several assumptions and uncertainties are associated with this approach, it has shown to 795 provide an interesting tool for reconstructing the BC concentrations derived experimentally. Such BC apportionment 796 performed for both cold and warm seasons solidified the conclusions reached through the OA source apportionment, with 797 almost half of BC_{ff} being of regional origin, with the Middle East playing an important role. This result is quite unexpected 798 given that local traffic emissions are usually considered the dominant contributor to BC_{ff} in urban background environments. 799 These conclusions have numerous implications related to PM regulation and the efficiency of local abatement strategies (in 800 particular regarding traffic emissions), health (combustion aerosols being considered as particularly adverse for human health),
- 801 and climate (major influence of light-absorbing aerosols from the Middle East fossil fuel emissions).
- More accurate OA and BC source apportionment i) with more co-located high-resolution measurements of specific trace metal and organic tracers, ii) better resolved OA mass spectra (e.g., from HR-ToF-AMS), iii) the use of various source-specific mass spectra fingerprints (e.g., from cooking or power plants), and iv) multi-site measurements (incl. both urban and regional background) will enable a more accurate estimation of local vs. regional fossil fuel emissions in Cyprus while better constraining the current regional efforts on air quality modelling and forecasting.
- 807

- 808 Data availability: All data used in this study can be accessed here: <u>https://doi.org/10.5281/zenodo.7802065</u>. More details on
- 809 the analyses are available upon request to the contact author Aliki Christodoulou (a.christodoulou@cyi.ac.cy).
- 810
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